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Is the Mediterranean Sea Circulation in a Steady State

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Abstract: Most global ocean models are based on the assumption of a “steady state” ocean. Here, we investigate the validation of this hypothesis for the anthropized Mediterranean Sea. In order to do so, we calculated the mixing coefficients of the water masses detected in this sea via an optimum multiparameter analysis referred to as the MIX approach, using data from the BOUM (2008) and MedSeA (2013) cruises. The comparison of the mixing coefficients of each water mass, between 2008 and 2013, indicates that some of their proportions have significantly changed. Surface water mass proportions did not change significantly ($\Delta 0.05-0.1$), while intermediate and deep water mass mixing coefficients of both Eastern and Western basins were significantly modified ($\sim \Delta 0.35$). This study clearly shows that the Mediterranean seawater is not in a “steady state”.

Keywords: Mixing Coefficient, MIX Approach, Water Masses, Seawater Circulation, Mediterranean Sea, Climate Change

1. Introduction

The Mediterranean Sea is a mid-latitude semi-enclosed sea. It witnesses two well-defined seasons, wet-cold winter and dry-warm summer, with short period of transition between them [1] [2]. An excess of evaporation over precipitation can be quantified by calculating the annual mean budget “evaporation minus precipitation” (E-P) over the whole Mediterranean Sea ($350-750 \text{ mm yr}^{-1}$) [1]. Thus, Atlantic Water (AW) inflows at the surface and Mediterranean Water outflows along the bottom. The Atlantic water entering into Gibraltar in the surface layer, after travelling to the easternmost Levantine Sub-basin, is transformed into one of the saltiest seawater masses through air-sea heat and moisture fluxes. This relatively cold and salty water, which crosses the entire Mediterranean Sea in the opposite direction of the surface Atlantic Water, finally exits from the Strait of Gibraltar at the mid-depths. In wintertime, a significant

negative heat budget [1] [3] causes a buoyancy loss, initiating deep and/or intermediate dense water formation. These events are frequent over the shelf areas and in the offshore regions, both in the Western and the Eastern basins [4] [5]. The Mediterranean Sea also exchanges water with the Black Sea through the Turkish Strait System (the Dardanelles, the Marmara Sea and the Bosphorus Strait) and receives significant amounts of freshwater from river discharge. In addition, this sea is always provided by groundwater discharges and sewage, which are likely, an important source of freshwater, nutrients, trace metals, alkalinity and other elements to the Mediterranean System [6] [7]. The Mediterranean Sea is composed mainly of two nearly equal size basins, the Western and the Eastern ones, connected by the relatively shallow Strait of Sicily (sill depth $\sim 500 \text{ m}$). In each basin a number of sub-basins are characterized by different water masses circulations, chemical and biological features and rough bottom topography. Note that we reserve

the term “sea” for the Mediterranean Sea and “basin” for the Eastern and Western parts of it; any smaller entity is a sub-basin.

The Mediterranean Sea is a particular system, characterized by a complex thermohaline, wind, and water flux-driven multi-scale circulation with interactive variability [8]. This land-locked sea is exporting warm and salty intermediate waters to the North Atlantic Ocean by the narrow and shallow Strait of Gibraltar [width ~ 13 km ; sill depth ~ 300 m] affecting the global thermohaline conveyor belt [9] [10]. Through pathways to the Atlantic Polar regions or through indirect mixing processes, the salty Mediterranean water preconditions the deep convection cells of the Polar Atlantic. There the North Atlantic Deep Water is formed and successively spreads throughout the world ocean constituting the core of the global thermohaline circulation [11] [12]. Thus, the salty water of Mediterranean origin may affect water formation processes and variability and even the stability of the global thermohaline equilibrium state. Hence, the study of the Mediterranean water masses (their formation, spreading, mixing, and impact on general circulation) is essential for a better understanding of the ocean circulation and variability. Moreover, the existence of a thermohaline cell (the Eastern Mediterranean “Conveyor Belt”), multiple scales of motion defining the general circulation (basin/sub-basin/mesoscale), and deep water mass formation processes, make this sea a “test basin” for general circulation studies.

Aiming to study the hydrographic situation in the Mediterranean Sea, many researchers have used an optimum multiparameter analysis ([13] in the Adriatic Sub-basin, and [14] in the Eastern Mediterranean Sub-basin; [15] in the Western Mediterranean Sub-basin; [16] in the entire Mediterranean Sea). However, a study of the evolution of water mass mixing coefficients in this semi-enclosed sea is needed to better understand the hydrographic system in this sea.

Reference [17] introduced the MIX approach based upon an optimum multiparameter mixing analysis. The use of this approach is particularly recommended for regional studies where the distribution of water masses can be clearly defined (this is the case of the Mediterranean Sea). In the present paper, our main objective is to evaluate the variation (if any) of the water mass circulation in the Mediterranean Sea based on their mixing coefficients calculated via the MIX approach using data collected in 2008 and 2013.

2. Methodology

2.1. Study Area

The BOUM cruise (Biogeochemistry from the Oligotrophic to the Ultra oligotrophic Mediterranean Sea [18]) was conducted during summer 2008, from 20 June to 22 July, on board the R/V L'Atalante [19]. It consists of a longitudinal transect (more than 3000 km long from the Levantine Sub-basin to the Northwestern Mediterranean Sea) of 27 short-term stations and 3 long-term stations referred as A, B, and C (Fig.1).

