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Concentration of Ions in the Topside Ionosphere as Measured onboard the *DEMETER* Satellite: Morphology and Dependence on Solar and Geomagnetic Activity

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Abstract—Variations in concentration of ions H^+ , He^+ , and O^+ are studied at a height of about 700 km using the data of continuous observations onboard the *DEMETER* satellite at the decline and in the minimum of solar activity from 2004 to 2008. Latitudinal distributions, seasonal behavior, and irregular variations in ion concentrations and their dependence on solar and geomagnetic activity are considered. Within this altitude range, for the first time an analysis is performed of the dataset of many-year continuous observations in both hemispheres from the equatorial to subauroral latitudes. This made it possible to describe the seasonal and irregular variations of the concentrations of main ion species with better time and spatial resolution than in the available empirical models. The dependence of concentrations of three types of ions on solar and geomagnetic activity is studied at time scales from several days to several years, and it is shown that the anti-phase change in concentrations of O^+ and light ions known from publications is partly a result of mutual dependence of solar and geomagnetic activity and is observed only at time scales beginning from several months. At time scales from several days to several weeks, variations in the concentration of O^+ and light ions are governed mainly by solar and geomagnetic activity, respectively.

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1. INTRODUCTION

The ionospheric plasma above the electron density maximum is intermediate from the ionospheric plasma where heavy ions prevail, first of all, O^+ , to the plasma typical for the plasmasphere, with domination of protons. At the transition of the oxygen-dominant to proton plasma, properties of the ionosphere change in a substantial way, so the height where the concentrations of O^+ and light ions (H^+ and He^+) become equal, called the ionospheric upper transition height, is an important parameter of ionospheric models (see, for example, [1]). The distributions of the thermal plasma density in the topside ionosphere are measured both onboard satellites [2–4] and with the help of radars. Results and constraints of early radar observations are presented in [5]. Depending on the height and inclination of the orbit, satellite measurements in the topside ionosphere provide either a vertical profile within a limited latitudinal and longitudinal range, or measurements almost at one and the same height within a wide range of latitudes. The first conclusions on the composition and main plasma parameters in the topside ion-

osphere were obtained in the early radar measurements [5] and onboard high-apogee satellites [6] which have measured vertical profiles of ion concentrations and established that the ionospheric upper transition height lies at various values of local time, latitude, and solar activity level at heights from 500 to 2000 km. Already in the first publications with analysis of the satellite data [7], the main features of vertical profiles of concentrations of main ion species were obtained, and the dependence of concentrations of heavy ions (direct) and light ions (inverse) on solar and geomagnetic activity was found. The analysis of measurements onboard the *OGO* series satellites [8] revealed a complicated character of the latitudinal dependence of concentrations in the topside ionosphere and strong longitudinal variation which are governed by a combination of the neutral wind and the $[E \times B]$ drift. At low latitudes, the influence of tidal effects is also important [9]. These factors and strong variability of ion concentrations with time and in space lead to difficulties in creation of a reliable empirical model of plasma in the topside ionosphere. For example, even the recent versions of the IRI model [1] differ substantially from the data of mea-

[†] Deceased.

measurements of light ion concentrations in the topside ionosphere, and values of ion concentrations in the topside ionosphere for different models differ several-fold. In a series of publications [10–13], the empirical model TTS (Truhlik, Triskova, Smilauer) for vertical profiles of relative concentrations of the main ion species and their variations in the solar activity cycle was proposed. The transition to absolute values was performed using the data on the total electron content. The model height of the upper ionospheric transition reaches 2000 km in the conditions of solar activity maximum and lies in the interval from 700 to 1100 km for solar minimum. The allowance for the seasonal variation is performed by splitting into three seasons centered on the June and December solstices and on equinoxes (without separation of the March and September ones). As a result, the latitudinal distribution of ion concentrations in the equinoxes turns out to be symmetrical about the equator, unlike the measurement data. It was shown in [13] that the H^+ concentration decreases with increasing geomagnetic activity described by the summated diurnal K_p index, but then the authors performed no separation between the influence of solar and geomagnetic activity. The only input parameter of the model taking into account factors of space weather is the solar flux index $F_{10.7}$. At some altitudes, the IRI and TTS models differ by an order of magnitude in estimation of absolute values of ion concentrations and relative role of H^+ and He^+ ions. According to [14], the role of He^+ becomes important and even dominating under high solar activity, while under low activity the main components of the ionospheric plasma above the electron concentration maximum are O^+ and H^+ [6].

The data of measurements onboard satellites of the *AE (Atmospheric Explorer)* series were compared in [15] with calculations by the model FLIP (Field Line Interhemispheric Plasma) [16] for four ion species (H^+ , He^+ , N^+ , and O^+). The experimental data were grouped in two seasons (summer and winter), in near-noon and near-midnight intervals of local time, and in two latitudinal bands (40° – 50° and 50° – 60° N). Since in that paper the data were analyzed within relatively short time interval, the dependence on solar and geomagnetic activity was not considered. The data were averaged also for two hemispheres with allowance made for the season. The noon ion and electron temperatures were fitting parameters of the model. The model and measured concentrations of O^+ and H^+ ions agree well at all altitudes, while the values for He^+ ions agree at heights up to 500 km. Thus, the FLIP model describes well enough large-scale variations in ion concentrations. As for description of the density variations at smaller spatial scales and times, it is a complicated task not only for theoretical description, but for construction of an empirical model as well.

