



**HAL**  
open science

## **The significance of the episodic nature of atmospheric deposition to Low Nutrient Low Chlorophyll regions**

C. Guieu, O. Aumont, A. Paytan, L. Bopp, C.S. Law, N. Mahowald, E. P. Achterberg, E. Marañón, B. Salihoglu, A. Crise, et al.

### ► **To cite this version:**

C. Guieu, O. Aumont, A. Paytan, L. Bopp, C.S. Law, et al.. The significance of the episodic nature of atmospheric deposition to Low Nutrient Low Chlorophyll regions. *Global Biogeochemical Cycles*, 2014, pp.1179-1198. <10.1002/2014GB004852>. <hal-01110987>

**HAL Id: hal-01110987**

**<https://hal.science/hal-01110987v1>**

Submitted on 28 Oct 2020

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



HAL Authorization



# Global Biogeochemical Cycles

## RESEARCH ARTICLE

10.1002/2014GB004852

### Key Points:

- Episodic nature of atmospheric deposition matters in LNLC productivity
- Atmospheric impact in LNLC area is not a simple “fertilization effect”
- Atmosphere supplies most of new N & Fe to the mixed layer in some LNLC regions

### Supporting Information:

- Readme
- Text S1
- Aumontetal\_in\_prep\_2014.pdf

### Correspondence to:

C. Guieu,  
guieu@obs-vlfr.fr

### Citation:

Guieu, C., et al. (2014), The significance of the episodic nature of atmospheric deposition to Low Nutrient Low Chlorophyll regions, *Global Biogeochem. Cycles*, 28, 1179–1198, doi:10.1002/2014GB004852.

Received 19 MAR 2014

Accepted 13 OCT 2014

Accepted article online 15 OCT 2014

Published online 10 NOV 2014

## The significance of the episodic nature of atmospheric deposition to Low Nutrient Low Chlorophyll regions

C. Guieu<sup>1</sup>, O. Aumont<sup>2</sup>, A. Paytan<sup>3</sup>, L. Bopp<sup>4</sup>, C. S. Law<sup>5,6</sup>, N. Mahowald<sup>7</sup>, E. P. Achterberg<sup>8,9</sup>, E. Mara<sup>10</sup>, B. Salihoglu<sup>11</sup>, A. Crise<sup>12</sup>, T. Wagener<sup>13</sup>, B. Herut<sup>14</sup>, K. Desboeufs<sup>15</sup>, M. Kanakidou<sup>16</sup>, N. Olgun<sup>17,18</sup>, F. Peters<sup>19</sup>, E. Pulido-Villena<sup>13</sup>, A. Tovar-Sanchez<sup>20,21</sup>, and C. Völker<sup>22</sup>

<sup>1</sup>Laboratoire d’Océanographie de Villefranche (LOV), UMR7093, CNRS-INSU-Université Paris 6, Villefranche sur Mer, France,

<sup>2</sup>Laboratoire de Physique des Océans, Institut Européen de la Mer, Centre IRD de Bretagne, Plouzané, France, <sup>3</sup>Institute of

Marine Sciences, University of California Santa Cruz, Santa Cruz, USA, <sup>4</sup>LSCE/IPSL, CNRS/CEA/UVSQ, Gif sur Yvette, France,

<sup>5</sup>National Institute of Water and Atmospheric Research (NIWA), Wellington, New Zealand, <sup>6</sup>Department of Chemistry,

University of Otago, Dunedin, New Zealand, <sup>7</sup>Department of Earth and Atmospheric Sciences, Cornell University, Ithaca,

USA, <sup>8</sup>Ocean and Earth Science, National Oceanography Centre Southampton, University of Southampton, Southampton,

UK, <sup>9</sup>Now at GEOMAR, Helmholtz Centre for Ocean Research Kiel, Kiel, Germany, <sup>10</sup>Departamento de Ecología y Biología

Animal, Facultad de Ciencias del Mar, Universidad de Vigo, Vigo, Spain, <sup>11</sup>Institute of Marine Sciences, METU, MERSİN,

Turkey, <sup>12</sup>Istituto Nazionale di Oceanografia e di Geofisica Sperimentale (OGS), Trieste, Italy, <sup>13</sup>Institut Méditerranéen

d’Océanologie (MIO), UMR 7294, Université d’AIX-Marseille-CNRS-IRD, Marseille, France, <sup>14</sup>Israel Oceanographic and

Limnological Research (IOLR), National Institute of oceanography, Haifa, Israel, <sup>15</sup>Laboratoire Interuniversitaire des Systèmes

Atmosphériques, UMR CNRS 7583, Université Paris Est Créteil, Université Paris Diderot, Institut Pierre Simon Laplace, Créteil,

France, <sup>16</sup>Environmental Chemical Processes Laboratory, Dept. of Chemistry, University of Crete, Heraklion, Greece,

<sup>17</sup>Helmholtz-Center for Ocean Research Kiel, GEOMAR, Kiel, Germany, <sup>18</sup>Now at Eurasia Institute of Earth Sciences, Istanbul

Technical University, Istanbul, Turkey, <sup>19</sup>Institut de Ciències del Mar, CMIMA (CSIC), Barcelona, Spain, <sup>20</sup>Department of Global

Change Research IMEDEA (CSIC-UIB), Instituto Mediterráneo de Estudios Avanzados, Spain, <sup>21</sup>Department of Ecology and

Coastal Management, Andalusian Institute for Marine Science, ICMAN (CSIC), Campus Universitario Río San Pedro, Puerto Real,

Cádiz, Spain, <sup>22</sup>Alfred Wegener Institute for Polar and Marine Research, Bremerhaven, Germany

**Abstract** In the vast Low Nutrient Low-Chlorophyll (LNLC) Ocean, the vertical nutrient supply from the subsurface to the sunlit surface waters is low, and atmospheric contribution of nutrients may be one order of magnitude greater over short timescales. The short turnover time of atmospheric Fe and N supply (<1 month for nitrate) further supports deposition being an important source of nutrients in LNLC regions. Yet, the extent to which atmospheric inputs are impacting biological activity and modifying the carbon balance in oligotrophic environments has not been constrained. Here, we quantify and compare the biogeochemical impacts of atmospheric deposition in LNLC regions using both a compilation of experimental data and model outputs. A metadata-analysis of recently conducted field and laboratory bioassay experiments reveals complex responses, and the overall impact is not a simple “fertilization effect of increasing phytoplankton biomass” as observed in HNLC regions. Although phytoplankton growth may be enhanced, increases in bacterial activity and respiration result in weakening of biological carbon sequestration. The application of models using climatological or time-averaged non-synoptic deposition rates produced responses that were generally much lower than observed in the bioassay experiments. We demonstrate that experimental data and model outputs show better agreement on short timescale (days to weeks) when strong synoptic pulse of aerosols deposition, similar in magnitude to those observed in the field and introduced in bioassay experiments, is superimposed over the mean atmospheric deposition fields. These results suggest that atmospheric impacts in LNLC regions have been underestimated by models, at least at daily to weekly timescales, as they typically overlook large synoptic variations in atmospheric deposition and associated nutrient and particle inputs. Inclusion of the large synoptic variability of atmospheric input, and improved representation and parameterization of key processes that respond to atmospheric deposition, is required to better constrain impacts in ocean biogeochemical models. This is critical for understanding and prediction of current and future functioning of LNLC regions and their contribution to the global carbon cycle.

## 1. Introduction

Oceans influence climate primarily due to their ability to take up and store heat, and their direct and indirect effects on global climate via regulation of atmospheric CO<sub>2</sub> and other radiatively active gases [Friedlingstein

*et al.*, 2006; *Le Quéré et al.*, 2009; *Le Quéré et al.*, 2013]. Sixty percent of the global oceans, primarily the central ocean gyres, are depleted in the primary macronutrients nitrate and phosphate, and consequently sustain low growth of phytoplankton and other marine organisms [*Antoine et al.*, 1996]; yet these “ocean deserts” represent ecosystems occupying a large proportion of Earth’s surface area. These oligotrophic regions are generally characterized by chlorophyll a (Chl a) concentrations  $<0.07 \text{ mg m}^{-3}$  and co-dominated by small phytoplankton and heterotrophic bacteria [*Uitz et al.*, 2010; *Cho and Azam*, 1990], and are referred to as Low Nutrient Low Chlorophyll (LNLC) areas (Figure 1a).

Wet and dry atmospheric deposition transport a range of compounds from a variety of natural and anthropogenic land sources to the ocean. The compounds include macro- and micronutrients (N, P, C, Si, Fe, and other metals) [*Duce et al.*, 1991] as well as potentially toxic elements (e.g., Cu and Pb) [*Paytan et al.*, 2009; *Jordi et al.*, 2012]. The main natural source of land-derived particles to the open ocean is wind-blown desert dust, which constitutes the primary atmospheric source of iron [*Jickells et al.*, 2005]. Atmospheric nitrogen is mainly derived from anthropogenic combustion or agricultural sources from densely populated regions throughout the world [*Duce et al.*, 2008], while phosphorus originates from both desert dust and anthropogenic sources [*Mahowald et al.*, 2008] (Figure 1b). Atmospheric supply of dissolved constituents to the surface ocean depends on particle concentration and size spectrum, and the solubility of the element-bearing phases in aerosols [*Trapp et al.*, 2010; *Baker and Jickells*, 2006] which is influenced by atmospheric processing during transport [*Krishnamurthy et al.*, 2009]. For example, the extent to which dust interacts with anthropogenic acids ( $\text{H}_2\text{SO}_4$  and  $\text{HNO}_3$ ) during transport increases the solubility of various elements [*Desboeufs et al.*, 2001] resulting in enrichment of nitrogen [*Geng et al.*, 2009], and enhanced supply of potentially bioavailable compounds to the surface ocean. Furthermore atmospheric deposition supplies N and P in both inorganic and organic forms, which are accessible to both heterotrophic organisms and autotrophic phytoplankton [*Cornell*, 2011; *Kanakidou et al.*, 2012; *Moore et al.*, 2013]. Deposition is dependent on the distance from source, with sites located between 10 and 1000 km from source areas receiving dust deposition of  $1.0\text{--}50 \text{ g m}^{-2} \text{ yr}^{-1}$ , and sites located  $>1000 \text{ km}$  receiving  $0.05\text{--}1.00 \text{ g m}^{-2}$  annually, as determined by a compilation of direct measurements of dust deposition [*Lawrence and Neff*, 2009]. According to the authors, this range represents the background rate of dust deposition, such that all sites throughout the world receive, at minimum, this deposition range each year.

The supply of new nutrients to the ocean from external sources such as atmospheric deposition has been extensively addressed in iron-limited High Nutrient-Low Chlorophyll (HNLC) regions [i.e., *Boyd et al.*, 2007], most of which receive low atmospheric inputs at the present time (Figure 1b). However, much less attention has been paid to the importance of atmospheric deposition to LNLC regions. Until recently, models considering atmospheric deposition to the ocean focused primarily on iron and typically regarded deposition as a continuous input, using mean deposition values, often without consideration of the highly episodic nature of such deposition. To our knowledge, only one study has considered the temporal variability of atmospheric deposition [*Aumont et al.*, 2008]. This study was restricted to iron deposition and showed that a significant variability in surface iron concentrations can be generated in high deposition regions, for instance in the subtropical North Atlantic Ocean. For the other nutrients taken independently or altogether, we are not aware of any equivalent studies. Yet it is widely recognized that atmospheric deposition, and in particular dust deposition, is by nature highly episodic. Indeed, daily dust deposition rates as high as 4 times the monthly mean flux have been measured in the North Atlantic [*Moxim et al.*, 2011]. This “pulsed character” is also well identified in the long (since 1965) aerosol sampling series in Barbados [*Prospero and Lamb*, 2003]. The episodic nature of deposition is also reported for the North Pacific where most of the mineral dust input to the ocean typically takes place during 3–5 events in the spring, each of which lasts 1–3 days [*Donaghay et al.*, 1991]. Wet deposition is also not continuous as rainfall events vary from small intense storms of a few square kilometers to large frontal systems that stretch for thousands of kilometers, and so wet atmospheric inputs are both episodic and spatially patchy [*Uematsu et al.*, 1985; *Donaghay et al.*, 1991]. In the Mediterranean Sea a few intense events provide the majority of the annual deposition [i.e. *Loje-Pilot and Martin*, 1996; *Guerzoni et al.*, 1999] with measured short (a few hours) event fluxes exceeding  $20 \text{ g m}^{-2}$  [*Bonnet and Guieu*, 2006; *Guieu et al.*, 2010a; *Ternon et al.*, 2010]. These observations confirm that episodicity is the norm for deposition over the ocean.

