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Stable atmospheric methane in the 2000s: key-role of emissions from natural wetlands

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Abstract

Two atmospheric inversions (one fine-resolved and one process-discriminating) and a process-based model for land surface exchanges are brought together to analyze the variations of methane emissions from 1990 to 2009. A focus is put on the role of natural wetlands and on the years 2000–2006, a period of stable atmospheric concentrations.

From 1990 to 2000, the two inversions agree on the time-phasing of global emission anomalies. The process-discriminating inversion further indicates that wetlands dominate the time-variability of methane emissions with 90% of the total variability. Top-down and bottom-up methods are qualitatively in good agreement regarding the global emission anomalies. The contribution of tropical wetlands on these anomalies is found to be large, especially during the post-Pinatubo years (global negative anomalies with minima between -41 and -19 Tgy^{-1} in 1992) and during the alternate 1997–1998 el-Niño / 1998–1999 la-Niña (maximal anomalies in tropical regions between $+16$ and $+22 \text{ Tgy}^{-1}$ for the inversions and anomalies due to tropical wetlands between $+12$ and $+17 \text{ Tgy}^{-1}$ for the process-based model).

Between 2000 and 2006, during the stagnation of methane concentrations in the atmosphere, total methane emissions found by the two inversions on the one hand and wetland emissions found by the process-discriminating-inversion and the process model on the other hand are not fully consistent. A regional analysis shows that differences in the trend of tropical South American wetland emissions in the Amazon region are mostly responsible for these discrepancies. A negative trend ($-3.9 \pm 1.3 \text{ Tgy}^{-1}$) is inferred by the process-discriminating inversion whereas a positive trend ($+1.3 \pm 0.3 \text{ Tgy}^{-1}$) is found by the process model. Since a positive trend is consistent with satellite-derived extent of inundated areas, this inconsistency points at the difficulty for atmospheric inversions using surface observations to properly constrain tropical regions with few available observations. A consequence is the need to revisit the large increase in anthropogenic emissions computed at the global scale by

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some inventories for the early 2000s, although process-based models have also their own caveats and may not take into account all processes.

1 Introduction

The growth rate of atmospheric methane (CH_4) has experienced large variations since the early 1990s: after a decade of decrease, interrupted by two peaks in 1997–1998 and 2002–2003 (Dlugokencky et al., 1998; Cunnold et al., 2002; Wang et al., 2004; Bousquet et al., 2006), the growth rate of atmospheric methane remained small from 1999 to 2006, only to increase again since 2007 (Rigby et al., 2008; Dlugokencky et al., 2009; Bousquet et al., 2011). As methane is emitted by a large variety of sources, the explanations for the observed atmospheric variations have generally implied changes in several source or sink types.

The sources involved in the variations of the 1990s are now well understood, although their relative magnitude may still be debated.

The 1991–1993 growth rate anomaly is linked to the Pinatubo eruption, which led to a decrease in methane loss, due to reduced tropospheric hydroxyl radical (OH) concentrations and stratospheric chemistry changes (Bândă et al., 2012), followed by a decrease in natural wetland emissions. The negative impact of this volcanic event on wetland CH_4 emissions is due to its effects on the climate, with both cooling (Dlugokencky et al., 1996) and precipitation anomalies in the Northern hemisphere (Walter et al., 2001b; Dlugokencky et al., 2011), likely slightly amplified by sulfur deposition (Gauci et al., 2008). The collapse of the economy of the former USSR and Eastern Europe has also led to decreased anthropogenic emissions starting in 1991 and spanning most of the 1990s (Dlugokencky et al., 2003).

The 1997–1999 large growth rate anomaly is explained by a positive anomaly in emissions from biomass burning, both in the Tropics and at high latitudes (van der Werf et al., 2006; Langenfelds et al., 2002), together with a negative anomaly in emissions from tropical wetlands in 1997–1998, linked to dry el-Niño conditions, followed

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by a positive anomaly in the same in 98-99, linked to wet la-Niña conditions (Bousquet et al., 2006; Chen and Prinn, 2006).

Overall, the role of wetlands in changes in atmospheric methane has often been identified as dominant (Chen and Prinn, 2006; Bousquet et al., 2006, 2011), linked to meteorological conditions of temperatures and precipitations (Dlugokencky et al., 2009; Bousquet et al., 2011). These conditions are thought to lead to changes in both the wetland extent (e.g. Ringeval et al., 2010) and the CH₄ flux per wetland area (e.g. Bloom et al., 2010). The impact of wetland emissions is often thought to be combined with changes in OH concentrations (Wang et al., 2004; Monteil et al., 2011), although it is now well accepted that inter-annual changes in OH concentrations are limited to 1-3% and can therefore explain only a limited part of the observed changes in the methane growth rate (Montzka et al., 2011).

The analysis of the variations of atmospheric methane after 1999 is still largely debated.