In order to construct an empirical model of ion concentrations in the topside ionosphere describing all main features of latitudinal distributions, seasonal

variations, and dependencies on space weather factors, one needs long-period continuous measurements on low-apogee satellites which are moving for the most part of time near the ionospheric upper transition height. But such measurements are relatively scarce. In [17] seasonal variations, dependencies on latitude, and changes in solar activity cycle of concentrations of the main ion species at a height of about 600 km at latitudes of $\pm 34^\circ$ over the period 1994–2001 (that is, approximately over a half of solar activity cycle from minimum to maximum) are studied. At relatively high solar activity, the concentrations of H^+ and He^+ ions were too low for detailed analysis, but high enough for estimation of the diurnal variation. During the period under study, the mean value of the H^+ concentration at the nighttime side was 10^4 cm^{-3} , which is an order of magnitude higher than the mean value at the daytime side. A high (0.83) correlation coefficient between the concentration of O^+ and solar activity index $F_{10.7}$ was obtained, as well as three-fold increase in the daytime concentration from 1995 to 2000 and inter-hemisphere asymmetry with the O^+ concentration maximum in the equinox periods at 15° N. The seasonal variations and longitudinal dependencies of the H^+ and O^+ concentrations at a height of 1100 km were considered in [3] for two years of observations at high solar activity in 1978–1980 at latitudes from 70° S to 70° N. For approximation of the seasonal behavior of the O^+ concentration we have used a harmonic function with one annual harmonics and amplitude depending linearly on solar activity level. In other words, in the final formula, only the most general features of the seasonal behavior and latitudinal dependence are taken into account. Thus, the available empirical models agree well in estimation of vertical profiles and of diurnal behavior of the O^+ concentration averaged over large intervals of time and latitude. The agreement is slightly worse for light ions (especially for He^+). But the available data are not sufficient for a description of details of the latitudinal distribution and seasonal behavior. In the problem on the dependence of ion concentration on external factors, first of all, on solar and geomagnetic activity, there are also unsolved problems. Though the existence of these dependencies was discovered more than 40 years ago, still no separation of the influence of solar and geomagnetic activity on concentration of each of ion species has been performed.

In order to solve these problems, in this paper the data of the *DEMETER* satellite are used. The measurements of the three ion species (H^+ , He^+ , and O^+) are continuously conducted onboard the satellite at a height of ~ 700 km at latitudes from 60° S to 60° N since 2004 [18]. In this paper, the main morphological properties, including the seasonal behavior and latitudinal distribution of concentrations of H^+ , He^+ , and O^+ are studied, and dependencies of the concentrations on solar and geomagnetic activity are analyzed. As far as the data set available for processing covers

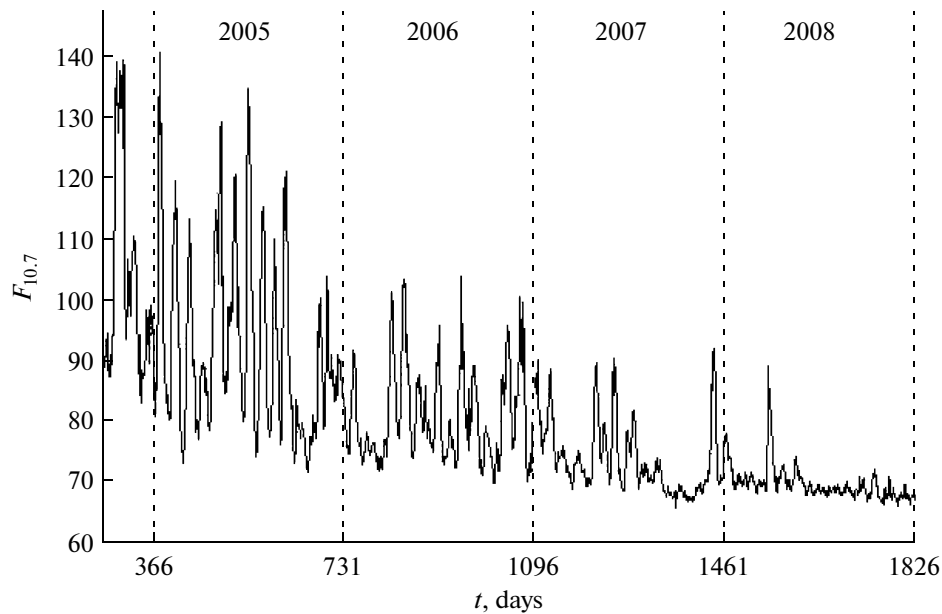


Fig. 1. Variations of the solar flux index $F_{10.7}$ in 2004–2008.

more than four years of continuous observations at approximately the same height, the ion concentration variations are studied with the better time resolution than in the previous publications. This makes it possible to reveal additional morphological features and to separate the influence of solar radiation and geomagnetic disturbances on concentrations of particular ion components.

In this paper, the results integrated over longitude are presented. A large number of papers has been dedicated to consideration of the longitudinal variation (see, for example, [9]). Ion concentrations at various longitudes in the topside ionosphere can differ several-fold due to the combination of the neutral wind, $[\mathbf{E} \times \mathbf{B}]$ drift, and tidal effects. Studies of longitudinal variation is beyond the scope of this paper, but using its results, one has to bear in mind the error related to averaging over longitude.

2. OBSERVATIONS AND DATA PROCESSING

Concentrations of the main components of plasma are measured onboard the *DEMETER* satellite continuously since 2004 (see, for example, [19]). The *DEMETER* satellite was launched on June 29, 2004 to a circular solar-synchronous orbit with an inclination of 98° and the altitude range 710–730 km. At the nighttime and daytime sides, the satellite orbit falls on the local time intervals 22:00–24:00 LT and 10:00–12:00 LT, respectively. Concentrations of three ion species (H^+ , He^+ , and O^+) are measured by the electrostatic thermal plasma analyzer developed in CETP [18]. The concentration of the main component O^+ is measured within the range 10^2 – $5 \cdot 10^5 \text{ cm}^{-3}$ with an accuracy of 5%. The concentration of light ions was

measured above the threshold value taken as 2% of the O^+ concentration. The early measurements onboard the *Ariel-I* satellite showed that isolines of the O^+ concentration lie approximately along the geomagnetic parallels [20]. In this paper, we also use the corrected geomagnetic coordinate system (CGM). The data for daytime and nighttime segments of the orbit are analyzed separately and the satellite orbit is split in geomagnetic latitude into bands 10° wide. The preliminary analysis includes selection of intervals with satisfactory quality of data and averaging over time and latitude. In order to weaken the impact of rare peak values, logarithms of concentrations are analyzed.