How the ocean responds to pulses of deposition, as either transient or long-term impacts on diversity of the natural assemblage and/or carbon export, is not obvious from in situ or remote sensing observations reported in the literature. For example, atmospheric dust concentrations measured at the Canary Islands concomitantly

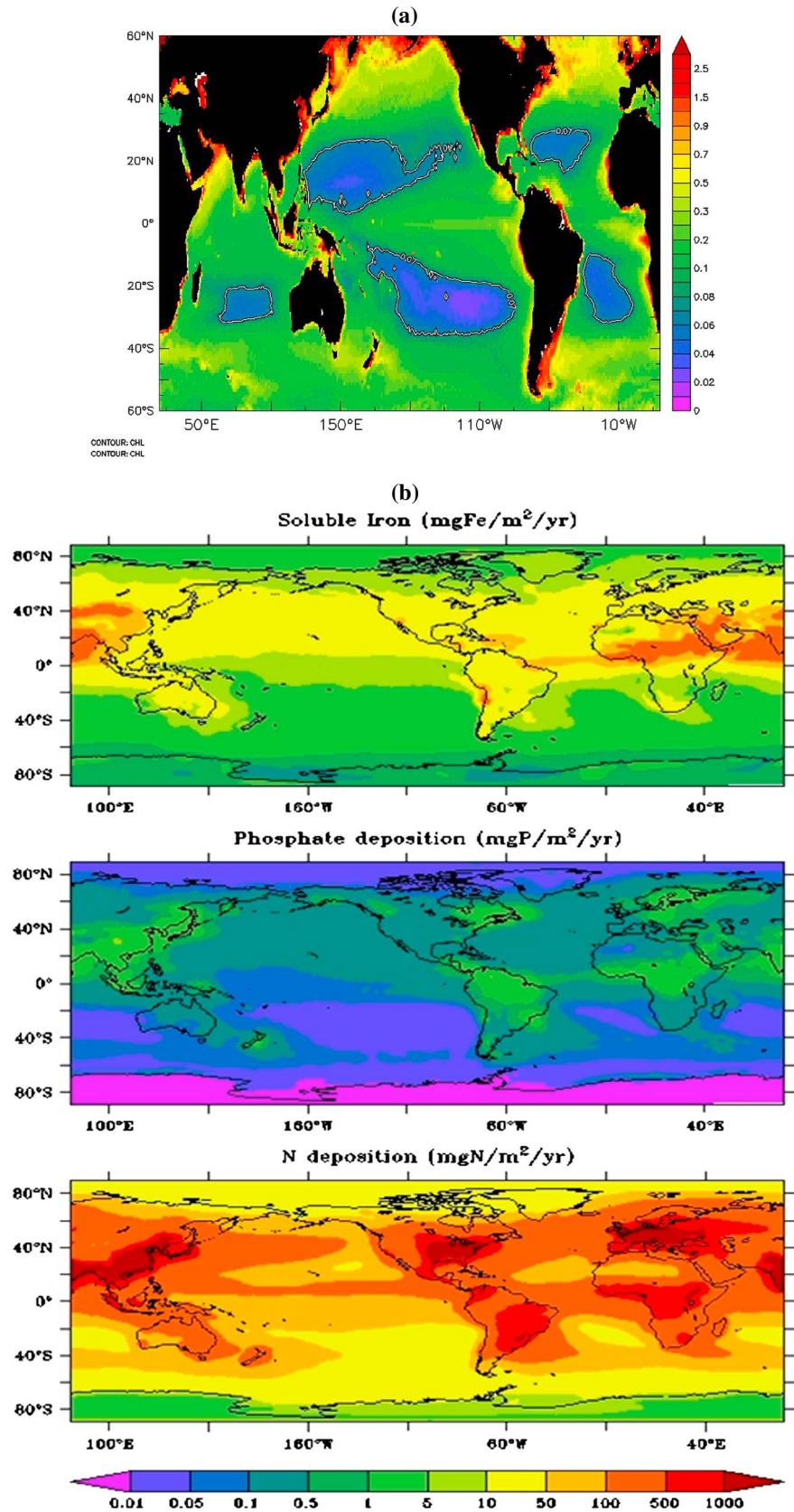
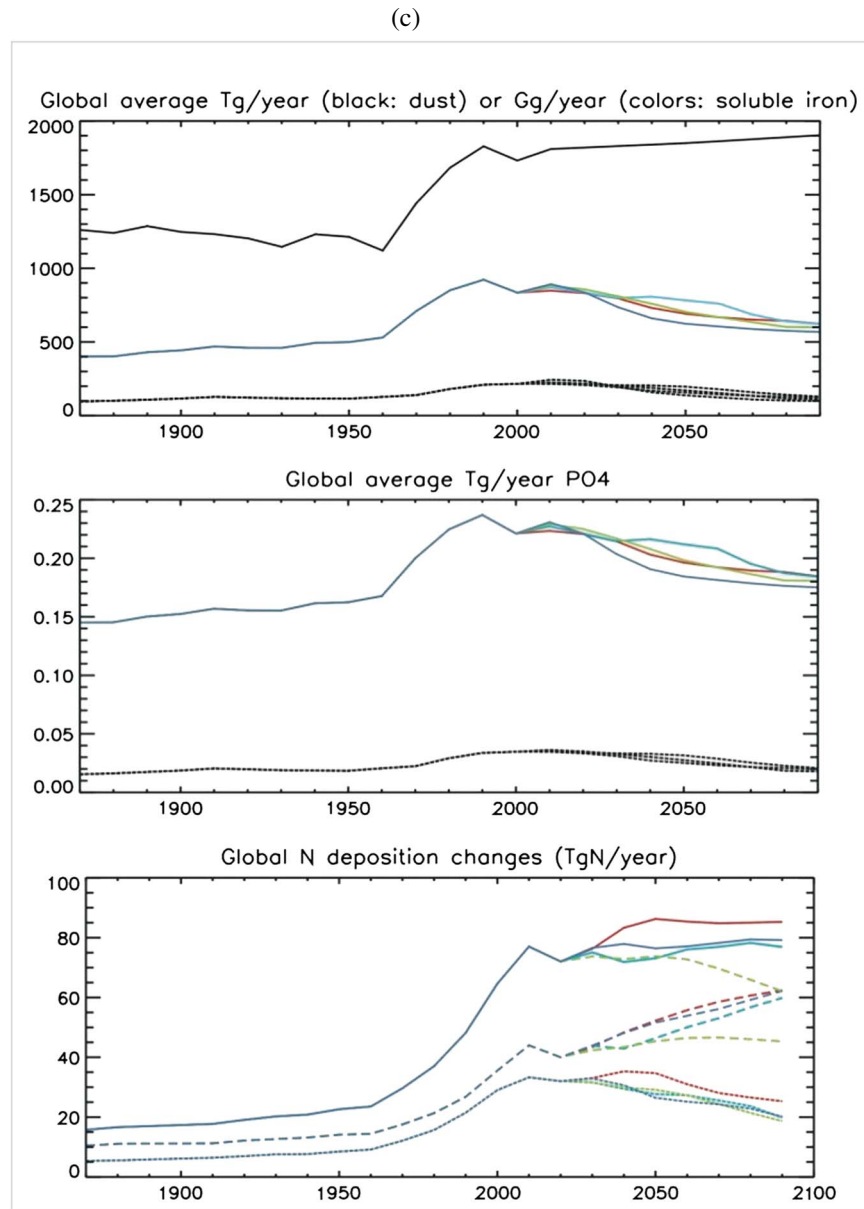


Figure 1



**Figure 1.** (a) Location of oligotrophic gyres (blue color) as derived from SeaWiFS climatology (1997–2007) with a criteria of Chl a  $< 0.07 \text{ mg} \cdot \text{m}^{-3}$ . (b) Global atmospheric dissolved fluxes derived from models for iron, phosphate and nitrogen (from Luo *et al.*, 2008; Mahowald *et al.*, 2008), Lamarque *et al.*, 2010. (c) Evolution of simulated deposition of dust (Tg/y) and soluble iron (Gg/y), phosphate (Tg/y), and nitrogen (Tg/y) over the period 1850–2100. For DFe and phosphate, dashed lines are deposition from combustion sources. For N deposition, solid lines are total N, dashed lines are NO<sub>y</sub>, and dotted lines are NH<sub>x</sub>. Uncertainties for deposition estimates are 40% for past changes, and 100% for future changes. Future estimates are from the four representational concentration pathways (RCPs) scenarios (RCP2.6 blue, RCP4.5 green, RCP6.0 light blue, and RCP8.5 red). Estimates based on Lamarque *et al.* [2010, 2011], Luo *et al.* [2008], Mahowald *et al.* [2008, 2009, 2010], and van Vuuren *et al.* [2011]; (see description in Methods).

with upper water biogeochemistry at the oligotrophic time series station ESTOC (European Station for Time series in the Ocean, Canary Island) over 2 years have shown that higher aerosol concentrations (continuously measured) were not accompanied by higher primary production (derived from Chl a) or export production [Neuer *et al.*, 2004]. The authors concluded that phytoplankton production remained unaffected by atmospheric nitrogen supply on annual timescales. In the Mediterranean Sea, Herut *et al.* [2005] observed a significant dust storm at sea (CYCLOPS cruise), which caused as a sharp reduction in Prochlorococcus abundance and a slight increase in Chl a and in bacterial activity. In the Pacific, three Asian dust storms stimulated the growth of Synechococcus but not

Prochlorococcus [Chung *et al.*, 2011]. Comparing aerosol optical thickness and chlorophyll derived from satellite color during a 4 year period in the Mediterranean Sea, Volpe *et al.* [2009] concluded that the methodology induces important biases arising from atmospheric correction linked to Saharan aerosols and so this could not be used to test a dust fertilization in oligotrophic systems such as the Mediterranean Sea. Yet, the authors conclude that dust inputs do not play a significant role in the phytoplankton dynamic in the Mediterranean Sea. In some of these studies, the role of grazers is suspected to play an important role in maintaining a low phytoplankton biomass but this has not been confirmed by observation and measurement. In conclusion available data indicate that the oligotrophic ocean exhibits a variety of responses to atmospheric deposition and no simple pattern can be drawn which denote the large variety of oligotrophic systems.

The published results from different models agree that increasing iron supply by atmospheric deposition stimulates marine productivity and export production in HNLC regions [Archer and Johnson, 2000; Moore *et al.*, 2002]. This may potentially increase the LNLC ocean area, due to the resulting reduction in lateral nutrient transport from HNLC regions [Dutkiewicz *et al.*, 2005; Aumont and Bopp, 2006]. Models also agree that an increase in Fe and P deposition to LNLC areas may enhance N<sub>2</sub> fixation in LNLC areas [Moore *et al.*, 2002; Krishnamurthy *et al.*, 2009; Mahowald *et al.*, 2011], potentially lowering atmospheric pCO<sub>2</sub> [Bopp *et al.*, 2003; Parekh *et al.*, 2006]. In the few modeling studies that have investigated the role of atmospheric deposition as a source of nutrients other than iron, atmospheric inputs of nitrogen were shown to have a very modest effect on marine productivity, export production, or carbon uptake on a global scale yet identified significant effects in LNLC regions [Krishnamurthy *et al.*, 2007, 2009, 2010]. Although phosphorus deposition accounted for only a very small fraction of export production [Krishnamurthy *et al.*, 2009, 2010] the contribution of atmospheric nitrogen deposition was significant. A further response was a decrease in N<sub>2</sub> fixation in LNLC regions, possibly due to P limitation.

To examine the impact of aerosol deposition and its temporal variability in LNLC systems, we first examined the turnover times relative to atmospheric deposition (TTADs) in the surface mixed layer for iron, nitrate, and phosphate, to assess the contribution of atmospheric deposition to nutrient stocks in the global surface ocean. We further evaluated the potential impact of new atmospheric nutrient inputs (iron, nitrate and phosphate) in LNLC regions on primary production, N<sub>2</sub> fixation, surface Chl *a* concentrations, and export production by applying atmospheric deposition to a coupled 3D ocean ecosystem-biogeochemical model. We then compared the model results with a compilation of published experimental responses of natural LNLC seawater to aerosol addition (see ref. in Table 1), to further examine the impacts of episodicity of aerosol deposition.

## 2. Methods

### 2.1. Turnover Times Relative to Atmospheric Deposition (TTADs) for Nitrate, Phosphate, and Dissolved Iron

The TTAD is defined as the time required to replace the surface mixed layer nutrient inventory solely by atmospheric deposition, and so TTADs (in years) were derived by dividing the vertically averaged nutrient concentrations (moles per cubic meter) in the surface mixed layer by the contribution of atmospheric deposition to the water volume of the mixed layer (moles per cubic meter per unit time) (see Figure 2a). Nitrate and phosphate in the surface mixed layer are obtained from the latest Levitus climatology (World Ocean Atlas 2009, <http://www.nodc.noaa.gov/OC5/indprod.html>). As a sensitivity analysis, we have alternatively used the CARS2009 climatology ([www.cmar.csiro.au/cars](http://www.cmar.csiro.au/cars)) to compute the TTADs. Results are almost identical (see Figure S1). For iron in the surface mixed layer, we used a global compilation of over 13,000 published measurements of dissolved iron [Tagliabue *et al.*, 2012]. The mixed layer depth was taken from the latest version of the global climatology [de Boyer Montégut *et al.*, 2004] (<http://www.lodyc.jussieu.fr/~cdblod/mldepth.html>), using a density criterion of 0.03 kg m<sup>-3</sup>. Atmospheric deposition fluxes for nitrogen were based on published emissions and simulations of the historical time period and Representative Concentration Scenarios (RCPs) [van Vuuren *et al.*, 2011; Lamarque *et al.*, 2010, 2011]. Estimates of P and soluble P are based on [Mahowald *et al.*, 2008], assuming that the combustion P follows the evolution of the black carbon in the same scenarios. Estimates of iron and iron solubility are based on [Mahowald *et al.*, 2009] including combustion iron, which follows the black carbon historical and RCP scenarios. Desert dust evolution is based for the historical time period on paleoclimate reconstructions [Mahowald *et al.*, 2010], with future evolution based on estimates of desert area change in the future [Mahowald, 2007] assuming no carbon dioxide fertilization. All fields have been linearly interpolated to the typical Levitus grid (1° × 1° horizontal resolution).

