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# Landslide inventory bimodality in volcanised tertiary basin of Puy-en-Velay (France)

## A geoindicator of climate change and geomorphologic evolution



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### 1. Background & Study area

The study of landslides aims often to understand the mechanisms and processes in order to reduce the hazard and the risk. The investigation of the past relations between climate and landslide is a common way to take account the impact of climate change on the landslide activity. By this way, landslides could also be used as a geoindicator of climate change and regional geomorphic sensitivity (Borgatti & Soldati, 2003; Canuti *et al.*, 2004; Damm & Terhorst, 2010). The tertiary basin of Puy-en-Velay is located Southeast of French Massif central in an area of 400km<sup>2</sup>. The tertiary sandy-clay material composing the basin is particularly sensitive to landslide and the relief inversion caused by volcanic shape create a complex slope system (fig.1).

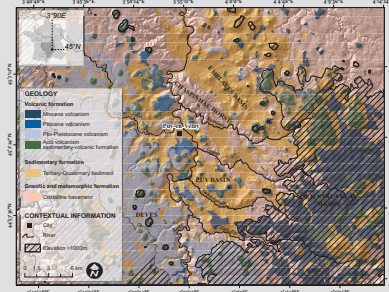


Figure 1 – Geological map of the Tertiary basin of Puy-en-Velay in the regional geologic context

The regional landslide activity is important and cause numerous injuries and damages on the infrastructures. If the present day landslides are essentially thin (<5m), they often correspond to a surficial reactivation of old deep-seated landslide. This justifies the PhD research on the landslide hazard on the tertiary basin of Puy-en-Velay. The first step of the hazard study is the historical inventory of landslide and its analysis. One interesting analysis approach is the frequency-magnitude for interpreting the structure of the population and proposing theoretical explanations.

This poster presents the frequency-magnitude analysis of the historical landslide inventory of the Puy basin and its geomorphological interpretation.

### 1. Results

More than 200 landslides were inventoried in the study area. The analysis of the inventory pointed out the existence of two clear populations i) the recent landslide (<100 years) et the ii) the old landslide (>1000 years).

#### LOCATION OF LANDSLIDES

The landslides are spatially structured in clusters, especially the old landslides (fig.3).

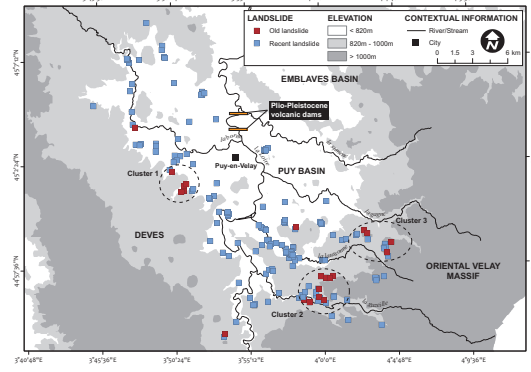


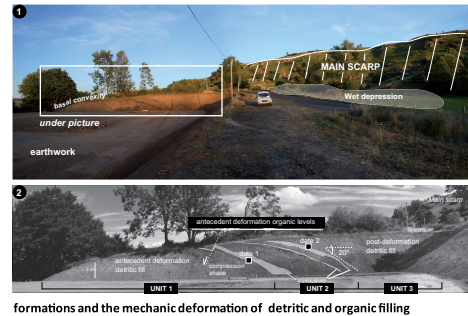
Figure 3 – Location map of inventoried landslides

The old landslides are located between 820m height and 1000m, corresponding to the elevation of Plio-Pleistocene volcanic dams located at the north of the Puy-en-Velay. These old landslides are numerous at the rims of the Oriental Velay massif (clusters 2 and 3) and are concentrated in the deep valleys cutting into the Devès plateau (cluster 1). This corresponds to the front of the Pleistocene erosion wave succeeding to the dams erosion.

#### TYPOLOGY OF LANDSLIDES AND AGE OF OLD LANDSLIDES

The recent landslide are organised in 3 different type of landslide while old landslide can be group in one type (tab.1).

