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Iron (III)-silica interactions in aqueous solution: Insights from X-ray absorption fine structure spectroscopy

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Abstract—The influence of aqueous silica on the hydrolysis of iron(III) nitrate and chloride salts in dilute aqueous solutions ($m_{\text{Fe}} \sim 0.01$ mol/kg) was studied at ambient temperature using X-ray absorption fine structure (XAFS) spectroscopy at the Fe K-edge. Results show that in Si-free iron nitrate and chloride solutions at acid pH (pH < 2.5), Fe is hexa-coordinated with 6 oxygens of H₂O- and/or OH-groups in the first coordination sphere of the metal, at an Fe-O distance of 2.00 ± 0.01 Å. With increasing pH ($2.7 < \text{pH} < 13$), these groups are rapidly replaced by bridging hydroxyls (-OH-) or oxygens (-O-), and polymerized Fe hydroxide complexes form via Fe-(O/OH)-Fe bonds. In these polymers, the first atomic shell of iron represents a distorted octahedron with six O/OH groups and Fe-O distances ranging from 1.92 to 2.07 Å. The Fe octahedra are linked together by their edges (Fe-Fe distance 2.92–3.12 Å) and corners (Fe-Fe distance $\sim 3.47 \pm 0.03$ Å). The Fe-Fe coordination numbers ($N_{\text{edge}} = 1-2$; $N_{\text{corner}} = 0.5-0.7$) are consistent with the dominant presence of iron dimers, trimers and tetramers at pH 2.5 to 2.9, and of higher-polymerized species at pH > 3.

At pH > 2.5 in the presence of aqueous silica, important changes in Fe(III) hydrolysis are detected. In 0.05-*m* Si solutions (pH $\sim 2.7-3.0$), the corner linkages between Fe octahedra in the polymeric complexes disappear, and the Fe-Fe distances corresponding to the edge linkages slightly increase (Fe-Fe_{edge} $\sim 3.12-3.14$ Å). The presence of 1 to 2 silicons at 3.18 ± 0.03 Å is detected in the second atomic shell around iron. At basic pH (~ 12.7), similar structural changes are observed for the iron second shell. The Fe-Si and Fe-Fe distances and coordination numbers derived in this study are consistent with (1) Fe-Si complex stoichiometries Fe₂Si₁₋₂ and Fe₃Si₂₋₃ at pH < 3; (2) structures composed of Fe-Fe dimers and trimers sharing one or two edges of FeO₆-octahedra; and (3) silicon tetrahedra linked to two neighboring Fe octahedra via corners. At higher Si concentration (0.16 *m*, polymerized silica solution) and pH ~ 3 , the signal of the Fe second shell vanishes indicating the destruction of the Fe-Fe bonds and the formation of different Fe-Si linkages. Moreover, ~ 20 mol.% of Fe is found to be tetrahedrally coordinated with oxygens in the first coordination shell ($R_{\text{Fe-O}} = 1.84$ Å). This new finding implies that Fe may partially substitute for Si in the tetrahedral network of the silica polymers in Si-rich solutions.

The results of this study demonstrate that aqueous silica can significantly inhibit iron polymerization and solid-phase formation, and thus increase the stability and mobility of Fe(III) in natural waters. The silica “poisoning” of the free corner sites of iron-hydroxide colloids should reduce the adsorption and incorporation of trace elements by these colloids in Si-rich natural waters. Copyright © 2003 Elsevier Ltd

1. INTRODUCTION

The motivation of this study is the improved quantification of the interactions of iron with silica in aquatic surface environments. Such information is crucial to better understand the effect of aqueous silica on the geochemical cycle of iron and related trace elements. Iron (III) oxides and hydroxides play an important role in a variety of environmental and industrial processes, such as chemical weathering, soil formation, and water treatment for purification and consumer use. Colloidal and amorphous Fe(III) oxy-hydroxides are important adsorbents for many toxic organic and inorganic substances in natural waters, and thus they significantly influence the fate and

toxicity of trace cations and anions in aquatic environments (e.g., Tessier et al., 1985; Manceau et al., 2000; Pokrovsky and Schott, 2002).

Silicic acid, Si(OH)₄, is a major component of natural waters, attaining concentrations as high as 100 mg/L. There is ample evidence of silica-iron association in both surficial and hydrothermal processes. Important amounts of silica are incorporated into natural Fe-hydroxide precipitates such as ferrihydrite (Carlson and Schwertmann, 1981; Schwertmann and Taylor, 1989; Konhauser and Ferris, 1996). A strong chemical affinity between Fe(III) and Si is demonstrated by the formation, under both natural and laboratory conditions, of iron silicates like nontronite, ferric chlorites/serpentes and smectites (e.g., Decarreau et al., 1987; Bailey, 1988; Toth and Fritz, 1997; Drief et al., 2001, and references therein). The presence of silica in solution significantly influences Fe(III) hydrolysis and precipitation. For example, aqueous silica inhibits goethite

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formation from solution and prevents colloidal Fe(III) hydroxides from coagulating (Schenk and Weber, 1968; Anderson and Benjamin, 1985; Vempati and Loeppert, 1989). The formation of ferrihydrite over lepidocrocite during oxidation of Fe^{2+} solutions is favored in the presence of silicic acid (e.g., Schwertmann and Thalmann, 1976; Mayer and Jarrel, 1996). Despite that several studies have been devoted to adsorption and incorporation of aqueous silica onto Fe(III) (oxy)hydroxides (McPhail et al., 1972; Sigg and Stumm, 1980; Hansen et al., 1994a, 1994b; Davis et al., 2002), little is known about the molecular structure of iron-silicate surface complexes and the position of silicon atoms in the iron (III) solids. Quantitative information about the stoichiometry and stability of Fe-Si complexes in aqueous solution is limited to $\text{FeSiO}(\text{OH})_3^{2+}$ that forms at acid pH between Fe^{3+} and monomeric $\text{Si}(\text{OH})_4$ (Weber and Stumm, 1965; Olson and O'Melia, 1973; Reardon, 1979). Structural aspects of the interactions between silicic acid and ferric iron hydroxide complexes during Fe(III) hydrolysis are, however, poorly known. The aim of the present work is to use X-ray Absorption Fine Structure (XAFS) measurements at the Fe K-edge to characterize the local environment around Fe in dilute aqueous solutions in the presence of dissolved silica over a wide range of pH.

In many studies dealing with iron hydrolysis in concentrated solutions ($>0.1\text{ m}$) and surface complexation of cations and anions with Fe(III) oxy-hydroxides and colloids (e.g., Combes et al., 1989, 1990; Waychunas et al., 1993; Bottero et al., 1994; Manceau and Charlet, 1994; Rose et al., 1997a; Manceau et al., 2000, and references therein), XAFS spectroscopy was the chosen method because it provides quantitative information at the atomic scale about the local structural environment of Fe and other elements in any system, including their aqueous and surface complexes (ligand's identity and number, and inter-atomic distances). Until now, however, the exact structural position of silicon in Fe-dominated solids or solutions could not be detected by XAFS spectroscopy, because the weak signal from the light silicon backscaterers at the Fe K-edge ($\sim 7100\text{ eV}$) is usually masked by the presence of heavier atoms (Fe) in concentrated iron solutions (e.g., Doelsh et al., 2000) or Si-bearing iron hydroxides (Manceau et al., 1995). Although limited information about the structural environment around silicon (and similar light elements like Al, P; e.g., Rose et al., 1997a; Ildefonse et al., 1998; Gehlen et al., 2002) can be also obtained at the Si K-edge ($\sim 1800\text{ eV}$), the low energy provides a short spectral range, and usually requires high Si concentrations and heroic efforts to diminish the absorption by the sample environment. In contrast, currently available third-generation synchrotron radiation sources allow XAFS analyses in the fluorescence mode at iron K-edge at Fe concentrations as low as 10^{-2} to 10^{-3} m . Although such concentrations are higher than those usually encountered for Fe in natural waters, they are low enough to significantly slow down Fe(III) polymerization and precipitation, and to allow accurate characterization of Fe-Si interactions (complexation) which can be easily masked by Fe-Fe condensation and solid-phase formation at higher Fe concentrations. Therefore, it is expected that the XAFS results obtained in this study should help in the refinement of aqueous and surface complexation models of $\text{Fe}(\text{III})_{\text{aq}}/\text{Fe}$ hydroxides with silica and other trace oxy-anions

which could be used for better understanding of the behavior of iron in natural waters and its role in the scavenging of inorganic and organic compounds.

2. MATERIAL AND METHODS

2.1. Sample Preparation

Iron(III)-bearing aqueous solutions for XAFS (X-ray Absorption Fine Structure) experiments were prepared by dissolving weighted amounts of hydrated ferric nitrate ($\text{Fe}(\text{NO}_3)_3 \cdot 9\text{H}_2\text{O}$, 99.99%, Aldrich Chemicals) or ferric chloride ($\sim 0.6\text{-m}$ FeCl_3 aqueous solution, Merck Chemicals) in doubly de-ionized water. The iron content of the initial solutions was checked by flame Atomic Absorption and was found to be close to the theoretical concentration within 2% of the total value. The pH of the solutions was adjusted by adding 1-*m* nitric (HNO_3) or hydrochloric (HCl) acid, or 1-*m* tetramethylammonium hydroxide ($\text{N}(\text{CH}_3)_4\text{OH}$) Titrisol solutions; it was measured before and after each experiment with a combination pH glass electrode (SCHOTT-H61) calibrated on the activity scale using the NIST tartrate, phthalate, phosphate and borate buffers ($\text{pH}_{20^\circ\text{C}} = 1.69, 4.00, 6.87$ and 9.21 , respectively) and 0.1-*m* HCl and 0.01-*m* NaOH solutions ($\text{pH}_{20^\circ\text{C}} = 1.11$ and 12.12 , respectively). Silicic acid (H_4SiO_4) was introduced into the Fe-bearing solutions as weighted amounts of a 0.67-*m* Si solution, prepared by equilibration of amorphous silica powder ($\text{SiO}_2 \cdot 2.2\text{H}_2\text{O}$) at ambient temperature in a 0.5-*m* tetramethylammonium hydroxide solution. Tetramethylammonium was chosen in this study as the major cation, because of its extremely low complexing affinity for metals in aqueous solution (Wesolowski et al., 1995) even in comparison to other commonly used background cations (Na^+ , K^+). The influence of $(\text{CH}_3)_4\text{N}^+$ on Fe hydrolysis and Fe-Si complexing in solution can thus be neglected when analyzing XAFS spectra. To avoid possible polymerization and/or precipitation during sample preparation, solutions were vigorously stirred using a magnetic stirrer during mixing and pH adjustment, and then immediately filtered through a 0.2- μm Millipore filter. This procedure was expected to eliminate possible heterogeneous nucleation sites and, thus, to delay significantly colloidal and solid-phase formation (see below). Measurements were started within 1 h, and completed usually within 5 h after solution preparation. Iron and silica concentrations in the experimental solutions used for XAFS experiments were 0.01 and 0.05 *m*, respectively (with exception of a single FeCl_3 solution containing 0.16 *m* silica). Such high silica concentrations were chosen to allow unambiguous detection of the light silicon backscaterers (see below). The pH of most solutions ranged from 1 to 3, which corresponds to H/Fe and OH/Fe hydrolysis total molal ratios from ~ 10 to ~ 0.1 , and from ~ 0.1 to ~ 4 (depending on Si concentration), respectively. At $\text{pH} > 3$ in both Si-free and Si-bearing solutions, rapid precipitation of Fe-hydroxide was observed; however, at strongly basic pH ($\text{pH} = 12\text{--}13$, $\text{OH/Fe} > 5\text{--}10$) this precipitate dissolved yielding a clear purple solution whose XAFS spectrum was also recorded (Tables 2–4). Monomeric silica, total Si and total Fe concentrations were determined using the molybdate blue colorimetric method (Koroleff, 1976), ICP-MS (Inductively Coupled Plasma Mass Spectrometry; see details in Pokrovski and Schott, 1998), and flame atomic absorption spectroscopy (Pokrovski et al., 2002a), respectively. Selected solutions before and after XAFS measurements were filtered through a 0.2- μm filter, diluted to 5 to 20 ppm of Fe and Si, and stabilized by 2 wt.% HNO_3 for analyses of total Fe and Si performed after the XAFS session. Monomeric silica was analyzed as a function of pH (2.00, 3.65 and 10.95) and time elapsed (from few minutes to 5 h) after preparation in 0.05-*m* Si solutions prepared exactly in the same manner as the experimental ones. Each of these solutions was diluted at constant pH to 7 to 10 ppm of Si with aqueous HNO_3 or $\text{N}(\text{CH}_3)_4\text{OH}$ and analyzed for $\text{Si}(\text{OH})_4$ with an Autoanalyser II Technicon colorimeter (molybdate blue method) from 2 min to 1 h after dilution. This procedure allows, even at alkaline pH where the kinetics of depolymerization is very fast, to measure only monomeric silicic acid (and, possibly, silicic acid dimers) present in the initial concentrated solution (Koroleff, 1976; Iler, 1979; Dietzel and Usdowski, 1995). Analyses performed on solutions before and after XAFS spectra acquisition showed no changes in total Si and Fe concentrations in the limit of the analytical uncertainties ($\pm 5\%$ of total concentration).

Table 1. Structural parameters of the Fe(III) atomic environment in iron-bearing solids derived in this study from fitting Fe K-edge EXAFS spectra, and comparison with X-ray diffraction (XRD) data.