**Table 1.** Compilation of Experimental Studies Adding Different Types of Aerosols to Surface Seawater, With Aerosol Type 1: Fine Fraction Surface Soil; Aerosol Type 2: Fine Fraction Surface Soil With Physico-chemical Treatment; Simulating Atmospheric Processing; Aerosol Type 3: Particulate Phase From Rain; Aerosol Type 4: Dust Collection During Dust Storm; Aerosol Type 5: Local Aerosol on Filters (not Restricted to Dust); Aerosol Type 6: Anthropogenic Aerosols (Urban Particulate Matter From NIST)<sup>a</sup>

Region	Tested Parameters	Physical Nature of the Experiment	Time Scale of the Study	Aerosol Type	Aerosol Type (Details)	Amount of Aerosol/Dust Added or Final Aerosol/Dust Conc in Bottle or Simulated Flux	Estimated Deposition Flux (g m <sup>-2</sup> ) Mimicked in the Experiments (*)	References
N Atlantic	Chl a, nano, and microphytoplankton, PP	1	6 d	1	Fine fraction of composite Saharan soils collected in Algeria***	1.34 mg.l <sup>-1</sup>	13	Blain et al., 2004
N Atlantic	BA, BP, Chl a, Syn-, Proc-, PP, and N <sub>2</sub> fix	1	2 d	1, 4	(1) fine fraction (<20 μm) of soils collected in Mali and (2) atmospherically processed dust collected in the N Atlantic	2 mg.l <sup>-1</sup>	20	Marañón et al. [2010]
N Atlantic	Chl a, PP, and N <sub>2</sub> fix	1	2 d	1	Fine fraction of composite Saharan soils collected in Algeria***	0.5–2 mg.l <sup>-1</sup>	5–20	Mills et al. [2004]
N Atlantic	Chl a, PP	1	2 d	1, 4	Fine fraction of composite Saharan soils collected in Algeria***	2 mg.l <sup>-1</sup>	20	Moore et al. [2006]
N Atlantic	Chl a	1	2 d	1	Dust collected in Barbados (atmospherically processed)	2 mg.l <sup>-1</sup>	20	Achterberg unpub. Data
Sargasso Sea	Syn-, proch-, and pico-euk	1	3 d	5	Locally collected aerosols		2–20	Mackey et al. [2012]
Coastal California	Chl a	1	6 d	5	Collected aerosols	1 mg.l <sup>-1</sup>	10	Mackey et al. [2010]
Southeast Pacific	Syn-, proch-, pico-euk, PP, and N <sub>2</sub> fix	1	2 d	1	Fine fraction of composite Saharan soils collected in Algeria***	0.25 mg.l <sup>-1</sup>	3	Bonnet et al. [2008]
SW Pacific	N <sub>2</sub> fix	1, 5	5 d	1, 7	20–25 mm size fraction obtained from sieving surface soils, one from Australia, one from Gobi Desert	0.84 mg.l <sup>-1</sup>	8.4	Law et al. [2011]
Tasman Sea	Syn- and proc-	1	5 d	1	20–25 mm size fraction obtained from sieving surface soils, one from Australia, one from Gobi Desert	0.84 mg.l <sup>-1</sup>	8.4	Ellwood et al. [2013]
Tasman Sea	N <sub>2</sub> fix	1	5 d	1	Aerosol dust collected during dust storm in Brisbane, Australia	0.84 mg.l <sup>-1</sup>	8.4	Law, unpub. Data
Red Sea	N <sub>2</sub> fix	1	2 d	5	Aerosol dust filters	0.71 mg.l <sup>-1</sup>	7	Foster et al. [2009]
Red Sea	Chl a	2	4 d	5	Aerosol dust collected	0.75 mg.l <sup>-1</sup>	8	Mackey et al. [2007]
Red Sea	Chl a, Syn-, Proc-, pico-euk	1	5 d	5	(1) European aerosols; (2) Saharan aerosols	0.75 mg.l <sup>-1</sup>	8	Paytan et al. [2009]
East Mediterranean	BA, BP, Chl a, syn-, proch-, PP	1, 5	4 d	4, 7	Collected dust dry deposition with and without pre-leaching in filtered SSW	0.2–4.9 mg.l <sup>-1</sup>	2	Herut et al. [2005]

**Table 1.** (continued)

Region	Tested Parameters	Physical Nature of the Experiment	Time Scale of the Study	Aerosol Type	Aerosol Type (Details)	Amount of Aerosol/Dust Added or Final Aerosol/Dust Conc in Bottle or Simulated Flux	Estimated Deposition Flux ( $\text{g}\cdot\text{m}^{-2}$ ) Mimicked in the Experiments (*)	References
Mediterranean Sea	BA, BP	1	2 d	1	Fine fraction of composite Saharan soils collected in Algeria***	$1.25\text{--}10\text{ mg}\cdot\text{l}^{-1}$	13	Ridame, 2001
Mediterranean Sea	$\text{N}_2$ fix	1	2 d	1	Fine fraction of composite Saharan soils collected in Algeria***	$1.1\text{ mg}\cdot\text{l}^{-1}$	11	Ridame et al. [2011]
Mediterranean Sea	BA, syn-, PP, and $\text{N}_2$ fix	1	2 d	2, 5	Saharan dust analog**	$1\text{ mg}\cdot\text{l}^{-1}$	10	Ternon et al. [2011]
W Mediterranean	Chl a, nano and microphytoplankton, BA, and PP	1	3 d	1, 6	Dust = fine fraction of Saharan soils collected in Algeria***; anthropogenic = NIST ref. material	dust = $0.25\text{ mg}\cdot\text{l}^{-1}$ . Anthropogenic particles = $0.01\text{ mg}\cdot\text{l}^{-1}$	0.1–2.5	Bonnet et al. [2005]
W Mediterranean	BA, syn-, pico, and nano-euks	3	8 d	2	Saharan dust analog**	$10\text{ g}\cdot\text{m}^{-2}$ (seeding over mesocosm)	10	Laghdass et al. [2011]
W Mediterranean	BA, BP, Chl a, and pico-euk	2	3 d	4	Particulate phase of rain	$50\text{--}500\text{ mg}\cdot\text{l}^{-1}$	500–5000 #	Lekunberri et al. [2010]
W Mediterranean	BA, BR	1, 5	3 d	1, 3	Saharan dust analog** and particulate fraction of a Saharan rain	$0.5\text{--}2\text{ mg}\cdot\text{l}^{-1}$	5–20	Pulido-Villena et al. [2008]
W Mediterranean	PP, $\text{N}_2$ fix	3	8 d	2	Saharan dust analog**	$10\text{ g}\cdot\text{m}^{-2}$ (seeding over mesocosm)	10	Ridame et al. [2013]
W Mediterranean	Chl a	1	8 d	2	Saharan dust analog**	$10\text{ g}\cdot\text{m}^{-2}$ (seeding over mesocosm)	10	Guieu et al. [2014a]
W Mediterranean	BA	3	8 d	2	Saharan dust analog**	$10\text{ g}\cdot\text{m}^{-2}$ (seeding over mesocosm)	10	Pulido-Villena unpubl. Data
W Mediterranean	BR	3	8 d	2	Saharan dust analog**	$10\text{ g}\cdot\text{m}^{-2}$ (seeding over mesocosm)	10	Pulido-Villena et al. [2014]
W Mediterranean	BR	3	8 d	2	Saharan dust analog**	$10\text{ g}\cdot\text{m}^{-2}$ (seeding over mesocosm)	10	Guieu et al. [2014b]

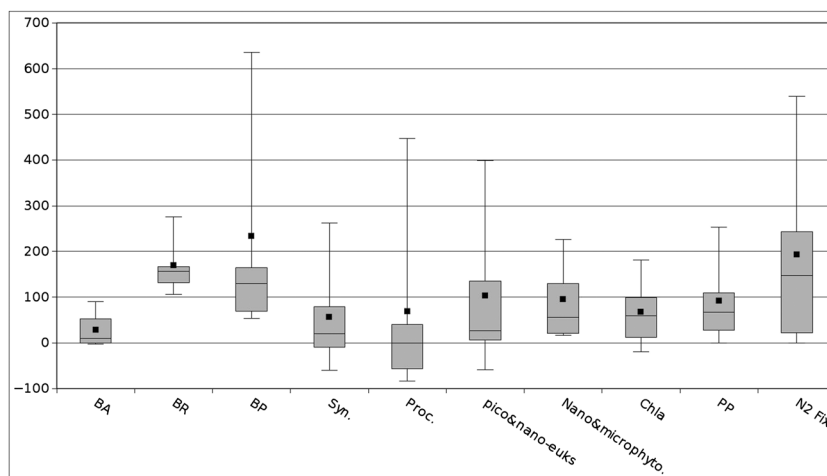
<sup>a</sup>Also reported in situ response to dust storm deposition (7).

Measured changes in bacteria abundance (BA), bacteria respiration (BR), bacteria production (BP), chlorophyll-a concentration (Chl a), organisms abundance (Syn. = Synechococcus; Proc. = Prochlorococcus; pico and nano-euks = pico- and nano-eukaryotes, nano and micro-phytoplankton), primary production (PP), and nitrogen fixation ( $\text{N}_2$  fix). The physical framework of the experiment is reported as follows: 1: bottles with volumes  $\leq 4$  L; 2: microcosms (volumes  $> 4$  L); 3: mesocosms ( $> 50\text{ m}^2$ ) and 4: in situ. Time scale of the studies in days (d).

\*\*If a homogeneous dilution of particles in 5–10 m surface mixed layer is considered.

\*\*\*Evapocondensed (EC) dust obtained from fine fraction of Saharan soils [Guieu et al., 2010b].

\*\*\*The same dust was used in those experiments. # these values do not correspond to any published fluxes and have not been taken into account in the calculation of the range of average values ( $0.1\text{--}20\text{ g}\cdot\text{m}^{-2}$ /event) given in the text (see section 2.4).



**Figure 2.** Box-Whisker plots (box portion = interquartile range (25<sup>th</sup> to 75<sup>th</sup> percentile) of the data set. Horizontal bar within the box = median value. Black square = mean value. Also represented maximum and minimum values) showing the responses of different biological variables to aerosol additions in LNLC waters: synthesized from available data from field and laboratory aerosol addition bioassay experiments and mesocosm experiments (Table 1). The responses are % changes in the aerosol treatment relative to the control after 2–8 days (Table 1), with zero indicating no difference between the aerosol treatment and the control, and a positive response indicating an increase in the parameter in the aerosol treatment relative to the control. Parameters: (BA) Bacteria Abundance, (BR) Bacteria Respiration, (BP) Bacteria Production, (Syn.) *Synechococcus* abundance, (Proc.) *Prochlorococcus* abundance, (pico and nano-euks) nano- and picoeukaryotes abundance, (nano and microphyto) nano- and micro-phytoplankton abundance, (Chl a) Chlorophyll-*a*, (PP) primary production, and (N2Fix) nitrogen fixation.

## 2.2. Contribution of Atmospheric Deposition to the Total Supply of Nutrient to the Surface Mixed Layer

Vertical dynamic supply of the N, P, and Fe are computed following an approach similar to *Fung et al.* [2000.] The upwelling rate was taken from the dynamic simulation produced by ORCA2-LIM, which is also used in our biogeochemical model experiments. Entrainment is considered to be zero when the mixed layer is shoaling or when nutrient concentrations just below the mixed layer are lower than in the mixed layer. Otherwise, it is computed as the amount of nutrients in excess of those in the mixed layer that are entrained into the mixed layer when the latter is deepening. Diffusion across the base of the mixed layer is set to  $10^{-5} \text{ m}^{-2} \text{ s}^{-1}$  [*Law et al.*, 2003; *Capone et al.*, 2005]. Nutrients supplied by lateral advection were not considered. Deposition fluxes, nutrient distributions, and the mixed layer depth are identical to those used to compute the turnover time maps.

## 2.3. Biogeochemical Model

The dynamic state of the ocean has been simulated using NEMO in its version 3.2 and in its global configuration ORCA2-LIM [*Madec*, 2008]. The spatial resolution is about  $2^\circ$  by  $2^\circ \cos(\phi)$  (where  $\phi$  is the latitude) with a focusing to  $0.5^\circ$  of the meridional resolution in the equatorial domain. The model has 30 vertical layers, increasing in thickness from 10 m at the surface to 500 m at 5000 m. The ocean model is driven by climatological atmospheric fields identical to those used in *Aumont and Bopp* [2006]. However, the resulting dynamics simulated by the ocean model is different as several new parameterizations and new algorithms have been included in ORCA2-LIM. Ocean biogeochemistry is simulated using PISCES [*Aumont and Bopp*, 2006] which is forced offline by the 5 days mean ocean physical fields produced by the ocean physical model. Some significant modifications have been added to PISCES (O. Aumont, et al., PISCES: An ocean biogeochemical model for carbon and ecosystem studies, *Geoscientific Model Development*, in preparation 2014). Nutrients are supplied to the ocean from four different sources: atmospheric dust and N deposition, rivers, sea-ice, and sediment mobilization, with river, sea-ice, and sediment sources described in *Aumont and Bopp* [2006] as modified by (O. Aumont et al., in preparation 2014). Atmospheric contributions of nutrients are the same as in Figure 1b. The coupled ocean biogeochemical model is spun up offline for 4000 years, so that a quasi-steady state is reached, with primary production and  $\text{CO}_2$  fluxes varying by less than  $0.01 \text{ GtC yr}^{-1}$ . A brief description of some new parameterizations of PISCES, relative to the previously published version of *Aumont and Bopp* [2006], is presented in the supporting information. This description is restricted to the processes that play a key role in this study, including  $\text{N}_2$  fixation and size variability of nanophytoplankton. A validation of the model behavior is also proposed in the supporting information (Figures S2–S10).