The MedSeA (Mediterranean Sea Acidification in a Changing Climate; [20] [21]) cruise occurred during spring 2013, on board of the Spanish R/V Angeles Alvariño, from 2 May to 2 June. The full cruise track (more than 8000 km long) consisted of two longitudinal legs where 23 stations along the Mediterranean Sea were sampled throughout the water column. During the first leg, samples were collected from Atlantic waters off Cadiz, Spain to the Levantine Sub-basin in the Eastern Mediterranean basin (3879 km long, 15 stations, 279 sampled points, maximum sampled depth = 3720 m). The second leg was conducted in the Northern part of the Mediterranean from Heraklion, Crete, Greece in the Eastern Mediterranean basin to Barcelona, Spain in the North Western Mediterranean basin (3232.5 km long, 8 stations, 183 sampled points, maximum sampled depth = 3000 m ; Fig.2).

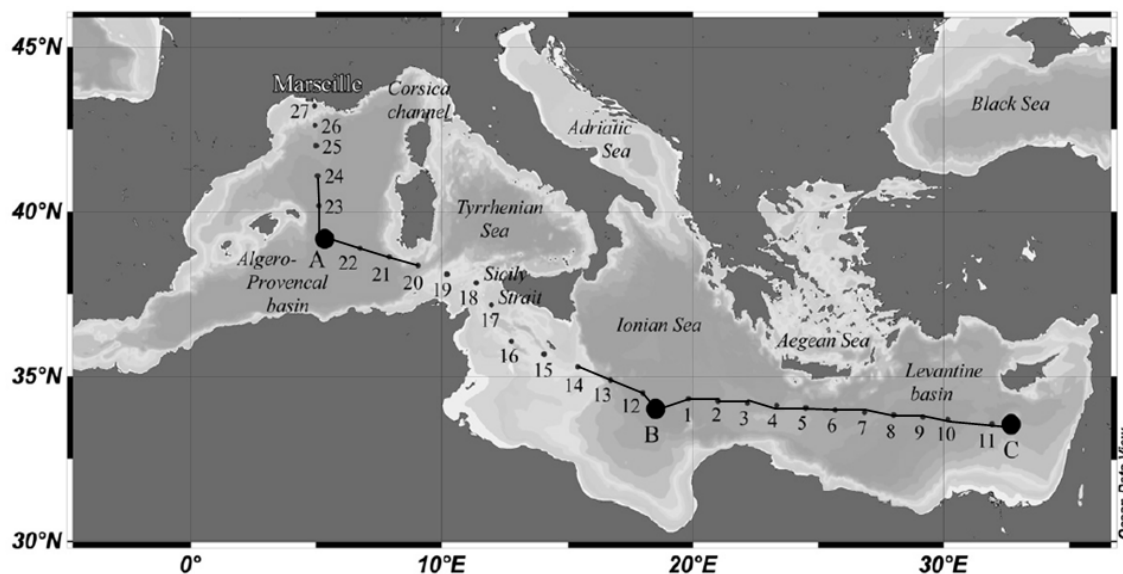


Figure 1. Map of the 2008 BOUM cruise in the Mediterranean Sea. Short-term stations are indicated by numbers (from 1 to 27); the three long-term stations are referred as A, B, and C (Touratier et al., 2012). The thick black lines referred to the sections used in this study.

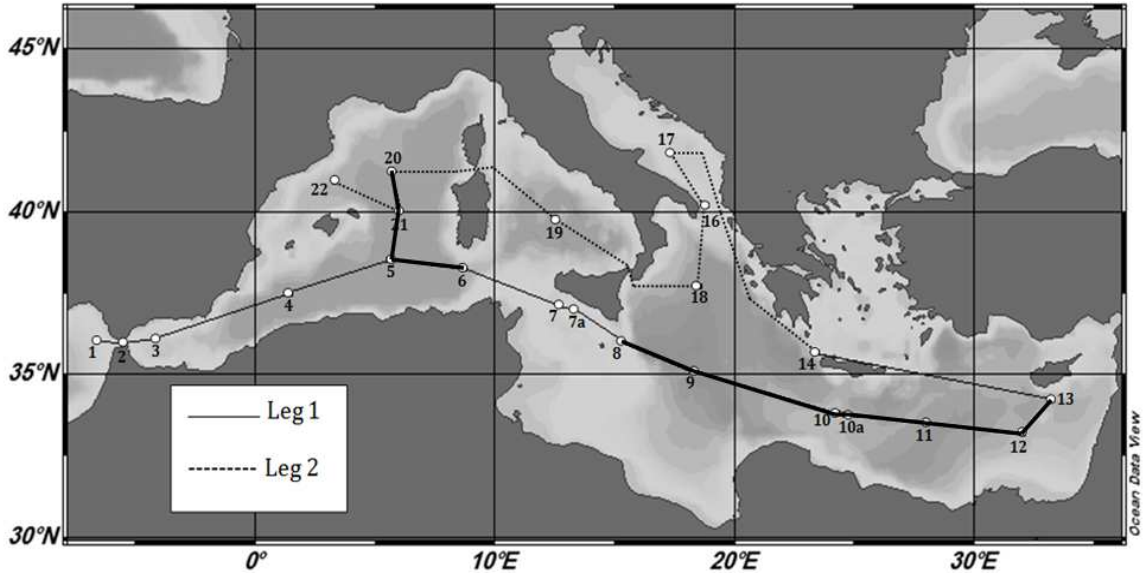


Figure 2. Map of the 2013 MedSeA cruise in the Mediterranean Sea. The numbers from 1 to 22 correspond to the sampled stations. The thick black lines referred to the sections used in this study.

2.2. Sampling and Measurements

In the present study, we use the following properties: potential temperature (θ ; °C), salinity (S), dissolved oxygen (O_2 ; $\mu\text{mol kg}^{-1}$), nitrates (NO_3 ; $\mu\text{mol kg}^{-1}$), and phosphates (PO_4 ; $\mu\text{mol kg}^{-1}$).

