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**Fault kinematics and active tectonics at the southeastern boundary of the eastern Alborz
(Abr and Khij fault zones): geodynamic implications for NNE Iran**

Bitá Javidfakhr^{1,*}, Olivier Bellier¹, Esmail Shabanian¹, Lionel Siame¹, Laëtítia Léanni¹,
Didier Bourlès¹, Seiran Ahmadian²

¹ *Centre Européen de Recherche et d'Enseignement de Géosciences de l'Environnement (CEREGE, UMR CNRS – Aix-Marseille Université, IRD et Collège de France), Université Paul Cézanne, BP 80, 13545 Aix-en-Provence Cedex 4, France. E-mail: bita_javidfakhr@yahoo.com*

² *Islamic Azad University, Shahr-e-Qods Branch, Iran.*

**) corresponding author*

Abstract:

The Alborz is a region of active deformation within the Arabia-Eurasia collision zone. The Abr and the Khij Faults are two NE-trending left-lateral strike-slip faults in the eastern Alborz that correspond to the Shahrud fault system extended through an area of about 95×55 km². Tectonic landforms typically associated with active strike-slip faults, such as deflected stream channels, offset ridges and fault scarps are documented along the mentioned faults. Detailed analyses of satellite images and digital topographic data accompanied by field surveys allowed us to measure horizontal offsets of about 420 ± 50 m and 400 ± 50 m for the Abr and Khij faults, respectively. A total of 8 quartz-rich samples were sampled and dated from two different fan surfaces using in situ-produced ¹⁰Be cosmogenic dating method. Minimum exposure ages for the abandonment of the alluvial fan surfaces of 115 ± 14 kyr along the Abr Fault and of 230 ± 16 kyr along the Khij Fault imply that both faults are active with slip rates of about 3-4 mm yr⁻¹ and 1-3 mm yr⁻¹, respectively. The results of our study provide the first direct quantitative geological estimates of slip rate along these two active

faults and place a new constraint on slip distribution between the faults in the eastern Alborz. Fault kinematic studies (from fault slip data) indicate a N35°E-trending maximum stress axis comprising a dominant strike-slip regime in agreement with the geomorphological analyses. The left-lateral strike-slip faulting along the Abr and Khij faults and their associated fault zones in the eastern Alborz can be due to the westward component of motion of the South Caspian Basin with respect to Eurasia and Central Iran.

Keywords: Alborz; active tectonics; Abr Fault; Khij Fault; cosmic ray exposure dating; slip rate

1. Introduction

Iran is located within the Alpine-Himalayan orogenic belt in a zone of continental convergence between Eurasia to the North and Arabia to the South which is mostly accommodated by shortening and strike-slip faulting in the mountain belts such as the Great Caucasus, Zagros, Alborz, Koppeh Dagh and also the active Makran subduction zone. Recent geological studies indicate that the Arabia-Eurasia collision in Iran occurred between late Eocene and late Oligocene time (e.g., Allen *et al.* 2004; Agard *et al.*, 2005; Vincent *et al.*, 2005; 2007; Allen and Armstrong, 2008; Fakhari *et al.*, 2008; Horton *et al.*, 2008; Boulton *et al.*, 2009). The Arabia-Eurasia convergence rate is reported to be about 22 mm yr⁻¹ at the longitude of Bahrain south of the Persian Gulf (Sella *et al.*, 2002; McClusky *et al.*, 2003; Vernant *et al.*, 2004a; 2004b; Reilinger *et al.*, 2006).

The Alborz accommodates the motion between the South Caspian Basin (SCB) and Central Iran (Figure 1A). The Alborz range has evolved from the Triassic to the present-day during the Cimmeride and Alpine Orogenies (Alavi, 1996). The last major tectonic event that led to widespread deformation in northern Iran and the SCB is coeval with middle Miocene to Recent collision-related compression (e.g., Sengor & Kidd, 1979; Berberian & Berberian,

1981). The south Caspian shorelines move northwest at a rate of about 5 mmyr^{-1} . The Central Iranian block is defined by internal deformation velocities smaller than 2 mm yr^{-1} (Vernant *et al.*, 2004b). The Central Iran northward velocities at the longitudes of Damghan and Jajarm have been reported to be of about 9 mm yr^{-1} and 7 mm yr^{-1} , respectively (Masson *et al.*, 2007). According to Masson *et al.* (2007), at the longitude of Tehran, the Alborz mountain range accommodates a motion rate of about 6 mm yr^{-1} which is in agreement with the results obtained from regional GPS network in the Alborz (Vernant *et al.*, 2004b). The left-lateral shear across the Alborz is constrained to $\sim 4 \text{ mm yr}^{-1}$ (Vernant *et al.*, 2004a). The large amount of geological, geodetic, geomorphological and geophysical data available for the Alborz allows characterizing and qualifying the deformation of this complex zone (*e.g.*, Axen *et al.*, 2001; Jackson *et al.*, 2002; Allen *et al.*, 2003a; Masson *et al.*, 2006; 2007; Guest *et al.*, 2006a; 2006b; 2007; Ritz *et al.*, 2006; Ritz, 2009; Tatar *et al.*, 2007; Ballato *et al.*, 2008; Landgraf *et al.*, 2009; Nazari *et al.*, 2009; Hollingsworth *et al.*, 2010; Solaymani Azad *et al.*, 2011). The N-S convergence of Central Iran together with the SW-ward motion of the SCB toward Central Iran results in a NNE-trending transpressional tectonic regime in the Alborz mountain range. This transpression could have initiated between 3 and 7 m.y. ago (Ritz *et al.*, 2006).

The Kopeh Dagh mountain range includes Mesozoic and Tertiary sediments, which were folded during the Oligo-Miocene orogenic movements (Stocklin, 1968; 1974; Afshar Harb, 1979; Lyberis & Manby, 1999). The right-lateral Main Kopeh Dagh Fault system (MKDF) is characterized as the northwestern boundary of the Kopeh Dagh range (Figure 1) and considered as a seismically active structure (Trifonov, 1978). Previous studies (Regard *et al.*, 2005; Authemayou *et al.*, 2006; 2009; Le Dortz *et al.*, 2009; Shabanian *et al.*, 2009a; 2009b) proposed estimates of long-term slip rates for some of the major active fault systems in Iran. A total slip rate of $9 \pm 2 \text{ mm yr}^{-1}$ is proposed for strike-slip faulting in the Kopeh Dagh,

resolving for western Kopeh Dagh as an average northward and westward slip rates of about 8 and 4 mm yr⁻¹, respectively (Shabanian *et al.*, 2009a).

The South Caspian Basin behaves as a relatively rigid aseismic block (Priestley *et al.*, 1994) and is suggested to be a trapped oceanic remnant (Berberian, 1983). This implies that uplift of the high Alborz and Talesh mountains loaded the SCB, causing accumulation of thick late Cenozoic deposits (Axen *et al.*, 2001; Allen *et al.*, 2003b). The SCB is surrounded by orogens including the Great Caucasus, Alborz, Kopeh Dagh, and the Great Balkhan. Based on the motion of a single GPS site on the Caspian shoreline, the SCB moves northwest with respect to Eurasia at a rate of 6 ± 2 mm yr⁻¹ and the motion across the central Alborz is suggested to be roughly 5 ± 2 mm yr⁻¹ of N-trending shortening and 4 ± 2 mm yr⁻¹ of left-lateral strike-slip (Vernant *et al.*, 2004b). A higher estimate of 11 ± 2 mm yr⁻¹ has been proposed for the South Caspian-Eurasia motion by Copley & Jackson (2006). Ritz *et al.* (2006) mentioned Pleistocene as the age of the SCB northward motion towards Eurasia and suggest that this may have caused not only the change from a compressional tectonic regime to a transpressional one but also the transtension in the internal parts of the range. They emphasize that the internal domain of central Alborz is affected by an active transtensional regime with a WNW-trending extensional axis.

In the current study we present a morphotectonic analysis along two NE-trending left-lateral strike-slip faults, i.e., the Abr and Khij faults (Figure 2) which belong to the Shahrud fault system affecting the eastern Alborz. The determination of the kinematics of the active deformation along the Shahrud fault system using fault slip vectors measured in the field combined with the determination of the abandonment age of two alluvial fan surfaces offset along the faults using in situ-produced cosmogenic ¹⁰Be exposure dating of quartz pebbles allow us to establish the rate and kinematics of the late Quaternary deformation in the southeastern limit of the eastern Alborz. At the regional scale, tectonic implications of the

new data and deduced results presented in this paper provide new geological constraints on active tectonics of the Alborz Mountains.

2. Seismo-tectonic setting and structural framework

2.1. Tectonics

The Alborz is located south of the SCB consisting late Precambrian to Quaternary deposits (Figure 1A). Tectonic evolution of the Alborz range results from a combination of thrust activity within the orogen and, based on the chronology of the foreland basin deposits (Ballato *et al.*, 2008), of frontal accretion. There are different ideas about the geometry and kinematics of faulting in the Alborz range (e.g., Jackson *et al.*, 2002; Allen *et al.*, 2003a; Guest *et al.*, 2006a; Ritz *et al.*, 2006; Zanchi *et al.*, 2006; Yassaghi *et al.*, 2008; Landgarf *et al.*, 2009). For instance, Ritz *et al.* (2006) indicated that the margins of Alborz are affected by a transpressional regime, while there is a general left-lateral shear inside the central Alborz range. They have interpreted the total kinematic view of the Alborz as a general strike-slip regime with a change in the position of σ_1 from a horizontally NNE-trending to a vertical one between the margins and the internal domain.

The compressive deformation has been active since Oligocene (Allen *et al.*, 2003a; 2003b). There are active faults along the southern part of the mountain belt and also along the Caspian coast (Bachmanov *et al.*, 2004). The present-day tectonic framework of the range is characterized by a NE-SW compression causing a lateral inversion of the sense of motion from right to left-lateral (Allen *et al.*, 2003a) as well as the activation of ESE-trending thrust faults. Late Cenozoic tectonics in the eastern Alborz is characterized by oblique and strike-slip faulting which reactivated inherited Mesozoic structures. The active tectonic framework of the eastern Alborz consists of major NE-trending left-lateral faults. One of the main structures in northern margin of the Alborz range is the Khazar Fault extended in western part

of our study area (Figure 1). In eastern part of the Alborz range, the Gorgan Schists are thrust against Quaternary deposits by the Khazar Fault.

2.2. Seismicity

A global record of strain distribution over the last century is available by instrumentally recorded earthquakes, but the interval times between seismic events on individual faults can be up to several thousands of years. Following, we summarize the most important historical and instrumental earthquakes in the eastern Alborz.

A relatively complete catalogue of the historical earthquakes in north-eastern Iran is available (Tchalenko, 1975; Ambraseys & Melville, 1982; Berberian & Yeats, 1999; 2001). The largest earthquakes in eastern Alborz are reported in 856, 1825, 1890, 1953 and 1985 with magnitudes ranging from 6.2 to 7.9 (based on the macroseismic data). 200,000 persons were killed in the catastrophic earthquake of 856. According to a recent paleoseismology study in eastern Alborz (Hollingsworth *et al.*, 2010), the Astaneh Fault's rupture is consistent with rupture during the historical 856 A.D. Qumis (Ambraseys & Melville, 1982) earthquake in NE Iran. Historical earthquakes are also reported in 874, 1436, 1470 and 1498 in the area of Gorgan and Gonbad-Kavous. The destructive earthquakes of 1890 in the Shahrud region could be concerned with the active faults of this study on the basis of the reported regions of maximum destruction (Ambraseys & Melville, 1982). The earthquake of 1944 caused major damage in Gorgan (Ambraseys & Melville, 1982). A left-lateral faulting has been involved in the Karnaveh earthquake of 1970 (Mw 6.4), based on the analysis of its aftershock zone (Jackson & Fitch, 1979). The earthquake of 1985 occurred in the northern part of the Abr Fault (Table 1, Figure 1B).