**Table 2.** Main Characteristics of the Model Experiments<sup>a</sup>

Model Experiment	Fe and P Deposition	N Deposition	Deposition Pulses dust ( $1 \text{ g m}^{-2} \text{ d}^{-1}$ ), N ( $2.8 \text{ mg N m}^{-2} \text{ d}^{-1}$ )	Duration (years)
Standard	Monthly	Annual	None	100
No "D"	None	None	None	100
"D"	Monthly, $\times 5$	Annual, $\times 5$	None	100
Pulse	Monthly	Annual	1 daily pulse, either on 15 of January or on 15 July	100, last year with 1 day pulses

<sup>a</sup>Additional experiments, which are not listed in this table have been performed with the model in which N, Fe, and P atmospheric depositions have been modified individually (set to 0 or current average deposition multiplied by 5). More details on the model setup are given in the Methods section. Atmospheric annual deposition is not constant over time as shown in Figure 1c. This was taken into consideration for example in the scenario "D" where monthly or annual deposition have been multiplied by 5: from year 1900 to year 2100, changes in atmospheric deposition have been calculated to vary by a factor of 2 to 5. The day of the pulse was chosen arbitrarily in January (summer conditions in southern hemisphere) and in July (summer conditions in northern hemisphere).

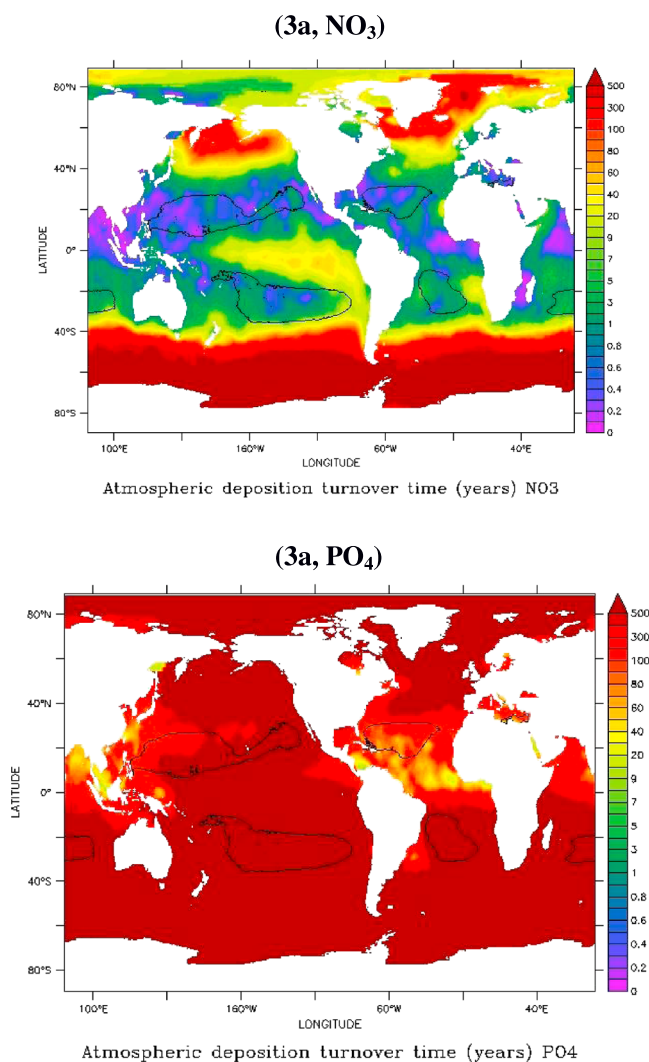
#### 2.4. Sensitivity Experiments

Using the quasi-steady state obtained from the spun-up ocean biogeochemical model, two different sets of sensitivity experiments are performed, as summarized in Table 2. In a first set of experiments, climatological monthly mean atmospheric depositions of P and Fe and climatological annual-mean atmospheric deposition of N ("Standard") are either multiplied by five ('D' scenarios) (based on the projection of the deposition fluxes presented on Figure 1c: from year 1900 to year 2100, changes in atmospheric deposition have been calculated to vary by a factor of 2 to 5) or set to zero (no "D" scenarios) for each nutrient independently and also for all three nutrients altogether. A total of 8 model runs were performed, 4 for the "D" scenarios and 4 for the no "D" scenarios. In each experiment, the model is integrated for 100 years. In a second set of experiments, the model has been integrated for 100 years using standard atmospheric deposition. During the last year (year 100), a strong pulse of dust ( $1 \text{ g m}^{-2} \text{ d}^{-1}$ ) and N ( $2.8 \text{ mg N m}^{-2} \text{ d}^{-1}$ ) is imposed arbitrarily on 15 January or on 15 July, everywhere in the LNL regions (derived from SeaWiFS climatology (1997–2007) with a criteria of Chl a  $< 0.07 \text{ mg.m}^{-3}$ , see Figure 1a). Fe and P input are related to dust deposition assuming a mean mass content of 3.5% [Jickells et al., 2005] and 0.07% [Guieu et al., 2002] and a solubility in surface seawater of 2% [Bonnet and Guieu, 2004] and 15% [Ridame and Guieu, 2002] for Fe and P, respectively. N deposition is assumed to be fully bioavailable as  $\text{NO}_3$  (a sensitivity study delivering N as ammonia did not produce significantly different results). Thus, the last year is run twice, with one pulse of deposition added to the standard deposition field. The magnitude of the pulse, both for dust and N, has been set to be of the same order of magnitude deposition range of observed episodic pulses (see references in introduction section) which were also used in aerosols/dust addition in bioassay experiments reported in Table 1 ( $0.1\text{--}20 \text{ g m}^{-2} \text{ event}^{-1}$ ). The pulsed model rate ( $1 \text{ g m}^{-2} \text{ d}^{-1}$ ) is representative of areas where moderate to strong episodic event is recorded (see Introduction), and represents the upper limit of the background rate of dust deposition [Lawrence and Neff, 2009]. Although less measurements are available for nitrogen deposition over the ocean,  $2.8 \text{ mg N m}^{-2} \text{ d}^{-1}$  is of the same order of magnitude as the strong deposition events observed at three stations in the North Atlantic [Prospero et al., 1996]. The mixing of dust with anthropogenic acids such as  $\text{HNO}_3$  (see Introduction) between emission and deposition regions will result in dust deposition enriched in nitrogen [Geng et al., 2009] which further confers a pulsed character in the fraction of N deposition associated with dust.

### 3. Results

#### 3.1. Impacts on Biota From Field and Laboratory Experiments: A Synthesis

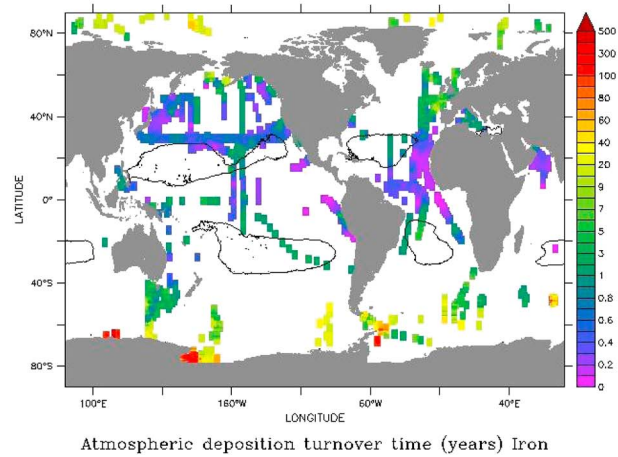
The dominant impacts of atmospheric deposition on biota in LNL surface ocean waters are summarized in a compilation of recently conducted dust/aerosol addition bioassay experiments (in situ, in vitro, and mesocosms, see Table 1). Response patterns of changes in standing stocks of organisms, their community structure, and metabolic rates, from comparison of aerosols treatments with the respective control incubations, are presented in Figure 2. Most data sets indicate positive responses to aerosol addition, with bacterial production and  $\text{N}_2$  fixation showing the strongest responses (average 234% and 193% increase, respectively). A mean 68% increase in Chl a is seen; however, differential responses among phytoplankton groups are also apparent,



**Figure 3.** Atmospheric deposition as a source of nutrients to the LNL ocean. (a) Turnover times (years) relative to atmospheric deposition (TTADs) for nitrate (NO<sub>3</sub>), phosphate (PO<sub>4</sub>), and dissolved iron (DFe). Areas with low values indicate that atmospheric supply plays an important role in maintaining nutrient concentrations. (b) Contribution of atmospheric deposition to the total vertical supply of nutrient to the surface mixed layer using both a model (NO<sub>3</sub> and PO<sub>4</sub>) and observational (DFe) data (see details in Methods section). Areas with high values indicate that atmospheric supply plays an important role in maintaining nutrient concentrations. Black contour: limits of the oligotrophic gyres (see Figure 1a).

with cyanobacteria *Synechococcus* and particularly *Prochlorococcus* showing weak responses to aerosol addition and nano- and micro-phytoplankton showing a similar increase to that of Chl-a, suggesting that these are related and that aerosol deposition may (temporally) support an increase in larger size class phytoplankton. Despite Chl-a increase in response to atmospheric deposition events in LNL waters, the natural tested waters typically remained close to oligotrophic conditions. This may help explain the variable and low response of satellite-derived chlorophyll signals to dust events [Volpe *et al.*, 2009]. Changes in standing stocks tend to be smaller than changes in metabolic rates as shown by comparison of Chl a vs primary production, and bacterial abundance vs bacterial respiration. It should be noted that most of the experiment studies presented in this compilation (Table 1) have been performed using desert dust and many of these have reported the composition of total nutrients content in the dust but few have considered their solubility and hence bioavailability. This is particularly important because desert dust from soils contain very little soluble nitrogen whereas “atmospherically processed dust” is rich in nitrogen. Differences in the source of the material used (e.g., soil vs locally collected aerosols), related composition, and solubility may partially explain the observed variability in response.

(3a, DFe)



(3b)

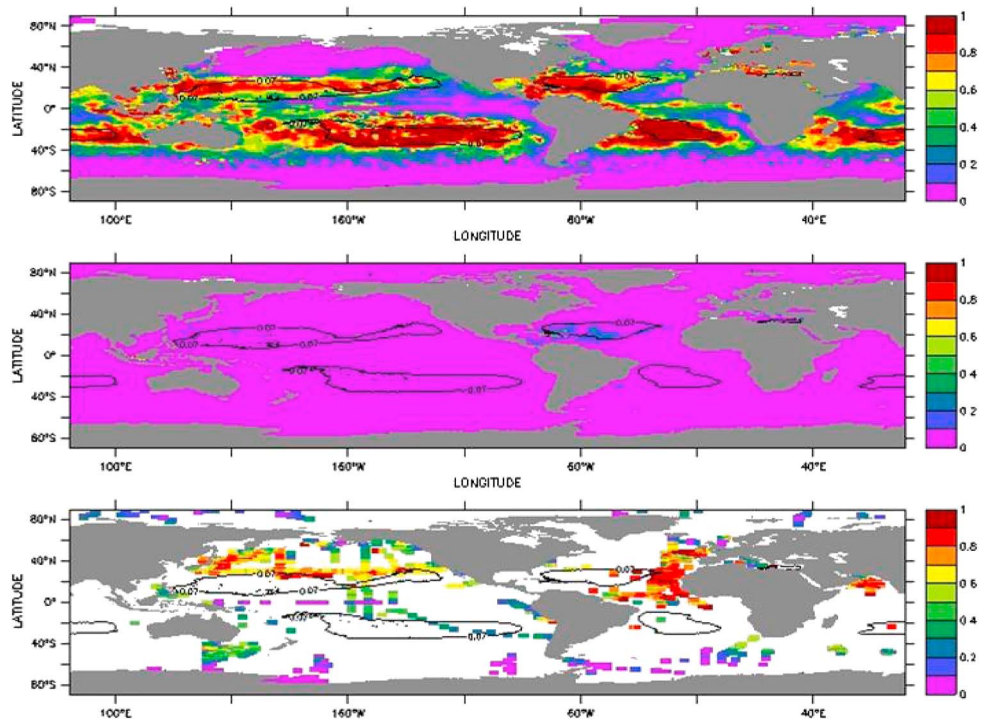
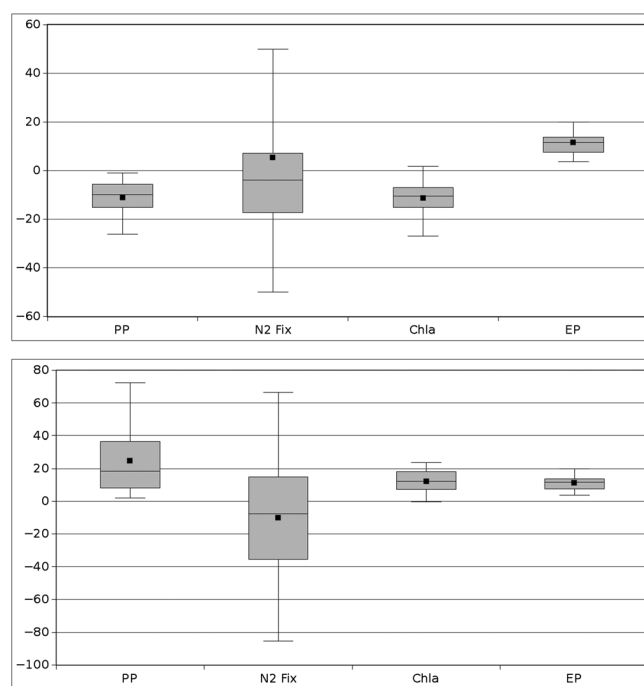


Figure 3. (continued)

**3.2. Contribution of Atmospheric Deposition to the Nutrient Reservoir in the Global Surface Ocean**

The mean TTAD in LNL regions are <1 month for nitrate; TTAD are orders of magnitude higher for phosphate and ~1 year for the few available iron values (Figure 3a). We note, however, that nutrient climatologies are poorly constrained in surface waters of LNL regions [reported concentrations often being set by the detection limits of conventional techniques; see for example Pulido-Villena *et al.*, 2010] leading to possible overestimation of the TTAD in particular for phosphate, as shown by Figure 3a. This identifies the need for more extensive databases for the deposition fluxes and surface nutrient concentrations; this will also allow to provide TTAD seasonal pattern.

