



HAL
open science

Auxiliary Variable-based Balancing (AVB) for source term treatment in open channel simulations

Carole Delenne, Vincent Guinot

► **To cite this version:**

Carole Delenne, Vincent Guinot. Auxiliary Variable-based Balancing (AVB) for source term treatment in open channel simulations. *Advances in Water Resources*, 2012, 44, pp.85-100. 10.1016/j.advwatres.2012.05.007 . hal-01196833

HAL Id: hal-01196833

<https://hal.science/hal-01196833>

Submitted on 10 Sep 2015

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

bed slope source term (DBF) [37], the quasi-steady wave propagation method [29], asymptotic balancing [12] or the source term projection technique in discontinuous Galerkin techniques [27]. Examples of the latter are the well-balanced approach [1, 2, 8, 24, 31, 35], flux and source term splitting [11], characteristics-based approximate-state and augmented Riemann solvers [10, 16, 18, 20, 30], the homogeneous approach [28] and other static equilibrium-preserving techniques [6, 19, 40].

Various solutions have also been proposed to enforce the C -property in finite volume-based discretizations. One of the earliest solutions, proposed in [33] for the solution of the SWE and later extended in [40], consists in replacing the water depth with the free surface elevation. This option can be extended to the open channel equations in arbitrary-shaped channels, as shown in the present paper. It has the drawback that simple flow configurations such as uniform flow over a constant slope cannot be computed accurately (see section 3.2). Another option is to approximate the variations in the cross-sectional area with a consistent estimate taken from the balance between the specific force and the source term in the momentum equation [6]. The estimate is defined in such a way that it is zero under steady state conditions. Very similar formulae to that of [6] have been obtained using completely different approaches in [28, 30]. The approaches [6, 28, 30] have the common point that the gradient in one of the flow variables is replaced with the gradient in another variable, called auxiliary variable hereafter. This gradient is zero under static conditions. That different approaches yield the same formulae lead to wonder whether a general methodology can be derived to define auxiliary variables.

In the present paper, the principle of Auxiliary Variable-based Balancing (AVB) is presented for one-dimensional free surface flow calculations in non-prismatic, trapezoidal channels. This is motivated by the fact that in industrial open channel packages, the cross-sectional geometry is broken into a set of trapezoidal elements. The AVB approach is used to derive flux formulae that allow non-static, steady state flow conditions to be preserved, even at low orders of discretization, that is, when first-order schemes are used.

The principle of the AVB method is presented for the water hammer and one-dimensional SWE in [25]. However, the one-dimensional shallow water equations are a very simplified description of free surface flows in natural channels. Besides, only one possible approach for source term discretization (a variant of source term upwinding) is considered in [25]. The applicability of the approach to more complex cross-sections and other source term discretization approaches is not investigated in [25]. The objectives of the present paper are (i) to present the methodology of Auxiliary Variable-based Balancing (AVB), (ii) to apply the AVB approach to the open channel flow equations in a well-balanced, finite volume framework, (iii) to provide the flux and source term discretizations for a variety of Riemann solvers, and (iv) to analyse the accuracy of the numerical solutions obtained using a number of various AVB-based discretizations. As mentioned above, first-order space discretizations are retained for the sake of computational rapidity.

The structure of the paper is as follows. Section 2 presents the governing equations and their discretisation. The AVB methodology is detailed in section 3 and its application to classical Riemann solvers presented in section 4. Section 5 provides computational examples, including steady-state configurations and transient test-cases as well as a convergence analysis for the classical dam-break problem (for which an analytical solution is available).

2. Governing equations and solution method

2.1. Governing equations

The purpose is to solve 2×2 hyperbolic systems of conservation laws in the form

$$\frac{\partial \mathbf{U}}{\partial t} + \frac{\partial \mathbf{F}}{\partial x} = \mathbf{S} \quad (1)$$

where \mathbf{U} , \mathbf{F} and \mathbf{S} are defined as

$$\mathbf{U} = \begin{bmatrix} A \\ Q \end{bmatrix}, \mathbf{F} = \begin{bmatrix} Q \\ M \end{bmatrix} = \begin{bmatrix} Q \\ \frac{Q^2}{A} + \frac{P}{\rho} \end{bmatrix}, \mathbf{S} = \begin{bmatrix} 0 \\ (S_0 - S_f)gA + \frac{R}{\rho} \end{bmatrix} \quad (2)$$

where A is the cross-sectional area, g is the gravitational acceleration, M is the specific force, P is the pressure force exerted on the wetted cross-sectional area, S_0 and S_f are respectively

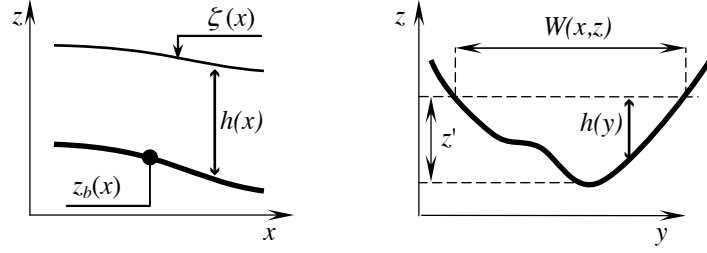


Figure 1: **Channel geometry.** Left) longitudinal view: bottom and water elevation. Right) transversal view: channel width and depth.

98 the bottom and energy slope, R is the x -component of the reaction of the walls onto the
 99 water (if the channel is non-prismatic) and ρ is the water density.

100 The forces P and R are derived from the assumption of a hydrostatic pressure distribution
 101 and obey the following definitions [14]:

$$\frac{P}{\rho} = \int_A (\zeta - z) g dA = \int_0^h (h - z') g W(z') dz' \quad (3)$$

$$\frac{R}{\rho} = \int_0^h (h - z') g \left(\frac{\partial W}{\partial x} \right)_{h-z'=\text{Const}} (z') dz' \quad (4)$$

102 where $W(z)$ is the width of the channel at the elevation z , h is the water depth (that is the
 103 distance between the lowest point in the cross-section and the free surface), $z' = z - z_b$ is
 104 the elevation above the bottom lowest point and ζ is the free surface elevation (Figure 1).

105 The energy slope is classically assumed to obey a turbulent-type friction law such as
 106 Manning's law:

$$S_f = n_M^2 u^2 R_H^{-4/3} \quad (5)$$

107 where n_M is Manning's friction coefficient, $u = Q/A$ is the flow velocity and R_H is the
 108 hydraulic radius, defined as the ratio of the cross-sectional area A to the wetted perimeter
 109 χ , yielding

$$S_f = n_M^2 Q^2 A^{-10/3} \chi^{4/3} \quad (6)$$

110 It is noted that the Jacobian matrix \mathbf{A} of \mathbf{F} with respect to \mathbf{U} is given by

$$\mathbf{A} = \frac{\partial \mathbf{F}}{\partial \mathbf{U}} = \begin{bmatrix} 0 & 1 \\ c^2 - u^2 & 2u \end{bmatrix} \quad (7)$$

111 where the speed c of the waves in still water is defined as

$$c^2 \equiv \frac{\partial(\frac{P}{\rho})}{\partial A} = \frac{gA}{b} \quad (8)$$

112 where $b = W(\zeta)$ is the top width of the channel. The matrix \mathbf{A} can be diagonalized into a
 113 matrix $\mathbf{\Lambda}$ defined as:

$$\mathbf{\Lambda} = \begin{bmatrix} \lambda^{(1)} & 0 \\ 0 & \lambda^{(2)} \end{bmatrix} \quad (9a)$$

$$\lambda^{(1)} = u - c \quad (9b)$$

$$\lambda^{(2)} = u + c \quad (9c)$$

114 The problem is assumed to be properly posed hereafter, that is, the initial and boundary
 115 conditions are specified such that Eq. (1) can be solved uniquely for \mathbf{U} at all points of a
 116 computational domain $[0, L]$ for all times $t > 0$.

117 *2.2. Finite volume discretization*

118 Eq. (1) is discretized using a finite volume formalism as

$$\mathbf{U}_i^{n+1} = \mathbf{U}_i^n + \frac{\Delta t}{\Delta x_i} \left(\mathbf{F}_{i-\frac{1}{2}}^{n+\frac{1}{2}} - \mathbf{F}_{i+\frac{1}{2}}^{n+\frac{1}{2}} \right) + \Delta t \mathbf{S}_i^{n+\frac{1}{2}} \quad (10)$$

119 where the subscript i denotes a cell average, subscripts $i \pm \frac{1}{2}$ denote estimates at the interfaces
 120 between the computational cells, the superscript n indicates that the variable is estimated
 121 a time level n , and the superscript $n + \frac{1}{2}$ denotes an average value between time levels n
 122 and $n + 1$, and where Δx_i is the width of the computational cell i . In explicit schemes, the
 123 variables with superscripts $n + \frac{1}{2}$ are computed using the known values at the time level n ;
 124 in implicit schemes, the unknown values at the time level $n + 1$ are used.

125 In what follows, non-prismatic, trapezoidal cross-sections are considered. The reason for
 126 this is that in all commercial open channel packages, the channel geometry is discretized into
 127 a series of trapezia. Consequently, the capability to deal with trapezoidal cross-sections is
 128 seen as an indispensable prerequisite to a generalisation of the method to arbitrary-shaped
 129 channels. Note that rectangular and triangular cross-sections are obtained as particular
 130 cases of the proposed approach, as illustrated by a number of computational examples in
 131 Section 5.

132 The geometric parameters of the cross-sections are defined at the interfaces between
 133 the computational cells. They are interpolated linearly within the cells. Consequently, the
 134 geometry is continuous at the cell interfaces. Assuming non-prismatic trapezoidal channel
 135 geometry, the width $W(x, z)$ at a given abscissa x and elevation z takes the form

$$W(x, z) = W_0(x) + (z - z_b(x)) W_1(x) \quad (11)$$

136 where $W_0(x)$ is the bottom width of the channel at the abscissa x , $W_1(x)$ is the derivative
 137 of W with respect to z and z_b is the bed elevation at the abscissa x . As mentioned above,
 138 W_0 , W_1 and z_b are assumed to vary linearly with x within the cells. In the cell i , one has:

$$W_0(x) = W_{0,i-\frac{1}{2}} + (x - x_{i-\frac{1}{2}}) W_{0,i}^{(x)} \quad (12)$$

$$W_1(x) = W_{1,i-\frac{1}{2}} + (x - x_{i-\frac{1}{2}}) W_{1,i}^{(x)} \quad (13)$$

$$z_b(x) = z_{b,i-\frac{1}{2}} + (x - x_{i-\frac{1}{2}}) z_{b,i}^{(x)} \quad (14)$$

139 where the superscript (x) denotes the derivative with respect to x :

$$W_{0,i}^{(x)} = \frac{W_{0,i+\frac{1}{2}} - W_{0,i-\frac{1}{2}}}{\Delta x_i} \quad (15)$$

$$W_{1,i}^{(x)} = \frac{W_{1,i+\frac{1}{2}} - W_{1,i-\frac{1}{2}}}{\Delta x_i} \quad (16)$$

$$z_{b,i}^{(x)} = \frac{z_{b,i+\frac{1}{2}} - z_{b,i-\frac{1}{2}}}{\Delta x_i} = S_{0,i} \quad (17)$$

140 The cross-sectional area A is given as the integral of W between the bottom level and
 141 the free surface elevation:

$$A(x) = \int_{z_b}^{\zeta} W(x, z) dz = \left[W_0(x) + W_1(x) \frac{h(x)}{2} \right] h(x) \quad (18)$$

