

# Statistical and empirical analyses of the triggers of coastal chalk cliff failure

Pauline Letortu, Stéphane Costa, Jean-Michel Cador, Cyril Coinaud, Olivier Cantat

### ▶ To cite this version:

Pauline Letortu, Stéphane Costa, Jean-Michel Cador, Cyril Coinaud, Olivier Cantat. Statistical and empirical analyses of the triggers of coastal chalk cliff failure. Earth Surface Processes and Landforms, 2015, 40 (10), pp.1371-1386. 10.1002/esp.3741. hal-01148291

HAL Id: hal-01148291

https://hal.science/hal-01148291

Submitted on 25 Dec 2020

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

- 1 Statistical and empirical analyses of the triggers of coastal chalk cliff failure
- 2 Pauline Letortu<sup>1,2\*</sup>, Stéphane Costa<sup>1</sup>, Jean-Michel Cador<sup>1</sup>, Cyril Coinaud<sup>3</sup>, Olivier
- 3 Cantat<sup>1</sup>
- <sup>4</sup> Laboratory LETG Caen-Géophen, University of Caen Basse-Normandie, Caen,
- 5 France.
- 6 <sup>2</sup> Laboratory LETG Brest-Géomer, University of Bretagne Occidentale, Brest, France.
- 7 <sup>3</sup> University Data Platform of the University of Caen, University of Caen Basse-
- 8 Normandie, Caen, France.
- 9 \* Correspondence to: P. Letortu, Laboratory LETG-Brest Géomer, University of
- 10 Bretagne Occidentale, IUEM-UBO, Technopôle Brest-Iroise, 29280 Plouzané,
- 11 France. Telephone: +33298498754. Fax: +33298498703. Email:
- 12 pauline.letortu@univ-brest.fr

- 14 ABSTRACT:
- 15 The triggering factors of rock falls remain unknown due to a lack of exhaustive,
- regular and accurate surveys. Based on an inventory of 331 rock falls collected
- 17 weekly between 2002 and 2009 from Veules-les-Roses to Le Treport (Upper
- Normandy), the relationships between coastal chalk cliff rock falls (dates and
- 19 geomorphological features) and external factors commonly agreed as triggering
- 20 (rainfall, temperature variations, tide and wind) are studied. The combination of
- 21 multivariate statistical and empirical analyses indicates that (1) "cold and dry
- weather" and "high rainfall and high wind" are the conditions most likely to trigger
- rock falls, (2) the main triggering factors of rock falls are effective rainfall (for rock
- falls mostly between 200 and 1,400 m<sup>3</sup> or larger than 10,000 m<sup>3</sup> and coming from the
- 25 whole cliff face), freeze/thaw cycles (especially for rock falls smaller than 200 m<sup>3</sup> and

coming from the foot and top of the cliff face) and marine roughness (rock falls mainly smaller than 200 m³ and coming from the cliff foot). However, the contribution of each factor to triggering is difficult to determine because of combinations of factors (85 % of 331 cases), relays of processes and hysteresis phenomena. In view of these first results, it is still presumptuous to predict the location and time of triggering of rock falls. However, the statistical and naturalistic approaches adopted and the observations made in this study are from an original database, and constitute a real starting point for the prediction and prevention of the hazard of coastal chalk cliff rock falls in Upper Normandy.

35

36

26

27

28

29

30

31

32

33

34

KEYWORDS: coastal chalk cliff rock falls, triggering factors, statistical analyses,

37 empirical analysis, Upper Normandy

38

39

#### Introduction

- Except for a consensus about the influence of structural features, external triggering
  factors of rock falls are still under scientific debate. In this paper, the term "rock fall" is
  used to describe movements of coherent rock (Varnes, 1978).

  Many studies emphasize the importance of sub-aerial actions over marine actions
  (Brossard and Duperret, 2004; Pierre and Lahousse, 2006), which are assumed to be
- just a transport agent. The role of rainfall through surface runoff, infiltration, and water
- table levels may reduce the stability of cliffs (Sunamura, 1992; Duperret et al., 2002;
- 47 Hénaff et al., 2002a; Young et al., 2009). Because of the porosity of coastal chalk
- 48 cliffs, they are particularly sensitive to processes induced by water (Hutchinson,
- 49 1972; May and Heeps, 1985; Duperret et al., 2002, 2004; Lahousse and Pierre,
- 50 2003; Pierre and Lahousse, 2006). Freezing action is also mentioned. Many

geomorphological studies have focused mainly on the freezing process (Robinson and Jerwood, 1987) whereas thermoclasty has been less studied (Hall, 1999). This freezing process has been extensively studied and sometimes quantified, especially in periglacial environments, and may be conducive to triggering rock falls (Fournier and Allard, 1992; Bernatchez and Dubois, 2008). The porosity of chalk leads to a high sensitivity to temperature variations, with freezing often being considered a triggering factor. For the coastal chalk cliffs studied, the porosity ranges between 32 % and 46 % from the Turonian to Campanian stages (Duperret et al., 2005). The characteristics (including its porosity range) and salt content (Robinson and Jerwood, 1987; Jerwood et al., 1990a, 1990b) of chalk identify it as a "frost-shattered" rock. It is the number of freeze/thaw cycles that is important rather than the intensity of freezing (Letavernier, 1984). Cryoclastism can be quite remarkable in terms of debris production: during the thaw of 31/12/1995, the volume of debris was estimated at 900 m<sup>3</sup> for a cliff face surface of 44,000 m<sup>2</sup> (Costa, 1997). Some authors consider that marine actions are too often overlooked and emphasize the prevalence of these actions in cliff retreat (Hoek and Bray, 1977; McGreal, 1979; Sunamura, 1982, 1992; Trenhaile, 1987; Trenhaile and Kanyaya, 2007). Their role as a factor in erosion is widely recognized (King, 1972; Emery and Kuhn, 1982; Sunamura, 1982; Trenhaile, 1987; Stephenson, 2000; Woodroffe, 2002; Costa et al., 2006a; Lee, 2008; Castedo et al., 2012). The potential of wave action to destabilize cliffs has been demonstrated by modeling (Emery and Kuhn, 1982; Sunamura, 1982; Trenhaile 2009, 2010) and in the field by microseismic sensors (Adams et al., 2002; Senfaute et al., 2009; Young et al., 2012) or with high resolution laser tools to view the effects of undercutting and extraction (Dewez et al., 2013). Moreover, waves may use gravels and boulders as projectiles to increase their attack power, causing

51

52

53

54

55

56

57

58

59

60

61

62

63

64

65

66

67

68

69

70

71

72

73

74

undercutting (Robinson, 1977; Costa et al., 2006b). On rocky chalk coasts, wave action has been increasingly studied as a triggering factor of rock falls (Brossard and Duperret, 2004; Costa, 2005; Costa et al., 2006a and b; Dornbusch et al., 2008). Although Brossard and Duperret (2004) stated that the impact frequency of waves without projectiles seems to be too low to trigger large rock falls, it appears that the triggering of rock falls related to marine actions often occurs during paroxysmal storm conditions. In the recent scientific literature, the triggering of rock falls seems increasingly the result of a combination of factors, as the coastal environment is a complex system. However, this may reflect the difficulty of identifying the contribution of each factor, which results from a lack of precise dating of rock falls over long periods, a lack of high frequency data about the factors involved in the triggering and the diversity of scales at which the data are recorded (Genter et al., 2004; Naylor et al., 2010). Moreover, the results are often valid only for the study sites because of specific contexts of morphostructural features and local meteorological and marine conditions. Collaboration between the LETG-Caen Geophen laboratory and the non-profit organization ESTRAN led to the creation of a weekly inventory of rock falls between 2002 and 2009 from Veules-les-Roses to Le Treport (37.5 km). Using this original database of 331 events over 7 years, the objective of this work is to participate in the debate about the triggering factors of rock falls with two main research questions: 1) which marine and meteorological conditions are conducive to rock falls? 2) what is/are the triggering factor(s) of rock falls? To answer these questions, this paper describes the application of two methods: statistical analyses to identify the meteorological and marine conditions conducive to

76

77

78

79

80

81

82

83

84

85

86

87

88

89

90

91

92

93

94

95

96

97

98

99

rock falls and an empirical analysis to give a hierarchy of factors leading to the triggering of rock falls. For each type of analysis, geomorphological features of rock falls (fallen volume and the part of the cliff face area affected by the rock fall) are integrated in our analyses to improve understanding of the relationships between triggering factors and the types of movements generated. In fact, this work aims to determine whether triggering factors generate various types of movements. If this is not the case, the geomorphological characterization of rock falls per triggering factor will improve the prediction of this hazard.