The contributions of atmospheric deposition to the total supply of new nutrients to the surface mixed layer (e.g., sum of vertical and atmospheric supplies) for nitrate, phosphate, and iron, using both a model and



**Figure 4.** Impact of modification of the mean magnitude of atmospheric deposition on primary production (PP), nitrogen fixation ( $N_2$  fix), surface chlorophyll- $a$  (Chla) between surface and 100 m, and export production at 100 m depth (EP), in LNL regions as computed with the NEMO-PISCES models. Box-Whisker plots (box portion = interquartile range (25<sup>th</sup> to 75<sup>th</sup> percentile) of the data set. Horizontal bar within the box = median value. Black square = mean value. Also represented maximum and minimum values) showing relative changes (as percentage) with respect to the standard run using standard climatological atmospheric deposition fields (Figure 1b) described in the Methods and Table 2. These relative changes are computed from the maximum daily response, simulated over the year, following the change in atmospheric deposition (see Methods) and include changes in all nutrients together (total set of 4 experiments): when zero deposition is considered (No “D” experiment; upper panel) and when atmospheric deposition has been multiplied by 5 (“D” experiment; lower panel).

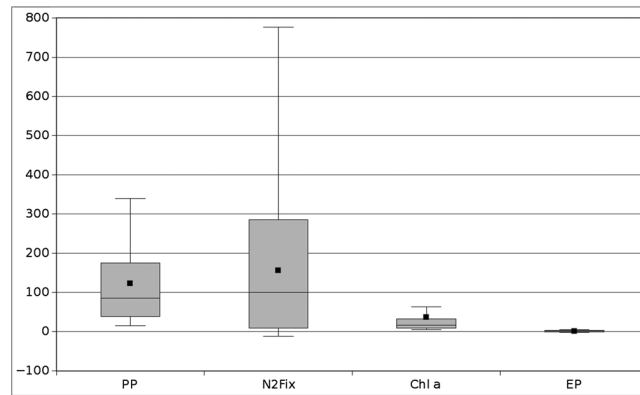
(relative change close to 0) and was modest for nitrogen fixation (maximum range % change  $-30 - +30\%$ ). However, when considering only LNL regions (Figure 4), primary production, export production, and  $N_2$  fixation were significantly impacted by atmospheric supply of nutrients (ranging between  $-40$  and  $+30\%$ ), although Chl  $a$  remained almost unchanged. Nevertheless, the simulations never reached the magnitude of responses observed in bioassay experiments (Figure 2) and in some cases showed responses of opposite sign, as with  $N_2$  fixation. This inconsistency between experiments and models cannot be explained by differences in temporal frequency of sampling, as we based our model analysis on daily outputs (see Methods), a similar sampling frequency to that of the bioassay experiments (see Table 1). Since the euphotic zone can be significantly deeper than 100 m in these LNL areas, we also did the analysis for export production at 200 m for the “D” and the “no D” experiments. Relative changes are almost identical to those computed from export production at 100 m (see Figure S12). We therefore refer to export production at the 100 m horizon in this study.

Results from the pulse experiment in which a strong pulse of dust and nitrogen deposition is imposed on 15 January or on 15 July, everywhere in the LNL regions, similar in magnitude to natural episodic high deposition events and to those used in the aerosols addition experiments (Table 1), were superimposed on the standard climatological atmospheric deposition fields in the LNL regions (dust =  $1 \text{ g m}^{-2} \text{ d}^{-1}$  and  $N = 2.8 \text{ mg N m}^{-2} \text{ d}^{-1}$ ; see Methods and Table 2). These indicate that the projected response to the combined deposition (e.g., present-day average deposition fields plus pulse; Figure 5) was typically much larger than in the “D” and no “D” models; the maximum response averaged over the LNL regions represented a change

observations are shown in Figure 3b. The calculated contribution is subject to the same shortcomings as the TTAD calculations with respect to the availability of data; however, a similar pattern is apparent indicating that atmospheric deposition could account for a significant proportion of the total input of new nutrients in the oligotrophic (LNL) gyres ( $>50\%$  for N;  $0-20\%$  for P and  $10-90\%$  for Fe) when lateral advection is not considered.

### 3.3. Evidence of Impacts on Biota in Models

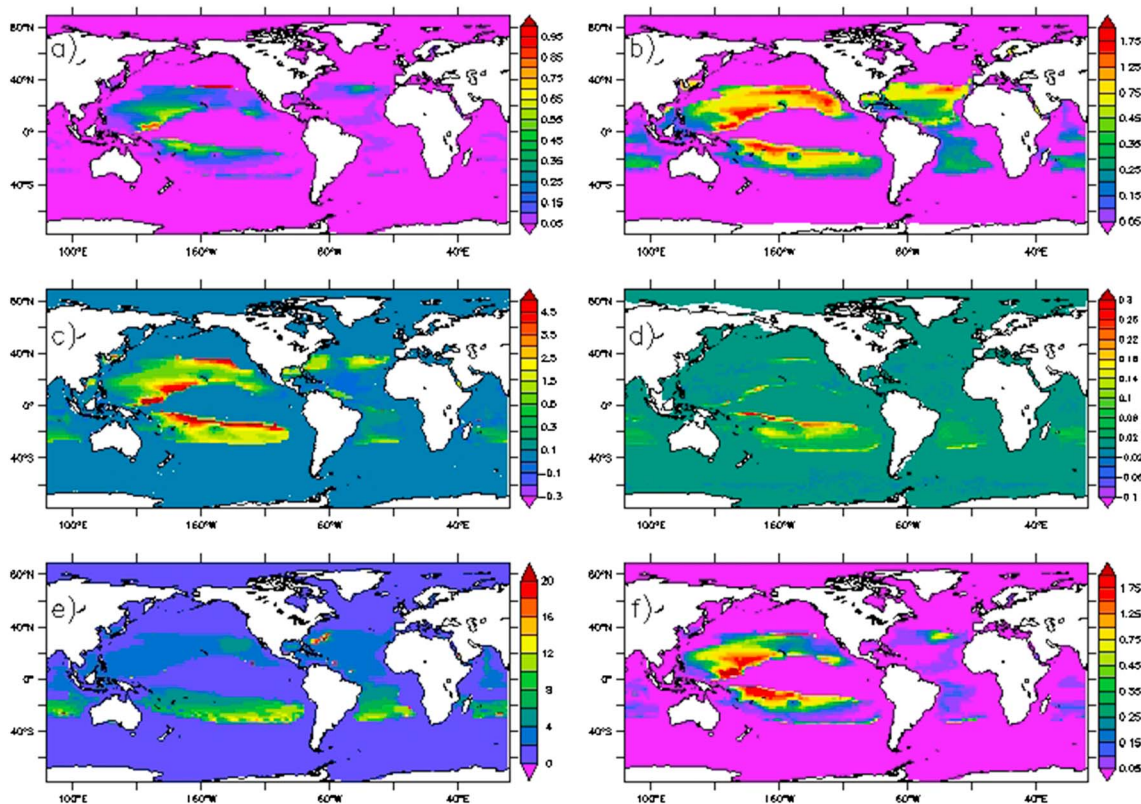
The sensitivity of ocean biogeochemistry to atmospheric deposition was investigated through a series of model experiments (Table 2 and Methods). The maximal daily anomaly simulated by the model over the last year (year 100) in response to the different scenarios was compared to the standard scenario (in which atmospheric deposition is not altered) by statistical analysis. In the first set of model experiments, atmospheric deposition of nitrate, phosphate, and iron was varied individually or in combination over a period of 100 years. For all experiments (both “D” and “no D” experiments) on a global scale (Figure S11), primary production, export production at 100 m, and Chl  $a$  did not differ significantly from the standard run



**Figure 5.** Impact of one pulse of dust and N (on 15 January or on 15 July) superimposed over the standard climatological atmospheric deposition fields, for primary production (PP), nitrogen fixation ( $N_2$  fix), surface chlorophyll a (Chl a) between surface and 100 m, and export production at 100 m depth (EP), in LNL regions computed using the NEMO-PISCES models (Pulse experiments, see Methods and Table 2). Box-Whisker plots (box portion = interquartile range (25<sup>th</sup> to 75<sup>th</sup> percentile) of the data set. Horizontal bar within the box = median value. Black square = mean value. Also represented maximum and minimum values) showing the maximum daily relative change averaged over the LNL regions (as percentage) relative to the Standard run.

of 123% for primary production, 154% for nitrogen fixation, and 36% for chlorophyll concentration. This also showed close agreement with the results of the bioassay experiments (Figure 5 vs Figure 2) for primary production (PP) and  $N_2$  fixation but still underestimated Chl a based biomass growth. It is perhaps intuitive that increasing the intensity of deposition on a short time scale may considerably increase the response. At the same time it is remarkable that those responses are close to the observations from bioassay experiments.

To better characterize the simulated responses to the pulses in our model, we show in Figure 6 the spatial structure of the surface changes induced by the deposition pulse imposed on 15 July (relative to the run without this pulse). A striking characteristic is the large



**Figure 6.** Surface changes (between surface and 100 m) induced by the pulse addition applied on 15 July (see Methods and Table 2). Maximum relative change in (a) surface Chl a; (b) surface NPP; (c) surface Nfix; (d) export at 100 m; and (e) Residence time (in days) of the most limiting nutrient. The residence time is defined as the duration of the perturbation of the considered nutrient and (f) maximum relative change in total grazing on phytoplankton.

spatial variability of the response, as inferred from the statistical analysis presented in Figure 5. (Figure 6; maximum response >100% for Chl a, >175% for PP), with more significant response in the oligotrophic gyres of the Pacific, reflecting the diverse local physical and biogeochemical conditions, as is apparent in bioassay experiments conducted at different locations or seasons (Table 1). The increase in PP is accompanied by a similar increase in grazing pressure. Interestingly, the strong pulse imposed in our model experiment is sufficient to trigger a strong increase in PP by small phytoplankton cells, but not by bigger species such as diatoms. Since small phytoplankton are kept in check by (micro-)zooplankton, these increased growth rates induce higher grazing rates and so do not result in accumulation of phytoplankton biomass, as illustrated by the changes in Chl a (Figure 6a).

Export production at 100 m remained generally unchanged (average 0% change) in these model experiments, despite some regionally significant changes in the South Pacific (Figure 6d). The relatively small response of export production to dust deposition may seem quite surprising, especially considering the increase in primary production and the sizable increase in chlorophyll. In fact, as supported by field and experimental observations reported in this paper (Table 1 and text), this is understandable in LNLC regions, due to the increased bacterial respiration and grazing.

The spatial variability highlighted in Figure 6 can be explained by the very diverse physical and biogeochemical conditions in LNLC regions. The responses to the pulse occur relatively rapidly and have a limited temporal extent, from a few days to about 2 weeks. They tend to be longer in winter time (in either hemisphere) because of the lower light levels, cooler surface temperature, and deeper mixed layer (the latter also resulting to lower mean available light for photosynthesis). Despite a generally strong local response, the pulses do not modify the system over the long term due to their short duration. Indeed, a very simple computation, over the mixed layer, assuming for instance a PP increase following a pulse of 100% for 7 days, the annual PP change would be  $100\% \times 7 / 365 \sim 2\%$ . The model indicates thus that, the “instantaneous” effect of a single pulse is strong but the annual effect is small.