According to morphotectonic and paleoseismologic studies (Hollingsworth *et al.*, 2010) frequent stream offsets along the trace of the Astaneh Fault (eastern Alborz) may record the

cumulative offset of the last two to five earthquakes in Holocene. We emphasize that eastern Alborz is more seismically active than western Alborz. Few earthquakes occurred in the SCB and it is thus considered seismically relatively quiet although it is surrounded by very active fault zones (e.g., Priestley *et al.*, 1994).

3. Shahrud Fault System

The 400-km-long Shahrud fault system is characterized by active NE-trending left-lateral strike-slip fault zones (Astaneh, Shahrud, Abr and Jajarm fault zones) extending at the SE flank of the eastern Alborz. The combined geomorphic characteristics suggest coeval strike-slip and reverse faulting mechanism for the ongoing activity of this fault system, characterizing a transpressional tectonic regime that is confirmed by the kinematic inversion results (Javidfakhr, 2010).

The left-lateral Astaneh fault zone (the westernmost part of Shahrud fault system) cuts across the topography north of Damghan city. This NE-trending fault zone runs across the southern boundary of eastern Alborz for over 150 km (Hollingsworth *et al.*, 2008; 2010). Western part of Shahrud fault system includes the Shahrud fault zone defined by two active faults, the 130-km-long Shahrud Fault and the 70-km-long Tazareh Fault (Figure 2). The geomorphic expression of the Shahrud Fault is characterized by abrupt fronts of the mountain range. The high-angle Shahrud Fault is a combination of the E- and NE-trending fault segments running along the eastern Alborz Mountains.

Eastern part of the Shahrud fault system consists of multiple fault zones: (1) the Jajarm fault zone which is an extended fault zone in this broad fault system, (2) the Cheshmeh-Nik fault zone which, at a first glance, seems to be the continuation of the Abr fault (Figure 2). However, this is not certain given the gap visible on satellite images of about 1 km between them at the eastern termination of the Abr Fault. the southern part of Shahrud fault system can

be described as the northern margin of Central Iran. The Jajarm Fault is a left-lateral strike-slip fault located in the north of Jajarm city running for about 150 km in the NE direction (e.g., Hollingsworth *et al.*, 2008; Javidfakhr, 2010). The active Jajarm Fault passes through Quaternary and Plio-Quaternary deposits in most of its trace. The Jajarm Fault is parallel to another left-lateral fault segment along its north side, i.e., Cheshmeh-Nik Fault (Figure 2). The area between the Jajarm and Cheshmeh-Nik faults is covered by Neogene continental red beds and Jurassic limestones as well as scattered Quaternary deposits. The Abr and Khij faults are two active boundary faults, located in the central part of the Shahrud fault system, that have separated the mountainous Alborz region from Central Iran.

3.1. The Abr fault zone

The geomorphic expression of the fault zone in the context of the obtained fault kinematic data suggests a reverse component of faulting in the Abr fault zone. The Abr Fault extends for about 95 km in the eastern Alborz and runs northeast between Nekarman and Tilabad villages (Figures 2B, 3A). The NE-trending Abr Fault is a left-lateral strike-slip fault with a reverse component of dip-slip that marks a distinct topographic boundary between mountainous regions and Quaternary basins (Figure 3E), displaying morphological evidence of active faulting (Figures 3C and 3D). The deflection of rivers across this segment of the Abr Fault suggests this segment has been active in the Quaternary (Figure 3C).

The Abr Fault is parallel to the Khij Fault in most parts of its length, with a spacing of ~20 km between them. The trend of the Abr Fault changes in its 40 km long eastern segment and turns to the East. In the termination zone of the Khij Fault, near Tilabad village, the spacing between the two faults decreased to 2 km due to a change in the orientation of the Abr fault trace (Figure 3A).

There are no geometric and/or structural changes at fault terminations. At the termination zone, relative fault displacement reduces to zero and is transferred to other structures, laterally or vertically, away from the fault end-point (Legg *et al.*, 2004). The process of fault termination depends on scale since faults interact with various features of different length scales and since also the length scale of a fault increases during fault growth (Alessio *et al.*, 2004). The trend of the Abr Fault changes noticeably in a relatively wide area of its western termination zone (Figure 3B). The Quaternary trace of the Abr Fault terminates in a cluster of distributed small faults within a horsetail structure arrangement (Figure 3B).

There is a 12-km-long active reverse fault in the south of Nekarman village parallel to the trace of the Abr Fault (Figure 4A). The outcrop of this fault was limited to few sites. Vertical offset was estimated from the profile perpendicular to the fault trace which is shown in Figure 4B. The vertical offset of this fault is estimated at about 24 m (Figure 4C) using the 1:20,000 topographic map.

3.2. The Khij fault zone

The left-lateral Khij Fault passes through Khij village (Figure 3A), and strikes NE, parallel to the eastern Alborz range for about 55 km. The trend of the Khij Fault remains constant all along its trace. At its both ends, the trace of the fault remains sharp. There are some nearly parallel small faults (10-15 km) in the central and western part of the fault. These faults offset the associated terraces left-laterally for several tens of meters (Figure 3D). They probably should be responsible for a negligible part of left-lateral slip across the fault zone. Some channels in the western part of the Khij Fault are deflected with left-lateral offsets of about 20-100 m (Figure 3D). Vertical deformations such as small scarps with reverse faulting arise along the Khij Fault, but most of the features are locally observed and result from strike-slip faulting.

4. Fault kinematic analysis

4.1. Methodology

The kinematic measurements were numerically processed to analyze the fault population data (Carey and Brunier, 1974; Carey, 1979), a mean best-fitted deviatoric stress tensor for a fault population being calculated through the minimization of the angular deviation between the measured striation and the resolved shear stress. This inversion method assumes that the slip established by the striation is in the same direction as the resolved shear. The orientation of the principal stress axes (σ_1 , σ_2 , σ_3) and the stress ellipsoid shape parameter (R), defined as $R = (\sigma_2 - \sigma_1)/(\sigma_3 - \sigma_1)$, are the results thus obtained through the applied inversion method (Carey and Brunier, 1974; Mercier *et al.*, 1991; Bellier and Zoback, 1995).

Relative chronology of the striations and their coherence with the regional tectonic events are considered to separate the distinct families of striation. We detected two slip generations of striations revealed by clear crosscutting striations on fault planes in most of the measurement sites. Just one slip direction on a given fault plane can be produced by a distinct stress deviator (σ_1 , σ_2 , σ_3 and R) due to the assumptions of the fault kinematic inversion methods. Three successive stress states are distinguished through analysis of distinct datasets after the separation of the slip generations (Javidfakhr, 2010).

4.2. Modern state of stress deduced from the fault kinematic analysis

The modern state of stress was analyzed using fault kinematics indicated by the last generation of striation sets measured mostly in Quaternary deposits (Table 2). Table 2 consists of the inversion results of the fault slip data accompanied by the lithology definition of different sites. It shows that R values (that must lie between 0 and 1) are in the range 0.63 to 0.98. All the tensors representative of the modern stress state indicate the same direction of

σ_1 stress axes despite different formation ages. For example, the trend of σ_1 at site number 1 comprising Cretaceous age limestone is similar to that at other sites measured in Quaternary or Neogene deposits. Figure 5 indicates the location of the measurement sites and the state of stress deduced from fault kinematics inversion. According to the inversion results for a major part of sites, the modern stress state is characterized by a NE-trending σ_1 ranging from N14°E to N46°E and indicates a relatively homogeneous stress field.

It is nevertheless possible to distinguish two specified regimes in the studied region. 9 out of 11 sites indeed show strike-slip regimes (with a minor reverse slip component) while two sites (Figure 6, sites number 1 and 8B) indicate mostly compressional stress regime. The inversion results obtained from five measurement sites along the Abr Fault trace characterized a relatively homogenous transpressional stress regime (Figure 5). At site number 8, two diagrams of inversion solution with σ_1 trending N66°E and N12°E for the diagram 8A and 8B, respectively (Figure 6), are obtained. However, there is a minor fault almost parallel with the Khij Fault in the vicinity of this number 8 site accompanied by an older fault (almost perpendicular to the trend of Khij Fault) cut by the Khij Fault. This intersection of faults may be the reason of this local variation. Structural framework of the region is thus simply explained by strike-slip and compressional stress deviators (Figure 6). Considering the fault geometry, the dominant left-lateral strike-slip faulting along the Abr and Khij faults is in agreement with an average computed stress deviator, N30°E trending horizontal σ_1 .

5. Geomorphic investigations

5.1. Methodology: mapping and site selection

The detailed structure of strike-slip fault zones and their evolution are basic subjects in understanding of the ongoing deformation in this region. Major fault segments have been mapped in order to recognize the structural and geomorphic characteristics of the fault

system. Different types of satellite images such as LANDSAT ETM+ (14 m resolution), SPOT5 images (5 m resolution) and SRTM data were used to improve the visualization of the traces of the fault systems and better constrain their geometry considering the structural linkage between the different fault segments.

The observation of the recent deposits (Pleistocene until present time) surface morphology allows quantifying the cumulative tectonic displacement which occurred during the last million years. In the current study, sites indicative of distinct fault activity were selected to estimate the fault slip rates. Firstly, the morphology of each site was precisely mapped by combining field observations with evidences from SPOT5 images in order to determine horizontal offsets through examination of the best match possible for the trend of streams incising Quaternary deposits. Study of the segments along the Abr and Khij faults offered evidences for late Quaternary geomorphic features associated with strike-slip faulting, such as stream offsets including tail-cut and head-cut streams and alluvial fan offsets.

5.2. Offsets

5.2.1. Offset of the Abr alluvial fan

The geomorphic expression along the Abr Fault is characterized by left-laterally offset features such as alluvial fans, stream channels and shutter ridges. Figure 7A shows the morphology of the fan system chosen because it is well preserved from surface outwash and fluvial erosion. In addition to a feeder channel clearly left-laterally offset by the fault, their distinct geomorphology and elevations allow distinguishing five main surfaces. From younger to older, they are classified as Q1 to Q5 (Figure 7B). The geomorphic relationship clearly indicates that the feeder channel was emplaced after the abandonment of the Q3 unit.

Figure 7C shows the geomorphic map of the Abr fan reconstructed to measure the cumulative offsets. With ~700 m of back-slip, the alignment between the fan's feeder channel

and the fan axis is restored. The fan's axial channel is beheaded and has a left-laterally offset with respect to its feeder channel. This reconstruction assumes that no rotation occurred. However, some rotation north of the fault zone would only increase the offset by about 20-30 m.

The amount of left-lateral offset ranges from ~60 m (Figure 7D) for the younger drainages incised in alluvial fans located in the west of the site area to ~700 m at the dating site (Figure 7C). It seems that the most likely amount of left-lateral offset at this site is ~420 m considering the active stream deflection (~280 m) shown in Figure 7E. However, sub-parallel faults locating in the area may accommodate additional left-lateral motion across the fault zone.

5.2.2. Offset of the Khij alluvial fan

The Khij Fault clearly cuts across the Quaternary surfaces incised by abundant streams in an extended area comprising the dating site (Figure 8A). The strike-slip component dominates the reverse component along the fault (Figures 8b, 9). Along its straight trace, the fault left-laterally offsets the Quaternary landforms such as fan surfaces and river channels (Figure 9).

The restoration of different fan surfaces in the Khij site is difficult due to effects of erosion in the area. Figure 9 illustrates our reconstruction of offsets across the fault zone using the high resolution Quickbird image exhibiting the obvious left-lateral displacement of streams. The actual offset is about 400 m because the trend of the active drainages is relatively similar in both the upstream and downstream segments (Figure 9). In this area, it is difficult to distinguish the original fan boundaries. We suggest that the uncertainty in offset across the Khij Fault is no more than 50 m because of the high similarity in trend of the offset features north and south of the fault zone. However, these uncertainties have been considered

in establishing bounds on the value of offset. Finally, we suggest 400 ± 50 m of back-slip to restore the offset, which is the best estimate deduced from the reconstruction of three major offset markers in the proximity of the dating site (Figures 9B, 9C). The displacement of the drainage system is surely younger than the displacement of the fan surfaces. Because the streams incised in the fan surfaces are younger than the fan abandonment, the cumulative offset value of 400 ± 50 m is considered as the minimum horizontal offset possible for this segment of Khij Fault. Moreover, digital topography maps with the scale of 1:20,000 were used to determine the vertical displacement across the Khij Fault. Vertical offsets were estimated from the profiles perpendicular to the fault trace after restoring the coeval lateral offset (Figure 8B). The dating site and its surrounding alluvial fans are shown in Figure 8A. The vertical displacement for this segment of the Khij Fault is about ~ 12 m (Figure 8B) which represent the prominent strike-slip character of the Khij Fault. The Quaternary surface is gently tilted over Neogene deposits in a site located in the vicinity of Khij dating site (Figure 10).