For the 2008 BOUM cruise, the profiles for θ , S, and O_2 were obtained using a Sea-Bird Electronics 911 PLUS CTD system. Each CTD cast was associated with a carousel of 24 Niskin bottles to collect seawater samples used to perform the analysis of the other chemical and biological properties. Concerning the nutrients (NO_3 and PO_4), the description of the methods used for the analysis are explained by [22] and [23].

For the 2013 MedSeA cruise, hydrologic properties [S and T (°C)] were measured *in situ* using a Sea-Bird Electronics CTD system (SBE 911plus) associated with a General Oceanic rosette sampler, equipped with twenty four 12 L Niskin bottles. The precision of measurements is ± 0.001 °C for T and ± 0.0003 for S. Water samples for dissolved oxygen determination were collected in calibrated BOD 60 ml bottles. Oxygen concentrations were measured using a Winkler iodometric titration [24] with a Mettler-Toledo. DL-21 potentiometric titrator with a Pt ring redox electrode for the determination of the equivalence point [25]. The analytical precision and accuracy are ± 1.5 $\mu\text{mol kg}^{-1}$. For the nitrates (NO_3) and phosphates (PO_4), water samples were pre-filtered over glass fiber filters (Whatman GF/F) with a pore size of 0.7 microns, immediately after sampling. Then, they were stored at -20°C in polyethylene bottles until their analysis. The samples were thawed and analyzed colorimetrically with a Bran + Luebbe 3 autoanalyzer according to [26], at OGS laboratory in Trieste, Italy.

2.3. Modeling: The MIX Approach

Aiming to quantify the contribution of the different water sources to the collected data, we used the MIX approach

developed by [17]. This approach uses a multi-parameter analysis. Optimum multiparameter methods [27] [28] are based on the assumption that observed water properties at a hydrographic station are the result of mixing among two or more “source waters”. They have been used to evaluate water mass properties in both the Southern Ocean [29] [30] and other regions [28] [31].

In the present study, four conservative tracers (S, θ , NO, and PO) are used in the MIX approach. Reference [32] defined the conservative tracers NO and PO as follows:

$$\text{NO} = O_2 + R_{\text{ON}}\text{NO}_3 \quad (1)$$

$$\text{PO} = O_2 + R_{\text{OP}}\text{PO}_4 \quad (2)$$

where NO and PO are composite tracers of the non conservative tracers O_2 (dissolved oxygen concentration in $\mu\text{mol kg}^{-1}$) and NO_3 (nitrate concentration in $\mu\text{mol kg}^{-1}$), and PO_4 (phosphate concentration in $\mu\text{mol kg}^{-1}$), respectively. According to the equation of [33] which describes the average photosynthesis and aerobic respiration in the interior of the ocean, the composite tracers NO and PO were built using the fact that the consumption of oxygen is balanced by the production of nutrients during the processes of respiration and decomposition. The two constants R_{ON} and R_{OP} are the ratios of the stoichiometric coefficients (ψ) involved in the Redfield equation ($R_{\text{ON}} = \psi O_2 / \psi \text{NO}_3$; $R_{\text{OP}} = \psi O_2 / \psi \text{PO}_4$).

The general conservation equation for a conservative tracer (Ω) is given by :

$$\Omega = \sum_{j=1}^n K_j \Omega_j \quad (3)$$

where K_j represents the contribution (also called “mixing coefficient”) of a water source j , n is the number of water sources in the system, and Ω_j is the value of the conservative tracer for the water source j .

For each seawater sample, where the conservative tracers Ω are either measured (S and θ) or calculated (NO and PO ; see equations 1 and 2), the contributions K_j are estimated after resolving the system of equations with the following constraints :

$$\sum_{j=1}^n K_j = 1 \text{ (mass conservation)} \quad (4)$$

$$\forall_j, 0 \leq K_j \leq 1 \quad (5)$$

It is a constrained linear least-squares problem, with one equality constraint and n inequality constraints. To solve it, we use a medium-scale optimization algorithm (with MATLAB) similar to that described by [34]. The difference in range between tracers leads to an ill-conditioned problem: the linear equation solution involves the inversion of the matrix containing the water sources characteristics, but this matrix is really badly scaled. So, prior to the inversion, rows and columns of the matrix are normalized.

The stability of the results (linked to the correct placement of sources) is then tested by adding thirty perturbations to the values of Ω and the Ω_j which provide at each point a mean

solution for each K_j [17]. These perturbations are independent Gaussian random variables with zero mean and given standard deviation. We entered the standard deviations for each tracer with a given value for the measurement points and another (larger) value for sources. These thirty Gaussian perturbations allowed us not only to test the stability of the results but also to take into account the differences between the errors of measurement based on the tracers.

2.4. Determination of the Water Sources

The Appropriate definition of water source properties is crucial to achieving physically meaningful results from the MIX approach. The different water sources j were typically identified using the conservative tracers S and θ , their corresponding θ/S diagrams (Fig. 3 and 6) and referring to the literature. In order to lower the number of water sources, the Mediterranean Sea was tested as two independent systems : the Western basin ($5^\circ \text{ E} \leq \text{longitude} \leq 9.5^\circ \text{ E}$) and the Eastern basin ($\text{longitude} \geq 15^\circ \text{ E}$). Consequently, the relatively shallow stations located in the Sicily Strait between 9.5° E and 15° E were ignored when the MIX approach was applied.

Table 1. Physical and chemical properties of water sources used by the MIX approach in the Western Mediterranean basin for BOUM 2008 and MedSeA 2013.