142 where $h(x)$ is the water depth at the abscissa x . Assuming that the free surface is horizontal
 143 in the cell i (which is true in the case of the first-order Godunov scheme), the average cell
 144 value A is given by

$$A_i(\zeta) \equiv \frac{1}{\Delta x_i} \int_{x_{i-\frac{1}{2}}}^{x_{i+\frac{1}{2}}} A(x) dx = \alpha_i h_{i-\frac{1}{2}}^2 + \beta_i h_{i-\frac{1}{2}} + \gamma_i \quad (19a)$$

$$h_{i-\frac{1}{2}} = (\zeta - z_b)_{i-\frac{1}{2}} \quad (19b)$$

$$\alpha_i = \frac{1}{2} W_{1,i-\frac{1}{2}} + \frac{\Delta x_i}{4} W_{1,i}^{(x)} \quad (19c)$$

145

146

$$\beta_i = W_{0,i-\frac{1}{2}} + \frac{\Delta x_i}{2} \left(W_{0,i}^{(x)} + S_{0,i} W_{1,i-\frac{1}{2}} \right) + \frac{\Delta x_i^2}{3} S_{0,i} W_{1,i}^{(x)} \quad (19d)$$

147

$$\gamma_i = \frac{\Delta x_i}{2} S_{0,i} W_{0,i-\frac{1}{2}} + \frac{\Delta x_i^2}{6} \left(2S_{0,i} W_{0,i}^{(x)} + S_{0,i}^2 W_{1,i-\frac{1}{2}} \right) + \frac{\Delta x_i^3}{8} S_{0,i}^2 W_{1,i}^{(x)} \quad (19e)$$

148 *2.3. Flux calculation*

149 Approximate Riemann solvers provide flux formulae that can be recast in the following
150 form, derived from [17]:

$$\mathbf{F} = a\mathbf{F}_L + (1-a)\mathbf{F}_R + \mathbf{D}(\mathbf{U}_L - \mathbf{U}_R) \quad (20)$$

151 where L and R denote respectively the left and right states of the Riemann problem, a is a
152 coefficient between 0 and 1 and \mathbf{D} is a diffusion matrix that contributes to stabilise the nu-
153 merical solution. The left and right states are obtained from an appropriate reconstruction,
154 the simplest possible option (the first-order Godunov scheme [23]) being to use the average
155 cell values. For $a = 1/2$, Eq. (20) is the sum of a centred flux and a so-called artificial vis-
156 cosity term. Both a and \mathbf{D} are functions of the wave speeds, in other words, the eigenvalues
157 of the Jacobian matrix \mathbf{A} of \mathbf{F} with respect to \mathbf{U} (Eq. 9).

158 How the left and right states for the Riemann problem are to be computed from the
159 average cell values is dealt with in Section 3.

160 *2.4. Source term discretization*

161 The momentum source term is discretized explicitly.

162 The friction source term is computed by applying explicit estimates to the terms in
163 Eq. (6):

$$(S_f)_i^n = n_M^2 \left(u^2 R_H^{4/3} \right)_i^n = \left(\frac{W_0 + \frac{1}{2}W_1 h}{W_0 + h(W_1^2 + 4)^{1/2}} \right)_i^n \quad (21)$$

$$(S_f)_i^n = n_M^2 \left(Q^2 A^{2/3} \chi^{-4/3} \right)_i^n = n_m^2 \left(Q^2 \left(W_0 + \frac{1}{2}W_1 h \right)^{2/3} \left(W_0 + h\sqrt{W_1^2 + 4} \right)^{-4/3} \right)_i^n \quad (22)$$

164 The geometric source term $gAS_0 + \frac{R}{\rho}$ is rather difficult to compute directly under the
165 assumption of varying W_0 , W_1 and z_b . However, it can be estimated cell-wise from simple
166 balance considerations. Consider static equilibrium conditions, i.e., $Q=0$ in all cells, then
167 Eqs. (1) and (10) yield

$$\left(gAS_0 + \frac{R}{\rho} \right)_i^n \Delta x_i = \left(\frac{P}{\rho} \right)_{i+\frac{1}{2}}^n - \left(\frac{P}{\rho} \right)_{i-\frac{1}{2}}^n \quad (23)$$

168 Given the definition (Eq 11) of the channel width, and the specific pressure force $\frac{P}{\rho}$
169 (Eq. 3), one has

$$\frac{P}{\rho} = g \int_0^h (h - z') (W_0(x) + z'W_1(x)) dz' = \frac{1}{2}gW_0(x)h^2 + \frac{1}{6}gW_1(x)h^3 \quad (24)$$

170 The value of the specific pressure force at the interface $i - \frac{1}{2}$ is then easily computed as

$$\left(\frac{P}{\rho} \right)_{i-\frac{1}{2}}^n = \frac{g}{2}W_{0,i-\frac{1}{2}}h_{i-\frac{1}{2}}^2 + \frac{g}{6}W_{1,i-\frac{1}{2}}h_{i-\frac{1}{2}}^3 \quad (25)$$

171 with the definition (19b) for $h_{i-\frac{1}{2}}$. Since the purpose is to estimate the source term in the
172 cell i , the free surface elevation to be used in Eq. (19b) is ζ_i^n . The same formulation can
173 be obtained at interface $i + \frac{1}{2}$ yielding the final estimate for the source term to be used in
174 Eq. (10):

$$\begin{aligned} \left(gAS_0 + \frac{R}{\rho} \right)_i^n \Delta x_i &= \frac{g}{2} \left(W_{0,i+\frac{1}{2}} + \frac{1}{3}W_{1,i+\frac{1}{2}}h_{i+\frac{1}{2},i} \right) h_{i+\frac{1}{2},i}^2 \\ &\quad - \frac{g}{2} \left(W_{0,i-\frac{1}{2}} + \frac{1}{3}W_{1,i-\frac{1}{2}}h_{i-\frac{1}{2},i} \right) h_{i-\frac{1}{2},i}^2 \end{aligned} \quad (26a)$$

175

$$h_{i-\frac{1}{2},i} = \zeta_i^n - z_{b,i-\frac{1}{2}} \quad (26b)$$

176

$$h_{i+\frac{1}{2},i} = \zeta_i^n - z_{b,i+\frac{1}{2}} \quad (26c)$$

177 *2.5. Balancing issues*

178 The discretization of the source term in the momentum equation usually poses no prob-
 179 lem. This issue has been dealt with abundantly in the literature, within a very wide variety
 180 of techniques [1, 2, 4, 6, 27, 28, 38, 40]. A remaining problem encountered in practical
 181 applications is related to the continuity equation and the difference often observed between
 182 the average cell values and the interface values for the volume discharge.

183 Consider a solution \mathbf{U} verifying steady state, $\frac{\partial \mathbf{U}}{\partial t} = 0$. The first component of Eq. (1)
 184 imposes that Q be equal to a constant Q_0 all throughout the computational domain. In
 185 particular, the discharge Q computed at the cell interfaces should be identical to that in the
 186 cell values. This, however, is not necessarily the case if Eq. (20) is used. Indeed, writing the
 187 first component of Eq. (20) leads to the following formula for the interface flux

$$Q_0 = (a + D_{12}) Q_L + (1 - a - D_{12}) Q_R + D_{11}(A_L - A_R) \quad (27)$$

188 where D_{11} and D_{12} are the components on the first row of the artificial viscosity matrix \mathbf{D} .
 189 Eq. (27) can be rewritten as

$$(a + D_{12}) Q_L + (1 - a - D_{12}) Q_R = Q_0 + D_{11}(A_R - A_L) \quad (28)$$

190 Assume that the discretized solution has reached steady state. If the geometry of the
 191 channel is arbitrary (non-constant bottom slope and/or non-prismatic channel), in general
 192 $A_L \neq A_R$. It is then obvious from Eq. (28) that at least one of the discharges Q_L and Q_R
 193 is different from the uniform discharge Q_0 . Consequently, a non-uniform discharge profile
 194 is obtained. In particular, if the initial situation is static ($Q_0 = 0$), non-zero discharges
 195 are computed. Artificial oscillations appear and propagate throughout the computational
 196 domain.

197 The ability of a numerical scheme to preserve static equilibrium conditions has been
 198 introduced as the C -property in [4]. Specifying the C -property exactly or approximately
 199 has proved to lead to efficient source term balancing techniques. The most widespread
 200 approach consists in adapting the discretization of the source term to the formulation of
 201 the flux so as to satisfy the C -property. In the Auxiliary Variable-based Balancing (AVB)
 202 approach, the opposite approach is followed: the formulation of the flux gradients is adapted
 203 to that of the source term.

204 **3. Auxiliary Variable-based Balancing method**205 *3.1. Principle*

206 The AVB method is based on the following requirements: (i) the artificial viscosity term
 207 in Eq. (20) should be modified in such a way that diffusion becomes zero when steady state is
 208 reached; (ii) the source term in the momentum equation should be discretized in such a way
 209 that it does not influence the calculation of the flux in the intermediate region of constant
 210 state. The second issue has been addressed in subsection 2.4 (source term discretization);
 211 the first issue is dealt with in the following subsections.

212 AVB uses an auxiliary variable \mathbf{V} in the expression of the artificial viscosity term:

$$\mathbf{F} = a\mathbf{F}_L + (1 - a)\mathbf{F}_R + \mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) \quad (29)$$

213 where \mathbf{V} is a function of both the variable \mathbf{U} and the parameter φ , $\mathbf{V} = \mathbf{V}(\mathbf{U}, \varphi)$. The dif-
 214 fusion matrix $\mathbf{D}_{\mathbf{V}}$ and the auxiliary variable \mathbf{V} are chosen such that the following conditions
 215 are verified:

216 **(C1):** under steady state conditions, $\mathbf{V}_L = \mathbf{V}_R$.

217 **(C2):** for $\varphi = \text{Const}$, $\mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) = \mathbf{D}(\mathbf{U}_L - \mathbf{U}_R)$

218 Condition (C1) is the so-called enhanced consistency condition for steady state flow, which is
 219 the desired property for scheme well-balancing. Condition (C2) means that the strengths of
 220 the artificial viscosity terms in Eqs. (20) and (29) are identical, thus preserving the stability
 221 properties of the numerical solution. The pending question is the determination of \mathbf{D}_V .

222 It is observed that the artificial viscosity terms in Eqs. (20) and (29) are approximations
 223 of the following derivatives:

$$\mathbf{D}_V(\mathbf{V}_L - \mathbf{V}_R) = -\Delta x \mathbf{D}_V \frac{\partial \mathbf{V}}{\partial x} + \text{HOT}(\Delta x) \quad (30)$$

$$\mathbf{D}(\mathbf{U}_L - \mathbf{U}_R) = -\Delta x \mathbf{D} \frac{\partial \mathbf{U}}{\partial x} + \text{HOT}(\Delta x) \quad (31)$$

224 with

$$\text{HOT}(\Delta x) \xrightarrow{\Delta x \rightarrow 0} 0 \quad (32)$$

225 Noticing that $\mathbf{V} = \mathbf{V}(\mathbf{U}, \varphi)$, the derivative of \mathbf{V} with respect to x is expressed as

$$\frac{\partial \mathbf{V}}{\partial x} = \frac{\partial \mathbf{V}}{\partial \mathbf{U}} \frac{\partial \mathbf{U}}{\partial x} + \frac{\partial \mathbf{V}}{\partial \varphi} \frac{\partial \varphi}{\partial x} \quad (33)$$

226 Substituting Eq. (33) into Eq. (30), comparing with Eq. (31) and imposing condition (C2)
 227 gives:

$$\mathbf{D}_V \frac{\partial \mathbf{V}}{\partial \mathbf{U}} = \mathbf{D} \quad (34)$$

228 Consequently, \mathbf{D}_V is given by

$$\mathbf{D}_V = \mathbf{D} \left(\frac{\partial \mathbf{V}}{\partial \mathbf{U}} \right)^{-1} \quad (35)$$