6. Cosmogenic dating and fault slip rates

Since the progress made over the last decades by the Accelerator Mass Spectrometry (AMS) technique, cosmogenic nuclides are extensively used to quantify the processes affecting and shaping the Earth's surface (see review in Gosse and Phillips, 2001; Dunai, 2010). In particular, absolute dating of landforms for deciphering climatic history or tectonic history has been more specifically developed and the concepts and principles of the Cosmic Ray Exposure (CRE) dating method relatively well established. Due to its half life of 1.39 ± 0.01 My (Korchinek *et al.*, 2009; Chmeleff *et al.*, 2010) and because its production is well known in quartz mineral, in situ-produced ^{10}Be is an adequate candidate to date surface

samples to 2 My in best conditions (Guralnik *et al.*, 2010) or more frequently from several kyr up to 1 My depending of the denudation rate affecting the samples.

Samples were prepared for AMS ^{10}Be measurements following chemical procedures adapted from Brown *et al.* (1991) and Merchel & Herpers (1999). Samples were first crushed and sieved. Pure quartz was obtained from the 250-500 μm fraction by repeated H_2SiF_6 - HCl etching; atmospheric ^{10}Be was subsequently eliminated by sequential dissolutions with diluted HF. Prior to complete HF digestion of silica, a weighted 100 μl of a 3025 ppm home-made carrier solution was added to allow for $^{10}\text{Be}/^9\text{Be}$ ratio calculation after AMS measurement. The spiked solutions were finally purified by solvent extractions of Be acetylacetonate in presence of EDTA followed by precipitations of $\text{Be}(\text{OH})_2$ at pH 8.5 and rinsing. The final precipitate, dissolved in a few drops of HNO_3 , is dried and heated at 900°C to obtain BeO . Beryllium oxide was mixed to 325 mesh niobium powder prior to measurement at ASTER, the French AMS national facility located at CEREGE, Aix en Provence. All ^{10}Be concentrations are normalized to $^{10}\text{Be}/^9\text{Be}$ SRM 4325 NIST reference material with an assigned value of $(2.79 \pm 0.03) \cdot 10^{-11}$. This standardization is equivalent to 07KNSTD within rounding error. The ^{10}Be half-life of $(1.39 \pm 0.01) \times 10^6$ years used is that recently recommended by Korschinek *et al.* (2009) and Chmeleff *et al.* (2010) according to their two independent measurements. Analytical uncertainties (reported as 1σ) include a conservative 0.5% external uncertainty based on long-term measurements of standards, a one sigma statistical error on counted ^{10}Be events, and the uncertainty associated with the chemical blanks correction. Two chemical blanks prepared with the samples yield similar $^{10}\text{Be}/^9\text{Be}$ ratio of $(3.3 \pm 0.8) \times 10^{-15}$ and $(3.4 \pm 0.7) \times 10^{-15}$, 2 orders of magnitude smaller than the minimum measured ratio. An additional 6% production rate uncertainty is considered for CRE age calculation (Stone, 2000).

A modern ^{10}Be production rate at sea-level and high-latitude of 4.5 ± 0.3 atoms/g/yr, computed for internal consistency from the data of Stone (2000) according to the conclusions of the recently published study on absolute calibration of ^{10}Be AMS standards by Nishiizumi *et al.* (2007), was used. This sea-level and high-latitude production rate has then been scaled for the sampling altitudes and latitudes using the scaling factors proposed by Stone (2000) because, using the atmospheric pressure as a function of altitude, they take into account the physical properties of cosmic ray particle propagation in the atmosphere and include an improved account for the muonic component in the total cosmogenic production. Corrections for local slope and topographic shielding due to surrounding morphologies following Dunne *et al.* (1999) appear negligible for both sampling sites.

Erosion is a major source of uncertainty in the surface exposure dating. Considering the very low erosion rate of 0.002 mm yr^{-1} estimated for the alluvial fan surfaces of Central Iran using ^{10}Be and ^{36}Cl cosmogenic nuclides (Le Dortz, 2010) and the geomorphic freshness of the Abr and Khij alluvial fans, erosion was reasonably considered as negligible throughout this work. This implies that the CRE ages calculated from the measured ^{10}Be concentrations are minimum exposure ages.

The ^{10}Be concentration in the sampled surface quartz boulders being directly related to alluvial fan abandonment, alluvial fans that have been preserved during left-lateral displacements along the studied faults have been dated in order to estimate their long-term slip rates. Long-term slip rates on strike-slip faults can indeed be obtained by dividing horizontal geomorphic offset by the CRE age of the surface considering the uncertainties associated to the different applied methods.

To minimize the possibility of rotation or rolling, boulders embedded into the ground surface with top surfaces as flat as possible were selected. Even if 6 boulders were collected on each fan, only 3 and 5 samples at the Abr and Khij sites, respectively, provided data

because the other samples failed through chemical processing due to their unsuitable lithology for ^{10}Be cosmogenic dating. The measured ^{10}Be concentrations and the calculated minimum CRE ages are presented with their associated uncertainties in Table 3.

6.1. Exposure ages and the slip rate of the Abr Fault

Along the Abr Fault, in situ-produced ^{10}Be concentration have been measured in samples collected from the Q3 fan surface which has been offset by 420 ± 50 m (Figure 7B). Three dated samples yield minimum exposure ages of 96 ± 6 kyr, 118 ± 8 kyr and 131 ± 9 kyr (Table 3). The observed low dispersion of the Q3 surface ages may result from different rates of exhumation to the surface of sampled boulders that were previously buried within the alluvial deposits. If this is the case, the oldest CRE age yields the more reliable fan surface abandonment age. If this low dispersion results from different erosion rates affecting the surface sampled boulders, the more ^{10}Be concentrated sample yielding the oldest CRE age is the less affected by erosion and thus the sample which again yields the more reliable fan surface abandonment age. Considering then the upper bound of the CRE age series as the more reliable fan surface abandonment age and the horizontal offset of 420 ± 50 m recorded by this surface yields a minimum slip rate of 3.2 ± 0.5 mm yr $^{-1}$ along the Abr Fault.

6.2. Exposure ages and the slip rate of the Khij Fault

The CRE ages calculated from the in situ-produced ^{10}Be concentrations measured in the five quartz-rich pebbles embedded in the fan surface at the Khij site are generally older than those calculated from the Abr site samples. They indeed range from 152 ± 10 to 364 ± 24 kyr. More precisely, three samples, statistically belonging to the same population, cluster around a weighted-mean CRE age of 165 ± 6 kyr while the two other samples are significantly older (Table 3). The cumulative offset of the fan surface is most likely between 350 and 450 m.

There is two ways to interpret this observed distribution. Either, as discussed above, erosion can reasonably be considered as negligible and the two older ages result from inherited ^{10}Be acquired during a previous exposure episode in the area. In that case, the weighted-mean ^{10}Be minimum CRE age of 165 ± 6 kyr is the best estimate for the fan surface abandonment age (Figure 11). Or, the negligible erosion rate assumption is not valid, and the oldest CRE age of 364 ± 24 kyr yields the more reliable fan surface abandonment age. In this case, the younger ages result from different erosion processes affecting the fan surface. Considering both interpretations, the cumulative 400 ± 50 m horizontal offset recorded by the Khij Fault constrains the Quaternary left-lateral slip rate on this segment of the fault to range between 2.4 ± 0.3 mm yr $^{-1}$ to 1 ± 0.2 mm yr $^{-1}$.

The dated fan has been tectonically uplifted ~ 12 m based on quantitative analysis of the slope profile (Figure 8). Taking into account the errors involved in the mean vertical offset measurement and the uncertainties associated to the CRE ages, the vertical slip rate along this segment of the Khij Fault is calculated to be less than 0.07 mm yr $^{-1}$.

7. Discussion and conclusion

Active deformation in the northeastern boundary of the Arabia-Eurasia collision zone is taken up by strike-slip faults, which is oblique to the convergence direction in this boundary zone (Vernant *et al.*, 2004b; Reilinger *et al.*, 2006). The whole compression in Alborz mountain belt seems to be roughly orientated north-south with a rate of about 8 ± 2 mm yr $^{-1}$ and the shortening rate absorbed by the Alborz and the SCB is about 14 ± 2 mm yr $^{-1}$ (Vernant *et al.*, 2004a).

It is essential to distinguish the present-day tectonic configurations in active deformation domains. This paper focused on a major fault system at the SE boundary of eastern Alborz. The major strike-slip faults are mainly located within the internal parts of the Alborz range

(e.g., Allen *et al.*, 2003a). Active left-lateral strike-slip faults trend ENE and WNW in the eastern (Allen *et al.*, 2003a; Ritz *et al.*, 2006) and western Alborz (Axen *et al.*, 2001; Guest *et al.*, 2006a; 2007), respectively. It is noticeable that there are much more evidences of left-lateral strike-slip faulting in the eastern Alborz (e.g., Hollingsworth *et al.*, 2010) than in the west. Allen *et al.* (2003a) proposed a model for late Cenozoic structural evolution of the Alborz. They suggested that the continued compressional deformation (probably initiated in Miocene) is accompanied by left-lateral strike-slip faulting along the length of the Alborz range during Plio-Quaternary accommodating the westward motion of the SCB basement with respect to Central Iran.

The ongoing deformation in the Tehran's piedmont (central Alborz) is accommodated along N070°E to N110°E trending left-lateral strike-slip faults and low angle dip thrusts (Abbassi & Farbod, 2009). Ritz *et al.* (2008) estimated a total left-lateral displacement along the active strike-slip faults in central Alborz of about 3-6 km and a cumulative left-lateral displacement of about 3 km along the central Mosha Fault (central Alborz) is estimated by Solaymani Azad *et al.* (2011). Based on their recent morphotectonic analysis within central Alborz (Solaymani Azad *et al.*, 2011), the Mosha and North Tehran faults are active left-lateral strike-slip faults displaying a left-stepping en-echelon pattern. Hollingsworth *et al.*, (2010) suggested that frequent stream offsets of 15-20 m along the active trace of the Astaneh Fault could record the cumulative offset of the last two to five earthquakes during Holocene.

The Abr and Khij faults strike parallel to the structural trend of the Alborz range. The trace of the Abr Fault corresponds to the contact between pre-Neogene rock units in north and Quaternary deposits in south. The major neighboring structures such as the Khazar, Shahrud and Jajarm fault zones can significantly influence the fault zone behavior and motion rate in the eastern Alborz. However, the left-lateral motion of the Shahrud fault system and the right-lateral one on the Main Kopeh Dagh Fault are simple explanations for the westward motion of

western Kopeh Dagh toward the SCB (e.g., Shabaniyan *et al.*, 2009b). This mechanism is due to the oblique convergence of this boundary zone. The Abr and Khij left-lateral strike-slip faults and their associated faults in the Shahrud fault system are involved in the northward convergence of Central Iran and the north-westwards motion of the SCB relative to Eurasia.

We concentrated on the analysis of geomorphic features to enrich present-day tectonic activity of the region. Alluvial fans can record cumulative displacements in tectonically active regions. Accurate dating of geomorphic offsets can lead us for a quantitative interpretation of long term slip rates on major active faults. It also helps to have a better understanding of the regional tectonics and to establish the deformation history of the area. We have presented evidence of late Quaternary activity along the Abr and Khij fault zones in eastern Alborz. Slip rates are calculated for the strike-slip Abr and Khij faults taking into account the geomorphic evidence and the applied ^{10}Be exposure dating method. A long term left-lateral slip rate of $3.2 \pm 0.5 \text{ mm yr}^{-1}$ along the NE-trending Abr Fault has an important implication to our understanding of the deformation in the eastern Alborz. Along the Khij Fault, the horizontal slip rate ranges between $2.4 \pm 0.3 \text{ mm yr}^{-1}$ and $1 \pm 0.2 \text{ mm yr}^{-1}$. The vertical slip rate of 0.07 mm yr^{-1} is much slower than the strike-slip rate of the Khij Fault, indicating that strike-slip is the dominant mechanism of the late Quaternary activity of this fault which is also supported by our geomorphic and fault kinematic data. The absence of clear vertical scarps in the fan deposits is also representative for the dominant component of strike-slip.