Water Sources	Depth	T (°C)	S	θ (°C)	NO ₃ ($\mu\text{mol kg}^{-1}$)	PO ₄ ($\mu\text{mol kg}^{-1}$)	O ₂ ($\mu\text{mol kg}^{-1}$)
BOUM 2008							
AW	50	17.21	37.269	17.20	0.00	0.002	240.8
EOW	50	15.12	38.163	15.12	0.05	0.032	259.0
LIW	496	14.15	38.853	14.07	5.69	0.192	181.8
TDW	2839	13.31	38.472	12.86	8.61	0.391	196.2
WMDW	1483	13.08	38.459	12.86	8.78	0.383	190.0
MedSeA 2013							
AW	50	16.17	36.607	16.16	0.63	0.058	234.8
EOW	50	15.12	38.163	15.12	0.05	0.032	259.0
LIW	496	14.15	38.853	14.07	5.69	0.192	181.8
TDW	2839	13.31	38.472	12.86	8.61	0.391	196.2
WMDW	1483	13.08	38.459	12.86	8.78	0.383	190.0

AW, Atlantic Water;

EOW, Effluent Outflow Water;

LIW, Levantine Intermediate Water;

TDW, Tyrrhenian Deep Water;

WMDW, Western Mediterranean Deep Water.

For the Western basin, we used five water sources ($n = 5$). Whereas, six ones were used for the Eastern basin ($n = 6$). The physical and chemical properties of these water sources, for both BOUM 2008 and MedSeA 2013 cruises, are shown in Tables 1 and 2 for the Western and the Eastern Mediterranean basins respectively. These water sources were identified based on the data of the two cruises and on the availability of the input parameters (S, θ , O₂, NO₃, PO₄) manipulated by MIX approach to calculate the mixing coefficients. The main water

sources are indicated on the corresponding θ/S diagrams of the Western and Eastern basins (Fig. 3 and 6). When the water sources of the MedSeA cruise have similar θ/S characteristics compared to the BOUM ones, we decided to take into consideration the characteristics of the BOUM water masses in our mixing coefficients calculation to maintain the same ratios previously calculated for the Mediterranean Sea based on the 2008 BOUM cruise.

Table 2. Physical and chemical properties of water sources used by the MIX approach in the Eastern Mediterranean basin for BOUM 2008 and MedSeA 2013.

Water Sources	Depth	T (°C)	S	θ (°C)	NO ₃ ($\mu\text{mol kg}^{-1}$)	PO ₄ ($\mu\text{mol kg}^{-1}$)	O ₂ ($\mu\text{mol kg}^{-1}$)
BOUM 2008							
MAW	73	15.73	37.618	15.72	0.00	0.001	239.7
EOW	174	15.59	38.436	15.56	2.37	0.102	205.9
LIW	248	16.25	39.163	16.21	0.96	0.013	220.6
CIW	173	15.17	39.045	15.14	3.34	0.090	204.9

Water Sources	Depth	T (°C)	S	θ (°C)	NO ₃ ($\mu\text{mol kg}^{-1}$)	PO ₄ ($\mu\text{mol kg}^{-1}$)	O ₂ ($\mu\text{mol kg}^{-1}$)
EMDW-Adr	2965	13.86	38.710	13.38	4.91	0.142	195.4
EMDW-Ag	2222	14.00	38.789	13.64	4.84	0.167	189.3
MedSeA 2013							
MAW	50	17.72	38.873	17.72	0.01	0.029	241.3
EOW	174	15.59	38.436	15.56	2.37	0.102	205.9
LIW	248	16.25	39.163	16.21	0.96	0.013	220.6
CIW	173	15.17	39.045	15.14	3.34	0.090	204.9
EMDW _{Adr}	950	13.18	38.702	13.04	1.65	0.087	219.7
EMDW _{Ag}	2222	14.00	38.789	13.64	4.83	0.167	189.3

MAW, Modified Atlantic Water ;

EOW, Effluent Outflow Water ;

LIW, Levantine Intermediate Water ;

CIW, Cretan Intermediate Water ;

EMDW_{Adr}, Eastern Mediterranean Water-Adriatic origin ;

EMDW_{Ag}, Eastern Mediterranean Deep Water-Aegean origin.

2.5. Determination of the Redfield Ratios R_{ON} and R_{OP}

Typical values given by [33] are $R_{ON} = 8.6$ and $R_{OP} = 138$ (molar ratios). However, these ratios cannot be representative for the Mediterranean waters since [33] never used observations from this region. In this paper, we use the R_{ON} and R_{OP} ratios previously estimated, specifically for the Mediterranean Sea by [35] using the 2008 BOUM database. The best results for the regressions between the nutrients and O₂ were obtained after considering two different layers for both the Western and the Eastern basins : the surface/intermediate layer (from 50 up to 750 m), and the deep layer (from 750 m to the bottom). We hypothesize that these layers are related to the traditional Mediterranean circulation scheme which is described by two cells [36] : the well ventilated surface cell which contains water masses like MAW, WIW, and LIW ; and the deep cells which contains the WMDW and TDW in the Western basin, and the EMDW_{Adr} and EMDW_{Ag} in the Eastern basin.

3. Results and Discussion

In order to assess the goodness of the solutions and to check possible errors due to the water-source description and to uncertainties associated with the measurements, the residual vector was calculated for each of the four conservative tracers. This vector corresponds to the

difference between the predicted and the measured data. For BOUM 2008 and MedSeA 2013, the MIX method performs well, producing unbiased small residuals (always less than 10%) for all measures.

Hereafter, we discuss the evolution of the main water sources mixing coefficients for the Western and Eastern Mediterranean basins.

3.1. Western Basin

3.1.1. Atlantic Water, AW

The mixing coefficients of the surface water masses did not change significantly since 2008 till 2013 (Fig.5). The Mediterranean Sea has an active water exchange with the Northern Atlantic through the Strait of Gibraltar. The surface Atlantic flow entering the Mediterranean Sea does not only fill the water deficit of 1 m per year, but it also replaces the Mediterranean deep outflow, which represents a loss of 20 m of water per year for the whole Mediterranean Sea [37]. As a consequence of the excess of evaporation over precipitation ($\sim 0.62 - 1.16 \text{ m year}^{-1}$) [38], heating and various physical phenomena (gyres, eddies,...), the proportion of the AW decreases progressively because its characteristics change while propagating Eastward to be referred as Modified Atlantic Water (MAW). This trend was comparable during the two oceanographic surveys (BOUM 2008 and MedSeA 2013).