3. RESULTS OF DATA PROCESSING

3.1. Variations of Ion Concentrations within the 11-year Solar Activity Cycle

During the observations (2004–2008), solar activity decreased (Fig. 1). Variations of the H^+ , He^+ , and O^+ concentrations are presented in Figs. 2, 3, and 4, respectively. Each panel shows results for one latitudinal band 10° wide. The concentration values were averaged in a sliding 30-day window. The results for daytime and nighttime parts of the orbit are shown in the left-hand and right-hand panels, respectively.

The concentration of light ions grows with decreasing solar activity. At all latitudes at the nighttime side, stronger dependence of the H^+ concentration on solar activity than at the daytime side is observed. For example, within the latitudinal band 30° – 60° N, the mean concentration of H^+ in summer is approximately $4.5 \cdot 10^2 \text{ cm}^{-3}$ in 2005, while it increases up to $1.3 \cdot 10^3 \text{ cm}^{-3}$ in 2008. In the same months, i.e., in local winter, in the

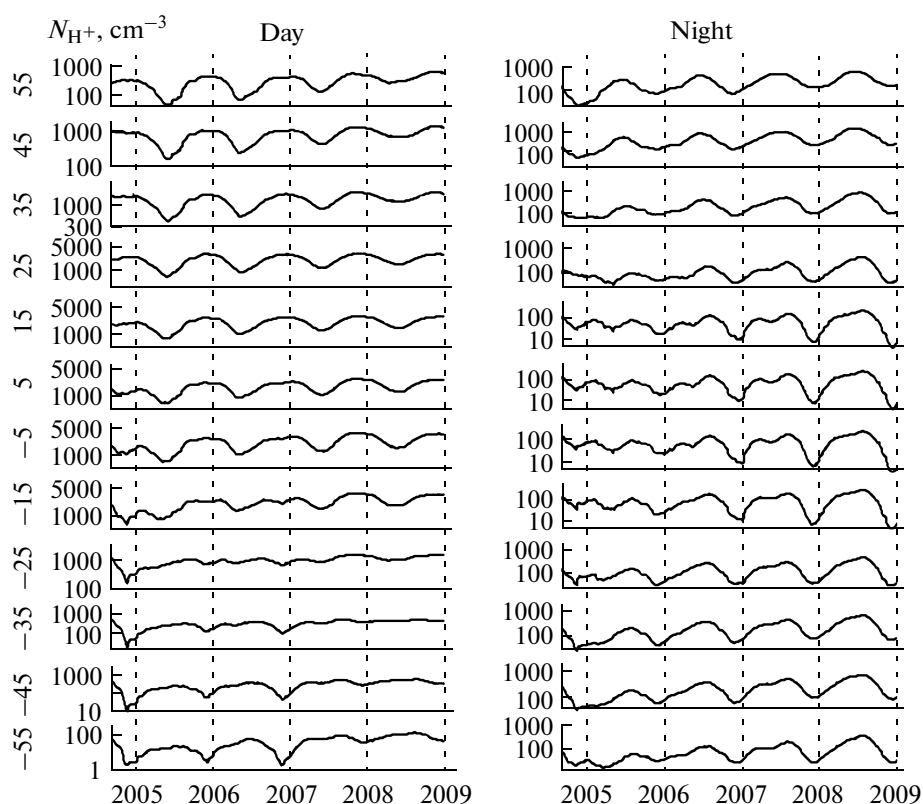


Fig. 2. Variations of the H^+ concentration in 2004–2008. The central latitude for each band is shown to the left from the ordinates in the left-hand panel.

Southern Hemisphere the H^+ concentration in 2008 is by a factor of 3 higher than in 2005: $8.4 \cdot 10^2$ and $2.9 \cdot 10^2 \text{ cm}^{-3}$, respectively. On the daytime side, the H^+ concentration also increases with decreasing solar activity. But the dependence of the concentration on solar activity is weaker: the maximal daytime concentration of H^+ in 2008 is approximately by a factor of two higher than in 2005. The variation in the He^+ concentration within the 11-year solar activity cycle (Fig. 3) coincides qualitatively with the variation in the H^+ concentration at all latitudes both at the daytime and nighttime sides, but the difference in the He^+ concentration in the period of minimal (2008) and moderate (2005) solar activity is less than for H^+ .

Unlike the concentrations of light ions, the O^+ concentration decreases in the period from 2004 to 2008 together with solar activity (Fig. 4). Depending on latitude, the O^+ concentration decreases from 2005 to 2008 by a factor of 3–5. Within the 30° – 60° N latitudinal band, the O^+ concentration was $1.6 \cdot 10^4$ and $5 \cdot 10^3 \text{ cm}^{-3}$ in 2005 and 2008, respectively. At low latitudes, the concentration decrease is even more substantial: from $2.0 \cdot 10^4 \text{ cm}^{-3}$ in 2005 to $5 \cdot 10^3 \text{ cm}^{-3}$ in 2008.

Thus, long-term (on time scales of several years) variations in the light ion concentrations show an anticorrelation with variations in the O^+ concentration

and solar activity. After the analysis of seasonal behavior of ion concentrations in the next subsection, we will return to correlations of ion concentrations with solar activity at shorter time scales for revealing physical factors governing ion concentrations in the topside ionosphere.

3.2. Seasonal Behavior at Various Latitudes and in Local Time Intervals

In addition to variations within solar activity cycle, the concentration of each type of ions depends systematically on latitude and local time of the year. Since within the 11-year cycle of solar activity, only annual mean values of the concentrations change substantially, while the relative seasonal variations in concentrations stay almost unchangeable, it is convenient to analyze the seasonal variations in the concentration logarithms averaged over the observational period.