In the study area, the active deformation north and east of Bastam city is principally taken up by the Abr and Khij left-lateral strike-slip faults (Figure 2). The Abr fault zone has generated three earthquakes (Mw 4.9-Mw 6) in the past thirty years (Table 1). However, moderate earthquakes could be judged as signals for considerable hazard if they are frequently reported in populated regions. It is fruitful to analyze the earthquake focal mechanisms considering the structural setting of the fault slip data in an extended region to establish more

precise seismotectonic framework of this boundary zone. The regional stress state (N47°E) is analyzed using the inversion of earthquake focal mechanisms in a vast area between central Alborz and western Kopeh Dagh in 52°-56°E (Javidfakhr, 2010). This result is in close agreement with the modern stress state deduced from inversion of fault kinematic data. The modern stress field is characterized by a regional NE-trending horizontal σ_1 (Figure 5) based on kinematic inversion results. All the earthquake focal mechanisms are not compatible with this regional stress state. However, we suggest that the complex dynamic characteristics of the study region enclosed between two deformation systems of Alborz and Kopeh Dagh seem to be the most effective factor to characterize the changes in the attitude of σ_1 axis.

Our results can be compared with GPS data that estimate a left-lateral strike-slip motion rate ranging between 2 and 6 mm yr⁻¹ along the faults in the Alborz (Vernant *et al.*, 2004), to have a better assess to the validity of the results within a regional active deformation framework. It should be noticed that the rate estimate provided by the GPS data, integrates strain over a much larger area.

Detailed geomorphic studies complemented by structural analysis in other parts of the Shahrud fault system will characterize a complete scheme for the geodynamic evolution of this tectonically active zone.

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Highlight

> 8 quartz-rich samples were dated from two fan surfaces using in situ-produced ^{10}Be cosmogenic dating method. > Left-lateral slip rate of $3.2 \pm 0.5 \text{ mm yr}^{-1}$ is calculated for the Abr Fault. > Along the Khij Fault, the horizontal slip rate ranges between $2.4 \pm 0.3 \text{ mm yr}^{-1}$ and $1 \pm 0.2 \text{ mm yr}^{-1}$. > The modern stress state is characterized by a NE-trending σ_1 indicating a relatively homogeneous stress field. >

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Figure captions:

Figure 1. A. Simplified tectonic map of NNE Iran superimposed on GTOPO30 topographic image. The Khazar Fault (KHF) exists in the boundary zone between the South Caspian Basin (SCB) and Alborz Mountains. The Main Kopeh Dagh Fault (MKDF) forms a topographic boundary between Eurasia and Kopeh Dagh Mountains. The black rectangle in the inset box indicates the location of Figure 1 in the Middle East Alpine Collision belt. White arrows and their associated numbers represent the velocity vectors in mm yr^{-1} with respect to Eurasia (Reilinger *et al.*, 2006). B. Fault map of the eastern Alborz and western Kopeh Dagh superimposed on 90 m SRTM (Shuttle-borne Radar Topography Mission) image of the area. Earthquake focal mechanisms are mostly based on the Harvard catalogue <http://www.globalcmt.org/CMTsearch.html> and McKenzie (1972) (see Table 1 for details).

Figure 2. A. Major faults in northeastern Iran superimposed on SRTM image showing left-lateral movement of NE-trending active faults in the Shahrud Fault System. The study area corresponds to the rectangle enlarged in Figure 3.

Figure 3. A. The locations of different sites are identified in this Landsat (14 m resolution) image. Yellow circles are the location of sites where slip rates have been determined by dating of the offset alluvial fans. The abbreviations are as follows: AF: Abr Fault, KHF: Khij Fault. B. SPOT5 (5 m resolution) images focused on areas along the Abr and Khij faults (B-E). This Figure shows the enlargement of the termination zone of the Abr Fault. There are multiple small faults splayed in a rather wide area at the termination of the Abr Fault. C. Left-lateral deflection of streams is indicative of left-lateral movement of the Abr Fault. Fan boundaries are shown by dotted lines. D. The sharp escarpment of the Khij Fault which runs between volcanic (mainly Andesites) rocks and Quaternary. E. Three dimension view of the area on the basis of digital topographic maps with the scale of 1/20,000. The squares indicate

the position of the sampling sites and the fault traces are shown by dotted lines. No altitude exaggeration is exerted in this image.

Figure 4. A. SPOT5 image of a 12-km-long active reverse fault in the south of Nekarman. The dotted white rectangle is enlarged in Figure B. The location and direction view of Figure 6D is shown by the symbol. The trace of the fault is shown with small arrows. B. Topographic map (10 m contours) of this area is superimposed on the SPOT5 image. The A-B line represents the position of the topographic profile. C. The vertical displacement of the fault is estimated about 25 m.

Figure 5. Structural map of the study area superimposed on the SRTM image, comprising the fault slip measurement sites. Trends of σ_1 (maximum stress axis) for the modern stress regime are presented. Sites are identified by numbers, see Table 2 for details. The inferred fault in the westernmost of the figure represents the Khazar Fault.

Figure 6. Lower hemisphere stereographic projections of the fault plane data together with inversion results determined by the inversion method proposed by Carey (1979). The labels refer to the site locations presented in Figure 5. The large arrows indicate the azimuth of σ_1 . A fixed solution (Bellier and Zoback, 1995) was applied to deal with the sites comprising less than four well-distributed fault directions (site number 6). One of the principal stress axes is fixed to remain purely vertical in the fixed inversion method. Measurement sites comprising two tectonic regimes in each state of stress are presented by adding “A” suffix to the identifying number of the site. The suffix “B” is added to the measurement site number in the cases that the data were not enough and/or well-distributed to be reliable as an inversion solution. Note that the trend of σ_1 is similar in site number 1 comprising the limestone of Cretaceous age with other sites measured in Quaternary or Neogene deposits (refer to Table 2).

Figure 7. A. SPOT5 image focused on the Abr fan. B. Geomorphic map of Quaternary alluvial fans along the Abr Fault. Distinct surfaces of different ages are shown for the fan, from younger to older (Q1 to Q5). Solid line indicates the fault trace. Small squares show sampling locations on the alluvial fan surfaces. The associated ages are presented for three samples (The other samples did not have enough quartz for the process of in situ exposure ^{10}Be dating). There is an offset between the fan's feeder channel and the actual position of the fan axis. C. A backward slip of about 700 m restores the alignment of the fan's feeder channel and the fan axis. D: The rectangle area located in the west of the dating site shows left-lateral offset of a smaller stream channel for about 60 m. E. True left-lateral offset is considered about 42 m taking into account the active deflection of the stream.

Figure 8. A. Topographic contour map (10 m contours) of Khij area displayed over SPOT5 image. A-B represents the position of the cross section. The small squares are indicative of collected samples. B. Topographic profile along the line A-B shows a vertical displacement of about 12 m along this segment of the Khij Fault. Sampling location and their associated in situ ^{10}Be ages are presented.

Figure 9. A. The Google Earth image (resolution 3 m) clearly shows left-lateral displacement along the Khij Fault through stream channel offsets. B and C. left-lateral offset of ~ 400 m is achieved for this segment of the Khij Fault. Two fault planes nearly parallel with this segment of Khij Fault are shown in the lower stereographic projections.

Figure 10. A. ENE looking view of the escarpment of the Khij Fault. B. NW looking view of the disconformity between the Quaternary and Neogene layers. The Quaternary surface is tilted and uplifted. This site is located in the vicinity of our dating site.

Figure 11. In situ-produced ^{10}Be exposure ages and their associated uncertainties for the samples collected from the fan surface along the Khij Fault assuming no erosion and no

inheritance, The black curve indicates the Gaussian age probability sum (P_{sum}) and thin curves represent the age probability for each sample.

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Table 1. Earthquake source parameters.

N°	Date (yyyymmdd)	Time (hhmm)	Lat. (°N)	Long. (°E)	Mb	Plane 1			Plane 2			Depth (km)	R	ID
						Azimuth	Dip	Rake	Azimuth	Dip	Rake			
1	19700730	0052	37.85	55.92	6.4	293	56	-150	185	66	-38	11	P	6
2	19710214	1627	36.62	55.74	5.3	336	39	93	152	51	87	4	P	7
3	19730802	2028	37.32	56.60	4.9	297	10	86	121	80	91	34	M	11
4	19740307	1136	37.60	55.83	5.1	23	89	-1	113	89	-179	21	JF	12
5	19790224	1356	37.22	56.66	4.7	278	41	56	140	57	116	33	M	13
6	19791209	0912	35.15	56.87	5.6	325	36	99	133	54	83	9	J	15
7	19851201	2031	37.63	56.62	4.6	281	54	69	134	41	116	-	M	18
8	19880311	0738	37.27	56.26	4.5	277	43	13	177	81	132	33	M	22
9	19880508	0650	35.33	55.94	4.7	219	87	4	129	86	177	53	M	23
10	19900818	0951	37.06	56.21	4.9	211	77	-4	302	86	-166	13	M	24
11	19810721	0445	39.41	53.22	5.1	290	87	-2	20	88	-177	42	CMT	35
12	19830325	1157	36.65	52.62	5.2	280	68	5	188	86	157	33	CMT	36
13	19830326	0407	35.88	52.01	5.4	104	61	17	6	75	150	10	CMT	37
14	19851029	1313	36.96	54.59	6.0	97	31	122	241	64	73	15	CMT	39
15	19870907	1132	39.13	54.87	5.5	312	14	106	116	77	86	29	CMT	41
16	19900120	0127	35.89	53.00	5.5	97	65	-8	190	82	-154	25	M	42
17	19890913	0701	37.28	54.25	5.1	262	23	-82	74	67	-92	38	M	43
18	19880113	0556	37.28	54.37	4.9	287	89	3	197	87	179	33	M	44
19	19880508	0650	35.33	55.94	4.7	219	87	4	128	86	177	53	M	45
20	19880823	0530	35.42	52.28	5.0	15	70	45	266	48	152	35	M	46
21	19880823	1058	35.34	52.34	4.6	163	64	129	226	11	323	38	M	47
22	19881024	1701	35.25	52.30	4.9	331	62	34	223	60	147	34	M	48
23	19700730	0052	37.85	55.94	5.7	33	57	5	300	86	146	-	M	49
24	19851029	1313	36.74	54.81	6.0	14	37	26	263	75	124	13	M	50
25	19840511	0958	36.71	55.00	4.9	261	57	-130	139	50	44	29	M	51
26	19840615	1850	37.21	55.39	4.4	277	29	40	150	71	113	-	M	52
27	19830326	0407	36.06	52.28	5.4	31	88	25	300	65	178	33	M	53
28	19820515	1736	35.57	54.11	4.5	96	66	48	342	48	146	13	M	54
29	19821025	1654	35.13	52.38	4.5	185	57	-32	294	63	-142	44	M	55
30	19810809	1308	36.92	55.29	4.9	256	60	-34	6	60	-143	-	M	56
31	19770502	1517	37.06	55.33	5.1	246	70	-24	345	67	-157	19	M	57
32	19760531	0807	38.91	55.85	4.7	43	84	-147	309	58	-6	7	M	58
33	19740307	1136	37.65	55.95	5.2	190	59	-30	297	64	-144	-	M	59
34	19710214	1627	36.62	55.74	5.3	231	51	8	136	84	141	4	M	60
35	19710809	0254	36.27	52.81	5.2	77	36	-70	234	56	-103	12	M	61
36	19680519	1649	36.61	53.35	4.6	254	15	0	346	90	-104	22	M	62

All angles are presented in degrees and date is as the usual convention (year/month/day). Figure 1B represents the focal mechanisms on the basis of these seismic data. The ID numbers refer to the label of each beachball in Figure 1B. It should be noted that the reported depths for individual earthquakes are not entirely reliable due to a poor azimuthal coverage of the worldwide station network.