229 3.2. Balancing option 1: free surface elevation

230 One of the earliest examples of the use of an auxiliary variable is found in [33] for the
 231 solution of the shallow water equations where the free surface elevation $\zeta = z_b + h$ is used
 232 in place of the water depth h . The rationale is that under static conditions, the free surface
 233 elevation is constant, consequently, both $d\zeta$ and dQ are zero at equilibrium. Note that the
 234 approach has been extended to the reconstruction technique in higher-order schemes in [40].
 235 This leads to the following possible definition for the auxiliary variable

$$d\mathbf{V}_1 = \begin{bmatrix} d\zeta \\ dQ \end{bmatrix} \quad (36)$$

236 Since $dA = bd\zeta$, one has from the definition of \mathbf{U} in Eq. (2):

$$\frac{\partial \mathbf{V}_1}{\partial \mathbf{U}} = \begin{bmatrix} b^{-1} & 0 \\ 0 & 1 \end{bmatrix}, \left(\frac{\partial \mathbf{V}_1}{\partial \mathbf{U}} \right)^{-1} = \begin{bmatrix} b & 0 \\ 0 & 1 \end{bmatrix} \quad (37)$$

237 This leads to the following artificial viscosity term:

$$\mathbf{D}_V(\mathbf{V}_{1L} - \mathbf{V}_{1R}) = \mathbf{D} \begin{bmatrix} b & 0 \\ 0 & 1 \end{bmatrix} \begin{bmatrix} \zeta_L - \zeta_R \\ Q_L - Q_R \end{bmatrix} = \mathbf{D} \begin{bmatrix} (\zeta_L - \zeta_R)b \\ Q_L - Q_R \end{bmatrix} \quad (38)$$

238 Note that in the case of the SWE, $b = 1$ and Nujic's [33] approach is retrieved.

239 This option has the drawback that steady state, uniform flow cannot be maintained
 240 exactly. Indeed, under uniform flow conditions, $Q_L = Q_R$ but the free surface elevations
 241 in two adjacent cells are not identical, $\zeta_L \neq \zeta_R$. Therefore, the artificial diffusion term in
 242 the continuity equation is non-zero and the interface flux is not equal to $Q_L = Q_R$. This is
 243 substantiated by the computational examples in Section 5.

244 *3.3. Balancing option 2: specific force*

245 This is the option explored in [6]. Similar formulae were obtained for the one-dimensional
 246 shallow water equations in [28, 30, 18], albeit from different considerations. The latter
 247 three approaches, however, focus on rectangular channels, while the proposed approach is
 248 applicable to arbitrary-shaped channels.

249 The specific force is used in place of the cross-sectional area in the first component of
 250 the auxiliary variable \mathbf{V} . The motivation is that under dynamic equilibrium (that is, under
 251 steady state flow conditions), the variations in the specific force are balanced exactly by
 252 the source terms in the momentum equation noted S_M . This leads to define the auxiliary
 253 variable \mathbf{V} in differential form as

$$d\mathbf{V} = \begin{bmatrix} dM - S_M \\ dQ \end{bmatrix} \quad (39)$$

254 Since $dM = (c^2 - u^2)dA + 2udQ$, one has

$$\frac{\partial \mathbf{V}}{\partial \mathbf{U}} = \begin{bmatrix} c^2 - u^2 & 2u \\ 0 & 1 \end{bmatrix}, \left(\frac{\partial \mathbf{V}}{\partial \mathbf{U}} \right)^{-1} = \begin{bmatrix} \frac{1}{c^2 - u^2} & -\frac{2u}{c^2 - u^2} \\ 0 & 1 \end{bmatrix} \quad (40)$$

255 This leads to the following artificial viscosity term:

$$\mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) = \mathbf{D} \left[\frac{M_L - M_R - 2u(Q_L - Q_R) - S_M \Delta x}{c^2 - u^2} \right]_{Q_L - Q_R} \quad (41)$$

256 where the source term S_M is computed in average between the centres of the left- and
 257 right-hand cells. It is simply estimated as the average of the cell values given by Eq. (26a).

258 Practical implementations [18, 30] indicate that in the neighbourhood of critical points,
 259 Eq. (41) induces a downwinding of the discharge and a discontinuous switch between sub-
 260 critical and supercritical flux formulae. Due to this, a different formula is proposed

$$d\mathbf{V} = \begin{bmatrix} dM - 2udQ - S_M \\ dQ \end{bmatrix} \quad (42)$$

261 This leads to

$$\frac{\partial \mathbf{V}}{\partial \mathbf{U}} = \begin{bmatrix} c^2 - u^2 & 0 \\ 0 & 1 \end{bmatrix}, \left(\frac{\partial \mathbf{V}}{\partial \mathbf{U}} \right)^{-1} = \begin{bmatrix} \frac{1}{c^2 - u^2} & 0 \\ 0 & 1 \end{bmatrix} \quad (43)$$

262 and the following artificial viscosity term is obtained

$$\mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) = \mathbf{D} \left[\frac{M_L - M_R - S_M \Delta x}{c^2 - u^2} \right]_{Q_L - Q_R} \quad (44)$$

263 This expression, however, remains invalid at critical points, for which $c^2 = u^2$. As
 264 proposed in [6], in the case of 1D SWEs on rectangular channel, the final estimate for $d\mathbf{V}$
 265 is the minmod of the estimates given by the specific force option and the original approach:

$$d\mathbf{V}_{\mathbf{2}} = \text{minmod}(d\mathbf{V}, d\mathbf{U}) \quad (45)$$

266 where $d\mathbf{V}$ is defined by Eq. (42) and the minmod operator by:

$$\text{minmod}(a, b) = \begin{cases} \min(|a|, |b|) & \text{if } ab \geq 0 \\ 0 & \text{if } ab < 0 \end{cases} \quad (46)$$

267 *3.4. Balancing option 3: hydraulic head*

268 In this option, presented in [25] (for 1D SWEs on rectangular channel), the hydraulic
 269 head $H = \zeta + \frac{u^2}{2g}$ is used as auxiliary variable:

$$d\mathbf{V} = \begin{bmatrix} dH - S_f \\ dQ \end{bmatrix} \quad (47)$$

270 Since $dH = \left(\frac{1}{b} - \frac{u^2}{gA} \right) dA + \frac{u}{gA} dQ = \frac{1}{b} [(1 - F^2) dA + \frac{F}{c} dQ]$, one has

$$\frac{\partial \mathbf{V}}{\partial \mathbf{U}} = \begin{bmatrix} \frac{1 - F^2}{b} & \frac{F}{bc} \\ 0 & 1 \end{bmatrix}, \left(\frac{\partial \mathbf{V}}{\partial \mathbf{U}} \right)^{-1} = \begin{bmatrix} \frac{b}{1 - F^2} & -\frac{F}{1 - F^2} \\ 0 & 1 \end{bmatrix} \quad (48)$$

271 This leads to the following artificial viscosity term:

$$\mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) = \mathbf{D} \left[\frac{(H_L - H_R - S_f \Delta x) b - \frac{F}{c} (Q_L - Q_R)}{1 - F^2} \frac{1}{Q_L - Q_R} \right] \quad (49)$$

272 For the same reasons as Option 2, the following variation is proposed for Option 3:

$$d\mathbf{V} = \begin{bmatrix} dH - S_f - \frac{F}{c} dQ \\ dQ \end{bmatrix} \quad (50)$$

273 The Jacobian matrix of \mathbf{V} with respect to \mathbf{U} is given by:

$$\frac{\partial \mathbf{V}}{\partial \mathbf{U}} = \begin{bmatrix} \frac{1-F^2}{b} & 0 \\ 0 & 1 \end{bmatrix}, \quad \left(\frac{\partial \mathbf{V}}{\partial \mathbf{U}} \right)^{-1} = \begin{bmatrix} \frac{b}{1-F^2} & 0 \\ 0 & 1 \end{bmatrix} \quad (51)$$

274 This leads to the following artificial viscosity term:

$$\mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) = \mathbf{D} \left[b \frac{H_L - H_R - S_f \Delta x}{1 - F^2} \frac{1}{Q_L - Q_R} \right] \quad (52)$$

275 As in option 2, this expression is not valid at critical points for which $F = 1$, the final
276 estimate for $d\mathbf{V}$ is thus

$$d\mathbf{V}_{\mathbf{3}} = \text{minmod}(d\mathbf{V}, d\mathbf{U}) \quad (53)$$

277 where $d\mathbf{V}$ is defined by Eq. (50).

278 Note that if a singular head loss ΔH_s is to be introduced, it can also be taken into
279 account in the artificial viscosity term:

$$\mathbf{D}_{\mathbf{V}}(\mathbf{V}_L - \mathbf{V}_R) = \mathbf{D} \left[b \frac{H_L - H_R - \Delta H_s - S_f \Delta x}{1 - F^2} \frac{1}{Q_L - Q_R} \right] \quad (54)$$

280 4. Application to classical approximate Riemann solvers

281 4.1. Application to the HLL solver

282 The HLL solver [26] can be written in the form (20) by defining a and \mathbf{D} as

$$a = \frac{\lambda^+}{\lambda^+ - \lambda^-} \quad (55a)$$

$$\mathbf{D} = -\frac{\lambda^- \lambda^+}{\lambda^+ - \lambda^-} \mathbf{I} \quad (55b)$$

283 where \mathbf{I} is the identity matrix and λ^- , λ^+ are respectively estimates of the fastest waves
284 $\lambda^{(1)}$ and $\lambda^{(2)}$ defined in Eqs. (9b, 9c) in the direction of negative and positive x [15, 17]:

$$\lambda^- = \min(u_L - c_L, u_R - c_R, 0) \quad (56a)$$

$$285 \lambda^+ = \max(u_L + c_L, u_R + c_R, 0) \quad (56b)$$

286 4.2. Application to Roe's solver

287 Roe's solver [34] can be written in the form (20) by setting

$$a = \frac{1}{2} \quad (57a)$$

$$\mathbf{D} = \frac{\tilde{\mathbf{A}}^\pm}{2} \quad (57b)$$

288 where $\tilde{\mathbf{A}}^\pm$ is the matrix generated by the absolute values of the eigenvalues of \mathbf{A} :

$$\tilde{\mathbf{A}}^\pm = \tilde{\mathbf{K}} \left| \tilde{\Lambda} \right| \tilde{\mathbf{K}}^{-1} \quad (58)$$

289 with

$$\tilde{\mathbf{K}} = \begin{bmatrix} 1 & 1 \\ \tilde{\lambda}^{(1)} & \tilde{\lambda}^{(2)} \end{bmatrix}, \quad \tilde{\mathbf{K}}^{-1} = \frac{1}{\lambda^{(2)} - \lambda^{(1)}} \begin{bmatrix} \tilde{\lambda}^{(2)} & -1 \\ -\tilde{\lambda}^{(1)} & 1 \end{bmatrix}, \quad \left| \tilde{\Lambda} \right| = \begin{bmatrix} |\tilde{\lambda}^{(1)}| & 0 \\ 0 & |\tilde{\lambda}^{(2)}| \end{bmatrix} \quad (59)$$

290 leading to the following expression for $\tilde{\mathbf{A}}^\pm$:

$$\tilde{\mathbf{A}}^\pm = \begin{bmatrix} a_{11} & a_{12} \\ a_{21} & a_{22} \end{bmatrix} \quad (60a)$$

$$a_{11} = \frac{\tilde{\lambda}^{(2)} |\tilde{\lambda}^{(1)}| - \tilde{\lambda}^{(1)} |\tilde{\lambda}^{(2)}|}{\tilde{\lambda}^{(2)} - \tilde{\lambda}^{(1)}} \quad (60b)$$

$$a_{12} = \frac{|\tilde{\lambda}^{(2)}| - |\tilde{\lambda}^{(1)}|}{\tilde{\lambda}^{(2)} - \tilde{\lambda}^{(1)}} \quad (60c)$$

$$a_{21} = -\tilde{\lambda}^{(1)} \tilde{\lambda}^{(2)} a_{12} \quad (60d)$$

$$a_{22} = \frac{\tilde{\lambda}^{(2)} |\tilde{\lambda}^{(2)}| - \tilde{\lambda}^{(1)} |\tilde{\lambda}^{(1)}|}{\tilde{\lambda}^{(2)} - \tilde{\lambda}^{(1)}} \quad (60e)$$

291 In Roe's approach [34], the eigenvalues $\tilde{\lambda}^{(1)} = (\tilde{u} - \tilde{c})$ and $\tilde{\lambda}^{(2)} = (\tilde{u} + \tilde{c})$ in the diagonal
292 matrix $\tilde{\mathbf{A}}$ are obtained from Roe's averages [21, 22] :

$$\tilde{c} = \left[\frac{g}{2} \left(\frac{A_L}{b_L} + \frac{A_R}{b_R} \right) \right]^{1/2} \quad (61a)$$

293

$$\tilde{u} = \frac{c_L u_L + c_R u_R}{c_L + c_R} \quad (61b)$$

294 4.3. Application to the Q-scheme

295 The Q-scheme uses the same formula as Roe's formula, except that the matrix $\tilde{\mathbf{A}}$ in
296 Eq. (58) is estimated from the average of the left and right-hand cells

$$\tilde{\mathbf{A}} = \tilde{\mathbf{A}} \left(\frac{\mathbf{U}_L + \mathbf{U}_R}{2} \right) \quad (62)$$

297 yielding to the following approximation for the eigenvalues:

$$\tilde{c} = g \left(\frac{A_{LR}}{\tilde{b}} \right)^{1/2} \quad (63a)$$

298

$$\tilde{u} = \frac{Q_{LR}}{A_{LR}} \quad (63b)$$

299 where $\tilde{b} = W_{0,i-\frac{1}{2}} + h_{LR} W_{1,i-\frac{1}{2}}$ and $X_{LR} = (X_L + X_R)/2$ ($X \in \{A, Q, h\}$).

