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Monique Fort, Regis Braucher, Didier Bourles, Valery Guillou, Lila Nath Rimal, et al.. Geomorphic impacts, age and significance of two giant landslide dams in the Nepal Himalayas: Ringmo-Phoksundo (Dolpo District) and Dhampu-Chhoya (Mustang District).. EGU General Assembly 2014, Apr 2014, Vienne, Austria. 1 p. hal-01262118

HAL Id: hal-01262118

https://hal.science/hal-01262118

Submitted on 26 Jan 2016

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# Geomorphic impacts, age and significance of two giant landslide dams in the Nepal Himalayas: Ringmo-Phoksundo (Dolpo District) and Dhampu-Chhova (Mustang District).

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### 1. INTRODUCTION: MAIN ISSUES

Large catastrophic slope failures have recently retained much attention in the northern dry Himalayas (Hewitt 2009). They play a prominent role in the denudation history of active orogens at a wide range of spatial and time scales (Korup & Clague, 2009), and they impact durably landforms and process evolution in upstream catchments. Their occurrence mostly results from three different potential triggers: earthquakes, post-glacial debuttressing, and permafrost melting. We focus on two examples of giant rock slone failures that occurred across and north of the Higher Himalaya of Nepal and assess their respective influence on the regional, geomorphic evolution.

Mass wasting failure		Volume (10 <sup>8</sup> m <sup>3</sup> )	Age	Geological context	Type of failure	Nature of material	Geomorphic impact	References
1	Phaksumdo- Ringmo rockslide (Suli Gad valley)	15	Presumably 30-40,000 yr (no dating)	AHTS (limestones)	In situ collapse	Disintegrated, recemented fragments and blocks	landslide dam lake, volume: 1500 Mm <sup>3</sup>	Yagi 1997; Ibetsberger & Weldinger 2000
2	Dhumpu-Kalopani rock-avalanche (Kali Gandaki)	30	29,674 +/- 1022 yr ( <sup>55</sup> Be); 4,1 +/- 0,6	HHC (augen and blockic gneisses)	Banc-to-banc rock-slide rock- avalanche	bedrock slabs + breccia and pulverized rocks	landslide dam lake of Marpha, volume: 4000 M m <sup>3</sup>	fort 2000; Zed all, 2009 + per unpubl. dates t D. Bourlès et al
3	Pokhera glant rock- avalanche/debris flow (Seti khola)	40	500 yr +/- 100 ( <sup>14</sup> C)	Crystallines (gneisses) + NHTS + tills	Catastrophic collapse of Higher Himalayan Thrust Front	Gigantic debris flow (channelized along the Seti khola gorges)	Damming of adjacent, bributary valleys (lakes)	Fort 1987, 198 Fort & Peulvast 1985
4	Manang rock- avalanche (Upper Marsyangdi khola)	150	Interstadial ? Pre-Last Glacial advance	NHTS (marly and hard limestones)	Catastrophic collapse of recumbent fold axis	diamicton (all clast sizes), recemented breccia	Inear, gully erosion + karstic like landforms	fort 1993; Weidinger 2000
s	Dhikur Pekhari rocksilde (Upper Marsyangdi khola)	10	Post Glacial	MHTS (limestones and schists)	In situ collapse	Brecciated cataclastic and recemented blocks	shallow, relict Dhikur pokhari (lake)	Karup & ali, 20
6	Latemrang landslide (Upper Marsyangdi khola)	> 55 (1 already removed)	5400 yr BP	HHC Crystalines (gneisses) + NHTS + tills	Sledge-like sliding	Brecciated cataclastic and recemented blocks	landslide dam lake, volume 160 M m <sup>3</sup>	Karup & all, 20
,	Tsergo Ri rockslide (Langtang khola)	100	34,2 +/- 10,4 (10 Be); 25-40 (TL) 51+/-13 (FT)	HHC (gneisses and intrusive granite)	Post glacial gravity instability	Microbreccia (nylonites and pseudotachylites)	Linear, gully erosion	Barnard & al., 2006; Weldinge all, 2002; Taks & all, 2007
8	Khumjung rock- avalanche (Upper Dudh Kesi)	21	Pre-Last Glacial advance	HHC (gneisses )	Post glacial gravity instability	Brecciated cataclastic, blocks		Heuberger, 198





