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## Validation of IASI FORLI carbon monoxide retrievals using FTIR data from NDACC

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**Abstract.** Carbon monoxide (CO) is retrieved daily and globally from space-borne IASI radiance spectra using the Fast Optimal Retrievals on Layers for IASI (FORLI) software developed at the Université Libre de Bruxelles (ULB). The IASI CO total column product for 2008 from the most recent FORLI retrieval version (20100815) is evaluated using correlative CO profile products retrieved from ground-based solar absorption Fourier transform infrared (FTIR) observations at the following FTIR spectrometer sites from the Network for the Detection of Atmospheric Composition Change (NDACC): Ny-Ålesund, Kiruna, Bremen, Jungfraujoch, Izaña and Wollongong. In order to have good statistics for the comparisons, we included all IASI data from the same day, within a 100 km radius around the ground-based stations. The individual ground-based data were adjusted to the lowest altitude of the co-located IASI CO profiles. To account for the different vertical resolutions and sensitivities of the ground-based and satellite measurements, the averaging kernels associated with the various retrieved products have been used to properly smooth coincident data products. It has been found that the IASI CO total column products compare

well on average with the co-located ground-based FTIR total columns at the selected NDACC sites and that there is no significant bias for the mean values at all stations.

### 1 Introduction

The main sources for CO in the atmosphere are incomplete combustion of biomass and fossil fuel, and oxidation of methane (CH<sub>4</sub>) and non-methane hydrocarbons. Vegetation and ocean contributions are less important. The total production is about 1230–3000 Tg of CO per year. CH<sub>4</sub> oxidation produces 400–800 Tg of CO per year, and oxidation of other hydrocarbons in the troposphere contributes a similar quantity of CO. The main sinks of CO are chemical loss by OH generating carbon dioxide (CO<sub>2</sub>), uptake by soil and removal in the stratosphere which amounts to a total of about 1600–3700 Tg CO yr<sup>-1</sup> (more details can be found in Brasseur and Solomon, 2005). CO has a lifetime of several weeks in the troposphere and is therefore often used as a tracer for pollution (Turquety et al., 2008).

Space-borne missions currently providing CO abundances from nadir infrared radiance measurements are the following: Measurements Of Pollution In The Troposphere (MOPITT) on Terra using a correlation radiometer (Drummond et al., 1995; Deeter et al., 2003; Fortems-Cheiney et al., 2011); Atmospheric Infrared Sounder (AIRS) on Aqua using a spectrometer (Aumann et al., 2003; McMillan et al., 2005, 2011); Tropospheric Emission Spectrometer (TES) on Aura using a Fourier transform spectrometer (FTS) (Beer et al., 2001; Rinsland et al., 2006); and Infrared Atmospheric Sounding Interferometer (IASI) on MetOp-A (Hilton et al., 2012; Clerbaux et al., 2009; George et al., 2009; Maddy et al., 2009). All these instruments measure in a sun-synchronous orbit.

Previous evaluations of IASI FORLI-CO retrievals have been performed using other satellite or aircraft data. George et al. (2009) have provided a first evaluation of IASI FORLI-CO distributions by comparison with MOPITT, AIRS and TES CO retrievals. They found that on a global scale 0.8 to 2.4 independent pieces of information are available in the IASI FORLI-CO retrievals (version 20081030). Comparisons between the CO columns of the different instruments show average discrepancies of about 7 % in the Northern Hemisphere and the equatorial region and of about 11 % south of 15° S. In case of strong CO concentrations (e.g. forest fires), the discrepancies can be as high as 17 % with IASI being larger.

Turquety et al. (2009) evaluate the quality of IASI FORLI-CO retrievals for highly polluted conditions due to wildfires. The coarse vertical resolution of the IASI retrievals does not allow the determination of exact plume heights of forest fires, but the large CO enhancements measured are consistent with respect to transport pathways and plume height, to the maximum aerosol backscatter coefficient measured from space by the CALIPSO lidar (Winkler et al., 2003).

Pommier et al. (2010) described in situ aircraft measurements of CO from the POLARCAT campaign in the framework of the International Polar Year (IPY) which they compared to IASI FORLI-CO. They found that IASI has the capability to detect long-range transport of (CO) plumes and high boundary layer CO signatures from forest fires in Siberia and that IASI CO retrievals are less accurate over sea ice and snow. They also compared IASI CO with aircraft data, completed above the aircraft profile with the CO climatology of the Atmospheric Chemistry Experiment Fourier Transform Spectrometer (ACE-FTS), and found good agreement (maximum differences of 17 % and 20 % for profiles in spring and in summer, respectively; 1 to 5 % and 12 % for total columns in spring and summer, respectively, with IASI being higher). Total column comparisons of IASI CO with these in situ aircraft measurements show correlations between 0.15 and 0.74 in spring and 0.26 and 0.84 in summer.

In this paper we focus on the comparison of IASI FORLI-CO retrievals with ground-based solar absorption FTIR measurements from selected NDACC stations. This is the first

study in which individual IASI CO data are compared to collocated high-quality ground-based data.