Various scenarios have been suggested to explain the stabilization of methane concentrations between 2000 and 2006: (i) reduced global fossil-fuel-related emissions, estimated from AGAGE and NOAA (Chen and Prinn, 2006) or from ethane emissions used as a proxy to fossil-fuel-related CH₄ emissions (Aydin et al., 2011; Simpson et al., 2012); (ii) compensation between increasing anthropogenic emissions (as inferred by inventories such as EDGAR4 (EDGAR 4, 2009)) and decreasing wetland emissions (Bousquet et al., 2006); (iii) a decrease in emissions due to rice-paddies, attributed to changes in agricultural practices (Kai et al., 2011); (iv) stable microbial and fossil-fuel-related emissions in the early 2000s (Levin et al., 2012) and/or (v) significant (Rigby et al., 2008) to small (Montzka et al., 2011) changes in OH concentrations.

The increase of methane concentrations observed since 2007 (Rigby et al., 2008; Dlugokencky et al., 2009; Susmann et al., 2012) is also still not understood with hints to possible increasing anthropogenic emissions (EDGAR 4, 2009) and increasing wetland emissions in relation to abnormal precipitations over tropical lands in the late 2000s (Bousquet et al., 2011; Dlugokencky et al., 2009).

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Methane emissions can be investigated using process-based models and inventories (bottom-up or B-U approach), and atmospheric inversions (top-down or T-D approach). Atmospheric inversions combine atmospheric observations of CH₄, an atmospheric chemistry and transport model, and prior information about sources and sinks. The flux estimates that give the best fit to the observed atmospheric concentrations are derived by optimization. However, atmospheric inversions provide a limited insight into the underlying biogeochemical processes controlling emissions, particularly over regions where several processes and sources overlap. B-U models computing wetland (Melton et al., 2012) or biomass-burning (van der Werf et al., 2010) methane emissions incorporate knowledge of small-scale processes but they need additional information and constraints to project their local emission estimates to larger scales compatible with the atmospheric signals. B-U emission inventories (EDGAR 4, 2009; EPA, 2011) are based on country-scale energy use and agricultural statistics and usually provide yearly to decadal estimates of anthropogenic emissions at global and national scales.

In this study, we present an analysis of methane fluxes for the 1990–2009 period. We specially investigate the relative contributions of the various emissions leading to the stable concentrations during the 2000–2006 period. For this, we bring together the T-D and B-U methodologies to analyze the variations of methane emissions and assess the particular role of emissions from natural wetlands. We study (i) the total methane emissions, as inferred by two different atmospheric inversions (T-D); (ii) the wetland emissions, as inferred by one inversion (the one which is able to discriminate the various sources or sinks) and by one process-based model of vegetation (B-U). The methodology of the process-based model and the two inversions is described in Sect. 2; results are given in Sect. 3 at the global (Sect. 3.1) and regional (Sect. 3.2) scales.

2 Methodology

We combine the results of a process-based model for natural wetland emissions, ORCHIDEE (Krinner et al., 2005; Ringeval et al., 2010, 2011), together with the CH₄ fluxes estimated by two different atmospheric inversions, variational (Chevallier et al., 2005; Pison et al., 2009) and analytical (Bousquet et al., 2005).

2.1 Process model for natural wetland emissions

The model of CH₄ emissions by natural wetlands used here (hereafter ORCHIDEE) is based on the ORCHIDEE global vegetation model (Krinner et al., 2005) which simulates land energy budget, hydrology, and carbon cycling. It has then been developed (Ringeval et al., 2010, 2011) to simulate:

- the CH₄-emitting wetland area dynamic, using some TOPMODEL concepts (Ringeval et al., 2012a). The wetland extents are here normalized to match Papa et al. (2010). A global, multi-year data set giving the monthly distribution of flooded areas at a ≈ 25 km resolution has been generated. It is based on multiple satellite observations which are optimized specifically for surface water detection (Papa et al., 2010; Prigent et al., 2012). It is built using a combination of satellite data including passive microwave observations and a linear mixture model to account for vegetation. Three wetland classes, differing by the value of the water table depth, are simulated for each grid cell and at each time step: saturated wetland and wetland with a mean water table at 3 and 9 cm below the soil surface.
- the flux density, which is computed following Walter et al. (2001a) and Ringeval et al. (2010) and results from three processes: production, oxidation and transport (via diffusion, ebullition and through plants).

A summary of the ORCHIDEE methodology used to compute CH₄ emissions is given in Wania et al. (2012).

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The wetland emissions are driven by the CRU-NCEP climate forcing data set (Viovy and Ciais, 2009). In this study, four scenarios of wetland emissions are given based on:

- accounting or not for the non-saturated wetlands (i.e. all classes of wetlands or saturated wetlands only)
- multiplying or not the simulated wetland area by peat-land map in boreal regions. This map is obtained by using soil organic carbon data from IGBP DIS at high resolution ($5' \times 5'$), by dividing each pixel of this database by 130 kg m^{-3} (which is the maximum soil carbon density of peat), and then by regridding the result at $1^\circ \times 1^\circ$ resolution (Lawrence and Slater, 2007). The hypothesis underlying the multiplication of the two products (map of peat-land cover and map of inundated areas) is that the inundated fraction is the same for an entire grid-cell as for a sub-grid peat-land into this grid-cell (see Bousquet et al., 2011, for more details).