#### Climate and rock fall database along the studied coastline

In north-west France, the Upper Normandy coast (N 50°0'0", E 1°0'0") is located along the English Channel (Figure 1). The environment is macrotidal with an average tidal range of 8 m. Swell is limited but the wind sea can reach a significant wave height of 4 m in Dieppe (annual return period) (Augris et al., 2004). Upper Normandy presents a marine temperate west coast climate. Winter temperatures are positive but an average of 26 daily freeze/thaw cycles is recorded per year (minimal temperature can reach -15°C). Rainfall is distributed over the year (≈800 mm) although fall and winter are the wettest seasons (min: 51 mm in August and max: 94 mm in November). Daily rainfall can reach 77 mm in October (Meteo-France, 1971-2000) (Table I).

Located at the northwestern end of the sedimentary Paris Basin, Upper Normandy coastal chalk cliffs consist of Upper Cretaceous chalk (from the Cenomanian to Campanian stages) (Figure 1). These cliffs recede with an average retreat rate of 0.15 m/y but with a high spatial variability (0.23 m/y between Saint-Valery-en-Caux and Dieppe (Costa et al., 2004, 2006b; Letortu et al., 2014a)). These values are not

126 representative of rock dynamics because, in reality, retreat occurs in jerks (rock falls). 127 These rock falls, currently unpredictable, threaten urbanization settled too close to 128 the coastline (Figure 2). 129 From 2002 to 2009, the non-profit organization ESTRAN monitored the cliff face 130 evolution between Veules-les-Roses and Le Treport (37.5 km) to provide an 131 inventory of rock falls. These cliffs are made of Turonian to Campanian chalk (Figure 132 1). For each of the 331 rock falls inventoried, photographs and measurements were 133 taken and recorded on an index card with information about the location, length, 134 height, width, and date of the rock fall (if the latter was unknown, the date of 135 observation was used) (Figure 3A). The rock falls have a variable distribution in time 136 and location (Figures 3B and 3C) (Letortu et al., 2014b). However, they occur 137 especially along the Cap d'Ailly (Varengeville/Sainte-Marguerite-sur-Mer; 53 % of 138 total fallen volume; 75 % of total number of rock falls) and in winter (53 % of total 139 fallen volume; 41 % of total number of rock falls). 140 Geomorphological features were added to this database (Table II). Using 141 photographs, it was possible to observe the departure area of material on the cliff 142 face (10 rock falls have an unknown departure area due to the lack of photographs). 143 The cliff face was divided into 3 parts (foot/middle/top and combinations): 32 % of 144 rock falls occur at the foot, < 1 % in the middle part, 6 % at the top, 42 % include the whole cliff face and 20 % the foot and middle part of the cliff face. 145 146 This spatial distribution is important because it may give some clues about the 147 triggering factor. For example, a rock fall with a cliff top departure is more probably 148 due to sub-aerial agents than marine agents. 149 From rock fall measurements made by ESTRAN, the volume was calculated (rock fall 150 shape as a straight prism with a rectangular base; overall bulking factor of 33 %

151 (Hénaff et al., 2002b)). Twelve rock falls have an unknown volume due to the lack of measurements. The fallen volume ranges from 1 m<sup>3</sup> to 236,000 m<sup>3</sup> but almost 40 % 152 153 of rock falls are smaller than 200 m<sup>3</sup> (Figure 4A). The biggest example (236,000 m<sup>3</sup> according to ESTRAN field measurements) occurred on the night of 13th to14th March 154 155 2008 in Dieppe. 156 Due to the statistical range of the fallen volumes, and in order to perform statistical 157 and empirical analyses, the population was divided into different classes, determined 158 from the volume frequency histogram between 2002 and 2009 (method of observed 159 thresholds). Four classes appeared (Figure 4B): 160 - rock falls with a volume less than 200 m<sup>3</sup> (126 rock falls, 39 %). Some may have 161 been evacuated in a few days (according to Hénaff et al. (2002b): about 40 m<sup>3</sup> per 162 day are removed by marine actions along the Upper Normandy coast) and have left 163 few traces for the weekly inventory of the ESTRAN organization. Thus, the analysis 164 and results of these small volumes should be treated with caution; 165 - rock falls with a volume between 200 and 1,400 m<sup>3</sup> (100 rock falls visible from one 166 week to another, 31 %). These mainly concern only one part of the cliff (foot, middle, 167 top); - rock falls with a volume between 1,400 and 10,000 m<sup>3</sup> (77 rock falls, 24 %). These 168 169 usually affect one part of the cliff face and sometimes the whole section; 170 - rock falls with a volume greater than 10,000 m<sup>3</sup> (16 cases, 5 %). These movements 171 mainly affect the whole cliff face (15 cases, only 1 case from the cliff top). 172 The originality of the database lies in (1) the length of the observation period (7) 173 years) and the coastline monitored (37.5 km), which enable statistical analyses of 174 331 rock falls, (2) the high periodicity of the readings (every week), (3) the precise 175 date of the discrete retreat events (from the 10-day period to the day), (4) the

horizontal point of view, which is the best one to follow all cliff changes (Young et al., 2009), especially small ones, (5) the measurements and photographs of the rock falls, which enable geomorphological characterization of the movements (volume, departure area of the rock fall). This database can thus contribute to the current debate about triggering factors of coastal chalk cliff rock falls.

181

182

183

184

185

186

187

188

189

190

191

192

193

194

195

196

197

198

176

177

178

179

180

#### Methods

The rock fall database offers great potential for statistical analyses due to the size of the population (331 individuals), its accuracy and reliability. Rock falls (number or fallen volume per decade) correspond to Y (dependent variables to explain) while the independent variables X (assumed explanatory) are meteorological and marine conditions (external factors influencing the occurrence of rock falls). This work is based solely on the analysis of external factors (marine and continental factors), the only homogeneously quantifiable elements across the sector considered. Obviously, internal factors are also fundamental, but these data (fracturing, karstification, lithostratigraphy) along 37.5 km of coastline are heterogeneous, at different scales, and not always available (Laignel, 2003; Duperret et al., 2004; Mortimore et al., 2004; Costa et al., 2006a; Hoyez, 2008). To choose external variables, many data are available (Meteo-France, SHOM (French Naval Hydrographic and Oceanographic Service)) but variables must: - have a meaning in terms of agents and processes potentially conducive to triggering (thermal variations such as freeze/thaw cycles, hydroclasty, marine roughness):

199 - be non-redundant (determined by a correlation matrix carried out with all factors) in 200 order not to alter statistical analyses (e.g. two variables about rainfall against one for 201 wind give a higher statistical weight for rainfall); 202 - be the most discriminating variable per type of factor (rainfall, temperature, wind, 203 tide). This is observable by the variable contributions from Principal Component 204 Analysis (PCA) performed with all meteorological and marine data (minimum 205 temperatures, maximum temperatures, sum of rainfall, sum of effective rainfall, 206 maximum tide coefficient, and mean tide coefficient). 207 After being selected, the meteorological and marine variables were used for 208 analyses. Some of these are more relevant with the mean value (or sum) or, in 209 contrast, with the extreme value: 210 1) MinTmin (temperature index): minimum daily temperature during a 10-day period 211 (°C) in Dieppe. We are interested in minimum temperatures because it is possible to 212 take into account freeze/thaw periods for a frost-shattered material such as chalk; 213 2) MReff (rainfall index): mean of daily effective rainfall over a 10-day period (mm) in 214 Dieppe. The mean of rainfall that actually affects the cliff is interesting because it 215 informs about runoff and infiltration (with widening of fissures), and thus about aquifer 216 recharge with delay. For these reasons, effective rainfall is chosen (equal to the 217 difference between total rainfall and actual evapotranspiration) and is calculated from 218 the hydric balance. The mean of effective rainfall gives the same information as the 219 sum of effective rainfall because the latter is divided by the ten days of the 10-day 220 period; 221 3) MaxWFmax (wind index): maximum of daily wind force over a 10-day period (m/s) 222 in Dieppe. Wind provides interesting data because it has an influence on swell and