Table 2. The results of fault kinematic inversions.

Site	Longitude (°E)	Latitude (°N)	σ_1 (trend/plunge)	σ_2 (trend/plunge)	σ_3 (trend/plunge)	N	R	Quality	Lithology	Age (Formation)
1	54.832	36.379	198/06	107/02	358/84	8	0.777	B	Limestone	Cretaceous
2	54.944	36.474	177/10	274/36	74/52	8	0.671	B	Conglomerate	PI-Q
3	54.977	36.445	231/24	138/06	34/65	9	0.659	A	Limestone	Pre-Pliocene
4	55.063	36.496	188/01	279/30	97/60	5	0.640	CF	Conglomerate	Quaternary
5	54.842	36.533	38/14	300/29	150/57	7	0.649	B	Conglomerate	Quaternary
6*	54.913	36.578	191/00	46/90	281/00	4	0.323	CF	Alluvial fan	Quaternary
7	55.076	36.703	208/06	337/79	117/08	22	0.763	A	Alluvial fan	Quaternary
8A	55.142	36.658	66/11	322/50	164/38	12	0.934	B	Limestone	Pre-Pliocene
8B	55.142	36.658	12/18	279/10	162/70	10	0.631	A	Limestone	Pre-Pliocene
9	55.283	36.723	14/3	107/45	280/45	6	0.806	B	Alluvial fan	Quaternary
10	55.358	36.839	46/03	140/50	313/40	7	0.805	B	Conglomerate	Quaternary
11	55.429	36.874	226/5	119/74	318/15	8	0.984	B	Conglomerate	Quaternary

Principal stress axes ($\sigma_1 > \sigma_2 > \sigma_3$) and the stress ratio (R), defined as $R = (\sigma_2 - \sigma_1) / (\sigma_3 - \sigma_1)$ are presented in this table. N is the number of fault slip planes involved in the stress calculation. Site number 6 is resulted through the mentioned “fixed” solution (see text) and is marked by an asterisk. The references for the site numbers are Figures 5 and 6.

Table 3. Sample characteristics and calculated minimum exposure ages.

Sample ID	Altitude (m)	Latitude (°N)	Longitude (°E)	^{10}Be (10^5 atoms/g)	Age (Kyr)
ABR1	1706	36.69	55.08	17.69 ± 1.1	118 ± 8
ABR4	1678	36.69	55.08	14.19 ± 0.92	96 ± 6
ABR5	1683	36.69	55.08	19.22 ± 1.2	131 ± 9
KHI1	1433	36.64	55.27	42.07 ± 2.7	364 ± 24
KHI2	1432	36.64	55.27	34.16 ± 2.8	291 ± 25
KHI3	1442	36.64	55.27	18.58 ± 1.2	152 ± 10
KHI4	1438	36.64	55.27	20.32 ± 1.3	167 ± 11
KHI5	1437	36.64	55.27	21.96 ± 1.4	181 ± 12

Note: ^{10}Be concentrations analytical uncertainties (reported as 1σ) include a conservative 0.5% external uncertainty based on long-term measurements of standards, a one sigma statistical error on counted ^{10}Be events, and the uncertainty associated with the chemical blanks correction. Two chemical blanks prepared with the samples yield similar $^{10}\text{Be}/^9\text{Be}$ ratio of $3.3 \pm 0.8 \times 10^{-15}$ and $3.4 \pm 0.7 \times 10^{-15}$. For CRE age calculations, a modern sea-level high-latitude ^{10}Be production rate at of 4.5 ± 0.3 atoms/g/yr has been scaled for the sampling altitudes and latitudes using the scaling factors proposed by Stone (2000). An additional 6% production rate uncertainty is considered for CRE age calculation (Stone, 2000).

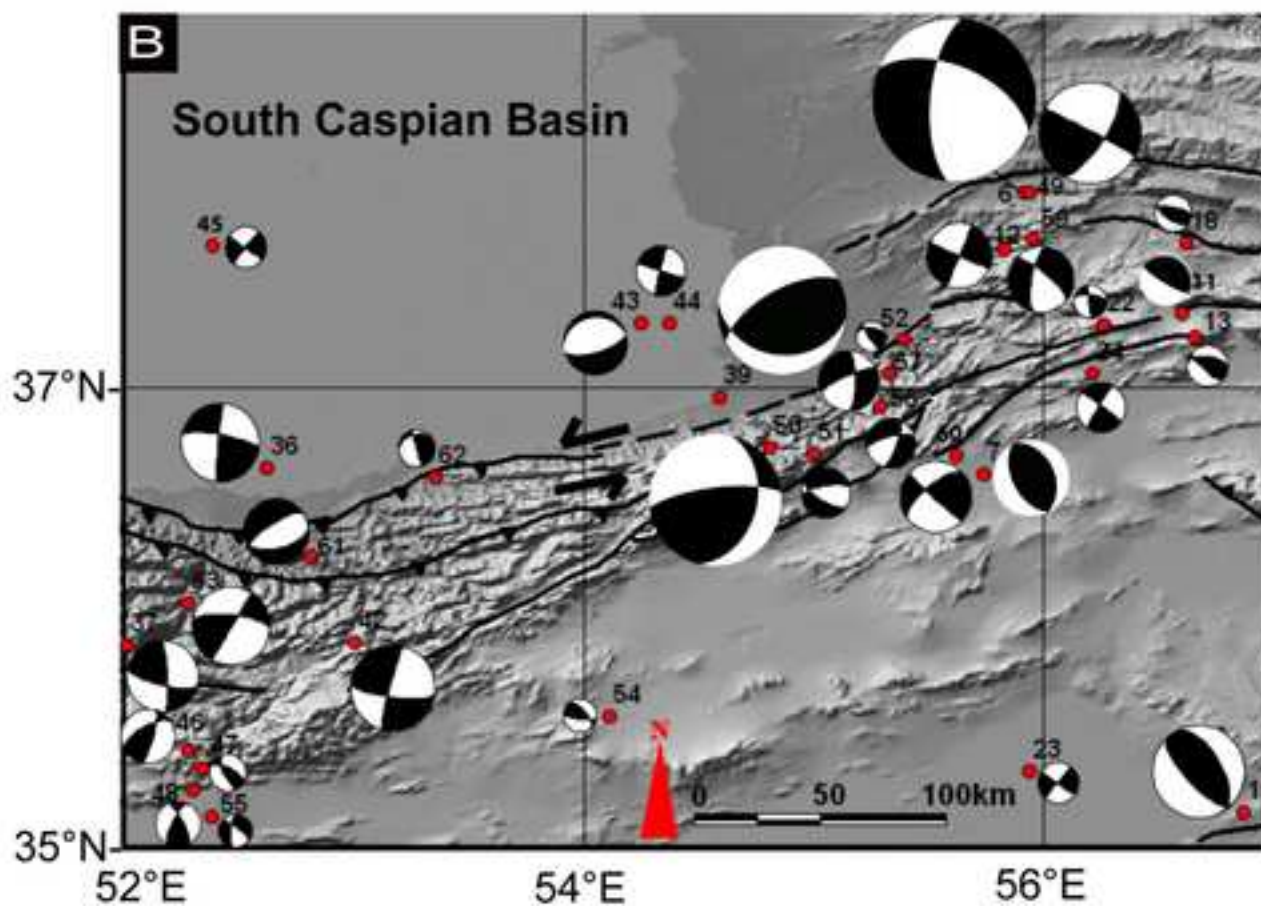
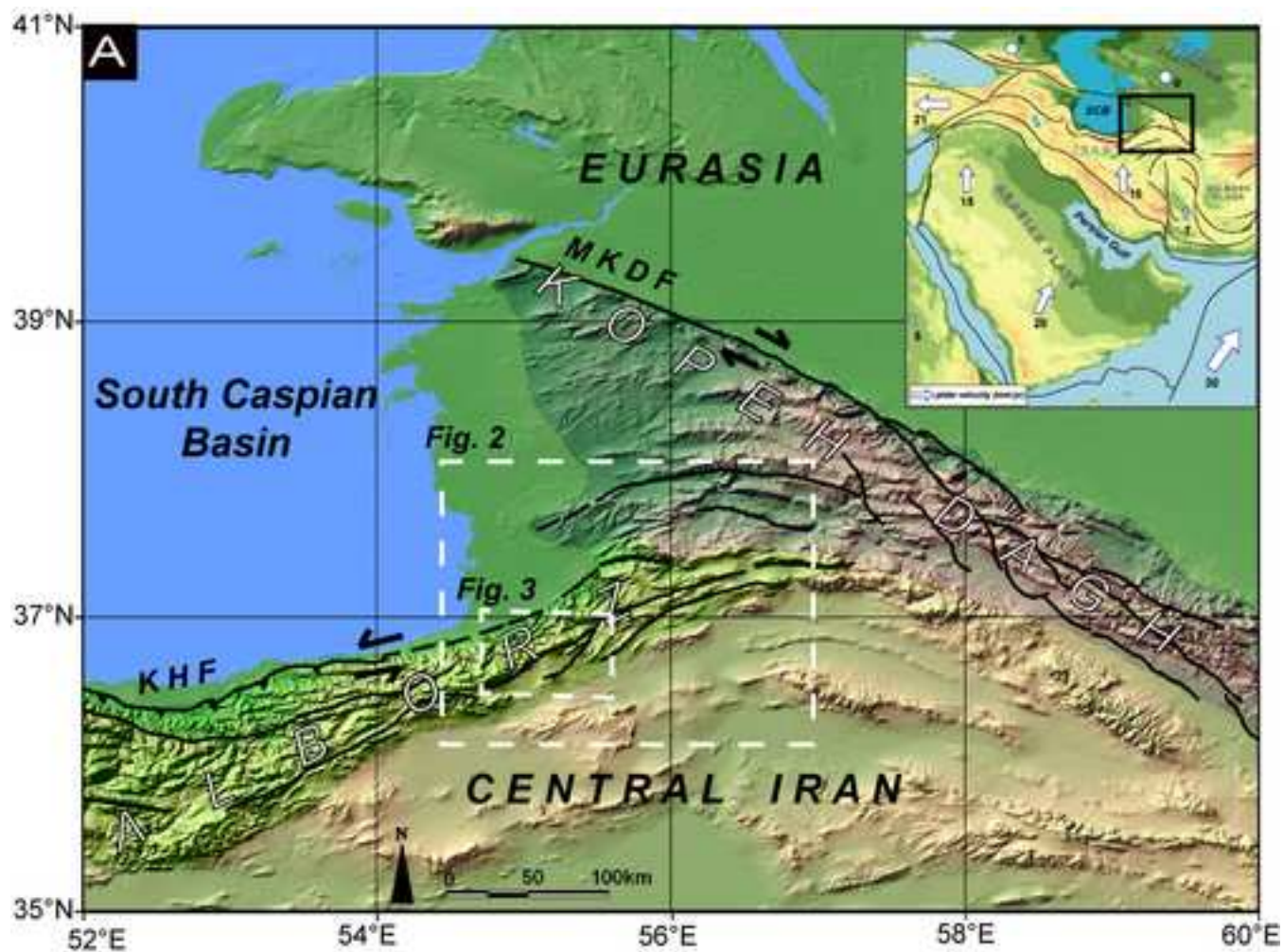