In the following Sect. 2 we shortly present the IASI instrument and the FORLI-CO retrievals and in Sect. 3 the ground-based FTIR data are described. Section 4 explains the comparison method, followed by the results of the comparisons in Sect. 5. A description of the error budget for the CO total column differences is given in Sect. 6.

## 2 The IASI instrument and the FORLI-CO retrievals

IASI was launched on 19 October 2006 on the satellite MetOp-A from the European Organisation for the Exploitation of Meteorological Satellites (EUMETSAT). It has a circular field of view (FOV) at nadir of 12 km, and its spectral range is from 645 to 2760  $\text{cm}^{-1}$  (3.62 to 15.5  $\mu\text{m}$ ) with an apodized spectral resolution of 0.5  $\text{cm}^{-1}$  and spectral sampling of 0.25  $\text{cm}^{-1}$  leading to 8461 spectral samples. It has a wide swath width of 2200 km providing global coverage twice a day. For more information about the IASI instrument, refer to Hilton et al. (2012).

The CO IASI data used here were retrieved from the nadir radiance spectra in the spectral range 2143 to 2181.25  $\text{cm}^{-1}$  using the retrieval code FORLI-CO from the Université Libre de Bruxelles (ULB) (Hurtmans et al., 2012). It uses pre-calculated lookup tables (LUTs) to make near real-time processing possible by avoiding line-by-line retrievals. The FORLI LUTs are absorption cross sections at various pressures (on a logarithmic grid), temperatures (on a linear grid) and humidities. For CO, these cross-sections are calculated in the spectral range 2128–2206  $\text{cm}^{-1}$ , larger than the one used in the retrieval, and with an oversampling of 10. More details are given in Hurtmans et al. (2012). FORLI-CO retrievals give CO profiles reported on 19 altitude levels using the optimal estimation method described by Rodgers (2000). The 19 altitude levels correspond to the altitude layers 0 to 1 km, 1 to 2 km, 2 to 3 km, ..., 17 to 18 km and 18 km to the top of the atmosphere. The product consists of a CO concentration profile retrieved on these vertical levels at each measurement point, as well as the associated errors and an averaging kernel matrix. A single a priori profile (in VMR) is used in the retrieval scheme. The temperature and humidity profiles used in FORLI are those from the the EUMETSAT Level 2 processor. The cloud fraction disseminated by EUMETSAT is used, and only those pixels in which the cloud fraction is below 25 % are processed. Hence, each observation has a different and already optimized set of meteorological profiles associated. One should further note that the water amount is systematically adjusted (as a total column), strongly minimizing the impact of the latter on the total retrieval error. For the emissivity, FORLI uses above continental surfaces the IASI climatology (Zhou et al., 2011) whenever possible (more detail can be found in Hurtmans et al., 2012).

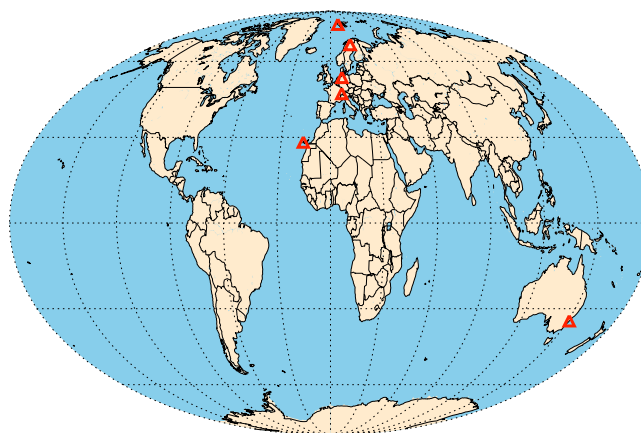
For the comparisons between the IASI FORLI-CO retrievals and selected NDACC FTIR column measurements in the year 2008, version 20100815 of IASI FORLI-CO data products was used. Comparisons with a previous version (20081030 studied in George et al., 2009) have also been performed, but since we find no significant differences between the results of the comparisons with the two FORLI retrieval versions, only the latest version will be presented here. The results are generally also applicable to the older version data.

We use here the IASI FORLI-CO products currently publicly available, which are the total column and the corresponding averaging kernel vector (estimated by adding the averaging kernels matrix rows). These products are available from the Ether website (<http://ether.ipsl.jussieu.fr>) for the whole IASI measurement period. Quality flags allow for filtering of the data. Therefore we are able to use here the most reliable data. In particular, cloud cover has to be less than 12% (cloud fraction from the EUMETSAT level 2 processor is used from a combination of instruments onboard MetOp A, following specific methods, detailed in August et al., 2012) and the averaging kernel vector must not include unrepresentative values (generally too high). This filtering reduces the number of data points used in the analysis by between 11 and 23% depending on the station analysed.

### 3 Ground-based Fourier transform infrared spectrometers

The IASI FORLI-CO data have been compared with total columns retrieved from solar absorption spectra recorded by ground-based FTIR spectrometers from six stations that are part of NDACC (<http://www.ndacc.org>). A list of the stations with their location is presented in Table 1; Fig. 1 shows the distribution of the FTIR sites. The participating stations cover latitudes from 34.5° S to 78.9° N. They provide measurements in the Northern Hemisphere from the subtropics (Izaña) to the polar regions (Ny-Ålesund and Kiruna); Wollongong is the only station in the Southern Hemisphere. These instruments make regular year-round measurements from which the abundances of several tropospheric and stratospheric chemical species, including CO, can be derived. The FTIR measurements require clear-sky daylight conditions, which excludes measurements at Ny-Ålesund and Kiruna during polar winter.