Compound, chemical formula	EXAFS ^a				XRD			Ref. ^b
	Atom	R_{EXAFS} (Å)	N_{EXAFS}	σ^2 (Å ²)	Atom	R_{XRD} (Å)	N_{XRD}	
Iron (III) nitrate, Fe(NO ₃) ₃ · 9H ₂ O	O	2.00	5.5	0.003	O1	1.974	2	HB
					O2	1.986	4	
Epidote, Ca ₂ Al _{2.17} Fe _{0.81} Si ₃ O ₁₂ OH	O1	1.88	1.7	0.002	O1	1.88	1	D
	O2	2.00	1.8	0.001	O2	1.95	1	
	O3	2.24	1.7	0.015	O3	2.00	2	
	Al	2.95	2f ^c	0.004	O4	2.24	2	
	Si1 + Ca1	3.28	4f	0.007	Al1	2.99	2	
	Si2	3.46	2f	0.002	Si1	3.26–3.29	3	
	Ca2	3.74	3f	0.015	Ca1	3.29	1	
					Si2	3.46	2	
				Ca2	3.64–3.86	3		
Goethite, α-FeOOH	O1	1.94	3.1	0.005	O1	1.953	3	S
	O2	2.08	2.5	0.006	O2	2.09	3	
	Fe1	3.03	2f	0.005	Fe1	3.01	2	
	Fe2	3.28	2f	0.006	Fe2	3.28	2	
	Fe3	3.44	4f	0.008	Fe3	3.46	4	
γ-LiAlO ₂ with 1 wt.% Fe ³⁺	O	1.86	3.1	0.0005	O	1.82 ^d	4 ^d	WR
	Al	3.15	4.2	0.007	Al	3.12	4	
Error of EXAFS fit		±0.01	±0.5	±30%				

^a R = Fe-neighbor mean distance, N = Fe coordination number, σ^2 = squared Debye-Waller factor (relative to $\sigma^2 = 0$ adopted in the calculation of reference amplitude functions by FEFF, see text).

^b References for XRD data: HB: Hair and Beattie (1977); D: Dollase (1971); S: Szytula et al. (1968); M: Marezio (1965), WR: Waychunas and Rossman (1983).

^c f = parameter was fixed to crystallographic value during fit.

^d parameters derived from EXAFS fitting.

Silicic acid remained essentially monomeric (~90% of total Si) in the alkaline solutions during experiments. This is in agreement with equilibrium calculations using the stabilities of silicate oligomers formed in basic solutions (Baes and Mesmer, 1976). In acid solutions, silica polymerization occurred within few minutes, and only 15 to 20% of total Si was found to be monomeric, depending on solution pH (2–3.65) and time (few minutes to 5 h). Nevertheless, the structural environment of Fe derived from the XAFS spectra was found to be insensitive to the degree of silica polymerization (see below). In a solution with the highest Si concentration (0.16 *m*) at pH 3, a silica gel was observed after measurements. No significant pH changes (± 0.1 –0.2 pH unit) were recorded after experiments for all solutions. Although the majority of solutions at pH > 2 were significantly supersaturated with respect to Fe oxy-hydroxides (α-FeOOH, amorphous Fe[OH]₃), the constancy of Fe and Si concentrations and pH, measured before and after experiments on solutions filtered through 0.2-μm membrane filters, indicates that no precipitation of Fe (or Fe-Si) solid phases occurred during measurements. However, the presence of nano-particles of such phases or colloids, which could have passed through the 0.2-μm filter, cannot be completely excluded on the bases of chemical analyses only. Nevertheless, XAFS analysis indicates that such particles were likely to be insignificant in our solutions. First, successive EXAFS scans, recorded in this study during at least 3 to 5 h for all solutions, were very similar for each given sample and were without suspicious noise and flaws, thus demonstrating that the local atomic environment of Fe remains the same in the experimental solutions. Second, as it will be shown below (section 3), the number of Fe-Fe neighbors derived from the EXAFS spectra of acid solutions corresponded to the dominant presence of small Fe polymeric species, and was much lower than what would be expected for Fe-bearing solid/colloidal particles. This indicates that no significant solid-phase formation occurred over at least the time of measurements (<5 h).

2.2. XAFS Spectra Acquisition and Data Reduction

XAFS spectra (including the X-ray absorption near edge structure region or XANES, and the extended X-ray absorption fine structure region or EXAFS) were collected at ambient temperature ($20 \pm 1^\circ\text{C}$) in the fluorescence mode at the Fe K-edge (~7130 eV) over the energy range 7000 to 8000 eV on the collaborative research group IF BM32 beamline at the European Synchrotron Radiation Facility (ESRF, Grenoble, France). The storage ring was operated at 6 GeV with a 200–150 mA current. The beam energy was selected using a Si(111) double crystal monochromator with dynamical sagittal focussing. The fluorescence spectra were collected using a Canberra 30-element solid-state detector. The solutions were placed in a special Teflon cell with Kapton-film windows. The cell handling and spectra acquisition procedure were similar to those described in Pokrovski et al. (2000, 2002b). To obtain the necessary signal-to-noise ratio for quantifying next-nearest neighbors (Fe, Si) around the absorbing Fe atoms, three scans (of ~50 min/scan data collection time) for each solution were collected to high energy values (up to $k = 14$ –15 Å⁻¹, where k is the photo-electron momentum). After acquisition, all scans for each sample were carefully inspected for each individual solid state detector and, if found reasonably free of beam intensity fluctuations and other flaws, were added together. Goethite, α-FeOOH, ferric iron nitrate, Fe(NO₃)₃·9H₂O, epidote, Ca₂(Al,Fe³⁺)₃(Si₂O₇)(SiO₄)(O,OH)₂, and LiAlO₂ doped with 1 wt.% of Fe³⁺, which can serve as model compounds for Fe local environment (see Wilke et al., 2001), were recorded in the transmission mode to exclude self-absorption effects.

Data analysis was performed with the XAFS 2.6 package (Winterer, 1997) using the “standard” procedure (Teo, 1986). Details about spectra reduction can be found in Pokrovski et al. (2002b). Briefly, the before-edge spectral region (7000–7100 eV) was fitted using a straight line and subtracting this from the whole spectrum. Then, the spectra were normalized for atomic absorption, based on the average absorp-

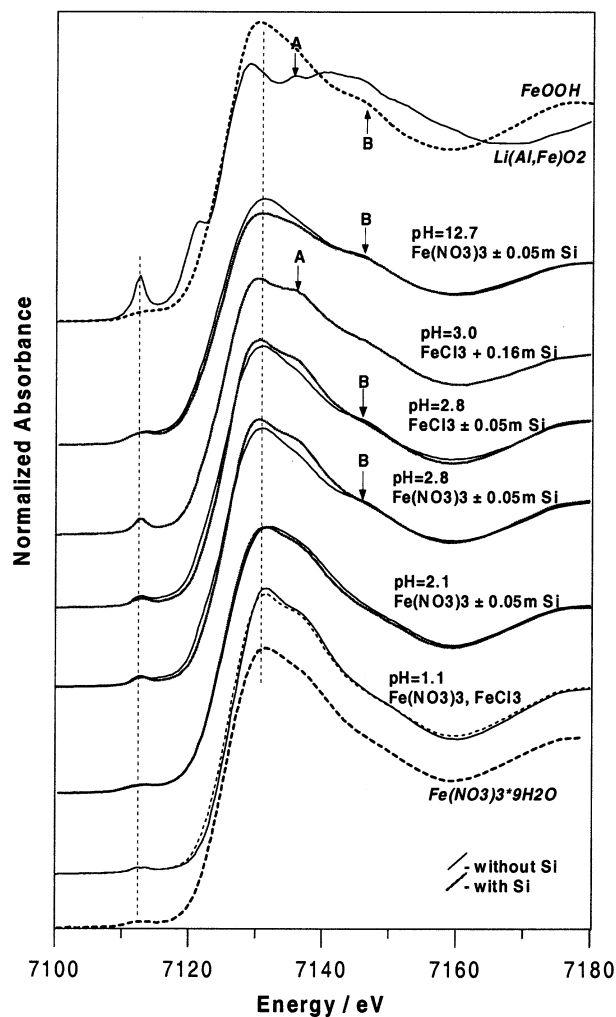


Fig. 1. XANES spectra at Fe K-edge of selected Fe-bearing solutions and solids. The vertical dashed lines indicate the pre-edge feature and main edge of iron. Arrows labeled A and B denote different main-edge features discussed in the text.

tion coefficient of the spectral region 7200 to 7400 eV. Pre-edge features were processed by subtracting a 2-degree polynomial fit to the absorption background, 2 eV from both sides of the prepeak, to determine the pre-edge intensity (normalized to the main-edge), peak area and half-height width. Background above the edge was determined by fitting cubic spline functions. Energies were recalculated into k -space (\AA^{-1}) with E_0 (i.e., the energy where k is zero) arbitrarily chosen at zero of the main edge first derivative. Normalized background-subtracted EXAFS spectra were weighted by k^3 , filtered over the k range from ~ 3 to $13\text{--}14 \text{\AA}^{-1}$, and Fourier-transformed (Kaiser-Bessel window with τ values of 2.5) to produce radial distribution functions (RDF) that isolate different atomic shells. To extract structural information, one or several RDF contributions were back-transformed into momentum space and modeled by least-square fitting. This gives the identity of the backscattering atoms (e.g., O, Si, or Fe), Fe-neighbor distance (R) and coordination number (N), and the Debye-Waller factor (σ^2) for a given scattering path. In addition to these structural parameters, a single nonstructural parameter, ΔE_0 , was varied in initial fits to account for its estimate made by FEFF (see below). It was found that the optimal value of ΔE_0 varied within $< \pm 2$ eV for most reference solids and aqueous samples. Raw EXAFS spectra were also fitted with multiple shells; they produced values of structural parameters similar to those extracted from fits of filtered signals for an individual shell. For

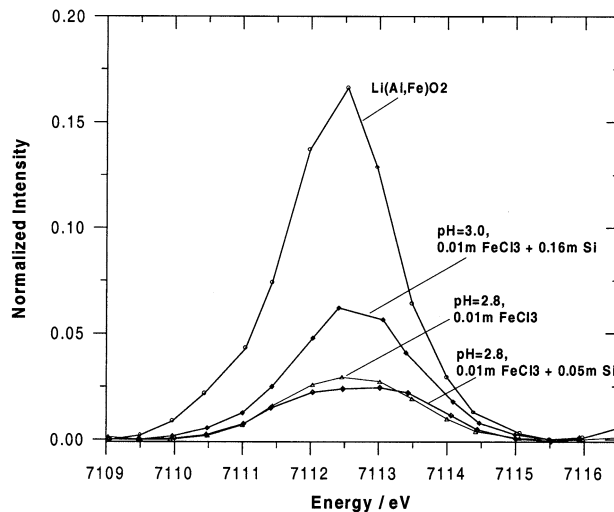


Fig. 2. Pre-edge features of selected 0.01- m FeCl_3 solutions and ferric lithium aluminate. Note a significant enhancement of the pre-edge intensity for the 0.16- m Si solution, suggesting the presence of tetracoordinated Fe like in $\text{Li}(\text{Al,Fe})\text{O}_2$.

most spectra which showed weak second shell contributions around Fe, fits of the total EXAFS spectrum were, however, less robust and exhibited larger uncertainties than those for the individual shell fits. Consequently, the results of the latter models were considered to be more reliable and they are reported in this study. Theoretical backscattering amplitude and phase-shift functions for Fe-O and Fe-Si, Fe-Fe, and Fe-Cl pairs were computed using the FEFF 8 ab initio code (Ankudinov et al., 1998), assuming the local structure around Fe in andradite garnet, $\text{Ca}_3\text{Fe}_2\text{Si}_3\text{O}_{12}$ (Hazen and Finger, 1989), α - FeOOH (Szytula et al., 1968), and FeOCl (Lind, 1970), respectively. The validity of the generated amplitude and phase functions was checked by fitting EXAFS spectra of Fe(III)-bearing solids. Interatomic distances and coordination numbers derived from fitting of model compounds are in close agreement with those obtained by X-ray Diffraction (XRD) in the limit of $\pm 0.02 \text{\AA}$ and ± 0.5 atoms for first and next-nearest atomic shells (Table 1). The influence of anharmonic disorder in determining structural parameters was checked using the cumulant expansion method (Crozier et al., 1988; Farges, 1996; Farges et al., 1996; Soldo et al., 1998). The typical values of third- and fourth-order cumulants (C_3 and C_4) found when fitting the filtered signal for the Fe first coordination shell were of the order of 10^{-4} and, consequently, too small to affect significantly R and N values derived from least-square fits without cumulants. As a result, to reduce the number of variables and thus the uncertainties on the derived parameters, we choose the harmonic approximation when extracting R , N and σ values for all Fe atomic shells. The influence of possible multiple scattering (MS) events within the Fe first and second coordination shells on the EXAFS spectra was also tested using the FEFF code, assuming local T_d and O_h geometries around Fe(III), as found in the model compounds investigated. However, we cannot exclude the possibility of unusual local geometries that can favor MS events. Nevertheless, with the exception of the most acid solutions ($\text{pH} \sim 1$), MS contributions were found to be negligible when fitting the next-nearest atomic shells.

3. RESULTS

3.1. XANES Analysis

3.1.1. Pre-Edge region

Normalized XANES spectra for selected Fe-bearing solutions and solids are presented in Figure 1. All spectra exhibit pre-edge features at 7112.5 ± 0.2 eV (Table 2, Figs. 1 and 2), which are caused by electronic transitions from the 1s to 3d

Table 2. Pre-edge and main-edge parameters from iron K-edge XANES spectra of Fe(III)-bearing solids and solutions measured in this study.