ORCHIDEE provides CH_4 emissions by natural wetlands as the product of a flux density by a wetland extent, at a 1 degree by 1 degree resolution for the time period 1990–2009.

Note that ORCHIDEE emissions are not used as prior emissions in INVANA nor INVVAR (Sect. 2.2).

2.2 Inverse methods

The two inversion models used here are based on the Bayesian formalism: they assimilate surface observations of CH_4 and methyl-chloroform (CH_3CCl_3 or MCF) concentrations (measurements from various surface monitoring networks: AGAGE (Prinn et al., 2000, accessed: 2012-03-31), CSIRO (Francey et al., 1999), EC (Worthy et al., 1998), ENEA (Artuso et al., 2007), AEMET (Gomez-Pelaez et al., 2010), JMA (Matsueda et al., 2004), LSCE (Schmidt et al., 2005), NIWA (Lowe et al., 1991), NOAA/ESRL (NOAA, access: 2012-03-31), SAWS (Brunke et al., 2001), UBA (UBA, 2013), as available on the World Data Center for Greenhouse Gases) into the LMDZ global circulation

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model (Hourdin et al., 2006) with prior information on the spatio-temporal distribution and uncertainties of CH₄ sources and sinks to estimate the magnitude and the uncertainties of optimized surface emissions. The two methods differ by the inversion setup and the resolution method: see Bousquet et al. (2005) and Pison et al. (2009).

The analytical scheme (hereafter INVANA) is described in more details in Bousquet et al. (2005, 2011). Briefly, it solves for monthly emissions for various types of sources and sinks in ten land regions plus one global ocean. The constraints are monthly mean CH₄ concentrations at 68 surface stations (CSIRO (Francey et al., 1999), LSCE (Schmidt et al., 2005), NOAA/ESRL (Dlugokencky et al., 1994, 2009)) with monthly uncertainties ranging from ±5 ppb to ±50 ppb, with a median of ±10 ppb. The transport model is LMDZt v3 offline nudged on analyzed horizontal winds (Uppala et al., 2005; Hourdin et al., 2006) with OH pre-optimized using methyl-chloroform concentrations (Bousquet et al. (2005) used MCF concentration measurements by AGAGE (Prinn et al., 2005) and NOAA/ESRL (Montzka et al., 2000, 2011)). The prior emissions are elaborated from various inventories by Matthews and Fung (1987), Olivier and Berdowski (2001) and van der Werf et al. (2006). Monthly uncertainties on fluxes are set at 150 % in each region; there are no error correlations but month-to-month changes are limited to ±250 % if the process follows a seasonal cycle or to ±50 % otherwise (Peylin et al., 2000, 2002).

The main advantage of this inversion scheme is the low cost of computation, which makes it possible to test various scenarios, varying OH concentrations, constraints, uncertainties. In this study, eleven scenarios are used, as described in Bousquet et al. (2011).

The variational scheme (hereafter INVVAR) is described in Chevallier et al. (2005) and Pison et al. (2009). Briefly, it solves for weekly net emissions at grid-point resolution (2.5° × 3.75°). The constraints are daily mean observations at continuous measurement stations by AGAGE for CH₄ (Rigby et al., 2008) and MCF (Prinn et al., 2005) and flask and 24-h averaged data for CH₄ and MCF by CSIRO (Francey et al., 1999), EC (Worthy et al., 1998), ENEA (Artuso et al., 2007), AEMET (Gomez-Pelaez et al.,

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2010), JMA (Matsueda et al., 2004), LSCE (Schmidt et al., 2005), NIWA (Lowe et al., 1991), NOAA/ESRL (Montzka et al., 2000, 2011; Dlugokencky et al., 1994, 2009), SAWS (Brunke et al., 2001), UBA (UBA, 2013). All methane data were re-scaled to the NOAA 2004 methane standard scale (Dlugokencky et al., 2005). The differences in MCF scales have been included in the errors associated to MCF observation error statistics i.e. in the variances attributed to the differences between measured and simulated MCF concentrations. The chemistry-transport model is LMDZt v4 – SACS (Pison et al., 2009) with prior OH fields provided by a full-chemistry run with the INCA scheme (Hauglustaine et al., 2004; Folberth et al., 2005). OH and CH₄ are then optimized simultaneously. Prior emissions are elaborated from inventories by Fung et al. (1991), Olivier and Berdowski (2001) and van der Werf et al. (2006) plus Montzka et al. (2000) re-scaled for MCF. The prior variances of CH₄ fluxes are set at ±100 % of the maximum flux over the grid cell and its eight neighbors; error correlations are modeled using correlation lengths (500 km on land, 1000 km on oceans) without time correlations (Chevallier et al., 2005). The main advantage of this inversion scheme is that there is no aggregation errors but it has a high computing cost.