223 thus on marine actions that affect the cliff foot. The maximum wind gives a good 224 indication of swell (linked to wind force); 225 4) MCoef (tide index): mean of tide coefficient over a 10-day period in Dieppe. Tide 226 has an influence on cliff stability. Combined with wind, it gives information about 227 marine roughness and thus instability at the cliff foot. Clearly, the tide coefficient 228 mean over a 10-day period might appear unrepresentative. However, 1) this period is 229 chosen to be consistent with the other factors; 2) despite this time gap, many rock 230 falls are distinguished by spring or neap tides. 231 The time scale chosen is the mobile 10-day period for three reasons: 232 - the frequency of the inventory is weekly, so ten days include this time lag between 233 two surveys; 234 - generally, ten days before a rock fall are often necessary to identify the triggering 235 factor because it may be due to a cumulative effect (e.g. rainfall reaching a threshold 236 with few but continuous rainfalls over many days); 237 - the period of 10 days is often used in meteorological data (e.g. evapotranspiration) 238 because it enables the inertia of the system to be taken into account. 239 The fixed 10-day period scale (used in many meteorological studies) was not used 240 because it is unsuited to the problem. For example, a rock fall identified on the first 241 day of the month is for the 10-day period of 1 to 10, while the conditions responsible 242 for triggering correspond to the previous fixed 10-day period. Thus, the time scale 243 used is a mobile 10-day period (with the day of the date of the rock fall and the 244 previous nine days) for each variable. 245

10

Preparatory analyses: descriptive and bivariate statistics

247 First, the descriptive statistics (quartiles and median) of "factor" and "rock fall" 248 variables (number and fallen volume) were calculated in order to determine some 249 mathematical properties of the data that can influence the choice of analyses to use: 250 asymmetry distributions, outliers, and non-linear relationships between variables. 251 Thus, one of the conditions necessary for the application of linear regression models 252 and appropriate tests (distribution of residuals according to a law close to a centered 253 normal distribution) was checked graphically (Figure 5). 254 The "factor" variables are generally well distributed on either side of the mean, except 255 those indicators describing the intensity of effective rainfall. This shows a skewed 256 distribution, characterized by an over-representation of high outliers. 257 The dependent variable "number of rock falls" is strongly influenced by the presence 258 of 10-day periods with a high number of rock falls (15-20). It therefore highlights an 259 asymmetric distribution, characterized by a large number of outliers whatever the 260 observation scale. For the fallen volume, the median seems much more suited to the 261 application of linear models. These parameters were included in the linear bivariate 262 regression models. 263 Secondly, we fitted simple bivariate linear regression models describing the evolution 264 of variables to explain Y (number and volume of rock falls) as a function of the 265 potential explanatory variables X (external factors), taken one by one. The purpose of 266 this regression is (Cornillon and Matzner-Lober, 2011): 267 - to fit a model to explain Y in terms of X; 268 - to predict the Y values for new values of X. 269 It could be interesting to observe the relationships between, for example, minimum 270 temperatures and median volume of rock falls over a 10-day period (Figure 6). For

each tested relationship, bivariate methods did not appear appropriate because

trigger origins seem complex, with many factors involved. Therefore, other methods need to be used. To identify marine and meteorological conditions conducive or not to rock falls, multivariate analyses are relevant. To identify the triggering factor(s) of rock falls, multivariate analyses cannot define a hierarchy in a combination of agents; this requires an empirical analysis.

Multivariate statistical analyses

In order to answer our first question about marine and meteorological conditions conducive to rock falls, multivariate analyses were performed with the four selected external factors on two populations:

- a first population with 159 10-day periods with rock falls (331) and 109 10-day periods without rock falls was chosen to identify which external factor(s) can discriminate a 10-day period with and without any events. In other words, this is to understand the combination of agents and processes likely or not to trigger rock falls; - a second population with only 10-day periods (159) with rock falls (331) in order to study precisely the rock falls and their characteristics as a function of the marine and meteorological conditions identified as conducive to rock falls. In other words, because there are only 10-day periods with rock falls, it is possible to identify relationships between combinations of factors and characteristics of rock falls (volume and area of departure).

In detail, the multivariate method used was divided into 3 steps from the rock fall inventory (Figure 7):

1) a PCA was performed to visualize the relationships between the variables and the existence of groups of individuals and groups of variables. This structure highlights the agents *a priori* most relevant to sum up meteorological and marine conditions that

characterize individuals (10-day periods). It is also possible to observe the seasonal distribution of variables and individuals; 2) the second step was a Hierarchical Cluster Analysis (HCA). Using the previous results of the PCA, the variables of individuals (10-day periods) are their PCA coordinates on the first three axes. This step differentiates individuals (10-day periods) as a function of their meteorological and marine agents and thus creates a typology (types of meteorological and marine conditions leading to rock falls). This typology is then used to describe the mean characteristics of meteorological and marine conditions leading to rock falls (average analysis); 3) lastly, the third step was the integration of the geomorphological component of rock falls (volume and departure area) within the 10-day periods. It was performed by the  $\chi^2$  test (independence test, p-value coupled with  $\alpha$  level of 5 %<sup>1</sup>) to check whether certain types of meteorological and marine conditions preferentially generate many rock falls. By integrating the type to each rock fall (and no longer to 10-day periods), it is possible to know whether certain conditions trigger specific classes of rock fall volume (V < 200 m<sup>3</sup>, 200 m<sup>3</sup>  $\leq$  V  $\leq$  1,400 m<sup>3</sup>, 1,400 m<sup>3</sup> < V  $\leq$  10,000 m<sup>3</sup>, V > 10,000 m<sup>3</sup>) and/or a specific departure area (cliff foot, middle part of the cliff, cliff top, whole cliff face).

315

316

317

318

319

320

297

298

299

300

301

302

303

304

305

306

307

308

309

310

311

312

313

314

Empirical analysis

This exploratory statistical approach (PCA, HCA and  $\chi^2$  test) provided major elements of knowledge by identifying the meteorological and marine conditions conducive to rock falls but it was impossible to determine the triggering factors. Indeed, in multivariate analyses, each "factor" variable is represented by a single

<sup>&</sup>lt;sup>1</sup> The risk of error is 5 % ( $\alpha$ ).

value that summarizes its behavior over the 10-day period. Thus, the detail of the factor's behavior on each day of the 10 days, which often informs about the triggering factor of the rock fall, remains hidden. To provide this information and answer our second question, we performed an empirical analysis (or human-supervised analysis) on 10-day periods with rock falls (159 cases). In many fields, such as environmental studies, the use of empirical statements is common and recognized (e.g. in the courts). Empirical methods and the experience of practitioners often yield excellent results in the studies of fluvial geomorphology, ecology, hydrology, and natural hazards. We believe that when statistical studies reach their limits (often due to the nature of the data and the assumptions they imply), experience, observation, common sense and intuition remain valid assessment instruments. Moreover, many statistical methods and spatial analyses need to be human-supervised. This empirical analysis (a subjective method of ranking variables) was based on the study of the daily values of sub-aerial and marine factors for each rock fall during 10day periods in order to identify the triggering factor (most active in frequency or intensity) (Figure 8). After this human-supervised analysis, it was possible to identify rock fall characteristics (volume, departure area) preferentially induced by respective factors.