Sample	Pre-edge			Main edge	
	Centroid position (eV)	Normalized height	Peak area	Half-width (eV)	Main-edge crest energy (eV)
<i>Solids</i>					
Fe(III) nitrate, Fe(NO ₃) ₃ · 9H ₂ O	7112.3	0.016	0.064	3.7	7131.5
Goethite, α-FeOOH	7112.7	0.020	0.068	3.4	7130.3
Epidote	7112.5	0.034	0.097	2.7	7127.6
LiAlO ₂ doped with 1 wt.% Fe ³⁺	7112.5	0.174	0.381	2.2	7129.0
<i>Si-free solutions (0.01-m Fe)^a</i>					
Fe(NO ₃) ₃ + 0.1 = m HNO ₃ , pH 1.15, H/Fe = 10	7112.8	0.020	0.063	3.2	7131.6
FeCl ₃ + 0.1-m HCl, pH 1.10, H/Fe = 10	7112.8	0.020	0.063	3.2	7131.5
Fe(NO ₃) ₃ , pH 2.10, H/Fe = 1	7112.6	0.020	0.070	3.5	7132.0
Fe(NO ₃) ₃ , pH 2.36, OH/Fe = 0	7112.5	0.022	0.068	3.2	7131.3
Fe(NO ₃) ₃ , pH 2.80, OH/Fe = 2.8	7112.5	0.032	0.078	2.5	7130.9
FeCl ₃ , pH 2.80, OH/Fe = 2.8	7112.5	0.032	0.074	2.5	7130.9
Fe(NO ₃) ₃ , pH 12.9, OH/Fe = 15	7112.5	0.026	0.101	3.9	7131.2
<i>Si-bearing solutions (0.01-m Fe)^a</i>					
Fe(NO ₃) ₃ + 0.05-m Si, pH 2.10, OH/Fe = 0	7112.6	0.023	0.095	4.0	7131.5
Fe(NO ₃) ₃ + 0.05-m Si, pH 2.80, OH/Fe = 2.0	7112.6	0.027	0.078	2.7	7130.5
FeCl ₃ + 0.05-m Si, pH 2.80, OH/Fe = 2.0	7112.6	0.024	0.078	2.7	7130.5
FeCl ₃ + 0.16-m Si, pH 2.95, OH/Fe = 4.0	7112.6	0.068	0.155	2.3	7130.3
Fe(NO ₃) ₃ + 0.05-m Si, pH 12.5, OH/Fe = 10	7112.5	0.029	0.117	4.0	7131.2
Error of deconvolution	±0.2 eV	±10%	±15%	±0.3 eV	±0.5 eV

^a H/Fe and OH/Fe stand for the molal ratio of iron salt concentration to the total added acid (HNO₃ or HCl) and base (N(CH₃)₄OH) concentration, respectively, during solution preparation (see text for details).

states in the iron atom (Apted et al., 1985; Combes et al., 1989, 1990; Wilke et al., 2001). The intensity, shape and position of the pre-edge features reflect the redox state and symmetry of Fe sites (Wilke et al., 2001). The normalized height, peak area, and half-width of these features for experimental solutions and reference solids are summarized in Table 2. Because of the relatively low spectral resolution provided by the Si(111) monochromator, these pre-edge features could not be modeled precisely into different components and, therefore, were treated as a single peak.

The pre-edge parameters of both Fe nitrate and chloride solutions at acid pH (pH < 2.2) are similar to that of Fe(NO₃)₃·9H₂O (Table 2), in which Fe³⁺ is octahedrally coordinated by six H₂O (Hair and Beattie, 1977). This confirms that a similar geometry around the Fe³⁺ cation exists in acid aqueous solution (Apted et al., 1985). With increasing pH, the intensity of the pre-edge feature grows slightly, most likely reflecting an increased distortion of centrosymmetry within the FeO₆ octahedron (although the appearance of minor amounts of four- and/or five-coordinated Fe(III) species cannot be excluded) (Combes et al., 1989; Wilke et al., 2001).

No significant differences in the pre-edge features were observed in the presence of 0.05-*m* Si at pH from 2.1 to ~13. By contrast, at higher silica concentration (0.16 *m*) at pH = 3, the pre-edge peak is enhanced by a factor of ~2 to 3 (Fig. 2). This enhancement indicates the presence of a tetrahedral environment for Fe in this solution, like in solid LiAlO₂ doped with Fe(III), in which Fe is tetracoordinated by oxygens (Waychunas and Rossman, 1983). Assuming that the pre-edge peak height and area are proportional to the percentage of tetrahedral Fe (Combes et al., 1989; Peterson et al., 1997; Wilke et al., 2001), and using the values of corresponding parameters for

Li(Al,Fe)O₂, the proportion of ¹⁴Fe in the 0.16-*m* Si solution can be estimated as ~22 ± 5 mol%. This value is confirmed by the EXAFS analysis (see below).

3.1.2. Main-Edge region

Spectra of 0.01-*m* Fe nitrate and chloride silica-free and silica-bearing acid solutions (pH < 2.5) are similar in shape and white-line positions (~7131.5 eV) to that of hydrated iron nitrate Fe(NO₃)₃·9H₂O suggesting the same Fe octahedral coordination. In Si-free solutions of higher pH (pH > 2.5), the edge-crests are slightly shifted to lower energies (~7131.0 eV); the spectra's shape is very similar to that of goethite (α-FeOOH) having a feature at ~7148 eV (identified by arrow B in Fig. 1). Because the goethite structure is characterized by distorted FeO₃(OH)₃ octahedra linked together by their edges and corners (Manceau and Combes, 1988), similar structures could be expected in the iron oxy-hydroxide oligomer complexes formed at pH > 2.5 in solution. In the presence of 0.05-*m* Si at pH of ~2.8, the goethite features are partially suppressed, and the spectra shape more resembles those for acid solutions. These qualitative changes in the presence of silica at pH > 2.5 might indicate a more symmetrical Fe atomic environment in Fe-Si complexes like that at more acid pH. At higher silica concentration (0.16-*m* Si polymerized solution, pH = 3), the main-edge shape and position are significantly different from those for the 0.05-*m* Si solutions. The white-line position (7130 eV) and the after edge-crest feature (at ~7138 eV, marked by arrow A in Fig. 1) are very close to those observed in the spectrum of Li(Al,Fe)O₂ (Fig. 1, Table 2). This similarity, together with the enhanced pre-edge feature (see

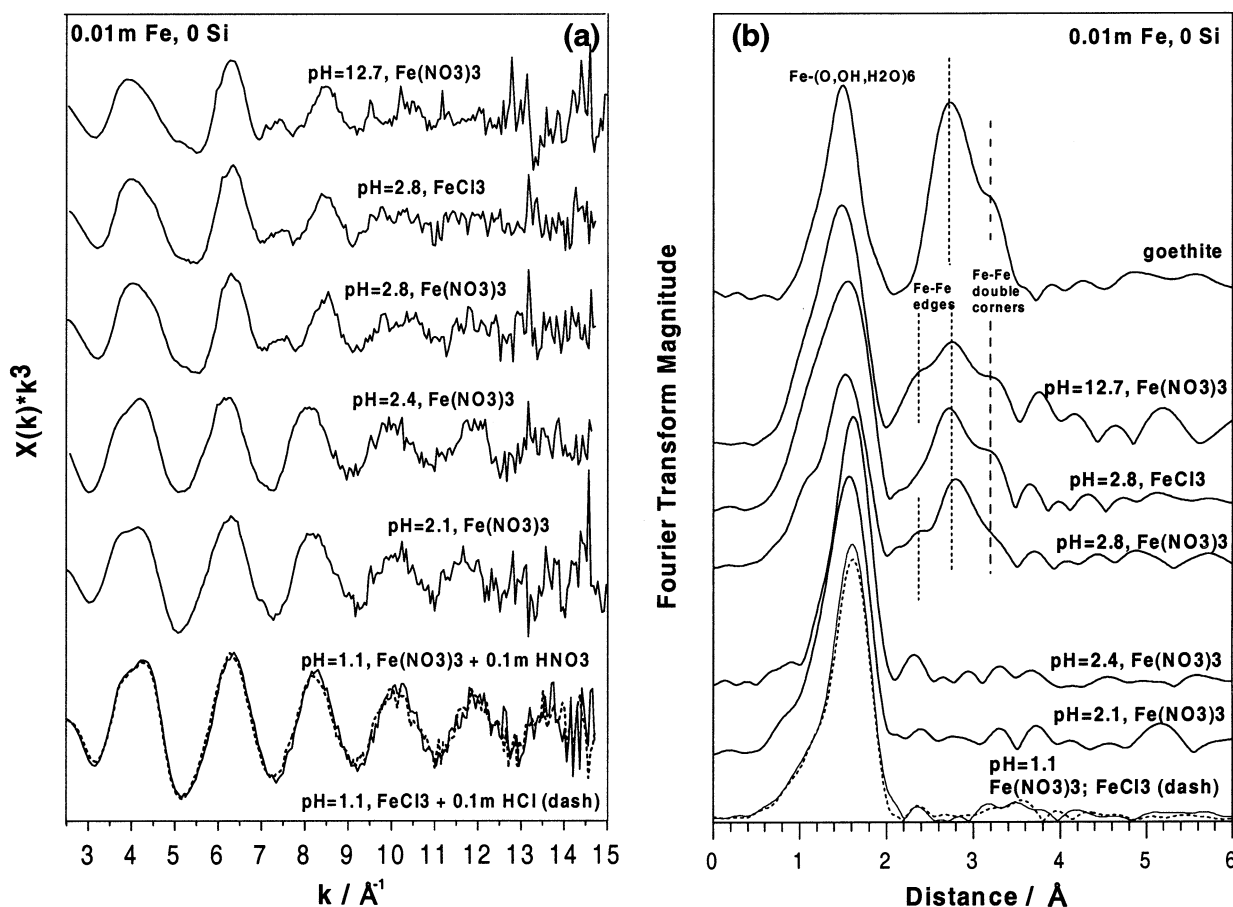


Fig. 3. k^3 -weighted EXAFS spectra at Fe-K edge of selected Si-free Fe nitrate and chloride solutions as a function of pH (a); and their corresponding Fourier Transforms (not corrected for phase shift) (b). The vertical lines on (b) indicate Fe-Fe edge and double-corner linkages in the polymerized Fe oxy-hydroxide complexes.

section 3.1.1), suggests the presence in this solution of iron in tetrahedral coordination.

3.2. EXAFS Analysis of Silica-Free Solutions

3.2.1. First atomic shell

Normalized k^3 -weighted EXAFS spectra and their Fourier Transforms (FT) of the studied Si-free solutions are presented in Figures 3a and 3b. In strongly acid chloride and nitrate solutions (pH < 2.4), both EXAFS and FT spectra exhibit a single contribution from a first shell of oxygens around Fe. Modeling this signal yields 6.0 ± 0.5 oxygens with a mean Fe-O distance of $2.00 \pm 0.01 \text{ \AA}$. This is in agreement with the EXAFS measurements of Combes et al. (1989) on highly concentrated Fe(III) hydrolyzed solutions ($m_{\text{Fe}} > 0.1 \text{ m}$), and consistent with the predominance of $\text{Fe}(\text{H}_2\text{O})_6^{3+}$ and $\text{Fe}(\text{H}_2\text{O})_5\text{OH}^{2+}$ (Baes and Mesmer, 1976). Including chlorine in the fit of the first shell contribution of the 0.01-*m* $\text{FeCl}_3 + 0.1\text{-m}$ HCl solution slightly improves the fit quality, suggesting the possible presence of $\sim 0.3 \pm 0.2$ Cl atoms (Table 3). No Cl atoms were detected in the first Fe shell of $\text{FeCl}_3 \pm \text{HCl}$ ($0.01 < m_{\text{Cl}} < 0.1$) solutions at higher pH.

At pH from 2.4 to 2.8 in both nitrate and chloride solutions, the number of neighbors remains the same within errors as in

the strongly acid solutions ($N_{\text{Fe-O}} = 5.5 \pm 0.5$), but the average Fe-O distance decreases by 0.03 \AA ($R_{\text{Fe-O}} = 1.97 \text{ \AA}$). The first coordination sphere of Fe deviates a little from a regular octahedron, as indicated by an increase of the DW factor (from 0.005 to ~ 0.01 , see Table 3).

At basic pH, fits of the first Fe shell are consistent with two different Fe-O distances (1.92 and 2.05 \AA), resembling closely the iron local environment in goethite (see Table 1). This is demonstrated in Figure 4. The EXAFS spectra of the first shell of both goethite and Fe basic solution are similar, but exhibit a distinct phase shift (indicated by the dashed line) in comparison to acid solutions.