We mainly focus our analysis on the inter-annual variations of anomalies of CH₄ fluxes: monthly anomalies are deseasonalized by computing 12-month running means and subtracting the deseasonalized mean for the reference period 1993–2007; yearly anomalies are computed by averaging fluxes over each year and subtracting the 1993–2007 average. 1993–2007 is chosen as the reference because satellite data used for additional tests (see Fig. 3) are available during this period so that it is the longest time-period covered by all our estimates. The anomalies of the net total fluxes (to the atmosphere) are computed from INVVAR and from INVANA (the fluxes per process categories are summed up into a net budget); the CH₄ wetland emission anomalies are computed from INVANA (using only the natural wetland category) and from ORCHIDEE. In the following, unless otherwise specified, the ± ranges given after INVANA (respectively ORCHIDEE) figures are actually the standard deviation over the 11 (re-

spectively 4) scenarios; since INVVAR consists in only one case, no such range is given.

3 Results

3.1 Global emissions

5 The total net emissions are distributed similarly in INVANA and INVVAR, with $58 \pm 2\%$ and 56% respectively South of 30° N, $26 \pm 2\%$ and 27% at the mid-latitudes ($30\text{--}50^\circ$ N) and $16 \pm 1\%$ and 17% at the high latitudes ($50\text{--}90^\circ$ N). The wetland emissions are also similarly distributed in INVANA and ORCHIDEE, with $67 \pm 4\%$ and $75 \pm 6\%$ of the wetland emissions South of 30° N in INVANA and ORCHIDEE respectively, only
10 $11 \pm 2\%$ and $14 \pm 2\%$ at the mid-latitudes and $21 \pm 3\%$ and $11 \pm 7\%$ at the high latitudes. Variations in the latitudinal distribution of wetlands in ORCHIDEE are significant ($\pm 7\%$) and due to the use or not of a soil carbon map overlaid to the wetland extent given by ORCHIDEE to reduce the simulated CH_4 emitting areas to peat-lands (see Sect. 2.1).

At the global scale, the inter-annual variations of the total net CH_4 fluxes of INVVAR and the various INVANAs are in good agreement, in the phasing through time and in the magnitude with time correlations ranging from 66% to 84% over the 19 common years (Fig. 1, top). Since these correlations are for deseasonalized anomalies, which are a relatively small signal, they are considered as good. This result is not simply a check of a correct mass balance as the two inversions use different OH fields and different
15 approaches to optimize the atmospheric sink of CH_4 (off-line in INVANA or simultaneously in INVVAR). INVVAR retrieves a lower variability than INVANA with standard deviations over all monthly anomalies of 12 and 18 Tgy^{-1} respectively. The standard deviation for wetlands in INVANA is 16 Tgy^{-1} , which indicates that wetlands explain about 90% of the variability of total methane emissions. In the following sections we
20 therefore focus on changes in natural wetland emissions.

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INVAR and INVANA show negative anomalies after the 1991 Pinatubo eruption (minima of -19 and -20 to -41 Tgy^{-1} in 1992 for respectively INVAR and INVANA, Fig. 1 top), in agreement with the literature. These anomalies are mainly due to the contribution of the tropical areas (minima for latitudes less than 30° N of -21 Tgy^{-1} for INVAR and -19 to -35 Tgy^{-1} for the 11 INVANA scenarios in 1992, Fig. 1). This is consistent with the negative impact of this volcanic event on wetland CH_4 emissions already noticed in the literature (Hogan and Harriss, 1994). The one INVANA scenario producing smaller changes in global and tropical emissions is the one assuming constant OH concentrations with time; it is closer to INVAR in the early 1990s. In INVANA, the optimized OH concentrations are sensitive to the errors on MCF emissions, which are proportional to MCF fluxes. Since MCF emissions are large during the 1980-90s, their errors are also large and INVANA find large year-to-year OH changes (up to -15% for the total column between 1996 and 1997 for one scenario). In INVAR, since the optimization is simultaneous, OH concentrations are constrained by both MCF and CH_4 measurements. Therefore, OH variability is smaller than in most INVANA scenarios, closer to the constant-OH one, and more consistent with Montzka et al. (2011).

A large positive anomaly of global total emissions is inferred by INVAR and INVANA in 1997–1999 (maxima at $+37$ and $+19$ to $+25$ Tgy^{-1} respectively, Fig. 1 top), related to the el-Niño/la-Niña events. Most of the signal comes from tropical regions (maxima at $+27$ and $+16$ to $+22$ Tgy^{-1} in INVAR and the INVANAs). That could be linked to fires in 1997–1998 (van der Werf et al., 2004) and exceptionally high CH_4 emissions from natural wetlands during la-Niña event in 1998 (Hodson et al., 2011). These high wetland emissions are actually consistent with 1998 natural wetland emissions as seen by ORCHIDEE, with anomalies of $+11$ to $+17$ Tgy^{-1} at the global scale (Fig. 1) and $+12$ to $+17$ Tgy^{-1} in the tropical regions (Fig. 1 bottom). Note that for the four ORCHIDEE scenarios, only two curves appear on Fig. 1 bottom since the differences obtained by varying boreal peat-lands do not impact the tropical emissions.