The seasonal behavior of logarithms of concentrations of three ion species averaged over all years of observations is shown in Fig. 5 for the latitudinal interval 40° – 50° both in the Northern and Southern Hemispheres. The seasonal behavior differs substantially for light ions and O^+ . The main qualitative features of the seasonal behavior of concentration of three ion species are presented in Table 1. The table

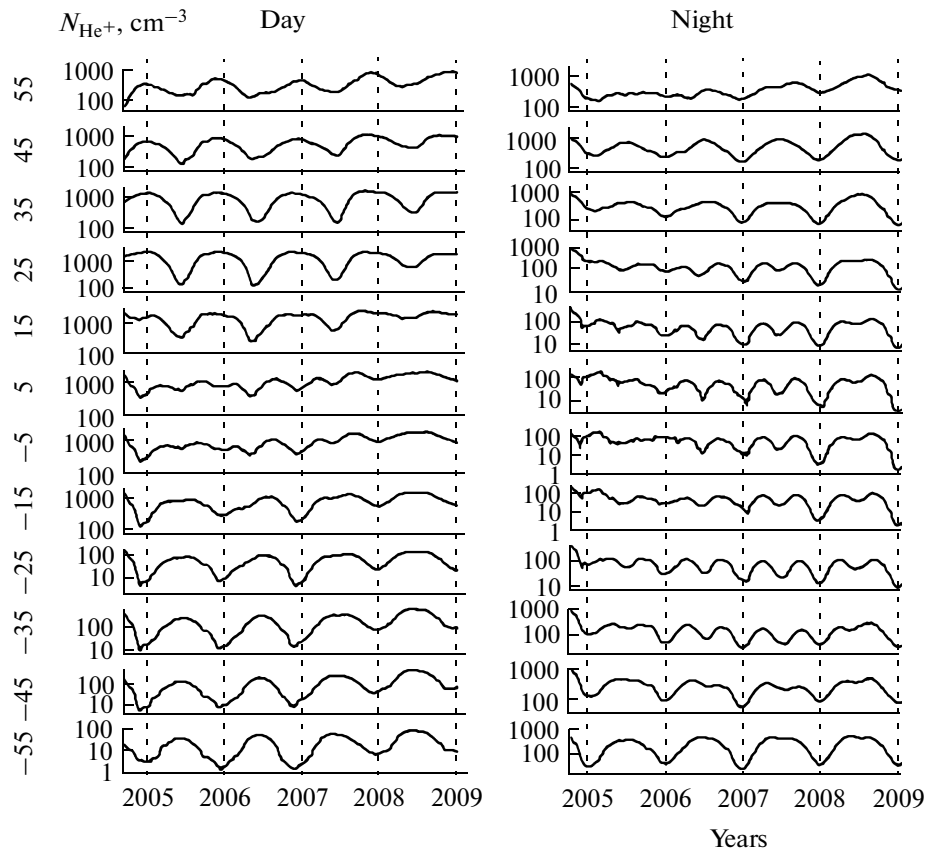


Fig. 3. Variations of the He^+ concentration. The figure format is the same as in Fig. 2.

presents the position of the yearly maximum concentration T_{\max} , maximal concentration N_{\max} , and ratio of the maximal and minimal yearly concentrations N_{\max}/N_{\min} . The Day/Night and N/S indices correspond to the daytime and nighttime sides and to the Northern and Southern Hemispheres, respectively. For all three types of ions at the daytime side in the Northern Hemisphere, larger values of the maximal yearly concentration, N_{\max} are observed, while in the Southern Hemisphere, larger relative amplitudes of the seasonal behavior N_{\max}/N_{\min} are observed. The main differences in the seasonal behavior between O^+ and light ions are as follows.

—seasonal behavior in its classical form is observed only for O^+ , while for light ions, in addition to season, global factors play an important role and this fact leads to different forms of the seasonal behavior curve in two hemispheres relative to local time of the year at the daytime side and to the existence of a global maximum of concentration in October on the nighttime side;

—a sharp inter-hemisphere asymmetry is observed in the latitudinal distribution of light ion concentrations: the maximal yearly concentration at the daytime side in the Northern Hemisphere is an order of magnitude higher than in the Southern Hemisphere, while

for O^+ the ratio of maximal yearly concentrations $N_{\text{O}, \max, \text{N}}/N_{\text{O}, \max, \text{S}} \approx 2$.

Seasonal variations in the ion concentrations in the near-equatorial region are shown in Fig. 6 in the same format as in Fig. 5. For all three ion species, the difference in the seasonal behavior between the hemispheres at low latitudes vanishes, but the seasonal behavior itself remains substantial, and some features appear in it, different from those observed at middle latitudes. The main features of the seasonal behavior in the near-equatorial region are presented in Table 2 (designations are the same as in Table 1). Since at the near-equatorial latitudes, the seasonal behavior in both hemisphere almost coincides, the Table 2 shows the parameter values averaged over the band $-10^\circ < \Phi < 10^\circ$. For all ion components at the daytime side, larger concentrations are observed, while the relative amplitude of the seasonal behavior N_{\max}/N_{\min} is higher at the night side. At the dayside where the relative amplitude of the seasonal behavior is small, the positions of maxima and minima in the seasonal behavior curve are different for all three types of ions. At the night side, there exist the following qualitative differences in the seasonal behavior of the concentrations of O^+ and light ions:

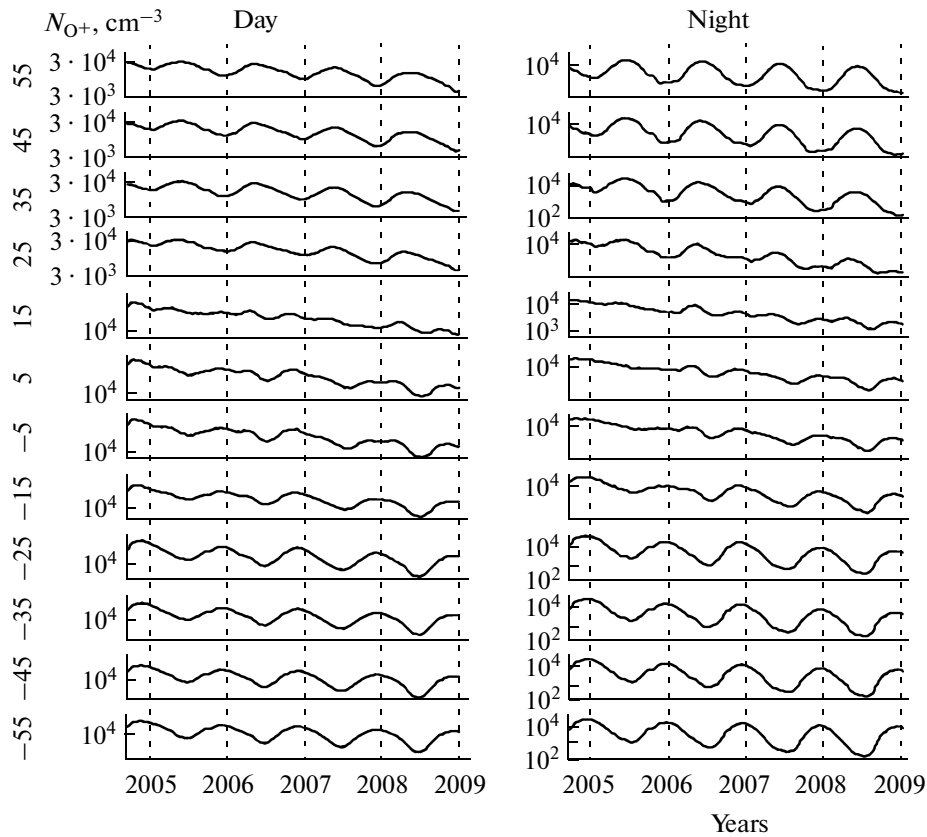


Fig. 4. Variations of the O^+ concentration. The figure format is the same as in Fig. 2.

– seasonal behavior curves for H^+ and He^+ have two maxima in April–May and in October, while the curve for O^+ has the only maximum in February;

– the relative amplitude of the seasonal behavior, N_{\max}/N_{\min} is substantially higher for light ions than for O^+ .

Thus, the seasonal behavior of the O^+ concentration is governed by the local time of the year, while global factors are important (and even dominating for some intervals of latitudes and local times) for variations in the concentrations of light ions during the year. In order to reveal the parameters governing changes in the concentrations of O^+ and light ions, in the next subsection we will study the dependencies of ion concentrations at the dayside separately from solar and geomagnetic activity on the time scales from several days to several weeks.

3.3. Variations in Ion Concentrations on Time Scales from Several Days to Several Weeks. Dependencies on Solar and Geomagnetic Activity

The above consideration has shown that the dependencies of ion concentrations in the topside ionosphere on time of the year and solar activity at time scales from several months to several years differ substantially for O^+ and light ions. The seasonal variation in its classical form exists only for the O^+ concentra-

tion. The summer yearly maximum in the concentration and positive correlation with solar activity within the 11-year cycle show that the O^+ concentration is governed mainly by solar radiation. Since on large time scales, the correlation of ion concentration with solar activity can be an artifact of the dependence on other parameters also correlating with solar activity, we chose for more detailed study one year, 2005, when the 27-day cycle related to the rotation of the Sun is distinctly seen in solar indices as it usually takes place at the decline phase of the 11-year solar cycle. Figure 7 shows the results of comparison of variations in the O^+ concentration and solar activity during nine months from January to September 2005. The top panel shows logarithm of the O^+ concentration at the daytime segments of the orbit. One point in the figure corresponds to a daily mean value of the concentration logarithm. The middle panel shows the daily mean values of the solar radiation flux $F_{10.7}$. In January–June, 5 maxima in the $F_{10.7}$ index with a distinct 27-day period are seen. After the phase break, three more maxima with the same period (the first of them falls on $t = 132$ days) are observed. The three last cycles in August–September have a changed period $T = 22$ days. In the bottom panel, the running 30-point correlation coefficient $C_{F_{10.7}-O^+}$ between the logarithm of the O^+ concentra-

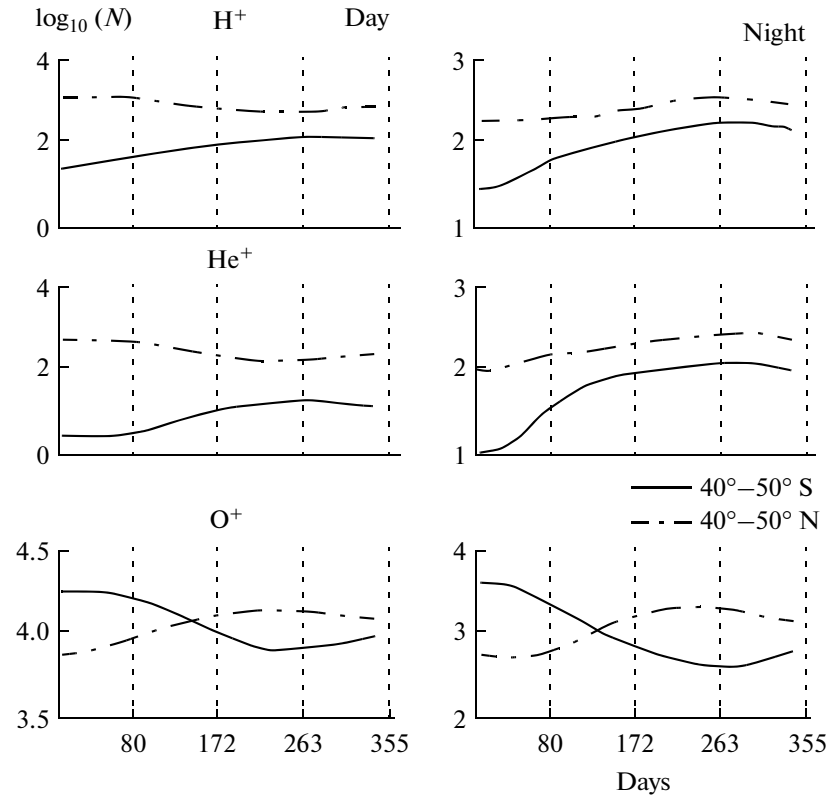


Fig. 5. Seasonal behavior of logarithm of the concentrations of three ion species at middle latitudes averaged over the entire observational period. The vertical dashed lines show the positions of equinoxes and solstices.