More information about the instruments, the retrieval methodologies and the measurements performed at each site can be found in the references provided in Table 1. All stations use Bruker FTIR spectrometers (Ny-Ålesund and Jungfraujoch: 120HR, Kiruna, Bremen, Izaña and Wollongong: 125HR); for these measurements the instruments use resolutions, i.e. the inverse of the maximum optical path difference ( $1/\text{maxOPD}$ ), given in Table 1, from 0.004 to 0.006  $\text{cm}^{-1}$ . In all cases, the microwindows 2057.70–2058.00  $\text{cm}^{-1}$ , 2069.56–2069.76  $\text{cm}^{-1}$  and



**Fig. 1.** Ground-based FTIR spectrometer locations used in the comparisons. From the north to the south, in red: Ny-Ålesund, Kiruna, Bremen, Jungfraujoch, Izaña, and Wollongong.

2157.50–2159.15  $\text{cm}^{-1}$  were fitted simultaneously to retrieve CO using either the SFIT2 code (Pougatchev and Rinsland, 1995; Pougatchev et al., 1995; Rinsland et al., 1998) or PROFFIT96 (Hase et al., 2004). Both algorithms implement the optimal estimation method (Rodgers, 2000) to retrieve vertical profiles from a statistical weighting between a priori information and information coming from the high-resolution spectral measurements. Spectroscopic line parameters from HITRAN 2004 were adopted for all ground-based retrievals. Other information required for the retrievals such as a priori profiles and covariances, treatment of instrument lineshape, and atmospheric temperature and pressure is optimized for each site as appropriate for the local conditions.

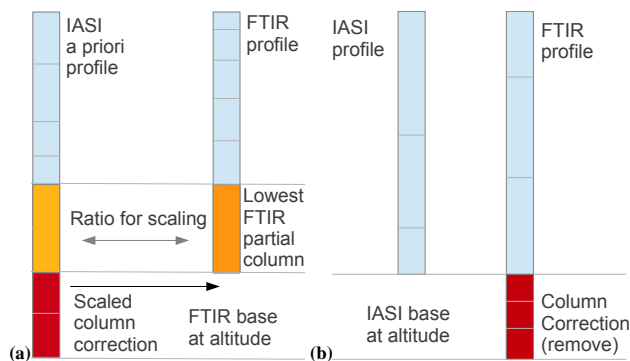
NDACC FTIR CO data have already proven useful in the past for the validation of various satellite data sets, e.g. in Barret et al. (2003, 2005); Sussmann and Buchwitz (2005); Dils et al. (2006); Senten et al. (2008); Clerbaux et al. (2008), and de Laat et al. (2010).

### 4 Comparison methodology

In the work presented here, we use IASI data from 1 January 2008 to 31 December 2008. We perform a careful comparison between IASI and NDACC total columns adjusting for altitude differences and accounting for the respective vertical resolutions of the measurements, within the limits of the co-location criteria and using strict quality control criteria for the filtering of the data. We took the same coincidence criteria for all the stations to define correlative measurements taking into account temporal and spatial variability, and aiming at a statistically relevant number of comparisons. A given FTIR measurement was compared to all IASI measurements that occurred on the same calendar day (UTC) and within 100 km of the ground-based station. No distinction was made between IASI measurement footprints over

**Table 1.** List of the FTIR stations that provided data for the analyses (Sect. 5). The latitude and longitude of each station are provided, together with the altitude above sea level in meters (m a.s.l.). The instrument, the maximum optical path difference (maxOPD) and the retrieval code (+version) used for the CO measurements are indicated for each station in columns 5, 6 and 7. References describing the stations, measurements and analyses are given in column 8.

Station	Location	Coordinates	Alt. [m a.s.l.]	Instrument	maxOPD [cm]	Retrieval Code (version)	Reference
Ny-Ålesund	Svalbard	78.9° N, 11.9° E	20	Bruker 120HR	180	SFIT2 (3.93)	Notholt et al. (1993)
Kiruna	Sweden	67.8° N, 20.4° E	419	Bruker 125HR	180	PROFFIT96	Blumenstock et al. (2006)
Bremen	Germany	53.1° N, 8.8° E	27	Bruker 125HR	180	SFIT2 (3.93)	Velazco et al. (2007)
Jungfrauoch	Switzerland	46.5° N, 8.0° E	3580	Bruker 120HR	175	SFIT2 (3.91)	Mahieu et al. (1997)
Izaña	Canary Islands	28.3° N, 16.5° W	2367	Bruker 125HR	180	PROFFIT96	Schneider et al. (2005)
Wollongong	Australia	34.5° S, 150.9° E	30	Bruker 125HR	257	SFIT2 (3.92)	Paton-Walsh et al. (2005)



**Fig. 2.** Altitude adjustment for the cases where the base of the ground-based FTIR spectrometer measurements is at a higher (a) or at a lower altitude (b) than the IASI ground pixel. The adjusting partial column is indicated in red; in (a) it is added to and in (b) subtracted from the FTIR column. The partial columns to calculate the scaling factor in (a) are indicated in yellow. The calculations were done using a combination of the FTIR and IASI altitude grids.

sea and land nor for IASI day or night measurements. The IASI location that was used in searching for coincidences was the latitude, longitude and time of the centre of the IASI footprint. In cases where several FTIR measurements from a site were co-located with one IASI measurement, they were all considered.