340

341

342

344

321

322

323

324

325

326

327

328

329

330

331

332

333

334

335

336

337

338

339

#### Results

- Multivariate statistical analyses
- 343 PCA, HCA and  $\chi^2$  test for 10-day periods with and without rock falls (first population)

The first analysis focuses on the structure of meteorological and marine conditions

345 (with the four variables mentioned above) over 10-day periods with rock falls (active

individuals in PCA) and without rock falls (supplementary individuals in PCA) (a total of 268 10-day periods) (Figure 9). For this population, axes F1 and F2 of the PCA represent 66 % of the total inertia (and 86 % with F3). The highest variable contribution on F1 (35 % of the total information) is the wind (52 % contribution), which opposes temperatures (contribution up to 30 %). Thus, the first axis characterizes wind and temperatures. where negative values represent calm wind and "warm" weather and the inverse for positive values. On the F2 axis (31 % of the total inertia), the tidal coefficient and effective rainfall are the most influential (37 % and 35 %, respectively) and evolve in the same direction. Thus, positive values of the F2 axis mean spring tide and rainy weather (Figure 9). On the F3 axis (21 %), the tidal coefficient and rainfall are located in opposite directions. Due to this distribution on the PCA, the diversity between 10day periods seems to be explained; first, by wind and temperature and secondly, by rainfall and tidal coefficient. Furthermore, the distribution of 10-day periods with rock falls is on either side of the variables. Therefore, the triggering of rock falls seems to be more the result of a combination of marine and sub-aerial factors than a single factor and occurs in many combinations (Figure 9). Regarding the temporal distribution between 10-day periods with and without rock falls, seasonality appears. Unsurprisingly, winter (rainy and "cold" weather) is the most favorable season for 10-day periods with rock falls while summer (low wind and "warm" temperatures) is the most favorable season for those without rock falls (Figure 9). Another result provided by the PCA is that "wind" and "temperature" factors differentiate 10-day periods with and without rock falls the most.

346

347

348

349

350

351

352

353

354

355

356

357

358

359

360

361

362

363

364

365

366

367

- The HCA and average analyses (histograms in Figure 10) provide a typology of four
- 370 classes of meteorological and marine conditions characterizing 10-day periods (in
- 371 descending order of the number of rock falls) (Figure 10):
- 372 1) Type A is characterized by "cold" (temperature mean under 0°C (histogram at the
- top of Figure 10) thus including freeze/thaw cycles) and dry weather (147 rock falls,
- 374 45 % within 100 10-day periods);
- 375 2) Type B is characterized by "warm" weather and calm wind (64 rock falls or 19 %
- within 75 10-day periods);
- 377 3) Type C is characterized by high rainfall and high wind (generating swell) (63 rock
- 378 falls, 19 % within 38 10-day periods);
- 379 4) Type D is characterized by spring tide and "warm" weather (the remaining 17 %
- 380 within 55 10-day periods).
- Lastly, the  $\chi^2$  test shows a statistically significant link between types of conditions and
- 382 10-day periods with and without rock falls (observed distance (13.0) higher than the
- 383 critical level (7.8) with a p-value of 0.005, so less than  $\alpha$ ):
- Type C "high rainfall and high wind" (generating swell) has an over-representation
- of 10-day periods with rock falls (79 % of 10-day periods have rock falls, 21 % of the
- remaining 10-day periods do not have rock falls);
- 387 Type B "warm weather and calm wind" (corresponding to summer weather) has an
- over-representation of 10-day periods without rock falls (52 %), so these are
- 389 conditions conducive to cliff stability.
- To go further in the analysis, a second  $\chi^2$  test was performed to measure
- independence between the number of rock falls (0, 1, 2 and more) and types of
- meteorological and marine conditions. There is a statistically significant link between

393 these variables (observed distance (18.3) higher than the critical level (12.6) with a p-394 value of 0.006, so less than  $\alpha$ ): 395 - Type C "high rainfall and high wind" has an over-representation of 10-day periods 396 with a minimum of 2 rock falls (47 %, against 21 % for 0 rock fall and 32 % for 1 rock 397 fall) whereas these are under-represented in Type B "warm weather and calm wind" (19 %). 398 399 To summarize, in the four types of meteorological and marine conditions identified as 400 responsible (or not) for triggering rock falls, Type C ("high rainfall and high wind") is 401 the best rock fall trigger over a 10-day period, especially in frequency (2 rock falls or 402 more) whereas Type B, characterized by "warm weather and calm wind", is the least 403 favorable for the occurrence of rock falls (but they might occur). 404 405 PCA, HCA and  $\chi^2$  test for 10-day periods only with rock falls (second population) 406 For this second population (159 10-day periods), relationships between the variables 407 and the existence of groups of variables are the same as the previous PCA (because 408 10-day periods without rock falls in the PCA were supplementary individuals). 409 HCA identifies four types of meteorological and marine conditions for 10-day periods. 410 It is interesting to note that, although "warm weather and calm wind" are identified 411 with the first population as conditions not conducive to rock falls, a group of 412 individuals (10-day periods with rock falls) is always present around these variables 413 (in fourth position, with 56 rock falls within 33 10-day periods) (Figure 11). Thus, rock 414 falls are possible in these summer conditions, when people may stay all day on the 415 beach, near the cliff foot. A  $\chi^2$  test was performed on rock falls (no longer 10-day

periods) to highlight whether some conditions trigger specific types of rock fall

- 417 (volume and departure area). The results are as follows (in descending order of
- 418 number of rock falls, Figure 11):
- 419 1) Type 1 conditions are "cold" (with a mean temperature histogram under 0°C, so
- 420 numerous freeze/thaw cycles) and dry weather (147 rock falls, 44 % within 65 10-day
- 421 periods). For these conditions, the  $\chi^2$  test highlights that rock falls with volumes less
- 422 than 200 m<sup>3</sup> are over-represented whereas rock falls with volumes greater than
- 423 10,000 m<sup>3</sup> are under-represented (49 % for V < 200 m<sup>3</sup>, against 27 % for 200 m<sup>3</sup>  $\leq$  V
- 424  $\leq$  1,400 m<sup>3</sup>, 21 % for 1,400 m<sup>3</sup> < V  $\leq$  10,000 m<sup>3</sup>, 2 % for V > 10,000 m<sup>3</sup> and 1 % for
- 425 unknown volume);
- 426 2) Type 2, characterized by high rainfall and high wind (70 rock falls or 21 % within
- 427 31 10-day periods), has an over-representation of rock falls with a departure area of
- 428 the whole cliff face (56 % of cases, against 21 % for cliff foot, 13 % for foot and
- 429 middle part of the cliff, 3 % for middle part of the cliff and 7 % for cliff top). This type
- of conditions has a high proportion of rock falls greater than 10,000 m<sup>3</sup> with 9 % of
- 431 cases (whereas these massive rock falls represent only 5 % of the total number of
- 432 rock falls in the ESTRAN inventory);
- 433 3) Type 3 is characterized by a spring tide and "warm" weather (57 rock falls, or 17 %
- 434 within 30 10-day periods). This type has an over-representation of rock falls with a
- 435 fallen volume between 1,400 and 10,000 m<sup>3</sup> (35 % of cases, against 32 % for V <
- 436 200 m<sup>3</sup>, 24 % for 200 m<sup>3</sup>  $\leq$  V  $\leq$  1,400 m<sup>3</sup>, 9 % for V > 10,000 m<sup>3</sup>);
- 437 4) Type 4, the rarest, is characterized by "warm" weather and calm wind conditions
- 438 (the remaining 17 % within 33 10-day periods). This type has an over-representation
- of (1) cliff foot rock falls (45 % of cases, against 28 % for the whole cliff face, 23 % for
- foot and middle part of the cliff, 2 % for cliff top and 2 % for unknown volume), and (2)
- rock fall volume between 200 and 1,400 m $^3$  (45 % of cases, against 29 % for V < 200

volume). Therefore, from this second population, we have complementary information about conditions likely to trigger rock falls. "Cold (freezing) and dry weather" conditions (Type 1) generate many rock falls, usually small in volume (less than 200 m³) without a specific departure area. In contrast, "high wind and high rainfall" conditions (Type 2) are more conducive to massive rock falls (more than 10,000 m³) from the whole cliff face. "Warm weather and calm wind" (Type 4) generate rock falls of modest size (200 to 1,400 m³) mainly coming from the cliff foot. More surprisingly, "spring tide and warm weather" conditions (Type 3) seem to be responsible for large rock falls (from 1,400 to 10,000 m³).