tion and solar activity index $F_{10.7}$ is shown. We remind that one point corresponds to the value of the concentration logarithm averaged over one day at the daytime segment of the orbit, that is, the 30-point correlation coefficient at the absence of gaps in the data corresponds to a month of observations. The correlation coefficient $C_{F_{10.7}-O^+}$ is positive everywhere and for the majority of intervals exceeds 0.5, and the positions of almost all maxima in the $F_{10.7}$ index and O^+ concentration coincide with an accuracy of one day. Note that the coincidence of positions of maxima is observed at changing phase and period. The value of the correlation coefficient averaged over nine months $\langle C_{F_{10.7}-O^+} \rangle = 0.69$. Figure 7 shows values of the concentration logarithms for one latitudinal band, but similar picture is observed along the entire profile: at all latitudes the positions of the maxima in O^+ concentration are constant and the correlation coefficient $C_{F_{10.7}-O^+}$ is substantially positive.

Figure 8 shows the results for the same time interval for the H^+ concentration within the latitudinal band $50^\circ-60^\circ$ N. Logarithm of the H^+ concentration and the solar activity index $F_{10.7}$ are shown in the top and middle panels, respectively. The bottom panel shows the running 30-point correlation coefficient $C_{F_{10.7}-H^+}$ between the $F_{10.7}$ index and logarithm of the H^+ concentration. On the average, the concentration of H^+

anti-correlates with the O^+ concentration and solar activity. The value of the correlation coefficient averaged over nine months $\langle C_{F_{10.7}-H^+} \rangle = -0.51$. Unlike the case of O^+ , the minima in $\log(N_{H^+})$ are often shifted by several days relative to the solar activity maxima. This can be caused by the fact that the concentration of H^+ depends on solar activity not directly, but through geo-

Table 1. Main features of the seasonal behavior at middle latitudes

	O^+	H^+	He^+
$T_{\max, N, Day}$, months	8	1–2	1–2
$T_{\max, S, Day}$, months	2	10	10
$T_{\max, N, Night}$, months	8	10	10
$T_{\max, S, Night}$, months	2	10	10
$N_{\max, N, Day}$, cm^{-3}	$1.8 \cdot 10^4$	10^3	560
$N_{\max, S, Day}$, cm^{-3}	$1.4 \cdot 10^4$	125	20
$N_{\max, N, Night}$, cm^{-3}	$4 \cdot 10^3$	350	260
$N_{\max, S, Night}$, cm^{-3}	$2.1 \cdot 10^3$	190	130
$N_{\max, N, Day}/N_{\min, N, Day}$	1.7	2	3.5
$N_{\max, S, Day}/N_{\min, S, Day}$	2.3	6	8.5
$N_{\max, N, Night}/N_{\min, N, Night}$	4	1.9	1.6
$N_{\max, S, Night}/N_{\min, S, Night}$	10	7	11

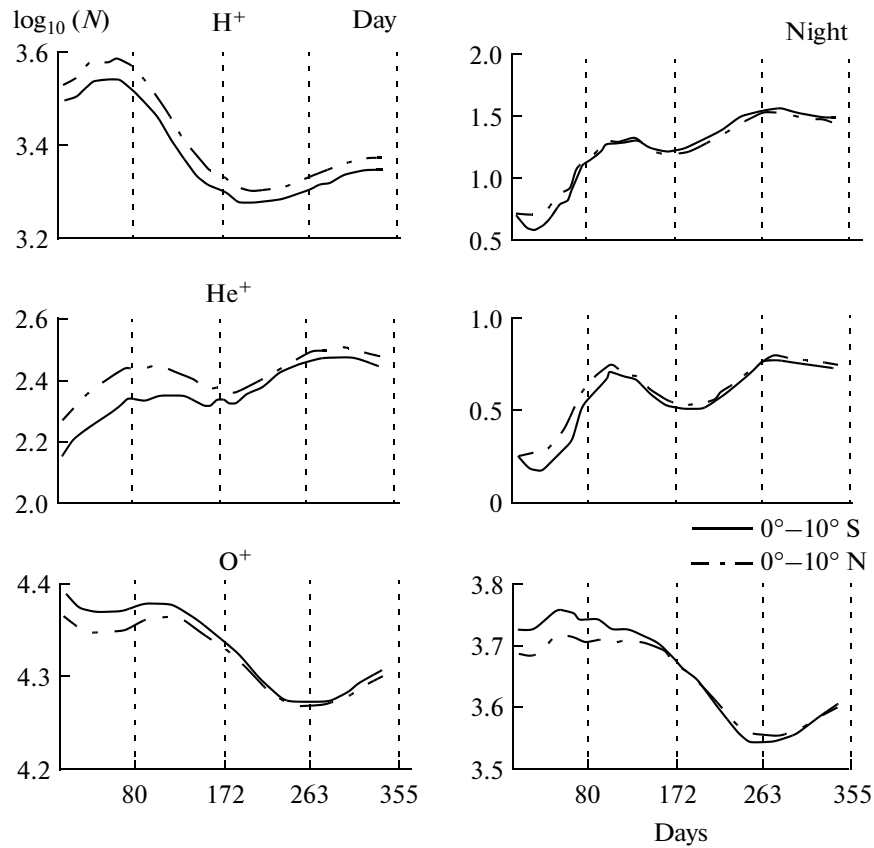


Fig. 6. Seasonal behavior of logarithm of the concentrations of three ion species at low latitudes averaged over the entire observational period. The figure format is the same as in Fig. 5.

magnetic activity. In order to check this assumption, we compare variations in the H^+ concentration and geomagnetic activity described by the D_{st} index (we remind that the negative values of the index correspond to high geomagnetic activity).