#### 4.1 Altitude adjustment of the FTIR measurements to the IASI ground level altitude

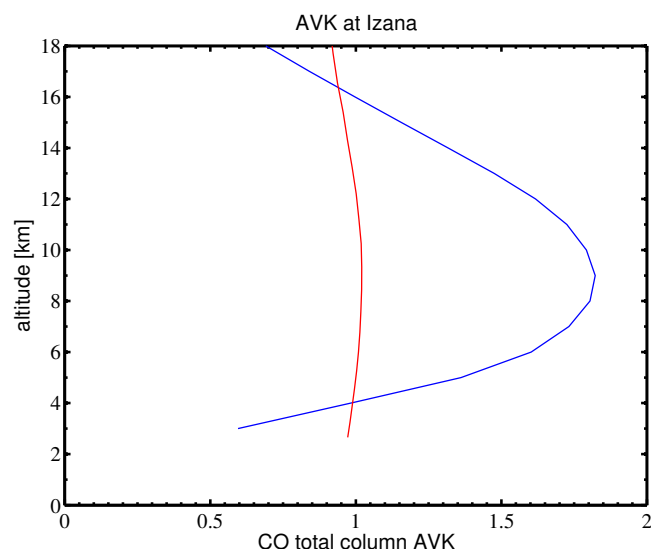
The IASI measurements are scattered within a 100 km radius around the FTIR instrument sites. They may therefore have a different ground level altitude than the ground-based measurements. Since we must compare IASI and FTIR total columns over the same altitude range, and since we want to smooth the FTIR profiles using the IASI total column averaging kernels, we need to adjust the FTIR profile in partial column units such that its total column corresponds to the IASI total columns with different ground level altitudes.

To do this the IASI a priori profiles in partial column units were used to calculate two partial columns (Fig. 2a indicated in red and in yellow): one partial column, the adjusting partial column, in the altitude range between the FTIR and the IASI ground level altitudes (Fig. 2a red) and another one in the altitude range of the lowest FTIR partial column (Fig. 2a yellow). The latter was used to calculate the ratio between the a priori partial columns and the FTIR partial column to scale the adjusting partial column. The scaled adjusting partial column was then added to the FTIR partial column profile when the IASI altitude range is larger than the FTIR altitude range. Because the IASI and FTIR data are on different grids, a combined grid has been used to do these calculations.

On the contrary, if the FTIR profile covers a larger altitude range than the IASI measurement, the FTIR profile is cut to fit the altitude range of IASI (Fig. 2b). If there are  $m$  IASI measurements available for one day, each of the  $n$  corresponding FTIR values was changed  $m$  times to correspond to the ground level altitude of the IASI measurements.

The ground level altitude of the IASI measurements is based on the coordinates of the centre of the IASI footprint. Since the 1-km vertical scale of the FORLI-CO profiles is coarser than the one of the FTIR CO profiles, the GLOBE (1999) digital elevation model from the National Geographical Data Center (National Oceanic and Atmospheric Administration – NOAA) is used to improve the vertical resolution for the altitude adjustment. The IASI ground level altitudes are therefore determined using either the GLOBE model or information gained from the averaging kernel values from the FORLI-CO IASI retrievals (IASI averaging kernel values are not defined at altitudes between mean sea level and the terrain altitude where the IASI footprint is situated).

Based on this altitude information, we can determine a minimum altitude for the IASI total column: if the IASI ground altitude, determined from the GLOBE database, is lower than the altitude determined using the minimum altitude associated with the IASI averaging kernel altitudes, then the IASI averaging kernel altitude is used; otherwise, we use the altitude from the GLOBE database if the difference is less than 1 km (the thickness of a FORLI retrieval layer). For



**Fig. 3.** Mean ground-based FTIR averaging kernels (red) compared to the mean IASI averaging kernel (blue) for Izaña. The averaging kernels are shown for altitudes above the station (2367 m).

differences of more than 1 km, we use the minimum altitude associated with the averaging kernel altitudes to which 999 m are added to reduce the altitude range of the first layer, but still allowing to determine the IASI ground altitude using the averaging kernel altitudes.