 $m^3$ , 20 % for 1,400  $m^3$  <  $V \le 10,000 \, m^3$ , 3 %  $V > 10,000 \, m^3$  and 3 % for unknown

#### Empirical analysis

subjective method of variable ranking, triggering factors are identified:

1) effective rainfall seems to be involved in 56 % of the number of rock falls and 72 % of the total fallen volume. Of these 185 rock falls probably due to rainfall, 79 cases come from the whole cliff face (43 %) and 34 % are between 200 and 1,400 m³. As a proportion of the 331 rock falls in the inventory, the characteristics of the 185 rock falls triggered by rainfall are mostly: (1) between 200 and 1,400 m³ or over 10,000 m³ (64 % and 56 % of the ESTRAN database, respectively), (2) from the whole cliff face (59 %) (Figure 12);

2) freeze/thaw cycles seem to be involved in 19 % of the number of rock falls and 10 % of the fallen volume. These rock falls are mostly from the whole cliff face (22 cases, 34 %) or the cliff foot (16 cases, 25 %) and smaller than 200 m³ (45 %).

The empirical analysis confirms and adds to the multivariate analyses. With this

467 However, as a proportion of the 331 rock falls in the database, the characteristics of 468 rock falls triggered by freeze/thaw cycles are mainly: (1) smaller than 200 m<sup>3</sup> (23 % 469 of the inventory), (2) from the cliff top (50 %) or cliff foot (16 %) (area where chalk is 470 preferentially saturated in water by runoff or infiltration) (Figure 12); 471 3) the next factor is marine roughness, a combination of spring tide (coefficient more 472 than or equal to 85<sup>2</sup>) and high wind (more than 8 m/s) that generates marine actions 473 at the cliff foot, with 15 % of the number of rock falls and 9 % of the fallen volume. 474 These rock falls are mostly from the whole cliff face (20 cases, 31 %) or 475 unsurprisingly from cliff foot (17 cases, 27 %) and smaller than 200 m<sup>3</sup> (42 %). As a 476 proportion of the 331 rock falls in the inventory, the characteristics of rock falls 477 triggered by marine roughness are mainly: (1) smaller than 200 m<sup>3</sup> (17 %), (2) from 478 the foot/middle part of the cliff (20 %) or the foot cliff (17 %) (Figure 12). This factor 479 needs complementary studies, as its influence is difficult to quantify due to rock falls 480 of small volume being quickly removed during a turbulent marine period (and that 481 could be missed by the ESTRAN organization). 482 Nevertheless, these identified triggering factors rarely act alone. Only 49 rock falls 483 (15 %) seem to be triggered by a single factor. These few cases occurred with only 484 one predominant parameter (while other marine and sub-aerial factors were of low 485 frequency or intensity). The vast majority of rock falls is due to a combination of the 486 triggering factors mentioned above (very common in mid-latitudes). This is the case 487 during unsettled weather (rainfall, marine roughness), which occurs approximately 488 2/3 of the year along the Normandy coast (sometimes with freeze/thaw cycles). 489 Consequently, it is more appropriate to identify three types of weather leading to the 490 triggering of rock falls: periods of high effective rainfall over a 10-day period, periods

<sup>&</sup>lt;sup>2</sup> With this value of 85, we are sure that the tide affects the foot of the cliff along the coastline studied.

of freeze/thaw cycles, types of weather characterized by atmospheric disturbance involving heavy rainfall and high wind (thus high waves).

493

494

495

496

497

498

499

500

501

502

503

504

505

506

507

508

509

510

511

512

513

514

515

492

491

#### **Discussion**

The ESTRAN database is unique and significant because of the weekly frequency of surveys (which track small and large rock falls), the precise dating of rock falls, the length of the coastline surveyed (37.5 km) and the duration of 7 years. There are other inventories of this coastal section (e.g. Duperret et al., 2004; Dewez et al., 2013) but a comparison of results is difficult. Their spatial and temporal representativeness are limited. For example, the inventory of Dewez et al. (2013) was carried out on 750 m and for 2.5 years. Dating of rock falls remains limited: of the order of the month to the season for Duperret et al. (2004). However, the types of favorable conditions for triggering rock falls can sometimes tend towards the same conclusions for massive rock falls, which are often linked to specific meteorological conditions. For example, Duperret et al. (2004) show two types of triggering meteorological conditions: heavy rainfall and sudden rainfall during dry periods. However, these authors point out that their results are questionable because of the small number of cases (55 including only 4 precisely dated). This participation in the scientific debate about determining the factors and processes likely to trigger rock falls is based on an exploratory statistical approach and an empirical analysis. These two methods are complementary. Whereas the statistical one provides the types of combination of meteorological and marine factors conducive or not to rock falls, the empirical one determines the triggering factor for each rock fall. Each of these approaches gives information about the modalities of retreat (rock fall departure area and volume). The statistical analyses underline that

"cold and dry weather" and "high rainfall and high wind" are the best conditions for triggering rock falls. The empirical analysis identifies effective rainfall as the most important triggering factor. It seems to generate the majority of both rock falls (56 %) and fallen volume (72 %). It is followed by freeze/thaw cycles and marine roughness. However, this subjective ranking of the triggering factors should be weighted in terms of the frequency of their occurrence per year. Whereas rainfall occurs about 130 days a year, freeze/thaw cycles are much less frequent: about 26 times a year over the seven years observed (with 183 registered cases). Based on these results, freeze/thaw cycles seem to be more "efficient" than rainfall, because when they occur, rock falls are nearly always observed. Contrary to what is often stated in the scientific literature, marine roughness does not appear to be only a transport and debris removal agent. This is the third factor explaining the triggering of rock falls. Unsurprisingly, marine roughness seems particularly effective at the cliff base, involving rock falls of small volume (but can cause mass movements affecting the whole cliff face). This result is consistent with fine quantitative results obtained by terrestrial laser surveys carried out along this coastline (Letortu et al., in press, accepted). However, in many cases, the triggering of rock falls seems to be the result of a combination of factors. The most common combination (56 rock falls out of 331) is marine roughness (linked to wind speed and during spring high tide) and high effective rainfall (instantaneous or cumulative). This corresponds to the very common types of disturbed weather in mid-latitudes, which dominate for approximately 2/3 of the year. Low pressure with two fronts crosses the English Channel and generates high rainfall in Normandy and a marine roughness conducive to triggering rock falls. If this type of weather is followed by a period of freeze/thaw cycles, cliff weathering and

516

517

518

519

520

521

522

523

524

525

526

527

528

529

530

531

532

533

534

535

536

537

538

539

cliff retreat will reach their peak. However, is this due to the inaccuracy of the rock fall dating (often around a 10-day period)? To go beyond this limit, we collected 74 correctly-dated rock falls from the local press and carried out the statistical and empirical analyses. Even if rock fall is dated to the day, the process of identifying the triggering factor(s) remains complex. High rainfall plays a role but is also combined with the other factors. This work confirms that identifying triggering factors is complex because the evolution of environments is not linear and not limited to direct causal relationships. The "purge" effect, relays or combinations of processes and/or hysteresis phenomena (which we tried to integrate by using a mobile 10-day period) complicate the analysis of relationships. In addition, cliffs inherit past environmental changes (Trenhaile, 2002) i.e. they are influenced by the size and location of previous rock falls (the "purge" effect). The purge effect means that a cliff face has been recently affected by a fall, called here a "purge". The cliff face is "new", so it must not be affected by another fall soon because it has to be "prepared" by preparatory processes for a certain time to be ready to fall. So, cliff retreat is a process with a "memory" (Lee et al., 2001). Whatever our understanding of current processes and their impacts, the evolution of the cliff is partially based on the long-term degradation of the rock (Griggs and Trenhaile, 1994; Swenson et al., 2006; Lim et al., 2011). Moreover, the dichotomy between weathering factors and triggering ones is not always obvious. This determination is complex because the relationships between the factors are non-linear. Several reasons may explain this: 1) the transition state between weathering and triggering factors can be inappropriate

541

542

543

544

545

546

547

548

549

550

551

552

553

554

555

556

557

558

559

560

561

562

563

564

for common statistical methods (e.g. the presence of discretized data). It could be