Figure 2

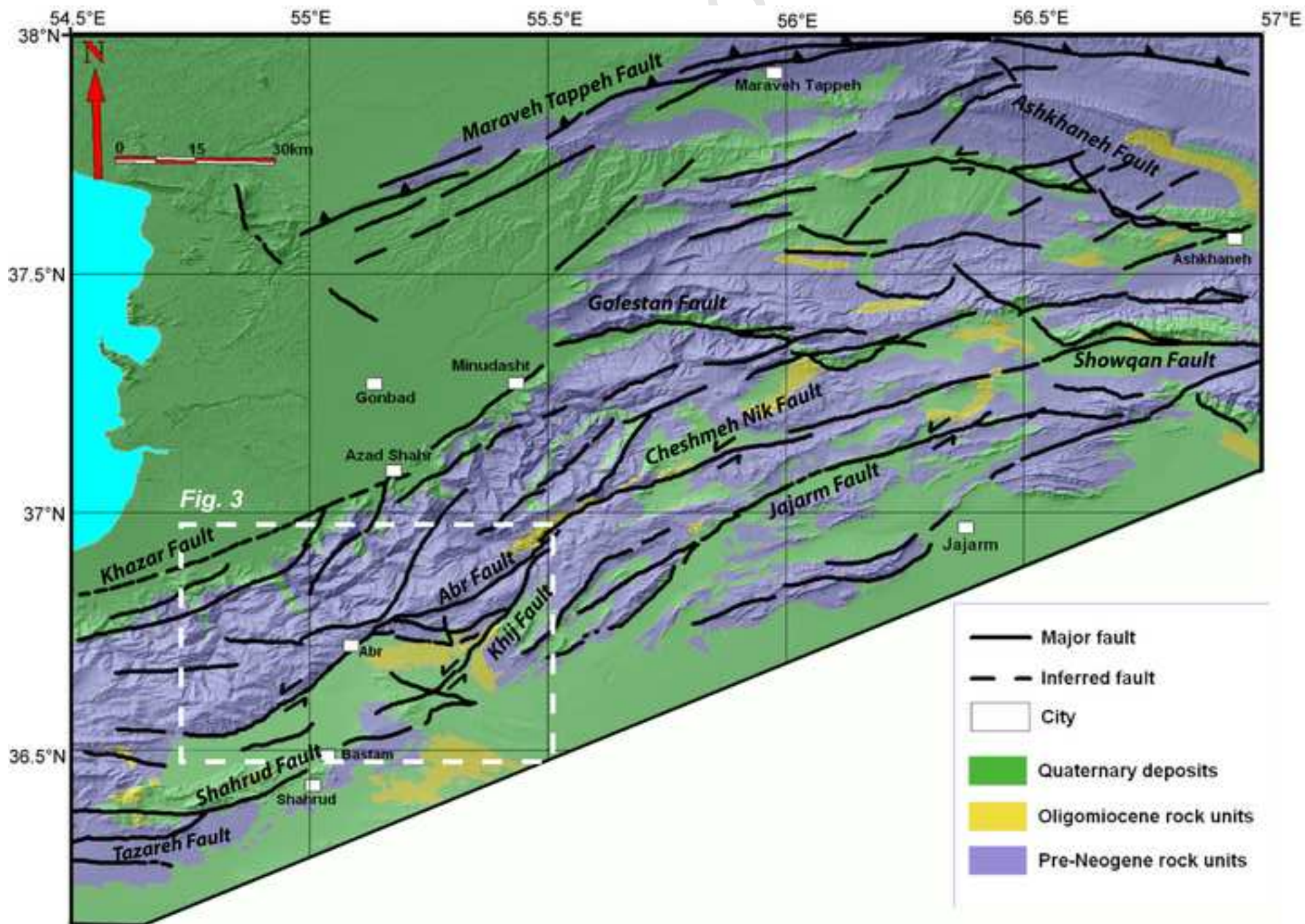
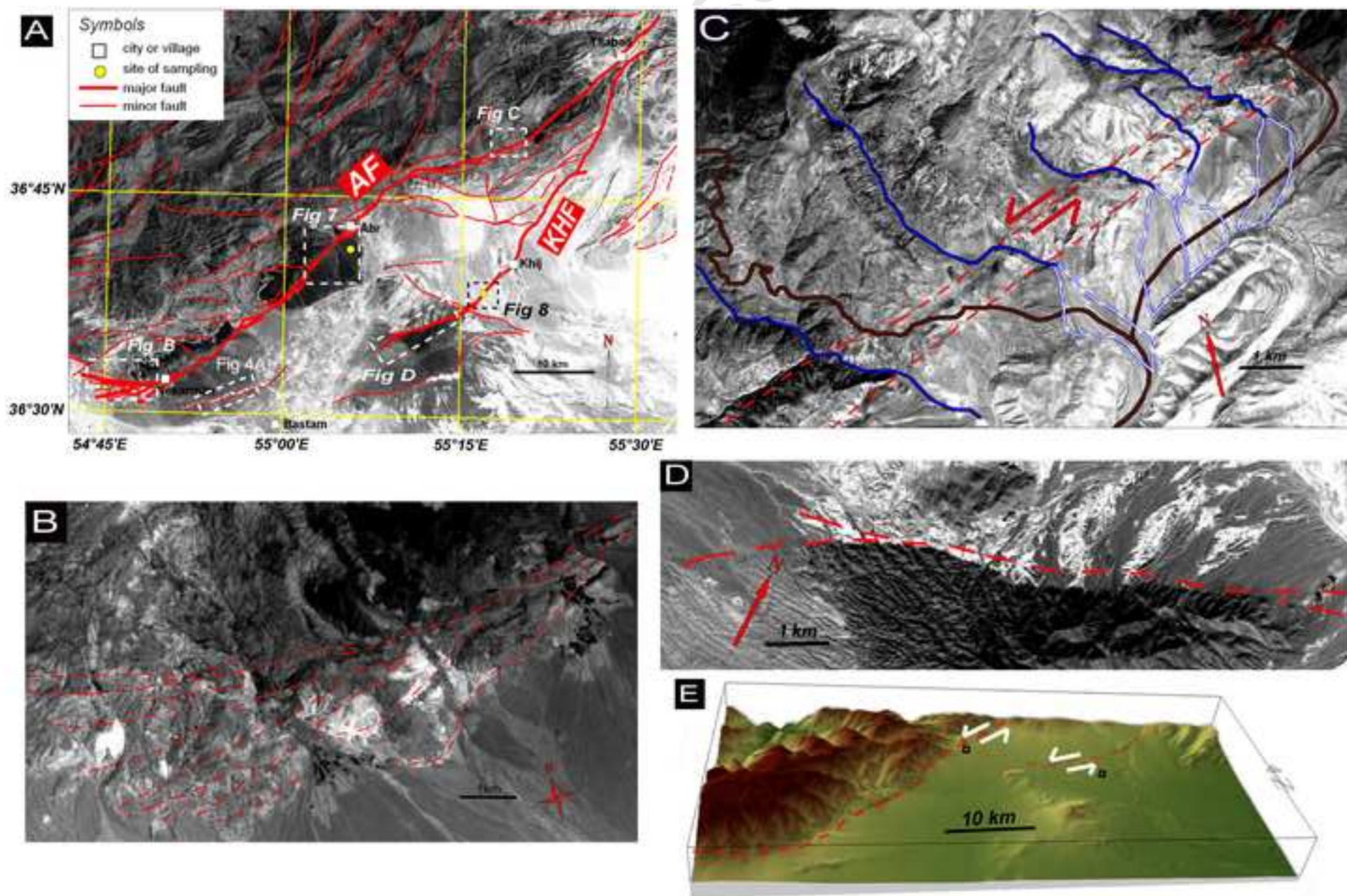