3.2.2. Second and third atomic shells

At pH < 2, the minor feature observed at $\sim 3.5 \text{ \AA}$ on the FT spectra (Fig. 3b, not corrected for phase shift) might correspond to both multiple scattering paths within the symmetrical FeO_6 octahedron and the second solvation shell around Fe^{3+} , by analogy with EXAFS and WAXS (wide-angle X-ray scattering) results on other trivalent cations like Cr^{3+} , In^{3+} , or Ga^{3+} (Lindqvist-Reis et al., 1998). At pH from 2 to 2.4, this contribution is hardly discernable, probably due to the distortion of the octahedral symmetry of the first Fe shell. The small lobes

Table 3. Structural parameters of the first atomic shell of iron obtained from fitting Fe K-edge EXAFS spectra of 0.01 molal Fe(III)-nitrate and -chloride solutions with and without aqueous silica.^a

Solution	pH _{20°C}	Atomic pair	R (Å)	N (atom)	σ^2 (Å ²)	chi ²
Si-free solutions						
Fe(NO ₃) ₃ , ~0.1- <i>m</i> HNO ₃	1.03	Fe-O	1.996 (0.002)	5.9 (0.3)	0.003 (0.0005)	1.36
Fe(NO ₃) ₃ , ~0.1- <i>m</i> HNO ₃	1.15	Fe-O	1.994 (0.002)	6.1 (0.2)	0.004 (0.0005)	0.60
FeCl ₃ , ~0.1- <i>m</i> HCl	1.10	Fe-O	2.002 (0.003)	6.0 (0.2)	0.004 (0.0005)	0.30
		Fe-Cl	2.30 (0.03)	0.3 (0.2)	0.002 (0.001)	
Fe(NO ₃) ₃	2.10	Fe-O	2.00 (0.01)	5.5 (0.3)	0.005 (0.001)	1.20
Fe(NO ₃) ₃	2.36	Fe-O	2.02 (0.01)	5.0 (0.5)	0.005 (0.001)	0.23
Fe(NO ₃) ₃	2.80	Fe-O	1.98 (0.01)	5.9 (0.3)	0.009 (0.002)	0.55
FeCl ₃	2.80	Fe-O	1.97 (0.02)	5.5 (0.4)	0.010 (0.003)	0.60
Fe(NO ₃) ₃	12.9	Fe-O1	1.92 (0.03)	2.5 (0.5)	0.006 (0.002)	0.70
		Fe-O2	2.05 (0.03)	2.8 (0.5)	0.006 (0.002)	
Si-bearing solutions						
Fe(NO ₃) ₃ , 0.05 · <i>m</i> Si	2.10	Fe-O	2.00 (0.01)	5.6 (0.4)	0.005 (0.001)	0.43
Fe(NO ₃) ₃ , 0.05 · <i>m</i> Si	2.80	Fe-O	1.995 (0.005)	5.9 (0.3)	0.009 (0.002)	0.57
FeCl ₃ , 0.05- <i>m</i> Si	2.80	Fe-O	1.99 (0.01)	6.2 (0.4)	0.010 (0.002)	0.46
FeCl ₃ , 0.16- <i>m</i> Si (colloidal)	2.95	Fe-O1	1.84 (0.02)	1.1 (0.2)	0.0005 (0.0002)	0.60
		Fe-O2	1.99 (0.02)	4.5 (0.5)	0.012 (0.002)	
Fe(NO ₃) ₃ , 0.05 · <i>m</i> Si	12.5	Fe-O1	1.93 (0.03)	3.7 (0.3)	0.005 (0.002)	0.76
		Fe-O2	2.07 (0.02)	1.7 (0.3)	0.006 (0.002)	

^a R = Fe-neighbor mean distance, N = Fe coordination number, σ^2 = squared Debye-Waller factor, $\text{chi}^2 = N/(F \times M) \times \sum (Y_{\text{exp}}^2 - Y_{\text{fit}}^2)^2$, where N = number of independent data, F = degrees of freedom, M = number of data points fitted, and $Y_{\text{exp}}^2 - Y_{\text{fit}}^2$ = difference between squared experimental and fitted EXAFS absorption coefficient ($X \cdot k^3$) of the filtered signal for each point (Press et al., 1986). Typical R -window range of FT used for back-transform to isolate the first shell is 0.6 to 2.2 Å. Values in parentheses represent uncertainty for each parameter.

visible at the FT's of these solutions above ~ 3 Å are likely to be caused by the spectral noise. No second-shell iron atoms ($N_{\text{Fe-Fe}} < \sim 0.1$) could be detected in our dilute solutions at pH < 2.4 .

At pH from 2.5 to 3, and from 12 to 13, the spectra exhibit an important second shell feature (Figs. 3a and 3b). This feature is, in part, similar to that of α -FeOOH, and corresponds to several Fe-Fe pairs with distances ranging from 2.9 to 3.5 Å. The appearance of Fe-Fe contributions clearly demonstrates the formation of polymeric iron-hydroxide complexes. For an FeCl₃ solution at pH 2.8, the second shell feature was successfully fitted using two Fe-Fe subshells with distances of 3.04 and 3.46 Å (Table 4). For iron nitrate solutions at pH 2.8 and 12.7, the best fits were achieved with three Fe-Fe subshells, at 2.94, ~ 3.09 , and 3.47 Å. These findings are in good agreement with the previous studies by XAFS spectroscopy of the hydrolysis of more concentrated iron chloride and nitrate solutions ($0.1 \leq m_{\text{Fe}} \leq 1.0$). Two Fe-Fe distances at 3.01 to 3.05 and 3.44 to 3.47 Å were detected during hydrolysis of FeCl₃ (e.g., Combes et al., 1989; Bottero et al., 1994), and at least three distances, at 2.88 to 3.05, 3.06 to 3.14, and 3.42 to 3.52 Å, in the case of Fe(NO₃)₃ hydrolysis (Rose et al., 1997b), in solutions with OH/Fe molal ratios > 1.5 . Four different Fe-Fe subshells, with distances ranging from 2.95 to 3.90 Å, were reported in a more recent study of iron hydroxide precipitates obtained by hydrolysis of 0.2-mol/L FeCl₃ solutions (Doelsh et al., 2000). Based on these studies and using the analogy with Fe oxy-hydroxide solids (Manceau and Combes, 1988, and references therein), whose structures consist of Fe-O/(OH)₆ octahedra sharing their faces ($R_{\text{Fe-Fe}} = 2.89$ Å in hematite), edges ($R_{\text{Fe-Fe}} = 2.95$ – 3.30 Å in goethite, hematite, akaganeite, lepidocrocite) and/or double corners ($R_{\text{Fe-Fe}} = 3.40$ – 3.50 Å in goethite, hematite), the Fe-Fe distances derived for the hydrolyzed solutions of the present study can be attributed to polymeric species with edge-

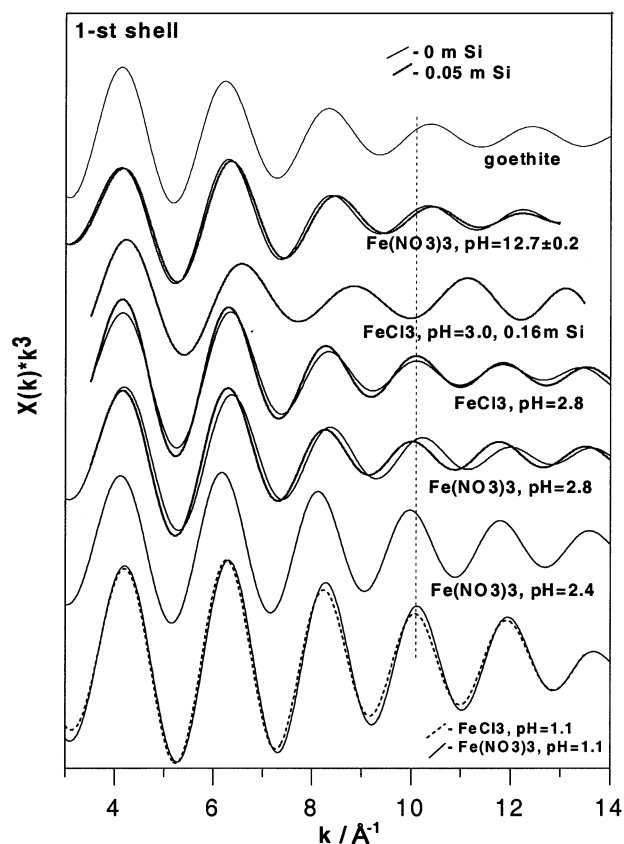


Fig. 4. EXAFS spectra of the first Fe atomic shell of selected 0.01-*m* Fe nitrate and chloride solutions at the indicated pH, with and without aqueous silica. The vertical line indicates phase shifts with increasing pH or Si concentration (see text for details).

Table 4. Structural parameters of the next-nearest atomic shells of iron obtained from fitting Fe K-edge EXAFS spectra of 0.01 molal Fe(III)-nitrate and -chloride solutions with and without aqueous silica.^a

Solution	pH _{20°C}	Atomic pair	R (Å)	N (atom)	σ^2 (Å ²)	chi ²
Si-free solutions						
Fe(NO ₃) ₃	2.8	Fe-Fe1	2.93 (0.03)	0.5 (0.2)	0.007 (0.003)	0.25
		Fe-Fe2	3.07 (0.02)	0.9 (0.4)	0.005 (0.003)	
		Fe-Fe3	3.46 (0.01)	0.5 (0.1)	0.004 (0.002)	
FeCl ₃	2.8	Fe-Fe1	3.04 (0.01)	1.3 (0.2)	0.010 (0.003)	0.28
		Fe-Fe2	3.46 (0.01)	0.5 (0.1)	0.005 (0.002)	
Fe(NO ₃) ₃	12.9	Fe-Fe1	2.95 (0.03)	0.8 (0.3)	0.006 (0.003)	0.30
		Fe-Fe2	3.12 (0.03)	1.2 (0.3)	0.006 (0.003)	
		Fe-Fe3	3.48 (0.04)	0.7 (0.3)	0.007 (0.002)	
Si-bearing solutions						
Fe(NO ₃) ₃ , 0.05- <i>m</i> Si	2.8	Fe-Fe	3.12 (0.03)	0.9 (0.2)	0.007 (0.002)	0.20
		Fe-Si	3.18 (0.03)	1.5 (0.5)	0.007 (0.002)	
FeCl ₃ , 0.05- <i>m</i> Si	2.8	Fe-Fe	3.12 (0.02)	1.3 (0.2)	0.007 (0.002)	0.16
		Fe-Si	3.17 (0.03)	2.0 (0.5)	0.007 (0.002)	
Fe(NO ₃) ₃ , 0.05- <i>m</i> Si	12.5	Fe-Fe1	2.93 (0.02)	0.3 (0.2)	0.007 (0.003)	0.30
		Fe-Fe2	3.11 (0.02)	1.5 (0.5)	0.007 (0.003)	
		±Fe-Si	3.17 (0.03)	2.0 (1.0)	0.007 f	

^a See footnotes for Tables 1 and 3.

Typical *R*-window range of FT used for back-transform to isolate the next-nearest shells is 2.2 to 3.5 Å for Si-free solutions, and 2.15 to 3.3 Å for Si-bearing solutions.

and double corner-sharing Fe-O₆ octahedra. Although the 2.94 ± 0.01 Å distance found for our nitrate solutions is close to that corresponding to Fe-Fe face-sharing in hematite (2.89 Å), the formation of edge-sharing structures during Fe hydrolysis at ambient temperature is unlikely (e.g., Combes et al., 1989; Jolivet et al., 1994). Hematite-like structures usually form from ferric gels at later stages of hydrolysis or after long aging times and at elevated temperatures that favors dehydration (e.g., Schwertmann and Taylor, 1989; Combes et al., 1990). The structures found in our dilute solutions are similar to those of Al(III), Ga(III), and Cr(III) hydroxide polymeric species formed in concentrated solutions (0.2–1.0 *m*) during the hydrolysis of nitrate or chloride salts of these metals (Jolivet et al., 1994; Michot et al., 2000; Pokrovski et al., 2002b). Note that beyond the first Fe shells, all hydrolyzed solutions investigated in this study were reasonably modeled using two to three Fe-Fe contributions, in accordance to the degree of freedom for our filtered second shell spectra which ranges between 5 and 9 (Stern, 1993). The addition of a fourth Fe-Fe shell (for example at ~3.3 or 3.9 Å) led to highly correlated and often unreasonable Debye-Waller factors and coordination numbers and decreased significantly the goodness of fit. Consequently, taking into account the signal-to-noise ratio and *k*-range of the spectra of the dilute solutions investigated in this study, a two- or three-shell model was kept throughout the analysis of Fe hydrolysis.

3.3. EXAFS Analysis of Silica-Bearing Solutions

3.3.1. First atomic shell

In the presence of 0.05-*m* aqueous silica in strongly acid (pH < 2–2.5) and basic (pH > 12) solutions, no changes in the first coordination shell of iron were detected in comparison with their Si-free analogs (Table 3), which is also in agreement with the XANES analysis (see above). In contrast, in the intermediate pH region (2.7 < pH < 3), Fe-O₆ distances in 0.05-*m* Si

chloride and nitrate solutions are distinctly longer than in their Si-free counterparts (2.00 instead of 1.97 Å, see Table 3 and Fig. 4) and close to those in more acidic solutions. This is also in line with the XANES analysis and can be interpreted by a partial destruction of the relatively short Fe-O-Fe linkages between FeO₆ octahedra and formation of longer Fe-OH or Fe-O-Si bonds.

In a more concentrated Si solution (0.16-*m* Si) at pH 3, the EXAFS spectrum is distinctly different, exhibiting important phase shifts and amplitude changes compared to the spectra of less concentrated Si solutions at similar pH (Figs. 4 and 5). Modeling of this spectrum is consistent with the presence of two distinct sets of Fe-O distances: ~1.84 Å (*N* = 1) and 1.99 Å (*N* = 4–5) (Table 3). The shorter Fe-O distance is very close to those found in tetrahedral sites of Fe³⁺ in Li(Al,Fe)O₂ (Table 1) and Fe(III)-bearing micas (e.g., Dyar et al., 2001; Giuli et al., 2001). The derived Fe-O distances and coordination numbers imply the presence of ~20% Fe in a tetrahedral coordination, probably substituting for Si in the tetrahedral framework of the silica polymers formed in this solution. Note that the percentage of ¹⁴⁴Fe deduced from the EXAFS analysis is in excellent agreement with that derived from the pre-edge data (see above).