565 more relevant to study these factors using fuzzy logic. This is why we tried to 566 incorporate temporal flexibility by working over mobile 10-day periods; 567 2) the triggering factor may be random. If the material is ready to fall, any factors, 568 whatever their intensity, are likely to generate a rock fall; 569 3) conversely, the factors identified as triggering (under certain conditions) may occur 570 without generating any departures of material due to a previous "purge" effect. The 571 material expected to fall is already on the ground. In other words, factors that can be 572 triggering are repeated but with no rock falls. They participate, like others, in lowering 573 the stability margin of the cliff to a critical value. In this case, the agents and 574 processes are just preparatory. 575 There is also the question of the time required to determine the triggering factors. Is 576 the 10-day period sufficient? Should the analysis period be increased when 577 meteorological and marine conditions are not dynamic in the 9 days preceding the 578 rock fall? This is the case of the collapse of 08/26/2001 (correctly-dated rock fall from 579 the local press) where no increase in the intensity or frequency of triggering agents 580 occurred (a priori) during the 10-day period, while on 08/02/2001 rainfall greater than 581 26 mm was recorded. Are these phenomena related? Are there other explanatory 582 factors (e.g. micro-earthquakes), processes such as hydroclasty, or decaying? Are 583 there hysteresis phenomena over intervals of several weeks? Lengthening the 584 observation period of agents and processes responsible for triggering rock falls would 585 make determining the contribution of each factor more complex, because of their 586 multiple occurrences over this longer time period. 587 This statistical approach and the empirical one are based on the analysis of external 588 conditions, but it is also essential to take internal factors into account (rock 589 mechanics, micro- and macro-cracking). Whereas, at first sight, the Upper Normandy

coastal chalk cliffs appear homogeneous in lithology, on a fine scale there are many differences (Mortimore et al., 2004; Lasseur, 2007). These limitations of this work must be taken into account in future developments to improve our results. For example, it would be interesting to focus on a type of chalk lithology (e.g. Santonian stage) with a specific departure area of rock falls. This will become possible when the database has been enriched with new inventory years. Furthermore, data of meteorological and marine parameters need to be on a finer temporal and spatial scale. Thus, determining the agents and processes responsible for triggering rock falls remains difficult due to the limited number of cases observed (especially precisely dated) and a lack of knowledge of the internal parameters in chalk cliffs and of meteorological and marine parameters on fine temporal and spatial scales.

#### **Conclusions**

geomorphological features) in Upper Normandy using multivariate statistical analyses (PCA, HCA) and empirical analysis. The results, deduced directly from these analyses, highlight innovative information about the conditions leading to rock falls and triggering factors:

1) the frequency of rock falls is linked to meteorological and marine conditions: unsurprisingly, high rainfall and high wind (winter weather with marine roughness) are the most favorable conditions for 10-day periods with many rock falls (2 or more) while warm weather and calm wind (summer weather) are the most favorable conditions for 10-day periods without rock falls (but rock falls might occur).

2) the volume of rock falls depends on triggering factors: freeze/thaw cycles are mainly responsible for debris falls (volume less than 200 m³) whereas high rainfall

The originality of this paper is the analysis of a large inventory of 331 rock falls (date,

triggers preferentially more massive rock falls (between 200 and 1,400 m³ or more than 10,000 m³). Marine roughness seems to trigger rock falls of small volume (less than 200 m³) in particular.

3) the cliff face location of rock falls is due to triggering factors: high rainfall mainly triggers rock falls coming from the whole cliff face while freeze/thaw cycles trigger rock falls coming from the top and foot of the cliff face. Unsurprisingly, marine roughness seems particularly effective at the cliff base.

Some difficulties remain in quantifying the proportion of marine and sub-aerial factors because of problems of combinations of factors (very common in mid-latitudes) or the relay of factors (85 % of 331 cases) and probably hysteresis phenomena. In view of these various findings, it is still presumptuous to predict the location and time of triggering of rock falls. However, the approaches adopted (statistical and naturalistic) and the observations made in this study are from an original inventory, and constitute a real starting point for the prediction and prevention of the hazard of rock falls in Upper Normandy.

#### Acknowledgements

The authors are grateful to the members of the ESTRAN organization and of Meteo-France who provided rock fall inventory and meteorological data. We thank reviewers for their detailed reviews of the manuscript and their helpful comments

#### References

Adams PN, Anderson RS, Revenaugh J. 2002. Microseismic measurement of waveenergy delivery to a rocky coast. *Geology* **30**(10): 895-898.

Augris C, Clabaut P, Costa S, Gourmelon F, Latteux B. 2004. Evolution morpho-639 640 sédimentaire du domaine littoral et marin de la Seine-Maritime. Editions Ifremer, 641 Bilans & Prospectives: Paris. 642 Bernatchez P, Dubois JMM. 2008. Seasonal quantification of coastal processes and 643 cliff erosion on fine sediment shorelines in a cold temperate climate, north shore of 644 the St. Lawrence maritime estuary, Québec. Journal of Coastal Research 24 (1A): 645 169-180. 646 Brossard J, Duperret A. 2004. Coastal chalk cliff erosion: experimental investigation 647 on the role of marine factors. In Coastal Chalk Cliff Instability, Mortimore RN, 648 Duperret A. (eds). Geological Society: London; 109-120. 649 Castedo R, Murphy W, Lawrence J, Paredes C. 2012. A new process-response 650 model of soft rock cliffs. Geomorphology 177: 128-143. 651 Cornillon PA, Matzner-Løber E. 2011. Régression avec R. Springer: Paris. 652 Costa S. 1997. Dynamique littorale et risques naturels: L'impact des aménagements, 653 des variations du niveau marin et des modifications climatiques entre la Baie de 654 Seine et la Baie de Somme. PhD thesis, Department of Geography, University of 655 Paris I. 656 Costa S. 2005. Falaises à recul rapide et plages de galets: de la guantification des 657 dynamiques d'un système complexe à la caractérisation des risques induits. 658 Accreditation to supervise research, Department of Geography, University of Brest. 659 Costa S, Laignel B, Hauchard E, Delahaye D. 2006a. Facteurs de répartition des

entonnoirs de dissolution dans les craies du littoral du Nord-Ouest du Bassin de

Paris. Zeitschrift für Geomorphologie **50**(1): 95-116.

660

- 662 Costa S, Hénaff A, Lageat Y. 2006b. The gravel beaches of North-West France and
- their contribution to the dynamic of the coastal cliff-shore platform system. Zeitschrift
- 664 für Geomorphology, Suppl.-Vol. **144**: 199-194.
- Costa S, Delahaye D, Freire-Diaz S, Davidson R, Di-Nocera LE. 2004. Quantification
- of the Normandy and Picardy chalk cliff retreat by photogrammetric analysis. In
- 667 Coastal Chalk Cliff Instability, Mortimore RN, Duperret A. (eds). Geological Society:
- 668 London; 139-148.
- Dewez TJB, Rohmer J, Regard V, Cnudde C. 2013. Probabilistic coastal cliff collapse
- 670 hazard from repeated terrestrial laser surveys: case study from Mesnil Val
- 671 (Normandy, northern France). Journal of Coastal Research, Special Issue 65: 702-
- 672 707.
- 673 Dornbusch U, Robinson DA, Moses CA, Williams RBG. 2008. Temporal and spatial
- variations of chalk cliff retreat in East Sussex, 1873 to 2001. *Marine Geology* **249**:
- 675 271-282.
- Duperret A, Genter A, Mortimore RN, Delacourt B, Pomerai MR. 2002. Coastal rock
- 677 cliff erosion by collapse at Puys, France: the role of impervious marl seams within
- 678 chalk of NW Europe. Journal of Coastal Research 18(1): 52-61.
- Duperret A, Taibi S, Mortimore RN, Daigneault M. 2005. Effect of groundwater and
- sea weathering on the strength of chalk rock from unstable coastal cliffs of NW
- France. Engineering Geology 78: 321-343.
- Duperret A, Genter A, Martinez A, Mortimore RN. 2004. Coastal chalk cliff instability
- 683 in NW France: role of lithology, fracture pattern and rainfall. In Coastal Chalk Cliff
- 684 Instability, Mortimore RN, Duperret A. (eds). Geological Society: London; 33-55.
- 685 Emery KO, Kuhn GG. 1982. Sea cliffs: their processes, profiles, and classification.
- 686 Geological Society of America Bulletin **93**: 644-654.