More challenging to understand is the mismatch between INVANA and ORCHIDEE wetland emissions between 2000 and 2006 (Fig. 1, 3rd graph from top). The median

time correlation between all ORCHIDEE and all INVANA scenarios is 45 % over 1990–2009 but reaches 68 % if the period 2000–2006 is excluded from the computation. Moreover, if the trends of both ORCHIDEE and INVANA are positive between 1992 and 1999 ($+4.7 \pm 1.0 \text{ Tgy}^{-1}$ and $+4.7 \pm 0.7 \text{ Tgy}^{-1}$ respectively), ORCHIDEE infers a stable to positive trend between 2000 and 2006 ($+2.0 \pm 0.4 \text{ Tgy}^{-1}$) whereas INVANA infers a negative trend over the same period ($-4.2 \pm 1.2 \text{ Tgy}^{-1}$). To assess the cause of this discrepancy between the two approaches, we studied the regional distributions of methane emissions within the Tropics.

3.2 Latitudinal and regional break-down of emissions

We find that the disagreement between ORCHIDEE and INVANA at the global scale after 2000 is mainly due to the anomalies south of 30 degrees North, an area where emissions are mainly due to tropical land regions (Fig. 1). Of the four ORCHIDEE scenarios described previously, only two are relevant in the tropics (the other two are obtained by varying boreal peat-lands) and are thus retained in the following: ORCHIDEE-all accounting for all classes of wetlands and ORCHIDEE-sat accounting only for saturated wetlands. In this latitudinal band, the signs of the anomalies from 1991 to 2000 for total emissions by the two inversions on the one hand and for wetland emissions by INVANA and ORCHIDEE on the other hand are broadly in agreement. From 2000, and particularly between 2003 and 2006, ORCHIDEE provides anomalies with a sign opposite to the one of INVANA. For example, in 2005 the anomaly at these latitudes is $+3$ to $+13 \text{ Tgy}^{-1}$ in ORCHIDEE versus -6 to -26 Tgy^{-1} in INVANA (Fig. 1 bottom).

Moreover, the trend over 2000–2006 in this band (Table 1) is $-4.1 \pm 0.9 \text{ Tg month}^{-1}$ for INVANA whereas the trend for the mean between ORCHIDEE-all (scenario accounting for all wetlands) and ORCHIDEE-sat (scenario with saturated wetlands only) is $+1.4 \pm 0.3 \text{ Tg month}^{-1}$ (Table 1).

In order to investigate the regional contributions to the flux anomalies and trends of wetland emissions in the Tropics, we computed the yearly flux anomalies for ORCHIDEE and the mean of the eleven scenarios of INVANA for eight key-regions South

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of 30° N (Fig. 2). Interestingly, ORCHIDEE and INVANA agree on the fact that variations in methane emissions from the Amazon basin (tropical South America in Fig. 2) drive the trend of tropical emissions, but with opposite signs. Indeed, the ratio between the slopes for tropical South America and the whole < 30° S region (Table 1) indicates that tropical South America explains 93 % of the trend for ORCHIDEE and between 89 and 95 % in INVANA (depending on the scenarios). However, the inferred trends for the Amazon region are of opposite signs with $-3.9 \pm 1.3 \text{ Tgy}^{-1}$ for INVANA and $+1.3 \pm 0.3 \text{ Tgy}^{-1}$ for ORCHIDEE (Table 1).

When an atmospheric inversion cannot provide insights about the underlying emission processes and the sensitivity of wetland emissions to the climate, the ORCHIDEE process-based model can. The monthly anomalies of CH₄ wetland emissions for the Amazon region for ORCHIDEE and INVANA, together with the precipitations and temperature changes are shown in Fig. 3. Not surprisingly, ORCHIDEE and INVANA phases of the inter-annual variations are not in agreement at this regional scale, although the magnitudes of the changes are close. As noted before, the trends between 2000 and 2006 are opposite for ORCHIDEE (positive) and INVANA (negative) (Fig. 3, top panel). Wetland extent is a key element in process-model algorithms (Melton et al., 2012). Therefore, in addition to ORCHIDEE-sat where some TOPMODEL concepts are used to compute the variability in time of the wetland extent (Ringeval et al., 2012a), we introduced an ORCHIDEE estimate where this variability is prescribed by the data set by Prigent et al. (2012). Whatever the choice of inter-annual variability (IAV) for wetland extent (i.e. computed or prescribed from remote sensing data), the IAV and the trend in the various ORCHIDEE scenarios are never close to INVANA's (Fig. 3, upper panel). Interestingly, changing the description of the wetland extent significantly impacts the phasing and the magnitude of the inter-annual variations inferred by ORCHIDEE (Fig. 3, e.g. the 1998–2003 period), but both descriptions lead to a consistent positive trend between 2000 and 2006 ($+1.7 \text{ Tgy}^{-1}$ for ORCHIDEE-all, $+1.1 \text{ Tgy}^{-1}$ for ORCHIDEE-sat and $+1.5 \text{ Tgy}^{-1}$ for “ORCHIDEE-sat;Prigent”), although the trend is

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close to zero for “ORCHIDEE-sat;Prigent” for the period 2002–2005 but remains positive for other ORCHIDEE simulations.