300 4.4. Summary of formulae - Algorithmic aspects

301 From an algorithmic point of view, the steps in the solution process are the following:

- 302
- 303 1. For each cell, compute the free surface elevations ζ_L and ζ_R from the left and right
304 states \mathbf{U}_L and \mathbf{U}_R , using the correspondence between A and h , Eq (19a). Use the free
305 surface elevations to compute the geometric source term from Eqs (26).
 - 306 2. For each interface, compute the flux \mathbf{F} using Eq (29) with the fluxes on both sides of
307 the interface ($\mathbf{F}_L, \mathbf{F}_R$), the auxiliary variables $\mathbf{V}_L, \mathbf{V}_R$ according to the AVB option
308 chosen and with a depending on the solver.
 - 309 3. Compute the friction source term as in Eq (21).
 - 310 4. Apply the balance equation (10) to compute the hydrodynamic variable at the next
time step.

Symbol	Meaning	Value
g	Gravitational acceleration	9.81 m s^{-2}
L	Length of the domain	3,000 m
S_0	Bottom slope	10^{-3}
W_0	Channel width	1 m
W_1	Derivative of the width with respect to z	0
h_0	Initial water depth	1 m
z_{ds}	Prescribed surface elevation downstream	1 m
Q_{up}	Prescribed discharge upstream	$1 \text{ m}^3 \text{ s}^{-1}$
n_M	Manning's friction coefficient	$0.025 \text{ m}^{-1/3} \text{ s}$
Δx	Computational cell width	1 m
Δt	Simulation time	20,000 s

Table 1: Test 1 - steady state flow in a prismatic, rectangular channel. Parameters of the test case.

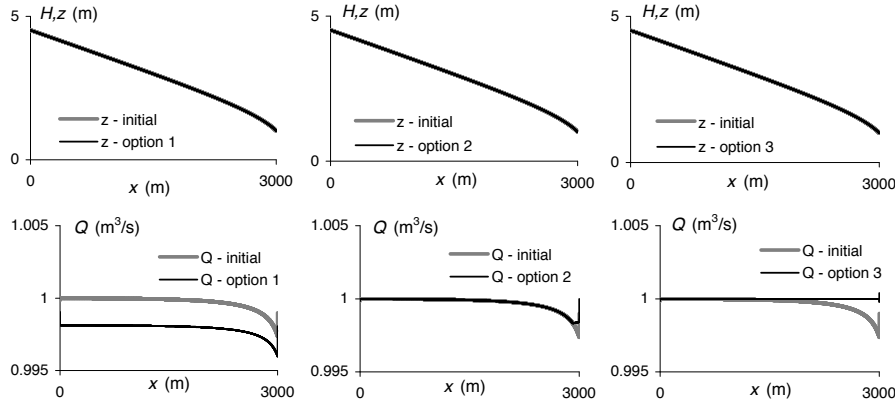


Figure 2: Test 1 - steady state flow in a prismatic, rectangular channel. Top: water elevation z and hydraulic head H , down: discharge Q obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options, using the HLL solver (the results obtained with the Roe's and the Q -scheme solvers are identical).

311 5. Computational examples

312 5.1. Steady state configurations

313 5.1.1. Test 1: steady state flow in a prismatic, rectangular channel

314 In this test case, the various AVB options are applied to steady state flow in a prismatic,
315 rectangular channel (i.e. with a constant value of W_0 and with $W_1 = 0$) including friction. A
316 transient simulation is carried out from an initial state at rest until steady state is obtained.
317 The parameters of the test case are given in Table 1.

318 Figure 2 shows the results obtained from the initial formulation *i.e.* with $\mathbf{V} = \mathbf{U}$ and
319 with the three different AVB options. Only the HLL solver is shown in this case because
320 the results obtained with the two other solvers are identical. The profiles of the free surface
321 elevation z and the hydraulic head H are identical regardless of the AVB option used (note
322 that the water elevation and the hydraulic head are nearly identical because of a small
323 velocity).

324 However, under steady-state conditions, the discharge Q is expected to be uniform over
325 the entire domain and equal to Q_{up} . The only option that provides the correct value of Q
326 over the whole domain but the downstream boundary, is the third one, *i.e.* based on the
327 hydraulic head.

328 Figure 3 shows results of the same test case but with the introduction of a singular
329 head loss in the middle of the channel. The head loss is computed using a classical Borda
330 relationship:

$$\Delta H_s = \alpha \frac{v^2}{2g}$$

331 where α is arbitrarily chosen to $\alpha = 5$ in this case, but can be estimated from any empirical
332 law. Figure 3a presents results using HLL solver. Each option provides a good estimate of

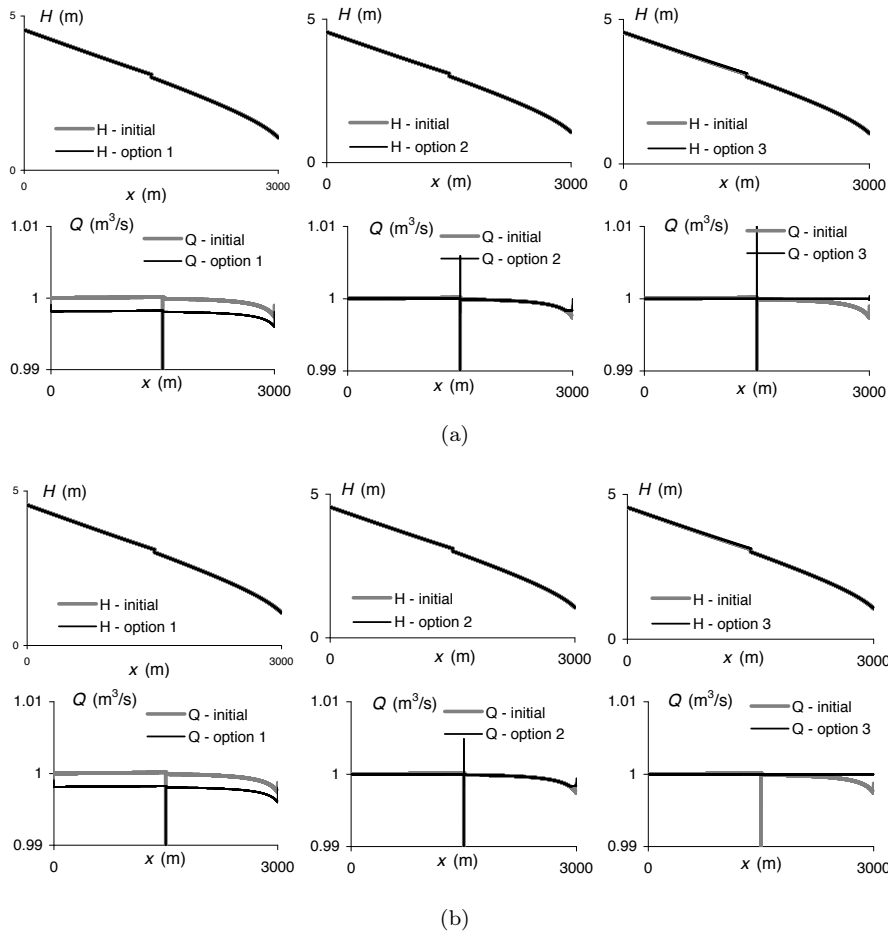


Figure 3: Test 1b - steady state flow in a prismatic, rectangular channel with an arbitrary singular head loss in the middle of the channel. Top: water elevation z and hydraulic head H , down: discharge Q obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options, using a) HLL solver, b) Roe's solver (identical to Q -scheme).

333 the hydraulic head and water elevation. In addition to the behaviour previously observed,
 334 the singular head loss triggers a spike in the discharge profiles when the HLL solver is used.
 335 Option 3, that explicitly takes into account the singular head loss in the flux computation,
 336 is the only one that provides a constant value $Q = Q_{\text{up}}$ with Roe's and Q -scheme solvers
 337 (Figure 3b).

338 5.1.2. Test 2 : frictionless steady state flow in a non-prismatic, rectangular channel

339 The channel profile is shown in Figure 4; it contains two consecutive narrowings of the
 340 cross-section: the first one due to the width narrowing (minimum width at 25% of the
 341 channel length) and the second one to a bump in the bottom elevation (maximum elevation
 342 at 75% of the channel length). This is a frictionless test case, the parameters of which are
 343 given in Table 2. As for the first test case, a transient simulation is run for a sufficiently
 344 long time, so that the transient regime vanishes and steady state is reached. The prescribed
 345 discharge and downstream water level are chosen such that the flow regime is subcritical
 346 upstream of both the narrowing and the bump, yielding two hydraulic jumps.

347 Figure 5 shows results obtained with the different AVB options and the three solvers
 348 (note that the results obtained with Roe's solver and Q -scheme are identical). The profiles
 349 obtained for the hydraulic head H and water elevation z with the different solvers and
 350 options bear similarities except for the points upstream the channel narrowing. In contrast,
 351 substantial differences can be observed for the discharge Q . As for the first test case, the
 352 steady state configuration theoretically implies a constant value for the discharge. It can
 353 be seen that the same profile is obtained using the initial formulation (i.e. $\mathbf{V} = \mathbf{U}$) for the
 354 three solvers, and that this profile is the most different from the constant value of $Q = Q_{\text{up}}$.
 355 Option 2 also give a strongly variable discharge in space when used with HLLC, but not
 356 with Roe's solver or Q -scheme. Option 3 gives better results: it is very close to $Q = Q_{\text{up}}$

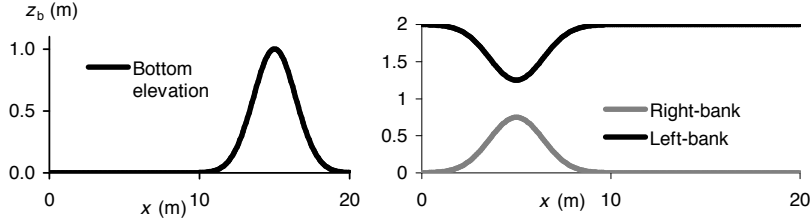


Figure 4: Test 2 - Frictionless steady state flow in a non-prismatic, rectangular channel. Channel profile: left, bottom elevation; right, left- and right-bank profile.