- 687 Fournier A, Allard M. 1992. Periglacial shoreline erosion of a rocky coast: George
- River Estuary, northern Quebec. *Journal of Coastal Research* **8**(4): 926-942.
- 689 Genter A, Duperret A, Martinez A, Mortimore RN, Vila JL. 2004. Multiscale fracture
- analysis along the French chalk coastline for investigating erosion by cliff collapse. In
- 691 Coastal Chalk Cliff Instability, Mortimore RN, Duperret A. (eds). Geological Society:
- 692 London; 57-74.
- 693 Griggs GB, Trenhaile AS. 1994. Coastal cliff and platform. In Coastal Evolution,
- 694 Carter RWG, Woodroffe CD (eds). Cambridge University Press: Cambridge, UK;
- 695 425-450.
- 696 Hall K. 1999. The role of thermal stress fatigue in the breakdown of rock in cold
- 697 regions. Geomorphology 31: 47-63.
- 698 Hénaff A, Lageat Y, Costa S, Plessis E. 2002a. Modalités du recul des falaises du
- 699 Pays de Caux. Actes du symposium international: from experts opinion to modelling,
- 700 a tribute to Pr. Jean-Claude Flageollet, European Center of Geomorphological
- 701 hazards, Strasbourg: 225-233.
- 702 Hénaff A, Lageat Y, Costa S, Plessis E. 2002b. Le recul des falaises crayeuses du
- 703 Pays de Caux: détermination des processus d'érosion et quantification des rythmes
- 704 d'évolution. Géomorphologie: relief, processus, environnement 2: 107-118.
- 705 Hoek E, Bray JW. 1977. Rock slope engineering. The Institution of Mining and
- 706 Metallurgy: London.
- 707 Hoyez B. 2008. Falaises du Pays de Caux. Lithostratigraphie des craies turono-
- 708 campaniennes. Publications des Universités de Rouen et du Havre: Mont-Saint-
- 709 Aignan.

- 710 Hutchinson JN. 1972. Field and laboratory studies of a rockfall in Upper Chalk cliffs at
- 711 Joss Bay, Isle of Thanet. In *Proceedings of the Roscoe Memorial Symposium*,
- 712 Cambridge, UK; 692-706.
- 713 Jerwood LC, Robinson DA, Williams RBG. 1990a. Frost and salt weathering of chalk
- 714 I. Earth Surface Processes and Landforms **15**: 611-624.
- 715 Jerwood LC, Robinson DA, Williams RBG. 1990b. Frost and salt weathering of chalk
- 716 II. Earth Surface Processes and Landforms **15**: 699-708.
- 717 King CAM. 1972. Beaches and Coasts. Arnold: London.
- 718 Lahousse P, Pierre G. 2003. The retreat of chalk cliffs at Cap Blanc-Nez (France):
- autopsy of an erosional crisis. *Journal of Coastal Research* **19**(2): 431-440.
- 720 Laignel B. 2003. Caractérisation et dynamique erosive de systèmes
- 721 géomorphologiques continentaux sur substrat crayeux. Exemple de l'Ouest du
- 722 Bassin de Paris dans le contexte nord-ouest européen. Accreditation to supervise
- research, Department of Geology, University of Rouen.
- 724 Lasseur E. 2007. La Craie du Bassin de Paris (Cénomanien-Campanien, Crétacé
- 725 supérieur). Sédimentologie de faciès, stratigraphie séquentielle et géométrie 3D. PhD
- thesis, Department of Geosciences, University of Rennes 1.
- Lee EM. 2008. Coastal cliff behavior: observations in the relationship between beach
- 728 levels and recession rates. *Geomorphology* **101**(4): 558-571.
- Lee EM, Moore R, McInnes RG. 1998. Assessment of the probability of landslide
- 730 reactivation: Isle of Wight Undercliff, UK. In Engineering Geology: The View from the
- 731 Pacific Rim, Moore D, Hungr O (eds). Balkema: Amsterdam.
- 732 Letavernier G. 1984. La gélivité des roches calcaires. Relations avec la morphologie
- 733 *du milieu poreux*. Thèse de doctorat, Department of Geography, University of Caen.

- 734 Letortu P, Costa S, Bensaid A, Cador JM, Quénol H. 2014a. Vitesses et modalités de
- 735 recul des falaises crayeuses de Haute-Normandie (France): méthodologie et
- variabilité du recul. Géomorphologie: relief, processus, environnement 2: 133-144.
- 737 Letortu P, Costa S, Bonnet E. 2014b. Spatial analysis of coastal chalk cliff falls in
- 738 Upper Normandy (France) from Veules-les-Roses to Le Treport (2002-2009). Revue
- 739 Internationale de Géomatique **24**(3): 335-354.
- Letortu P, Costa S, Delacourt C, Maquaire O, Augereau E, Davidon R, Suanez S., in
- 741 press. Taux d'ablation des falaises crayeuses haut-normandes: l'apport du scanner
- 742 laser terrestre. Géomorphologie, relief, processus et environnement.
- Letortu P, Costa S, Delacourt C, Maquaire O, Augereau E, Davidon R, Suanez S,
- Nabucet J., accepted. Retreat rates, modalities and agents responsible for erosion
- 745 along the coastal chalk cliffs of Upper Normandy: the contribution of terrestrial laser
- 746 scanning. Geomorphology.
- Lim M, Rosser NJ, Petley DN, Keen M. 2011. Quantifying the controls and influence
- of tide and wave impacts on coastal rock cliff erosion. Journal of Coastal Research
- 749 **27**(1): 46-56.
- 750 May VJ, Heeps C. 1985. The nature and rates of change on chalk coastlines.
- 751 Zeitschrift für Geomorphologie, Supplementband **57**: 81-94.
- 752 McGreal WS, 1979. Marine erosion of glacial sediments from a low-energy cliffline
- environment near Kilkeel, Northern Ireland. *Marine Geology* **32**(1): 89-103.
- 754 Mortimore RN, Duperret A. 2004. Coastal Chalk Cliff Instability, Geological Society:
- 755 London.
- 756 Mortimore RN, Stone KJ, Lawrence J, Duperret A. 2004. Chalk physical properties
- and cliff instability. In Coastal Chalk Cliff Instability, Mortimore RN, Duperret A. (eds).
- 758 Geological Society: London; 75-88.

- Naylor LA, Stephenson WJ, Trenhaile AS. 2010. Rock coast geomorphology: Recent
- advances and future research directions. *Geomorphology* **114**(1-2): 3-11.
- Pierre G, Lahousse P. 2006. The role of groundwater in cliff instability: an example at
- 762 Cape Blanc-Nez (Pas-de-Calais, France). Earth Surface Process and Landforms 31:
- 763 31-45.
- Robinson LA. 1977. Marine erosive processes at the cliff foot. *Marine Geology* 23:
- 765 257-271.
- Robinson LA, Jerwood LC. 1987. Frost and salt weathering of chalk shore platforms
- near Brighton, Sussex, UK. Transactions of the Institute of British Geographers 12:
- 768 217-226.
- Senfaute G, Duperret A, Lawrence JA. 2009. Micro-seismic precursory cracks prior to
- rock-fall on coastal chalk cliffs: a case study at Mesnil-Val, Normandie, NW France.
- 771 Natural Hazards and Earth System Sciences 9: 1625-1641.
- 772 Stephenson WJ. 2000. Shore platforms: A neglected coastal feature? *Progress in*
- 773 Physical Geography **24**: 311-327.
- 774 Sunamura TA. 1982. Predictive model for wave-induced cliff erosion, with application
- to Pacific coast of Japan. *Journal of Geology* **90**: 167-178.
- 776 Sunamura TA. 1992. *Geomorphology of rocky coasts*. John Wiley: New York.
- 777 Swenson MJ, Wu CH, Edil TB, Mickelson DM. 2006. Bluff recession rates and wave
- impact along the Wisconsin coast of Lake Superior. *Journal of Great Lakes Research*
- 779 **32**: 512-530.
- 780 Trenhaile AS. 1987. *The geomorphology of rock coast*. Oxford University Press:
- 781 Oxford.
- 782 Trenhaile AS. 2002. Rock coasts, with particular emphasis on shore platforms.
- 783 Geomorphology **48**: 7-22.