In tropical South America, the trend in ORCHIDEE wetland emissions is driven by precipitations (Fig. 3, middle panel), which show a positive trend of 0.05 mm d^{-1} between 2000 and 2006 over the Amazon. When the IAV of wetland extent is removed (black curve in Fig. 3, bottom panel), the IAV of emissions over 1990–2009 decreases from 3.8 (ORCHIDEE-sat) to 2.1 Tgy^{-1} and appears to be driven by air temperature with a strong correlation of 0.86 over 1990–2009 (Fig. 3, bottom panel). Temperature could have an effect on both methanogenesis rate and substrate supply but large uncertainties remain on the contribution of each process (e.g. White et al., 2008, for northern peat-lands).

Overall, ORCHIDEE infers an increase of CH_4 emissions in the Amazon between 2000 and 2006, which is opposite to the decreasing emissions inferred by INVANA inversions. The difference of IAV between ORCHIDEE-simulated emissions when the time variability of wetland extent is either prescribed or computed (Fig. 3, upper panel) underlines the difficulty to capture a good IAV of the processes involved in the extension/retraction of wetland areas. A way of improvement would be to implement the floodplains into the model (e.g. like Decharme et al., 2008, 2012), the relevant processes leading to wetland formation in the Amazon basin (Hess et al., 2003). Uncertainties linked to precipitations (magnitude and spatial distribution) in the Amazon basin and their huge effect on the floodplains extent simulated by land surface models (Guimberteau et al., 2012) is also problematic.

The lack of wetland Plant Functional Type into the ORCHIDEE model and the use of the mean grid-cell simulated labile carbon as methanogenesis substrate’s proxy could also lead to an overestimated sensitivity of CH_4 emissions to precipitations (Ringeval et al., 2012b). However, in our study, the sensitivity of CH_4 emissions to precipitations seems to mainly happen through the wetland extent (Fig. 3) and make the lack of floodplain representation the main caveat of the B-U approach used here.

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The IAV of the remote-sensed inundated extent in the Amazon basin over 2000–2006 and of the in-situ river discharge are in agreement and do not exhibit a clear negative trend (see Fig. 10a in Papa et al., 2010). This points out at a possible issue with INVANA global inversions. Indeed, the Amazon region is poorly constrained by the surface networks, which only provide routine observations over the neighboring oceans. This lack of nearby observations means that large changes in emissions can be tolerated in the inversion when assimilating these observations. Also, using large regions may induce aggregation errors as explained in Kaminski et al. (2001). It should be noted that for this area INVVAR (which works at the grid cell's scale but only optimizes total net emissions) obtains emissions which are close to the prior; the only differences appear in the seasonal time profile, with higher emission peaks, particularly in August–September 2005 (a dry year (Frappart et al., 2012) with exceptional fires and September 2007–2008). This seems to confirm that the changes made by INVANA in the large region are due to constraints which are not directly relevant to the Amazon basin.

Overcoming the lack of observations requires to assimilate more regional data in the Amazon area (such as aircraft data (Miller et al., 2007), satellite data (for example SCIAMACHY (Frankenberg et al., 2008), IASI (Crevoisier et al., 2009), GOSAT (Morino et al., 2011) (C. Cressot, personal communication, 2012))). It may then be possible to solve the fluxes at the model's resolution (in INVANA's set-up) or to solve for the different source types in the variational approach (with INVVAR method).

In INVANA, if not constrained by nearby measurements, the changes in emissions from the Amazon region are indirectly constrained by the atmospheric growth rate. Global observed concentrations are stable between 2000 and 2006 (Dlugokencky et al., 2009). Bottom-up inventories for anthropogenic emissions give positive trends at the global scale over 2000–2006. For instance, the EDGAR inventory computes a $+8 \text{ Tgy}^{-1}$ increase in anthropogenic emissions during this period (EDGAR 4, 2009). Consistent with this bottom-up inventory albeit smaller, INVANA infers a $+4 \pm 2 \text{ Tgy}^{-1}$ increase in anthropogenic emissions for the same period (Bousquet et al., 2011) and, to

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match the stable global concentration, infers a $-4 \pm 1 \text{ Tgy}^{-1}$ decrease, mostly in South American wetland emissions. This result might be sensitive to the large scale transport between the Northern latitudes and the tropics as stated in Stephens et al. (2007) for carbon dioxide. If the increase in anthropogenic CH_4 emissions is real, a decreasing source or an increasing sink must be identified to match the stable concentrations. There is no indication of a positive trend in OH concentrations after 2000 that could have induced an abnormally increasing methane loss in the troposphere (Rigby et al., 2008; Montzka et al., 2011). The increasing wetland emissions found by ORCHIDEE indicate that South America might only be an opportunistic candidate for reduced emissions and that another region and/or another process is indeed decreasing. The analyses of ^{13}C in CH_4 recently brought elements to the debate. The possible decrease in microbial-related emissions in the Northern hemisphere since the 1990s inferred by Kai et al. (2011) is contested by the analysis of another more consistent ^{13}C data set, which infers no trend in the inter-hemispheric difference of ^{13}C (Levin et al., 2012). This last result questions the validity of the large increase in anthropogenic methane emissions computed by the EDGAR inventory for the early 2000s. Analysis of ethane emissions (Aydin et al., 2011; Simpson et al., 2012) or another inventory (IIASA, Lamarque et al., 2010) also point in the direction of non-increasing fossil-fuel-related emissions in the past decades. If anthropogenic emissions are not increasing between 2000 and 2006, one could prescribe stable anthropogenic emissions in INVANA. Then, no negatively trending source would be necessary to match the stable CH_4 concentrations in INVANA, which could reconcile INVANA and ORCHIDEE.