3.3.2. Second and third atomic shells

In the spectra of acid solutions in the presence of 0.05-*m* Si (pH < 2.4), no second-shell Fe-Fe or Fe-Si contributions were detected in the limit of the spectral resolution (*N*_{Fe-Fe} < ~0.1; *N*_{Fe-Si} < ~0.5 atoms).

At pH 2.8 in contrast, important changes occur in the spectra of both 0.01-*m* ferric nitrate and chloride solutions in the presence of 0.05-*m* aqueous silica. This is manifested by the disappearance, in the total EXAFS spectrum, of the feature at ~7.5–8 Å⁻¹ which corresponded to the Fe-Fe contributions in Si-free solutions, and by distinct phase shifts in the region 7 to

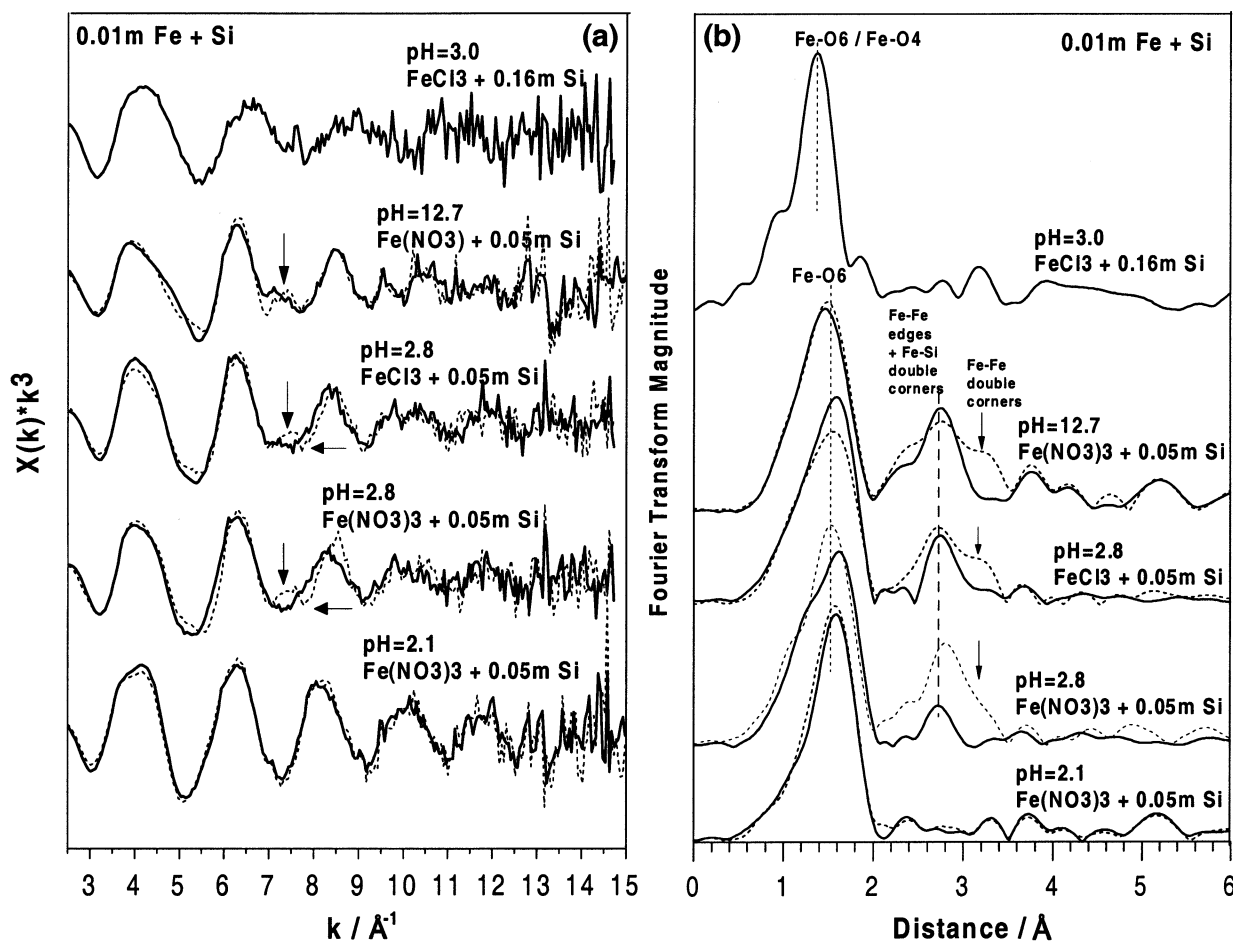


Fig. 5. k^3 -weighted EXAFS spectra at the Fe-K edge of 0.01-*m* Fe nitrate and chloride solutions in the presence of aqueous silica at different pH (a); and their Fourier Transforms (not corrected for phase shift) (b). Solid curves = Si-bearing solutions with indicated concentrations and pH; dashed curves = spectra of the corresponding Si-free solutions at the same pH (see also Fig. 3). The arrows indicate changes in the presence of aqueous silica: phase shifts (a); suppression of the double-corner Fe-Fe contributions in the second Fe atomic shell (b). The vertical lines on (b) stand for octahedral or tetrahedral first shell coordination of Fe, and Fe-Fe edge and Fe-Si double corner linkages in the second shell of iron in its silicate complexes.

9 \AA^{-1} (Fig. 5a). The contributions from the Fe-Fe edges at $\sim 2.95 \text{ \AA}$ for nitrate solutions, and Fe-Fe double-corner linkages for both nitrate and chloride solutions, apparent in the RDF's of the Si-free samples at these pH values, disappear in the presence of silica (Fig. 5b). The filtered second shell signal exhibits an important amplitude decrease and phase shift in the presence of silica (Fig. 6). These observations strongly suggest a destruction of the Fe-Fe linkages of iron hydroxide polymers, and the entering of Si in the second coordination shell of Fe. The analysis of the iron second shell for these samples was performed using either two Fe-Fe subshells or one Fe-Fe + one Fe-Si subshell. Figure 7 shows such fits for Fe nitrate and chloride solutions at pH 2.8 with 0.05-*m* Si. It can be seen in this figure that the presence of iron atoms only cannot account for the observed second shell signal. Including light neighbors like oxygens or Cl/NO₃ do not affect the fit; adding a third Fe-Fe subshell resulted in less robust models, yielding highly correlated and unreasonable N and σ^2 values. In contrast, significant good-quality fits were obtained for both solutions

using 0.9 to 1.3 Fe atoms at $3.12 \pm 0.03 \text{ \AA}$, and $\sim 2 \pm 0.5$ Si atoms at $3.18 \pm 0.03 \text{ \AA}$ (Fig. 7, Table 4). The uncertainties on these N and R values stem mostly from possible variations of the ΔE_0 parameter ($\pm 2 \text{ eV}$, see section 2.2). Because Fe-Fe and Fe-Si individual electronic waves are shifted by $> \pi/2$ and thus are almost out of phase, the amplitude of the resulting spectrum for the second shell is quite sensitive to the value chosen for ΔE_0 . Nevertheless, fits of the Fe-Fe/Si contributions at pH 2.8 were robust and always converged to the parameters reported in Table 4 whatever the starting values of R , N , and ΔE_0 . These results confirm quantitatively the observations of RDF's of these solutions, which indicate the destruction of Fe-Fe edge and double-corner bonds in the iron polymeric species in the presence of Si (Fig. 5b).

At pH 12.7 in the presence of 0.05-*m* Si, similar changes in the EXAFS spectrum were observed (Figs. 5 and 6). The suppression of Fe-Fe edge and corner contributions is apparent on the Fourier Transform (Fig. 5b). A fit of the second-shell signal confirms these observations, yielding two Fe-Fe edge

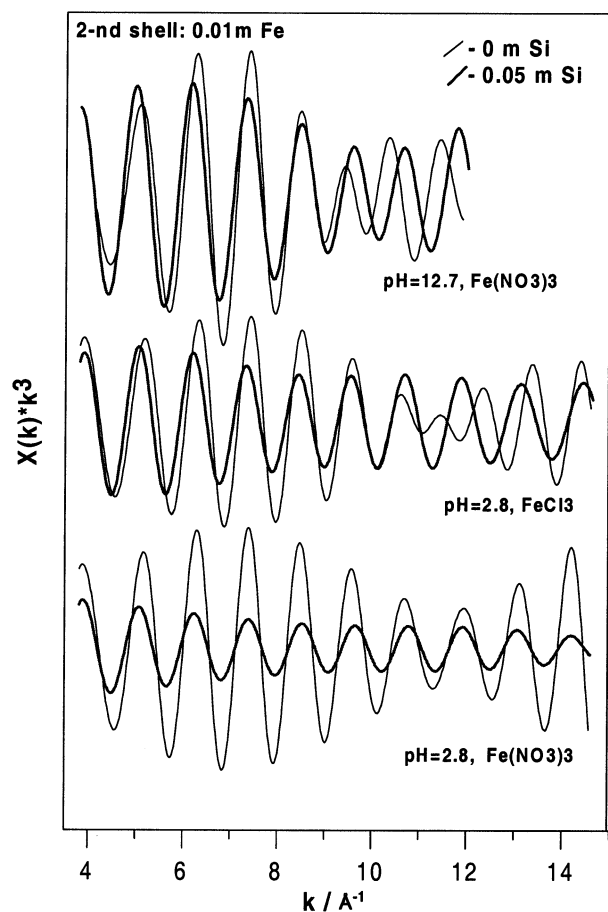


Fig. 6. Comparison of second-shell filtered EXAFS signal of Si-free and Si-bearing 0.01-*m* Fe solutions. Note important amplitude decrease and phase shifts in the presence of Si, suggesting a partial destruction of Fe-Fe bonds and formation of Fe-Si linkages.

contributions at 2.93 and 3.11 Å (Table 4). The amplitude decrease of the second shell Fe-Fe signal was, however, much weaker than for less hydrolyzed solutions at acid pH (Fig. 6). Consequently, fits were less sensitive to the presence of silicon. It was found that addition of 2 ± 1 Si atoms at 3.17 Å improves only slightly the fit quality. This is in agreement with previous EXAFS studies of Fe hydroxide precipitates formed in the presence of Si (Doelsh et al., 2000), and hydrous iron silicate scale deposits (Manceau et al., 1995) that failed to detect the light Si in the second iron coordination shell dominated by Fe-Fe contributions. At earlier hydrolysis stages ($\text{pH} < 3.0$) of our experimental solutions of low Fe concentration, the Fe-Fe contributions to the EXAFS signal are significantly lower and thus allow accurate detection of light silicon atoms in the second shell.

With increasing Si concentration (0.16-*m* Si, $\text{pH} = 3$), Fe-Fe second shell contributions vanish indicating the almost complete destruction of iron polymers. The detection of Si was also impossible, probably because of an important disorder associated with various $^{14}\text{Fe-O-Si}$ and $^{16}\text{Fe-O-Si}$ contributions. More measurements in Si-rich solutions are required to distinguish between these different linkages.

4. DISCUSSION

4.1. Iron Hydrolysis and Polymerization in Dilute Solutions

The results of this study demonstrate that the Fe^{3+} cation rapidly hydrolyses in dilute chloride and nitrate aqueous solutions via formation of Fe oxy-hydroxide polynuclear species composed of $\text{Fe}(\text{O}/\text{OH}/\text{H}_2\text{O})_6$ octahedra linked together by their edges and double corners. This is in good agreement with previous XAFS studies of Fe hydrolysis in concentrated solutions (>0.1 -*m* Fe) that have shown the formation of large Fe polymers in chloride and nitrate solutions with increasing OH/Fe ratios or aging time (e.g., Combes et al., 1989, 1990; Bottero et al., 1994; Rose et al., 1997b). The number of edge (2.90–3.05 Å) and double-corner (3.40–3.50 Å) linkages derived in those studies increase, respectively, from 1 to 3.5 and from 1 to 2 in solutions with OH/Fe ratios > 1.5 ($2 < \text{pH} < 5$), suggesting the presence of large Fe oxy-hydroxide entities (e.g., “ Fe_{24} ,” Bottero et al., 1994). Such species are regarded as precursors for goethite, akaganeite or hematite (depending on solution composition, temperature and time of aging). In the more dilute solutions of the present study, the number of edge- and corner-sharing neighbors at pH of ~ 2.8 is consistent with the formation of smaller Fe polymers, most likely a mixture of dimers, trimers and tetramers. At basic pH (OH/Fe > 5), the Fe-Fe coordination numbers ($N_{\text{edge}} = 2.0$, $N_{\text{corner}} = 0.7$; Table 4) correspond to tetramers or/and bigger polymeric species (e.g., “ Fe_{24} ”), in agreement with previous studies. The schematic structures of the main hydroxide polymers likely to form during Fe(III) hydrolysis are presented in Figure 8.