## 4. Discussion

### 4.1. Significance of Atmospheric Nutrients in LNLC Regions

Atmospheric deposition supplies most of the new N and Fe to the mixed layer in some LNLC regions, based upon the comparison of atmospheric deposition with vertical supply from sub-surface waters in both model data and observations (Figure 3b). In the case of P, the atmospheric contribution is low and only significant in the North Atlantic subtropical gyre, and the Eastern Mediterranean Sea. However, surface water phosphate concentrations are overestimated in some regions due to lack of measurements, which leads to incorrect estimation of the vertical supply from below. Our analysis is only intended to be qualitative as it suffers from important shortcomings. For example lateral transport of dissolved inorganic and organic nutrients is not accounted for, although several studies have shown that this can be important, especially at the boundaries of the LNLC regions [i.e., *Williams and Follows*, 1998]. Also, the dynamic output from the coarse-resolution model does not resolve mesoscale and submesoscale dynamic processes, which may contribute a significant fraction of the vertical and lateral supply of new nutrients [i.e., *Williams and Follows*, 1998; *Pelegri et al.*, 2006; *Lévy*, 2008]. Our results are different from those by *Krishnamurthy et al.* [2010]; see their Figure 2), where atmospheric input is compared to the downward export by sinking particles at 103 m. Here we compare the atmospheric input of N, P, and Fe to the vertical input of these nutrients into the mixed layer by mixing, entrainment, and vertical advection. Both represented ratios do not display the same diagnostic and thus differ. In fact, the vertical supply of nutrients by the ocean dynamics can roughly equal the export at a certain depth only if long time scales are considered, if no nutrients are supplied as dissolved organic materials, and if lateral transport is negligible which is barely never the case in the ocean. An additional explanation to the differences with *Krishnamurthy et al.* [2010] is that our diagnostic is largely based on data whereas they mainly used model outputs. When using our model output to compute the ratio between the nutrient input by atmospheric deposition and the export at 100 m (Figure S13), as expected, it looks very similar to Figure 2 in *Krishnamurthy et al.* [2010].

### 4.2. Variability in Biological Response to Dust Addition in LNLC Regions

The compilation of the experimental results shows a strong variability in the response to aerosol inputs (Figure 2). A similarly high spatial and temporal variability is also simulated by our biogeochemical model (Figures 5 and 6).

Such variability is explained in the model by the heterogeneity in the physical, biogeochemical and ecosystem characteristics of the surface ocean. Nevertheless, additional processes that are not represented in models may generate variable responses in experiments and in the field. For example, the composition of aerosol deposition is spatially variable (see Figure 1b), and atmospheric inputs may provide the limiting nutrients to sustain productivity but may have additional stoichiometric effects through variability in N:P and Fe:P ratios [Moore *et al.*, 2013], and also the supply of microbes, other trace metals than iron, toxic elements, and other pollutants [i.e., Paytan *et al.*, 2009; Jordi *et al.*, 2012; Després *et al.*, 2012]. The organic nutrient content of atmospheric deposition, for which fewer data exist compared with the inorganic fraction [Kanakidou *et al.*, 2012], may also contribute to the observed variability.

Another cause of the observed variability is related to differences in the natural assemblages of organisms initially present in the bioassay incubation experiments as different LNLC areas do not exhibit the same nutrient limitation or co-limitation [Moore *et al.*, 2013]. This has been shown by Giovagnetti *et al.* [2013] who observed different changes in the composition and structure of the phytoplankton community and physiological state of the communities whether one dust seeding or a succession of dust seedings were performed in large mesocosm experiments. This confirmed that initial seawater conditions (and on-going nutrient availability) are in part controlling the response of the natural assemblage. In areas where elements in atmospheric deposition—such as copper—reach toxic levels, there could be a selective sensitivity to toxins in the deposited material that may cause a shift in species dominance [Paytan *et al.*, 2009] or even a decline in phytoplankton biomass over large areas of the ocean [Jordi *et al.*, 2012].

A striking result from both the experimental and field observations and the model experiments is the relatively larger changes in metabolic rates compared to changes in standing stocks. This can be attributed to the effect of grazing by zooplankton on phytoplankton and bacteria, which may increase turnover at the expense of stocks and also rapidly propagate responses through the whole food web [Bonnet *et al.*, 2005; Herut *et al.*, 2005; Marañón *et al.*, 2010]. This could result in experimental artifacts in that chlorophyll-*a* concentration may increase in experiments where grazers are filtered, but remain unchanged in natural waters. Phosphate addition to surface waters in the Eastern Mediterranean Sea during a Lagrangian experiment caused a negative Chl *a* response and an increase in abundance of heterotrophs (both bacteria and zooplankton) [Thingstad *et al.*, 2005]. Thus, in addition to the role of the grazers, the potential competitive advantage of bacteria may also explain the small increase of Chl *a* biomass. The larger increase in bacterial production compared to that of primary production indicates the intrinsically faster metabolism of nutrient-limited bacteria [Marañón *et al.*, 2010], and indicates an increase in organic matter remineralization and a corresponding reduction in carbon export.

The large increase in N<sub>2</sub> fixation indicates that diazotrophs may disproportionately benefit from a pulsed increase in dissolved iron and phosphorus, as observed in the eastern tropical North Atlantic [Mills *et al.*, 2004] or alternatively the contribution of other elements by dust, as observed in the Central Mediterranean Sea [Ridame *et al.*, 2011]. As the same dust was used in both experiments, then the apparent discrepancy between their results may reflect differences in the nutrient or biological status of the water at the start of the respective experiments.

### 4.3. Potential Underestimation of Atmospheric Impacts by Models

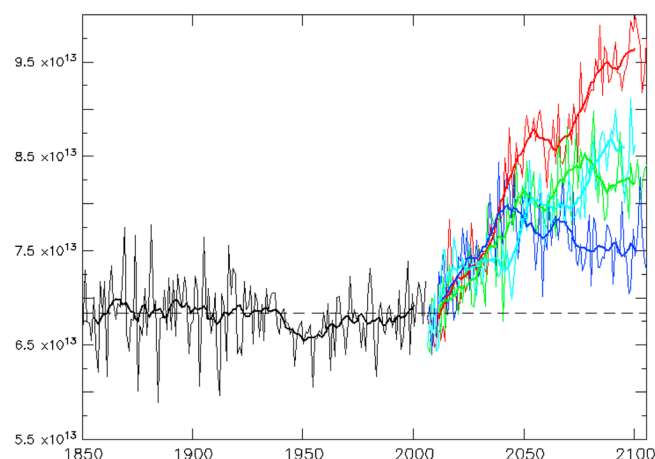
Experimental studies performed over short time periods, and thus representative of pulsed inputs, have highlighted the important role of atmospheric deposition for LNLC ecosystem functioning and in particular the activity of heterotrophic bacteria and diazotrophs. Biogeochemical modeling allows extrapolation of the impact of atmospheric deposition over larger space and time scales than laboratory/field experiments. In particular, modeling allows investigations into “cascading effects” (or feedbacks) between regions with contrasting responses and temporal shifts in limitation by different nutrients, and so the net impacts on nutrient cycling and carbon export. In the model scenarios using average atmospheric deposition on yearly or monthly time scales (Figure 4 “no D scenario”) and also fivefold increases (Figure 4 “D scenario”), the simulated responses on all timescales never reach the magnitudes observed in field and laboratory aerosol addition bioassay experiments (Figure 2) and at times, responses were even in an opposite direction, e.g., for N<sub>2</sub> fixation. In addition to highlighting possible problems in model structure or parameterizations (including deficiencies in the modeled dynamics such as the absence of meso- to submesoscale processes), the disagreement between models and observations may be due to feedbacks that operate at large spatial

and temporal scales that are considered in models but excluded in short-duration field or laboratory experiments. Another possibility is that models driven by monthly mean or annual-mean atmospheric deposition fields do not capture the highly episodic nature of atmospheric deposition. Indeed, a better agreement between experimental data and models is obtained when a strong pulse of deposition, similar in magnitude to that observed in episodic deposition events (see ref. in introduction) and simulated during field and laboratory aerosol addition bioassay experiments, was superimposed over the standard climatological atmospheric deposition fields in the LNLC regions (Figure 5). The observed agreement in the responses of primary productivity and  $N_2$  fixation following deposition events suggests that atmospheric impacts have so far been strongly underestimated by models, at least on synoptic timescales. These results have direct implications for our understanding of productivity in LNLC regions (e.g., dominant role of atmospheric pulse events).

## 5. Conclusions and Recommendations

Our new model analysis suggests that at least some of the discrepancies in the biogeochemical response to atmospheric deposition between published global model studies (which are similar in their setup to our standard and “D”/no “D” model experiments) and field and laboratory aerosol addition bioassay experiments may be due to differences in the time-scale and the mode of the atmospheric supply. The field and laboratory experiments document important aspects that may not have been captured properly by models. This stresses the need to improve the representation of key processes brought into play by atmospheric deposition in ocean biogeochemical models. Indeed, experimental studies show that the effects of atmospheric deposition on surface ocean productivity in LNLC areas appear to be more complex than a simple, overall “fertilization effect” of increasing phytoplankton biomass that is typical of HNLC regions. The term “fertilization” is often associated with the a priori belief that dust deposition should increase chlorophyll biomass and carbon sequestration (and thus increases atmospheric  $CO_2$  drawdown). Recent experimental studies [Marañón *et al.*, 2010; Guieu *et al.*, 2014b] have changed the way we understand dust deposition to the oligotrophic ocean by showing that fertilization predominantly enhances heterotrophic bacterial growth, and thus dust deposition induces the remineralization of DOC, and so reduces atmospheric  $CO_2$  drawdown. Consequently it is not surprising that dust deposition is not typically followed by phytoplankton bloom or significant carbon export in LNLC regions. This is apparent even without the consideration of grazing, as demonstrated by the model. Specifically, the strong sensitivity of bacterial production to aerosol addition, and associated competitive interactions between phytoplankton and heterotrophic bacteria, needs to be better represented in models. Another recommendation is to increase the number of modeled functional groups, given that different phytoplankton groups show differential responses to aerosol addition in incubation experiments [ie Paytan *et al.*, 2009; Giovagnetti *et al.*, 2013] (Figure 2). Furthermore, models need to include variable nutrient elemental ratios in atmospheric deposition, in the water column and in organic matter [Krishnamurthy *et al.*, 2010; Moore *et al.*, 2013]. Remaining disagreement in the predicted magnitude of responses between the bioassays and pulsed model may be attributed to processes that are not represented in models. For example, inclusion of relevant processes, such as the impact of aerosol deposition on aggregation and vertical flux, may result in an increase in export in pulsed models. Indeed, experiments and observations have previously indicated that atmospheric deposition can enhance export of particulate organic carbon (POC) to the deep ocean in LNLC regions by facilitating aggregation processes and providing ballast [Armstrong *et al.*, 2002; Ternon *et al.*, 2010; Bressac *et al.*, 2011; Bressac and Guieu, 2013]. This can induce a strong and rapid POC export that is independent of a fertilization effect [Ternon *et al.*, 2010; Bressac and Guieu, 2013; Bressac *et al.*, 2014] that also results in scavenging of trace metals [Wagener *et al.*, 2010; Wuttig *et al.*, 2013; Bressac and Guieu, 2013]. In addition models should be carefully compared to short-duration experiments to test whether they properly represent the key processes brought into play by aerosol addition on all relevant time scale and modes of addition.

These fundamental differences strongly argue in favor of coordinated efforts between modelers and experimentalists to improve models, controlled experiments, and field data, and to test model results in experimental design. In particular, we did not address in the study the annual effect of a series of pulses: only one pulse was superimposed over the climatological atmospheric deposition; doing a series of pulses in the model would be the next step, based for example on statistics from atmospheric data collected at ie BATS, CVOO, and DYFAMED time series.



**Figure 7.** Evolution of the size of oligotrophic areas over 1850–2100, for the historical period (black) and four representational concentration pathways (RCPs) scenarios (RCP2.6 blue, RCP4.5 green, RCP6.0 light blue, and RCP8.5 red).

scenarios and recent satellite observations suggest that anthropogenic global warming may induce an increase in the size of the oligotrophic gyres in all ocean basins [Henson *et al.*, 2010; Steinacher *et al.*, 2010; Polovina *et al.*, 2008] (Figure 7). In most current models, atmospheric deposition of nutrients is kept constant in time and space over the course of the model simulations, and the modeled changes primarily reflect alterations in ocean physics (enhanced stratification, reduced mixed layer depth, and slowed circulation) that reduce the supply of macronutrient (N and P) from below the thermocline. The combination of changes in atmospheric deposition and expansion of the LNLC areas, and associated decrease in supply from below the mixed layer, could result in a shift in the relative availability of different nutrients increasing the relative importance of the atmospheric inputs for ocean biogeochemistry.

We have considered the impact of short variability in aerosol supply to LNLC regions by comparing the response to pulsed addition of aerosols using two different approaches: model and bioassay experiments. These show similar magnitude responses for a number of parameters that are an order of magnitude greater than mean deposition-based models. This confirms that the episodic nature of atmospheric deposition needs to be considered to understand and model temporal variability in LNLC productivity. Improved representation of the effects of deposition in models is clearly important, particularly as (a) atmospheric deposition and the oligotrophic ocean area will change in the future and (b) the episodic nature of atmospheric deposition may have further significant effects that are currently not considered.