- 784 Trenhaile AS. 2009. Modelling the erosion of cohesive clay coasts. Coastal
- 785 Engineering **56**: 59-72.
- 786 Trenhaile AS. 2010. Modelling cohesive clay coast erosion and response to climate
- 787 change. Marine Geology 277: 11-20.
- 788 Trenhaile AS, Kanyaya JI. 2007. The role of wave erosion on liping and horizontal
- 789 shore platforms in macro- and mesotidal environments. Journal of Coastal Research
- 790 **23**(2): 298-309.
- 791 Varnes DJ. 1978. Slope movement types and processes. In Special Report
- 792 176: Landslides: Analysis and Control, Schuster RL, Krizek RJ. (eds).
- 793 Transportation and Road Research Board, National Academy of Science:
- 794 Washington; 11-33.
- 795 Woodroffe CD. 2002. Coasts: form, process and evolution. Cambridge University
- 796 Press: Cambridge.

- 797 Young AP, Flick RE, Gutierrez R, Guza RT. 2009. Comparison of short-term seacliff
- retreat measurement methods in Del Mar, California. *Geomorphology* **112**: 318-323.
- 799 Young AP, Guza RT, Adams PN, O'Reilly WC, Flick RE. 2012. Cross-shore decay of
- 800 cliff top ground motions driven by local ocean swell and infragravity waves. Journal of
- 801 Geophysical Research 117, C06029.

Dieppe (1971-2000)	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec	Total
Maximum temperatures (°C)	7.2	7.7	9.9	11.5	14.9	17.2	19.5	20.3	18.5	15.2	10.7	8.2	13.4
Mean temperatures (°C)	4.9	5.0	7.0	8.5	11.8	14.2	16.4	16.9	15.0	12.0	8.1	5.9	10.5
Minimum temperatures (°C)	2.5	2.3	4.1	5.4	8.6	11.3	13.4	13.5	11.5	8.8	5.5	3.6	7.5
Mean number of days with	7.7	8.3	3.8	1.3	_	_	_	_		0.3	3.4	7	31.8
minimum temperatures <= 0°C	/./	8.3	3.6	1.5	_	_	_	_		0.3	3.4		31.6
Mean number of days with	2.2	1.4	0.3	-	-	-	-	-	-	-	0.2	0.8	4.9
minimum temperatures <= -5°C	2.2												
Mean number of days with	0.4	0.2	_	_	_	_	_	_	_	_		0.0	0.6
minimum temperatures <= -10°C	0.4	0.2			_					_		0.0	0.0
Maximum daily height of rainfall	29.5	38.5	32.0	27.0	52.5	40.6	58.5	100.8	56.1	77.5	46.1	42.0	-
Height of rainfall (mean in mm)	64.0	51.8	57.6	56.7	59.2	58.6	53.3	50.6	80.4	90.8	93.8	87.4	-
Mean number of days with	11.7	9.9	11.9	10.3	10.3	9.6	8.3	8.5	10.8	11.4	13.8	13.2	129.6
rainfall >= 1 mm	11.7	9.9	11.5	10.5	10.5	9.0	6.5	6.5	10.8	11.4	13.6	13.2	125.0
Mean number of days with	an number of days with 4.6		3.8	4.2	4.1	4.0	3.1	3.5	5.6	5.8	6.8	6.3	55.3
rainfall >= 5 mm	4.0	3.5	3.0	4.2	4.1	4.0	3.1	3.3	5.0	5.6	0.0	0.5	33.3
Mean number of days with	1.7	1.1	1.1	1.3	1.5	1.5	1.6	1.2	2.5	2.9	3.4	2.8	22.5
rainfall >= 10 mm	1.7	1.1	1.1	1.5	1.5	1.5	1.0	1.2	2.5	2.9	3.4	2.0	22.5
Mean number of days with	13.6	10.4	11.3	9.0	6.7	5.6	5.1	_	7.9	10.5	10.9	11.8	102.8
gust >= 16 m/s		10.4	11.5	5.0	0.7	3.0	5.1		,.5	10.5	10.5	11.0	102.0
Mean number of days with	1.2	1.4	0.7	0.3	0.1	_	0.2	_	0.2	0.7	0.5	1.4	6.7
gust >= 28 m/s	1.2	1.4	0.7	0.5	0.1	_	0.2	_	0.2		rce of dat		

803 Source of data: Meteo-France

## Table I: Meteorological data in Dieppe (Meteo-France, 1971-2000)

Rock fall date	Rock fall number	Town	Coordinate x (RGF93-L93)	Coordinate y (RGF93-L93)	Volume (m³)	Departure area of rock fall	
02/06/02	1	Varengeville	556,948	6,981,584	1,372	cliff foot	
02/18/02	2	Varengeville	554,542	6,981,917	942	whole cliff face	
02/18/02	3	Varengeville	554,661	6,981,894	20,100	whole cliff face	
02/23/02	4	Dieppe	559,928	6,982,139	92	whole cliff face	
03/04/02	5	Sainte-Marguerite	553,865	6,981,975	92	whole cliff face	
03/04/02	6	Sainte-Marguerite	554,030	6,981,961	976	whole cliff face	
03/25/02	7	Varengeville	555,028	6,981,826	1,508	whole cliff face	
:	:	 	;	!	!	:	
12/14/09	325	Varengeville	555,344	6,981,751	93	cliff foot	
12/22/09	326	Sainte-Marguerite	553,626	6,982,019	202	whole cliff face	
12/22/09	327	Sainte-Marguerite	553,683	6,982,013	1,581	whole cliff face	
12/22/09	328	Varengeville	554,798	6,981,867	448	whole cliff face	
12/22/09	329	Varengeville	555,137	6,981,789	478	whole cliff face	
12/22/09	330	Varengeville	555,157	6,981,785	259	cliff foot	
12/23/09	331	Dieppe	563,951	6,983,619	31	cliff foot	

Table II: Extract of enriched inventory (volume, departure area of rock fall) used for the next analyses

809 Figure 1: Presentation of the studied area (from Veules-les-Roses to Le Treport) 810 showing a cross-section with lithostratigraphy 811 Figure 2: Example of a rock fall (then a slide of tertiary strata) on the west side of 812 Dieppe that threatens urbanization (18/12/2012, Les informations Dieppoises) 813 Figure 3: Rock fall inventory (A: index card made by the ESTRAN organization for 814 rock fall n°123; B: spatial distribution of rock falls; C: temporal distribution of rock 815 falls) 816 Figure 4: Rock fall distribution (number and frequency) as a function of volume (A: 817 total distribution of the number of rock falls as a function of fallen volume; B: 818 thresholds observed to distinguish classes of volume as a function of rock fall 819 frequency) 820 Figure 5: Descriptive statistics of variables "factor" and "rock fall" over a 10-day 821 period 822 Figure 6: Bivariate regression between minimal temperatures and median fallen 823 volume over a 10-day period 824 Figure 7: Method of multivariate statistical analyses (with 3 steps) 825 Figure 8: Method of empirical analysis with the example of the rock fall of 02/08/02 826 Figure 9: PCA for 10-day periods with and without rock falls (distribution of variables 827 in insert) and seasonal distribution Figure 10: Results of the typology (4 types deduced by the HCA) and their 828 829 meteorological and marine conditions per type (average analysis with histograms on 830 the right), superimposed on the PCA (previously performed on the 268 10-day 831 periods with and without rock falls) 832 Figure 11: Results of the typology (4 types deduced by the HCA) and their 833 meteorological and marine conditions per type (average analysis with histograms on the right), superimposed on the PCA (previously performed on the 159 10-day
periods with rock falls)

Figure 12: Results from the empirical analysis with the first three triggering factors
(histograms of number of rock falls, sector diagrams of distribution of rock falls per
class of volume, and geomorphological specificities of rock fall per triggering factor)