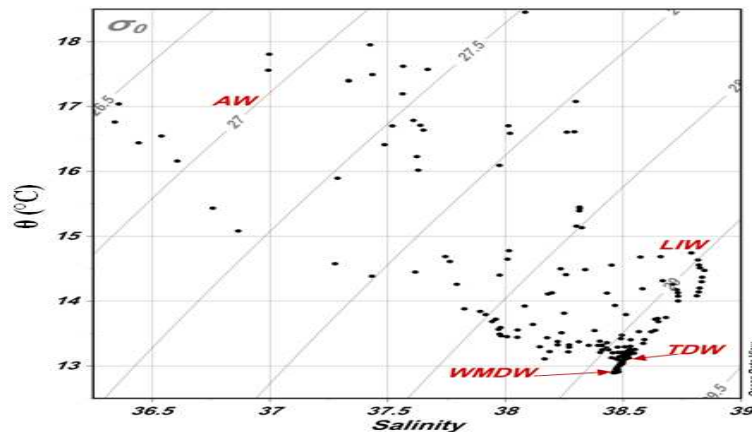


Figure 3. θ/S diagram illustrating the main water masses detected in the Western basin during the MedSeA cruise in May 2013 (AW: Atlantic Water, LIW: Levantine Intermediate Water, TDW: Tyrrhenian Deep Water, WMDW: Western Mediterranean Deep Water).

3.1.2. Levantine Intermediate Water, LIW

The amount of this water mass has increased during the study period (2008-2013). For example, its proportion increased from 0.55 in 2008 to 0.9 in 2013 within the intermediate layers (~ 300 m) of the Sardinia Strait (Fig.5). The LIW contributes predominately to the non-returning efflux, mixed with both EMDW and WMDW in the Strait of Gibraltar, and to the outflow into the Atlantic Ocean [2] [8]

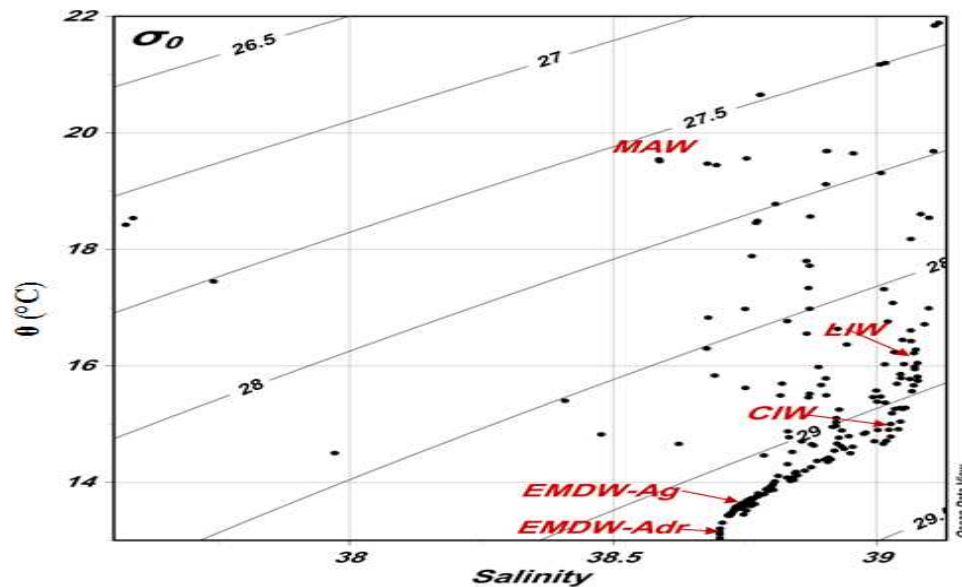


Figure 4. θ/S diagram illustrating the main water masses detected in the Eastern basin during the MedSeA cruise in May 2013 (MAW: Modified Atlantic Water, LIW: Levantine Intermediate Water, CIW: Cretan Intermediate Water, EMDW-Ag: Eastern Mediterranean Deep Water originated from the Aegean Sub-basin, EMDW-Adf: Eastern Mediterranean Deep Water originated from the Adriatic Sub-basin).

3.1.3. Tyrrhenian Deep Water, TDW

The proportion of this water mass has clearly decreased below 1000 m in the Western basin. Its amount during 2013 became less abundant in the Southern sub-basins compared to the one recorded for 2008 (Fig.5). This water, which is a product of the mixing between Eastern (LIW and EMDW) and western waters (WMDW), fills the Tyrrhenian Sub-basin down to the bottom [44]. Thus, this decrease could be attributed to the modification in the EMDW formation, since the Adriatic Sub-basin seems to retrieve its role as a main contributor in the deep water mass formation in the Eastern basin instead of the Aegean during the EMT (Eastern Mediterranean Transient ; data of the MedSeA cruise 2013). However, [45] argue that the TDW might result from a dense water formation process occurring within the Tyrrhenian itself, in the East of the Bonifacio Strait. This means that the decrease in the mixing proportion could also be attributed to climatic changes affecting temperature and salinity, thus the deep water formation.

Nevertheless, the mixing coefficients proportions of the TDW have slightly increased from 0.05 in 2008 to 0.2 in 2013, exclusively in the North of the Western basin between 0 and 500 m. This fact could be firstly related to the harsh winter convective mixing in this wind-driven area [46]. Secondly, this smooth proportions increase could be also connected to

[39]. Our results show recent changes in the Mediterranean outflow. These increasing proportions could be caused by the excessive evaporation and the decrease in precipitation and freshwater supplies to the Eastern Mediterranean Sea [40] which can be direct consequences of the unequivocal global warming [41]. Decreasing trends of river discharges in the Eastern Mediterranean Sea [42] has been observed, as in the Adriatic Sub-basin [43].

the mixing caused by heavy freshwater inputs from the Northern rivers during May (time of the snow melting) [47].