#### Acknowledgments

This article was initiated in Istanbul, Turkey during a 2 day workshop financed by the European CoOperation in Science and Technology (COST) Action 735 “Tools for assessing global air-sea fluxes of climate and air pollution relevant gases”. It is a contribution to the The International Surface Ocean - Lower Atmosphere Study (SOLAS) project. We thank L. Legendre and P. Boyd for insightful comments on an earlier version of this manuscript. We also thank the two anonymous reviewers for their helpful review. Model output can be obtained from Olivier Aumont (Olivier.Aumont@ird.fr). The corresponding code can be downloaded from the official NEMO website (<http://www.nemo-ocean.eu/>).

#### References

- Antoine, D., J. M. André, and A. Morel (1996), Oceanic primary production: 2. Estimation at global scale from satellite (coastal zone color scanner) chlorophyll, *Global Biogeochem. Cycles*, *10*(1), 57–69.
- Archer, D. E., and K. Johnson (2000), A model of the iron cycle in the ocean, *Global Biogeochem. Cycles*, *14*, 269–279, doi:10.1029/1999GB900053.
- Armstrong, R. A., C. Lee, J. I. Hedges, S. Honjo, and S. G. Wakeham (2002), A new, mechanistic model for organic carbon fluxes in the ocean based on the quantitative association of POC with ballast minerals, *Deep Sea Res., Part II*, *49*, 219–236.
- Aumont, O., and L. Bopp (2006), Globalizing results from ocean in situ iron fertilization studies, *Global Biogeochem. Cycles*, *20*, GB2017, doi:10.1029/2005GB002591.
- Aumont, O., L. Bopp, and M. Schultz (2008), What does temporal variability in aeolian dust deposition contributes to iron and chlorophyll distributions?, *Geophys. Res. Lett.*, *35*, L07607, doi:10.1029/2007GL031131.
- Baker, A. R., and T. D. Jickells (2006), Mineral particle size as a control on aerosol iron solubility, *Geophys. Res. Lett.*, *33*, L17608, doi:10.1029/2006GL026557.
- Blain, S., C. Guieu, H. Claustré, K. Leblanc, T. Moutin, B. Quéguiner, and G. Sarthou (2004), Availability of iron for phytoplankton in the north-east Atlantic Ocean, *Limnol. Oceanogr.*, *49*, 2095–2104.
- Bonnet, S., and C. Guieu (2004), Dissolution of atmospheric iron in seawater, *Geophys. Res. Lett.*, *31*, L03303, doi:10.1029/2003GL018423.
- Bonnet, S., and C. Guieu (2006), Atmospheric forcing on the annual iron cycle in the Mediterranean Sea. A one-year survey, *J. Geophys. Res.*, *111*, C09010, doi:10.1029/2005JC003213.
- Bonnet, S., C. Guieu, J. Chavierini, J. Ras, and A. Stock (2005), Effect of atmospheric nutrients on the autotrophic communities in a low nutrient, low chlorophyll system, *Limnol. Oceanogr.*, *50*, 1810–1819.
- Bonnet, S., et al. (2008), Nutrient limitation of primary productivity in the Southeast Pacific (BIOSEPE cruise), *Biogeosciences*, *5*, 215–22.
- Bopp, L., K. E. Kohfeld, C. Le Quéré, and O. Aumont (2003), Dust impact on marine biota and atmospheric CO<sub>2</sub> during glacial periods, *Paleoceanography*, *18*, 1046, doi:10.1029/2002PA000810.

Improving our knowledge and understanding of the impacts of atmospheric deposition in LNLC regions and their accurate representation in biogeochemical models is critical as climate models predict changes in both the magnitude and distribution of atmospheric nutrient deposition, and in the size and intensity of LNLC regions. Indeed, deposition of iron, nitrogen, phosphorus, and organic matter has increased considerably since preindustrial times [Duce *et al.*, 2008; Mahowald *et al.*, 2008, 2010], and nitrogen could further slightly increase in the future (Figure 1c). At the same time climate model simulations forced by IPCC-type

- Boyd, P. W., et al. (2007), A synthesis of mesoscale iron enrichment experiments 1993–2005: Key findings and implications for ocean biogeochemistry, *Science*, *315*, 612–617.
- Bressac, M., and C. Guieu (2013), Post-depositional processes: What really happens to new atmospheric iron in the ocean surface?, *Global Biogeochem. Cycles*, *27*, 859–870, doi:10.1002/gbc.20076.
- Bressac, M., C. Guieu, D. Doxaran, F. Bourrin, G. Obolensky, and J. M. Grisoni (2011), A mesocosm experiment coupled with optical measurements to observe the fate and sinking of atmospheric particles in clear oligotrophic waters, *Geo Mar. Lett.*, doi:10.1007/s00367-011-0269-4.
- Bressac, M., C. Guieu, D. Doxaran, F. Bourrin, K. Desboeufs, N. Leblond, and C. Ridame (2014), Quantification of the lithogenic carbon pump following a simulated dust-deposition event in large mesocosms, *Biogeosciences*, *11*, 1007–1020.
- Capone, D. G., J. A. Burns, J. P. Montoya, A. Subramaniam, C. Mahaffey, T. Gunderson, A. F. Michaels, and E. J. Carpenter (2005), N<sub>2</sub> fixation by *Trichodesmium* spp.: An important source of new nitrogen to the tropical and subtropical North Atlantic Ocean, *Global Biogeochem. Cycles*, *19*, GB2024, doi:10.1029/2004GB002331.
- Cho, B. C., and F. Azam (1990), Biogeochemical significance of bacterial biomass in the ocean's euphotic zone, *Mar. Ecol. Prog. Ser.*, *63*, 253–259.
- Chung, C. C., J. Chang, G. C. Gong, S. C. Hsu, K. P. Chiang, and C. W. Liao (2011), Effects of Asian dust storms on *Synechococcus* populations in the subtropical Kuroshio Current, *Mar. Biotechnol.*, *13*(4), 751–763.
- Cornell, S. E. (2011), Atmospheric nitrogen deposition: Revisiting the importance of the organic component, *Environ. Pollut.*, *159*, 2214–2222.
- de Boyer Montégut, C., G. Madec, A. S. Fischer, A. Lazar, and D. Iudicone (2004), Mixed layer depth over the global ocean: An examination of profile data and a profile-based climatology, *J. Geophys. Res.*, *109*, C12003, doi:10.1029/2004JC002378.
- Desboeufs, K. V., R. Losno, and J. L. Colin (2001), Factors influencing aerosol solubility during cloud processes, *Atmos. Environ.*, *35*, 3529–3537.
- Després, V. R., et al. (2012), Primary biological aerosol particles in the atmosphere: A review, *Tellus, Ser. B*, *64*, 15,598, doi:10.3402/tellusb.v64i0.15598.
- Donaghay, P. L., P. S. Liss, R. A. Duce, D. R. Kester, A. K. Hanson, T. Villareal, N. W. Tindale, and D. J. Gifford (1991), The role of episodic atmospheric nutrient inputs in the chemical and biological dynamics of oceanic ecosystems, *Oceanography*, *4*, 62–70.
- Duce, R. A. J., et al. (2008), Impacts of atmospheric anthropogenic nitrogen on the open ocean, *Science*, *320*, 893–897.
- Duce, R. A., et al. (1991), The atmospheric input of trace species to the world ocean, *Global Biogeochem. Cycles*, *5*, 193–259, doi:10.1029/91GB01778.
- Dutkiewicz, S., M. J. Follow, and P. Parekh (2005), Interactions of the iron and phosphorus cycles: A three-dimensional model study, *Global Biogeochem. Cycles*, *19*, GB1021, doi:10.1029/2004GB002342.
- Ellwood, M. J., C. S. Law, J. Hall, E. M. S. Woodward, R. Strzepek, J. Kuparinen, K. Thompson, S. Pickmere, P. Sutton, and P. W. Boyd (2013), Relationships between nutrient stocks and inventories and phytoplankton physiological status along an oligotrophic meridional transect in the Tasman Sea, *Deep Sea Res., Part I*, *72*, 102–120.
- Foster, R. A., A. Paytan, and J. P. Zehr (2009), Seasonality of N<sub>2</sub> fixation and nifH gene diversity in the Gulf of Aqaba (Red Sea), *Limnol. Oceanogr.*, *54*, 219–233.
- Friedlingstein, P., et al. (2006), Climate–carbon cycle feedback analysis: Results from the C4MIP model intercomparison, *J. Clim.*, *19*, 3337–3353.
- Fung, I. Y., S. K. M. Tegen, S. C. Doney, J. G. John, and J. K. B. Bishop (2000), Iron supply and demand in the upper ocean, *Global Biogeochem. Cycles*, *14*, 281–295, doi:10.1029/1999GB900059.
- Geng, H., Y. Park, H. Hwang, S. Kang, and C. U. Ro (2009), Elevated nitrogen-containing particles observed in Asian dust aerosol samples collected at the marine boundary layer of the Bohai Sea and the Yellow Sea, *Atmos. Chem. Phys.*, *9*, 6933–6947.
- Giovagnetti, V., C. Brunet, F. Conversano, F. Tramontano, I. Obernosterer, C. Ridame, and C. Guieu (2013), Assessing the role of dust deposition on phytoplankton ecophysiology and succession in a low-nutrient low-chlorophyll ecosystem: A mesocosm experiment in the Mediterranean Sea, *Biogeosciences*, *10*, 2973–2991.
- Guerzoni, S., et al. (1999), The role of atmospheric deposition in the biogeochemistry of the Mediterranean, *Sea, Prog. Oceanogr.*, *44*, 147–190.
- Guieu, C., M.-D. Loÿe-Pilot, C. Ridame, and C. Thomas (2002), Chemical characterization of the Saharan dust end-member; some biological implications for the western Mediterranean, *J. Geophys. Res.*, *107*(D15), 4258, doi:10.1029/2001JD000582.
- Guieu, C., et al. (2010a), Large clean mesocosms and simulated dust deposition: A new methodology to investigate responses of marine oligotrophic ecosystems to atmospheric inputs, *Biogeosciences*, *7*, 2765–2784, doi:10.5194/bg-7-2765-2010.
- Guieu, C., M.-D. Loÿe-Pilot, L. Benyaya, and A. Dufour (2010b), Spatial variability of atmospheric fluxes of metals (Al, Fe, Cd, Zn and Pb) and phosphorus over the whole Mediterranean and from a one year monitoring experiment: Biogeochemical implications, *Mar. Chem.*, *120*, 164–178.
- Guieu, C., F. Dulac, C. Ridame, and P. Pondaven (2014a), Introduction to project DUNE, a DUst experiment in a low Nutrient, low chlorophyll Ecosystem, *Biogeosciences*, *11*, 425–442, doi:10.5194/bg-11-425-2014.
- Guieu, C., C. Ridame, E. Pulido-Villena, M. Bressac, K. Desboeufs, and F. Dulac (2014b), Impact of dust deposition on carbon budget: A tentative assessment from a mesocosm approach, *Biogeosciences*, *11*, 5621–5635.
- Henson, S. A., J. L. Sarmiento, J. P. Dunne, L. Bopp, I. D. Lima, S. C. Doney, J. John, and C. Beaulieu (2010), Detection of anthropogenic climate change in satellite records of ocean chlorophyll and productivity, *Biogeosciences*, *7*, 621–640.
- Herut, B., T. Zohary, M. D. Krom, R. F. Mantoura, P. Pitta, S. Psarra, F. Rassoulzadegan, T. Tanaka, and T. F. Thingstad (2005), Response of East Mediterranean surface water to Saharan dust: On-board microcosm experiment and field observations, *Deep Sea Res., Part II*, *52*, 3024–3040.
- Jickells, T. D., et al. (2005), Global iron connections between desert dust, ocean biogeochemistry, and climate, *Science*, *308*, 67–71.
- Jordi, A., G. Basterretxea, A. Tovar-Sánchez, A. Alastuey, and X. Querol (2012), Copper aerosols inhibit phytoplankton growth in the Mediterranean Sea, *Proc. Natl. Acad. Sci. U.S.A.*, *109*, 21,246–21,249.
- Kanakidou, M. R. A., et al. (2012), Atmospheric fluxes of organic N and P to the global ocean, *Global Biogeochem. Cycles*, *26*, GB3026, doi:10.1029/2011GB004277.
- Krishnamurthy, A., J. K. Moore, C. S. Zender, and C. Luo (2007), Effects of atmospheric inorganic nitrogen deposition on ocean biogeochemistry, *J. Geophys. Res.*, *112*, G02019, doi:10.1029/2006JG000334.
- Krishnamurthy, A., J. K. Moore, N. Mahowald, C. Luo, S. C. Doney, K. Lindsay, and C. S. Zender (2009), Impacts of increasing anthropogenic soluble iron and nitrogen deposition on ocean biogeochemistry, *Global Biogeochem. Cycles*, *23*, GB3016, doi:10.1029/2008GB003440.
- Krishnamurthy, A., J. K. Moore, N. Mahowald, C. Luo, and C. S. Zender (2010), Impacts of atmospheric nutrient inputs on marine biogeochemistry, *J. Geophys. Res.*, *115*, G01006, doi:10.1029/2009JG001115.
- Laghdass, M., S. Blain, M. Besseling, P. Catala, C. Guieu, and I. Obernosterer (2011), Effects of Saharan dust on the microbial community, during a large in situ mesocosm experiment in the NW Mediterranean Sea, *Aquat. Microb. Ecol.*, *62*, 201–213.
- Lamarque, J. F., G. P. Kyle, M. Meinshausen, K. Riahi, S. Smith, D. P. van Vuuren, A. Conley, and F. Vitt (2011), Global and regional evolution of short-lived radiatively-active gases and aerosols in the Representative Concentration Pathways, *Clim. Change*, *109*, 191–212.
- Lamarque, J.-F., et al. (2010), Historical (1850–200) gridded anthropogenic and biomass burning emissions of reactive gases and aerosols: Methodology and application, *Atmos. Chem. Phys.*, *10*, 7017–7039.