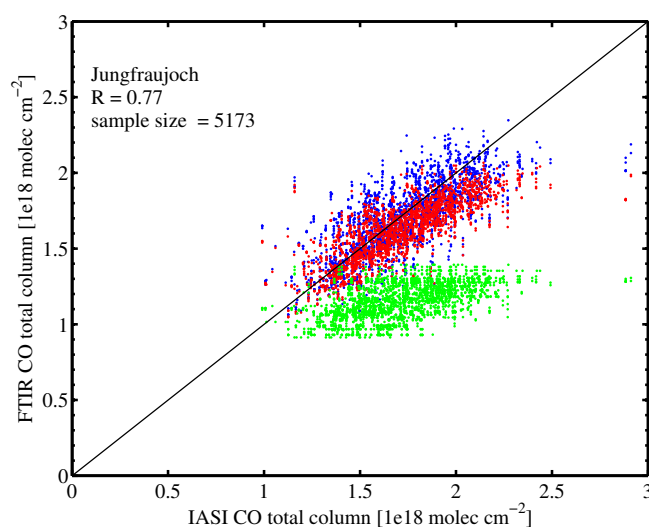
#### 4.2 Smoothing of the FTIR data with the IASI CO averaging kernel

As can be seen in Fig. 3 for the example of Izaña, the IASI mean total column averaging kernel from FORLI-CO has a maximum in the upper troposphere that is larger than one (shown in blue), whereas the one for the FTIR retrievals (shown in red) is constant at about one.

After correction for the altitude differences between the IASI and FTIR measurements, the FTIR profiles in partial column units are interpolated on the altitude grid of the IASI retrievals. Then the altitude-corrected FTIR profiles (in partial column units) are smoothed with the total column averaging kernels of the IASI retrievals according to Eq. (4) of Rodgers and Connor (2003):  $x_s = x_a + A_{IASI}(x - x_a)$ , where  $x_s$  is the smoothed version of  $x$  with  $x$  being the altitude-corrected FTIR profile in partial column units,  $x_a$  is the IASI a priori profile in partial column units and  $A_{IASI}$  is the IASI total column averaging vector.

Figure 4 shows the effectiveness of the smoothing and altitude correction: shown are total column values of uncorrected, unsmoothed columns (green), altitude-corrected columns (red) and altitude-corrected, smoothed columns (blue) for Jungfraujoch.

We now have an averaging kernel smoothed, altitude-adjusted FTIR ground-based CO total column, called



**Fig. 4.** FTIR CO total columns against the IASI CO for the uncorrected, unsmoothed columns (green), the altitude-adjusted (red) and altitude-adjusted, smoothed columns (blue) for Jungfraujoch. The correlation coefficient of the comparison between the altitude-adjusted, smoothed FTIR CO and IASI CO (blue) is shown. For clarity, only every second data point is plotted.

as-IASI-FTIR CO total column in the following, corresponding to the IASI CO total column. Some of the IASI a priori information, however, has been added to the FTIR column by adjusting for the differences between the base altitudes of the FTIR and IASI measurements. Since the IASI averaging kernels are small towards the surface, which implies that the IASI retrieved profile will tend towards the a priori, the inclusion of some IASI a priori information onto the FTIR column has only a minor influence on the FTIR-IASI comparison. The alternative would be to remove the added IASI information from the FTIR column after smoothing, and likewise a corresponding section from the IASI column itself. However, given the limited number of pieces of information in the IASI column, removing such a section would induce errors as well.

#### 4.3 Comparisons

We now have, for each of the  $m$  daily IASI CO measurements,  $n$  as-IASI-FTIR ground-based CO total columns to compare with ( $m$  and  $n$  change for each day).

Comparisons of the total columns are performed for each station, leading to relative differences and standard deviations over the entire time series. Since the different comparison pairs have their total columns calculated over different altitude ranges, we have to compare relative total column differences. The relative values are calculated with respect to the validation instrument (FTIR); e.g. the total mean relative difference was calculated as follows:

**Table 2.** Results of the CO total column comparisons between IASI and the as-IASI-FTIR data. The mean and median differences and the 1- $\sigma$  standard deviation are indicated in columns 2 to 4, the number of individual comparisons, of unique FTIR values and of unique IASI values in columns 5 to 7, the number of days in column 8, the correlations,  $R$ , between the individual column values in column 9 and the maximum temporal coincidence (hours and minutes),  $\max \Delta t$ , in column 10.

FTIR station	Mean diff. [%]	Median diff. [%]	1-Std diff. [%]	# of pairs	# of FTIR	# of IASI	# of days	$R$	Max $\Delta t$
Ny-Ålesund	-1.5	-3.3	15.3	4514	53	17	14	0.60	9 h 21'
Kiruna	1.3	-0.9	19.3	2519	66	34	42	0.60	11 h 43'
Bremen	10.8	9.3	13.7	952	40	14	20	0.32	9 h 16'
Jungfrauoch	2.3	2.0	12.5	5173	309	257	97	0.77	15 h 43'
Izaña	-2.4	-2.4	9.9	2466	109	81	67	0.73	13 h 39'
Wollongong	-4.5	-5.3	12.9	13 402	611	244	77	0.77	16 h 27'

$$\Delta_{\text{rel}} = 100\% \times \frac{1}{N} \sum_{i=1}^N \frac{\text{IASI}_i - \text{FTIR}_i}{\text{FTIR}_i}, \quad (1)$$

where  $N$  is the total number of comparison pairs over all days,  $\text{IASI}_i$  is the IASI total column for the  $i$ -th coincident pair, and  $\text{FTIR}_i$  is the corresponding total column for the FTIR instrument.

The relative differences and standard deviations can be interpreted taking into account the errors on these differences, due to the errors on the individual FTIR and IASI data. This calculation has been performed for one representative station only, as described in Sect. 6. The results of the FTIR total column comparisons are shown in Sect. 5.