3.1.4. Western Mediterranean Deep Water, WMDW

The mixing coefficients schemes of this water mass (Fig.5) show that its amount decreased slightly below 1000 m, between 2008 and 2013, particularly in the South of the Western basin. The WMDW remains well distributed in the Western basin with higher proportions in the North of this basin, where it is originated (Fig.5). The observed decrease could be due to the significant warming and salinification, widely mentioned in the literature [48] [49] [50] [51] [52]. However, at a station in the south of the Tyrrhenian Sub-basin sampled in September 1999 during the PROSOPE cruise [53], the total alkalinity in deep water was higher than in deep water at the Dyfamed site [54]. If it is confirmed that the Tyrrhenian deep water contributes to the deep water in the Liguro-Provençal Sub-basin, this observation could explain the modifications in the WMDW proportions which could also be attributed to the changes in the TDW linked to the Eastern deep water formation. Based on 23 years eddy-permitting reanalysis, [55] argued that the largest water mass formation event of the past 23 years occurred in the Western Mediterranean basin in 2005-2006. This event was presumably preconditioned by the EMT which modified the characteristics of the LIW crossing the Sicily Strait.

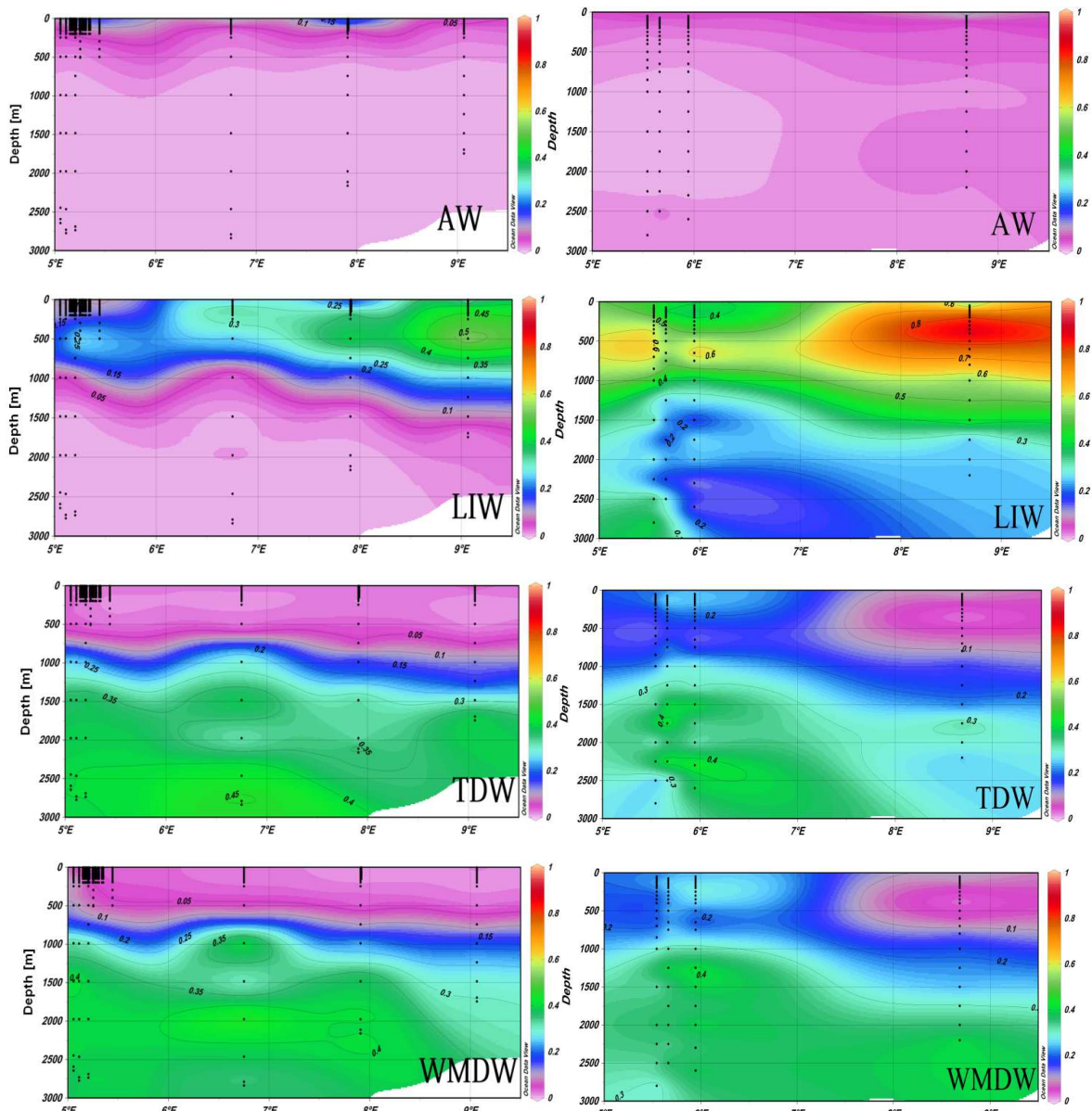
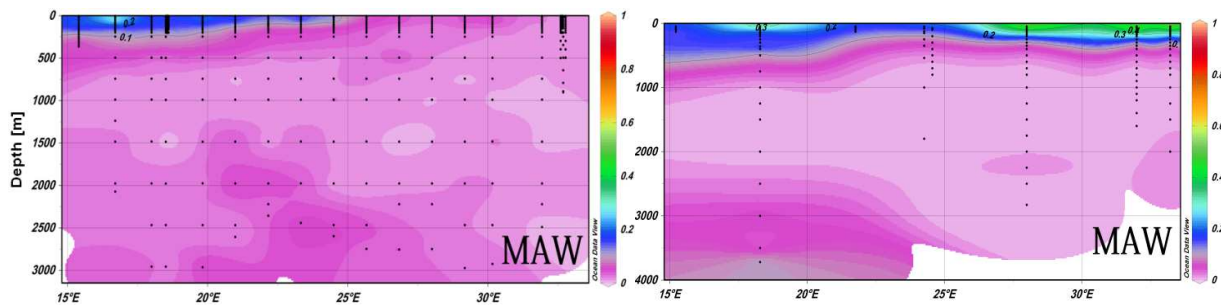