- Law, C. S., E. R. Abraham, A. J. Watson, and M. I. Liddicoat (2003), Vertical diffusion and nutrient supply to the surface mixed layer of the Antarctic Circumpolar Current, *J. Geophys. Res.*, *108*(C8), 3272, doi:10.1029/2002JC001604.
- Law, C. S., E. M. S. Woodward, M. J. Ellwood, A. Marriner, S. J. Bury, and K. A. Safi (2011), Response of surface nutrient inventories and nitrogen fixation to a tropical cyclone in the South-West Pacific, *Limnol. Oceanogr.*, *56*, 1372–1385.
- Lawrence, C. R., and J. C. Neff (2009), The contemporary physical and chemical flux of aeolian dust: A synthesis of direct measurements of dust deposition, *Chem. Geol.*, *267*, 46–63.
- Le Quéré, C., et al. (2009), Trends in the sources and sinks of carbon dioxide, *Nat. Geosci.*, *2*, 831–836, doi:10.1038/ngeo689.
- Le Quere, C., et al. (2013), The global carbon budget 1959–2011, *Earth Syst. Sci. Data*, *5*, 165–185.
- Lekunberri, I., T. Lefort, E. Romero, E. Vázquez-Domínguez, C. Romera-Castillo, C. Marrasé, F. Peters, M. Weinbauer, and J. M. Gasol (2010), Effects of a dust deposition event on coastal marine microbial abundance and activity, bacterial community structure and ecosystem function, *J. Plankton Res.*, *32*, 381–396.
- Lévy, M. (2008), The modulation of biological production by oceanic mesoscale turbulence, Transport and Mixing in Geophysical Flows, *Lect. Notes Phys.*, *744*, 219–261, doi:10.1007/978-3-540.
- Lojze-Pilot, M. D., and J. Martin (1996), Saharan dust input to the Western Mediterranean: An eleven years record in Corsica, in *The Impact of Desert Dust Across the Mediterranean*, pp. 191–199, Springer, Netherlands.
- Luo, C., N. Mahowald, T. Bond, P. Y. Chuang, P. Artaxo, R. Siefert, Y. Chen, and J. Schauer (2008), Combustion iron distribution and deposition, *Global Biogeochem. Cycles*, *22*, GB1012, doi:10.1029/2007GB002964.
- Mackey, K. R. M., R. G. Labiosa, M. Calhoun, J. H. Street, A. Post, and A. Paytan (2007), Phosphorus availability, phytoplankton community dynamics, and taxon-specific phosphorus status in the Gulf of Aqaba, Red Sea, *Limnol. Oceanogr.*, *52*, 873–885.
- Mackey, K. R. M., G. L. van Dijken, S. Mazloom, A. M. Erhardt, J. Ryan, K. R. Arrigo, and A. Paytan (2010), Influence of atmospheric nutrients on primary productivity in a coastal upwelling region, *Global Biogeochem. Cycles*, *24*, GB4027, doi:10.1029/2009GB003737.
- Mackey, K. R., K. N. Buck, J. R. Casey, A. Cid, M. W. Lomas, Y. Söhrin, and A. Paytan (2012), Phytoplankton responses to atmospheric metal deposition in the coastal and open-ocean Sargasso Sea, *Front. Microbiol.*, *3*, doi:10.3389/fmicb.2012.00359.
- Madec, G. (2008), "NEMO ocean engine" Note du Pole de modélisation, Institut Pierre-Simon Laplace (IPSL), France, No 27 ISSN No 1288–1619.
- Mahowald, N. (2007), Anthropocene changes in desert area: Sensitivity to climate model predictions, *Geophys. Res. Lett.*, *34*, L18817, doi:10.1029/2007GL030472.
- Mahowald, N., et al. (2008), The global distribution of atmospheric phosphorus deposition and anthropogenic impacts, *Global Biogeochem. Cycles*, *22*, GB4026, doi:10.1029/2008GB003240.
- Mahowald, N., et al. (2009), Atmospheric Iron deposition: Global distribution, variability and human perturbations, *Annu. Rev. Mar. Sci.*, *1*, 245–278, doi:10.1146/annurev.marine.010908.163727.
- Mahowald, N., et al. (2010), Observed 20th century desert dust variability: Impact on climate and biogeochemistry, *Atmos. Chem. Phys.*, *10*, 10,875–10,893.
- Mahowald, N., K. Lindsay, D. Rothenberg, S. C. Doney, J. K. Moore, P. Thornton, J. Randerson, and C. Jones (2011), Desert dust and anthropogenic aerosol interactions in the Community Climate System Model coupled-carbon-climate model, *Biogeosciences*, *8*, 387–414.
- Marañón, E., et al. (2010), Degree of oligotrophy controls the response of microbial plankton to Saharan dust, *Limnol. Oceanogr.*, *55*(6), 2339–2352.
- Mills, M., C. Ridame, M. Davey, J. La Roche, and R. J. Geider (2004), Iron and phosphorus co-limit nitrogen fixation in the eastern tropical North Atlantic, *Nature*, *429*, 292–294.
- Moore, C. M., M. M. Mills, A. Milne, R. Langlois, E. P. Achterberg, K. Lochte, R. J. Geider, and J. La Roche (2006), Iron limits primary productivity during spring bloom development in the central North Atlantic, *Global Change Biol.*, *12*, 626–634.
- Moore, C. M., et al. (2013), Processes and patterns of oceanic nutrient limitation, *Nat. Geosci.*, doi:10.1038/ngeo1765.
- Moore, J. K., S. C. Doney, and K. Lindsay (2002), Upper ocean ecosystem dynamics and iron cycling in a global three-dimensional model, *Global Biogeochem. Cycles*, *18*, GB4028, doi:10.1029/2004GB002220.
- Moxim, W. J., S. M. Fan, and H. Levy (2011), The meteorological nature of variable soluble iron transport and deposition within the North Atlantic Ocean basin, *J. Geophys. Res.*, *116*, D03203, doi:10.1029/2010JD014709.
- Neuer, S., M. E. Torres-Padrón, M. D. Gelado-Caballero, M. J. Rueda, J. Hernández-Brito, R. Davenport, and G. Wefer (2004), Dust deposition pulses to the eastern subtropical North Atlantic gyre: Does ocean's biogeochemistry respond?, *Global Biogeochem. Cycles*, *18*, GB4020, doi:10.1029/2004GB002228.
- Parekh, P., S. Rutkiewicz, M. J. Follows, and T. Ito (2006), Atmospheric carbon dioxide in a less dusty world, *Geophys. Res. Lett.*, *33*, L03610, doi:10.1029/2005GL025098.
- Paytan, A., K. R. M. Mackey, Y. Chen, I. D. Limac, S. C. Doney, N. Mahowald, R. Labiosa, and A. F. Postf (2009), Toxicity of atmospheric aerosols on marine phytoplankton, *Proc. Natl. Acad. Sci. U.S.A.*, doi:10.1073/pnas.08114868106.
- Pelegri, J. L., A. Marrero-Díaz, and A. W. Ratsimandresy (2006), Nutrient irrigation of the North Atlantic, *Prog. Oceanogr.*, *70*, 366–406.
- Polovina, J. J., E. A. Howell, and M. Abecassis (2008), Ocean's least productive waters are expanding, *Geophys. Res. Lett.*, *35*, L03618, doi:10.1029/2007GL031745.
- Prospero, J. M., and P. J. Lamb (2003), African droughts and dust transport to the Caribbean: Climate change implications, *Science*, *302*, 1024, doi:10.1126/science.1089915.
- Prospero, J. M., K. Barrett, T. Church, F. Dentener, R. A. Duce, J. N. Galloway, and P. Quinn (1996), Atmospheric deposition of nutrients to the North Atlantic Basin, in *Nitrogen Cycling in the North Atlantic Ocean and its Watersheds*, pp. 27–73, Springer, Netherlands.
- Pulido-Villena, E., T. Wagener, and C. Guieu (2008), Bacterial response to dust pulses in the western Mediterranean: Implications for carbon cycling in the oligotrophic ocean, *Global Biogeochem. Cycles*, *22*, GB1020, doi:10.1029/2007GB003091.
- Pulido-Villena, E., V. Rerolle, and C. Guieu (2010), Transient fertilizing effect of dust in P-deficient LNLC surface ocean, *Geophys. Res. Lett.*, *37*, L01603, doi:10.1029/2009GL041415.
- Pulido-Villena, E., A.-C. Baudoux, I. Obernosterer, M. Landa, J. Caparros, P. Catala, C. Georges, J. Harmand, and C. Guieu (2014), Microbial food web dynamics in response to a Saharan dust event: Results from a mesocosm study in the oligotrophic Mediterranean Sea, *Biogeosciences*, *11*, 5607–5619.
- Ridame, C. (2001), Rôle des apports atmosphériques d'origine continentale dans la biogéochimie marine: Impact des apports sahariens sur la production primaire en Méditerranée (Doctoral dissertation), Univ. Paris 6, 297 pp. Villefranche sur Mer, France.
- Ridame, C., and C. Guieu (2002), Saharan input of phosphorus to the oligotrophic water of the open western Mediterranean, *Limnol. Oceanogr.*, *47*(3), 856–869.
- Ridame, C., M. Le Moal, C. Guieu, E. Ternon, I. Biegala, S. L'Helguen, and M. Pujo-Pay (2011), Nutrient control of N<sub>2</sub> fixation in the oligotrophic Mediterranean Sea and the impact of Saharan dust events, *Biogeosciences*, *8*, 2773–2783.

- Ridame, C., C. Guieu, and S. L'Helguen (2013), Strong stimulation of N<sub>2</sub> fixation to contrasted Saharan dust events in a Low Nutrient-Low Chlorophyll environment: Results from dust addition in large mesocosms, *Biogeosciences*, *10*, 7333–7346.
- Steinacher, M., et al. (2010), Projected 21st century decrease in marine productivity: A multi-model analysis, *Biogeosciences*, *7*, 919–1005.
- Tagliabue, A., T. Mtshali, O. Aumont, A. R. Bowie, M. Klunder, A. N. Roychoudhury, and S. Swart (2012), A global compilation of over 13,000 dissolved iron measurements: Focus on distribution and processes in the Southern Ocean, *Biogeosciences*, *9*, 2333–2349, doi:10.5194/bg-9-2333-2012.
- Ternon, E., C. Guieu, M.-D. Loye-Pilot, N. Leblond, E. Bosc, B. Gasser, J.-C. Miquel, and J. Martin (2010), The impact of Saharan dust on the particulate export in the water column of the North Western Mediterranean Sea, *Biogeosciences*, *7*, 809–826, doi:10.5194/bg-7-809-2010.
- Ternon, E., C. Guieu, C. Ridame, S. L'Helguen, and P. Catala (2011), Longitudinal variability of the biogeochemical role of Mediterranean aerosols in the Mediterranean Sea, *Biogeosciences*, *8*, 1067–1080.
- Thingstad, T. F., et al. (2005), Nature of P limitation in the ultraoligotrophic Eastern Mediterranean, *Science*, *309*(5737), 1068–1071.
- Trapp, J. M., F. J. Millero, and J. M. Prospero (2010), Temporal variability of the elemental composition of African dust measured in trade wind aerosols at Barbados and Miami, *Mar. Chem.*, *120*, 71–82.
- Uematsu, M., R. A. Duce, and J. M. Prospero (1985), Deposition of atmospheric mineral particles in the North Pacific Ocean, *J. Atmos. Chem.*, *3*(1), 123–138.
- Uitz, J., H. Claustre, B. Gentili, and D. Stramski (2010), Phytoplankton class-specific primary production in the world's oceans: Seasonal and interannual variability from satellite observations, *Global Biogeochem. Cycles*, *24*, GB3016, doi:10.1029/2009GB003680.
- van Vuuren, D. P., et al. (2011), The representative concentration pathways: An overview, *Clim. Change*, *109*, 5–31.
- Volpe, G., V. F. Banzon, R. H. Evans, R. Santoleri, A. J. Mariano, and R. Sciarra (2009), Satellite observations of the impact of dust in a low-nutrient, low chlorophyll region: Fertilization or artifact?, *Global Biogeochem. Cycles*, *23*, GB3007, doi:10.1029/2008GB003216.
- Wagener, T., C. Guieu, and N. Leblond (2010), Effects of dust deposition on iron cycle in the surface Mediterranean Sea: Results from a mesocosm seeding experiment, *Biogeosciences*, *7*, 3769–3781.
- Williams, R. G., and M. J. Follows (1998), The Ekman transfer of nutrients and maintenance of new production over the North Atlantic, *Deep Sea Res., Part 1*, *45*, 461–489.
- Wuttig, K., T. Wagener, M. Bressac, A. Dammshäuser, P. Streu, C. Guieu, and P. L. Croot (2013), Impacts of dust deposition on dissolved trace metal concentrations (Mn, Al and Fe) during a mesocosm experiment, *Biogeosciences*, *10*, 2583–2600.