Table 1 – Landslide type inventoried in the study area		
	Surficial translational slide	Planar landslide with less than 2m deep
Recent landslide	Surficial or deep rotational slide	Rotational slide with deep between 2 and 5m
	Flowage	Soil deformation with a viscoplastic compartment
Old landslide	Like Complex or flow-slide	Basaltic cornice dismantling and flowage



Four old landslides have been dated by C14, with peat formation in the reverse slope cause by a rotational component of the high part of this kind of landslide (fig.4). Three landslides were dated between 4469±43 Cal.BP and 2820±42 Cal.BP, so during the Subboreal period. The last one (fig.4) presents two phases of activation, the first one before 25,975±223 Cal.BP (date 1) and the second one after 20,457±115 Cal.BP (date 2). The second phase might have occurred during the Subboreal so.

Figure 4 – Example of a palaeolandslide. 1) general context of the landslide; 2) details of the formations and the mechanic deformation of detritic and organic filling

### 2. Methodology

The inventory of landslides was based on two classical approaches:

1. A photo interpretation of aerial photographs obtained by Google Earth®
2. A field survey to verify the photo interpretation and to prospect under forest cover

- 1. All the landslides were digitalised on a GIS software. The colour aerial photography had a 50cm resolution and the digitalisation was carried off to 1/1000. When it was possible, we allocate a landslide type for each polygon.
- 2. Each polygon was verified in the field to determine the type of landslide, the geologic material, the possible causes of rupture, etc.

In addition, more detailed investigations on some representative landslides have been made (fig.2):



Figure 2 – Example of field investigation (DGPS monitoring, peat extraction for C14 datation, geophysics, Pressler extraction on a tree, borehole)

1. Geophysics – 2 seismic tomography cross-sections and 15 boreholes,
2. Monitoring – centimetric DGPS monitoring with a monthly frequency during one year for 4 landslides
3. Dendrogeomorphology – 12 trees investigating on 2 landslides
4. Datation – C14 datation of 4 palaeolandslides

All these investigations have allowed us to specify for each landslide the type and the processes.

#### MORPHOMETRIC ANALYSIS

The landslides present a clear bimodality in their dimensions (fig.5) and their frequency-magnitude relation (fig.6)

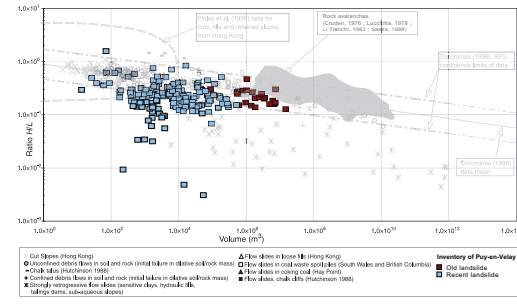
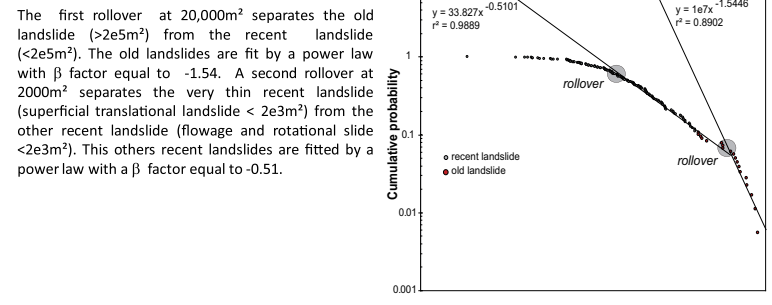


Figure 5 – Tan $\alpha$ /Volume relationship of the landslides inventory compare to other referential inventories

The old landslides form a group of high volume and low Tan $\alpha$  compare to the recent landslides which have higher values of Tan $\alpha$  but lower in volume. Those criteria define a different physical compartment and initiation processes. The recent landslides are defined by soil rupture initiation or like-flowslide mechanical compartment. The old landslides are closed to the flowslide or debris avalanche compartment.

This difference of apparent mechanical compartment, with the same geological material, can be caused by the triggering factors or the physiography. The frequency-magnitude relation gives us some elements.

Figure 6 – Frequency-magnitude distribution of landslides of Puy-en-Velay basin



The first rollover at 20,000m<sup>2</sup> separates the old landslide (>2e5m<sup>2</sup>) from the recent landslide (<2e5m<sup>2</sup>). The old landslides are fit by a power law with  $\beta$  factor equal to -1.54. A second rollover at 2000m<sup>2</sup> separates the very thin recent landslide (superficial translational landslide < 2e3m<sup>2</sup>) from the other recent landslide (flowage and rotational slide <2e3m<sup>2</sup>). These others recent landslides are fitted by a power law with a  $\beta$  factor equal to -0.51.