Figure 5. Profiles of the mixing coefficients of water masses in the Western Mediterranean basin during the 2013 MedSea cruise (right column) and the 2008 BOUM cruise (left column)

3.2. Eastern Basin

However, similarly to the TDW, a slight rise of the mixing coefficients has been registered within the layer above 1000 m in the Northern part of the Western basin. This boost could

also be attached to the vertical mixing related to the tough meteorological winter conditions in the Gulf of Lions, as well as in Balearic and Liguro-Provençal Sub-basins [46] [56].



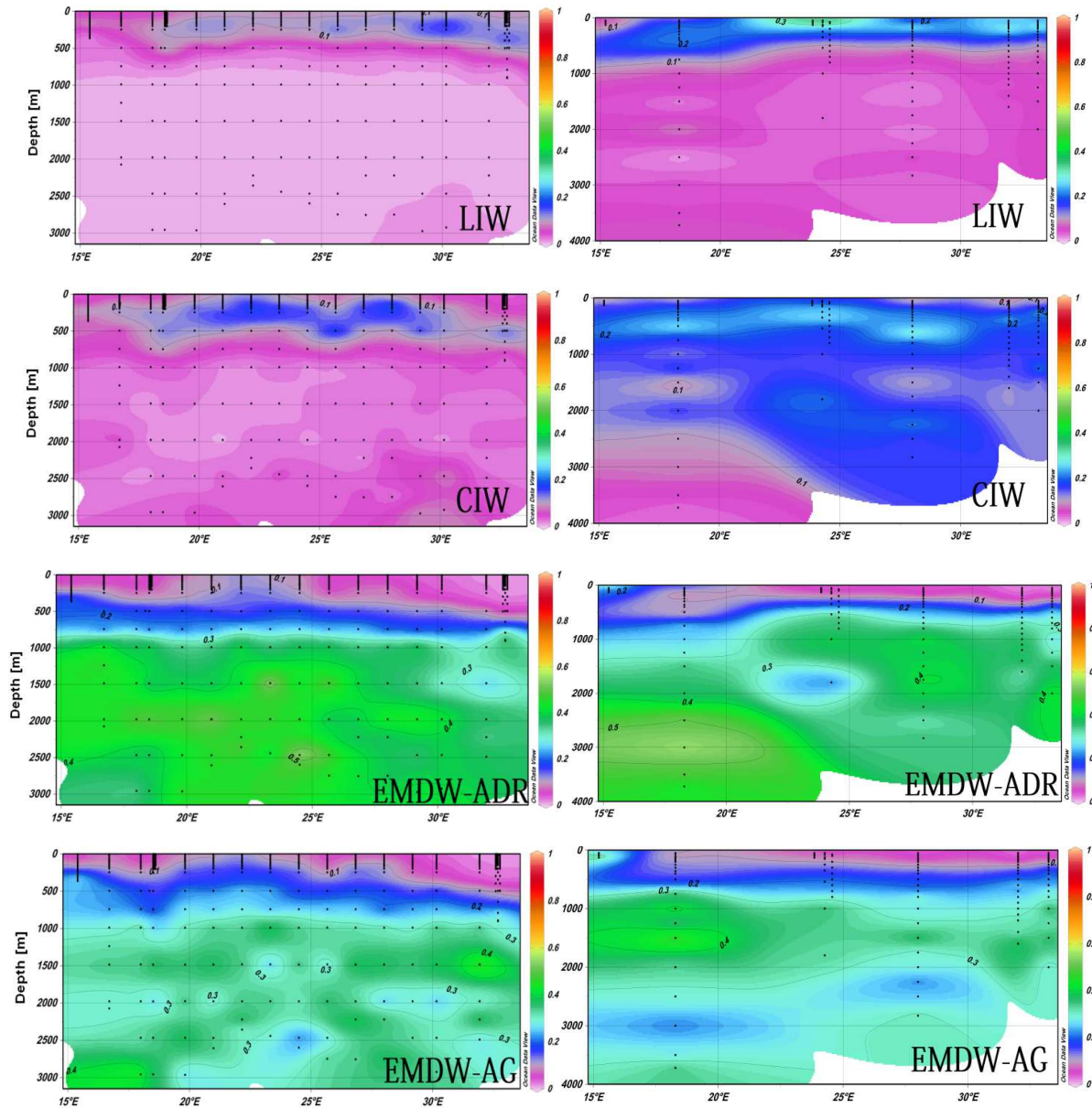


Figure 6. Profiles of the mixing coefficients of water masses in the Eastern Mediterranean basin during the 2013 MedSea cruise (right column) and the 2008 BOUM cruise (left column).