## 5 Results for the ground-based inter-comparisons of CO from FTIR spectrometers

The comparisons of the as-IASI-FTIR columns with the IASI data are presented in Table 2. Shown are mean differences, median differences, the 1- $\sigma$  standard deviation of the mean, the number of comparison pairs, the number of comparison days, and the correlation coefficients of the IASI total columns and the corresponding as-IASI-FTIR total columns.

It can be seen that the standard deviations are generally larger than the mean or median differences, indicating that the differences are not significant compared to the variability.

Bremen has large CO total columns as can be seen from Table 3. It also shows the largest IASI-FTIR differences, the reason of which is not clear to us. We verified that it is not due to the co-location criteria chosen. A reduction of the spatial coincidence criterion to 50 km slightly improves the differences for Ny-Ålesund, Jungfrauoch and Izaña, but the variability increases. The differences remain the same for Kiruna, Bremen and Wollongong. Analyses using tighter temporal coincidences did not show a general reduction of the differences for Ny-Ålesund, Bremen and Jungfrauoch. Kiruna, Izaña and Wollongong show smaller differences when 6 h are used for the coincidence criterion;

**Table 3.** Median CO vertical column densities (VCD) for IASI and the as-IASI-FTIR data.

FTIR station	Median CO VCD IASI [ $\times 10^{18}$ molec $\text{cm}^{-2}$ ]	Median CO VCD as-IASI-FTIR [ $\times 10^{18}$ molec $\text{cm}^{-2}$ ]
Ny-Ålesund	1.99	2.10
Kiruna	1.74	1.75
Bremen	2.47	2.27
Jungfrauoch	1.55	1.53
Izaña	1.78	1.86
Wollongong	1.39	1.51

for even smaller temporal coincidences (3 h), the mean differences (and the variabilities) become larger. For a 12 h coincidence criterion, there is no change for any of the stations, because it roughly coincides with our chosen criterion.

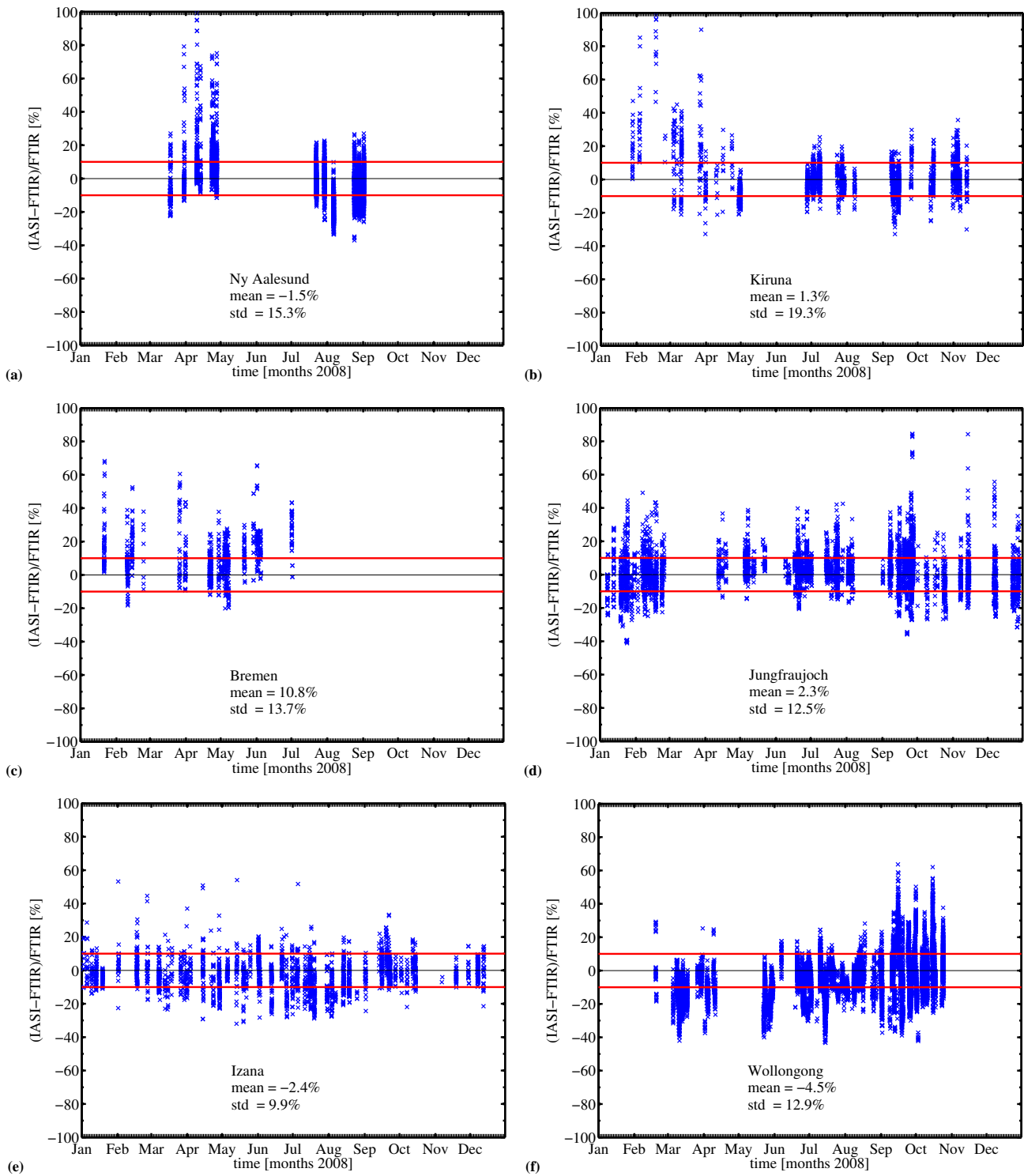
Figure 5 shows the relative differences of the as-IASI-FTIR CO total column and the IASI CO total column for all the comparison pairs for all stations. It can be seen that there is large intra-day variability.