ectonic event



The Pokhara gravels fill (#3): result of a gigantic rockavalanche 500-years old. Mechanism? Catastrophic collapse of the Annapurna IV (thrust) front, possibly seismically triggered in 1505 (Fort 1987; 2010)

### 2. DATING METHODS: COSMIC RAY EXPOSURE DATING

AMS measurements of the 36Cl and 10Be concentrations were performed at the 5 MV nation AMS facility ASTER located at CEREGE, Aix en Provence, France, All 36Cl concentrations were normalized to the KNSTD1600 calibration material (36Cl/35Cl =1.6\*10-12 provided by Nishiizumi) and all <sup>10</sup>Be concentrations were normalized to the SRM 4325 NIST reference material with an assigned 10Be/9Be value of (2.79±0.03)\*10-11. The 36Cl decay constant of (2.303  $\pm 0.016$ )\* $10^{-6}$  a<sup>-1</sup> corresponds to a half-life ( $T_{1/2}$ ) of (3.014 $\pm 0.021$ )\* $10^{5}$  years and the <sup>10</sup>Be deca constant of (4.987±0.036)\*10<sup>-7</sup> a<sup>-1</sup> to a half-life of (1.39±0.01)\*10<sup>6</sup> years. Analytical uncertainties include the counting statistics, machine stability (~0.5%) and blank correction the associated 36Cl/35Cl blank ratio being ~10-14 and the 10Be/9Be blank ratio ~2\*10-15. Cosmi ray exposure ages derived from these cosmogenic nuclide concentrations were calculated usin production rates corrected for topographic shielding following Dunne et al. (1999) and scaled to the sites position (latitude, altitude) using the Stone scheme (Stone, 2000) from a sea leve and high latitude production rate of 4.5±0.3 atoms of 10Be g(S(O2))-1.yr-1 for 10Be and integrating all production pathways (spallation, including a spallation production rate of 42.0  $\pm$  2.0 atom of <sup>36</sup>Cl g<sup>-1</sup><sub>Ca</sub><sup>-1</sup> yr<sup>-1</sup> and at sea level and high latitude, thermal neutron, muon captures and radiogenic production) and the rock's chemical composition for 36Cl.

# 3. RINGMO-PHOKSUNDO SITE

rockslides bodies derived from different parts of the rocky face bounding the lake in its SE part.



The Phoksumdo lake (3600 m asl.: area of 4.5-to-5 km²) is the

second largest lake of Nepal: it owes its origin to the damming of the Suli Gad River by the large (4.5 km3) collapse of a mountain wall (Dhaulagiri limestones) culminating at 5148 m, SE of the lake (Fort & Rimal 2010). According to Yaqi (1997). the collapse may have occurred 30 to 40 ka, an interpretation revised by this study.



The landslide morphology presents a series of mounds (metres to tens of metres high) and depressions of varying size. Its mass is composed of large blocks of massive dolomites found as far as west of Ringmo Village and down to the southern rim of the landslide dam

Most of the Suli scarp is uphold by limestones outcrops, but along the central part, the former Phoksumdo gorge is filled in by dark grey, fine sediments (till material) and overlain by orange conglomerates (including dolomites boulders), then by yellowish, unsorted conglomeratic material blocks and calcareous silty matrix). The presence of glacial till at the very base of the sequence suggests the rockslide event is post-glacial.