The results on Fe-hydrolysis obtained in this study can be compared with speciation calculations using the available stability constants for the iron dimer, $\text{Fe}_2(\text{OH})_2^{4+}$, and trimer, $\text{Fe}_3(\text{OH})_4^{5+}$, obtained from potentiometric measurements in dilute Fe nitrate or chloride solutions (e.g., Baes and Mesmer, 1976; Martell et al., 1998, and references therein). The distribution of Fe hydroxide species in a 0.01-*m* Fe(III) nitrate solution at 25°C and a 0.15-mol/L ionic strength is shown as a function of pH in Figure 9a. At $\text{pH} < 2.4$, calculations predict $< 20\%$ of total Fe is present as dimers, and $< 5\%$ as trimers, which corresponds to ~ 0.2 Fe atoms in the second Fe shell. This is close to the detection limit of Fe-Fe second neighbors of our EXAFS spectra. At higher pH ($2.6 < \text{pH} < 3.0$), calculations indicate that the dimer and trimer account for ~ 20 to 30% of total Fe each. This corresponds to $N_{\text{edge}} = \sim 0.4$ to 0.5 and $N_{\text{corner}} = \sim 0.3$ Fe atoms, assuming an edge-sharing dimer and a double-corner trimer (Bottero et al., 1994; Jolivet et al., 1994). Our EXAFS analysis yields significantly higher coordination numbers at these pH values, confirming the formation of bigger Fe polymers, likely tetramers (Fig. 8). Note that tetrameric species with edge and double-corner linkages were shown to form during Cr^{3+} and Ga^{3+} hydrolysis; they are also considered as precursors for the nucleation of Al and Fe(III) oxy-hydroxide phases (Jolivet et al., 1994, and references therein; Michot et al., 2000). The very high lability of Fe^{3+} in comparison to Al^{3+} or Cr^{3+} makes it difficult to stabilize small Fe polymers in supersaturated solution in the absence of strongly complexing ligands (Jolivet et al., 1994) which can explain the absence of thermodynamic data for Fe-tetramers

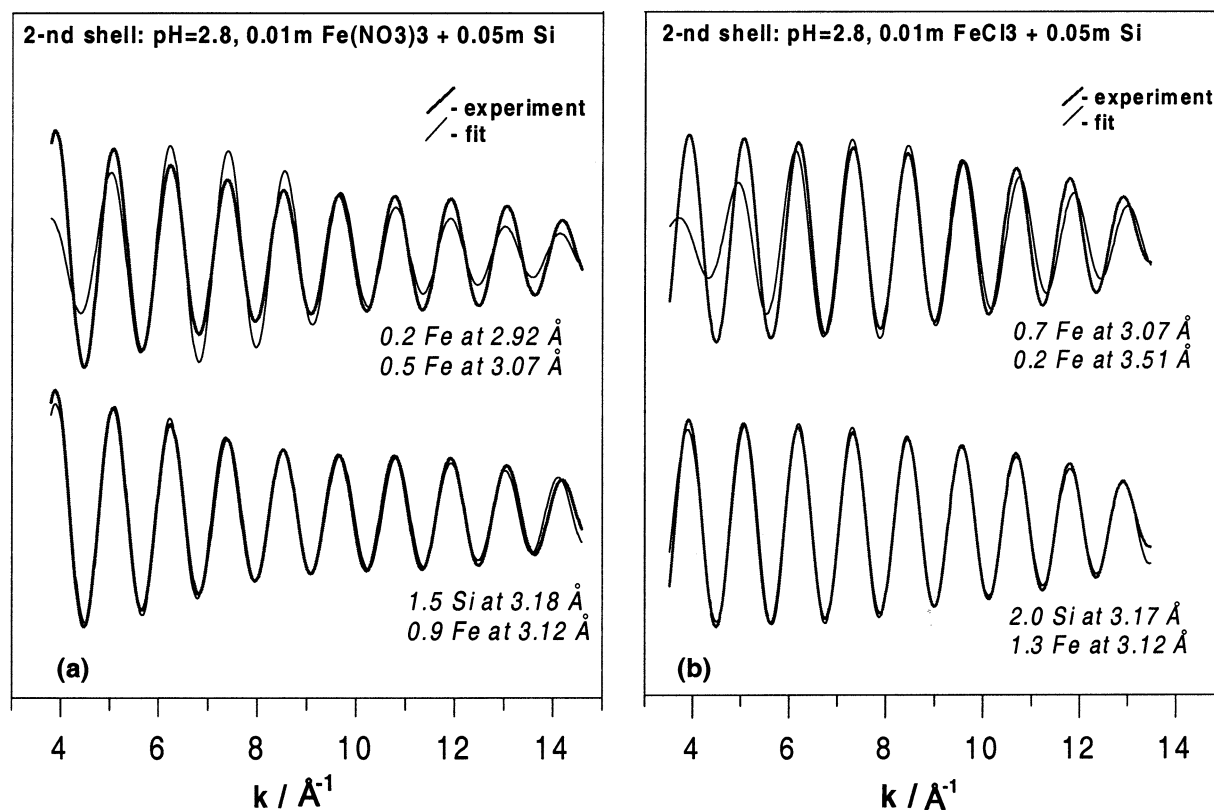


Fig. 7. Examples of EXAFS fits of the filtered second shell signal of 0.01-*m* ferric nitrate (a) and chloride (b) solutions at pH 2.8 in the presence of 0.05-*m* Si. It can be seen that Fe-Fe contributions only cannot account for the observed spectrum. Including a Fe-Si shell yields robust good-quality fits in a wide *k*-range (see text for details).

and the “Fe₂₄” polycation.¹ Consequently, the equilibrium calculations presented here give a simplified image of iron speciation especially at pH > 3–4, which is likely dominated by tetrameric and/or bigger species.

4.2. Iron-Chloride Complexing

The presence of chlorine neighbors was not detected in the first and second Fe atomic shell in hydrolyzed 0.01-*m* FeCl₃ solution at pH 2.8. This is in agreement with equilibrium calculations carried out using the available stability constants for Fe(III) chloride species (taken from the compilation of Tagirov et al., 2000), suggesting that only 6% of total Fe should be complexed in the form of FeCl²⁺ and FeCl₂⁺ species in this solution. In strongly acid non-hydrolyzed solution (pH ~ 1, 0.1-*m* Cl⁻, 0.01-*m* FeCl₃), less than 0.3 Cl atoms could be detected in the Fe first shell from analysis of its EXAFS spectrum. Speciation calculations for this solution composition indicate that FeCl²⁺ and FeCl₂⁺ represent respectively ~40 and ~10% of total Fe. This corresponds to an average Fe-Cl coordination number of 0.6, if one assumes inner-sphere complexation. The lower first shell coordination number obtained in

this study ($N_{\text{Fe-Cl}} \sim 0.3$, Table 3) suggests a portion of Fe-Cl species may be in the form of outer-sphere complexes which could not be detected by EXAFS analysis in this study. Thus, the presence of Cl atoms weakly bonded to Fe hydroxide species in outer coordination shells during iron hydrolysis cannot be completely excluded. More ordered second shell Fe-Fe structures observed in this study at pH 2.8 for 0.01-*m* chloride (two Fe-Fe subshells) compared to nitrate (three Fe-Fe subshells) solutions, imply that chloride can probably intervene in Fe hydrolysis in dilute solutions, as it was suggested for Cl-rich solutions (0.2–1.0-*m* Cl; Combes et al., 1989; Doelsh et al., 2000). However, the structural role of Cl⁻ during the hydrolysis of dilute Fe(III)-chloride solutions remains unclear. At ambient temperature, the negatively charged Cl⁻ could weakly interact with the outer shells of positively charged Fe polymeric oxy-hydroxide species and affect polymerization/nucleation kinetics. However, as it is demonstrated in this study, the role of Cl⁻ in Fe hydrolysis in dilute solutions is negligible as compared to that of silica.

4.3. Structure and Stoichiometry of Fe-Si Aqueous Complexes

XAFS measurements reported in this study unambiguously demonstrate that aqueous silica inhibits Fe³⁺ hydrolysis via formation of stable complexes with polymeric ferric oxy-hydroxide species. Fe-Si and Fe-Fe distances and coordination

¹ The value of the first-order rate coefficient for the exchange of water molecules from the bulk solution into the inner coordination sphere of the hexacoordinated aquo ion is 1.6×10^2 , 1.3 and $2.4 \times 10^{-6} \text{ s}^{-1}$ for Fe³⁺, Al³⁺ and Cr³⁺, respectively (Merbach and Akitt, 1990).

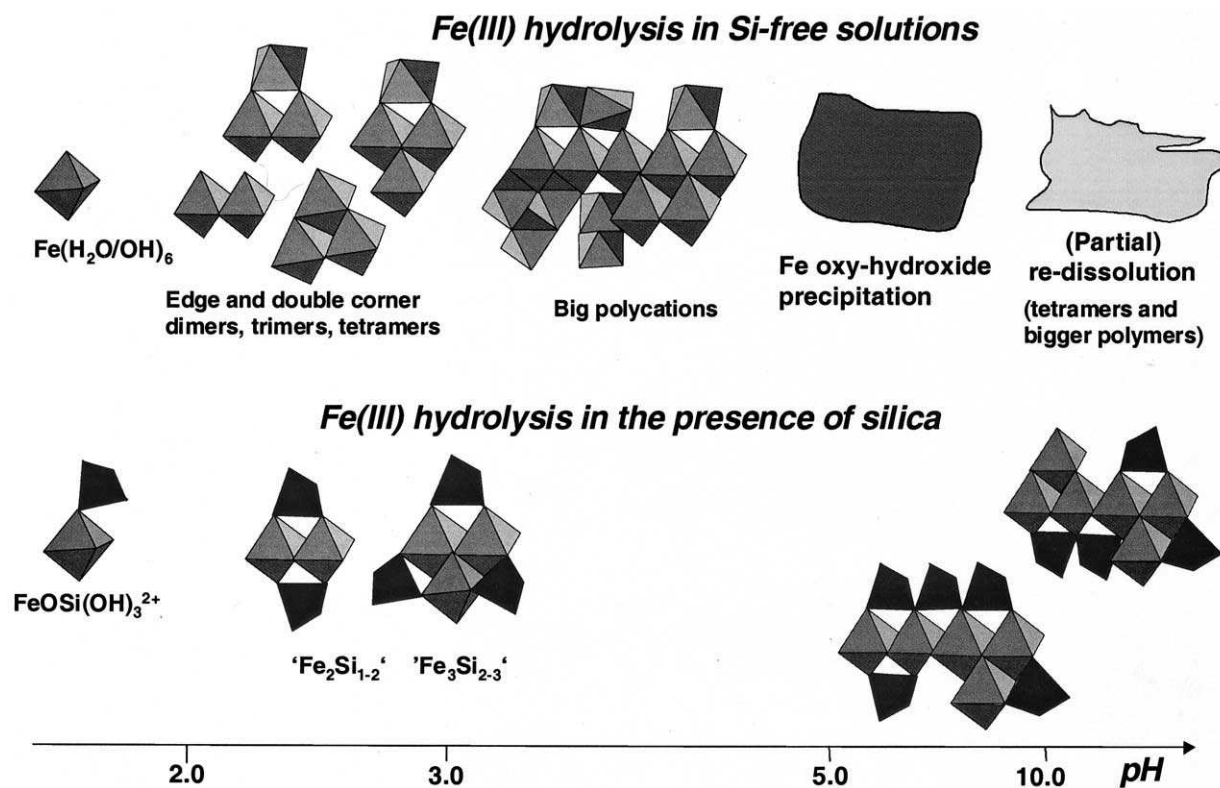


Fig. 8. Schematic illustration of Fe(III) hydrolysis in Si-free aqueous solution, and structures of Fe-Si complexes proposed in this study for moderately concentrated Si-solutions. Octahedra represent the first $\text{Fe}(\text{O}/\text{OH}/\text{H}_2\text{O})_6$ atomic shell; tetrahedra stand for the silicate moieties $\text{Si}(\text{O}/\text{OH})_4$ of monomeric silicic acid or silica polymers.

numbers derived in this study indicate that SiO_4 -tetrahedra substitute for double-corner FeO_6 -octahedra in the Fe(III)-polymers. The iron-silicon distance of 3.18 Å found in both nitrate and chloride solutions is close to those for Fe-As and Fe-Se in ${}^2\text{C}$ -type surface complexes formed by slightly bigger arsenate and selenate anions with hydrous ferric oxides (~3.25 Å, Manceau and Charlet, 1994; Manceau, 1995; Waychunas et al., 1993, 1995), and those for Fe-P found in Fe(III) phosphate minerals and aqueous solution (~3.20 Å, Rose et al., 1996, 1997a). The iron-iron coordination numbers corresponding to the Fe-Fe edge linkages are affected only slightly in the presence of aqueous silica both at acid and basic pH (Table 4). This suggests that at the early stages of hydrolysis, after the Fe-Fe dimer formation (Jolivet et al., 1994), aqueous silica is not capable of destroying these bonds significantly. The iron-iron edge distances in the silicate complexes found in this study (3.12 Å) are the same in both nitrate and chloride solutions at pH 2.8, suggesting that the edge-linked Fe polymers accommodate silica tetrahedra in a very similar way, independently of the nature of the ligand (Cl^- versus NO_3^-). This confirms that neither chloride nor nitrate influence significantly Fe-Si interactions in dilute solutions typical of natural waters.