4 Conclusions

In this study, we bring together two “top-down” and one “bottom-up” approaches. The emissions obtained by the various methods have been compared. They agree on the time-phasing of the anomalies of global methane fluxes from 1990 to 2000 and show that wetlands dominate the time-variability of methane emissions. As in previous stud-

ies, particular events can be identified: the post-Pinatubo years illustrate the impact of tropical wetlands through negative anomalies; the alternate 1997–1998 el-Niño / 1998–1999 la-Niña illustrate the qualitative agreement of all methods with a large impact of wetlands.

5 A period of poor agreement between T-D and B-U approaches is found between 2000 and 2006, during the stagnation of methane atmospheric concentrations. A regional analysis indicates that this discrepancy is mostly due to tropical South American wetland emissions. The negative trend in wetland emissions inferred by one set of inversions is not supported by the ORCHIDEE process-based model, which computes
10 a positive trend in wetland emissions due to increased wetland extent, consistent with the inundated areas retrieved by satellite. This result questions the ability of inversions based on surface observations to properly infer emissions in regions where atmospheric constraints are sparse, such as South America. It also points to the need to revisit the large increase in anthropogenic emissions computed by some inventories
15 for the early 2000s, which somehow implies to infer a decreasing source to compensate and match the stable observed CH₄ concentrations. B-U models have also their own caveats in tropical regions, such as the non representation of floodplain processes. Regarding the T-D approach, the variational method allows us to assimilate instantaneous data of any kind. After 2003, satellite data, such as SCIAMACHY, IASI (after
20 2006) or GOSAT (after 2009), can then be used to constrain methane fluxes in remote areas where surface stations are sparse, particularly at low latitudes (Bergamaschi et al., 2009). Future active missions may also add data at higher latitudes with possibly less biases than former passive retrievals (Heimann and Marshall, 2011).

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30 des Sciences du Climat et de l'Environnement), NIWA (National Institute of Water and Atmo-

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spheric Research), NOAA/ESRL (National Oceanic & Atmospheric Administration/Earth System Research Laboratory), SAWS (South African Weather Service), UBA (Federal Environment Agency Austria). The authors contacted all data PIs and in particular thank A. J. Gomez-Pelaez (AEMET); R. G. Prinn, R. F. Weiss, H. J. R. Wang (AGAGE); P. Krummel, R. Langenfelds, P. Steele (CSIRO); D. Worthy (EC); S. Piacentino (ENEA); T. Kawasato (JMA); M. Ramonet, M. Schmidt (LSCE); G. Brailsford (NIWA); E. Dlugokencky, S. A. Montzka, G. Dutton (NOAA); E. Brunke (SAWS); K. Uhse (UBA).

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Table 1. Trends for ORCHIDEE and INVANA. Slope in Tg/year; std = standard deviation; ratio S = ratio of the trends over 2000–2006 for tropical South America to the whole less than 30° N (in %).

Scenarios	Mean of ORCHIDEE-all and ORCHIDEE-sat			Mean of all 11 INVANAs		
Region	slope	R^2	ratio S	slope	R^2	ratio S
global	1.4 (std = 0.3)	0.1 (std = 0.0)	/	-4.4 (std = 1.1)	0.8 (std = 0.1)	/
< 30° N	1.4 (std = 0.3)	0.2 (std = 0.0)	/	-4.1 (std = 0.9)	0.7 (std = 0.1)	/
Amazon region only	1.3 (std = 0.3)	0.2 (std = 0.0)	93 (std = 0.9)	-3.9 (std = 1.3)	0.8 (std = 0.0)	95 (std = 22.3) excl. scenario based on Kaplan's prior: 89 (std = 11.3)

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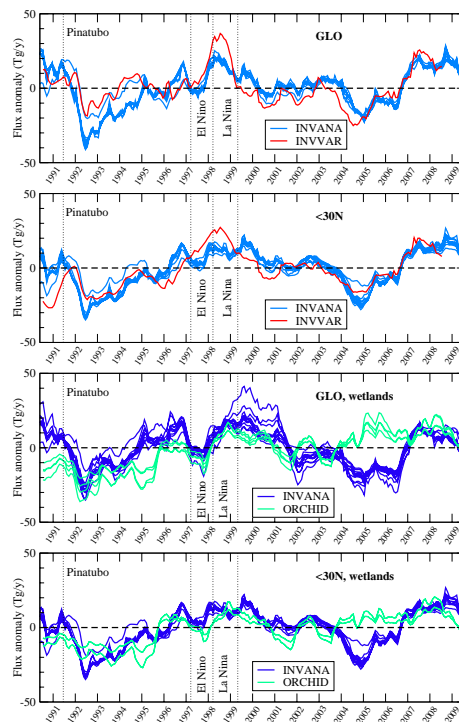


Fig. 1. Deseasonalized monthly anomalies for methane emissions. Total net emissions are given by INVVAR and INVANA. Natural wetland emissions are given by INVANA and ORCHIDEE. Anomalies are shown for the whole globe (“GLO”) and for the latitudes less than 30° N only (“< 30° N”). As detailed in Sect. 2, there are 11 scenarios for INVANA (in light and dark blue for total and wetland emissions) and 4 scenarios for ORCHIDEE (in green).