## 6 Representative error estimates calculated for Kiruna

Error budget calculations for the comparisons are performed following Rodgers and Connor (2003). Assuming that the ground-based averaging kernels are almost ideal (i.e. they are about one at all altitudes), Eq. (27) from Rodgers and Connor (2003) can be reduced to

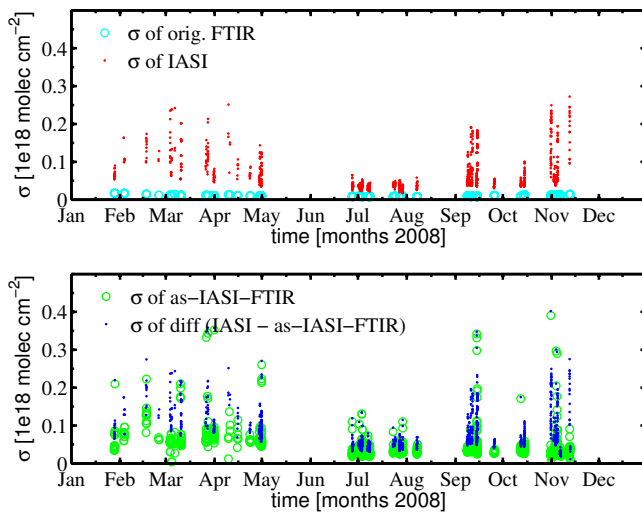
$$\sigma^2 = \sigma_{\text{IASI}}^2 + \mathbf{A}_{\text{IASI}}^T \mathbf{S}_{\text{FTIR}}^{\text{pc}} \mathbf{A}_{\text{IASI}} \quad (2)$$

where  $\mathbf{S}_{\text{FTIR}}^{\text{pc}}$  is the FTIR error covariance matrix in partial column units,  $\mathbf{A}_{\text{IASI}}$  is the IASI total column averaging kernel given in partial columns units,  $\sigma$  is the random error of the differences between the IASI total columns and the as-IASI-FTIR total columns,  $\sigma_{\text{IASI}}$  is the random error on the IASI total column and  $\mathbf{A}_{\text{IASI}}^T \mathbf{S}_{\text{FTIR}}^{\text{pc}} \mathbf{A}_{\text{IASI}} = \sigma_{\text{FTIR}}^2$  is the square of the random error of the as-IASI-FTIR total column.



**Fig. 5.** Relative differences of the smoothed and altitude-corrected FTIR CO total column and the IASI CO total column of all the comparison pairs for (a) Ny-Ålesund, (b) Kiruna, (c) Bremen, (d) Jungfraujoch, (e) Izaña and (f) Wollongong. The red lines indicate  $\pm 10\%$  difference.



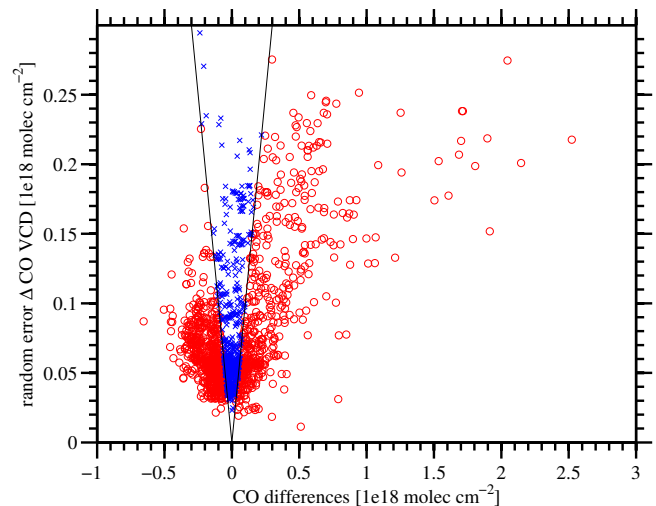


**Fig. 6.** Carbon monoxide random errors,  $\sigma$ , for the quality-filtered data against time at Kiruna. Shown are the errors for the original FTIR total columns (cyan open circles), the IASI total columns (red dots), the as-IASI-FTIR total columns (green open circles) and the random errors of the differences between IASI and as-IASI-FTIR total columns (blue dots).