Figure 3



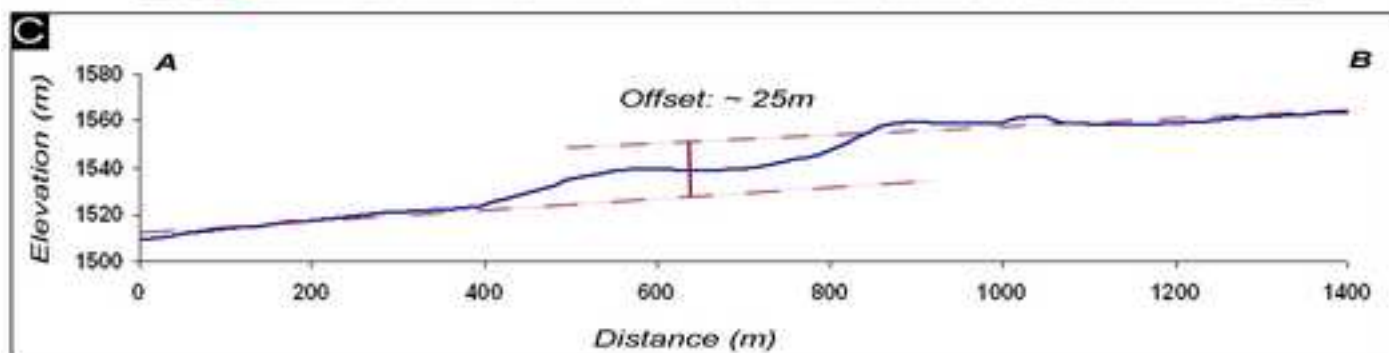
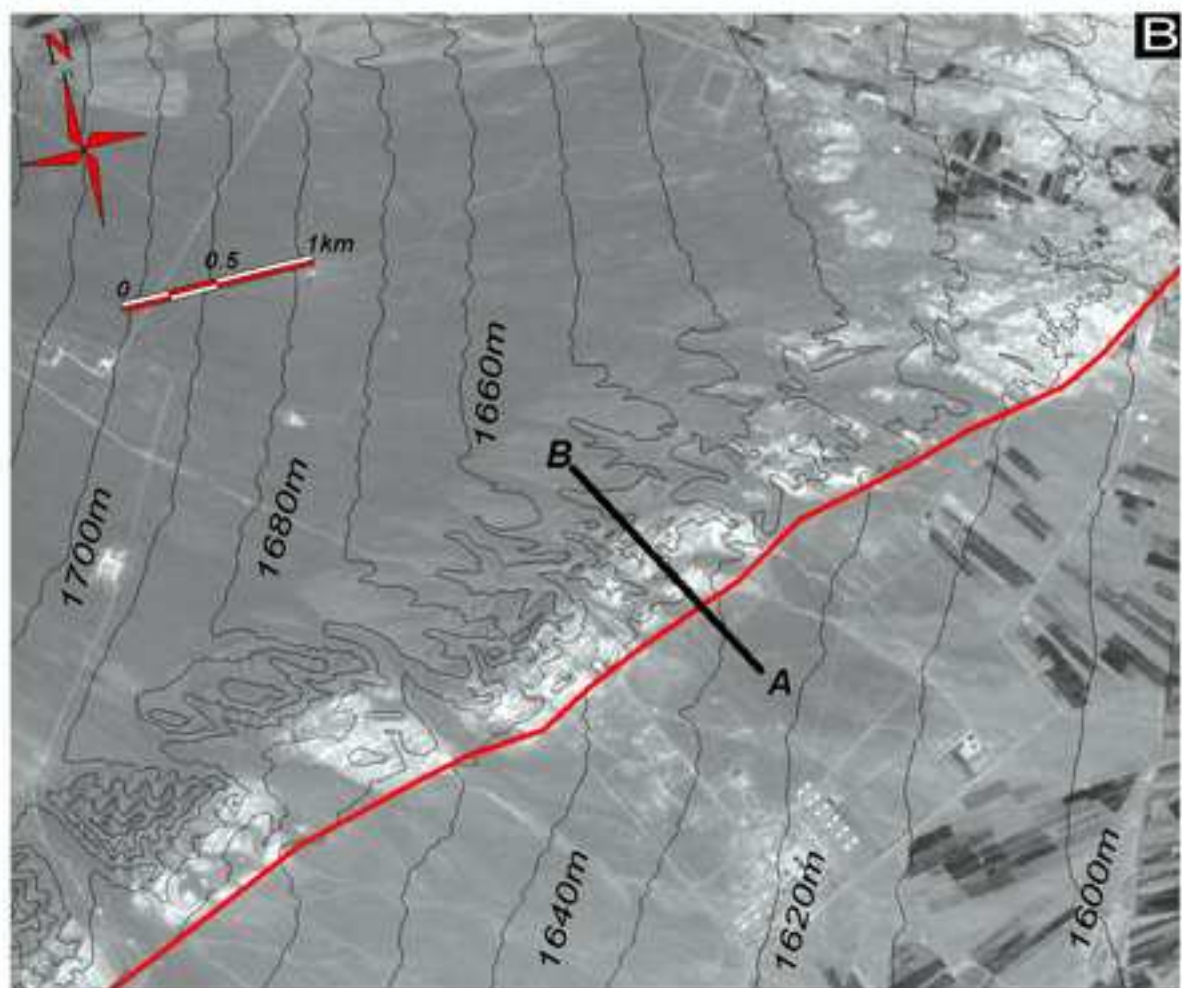
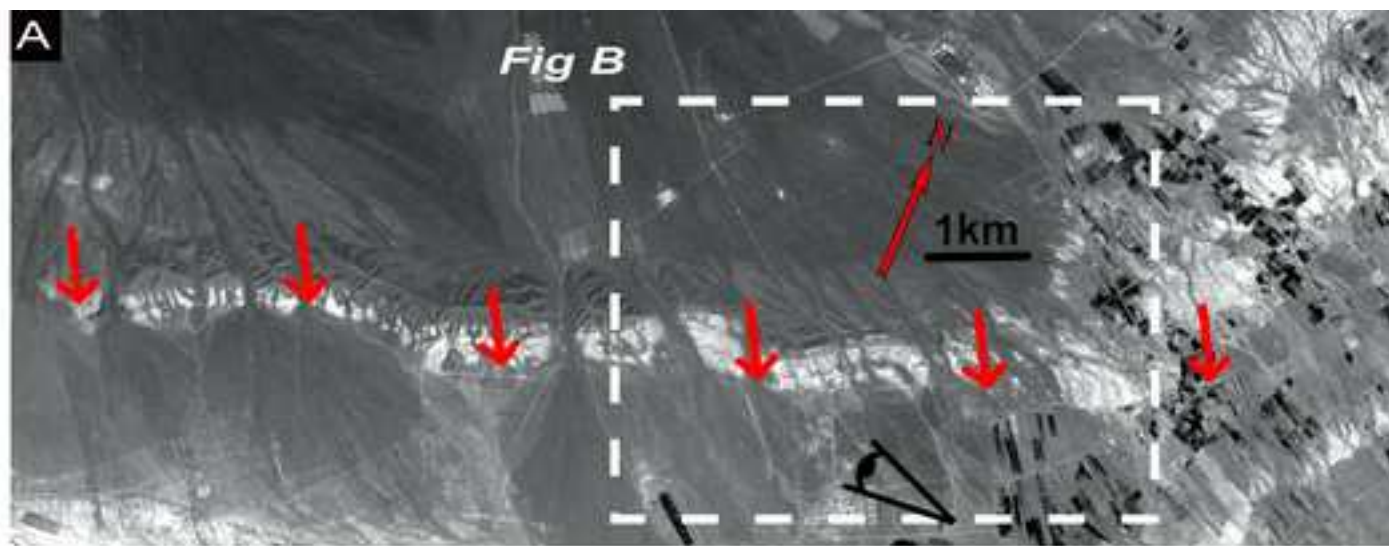
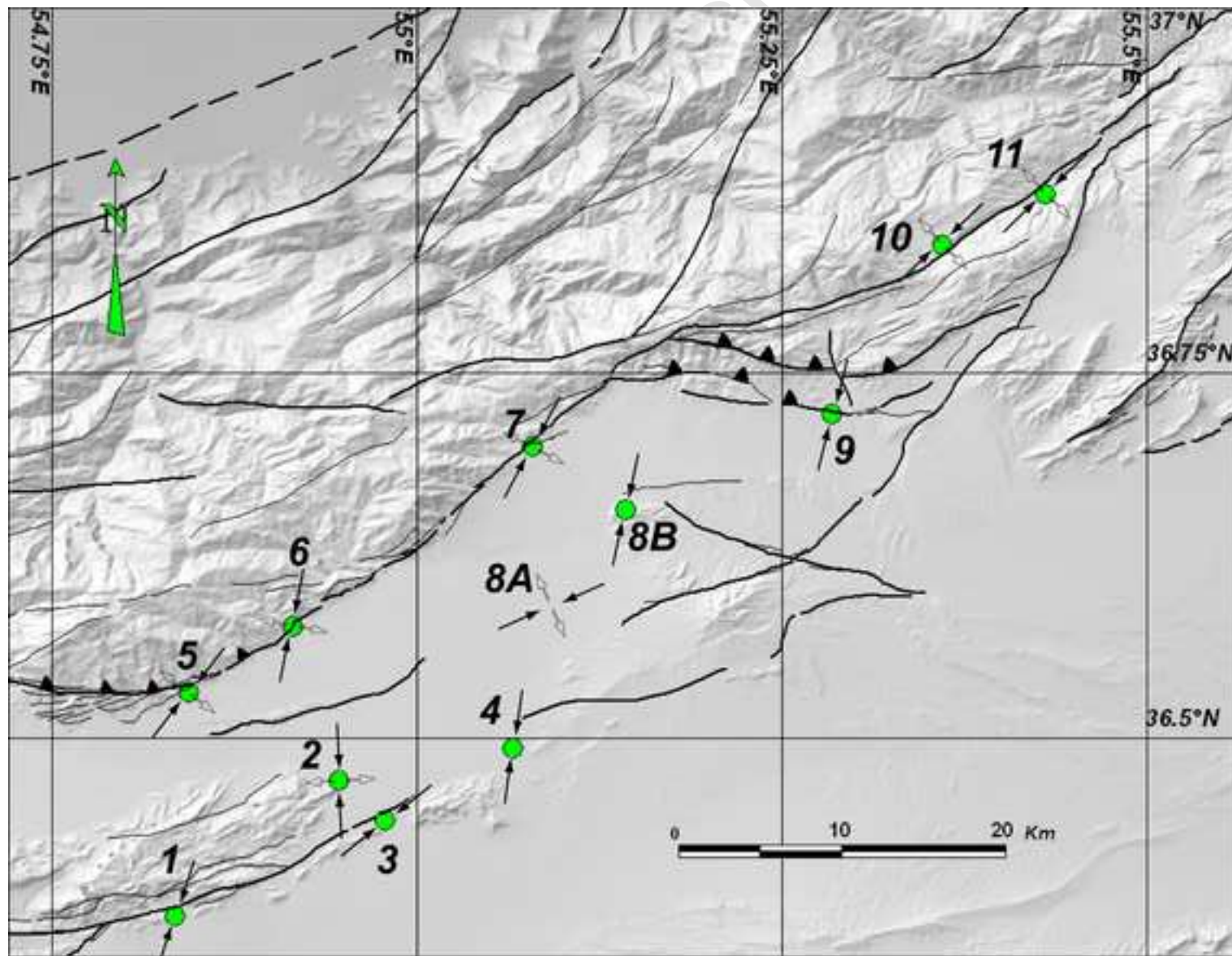