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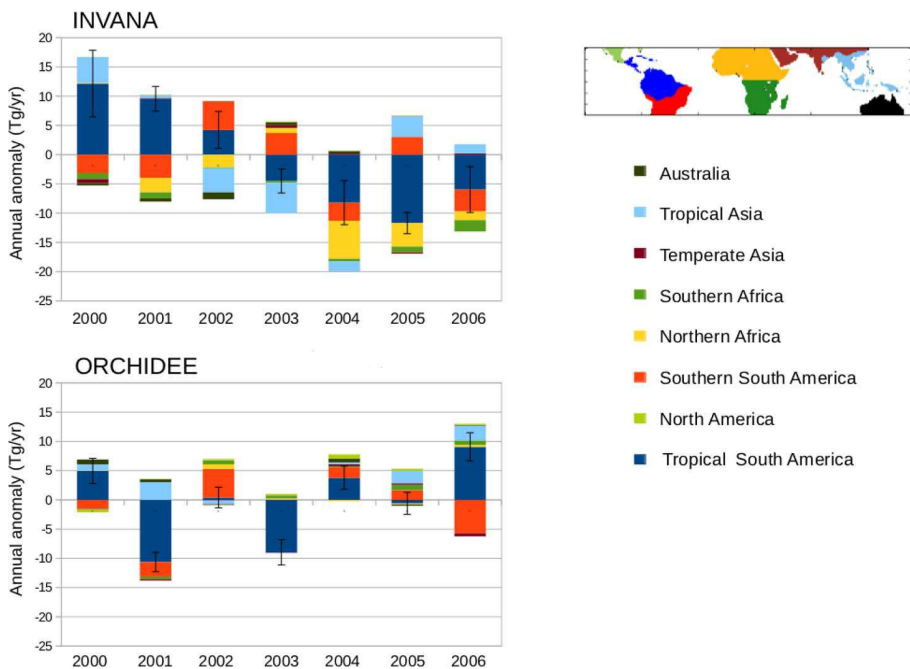


Fig. 2. Yearly anomalies of methane wetland emissions from 2000 to 2006 in eight areas in the Tropics. The distributions are shown for ORCHIDEE and the mean of the 11 INVANA scenarios. The “error bars” on the tropical South America region are computed from (i) the difference between the 11 scenarios for INVANA and (ii) the difference between the scenario with all wetlands and the one with saturated wetlands only for ORCHIDEE.



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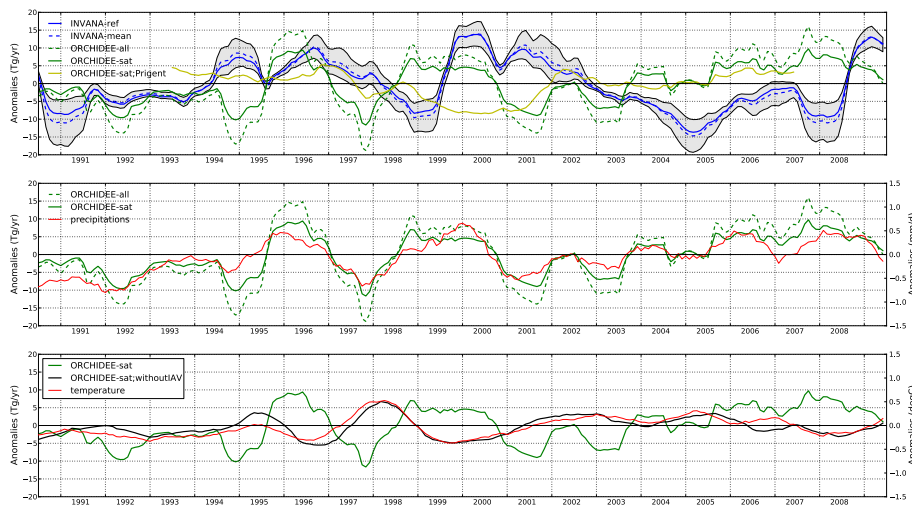


Fig. 3. Monthly anomalies of CH_4 wetland emissions, precipitations and temperature in tropical South America (see map in Fig. 2). ORCHIDEE-all = ORCHIDEE scenario accounting for all classes of wetlands; ORCHIDEE-sat = ORCHIDEE scenario with only saturated wetland; “ORCHIDEE-sat;Prigent” = ORCHIDEE scenario where the inter-annual variability of the ORCHIDEE-computed wetland extent is replaced by the one by Papa et al. (2010); “ORCHIDEE-sat;without IAV” = ORCHIDEE scenario where the inter-annual variability of the wetland extent is removed.

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