Logarithm of the H^+ concentration within the latitudinal band 50° – 60° N and geomagnetic activity index D_{st} are shown in Fig. 9. Logarithm of the H^+ concentration is shown in the top panel, the D_{st} index

is shown in the middle panel, and the bottom panel shows the running 30-point correlation coefficient between D_{st} and $\lg(N_{H^+}) - C_{D_{st}-H^+}$. The correlation coefficient averaged over nine months $\langle C_{D_{st}-H^+} \rangle = 0.71$ and the positions of all pronounced minima in D_{st} agree with a good accuracy with the positions of minima in the H^+ concentration. At lower latitudes the absolute value of the correlation coefficient decreases, but the coincidence of the minima positions remains good. Low values of the correlation coefficient $C_{D_{st}-H^+}$ are observed during approximately a month from July 10, 2005 to August 10, 2005, when geomagnetic activity was low and there was no magnetic storms with $D_{st} < -100$ nT. Note that at the same time, the absolute values of the correlation coefficient $C_{F10.7-H^+}$ exceed the mean values over the entire observational period, that is, the anti-correlation between the H^+ concentration and solar activity known from publications reveals itself. For the remaining intervals and on the average over the period considered, geomagnetic activity influences the H^+ concentration stronger than solar activity.

Table 2. Main features of the seasonal behavior at low latitudes

	O^+	H^+	He^+
$T_{\max, Day}$, months	1 and 4	2–3	3–4 and 11–12
$T_{\max, Night}$, months	2	4–5 and 10	4–5 and 10
$N_{\max, Day}$, cm^{-3}	$2.5 \cdot 10^4$	$3.8 \cdot 10^3$	310
$N_{\max, Night}$, cm^{-3}	$5.7 \cdot 10^3$	35	6
$N_{\max, Day}/N_{\min, Day}$	1.4	1.9	1.7
$N_{\max, Night}/N_{\min, Night}$	1.7	9	4

4. DISCUSSION AND CONCLUSIONS

The performed analysis of the measurements onboard the *DEMETE*R satellite made it possible to

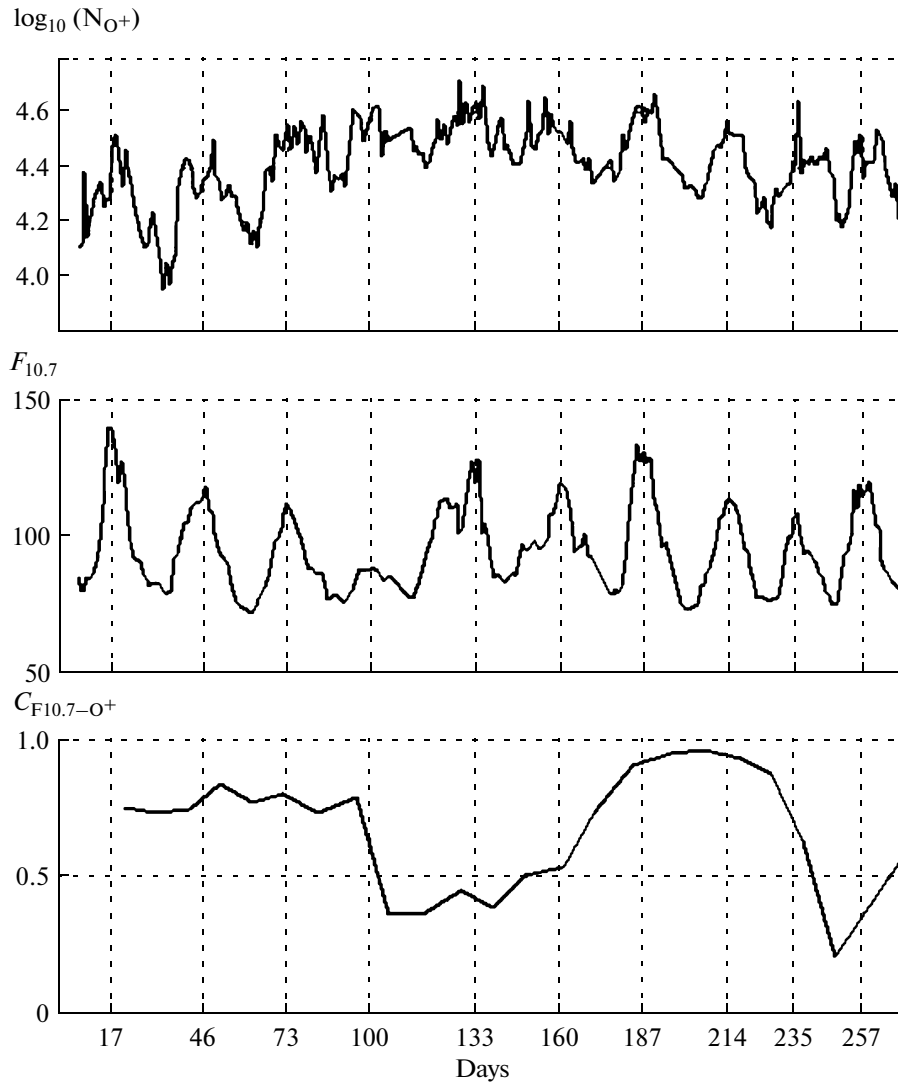


Fig. 7. Logarithm of the O^+ concentration within the latitudinal band 20° – 30° N, solar activity index $F_{10.7}$, and running 30-point correlation coefficient $C_{F10.7-O^+}$ between the $F_{10.7}$ index and logarithm of the concentration. Positions of $F_{10.7}$ index maxima are shown by thin dashed vertical lines.

reveal new qualitative features in the seasonal behavior and latitudinal distribution of ion concentrations in the topside ionosphere relative to the IRI model [1].