Symbol	Meaning	Value
g	Gravitational acceleration	9.81 m s^{-2}
L	Length of the domain	20 m
W_0, z_b	Channel width and bottom elevation	Figure 4
W_1	Derivative of the width with respect to z	0
z_0	Initial free surface elevation	1.1 m
z_{ds}	Prescribed surface elevation downstream	1.1 m
Q_{up}	Prescribed discharge upstream	$2 \text{ m}^3 \text{ s}^{-1}$
n_M	Manning's friction coefficient	$0 \text{ m}^{-1/3} \text{ s}$
Δx	Computational cell width	0.1 m
Δt	Simulation time	400 s

Table 2: Test 2 - Frictionless steady state flow in a non-prismatic, rectangular channel. Parameters of the test case.

357 over the whole domain except in the immediate vicinity of the hydraulic jump ($x \approx 17\text{m}$).

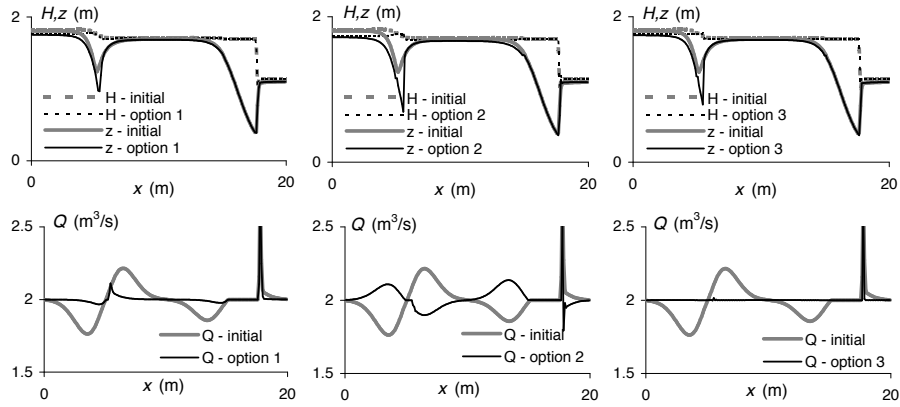
358 5.1.3. Test 3: frictionless steady-state flow in a non-prismatic trapezoidal channel

359 The channel profile is shown in Figure 6. It presents two simultaneous reductions of the
360 cross-section (bump and width narrowing), located at the same abscissa. The channel is not
361 prismatic with a variable bank slope yielding a transition from a trapezoidal shape at the
362 boundaries to a rectangular shape at half length. The parameters used for this steady-state,
363 frictionless test case are given in Table 3.

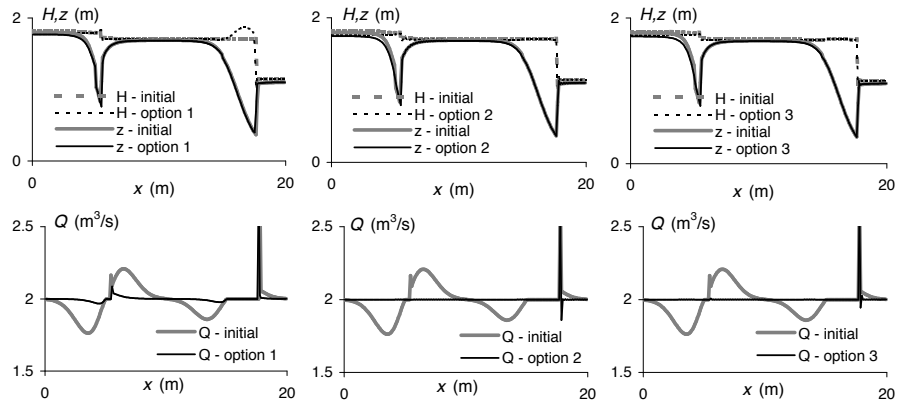
364 The simulated free surface elevation z , hydraulic head H and discharge Q , obtained
365 with the three AVB options and the three solvers are given in Figure 7. In this case again,
366 all three AVB options provide improved solutions compared to that given by the initial
367 formulation, for which the transition from subcritical to supercritical conditions (and vice-
368 versa) is observed to induce strong variations in the estimation of the discharge. This
369 statement however is to be moderated concerning Option 2 combined with HLLC solver
370 that also yields such variations. In a largely lesser extent, option 1 also exhibits some small
371 variations in the discharge. Moreover, it can be seen that the abscissa of the hydraulic jump
372 is not exactly located using Option 1 with Roe's solver or Q -scheme, with an increase in
373 hydraulic head upstream the jump.

Symbol	Meaning	Value
g	Gravitational acceleration	9.81 m s^{-2}
L	Length of the domain	20 m
W_0, W_1	Channel width and its derivative with respect to z	Figure 6
z_b	Bottom elevation	Figure 6
z_0	Initial free surface elevation	1.2 m
z_{ds}	Prescribed surface elevation downstream	1.2 m
Q_{up}	Prescribed discharge upstream	$4 \text{ m}^3 \text{ s}^{-1}$
n_M	Manning's friction coefficient	$0 \text{ m}^{-1/3} \text{ s}$
Δx	Computational cell width	0.1 m
Δt	Simulation time	200 s

Table 3: Test 3 - Frictionless steady state flow in a non-prismatic trapezoidal channel. Parameters of the test case.



(a) HLL



(b) Roe (identical to Q -scheme)

Figure 5: Test 2 - Frictionless steady state flow in a non-prismatic, rectangular channel. Up: water elevation z and hydraulic head H , down: discharge Q obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options, using a) HLL solver, b) Roe's solver. The Q -scheme gives similar results to the Roe's solver.

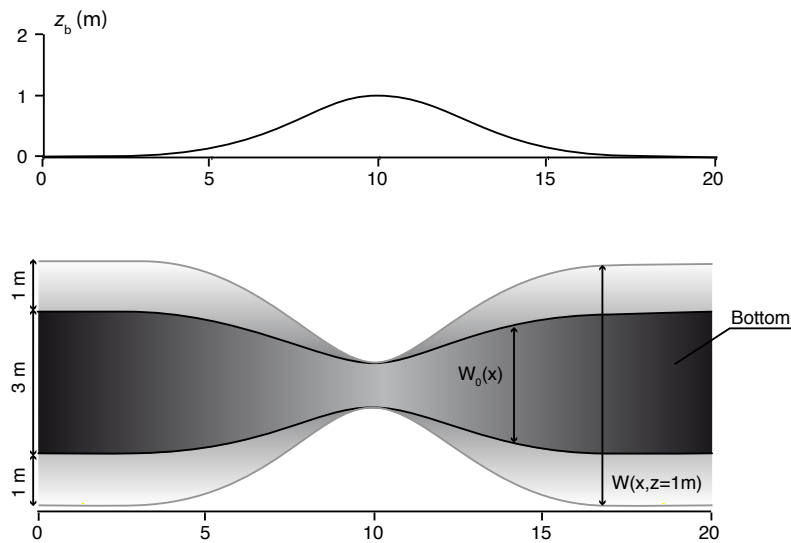


Figure 6: Test 3 - Frictionless steady state flow in a non-prismatic trapezoidal channel. Channel profile: top, bottom elevation; down, $W_0(x)$ and $W(x, z = 1)$.

374 Option 3 gives a uniform value for the discharge everywhere, except across the hydraulic
 375 jump, regardless the solver used. Once again, Option 3 is thus deemed more suitable to deal
 376 with transcritical flows.

377 5.1.4. Test 4: steady state flow in a Venturi flume

378 This test case involves the simultaneous presence of all source terms: friction, bottom
 379 slope and width variation. It is a real world test case for which experiment validation has
 380 been carried out in a channel of 67 cm wide. The Venturi flume used is 2.5 m long with
 381 narrow section of 10 cm wide (Figure 8). The flume is made of aluminium plates, with a
 382 Manning friction coefficient $n_M = 10^{-2} \text{ m}^{-1/3} \text{ s}$ calibrated from experiments in a straight
 383 channel made of the same material.

384 In the experiment, steady state was obtained under a discharge of 40 litres per second.
 385 The elevation of the free surface along the walls and axis of the channel was measured every
 386 5 cm. Figure 9 shows numerical results obtained with Roe’s solver and the different AVB
 387 options. The three AVB options give similar results. The unit-discharge is better estimated
 388 upstream than with the initial formulation. Option 2 and 3 give erroneous results with the
 389 HLL solver and the Q -scheme. Figure 10 shows the longitudinal profiles of the measured
 390 and simulated free surface. As can be seen from the figure, the simulation agrees well with
 391 the measurement upstream and downstream of the narrowing. In contrast, the free surface
 392 elevation is overestimated by the numerical model in the narrow section of the Venturi flume.
 393 Besides, the curvature of the simulated free surface profile is wrong. These results invalidate
 394 the shallow water assumption of a hydrostatic pressure distribution but this is beyond the
 395 scope of the present paper.

396 5.2. Transient test cases

397 There is no guarantee that an accurate well-balanced approach for steady state flows,
 398 gives correct results on transient configurations. The following transient test cases are thus
 399 performed.

400 5.2.1. Test 5: frictionless dam-break problem in a rectangular channel with flat bottom

401 The dam-break problem is an initial-value problem in which the water is initially at rest
 402 and the water levels are different on both sides of the dam. The solution of the dam-break
 403 problem in rectangular channels is similar to that of the one-dimensional shallow water
 404 equations. The properties of the analytical solution are presented in [36]. The dam-break
 405 problem is a Riemann problem defined as:

$$h(x, 0) = \begin{cases} h_L & \text{for } x \leq x_0 \\ h_R & \text{for } x > x_0 \end{cases} \quad (64a)$$

$$q(x, 0) = 0 \forall x \quad (64b)$$

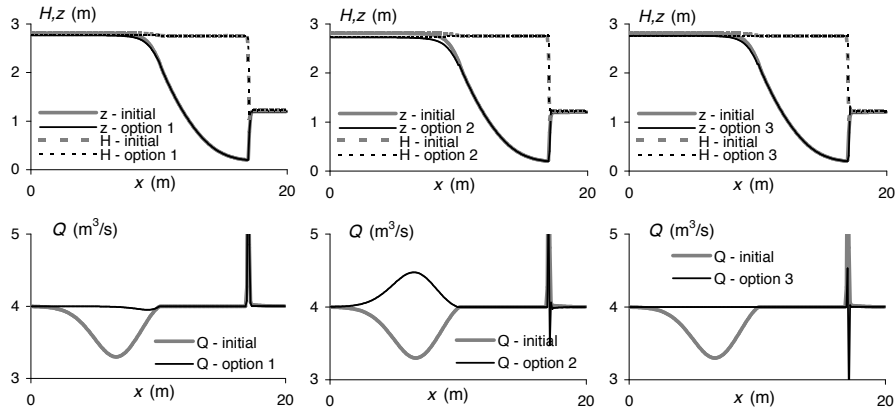
406 The solution is made of a rarefaction wave and a moving shock separated by a region of
 407 constant state. For the dam-break problem without source terms (friction or bottom slope),
 408 the profile obeys the following equations in the rarefaction wave

$$u(x, t) = \frac{2}{3} \left(c_L + \frac{x}{t} \right) \quad (65a)$$

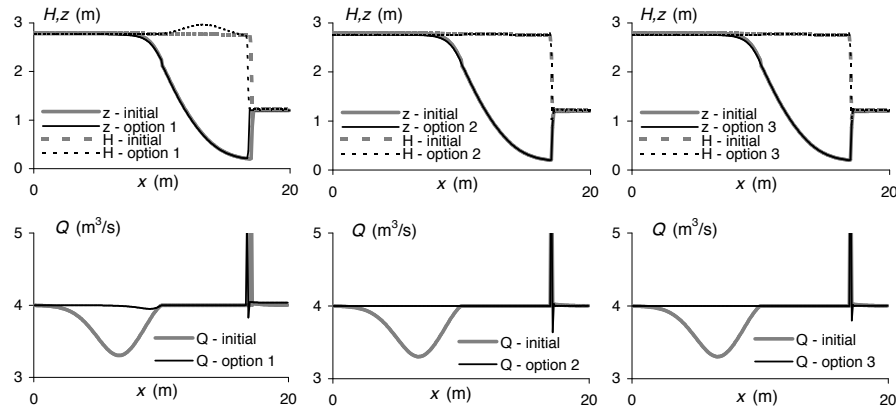
$$c(x, t) = \frac{1}{3} \left(2c_L - \frac{x}{t} \right) \quad (65b)$$

410 from which the expression of the flow solution \mathbf{U} is straightforward using $A = c^2/g$ and
 411 $Q = uA$. In the other parts of the domain, the profile is piecewise constant (see [36] for
 412 more details).