It can be seen that the altitude range of the FTIR error covariance matrix,  $\mathbf{S}_{\text{FTIR}}$ , has to be adjusted to the altitude range of the corresponding IASI averaging kernels, similarly to the altitude correction of the profile itself. So either we have to cut off matrix entries corresponding to the altitude range that was removed, or we have to add entries corresponding to the extended altitude range. In the latter case, we added a diagonal matrix element which was determined from the difference between the FTIR partial column at the lowest FTIR layer and the IASI a priori partial column corresponding to the FTIR layer. Also, the error covariance matrix for the ground-based data,  $\mathbf{S}_{\text{FTIR}}$ , has to be regridded on the IASI total column averaging kernel grid,  $\mathbf{A}_{\text{IASI}}$ , using the method described in Calisesi et al. (2005) and Vigouroux et al. (2009). With this information the random errors can be calculated.

Figure 6 shows these random errors for the quality-filtered data. The average random errors of the CO columns are about  $0.01 \times 10^{18} \text{ molec cm}^{-2}$  (cyan open circles) and  $0.06 \times 10^{18} \text{ molec cm}^{-2}$  (red dots), for the FTIR and IASI data, respectively. The FTIR random error includes the random contributions from the measurement error, the error of the model parameters and the temperature error. The smoothing error is excluded. The random errors of the FTIR data after altitude adjustment and smoothing using the above formalism are of order  $0.05 \times 10^{18} \text{ molec cm}^{-2}$  (green open circles), and the random error in the differences is  $0.08 \times 10^{18} \text{ molec cm}^{-2}$  (blue dots).

These error calculations were performed for the comparison site Kiruna. The resulting random errors on the



**Fig. 7.** Carbon monoxide random error of the differences against differences of the IASI-FTIR total column for the quality-filtered data for Kiruna. For the values indicated in blue (red), the absolute values of the differences are smaller (larger) than the random errors (VCD = vertical column density).

differences have been compared to the standard deviations of the differences and to the differences themselves. The mean value of the random errors on the differences for Kiruna is 5.1 % (median 3.5 %). It can be seen also that, on average, this error of the differences is smaller than the standard deviation of the differences for Kiruna shown in Table 2. So even if there is no significant bias on average between the IASI data and the ground-based data, there are many cases where the differences between individual IASI total columns and the corresponding FTIR columns are significant.

This is further illustrated in Fig. 7, in which we plot individual values of the random error on the differences versus the differences themselves, for the quality-filtered data at Kiruna. The red dots correspond to the cases where the absolute value of the difference between the IASI and FTIR ground-based CO columns is larger than the error associated with it, in other words, where the difference is significant. The blue dots are the cases where the opposite is true, in other words, where the difference is not significant. This evaluation indicates that the differences are significant in 69 % of the cases.

One may wonder whether this large number is due to the choice of the co-location criteria. If we impose more stringent co-location criteria, e.g. by reducing the co-location distance from 100 km to 50 km, we find that the percentage of cases in which the differences are significant does not change – if instead 1 h is chosen for the temporal coincidence criterion, it increases slightly to 71 %.

In conclusion, we can say that there is an important number of cases for which the differences between the IASI and the FTIR ground-based CO columns are significant,

independently of the co-location criteria; the reported IASI error may therefore be underestimated. Nevertheless, these differences are random and there is no bias on average between the IASI and FTIR ground-based data.

For the individual comparisons one should not forget that we are looking at two different sounding geometries (IASI has a large range of viewing geometries; FTIR instruments follow the sun vector for their air mass) and, besides co-location, IASI has large pixel sizes and the two instruments are thus unlikely to probe exactly the same air mass (within a given pixel IASI even sees an average of concentrations). Therefore, the differences of individual comparisons may still be partly attributed to real CO variability.

## 7 Summary and conclusions

This paper presents first validation results of IASI FORLI-CO total columns using CO retrievals from six ground-based FTIR NDACC sites for measurements from January to December 2008. These stations are Ny-Ålesund, Kiruna, Bremen, Jungfraujoch, Izaña and Wollongong; they span a latitude band from 79° N to 34° S.

Compared to the FTIR CO total columns, the IASI total columns generally show an equal distribution of under- and overestimation. The mean differences, which vary between -4.5 % for Wollongong and 10.8 % for Bremen, have been shown not to be significant since the standard deviations of the differences are greater than the differences themselves. No significant impact of the co-location criteria (50 km to 100 km; 1 h to 24 h) on the differences could be observed.

For Kiruna we have performed a complete error analysis which shows that individual differences between IASI and ground-based FTIR total columns of CO are significant in 69 % of the cases studied. From Eq. (2) the possibility exists that either the errors of the FTIR validation data set or the IASI random errors are underestimated. Assuming that the NDACC FTIR data and associated error budgets are a reliable reference, and that the co-location criteria have been chosen appropriately, we conclude that on individual IASI retrievals the reported IASI error may be underestimated. It has to be kept in mind that IASI might not probe exactly the same air mass as the FTIR instruments.

On average we find that there is no bias between the IASI data and the ground-based NDACC data for the CO total columns above the 6 stations that have been investigated in this study. Further work, and in particular detailed comparisons of the FORLI-CO retrieved profiles along with the NDACC profiles, should help to identify the source of discrepancies observed on individual measurements, such as for instance specific contaminations due to fire or pollution plumes.

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