The numbers of Fe and Si neighbors around iron in the silicate complexes formed at acid pH ($N_{\text{Fe-Fe}} = 0.9\text{--}1.3$; $N_{\text{Fe-Si}} = 1.0\text{--}2.0$) is consistent with stoichiometries $\text{Fe}_2\text{Si}_{1-2}$ and $\text{Fe}_3\text{Si}_{2-3}$, corresponding to a mixture of Fe-Fe dimers and planar trimers linked to one to three silicate anions by their free double corners (Fig. 8). At basic pH, the Fe-Fe and Fe-Si

coordination numbers found in this study imply that more polymerized Fe-Fe species, such as edge-sharing tetramers or bigger, are likely to form similar structures with silica. Note that the speciation of silica in 0.05-*m* Si solutions investigated in the present work is dominated by polymeric species (~80% of total Si) at acid pH and by monomeric silicate anions (>90% of total Si) at basic pH (see section 2.1). Nevertheless, the structural position of silicon atoms in the Fe-Si complexes appears to be similar at both acid and basic pH. This might reflect the preferential interaction of hydrolyzed Fe-polymers with monomeric $\text{Si}(\text{OH})_4$. Indeed, reactions with big silicate polymers are likely to be hampered by slow kinetics and steric constraints (e.g., Swaddle, 2001). However, it cannot be excluded that Fe might interact with certain configurations of silicic acid polymers in acid solution. The structures of Fe-Si complexes proposed in this study are summarized in Figure 8.

By analogy with other tetrahedral anions like arsenate, selenate or phosphate, silicate can be expected to form also other types of bonds with Fe-octahedra, such as edge (type ${}^1\text{E}$, $R_{\text{Fe-Si}} \sim 2.8\text{--}2.9$ Å) and single-corner (type ${}^1\text{C}$, $R_{\text{Fe-Si}} \sim 3.4\text{--}3.6$ Å) linkages (e.g., see Manceau and Charlet, 1994; Manceau, 1995, for the linkage terminology). Nevertheless, none of these contributions was detected in this study. The ${}^1\text{E}$ -linkages are not known in Fe(III) silicate and phosphate minerals. Thus, the edge-edge linkages between FeO_6 and SiO_4 are supposed to be negligible in solution, at least at the conditions of the present work. The existence of ${}^1\text{C}$ bonds in the Si-rich solutions investigated in this study, however, cannot be

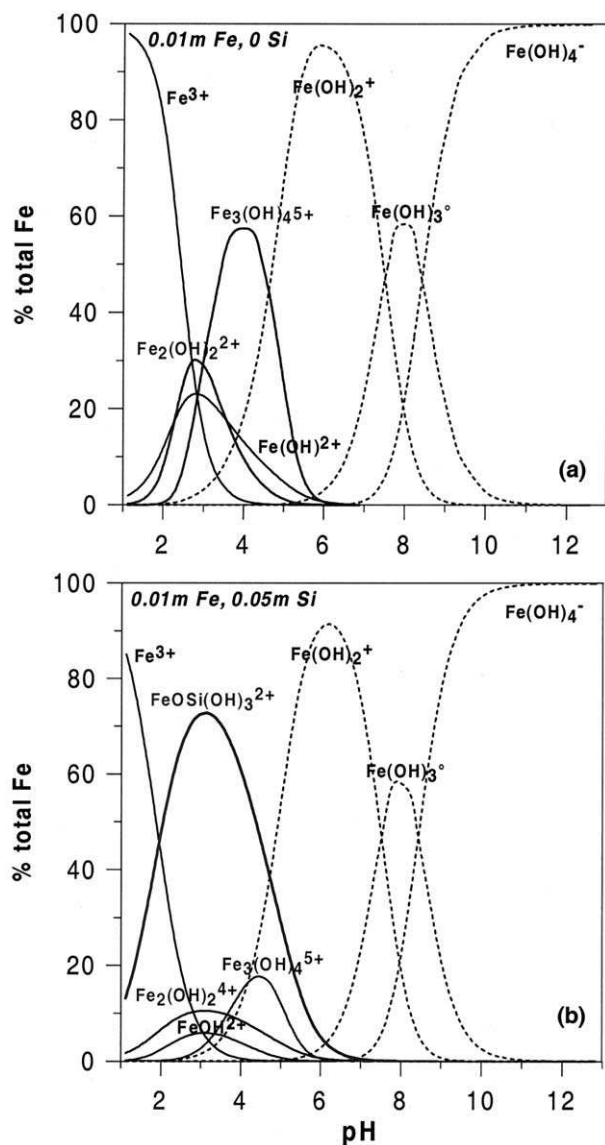


Fig. 9. Predicted iron speciation in a 0.01-*m* ferric nitrate Si-free solution (a), and in the presence of 0.05-*m* monomeric silica (b) at 25°C as a function of solution pH. Calculations were performed using the stability constants for Fe hydroxide species reported by Baes and Mesmer (1976), and that for the iron-silica complex as the average value from Weber and Stumm (1965), Olson and O'Melia (1973), and Reardon (1979): $\text{Fe}^{3+} + \text{Si}(\text{OH})_4(\text{aq}) = \text{FeOSi}(\text{OH})_3^{2+} + \text{H}^+$, $\log_{10} K (25^\circ\text{C}, I = 0) = 0.03 \pm 0.10$. Ionic strength was fixed at 0.15, and the activity coefficients of aqueous species were approximated by the extended Debye-Hückel equation with $\hat{a} = 4.5 \text{ \AA}$ for all charged species. Note that because of the lack of stability constants for iron oligomers bigger than trimers, this diagram represents a simplified image of the iron speciation at $\text{pH} > 3-4$. Consequently, the monomeric species at $\text{pH} > 4$ (dashed lines) are likely not dominant, and they are shown only for consistency with the calculations.

ruled out. Because of the expected higher degree of freedom associated with such single-corner bonds, their detection by EXAFS spectroscopy, which is very sensitive to disorder, is likely to be difficult. This is confirmed by the fact that no Si could be detected around the iron atom in the present study in a Si-rich solution (0.16-*m* Si) at pH 3. The presence of single

corner Fe-O-Si bonds is also expected in the monomer complex $\text{FeOSi}(\text{OH})_3^{2+}$ reported in acid solutions (Weber and Stumm, 1965; Olson and O'Melia, 1973; Reardon, 1979). Equilibrium calculations using the stability constant for this species derived from the studies cited above suggest that this complex should account for $50 \pm 10\%$ of total Fe at $\text{pH} \sim 2.1$ in the presence of 0.05-*m* $\text{Si}(\text{OH})_4$ (Fig. 9b). This percentage corresponds to an average number of Si neighbors of ~ 0.5 in the second coordination sphere of Fe at this pH. However, no Si neighbors could be found in the spectrum of a 0.05-*m* total Si solution at pH 2.1, in which silica was partially polymerized (see above). This might reflect both the difficulties to detect, by EXAFS spectroscopy, the light silicon in the disordered second shell environment of iron, and the lower affinity of Fe^{3+} cation for polymeric silicate species than for monomeric silicic acid (see above). More measurements at different Si concentrations are needed to confirm this hypothesis.

5. CONCLUDING REMARKS

The results of this study demonstrate that aqueous silica affects the hydrolysis of Fe(III) in dilute solution via the formation of iron silicate aqueous complexes. In these structures, silica substitutes for double-corner FeO_6 -octahedra in iron oxy-hydroxide polymeric complexes existing at the early stages of Fe(III) hydrolysis, likely by forming ^2C -type (double corner) complexes with small Fe oxy-hydroxide polymers whose structure consists of FeO_6 -octahedra linked together by common edges. To our knowledge, this is the first quantitative evidence of the silica atomic environment in ferric silicate aqueous complexes.

The formation of soluble Fe-Si complexes delays the nucleation and growth of solid ferric oxy-hydroxides from aqueous solution. The significant silica concentrations found in these minerals in nature (e.g., Carlson and Schwertmann, 1981; Konhauser and Ferris, 1996) are likely to be a consequence of the aqueous Fe-Si complexing. However, the effect of silica on Fe(III) hydrolysis is significantly weaker than on Al^{3+} or Ga^{3+} hydrolysis. Due to the high affinity of Al and Ga for the tetrahedral coordination, these elements form much stronger covalent-bonded silicate complexes in which the metal is tetra-coordinated even at acid pH and low silica concentrations (< 0.01 -*m* Si) (Pokrovski et al., 2002b). Ferric iron, in contrast, keeps an octahedral environment at moderate Si concentrations, and takes a tetrahedral coordination only in solutions highly enriched in silica (> 0.1 -*m* Si), in which Fe^{3+} may substitute for Si in the Si-polymer network. The observation, in high-temperature hydrothermal and metamorphic environments, of micas and amphiboles having significant proportions of Fe^{3+} in tetrahedral coordination (e.g., Dyar et al., 2001; Giuli et al., 2001) could thus be indicative of high silica activities during the formation of such minerals. In near-surface conditions, which are characterized by much lower aqueous silica activities, Fe^{3+} remains essentially hexa-coordinated in low-temperature oxy-hydroxides and clays (e.g., Sherman and Vergo, 1988; Manceau et al., 1998).

The geometry and structure of the silicate complexes formed at the surfaces of iron oxy-hydroxide minerals and colloids are expected to be similar to those found for the Fe-Si aqueous solutions investigated in this study. Via its linking to the free

corners of FeO₆-octahedra, aqueous silica can decrease significantly the number of available sorption sites in Si-rich natural waters, and thus reduce both the incorporation of other trace anions and cations into iron hydroxides, and the bioavailability of their surfaces for Fe-metabolizing microorganisms (e.g., Meng et al., 2000; Santelli et al., 2001). The results on Fe-Si complex stoichiometries and structures obtained in this work may help improve models of silicate sorption by iron hydroxide phases, and abiotic or microbially catalyzed dissolution of iron silicate minerals.

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REFERENCES

- Anderson P. R. and Benjamin M. M. (1985) Effects of silicon on the crystallization and adsorption properties of ferric oxides. *Environ. Sci. Technol.* **19**, 1048–1053.
- Ankudinov A. L., Ravel B., Rehr J. J., and Conradson S. D. (1998) Real-space multiple-scattering calculation and interpretation of X-ray-absorption near-edge structure. *Phys. Rev. B* **58**, 7565–7576.
- Apted M. J., Waychunas G. A., and Brown G. E. Jr (1985) Structure and specification of iron complexes in aqueous solutions determined by X-ray absorption spectroscopy. *Geochim. Cosmochim. Acta* **45**, 421–429.
- Baes C. F. Jr. and Mesmer R. E. (1976) *The Hydrolysis of Cations*. John Wiley, New York.
- Bailey S. W. (1988) Odinite, a new dioctahedral-trioctahedral Fe³⁺-rich 1:1 clay mineral. *Clay Miner.* **23**, 237–247.
- Bottero J.-Y., Manceau A., Villieras F., and Tchoubar D. (1994) Structure and mechanisms of formation of FeOOH(Cl) polymers. *Langmuir* **10**, 316–319.
- Carlson L. and Schwertmann U. (1981) Natural ferrihydrite in surface deposits from Finland and their association with silica. *Geochim. Cosmochim. Acta* **45**, 421–429.
- Combes J. M., Manceau A., Calas G., and Bottero J.-Y. (1989) Formation of ferric oxides from aqueous solutions: A polyhedral approach by X-ray absorption spectroscopy: I. Hydrolysis and formation of ferric gels. *Geochim. Cosmochim. Acta* **53**, 583–594.
- Combes J. M., Manceau A., and Calas G. (1990) Formation of ferric oxides from aqueous solutions: A polyhedral approach by X-ray absorption spectroscopy: II. Hematite formation from ferric gels. *Geochim. Cosmochim. Acta* **54**, 1083–1091.
- Crozier E. D., Rehr J. J., and Ingalls R. (1988) Amorphous and liquid systems. In *X-Ray Absorption: Principles, Applications, Techniques of EXAFS, SEXAFS and XANES* (eds. D. C. Koningsberger and R. Prins), pp. 373–442. Wiley-Interscience, New York.
- Davis C. C., Chen H. W., and Edwards M. (2002) Modeling silica sorption to iron hydroxide. *Environ. Sci. Technol.* **36**, 582–587.
- Decarreau A., Bonnin D., Badaut-Trauth D., Couty R., and Kaiser P. (1987) Synthesis and crystallogenesis of ferric smectite by evolution of Si-Fe coprecipitates in oxidizing conditions. *Clay Miner.* **22**, 207–223.
- Dietzel M. and Uzdowski E. (1995) Depolymerization of soluble silicate in dilute aqueous solutions. *Colloid Polym. Sci.* **273**, 590–597.
- Doelsch E., Rose J., Masion A., Bottero J. Y., Nahon D., and Bertch P. M. (2000) Speciation and crystal chemistry of iron (III) chloride hydrolyzed in the presence of SiO₄ ligands. 1. An Fe K-edge EXAFS study. *Langmuir* **16**, 4726–4731.
- Dollase W. A. (1971) Refinement of the crystal structures of epidote, allanite and hancockite. *Am. Miner.* **56**, 447–464.
- Drief A., Nieto F., and Sanchez-Navas A. (2001) Experimental clay-mineral formation from a subvolcanic rock by interaction with 1M NaOH solution at room temperature. *Clays Clay Miner.* **49**, 92–106.
- Dyar M. D., Delaney J. S., and Sutton S. R. (2001) Fe XANES spectra of iron-rich micas. *Eur. J. Mineral.* **13**, 1079–1098.
- Farges F. (1996) Does Zr-F complexation occur in magmas? *Chem. Geol.* **127**, 253–268.
- Farges F., Brown G. E. Jr., Navrotsky A., Gan H., and Rehr J. J. (1996) Coordination chemistry of Ti(IV) in silicate glasses and melts: III. Glasses and melts from ambient to high temperatures. *Geochim. Cosmochim. Acta* **60**, 3055–3065.
- Gehlen M., Beck L., Calas G., Flank A. M., van Bennekom A. J., and van Beusekom J. E. E. (2002) Unravelling the atomic structure of biogenic silica: Evidence of the structural association of Al and Si in diatom frustules. *Geochim. Cosmochim. Acta* **66**, 1601–1609.
- Giuli G., Paris E., Wu Z., Brigatti F., Cibin G., Montana A., and Marcelli A. (2001) Experimental and theoretical XANES and EXAFS study of tetra-ferriphlogopite. *Eur. J. Miner.* **13**, 1099–1108.
- Hair N. J. and Beattie J. K. (1977) Structure of hexaqua(iron(III)) nitrate trihydrate. Comparison of iron(II) and iron(III) bond lengths in high-spin octahedral environments. *Inorg. Chem.* **16**, 245–250.
- Hansen H. C. B., Raben-Lange B., Raulund-Rasmussen K., and Borggaard O. K. (1994a) Monosilicate adsorption by ferrihydrite and goethite at pH 3–6. *Soil Sci.* **158**, 40–46.
- Hansen H. C. B., Wetche T. P., Raulund-Rasmussen K., and Borggaard O. K. (1994b) Stability constants for silicate adsorbed to ferrihydrite. *Clay Miner.* **29**, 341–350.
- Hazen R. M. and Finger L. W. (1989) High-pressure crystal chemistry of andradite and pyrope: Revised procedures for high-temperature diffraction experiments. *Am. Miner.* **74**, 353–359.
- Ildefonse P., Cabaret D., Saintavrit P., Calas G., Flank A. M., and Lagarde P. (1998) Aluminium X-ray absorption near edge structure in model compounds and Earth's surface minerals. *Phys. Chem. Miner.* **25**, 112–121.
- Iler R. K. (1979) *Chemistry of Silica—Solubility, Polymerisation, Colloid and Surface Properties and Biochemistry*. John Wiley, New York.
- Jolivet J.-P., Henry M., and Livage J. (1994) *De la Solution à l'Oxyde*. InterEditions and CNRS Editions, Paris, France.
- Konhauser K. O. and Ferris F. G. (1996) Diversity of iron and silica precipitation by microbial mats in hydrothermal water, Iceland: Implications for Precambrian iron formations. *Geology* **24**, 323–326.
- Koroleff F. (1976) Determination of silicon. In *Methods of Seawater Analysis*. (ed. K. Grasshoff). Springer-Verlag, Berlin, Germany.
- Lind M. D. (1970) Refinement of the crystal structure of iron oxychloride. *Acta Cryst.* **B26**, 1058–1062.
- Lindqvist-Reis P., Munoz-Paez A., Diaz-Moreno S., Pattanaik S., Parson I., and Sandstom M. (1998) The structure of the hydrated Ga(III), In(III) and Cr(III) ions in aqueous solutions. A large angle X-ray scattering and EXAFS study. *Inorg. Chem.* **37**, 6675–6683.
- Manceau A. (1995) The mechanism of anion adsorption on iron oxides: Evidence for the bonding of arsenate tetrahedra on free Fe(O,OH)₆ edges. *Geochim. Cosmochim. Acta* **59**, 3647–3653.
- Manceau A. and Combes J. M. (1988) Structure of Mn and Fe oxides and oxyhydroxides: A topological approach by EXAFS. *Phys. Chem. Miner.* **15**, 283–295.
- Manceau A. and Charlet L. (1994) The mechanism of selenate adsorption on goethite and hydrous ferric oxide. *J. Colloid Interface Sci.* **168**, 87–93.
- Manceau A., Ildefonse P., Hazemann J. L., Flank A. M., and Gallup D. (1995) Crystal chemistry of hydrous iron silicate scale deposits at the Salton Sea geothermal field. *Clays Clay Miner.* **43**, 304–317.
- Manceau A., Chateigner D., and Gates W. P. (1998) Polarized EXAFS, distance-valence least-squares modeling (DVLS), and quantitative texture analysis approaches to the structural refinement of Garfield nontronite. *Phys. Chem. Miner.* **25**, 347–365.