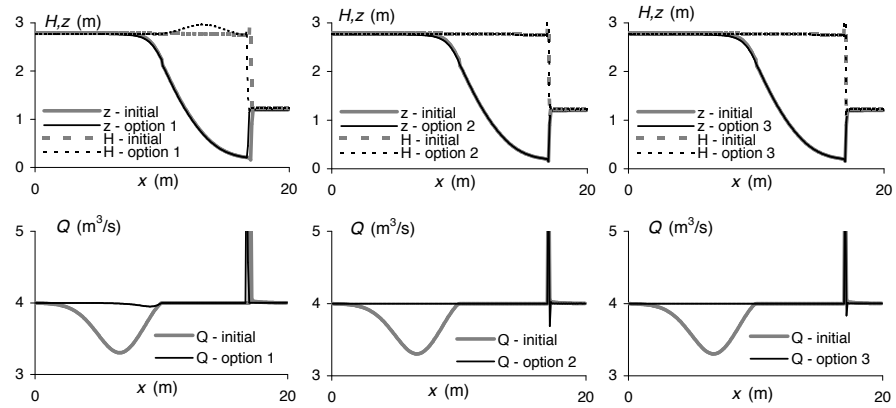
413 The parameters used in this test case are given in Table 4. Profiles of free surface
 414 elevation, hydraulic head and discharge, obtained with the initial formulation and the three
 415 AVB options are given in Figure 11 for the three solvers. Contrarily to previous test cases,
 416 the discharge Q is correctly estimated by each option included initial formulation except
 417 for the combinations HLL/Option 2 (Figure 11b) and Q -scheme/Option 3 (Figure 11d).
 418 For these latter, the free surface elevation and hydraulic head are discontinuous across the
 419 critical point (Note that this problem was pointed out in [30, 18] where a specific treatment
 420 of the critical point was proposed).



(a) HLL



(b) Roe



(c) Q-scheme

Figure 7: Test 3 - Frictionless steady state flow in a non-prismatic trapezoidal channel. Top: water elevation z and hydraulic head H , down: discharge Q obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options, using a) HLL solver, b) Roe's solver, c) Q-scheme.

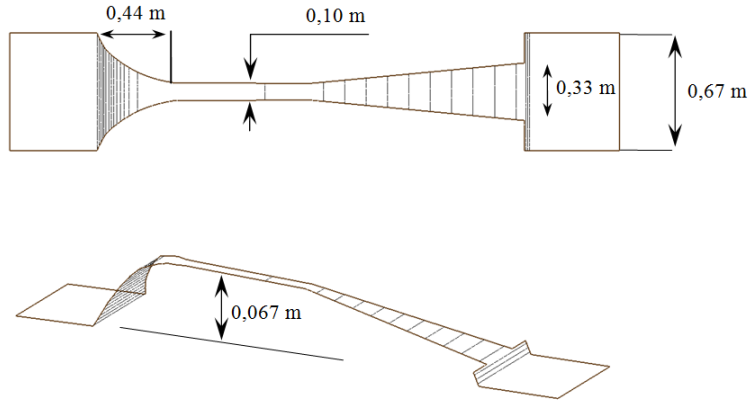


Figure 8: Test 4 - Dimensions of the Venturi flume used in the experiment. Top: plan view. Bottom: bird eye's view with a vertical scale magnified by a factor 5.

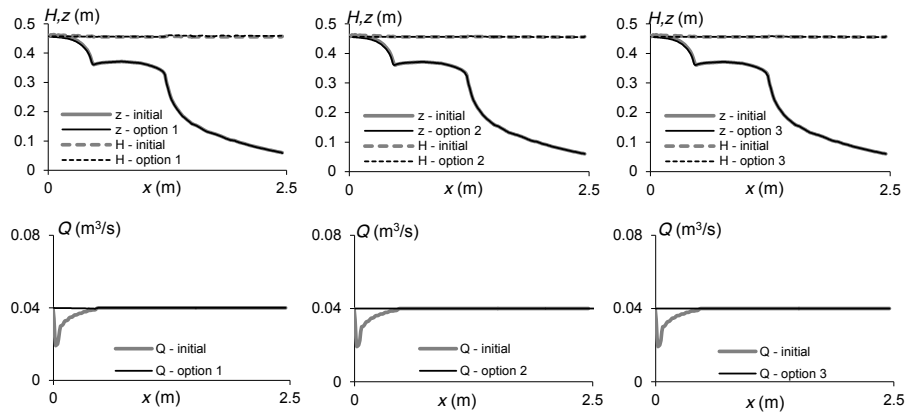


Figure 9: Test 4 - Steady state flow in a Venturi flume. Top: water elevation z and hydraulic head H , down: discharge Q obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options, using Roe's solver.

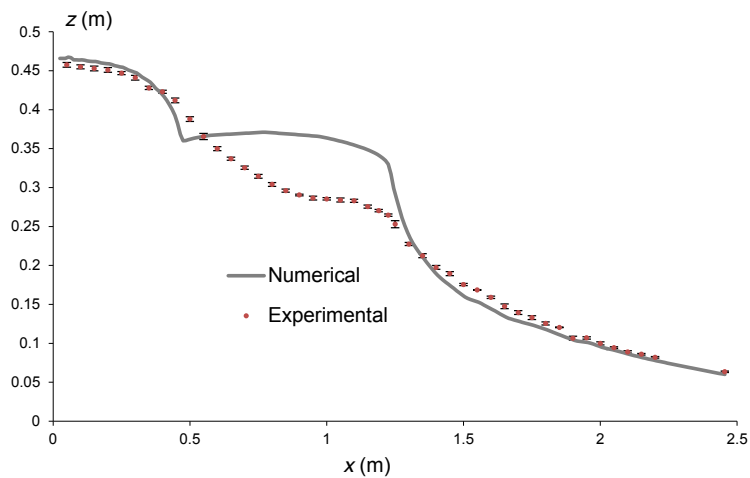


Figure 10: Test 4 - Steady state flow in a Venturi flume. Comparison between numerical results and experimental data.

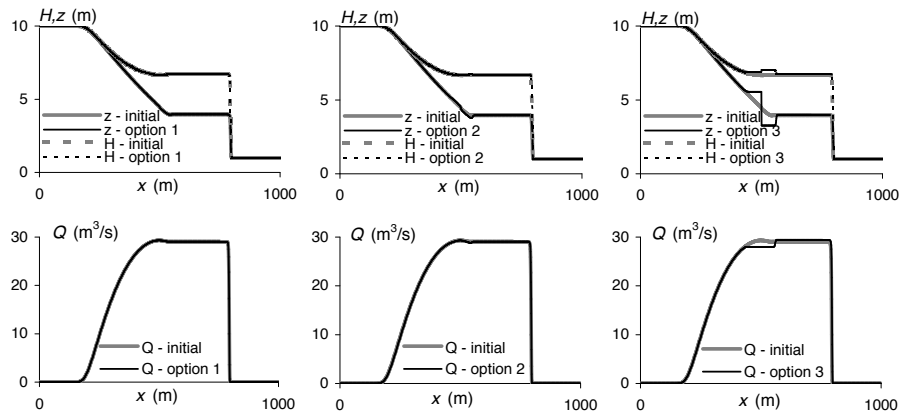
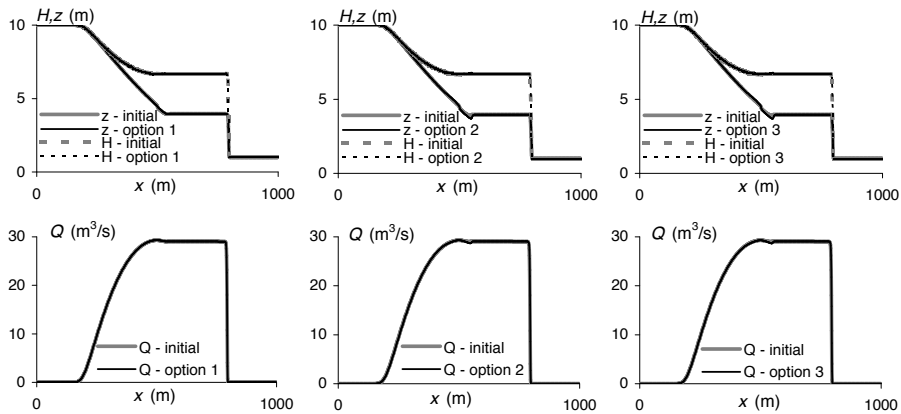
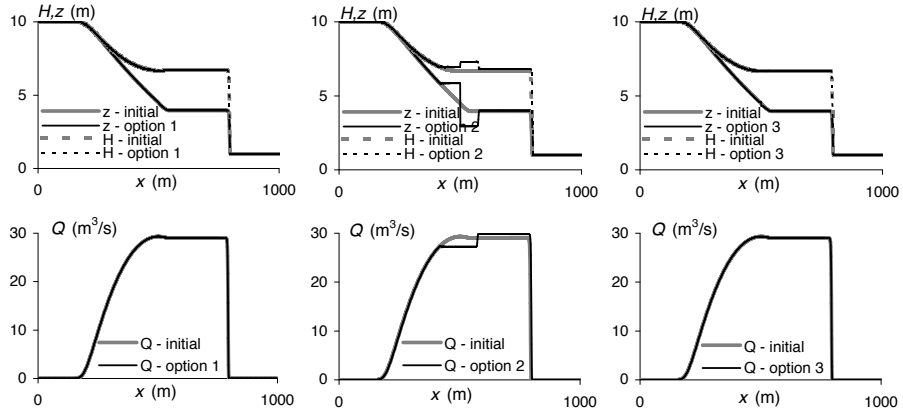
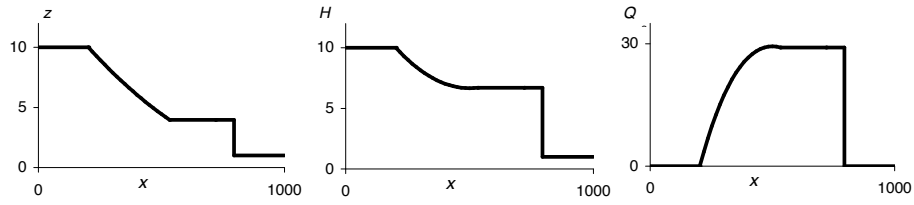


Figure 11: Test 5 - Dam-break problem in a rectangular channel. a) analytical solution; b) HLL solver; c) Roe's solver; d) Q -scheme. Top: water surface elevation z and hydraulic head H , bottom: discharge Q , obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options.

Symbol	Meaning	Value
g	Gravitational acceleration	9.81 m s^{-2}
L	Length of the domain	1,000 m
W_0	Channel width	1 m
W_1	Derivative of the channel width with respect to z	0
z_b	Bottom elevation	0 m
h_L	Initial free surface elevation on the left-hand side of the dam	10 m
h_R	Initial free surface elevation on the right-hand side of the dam	1 m
n_M	Manning's friction coefficient	$0 \text{ m}^{-1/3} \text{ s}$
Δx	Computational cell width	1 m
Δt	Simulation time	30 s

Table 4: Test 5 - dam-break problem in a rectangular channel. Parameters of the test case.