#### Interpretations & Conclusions

The landslide population of the tertiary Puy basin is composed of two sub-groups: i) recent landslides, majority surficial and often due to local reactivation of old landslide and ii) old landslides, deep and complex, with real effects on relief (basaltic cornice dismantling) and with a timely convergence on the Subboreal period. This period is also known by other authors in Europe (Dapples, 2002 ;Soldati *et al.*, 2004; Hradecky *et al.*, 2007; Van-den-Eeckhaut *et al.*, 2007; Panek *et al.*, 2008a; Gioia *et al.*, 2010).

Concerning the different power law  $\beta$  factor, we can interpret it as a physiographic limit as Guthrie & Evans (2004) but the time separates the two populations is not so long to obtain a real physiographic difference. We prefer to bring the findings of Van Den Eeckhaut *et al.* (2007) in the Flemish Ardennes and consider that the  $\beta$  factor corresponding to a different sensitivity of the geomorphologic system, which was submitted to triggering factors with the highest magnitude in the past (highest  $\beta$  factor) than the present day (lowest  $\beta$  factor). In our study area, we consider that the Subatlantic conditions correspond to a  $\beta$  factor of 0.51 while the Subboreal period corresponds to a  $\beta$  factor of 1.54. This magnitude difference equals  $p^{1.36}$ .

All these conclusions are in agreement with literature (Buma & Dehn, 1998; Dehn *et al.*, 2000; Soldati *et al.*, 2004) that considers the Subboreal period is marked by a largely positive water balance. This positive water balance is also in agreement with old landslide with low Tan $\alpha$  values and close to flow-slide or debris avalanche mechanical compartment.

Finally, this two landslides populations can be considered as witnesses of climatic conditions evolution and geoindicators.

### References

Borgatti, L. and M. Soldati (2003) - Landslide events as indicators of landscape sensitivity to climate change. *Geomorphological sensitivity and system response*, Camerino - Modena Apennines, Italy, 35-46  
 Buma, J. and M. Dehn (1998) - A method for predicting the impact of climate change on slope stability. *Environmental Geology*, 35, (2-3), 190-196.  
 Canuti, P., N. Casagli, L. Ermini, R. Fantini and P. Farina (2004) - Landslide activity as a geoindicator in Italy: significance and new perspectives from remote sensing. *Environmental Geology*, 45, 907-915.  
 Damm, R. and B. Terhorst (2010) - A model of slope formation related to landslide activity in the Eastern Prealps, Austria. *Geomorphology*, 122, 338-350.  
 Dapples, F. (2002) - Instabilités de terrain dans les Préalpes Fribourgeoises (Suisse) au cours du Tardiglaciaire et de l'Holocène: influence des changements climatiques, des fluctuations de la végétation et de l'activité humaine. Thèse de doctorat, Faculté des Sciences de l'Université de Fribourg (Suisse), 158 p.  
 Dehn, M., G. Bürger, J. Buma and P. Gasparetto (2000) - Impact of climate change on slope stability using expanded downscaling. *Engineering Geology*, 55, 193-204.  
 Gioia, D., P. Di Leo, S. I. Giano and M. Schiattarella (2010) - Chronological constraints on a Holocene basin in an intermontane basin of the southern Apennines, Italy: Morphological evolution and palaeoclimatic implications. *The Holocene*, 21, (2), 263-273.  
 Guthrie, R. H. and S. G. Evans (2004) - Analysis of landslide frequencies and characteristics in a natural system, coastal British Columbia. *Earth Surface Processes and Landforms*, 29, 1321-1339.  
 Hradecky, J., T. Hradecky, V. Smolova and K. Silhan (2007) - Landslide complex in the northern part of the Silesian Beskydy Mountains (Czech Republic). *Landslides*, 4, 53-62.  
 Panek, T., J. Hradecky, V. Smolova and K. Silhan (2008b) - Giant ancient landslide in the Alma water gap (Crimean Mountains, Ukraine): notes to the predisposition, structure, and chronology. *Landslides*, 5, 367-378.  
 Soldati, M., A. Corsini and A. Pasuto (2004) - Landslides and climate change in the Italian Dolomites since the Late glacial. *Cotena*, 55, 141-161.  
 Van-Den-Eeckhaut, M., J. Poesen, G. Govers, G. Verstraeten and A. Demoulin (2007b) - Characteristics of the size distribution of recent and historical landslides in a populated hilly region. *Earth and Planetary Science Letters*, 256, 588-603.