- Manceau A., Schlegel M. L., Musso M., Sole V. A., Gauthier C., Petit P. E., and Trolard F. (2000) Crystal chemistry of trace elements in natural and synthetic goethite. *Geochim. Cosmochim. Acta* **64**, 3643–3662.
- Marezio M. (1965) Crystal structure and anomalous dispersion of γ -LiAlO₂. *Acta Cryst.* **19**, 396–400.
- Martell A. E., Smith R. M., Motekaitis R. J. (1998) NIST Critically Selected Stability Constants of Metal Complexes Database. Version 5.0, Users' Guide. U.S. Department of Commerce, NIST, Gaithersburg, MD.
- Mayer T. D. and Jarrell W. M. (1996) Formation and stability of iron(II) oxidation products under natural concentrations of dissolved silica. *Water Res.* **30**, 1208–1214.
- McPhail M., Page A. L., and Bingham F. T. (1972) Adsorption interactions of monosilicic and boric acid on hydrous oxides of iron and aluminum. *Soil. Sci. Am. Proc.* **36**, 510–514.
- Meng X., Bang S., and Korfiatis G. P. (2000) Effects of silicate, sulfate, and carbonate on arsenic removal by ferric chloride. *Water Res.* **34**, 1255–1261.
- Merbach A. E. and Akitt J. W. (1990) High-resolution variable pressure NMR for chemical kinetics. *NMR Basic Principles Progr.* **24**, 190–232.
- Michot L. J., Montargès-Pelletier E., Lartiges B. S., d'Espinose de la Caillière J.-B., and Briois V. (2000) Formation mechanism of the Ga₁₃ Keggin ion: A combined EXAFS and NMR study. *J. Am. Chem. Soc.* **122**, 6048–6056.
- Olson L. L. and O'Melia C. R. (1973) The interactions of Fe(III) with Si(OH)₄. *J. Inorg. Nucl. Chem.* **35**, 1977–1985.
- Peterson M.L., Brown G.E. Jr., Parks G.A., and Stein C.L. (1997) Differential redox and sorption of Cr(III/VI) on natural silicate and oxide minerals: EXAFS and XANES results. *Geochim. Cosmochim. Acta* **61**, 3399–3412.
- Pokrovski G. S. and Schott J. (1998) Experimental study of the complexation of silicon and germanium with aqueous organic species. Implications for Ge and Si transport and the Ge/Si ratio in natural waters. *Geochim. Cosmochim. Acta* **62**, 3413–3428.
- Pokrovski G. S., Martin F., Hazemann J.-L., and Schott J. (2000) An X-ray absorption fine structure spectroscopy study of germanium-organic ligand complexes in aqueous solution. *Chem. Geol.* **163**, 151–165.
- Pokrovski G. S., Kara S., and Roux J. (2002a) Stability and solubility of arsenopyrite, FeAsS, in crustal fluids. *Geochim. Cosmochim. Acta* **66**, 2361–2378.
- Pokrovski G. S., Schott J., Hazemann J.-L., Farges F., and Pokrovsky O. S. (2002b) An X-ray Absorption Fine Structure (XAFS) and Nuclear Magnetic Resonance (NMR) spectroscopy study of gallium-silica complexes in aqueous solution. *Geochim. Cosmochim. Acta* **66**, 4203–4322.
- Pokrovsky O. S. and Schott J. (2002) Iron colloids/organic matter associated transport of major and trace elements in small boreal rivers and their estuaries (NW Russia). *Chem. Geol.* **190**, 141–179.
- Press W. H., Flannery B. P., Teukolsky S. A., and Vetterling W. T. (1986) *Numerical Recipes—The Art of Scientific Computing*. Cambridge University Press, Cambridge, UK.
- Reardon E. J. (1979) Complexing of silica by iron(III) in natural waters. *Chem. Geol.* **25**, 339–345.
- Rose J., Manceau A., Bottero J. Y., Masion A., and Garcia F. (1996) Nucleation and growth mechanisms of Fe oxyhydroxide in the presence of PO₄ ions. 1. Fe K-edge EXAFS study. *Langmuir* **12**, 6701–6707.
- Rose J., Flank A. M., Masion A., Bottero J. Y., and Elmerich P. (1997a) Nucleation and growth mechanisms of Fe oxyhydroxide in the presence of PO₄ ions. 2. P K-edge EXAFS study. *Langmuir* **13**, 1827–1834.
- Rose J., Manceau A., Masion A., and Bottero J. Y. (1997b) Structure and mechanisms of formation of FeOOH(NO₃) oligomers in the early stages of hydrolysis. *Langmuir* **13**, 3240–3246.
- Santelli C. M., Welch S. A., Westrich H. R., and Banfield J. F. (2001) The effect of Fe-oxidizing bacteria on Fe-silicate mineral dissolution. *Chem. Geol.* **180**, 99–115.
- Schenk J. E., Weber W. J., Jr. (1968) Chemical interactions of dissolved silica with iron(II) and iron(III). *J. Am. Water Works Assoc.* **60**:199–212.
- Schwertmann U. and Thalmann H. (1976) The influence of Fe(II), Si, and pH on the formation of lepidocrocite and ferrihydrite during oxidation of aqueous FeCl₂ solutions. *Clay Miner.* **11**, 189–200.
- Schwertmann U., Taylor R. M. (1989) Iron oxides. In *Minerals in Soil Environments*. (eds. J. B. Dixon and S. B. Weed). SSSA, Madison, Wisconsin.
- Sherman D. M. and Vergo N. (1988) Optical (diffuse reflectance) and Mössbauer spectroscopic study of nontronite and related Fe-bearing smectites. *Am. Miner.* **73**, 1346–1354.
- Sigg L. and Stumm W. (1980) The interaction of anions and weak acids with the hydrous goethite (α -FeOOH) surface. *Colloids Surf.* **2**, 101–117.
- Soldo Y., Hazemann J. L., Aberdam D., Inui M., Tamura K., Raoux D., Pernot E., Jal J. F., and Dupuy-Philon J. (1998) Semiconductor-to-metal transition in fluid selenium at high pressure and temperature: An investigation using X-ray absorption spectroscopy. *Phys. Rev. B: Condens. Matter* **57**, 258–268.
- Stern E. A. (1993) Number of relevant independent points in x-ray-absorption fine-structure spectra. *Phys. Rev. B: Condens. Matter* **48**, 9825–9827.
- Swaddle T. W. (2001) Silicate complexes of aluminum (III) in aqueous systems. *Coord. Chem. Rev.* **219–221**, 665–686.
- Szytula A., Burewicz A., Dimitrijevic Z., Krasnicki S., Rzany H., Todorovic J., Wanic A., and Wolski W. (1968) Neutron diffraction studies of α -FeOOH. *Phys. Status Solidi* **26**, 429–434.
- Tagirov B. R., Diakonov I. I., Devina O. A., and Zotov A. V. (2000) Standard ferric-ferrous potential and stability of FeCl²⁺ to 90°C. Thermodynamic properties of Fe³⁺_(aq) and ferric-chloride species. *Chem. Geol.* **162**, 193–219.
- Teo B. K. (1986) *EXAFS: Basic Principles and Data Analysis*. Inorganic Chemistry Concepts 9. Springer-Verlag, Berlin, Germany.
- Tessier A., Rapin F., and Carignan R. (1985) Trace metals in oxic lake sediments: possible adsorption onto iron oxyhydroxides. *Geochim. Cosmochim. Acta* **49**, 183–194.
- Toth T. A. and Fritz S. J. (1997) An Fe-berthierine from a cretaceous laterite. Parts I and II. *Clays Clay Miner.* **45**, 564–586.
- Vempati R. K. and Loeppert R. H. (1989) Influence of structural and adsorbed Si on the transformation of synthetic ferrihydrite. *Clays Clay Miner.* **37**, 273–279.
- Waychunas G. A. and Rossman G. R. (1983) Spectroscopic standard for tetrahedrally coordinated ferric iron: γ -LiAlO₂:Fe³⁺. *Phys. Chem. Miner.* **9**, 212–215.
- Waychunas G. A., Rea B. A., Fuller C. C., and Davis J. A. (1993) Surface chemistry of ferrihydrite: Part I. EXAFS studies of the geometry of coprecipitated and adsorbed arsenate. *Geochim. Cosmochim. Acta* **57**, 2251–2269.
- Waychunas G. A., Davis J. A., and Fuller C. C. (1995) Geometry of sorbed arsenate on ferrihydrite and crystalline FeOOH: Re-evaluation of EXAFS results and topological factors in predicting sorbate geometry, and evidence for monodentate complexes. *Geochim. Cosmochim. Acta* **59**, 3655–3661.
- Weber W. and Stumm W. (1965) Formation of a silicate-iron (III) complex in dilute aqueous solutions. *J. Inorg. Nucl. Chem.* **27**, 237–239.
- Wesolowski D. J., Palmer D. A., and Mesmer R. E. (1995) Measurements and control of pH in hydrothermal solutions. In *Water-Rock Interaction* (eds. Y. K. Kharaka and O. V. Chudakov), pp. 51–55. Balkema, Rotterdam, the Netherlands.
- Wilke M., Farges F., Petit P. E., Brown G. E. Jr., and Martin F. (2001) Oxidation state and coordination of Fe in minerals: An Fe K-XANES spectroscopic study. *Am. Mineral.* **65**, 713–730.
- Winterer M. (1997) XAFS—A data analysis program for material sciences [abstract]. In *Proceedings of the 9th International Conference on X-Ray Absorption Fine Structure (XAFS-IX): Grenoble, 1996* (eds. J. Goulon, C. Goulon-Ginet, and N. B. Brookes). J. Physique IV 7:C2-243.