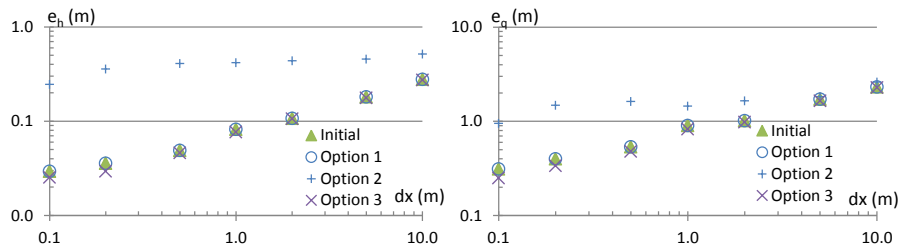


Figure 12: Test 5 - Dam-break problem in a rectangular channel. Convergence analysis using HLL solver and the three AVB options. L_2 -norm between the computed output water depth (left) or unit discharge (right) and analytical solution.

421 Since the analytical solution is available for the dam-break problem, a convergence
422 analysis is performed on this test case using the three AVB options and HLL solver. Figure 12
423 shows that options 1 and 3 have almost the same convergence as initial formulation (slightly
424 faster for Option 3), and confirms the non-convergence of Option 2 used with HLL solver.

425 5.2.2. Test 6: frictionless dam-break problem in a triangular channel with flat bottom

426 This test case is identical to the previous one (dam-break problem in a rectangular
427 channel, without bottom slope or friction) except that the cross-section of the channel has
428 a triangular shape. The parameters of the test case are the same as given in Table 4 for
429 Test 5 except that $W_0 = 0$ and $W_1 = 2$ in the whole domain.

430 The analytical solution is given by [25]:

$$431 \quad u_* + 4c_* = u_L + 4c_L \quad (66a)$$

$$432 \quad Q_* - Q_R = (A_* - A_R) c_s \quad (66b)$$

$$433 \quad \left(\frac{Q^2}{A} + \frac{gA^2}{2} \right)_* - \left(\frac{Q^2}{A} + \frac{gA^2}{2} \right)_R = (Q_* - Q_R) c_s \quad (66c)$$

434 where the subscript $*$ denotes the intermediate region of constant state.

435 Equation (66a) expresses the invariance of the Riemann invariant ($u + 4c$) across the
436 rarefaction wave. Equations (66b) and (66c) are the jump relationships across the shock
437 moving at the speed c_s . The unknown shock speed can be eliminated from the system by
438 combining the second and third equations. The system can then be solved iteratively to
439 find the values of A and Q in the intermediate region of constant state using $A = c^2/g$ and
 $Q = uA$. Across the rarefaction wave, u and c verify:

$$440 \quad u + 4c = u_L + 4c_L \quad (67a)$$

$$441 \quad u - c = \frac{x}{t} \quad (67b)$$

yielding the following profile for u and c in the rarefaction wave:

$$442 \quad u(x, t) = \frac{4}{5} \left(c_L + \frac{x}{t} \right) \quad (68a)$$

Symbol	Meaning	Value
L	Length of the domain	1,000 m
W_0	Channel width	1 m
W_1	Derivative of the channel width with respect to z	0
z_{bL}	Bottom elevation on the left-hand side of the dam	0 m
z_{bR}	Bottom elevation on the right-hand side of the dam	5 m
h_L	Initial free surface elevation on the left-hand side of the dam	15 m
h_R	Initial free surface elevation on the right-hand side of the dam	1 m
n_M	Manning's friction coefficient	$0 \text{ m}^{-1/3} \text{ s}$
Δx	Computational cell width	1 m
Δt	Simulation time	30 s

Table 5: Test 5 - dam-break problem in a rectangular channel. Parameters of the test case.

442

$$c(x, t) = \frac{1}{5} \left(4c_L - \frac{x}{t} \right) \quad (68b)$$

443

from which A and Q profiles can be determined.

444

Results of water elevation z , hydraulic head H and discharge Q , obtained with the three AVB options and the three solvers are given in Figure 13. In this case again, Roe's solver gives satisfactory results with the 3 options as well as the initial formulation. However, very strong discontinuities at the critical point can be seen with Option 2 and 3 combined with HLL solver and Q -scheme, yielding to an underestimation of the maximum discharge. Moreover, the shock is incorrectly located with Option 2/HLL.

445

450

5.2.3. Test 7: frictionless dam-break problem on a bottom step

451

The parameters of this test case are given in Table 5. A bottom step of 5 m is located at the same abscissa as the initial water depth discontinuity. The analytical solution (that can be found for example in [1, 5]) as well as results obtained with the three AVB options and Roe's solver are given in Figure 14. HLL solver and Q -scheme provide erroneous solutions with Option 2 and 3.

452

456

6. Discussion - Conclusions

457

In practical engineering applications, geometrical source terms arising e.g. from bottom slope or the non prismatic character of the channel are to be accounted for in the governing equations. These source terms can in general not be discretized independently of the conservation part. Riemann-solver based techniques compute the fluxes from the average cell values on the left and right hand of the interface. The flux can be seen as a combination of the average cell fluxes, augmented with a diffusion term involving the gradient in the conserved variable. Artificial oscillations may appear in the computed profiles if the gradients (and hence the diffusive part of the flux) is not estimated properly. The Auxiliary Variable-Based balancing, consists of using an "auxiliary" variable instead of the conserved one in the flux function, defined so as to allow the steady-state condition (of which static equilibrium is only a particular case) to be preserved. It is applied to the one-dimensional open channel equations in the present paper.

458

Three different options of AVB have been tested in this paper in addition to the classical flux formulation that uses the gradient of the conserved variables: 1) free surface elevation; 2) specific force and 3) hydraulic head. The application of the method to three classical approximate Riemann solvers (HLL, Roe and Q -scheme) is also presented.

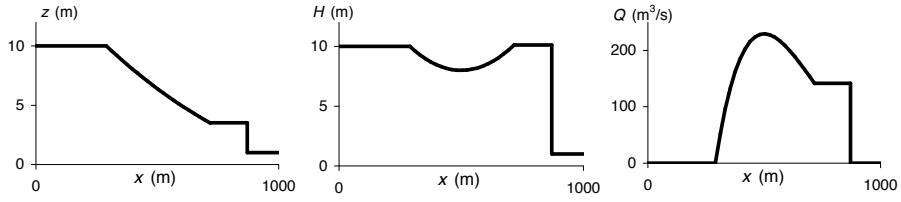
459

460

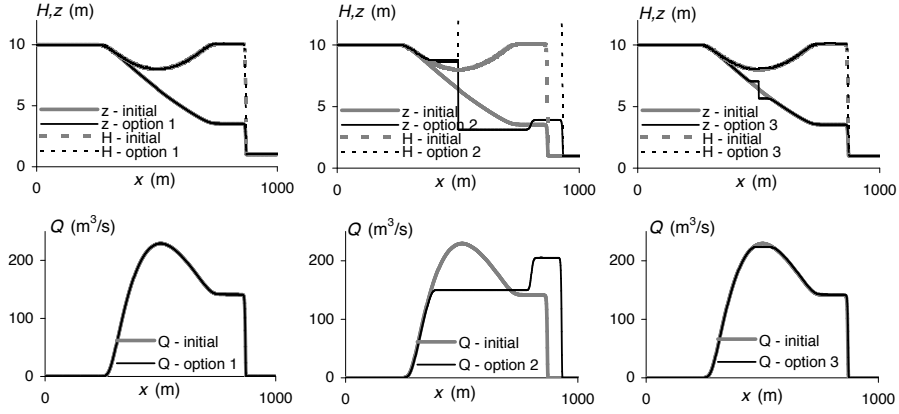
Various steady-state test cases including singular head losses, friction, bottom and width variation, non-prismatic configurations, have been implemented to assess the ability of the AVB approach to deal with transcritical flows, the critical points being well-known to introduce instabilities. In the steady-state test cases, the three options generally gives a better estimate of the uniform discharge than the initial formulation. Option 3, based on the hydraulic head, is the one that gives a uniform discharge equal to the prescribed one with the best accuracy, for each steady-state test case and in the whole domain, except across hydraulic jumps where a small spike remains. It is important to check the validity of these approaches for unsteady states configurations. Indeed, some examples in the literature that give correct results in steady state configurations (such as [28]) have revealed incorrect on

461

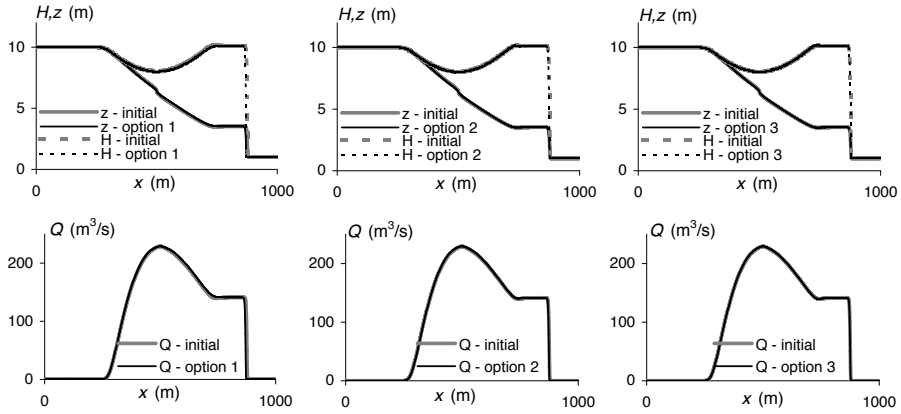
462



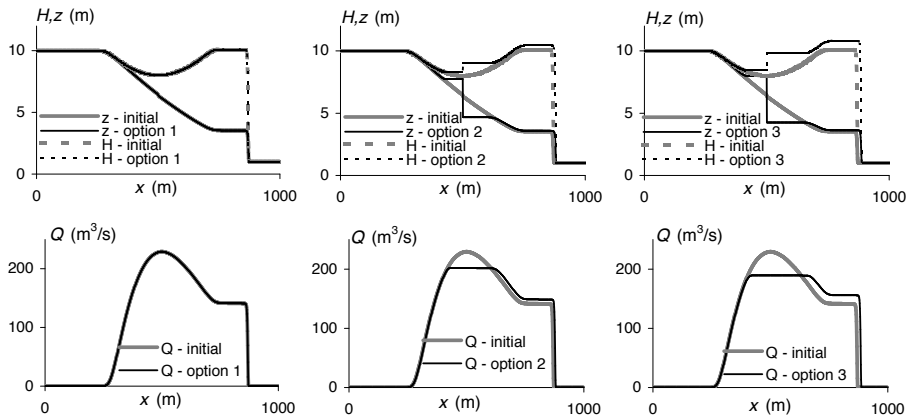
(a) Analytical solution



(b) HLL

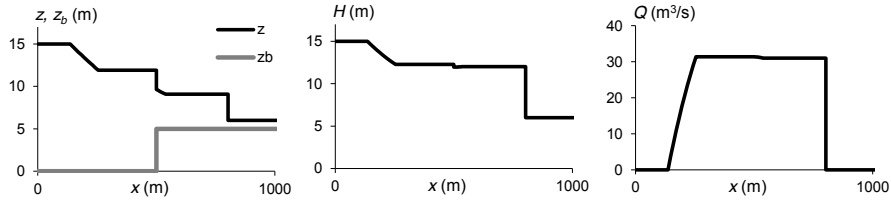


(c) Roe

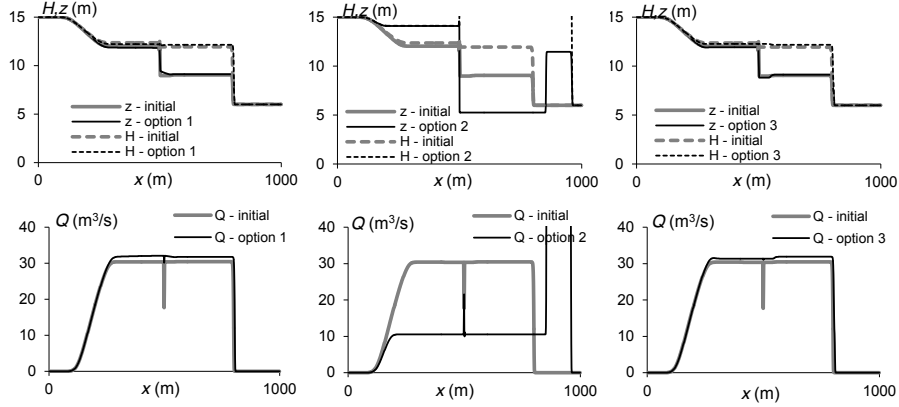


(d) Q -scheme

Figure 13: Test 6 - Dam-break problem in a triangular channel. a) Analytical solution, b) HLL solver, c) Roe's solver and d) Q -scheme. For each sub-figure, top: hydraulic head H and water elevation z , bottom: discharge Q , obtained with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options.