Two consistent 36Cl ages of 20,885 ±1675 argue for a single, massive event of paraglacial origin Field assumptions are confirmed by cosmic ray exposure dating:

the landslide dam was formed in response to the collapse of the limestone cliff in one single, massive event. The 36Cl dates confirm the paraglacial/post-glacial origin of the failure. an event that fits well with the last chronologies available on the Last Glacial Maximum in the Nepal Himalaya (Owen & Dortch 2014). IRSL dating of silts suggests a progressive trend to dryness of these Northern Himalayan regions from Late Pleistocene to



The IRSL samples, PHOK 2 and PHOK 1, were collected respectively at the top and the bottom of the silty deposit, below the surface soil. IRSL measu were performed at Montreal laboratory by S. Balescu, S. Huot and M. Lamothe. The corrected IRSL ages of PHOK 2 and PHOK 1, respectively 4.7 ± 0.3 ka and 12.6 ± 0.7 ka, are stratigraphically consistent (Fort et al., 2013).



Both the rock avalanche and glaciers controlled the existence and level of the "Marpha Lake" (lacustrine deposits up to

The Dhampu-Chhova rock avalanche (109 m3, area extent 10 km2) was derived from the northward failure of the Kaiku ridge uphold by north-dipping, upper crystal lines of the Higher Himalava (Fort 1974: 2000). It dammed the Kali Gandaki River, with complex interactions with the Late Pleistocene ice tongues derived from the Dhaulagiri (8167 m) and Nilgiris (7061 m

> Be-10 cosmic ray exposure ages (CEREGE, Gribenski 2010) i09-01 : 29 674 ± 1 022 ka Ti09-02 · 24 844 + 628 ka Ti09-03: 29 694 + 1 011 ka

The discrepancy in the dating of both lacustrine and failure events raise the issue of the nature of the lacustrine dam: (i) rockght (Fort 1974: 1980: Baade 2000) or (ii) possibly a glacial tongue controlling the lake level (Fort and Cossart 2013); (iii) repeated failures cannot be ruled out (Zech et al. 2010).



geomorphological investigations was given by the Dpt of National Parks, Ministry of Forests, and free entrance in the Shey-Phoksundo National Park was delivered by Home Ministry, Government of Nepal. Our very deep appreciation goes to Dr. Som Sapkota (Dpt of Mines and Geology) for his help.

The fact that both Zech et al. (2009) and our team (Fort et al., in prep.) have obtained consistent exposure duration to cosmic ray, from different samples and different laboratories, provides very good evidence of a chronologically, well constrained Dhumpu-Chhoya rock-avalanche event, i.e. about 30,000 years ago. Conversely, dating of lacustrine sediments has not been crosschecked to date, and some uncertainty cannot entirely be ruled out.

Note on Infrared Stimulated Luminescence (IRSL) dating

## 5. COMPARISONS and CONCLUDING REMARKS

### Phoksundo-Ringmo

- Efficient drainage blockage, but no backwash sedimentation (predominant carbonates + low discharge)
- Prominant knick point: karst + boulder pavement reducing incision => prominent hanging valley
- Triggering factor: glacial retreat («paraglacial collapse»), cf. <sup>36</sup>Cl dates

- Efficient drainage blockage, and extensive lacustrine sedimentation (glacial
- sediment reworking + Spiti shales + glacial- and snow-melt discharges) - Knick points and associated epigenetic gorges + braided Kali Gandaki
- valley: delayed incision => relatively hanging valley - No climate forcing, cf. 10Be dates => Triggering factor related to the North
- Himalavan Detachment Fault + foliation dip of Upper Himalayan Gneisses - Complex interplay between glaciers and mountain wall collapses (Fort 2000)
- Both examples confirm that giant landslides play a significant role in the:

# Destruction of Himalayan topography (Fort 1988),

- Bedrock protection from river incision (Korup et al. 2010).
- Delay of sediment transport outward from the mountain zone (cf.
- sediment storages & budgets; Fort & Cossart 2013; Blöthe J. & Korup O., 2013)
- Local controls exerted by topography and climate (eg. magnitude of glaciation + snow <--> discharge) + lithology (eg. sediment supply <--> storage)