3.2.1. Modified Atlantic Water, MAW

The proportion of the MAW increased from 2008 to 2013. During 2013, it is clear that its amount increases while propagating Eastward from the Sicily Strait (Fig.6). This water mass is initially made as a result of the mixing of comparatively fresh Atlantic water ($S < 36.5$) flowing via the Strait of Gibraltar into the Mediterranean Sea with the surface waters of the Alboran Sub-basin [57]. The incoming MAW is continuously modified by interactions with the atmosphere and mixing with older surface waters and with the waters underneath. All along its course, it is seasonally warmed or cooled, but overall its salt content increases and it becomes denser [58]. Thus, its increasing proportion could be attributed to the excessive evaporation trend in the context of global warming occurring in the Mediterranean Sea [59] [60]. Satellite observations from 1985–2006 indicate that in the last

two decades the temperature in the upper layer of the Mediterranean Sea has been increasing at an average (\pm SD) rate of $0.03 \pm 0.008^\circ\text{C yr}^{-1}$ for the Western basin and $0.05 \pm 0.009^\circ\text{C yr}^{-1}$ for the Eastern basin [60]. Nevertheless, these changes could also be due to the modifications in circulation patterns leading to blocking situations during the EMT (i.e. gyres, eddies, ...), to the Adriatic-Ionian Bimodal Oscillating System (BiOS) [61] and/or to variations in the fresher water of Black Sea origin input through the Strait of Dardanelles [8].

3.2.2. Levantine Intermediate Water, LIW

The amount of this water mass has increased during the study period (2008–2013; Fig.6). Its proportion increased from 0.27 in 2008 to 0.37 in 2013 within the intermediate layers (~ 175 m) in the South of Cyprus. The increasing proportion could be related to the occurrence of high amount of MAW in this area which becomes denser. Thus, this latter sinks more

quickly to participate in the formation of the LIW. This fact could be caused by the excessive evaporation and the decrease in precipitation and freshwater supplies to the Mediterranean Sea. This can be direct consequences of the global warming, well described in literature [62] [63] [64].

3.2.3. Cretan Intermediate Water, CIW

The proportion of this intermediate water mass has increased in the entire Eastern basin, particularly in the South of Crete where it has the highest proportions (~ 0.35 at ~ 170 m; Fig. 6). This fact is due to changes in circulation patterns leading to blocking situations concerning the MAW and the LIW, and to variations in the fresher water of Black Sea origin input through the Strait of Dardanelles [8] [42]. These results could also be attributed to specific atmospheric conditions which created large buoyancy fluxes from the Aegean Sub-basin, similarly to the ones mentioned by [65] during winters 1991/1992 and 1992/1993 (the "enhanced EMT winters"), intensifying intermediate and deep water production, although with modified characteristics [66].

3.2.4. Eastern Mediterranean Deep Water Originated in the Adriatic Sub-Basin, $EMDW_{Adr}$

During the study period, it is clear that the mixing coefficients of the $EMDW_{Adr}$ decreased in the Levantine Sub-basin from a range of 0.33-0.4 to 0.2-0.39 in deep layers (below 1000 m ; Fig. 6). However, its mixing proportion remains high in the South of the Ionian Sub-basin (~ 0.54). These results indicate that the $EMDW_{Adr}$ stays a minor deep water mass in the Levantine Sub-basin. Moreover, it shows that the $EMDW$ originated from the Adriatic Sub-basin is reaching the South of the Ionian Sub-basin but maybe physical phenomena (like the more defined open ocean, free jet intensified structures mentioned by [55] : the Atlantic-Ionian Stream and the Mid-Mediterranean Jet in the Eastern basin) are disabling its dissemination all over the Eastern basin.

3.2.5. Eastern Mediterranean Deep Water Originated in the Aegean Sub-Basin, $EMDW_{Ag}$

The mixing coefficients of this water mass show a remarkable increase, particularly in the Levantine Sub-basin (from 0.29 to 0.54 at 1500 m ; Fig.6). The steady occurrence of this dominant deep water mass in the Levantine Sub-basin, instead of $EMDW_{Adr}$, could be due to many factors : 1- the waters formed in the Aegean Sub-basin still sufficiently denser than those originated in the Adriatic one, hence, it is able to enter the deep layers of the Eastern basin and not address the intermediate water range depth [67], 2- the deep water formation in the Adriatic Sub-basin seems to be significantly impacted by the EMT [68], thus it is not able to match its previous distribution in the Eastern basin, 3- the atmospheric and physical conditions favorable for the formation of this dense water still dominant in this area, 4- our results could also be attributed to the long residence time of deep water masses in the Levantine Sub-basin due the robust topography which could trap the deep waters for a long period (~ 100 years) [69] [70].

4. Conclusion

Overall, based upon data from the 2008 BOUM and 2013 MedSeA cruises, the present study shows a significant evolution of some water masses mixing coefficients, calculated via the MIX approach. Surface water mass proportions did not change significantly. However, intermediate and deep water mass mixing coefficients of both Eastern and Western basins were noticeably modified. Moreover, these results indicate that the mixing coefficient of the $EMDW_{Adr}$ is always high in the Ionian Sub-basin, while it remains low in the Levantine Sub-basin compared to the $EMDW_{Ag}$. Furthermore, the observed decrease of the mixing coefficient in the deep Western water masses (TDW, WMDW) is mainly attributed to changes in the deep Eastern water masses circulation. This study proves that the hypothesis of "steady state" situation for the Mediterranean Sea is far from being validated. This sea witnesses continuous and significant water masses changes [ex. The Eastern Mediterranean Transient, EMT [66] [71] [72] [73] [74], changes in circulation pattern in the Ionian Sub-basin which affect the contiguous sub-basins [61] [75]. Therefore, this work could be an incentive for further studies to innovate new oceanic models that take into considerations the unsteady state situation of the water masses circulation in the Mediterranean Sea (and probably in other oceanic areas). In the context of the global warming, further measurements of these water masses properties are necessary, to assess their evolutions and to evaluate the consequences of any modification on the global circulation.

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