Figure 5



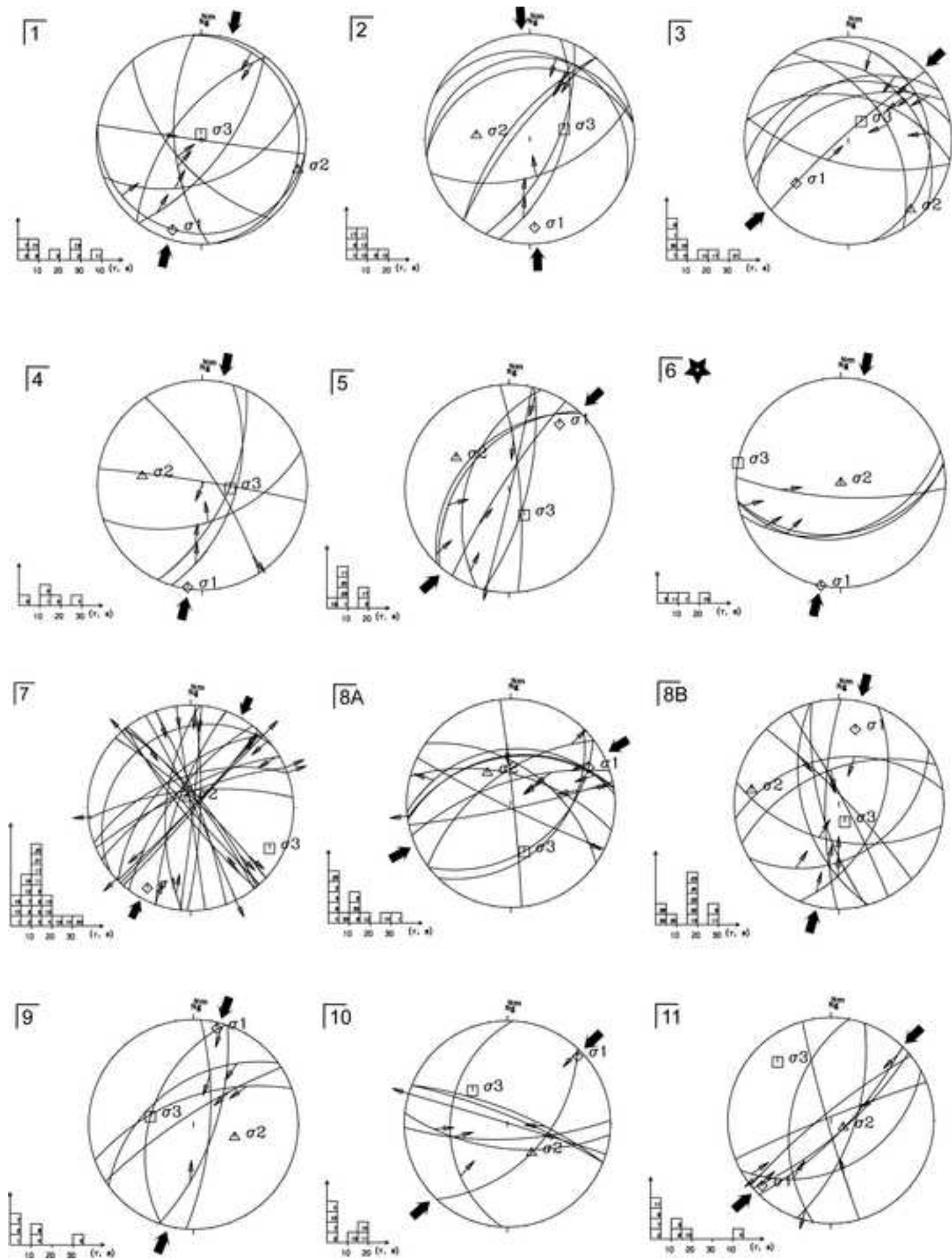
— Major fault

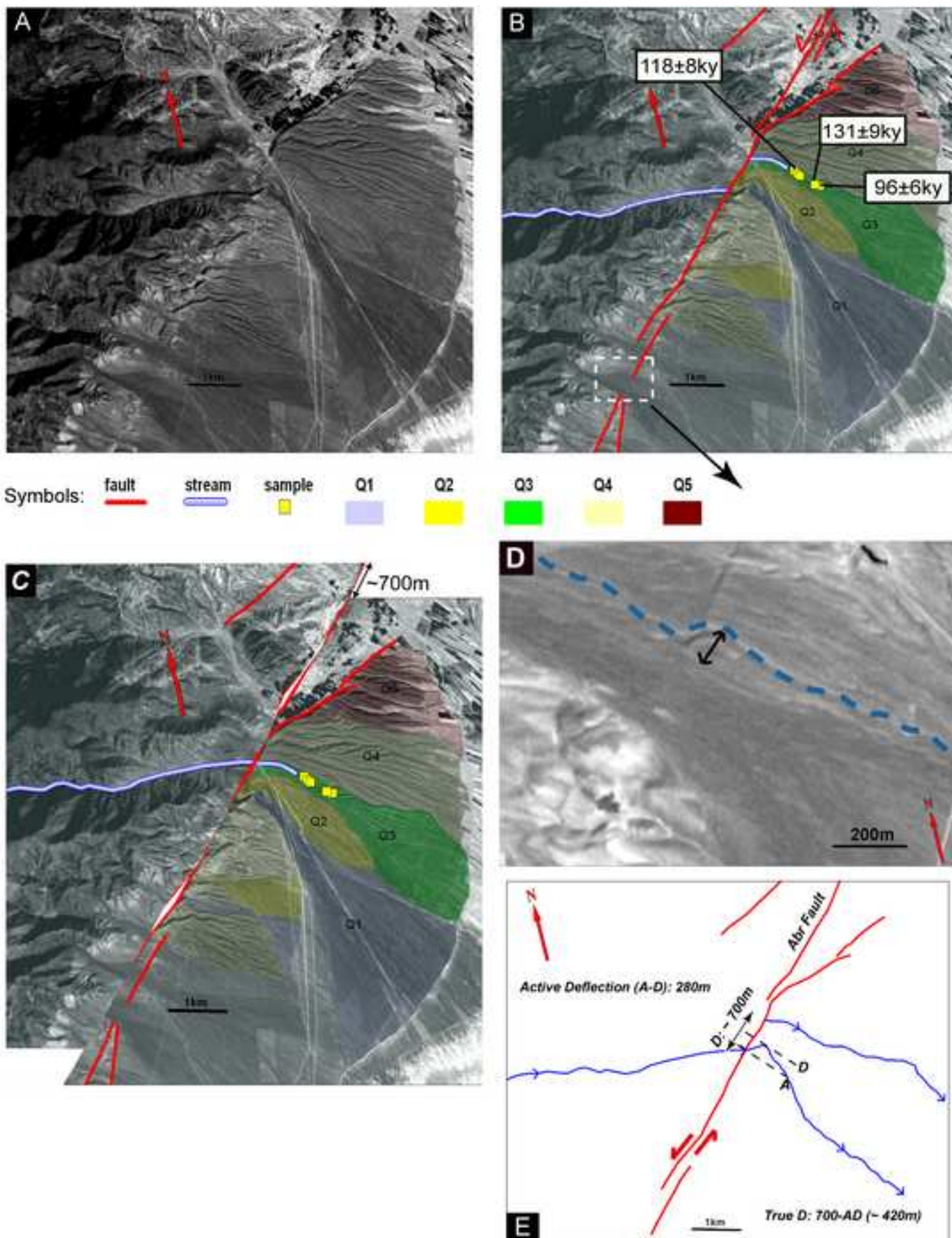
- - - Inferred fault

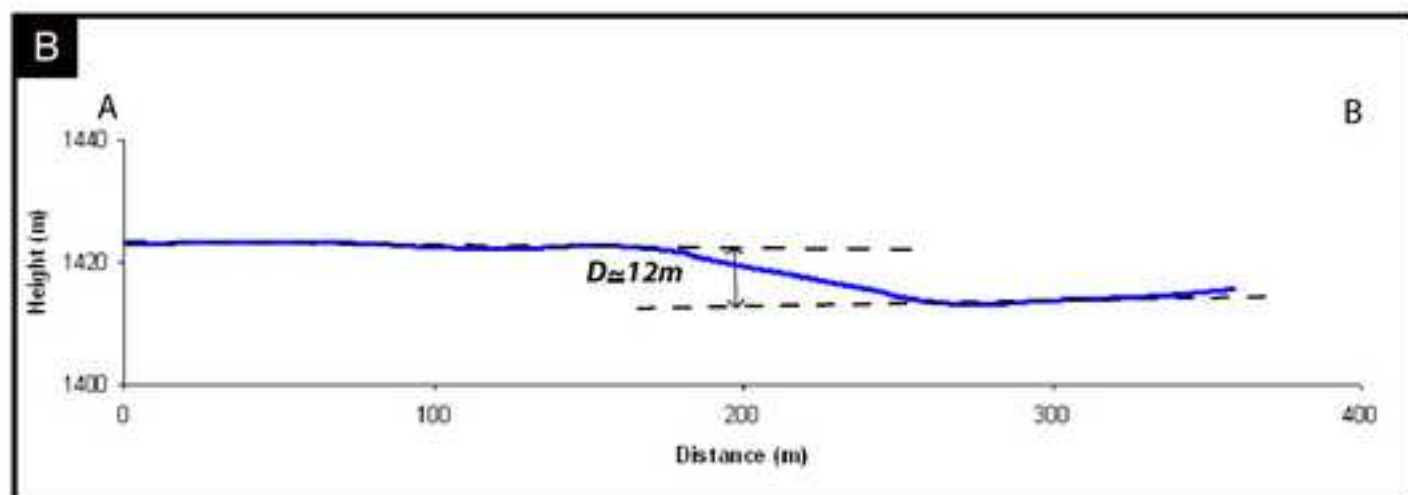
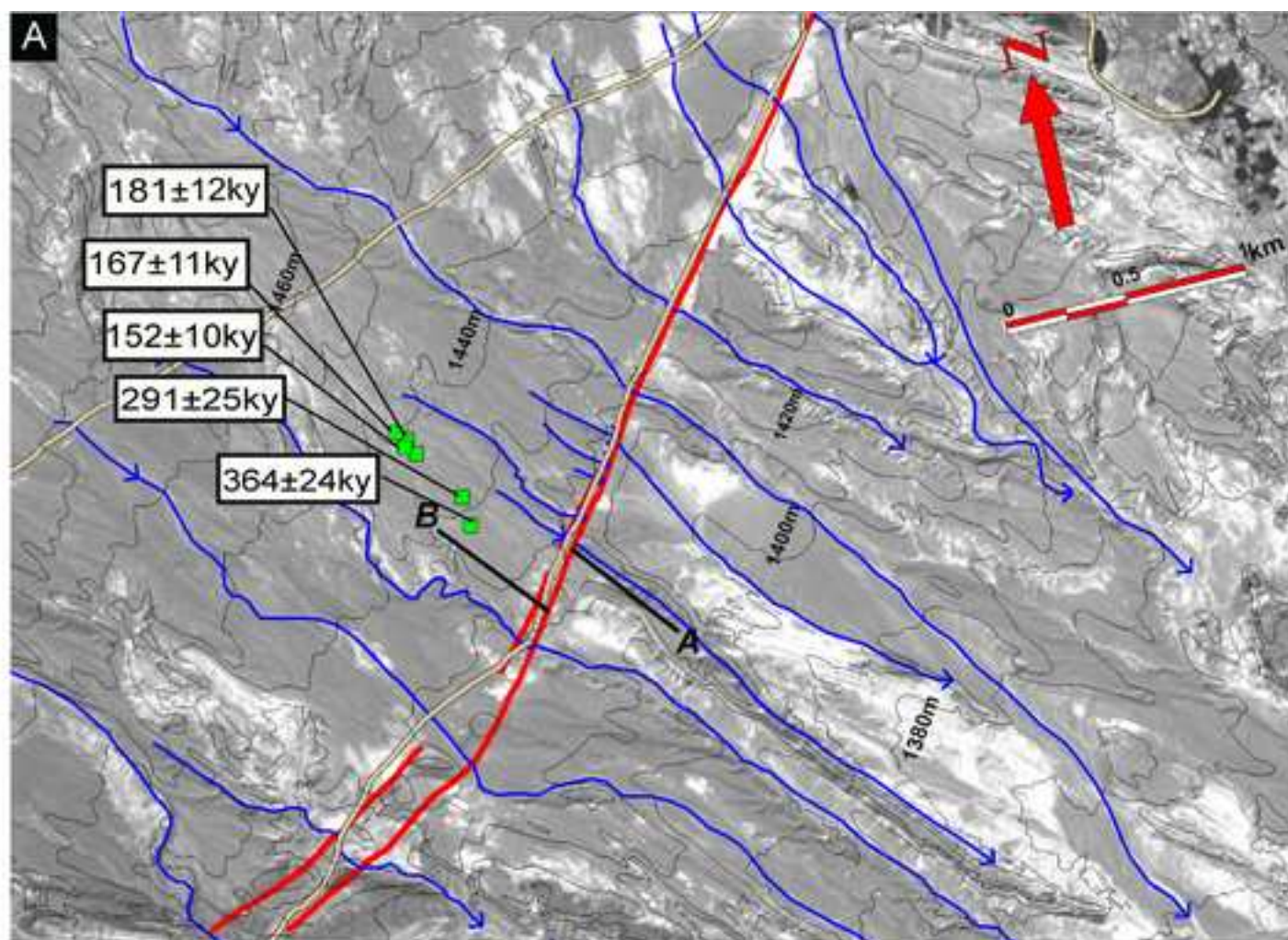
— Minor fault

● Striation measurement site

Stress Regime:
→ ← Compressional
→ ← Strike slip







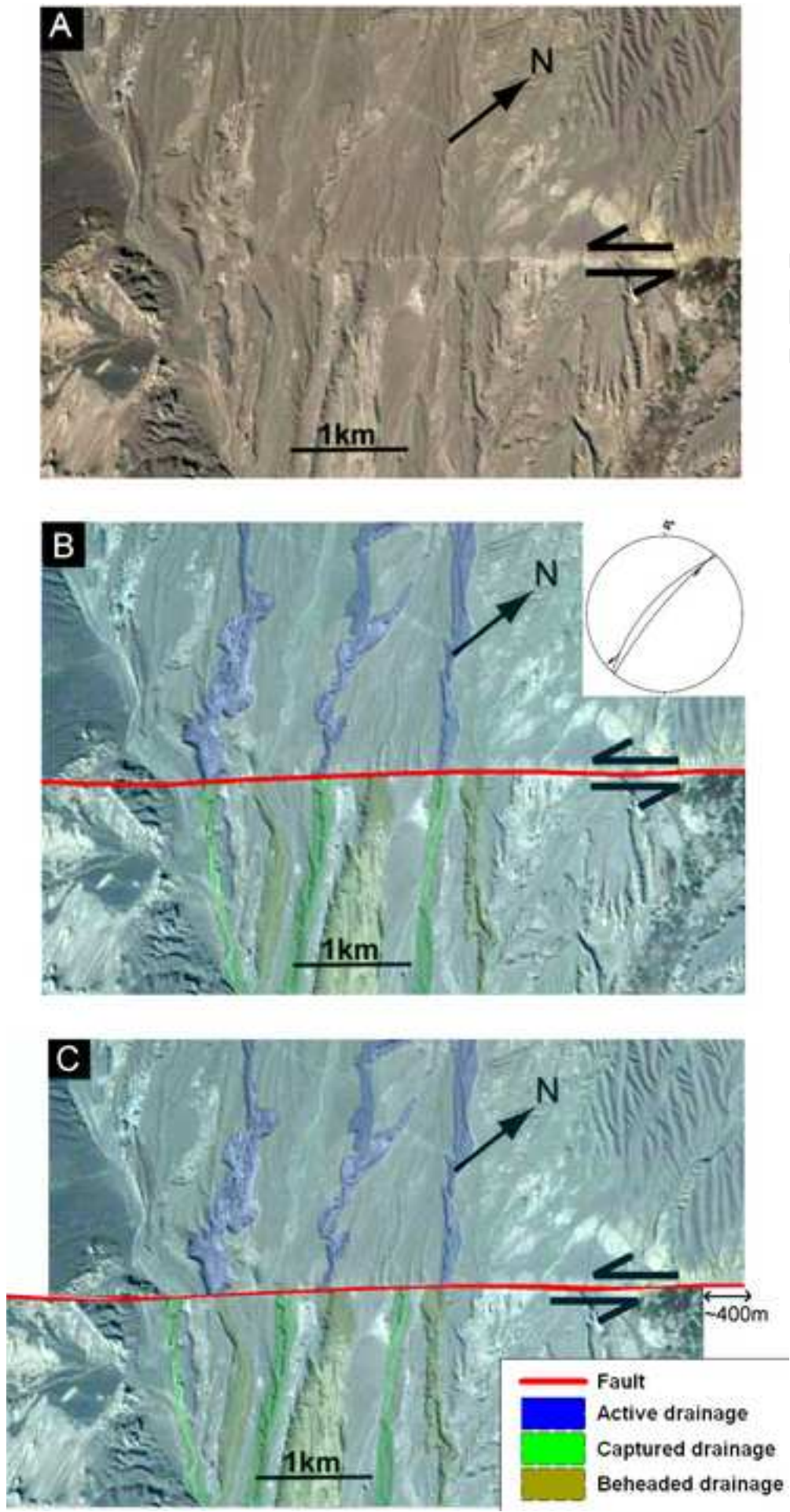


Figure 10



Figure 11

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