(a) Analytical solution



(b) Roe

Figure 14: Test 07 - Dam-break problem on a bottom step. a) Analytical solution; b) top: water surface elevation z and hydraulic head H , bottom: discharge Q , obtained using Roe's solver with $\mathbf{V} = \mathbf{U}$ (Initial) and with the three different AVB options.

483 transient test cases, such as shown in the three unsteady configurations presented (dam-
 484 break in rectangular or triangular channels and over a bottom step) where Option 2 gives
 485 bad results when used with HLL or Q -scheme.

486 Finally, when used with Roe's solver, Option 3 is the only one that produces correct
 487 results for all the test cases. An interesting feature of this option is that it allows head loss
 488 functions (stemming from e.g. bridges or other singularities) to be accounted for directly
 489 within the discretized equations. In contrast with what is classically done in commercially
 490 available river packages, the AVB method eliminates the need for internal boundaries across
 491 hydraulic singularities.

492
 493 In the present paper, the geometric source term is accounted for by integrating the
 494 bottom slope over the surface of the (non-horizontal) computational cell. This procedure is
 495 rather easy to carry out when the cell is full. But in the case of wetting/drying, the water
 496 does not occupy the full length of the cell. Computing the integral of the term ghS_0 over
 497 only part of the cell becomes a very complex and time-consuming task. The approach to
 498 source term estimation proposed in the present paper is thus not the best possible option
 499 to the discretization of real-world geometries and practical river problems. An alternative
 500 option is currently under study. It consists in considering each cell as prismatic and lumping
 501 the geometric source term at the cell interfaces. This approach, however, requires that a
 502 proper splitting of the lumped source terms between the adjacent cells to the interfaces be
 503 devised. This point is currently under study.

504 Finally, this paper deals with finite volume Godunov-type discretizations, but if higher-
 505 order schemes are to be designed, the AVB method may be applied by reconstructing the
 506 auxiliary variable.

507 References

- 508 [1] F. Alcrudo and F. Benkhaldoun. Exact solutions to the Riemann problem of the shallow
 509 water equations with a bottom step. *Computers & Fluids*, 30:643–671, 2001.
- 510 [2] E. Audusse and M.-O. Bristeau. A well-balanced positivity preserving "second-order"

- 511 scheme for shallow water flows on unstructured meshes. *Journal of Computational*
512 *Physics*, 206:311–333, 2005.
- 513 [3] F. Benkhaldoun, I. Elmahi, and M. Seaid. A new finite volume method for flux-gradient
514 and source-term balancing in shallow water equations. *Computer Methods in Applied*
515 *Mechanics and Engineering*, 199:3324–3335, 2010.
- 516 [4] A. Bermudez and M.E. Vazquez. Upwind methods for hyperbolic conservation laws
517 with source terms. *Computers & Fluids*, 23:1049–1071, 1994.
- 518 [5] R. Bernetti, V.A. Titarev, and E.F. Toro. Exact solution of the Riemann problem for the
519 shallow water equations with discontinuous bottom geometry. *Journal of Computational*
520 *Physics*, 227:3212–3243, 2008.
- 521 [6] J. Burguete and P. Garcia-Navarro. Improving simple explicit methods for unsteady
522 open channel and river flow. *International Journal for Numerical Methods in Fluids*,
523 45:125–156, 2004.
- 524 [7] V. Caleffi, A. Valiani, and A. Bernini. Fourth-order balanced source term treatment in
525 central weno schemes for shallow water equations. *Journal of Computational Physics*,
526 218(1):228 – 245, 2006.
- 527 [8] A. Canestrelli, M. Dumbser, A. Siviglia, and E.F. Toro. Well-balanced high-order
528 centered schemes on unstructured meshes for shallow water equations with fixed and
529 mobile bed. *Advances in Water Resources*, 33(3):291–303, MAR 2010.
- 530 [9] H. Capart, T.I. Eldho, S.Y. Huang, D.L. Young, and Zech Y. Treatment of natural
531 geometry in finite volume river flow computations. *Journal of Hydraulic Engineering -*
532 *ASCE*, 129:385–393, 2003.
- 533 [10] M.J. Castro, J.M. Gallardo, and C. ParÈs. High-order finite volume schemes based on
534 reconstruction of states for solving hyperbolic systems with nonconservative products.
535 applications to shallow-water systems. *Mathematics of Computation*, 75:1103–1134,
536 2006.
- 537 [11] T. Chacon Rebollo, A. Dominguez Delgado, and E.D. Fernandez Nieto. A family of
538 stable numerical solvers for the shallow water equations with source terms. *Computer*
539 *Methods in Applied Mechanics and Engineering*, 192:203–225, 2003.
- 540 [12] T. Chacon Rebollo, A. Dominguez Delgado, and E. Fernandez Nieto. Asymptotically
541 balanced schemes for non-homogeneous hyperbolic systems - application to the shallow
542 water equations. *Comptes Rendus de l'Académie des Sciences Serie I*, 338:85–90, 2004.
- 543 [13] Nelida Crnjaric-Zic and Bojan Crnkovic. High order accurate semi-implicit WENO
544 schemes for hyperbolic balance laws. *Applied mathematics and computation*, 217(21):
545 8611–8629, JUL 1 2011.
- 546 [14] J.A. Cunge, F.M. Jr Holly, and A. Verwey. *Practical Aspects of River Computational*
547 *Hydraulics*. Pitman, 1980.
- 548 [15] S.F. Davis. Simplified second-order Godunov-type methods. *SIAM Journal of Scientific*
549 *and Statistical Computing*, 9:445–473, 1988.
- 550 [16] M. Dumbser and E.F. Toro. A simple extension of the osher riemann solver to non-
551 conservative hyperbolic systems. *Journal of Scientific Computing*, 48:70–88, 2011.
- 552 [17] B. Einfeldt. On Godunov-type methods for gas dynamics. *SIAM Journal of Numerical*
553 *Analysis*, 25:294–318, 1988.
- 554 [18] P. Finaud-Guyot, C. Delenne, J. Lhomme, V. Guinot, and C. Llovel. An approximate-
555 state Riemann solver for the two-dimensional shallow water equations with porosity.
556 *International Journal for Numerical Methods in Fluids*, 62:1299–1331, 2010.
- 557 [19] T. Gallouet, J.M. Herard, and N. Seguin. On the use of symmetrizing variables for
558 vacuums. *CALCOLO*, 40:163–194, 2003.

- 559 [20] T. Gallouet, J.M. HÈrard, and N. Seguin. Some approximate godunov schemes to
560 compute the shallow-water equations with topography. *Computers & Fluids*, 32:479–
561 513, 2003.
- 562 [21] P. Garcia-Navarro and M.E. Vazquez-Cendon. On numerical treatment of the source
563 terms in the shallow water equations. *Computers & Fluids*, 29:951–979, 2000.
- 564 [22] P. Glaister. Approximate Riemann solutions of the shallow water equations. *Journal*
565 *of Hydraulic Research*, 26:293–306, 1988.
- 566 [23] S.K. Godunov. A difference method for calculation of discontinuous solutions of hydro-
567 dynamics. *Matematicheski Sbornik*, 47:271–306, 1959.
- 568 [24] J.M. Greenberg and A. Leroux. A well-balanced scheme for the numerical processing of
569 source terms in hyperbolic equations. *SIAM Journal of Numerical Analysis*, 33:1–16,
570 1996.
- 571 [25] V. Guinot. *Wave Propagation in Fluids. Models and numerical techniques. 2nd edition.*
572 Wiley-ISTE, 2010.
- 573 [26] A. Harten, P.D. Lax, and B. Van Leer. On upstream differencing and Godunov-type
574 schemes for hyperbolic conservation laws. *Journal of Computational Physics*, 50:235–
575 269, 1983.
- 576 [27] G. Kesserwani and Q. Liang. Well-balanced RKDG2 solutions to the shallow water
577 equations over irregular domains with wetting and drying. *Computers and Fluids*, 39:
578 2040–2050, 2010.
- 579 [28] S.H. Lee and N.G. Wright. Simple and efficient solution of the shallow water equations
580 with source terms. *International Journal for Numerical Methods in Fluids*, 63:313–340,
581 2010.
- 582 [29] R.J. LeVeque. Balancing source terms and flux gradients in high-resolution Godunov
583 methods: the quasi-steady wave propagation algorithm. *Journal of Computational*
584 *Physics*, 146:346–365, 1998.
- 585 [30] J. Lhomme and V. Guinot. A general approximate-state solver for hyperbolic systems
586 of conservation laws with source terms. *International Journal for Numerical Methods*
587 *in Fluids*, 53:1509–1540, 2007.
- 588 [31] Q. Liang and F. Marche. Numerical resolution of well-balanced shallow water equations
589 with complex source terms. *Advances in Water Resources*, 32:873–884, 2009.
- 590 [32] J. Murillo, P. García-Navarro, J. Burguete, and P. Brufau. The influence of source terms
591 on stability, accuracy and conservation in two-dimensional shallow flow simulation using
592 triangular finite volumes. *International Journal for Numerical Methods in Fluids*, 54
593 (5):543–590, 2007.
- 594 [33] M. Nujic. Efficient implementation of non-oscillatory schemes for the computation of
595 free-surface flows. *Journal of Hydraulic Research*, 33:101–111, 1995.
- 596 [34] P.L. Roe. Approximate Riemann solvers, parameter vectors and difference schemes.
597 *Journal of Computational Physics*, 43:357–372, 1981.
- 598 [35] L. Song, J. Zhou, J. Guo, Q. Zou, and Y. Liu. A robust well-balanced finite volume
599 model for shallow water flows with wetting and drying over irregular terrain. *Advances*
600 *in Water Resources*, 34(7):915–932, JUL 2011.
- 601 [36] J.J. Stoker. *Water Waves*. Interscience, 1957.
- 602 [37] A Valiani and L Begnudelli. Divergence form for bed slope source term in shallow water
603 equations. *Journal of Hydraulic Engineering - ASCE*, 132(7):652–665, JUL 2006.
- 604 [38] M.E. Vazquez-Cendon. Improved treatment of source terms in upwind schemes for the
605 shallow water equations in channels with irregular geometry. *Journal of Computational*
606 *Physics*, 148:497–526, 1999.

- 607 [39] Y. Xing and C.-W. Shu. High order well-balanced finite volume weno schemes and
608 discontinuous galerkin methods for a class of hyperbolic systems with source terms.
609 *Journal of Computational Physics*, 214(2):567 – 598, 2006.
- 610 [40] J.G. Zhou, D.M. Causon, C.G. Mingham, and D.M. Ingram. The surface gradient
611 method for the treatment of source terms in the shallow water equations. *Journal of*
612 *Computational Physics*, 168:1–25, 2001.