Variations of the concentrations of H^+ , He^+ , and O^+ ions within the 11-year cycle of solar activity agree with the variations predicted by the IRI [1] and TTS [10] models: at decreasing solar activity, the O^+ concentration decreases, while the concentrations of light ions increase.

Even qualitatively, the seasonal behavior obtained on the basis of the *DEMETER* measurements coincides with the IRI results only partially. The most substantial differences are observed for light ions. In Fig. 10, the seasonal behavior of logarithm of the H^+ concentration obtained in measurements onboard the *DEMETER* satellite is compared with the results of the IRI model within two latitudinal intervals in both

hemispheres at the daytime and nighttime sides. Since IRI predicts absolute values of the H^+ concentrations exceeding several times the experimental values, Fig. 10 shows normalized results of calculations by IRI. The normalizing coefficients for each latitude and local time are given in the caption to Fig. 10. The main differences in the seasonal behavior of light ion concentrations according to the *DEMETER* measurements and IRI are the following:

- according to the *DEMETER* data, the substantial inter-hemisphere asymmetry in the daytime concentration of light ions exists at middle latitudes even near equinoxes;
- according to the *DEMETER* data, the concentrations of light ions at middle latitudes at the night-side have a common October maximum in both hemi-

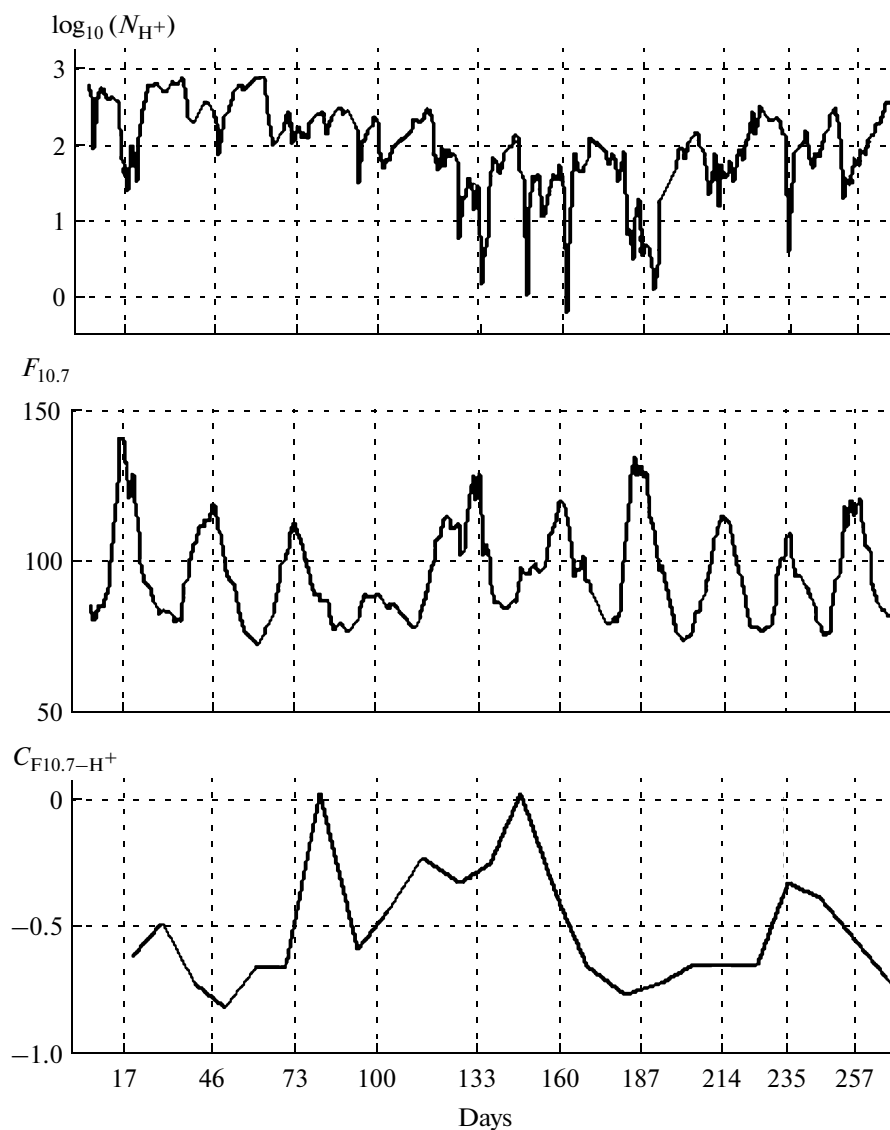


Fig. 8. Logarithm of the H^+ concentration within the latitudinal band $50^\circ\text{--}60^\circ\text{ N}$, solar activity index $F_{10.7}$, and running 30-point correlation coefficient $C_{F_{10.7}-H^+}$ between the $F_{10.7}$ index and logarithm of the concentration.

spheres, unlike IRI which predicts two equivalent maxima near equinoxes;

- in the near-equatorial region, according to the *DEMETER* data, there is a qualitative difference between the seasonal behavior at the daytime and nighttime sides (the H^+ concentration has one maximum at the dayside and two maxima at the nightside), while IRI predicts a seasonal behavior with two maxima at low latitudes both at the daytime and nighttime sides;

- according to the *DEMETER* data, the positions of maxima and minima in the seasonal behavior are shifted by 1–1.5 months relative to their position according to IRI.

The annual mean values of ion concentrations according to the *DEMETER* data are closer to the values from the TTS model than to the IRI model, but averaging over time performed in the TTS model makes it impossible to take into account the qualitative features in the seasonal behavior described above.

The analysis of variations in ion concentrations on the scales from several days to several weeks made it possible to separate the impact of solar and geomagnetic activity on concentration of each type of ions. The conclusion (known from publications) on anti-phase change in the concentrations of O^+ and light ions in the dependence on solar activity is true for the 11-year cycle, but is invalid for variations with shorter time scales, in particular for the 27-day solar cycle

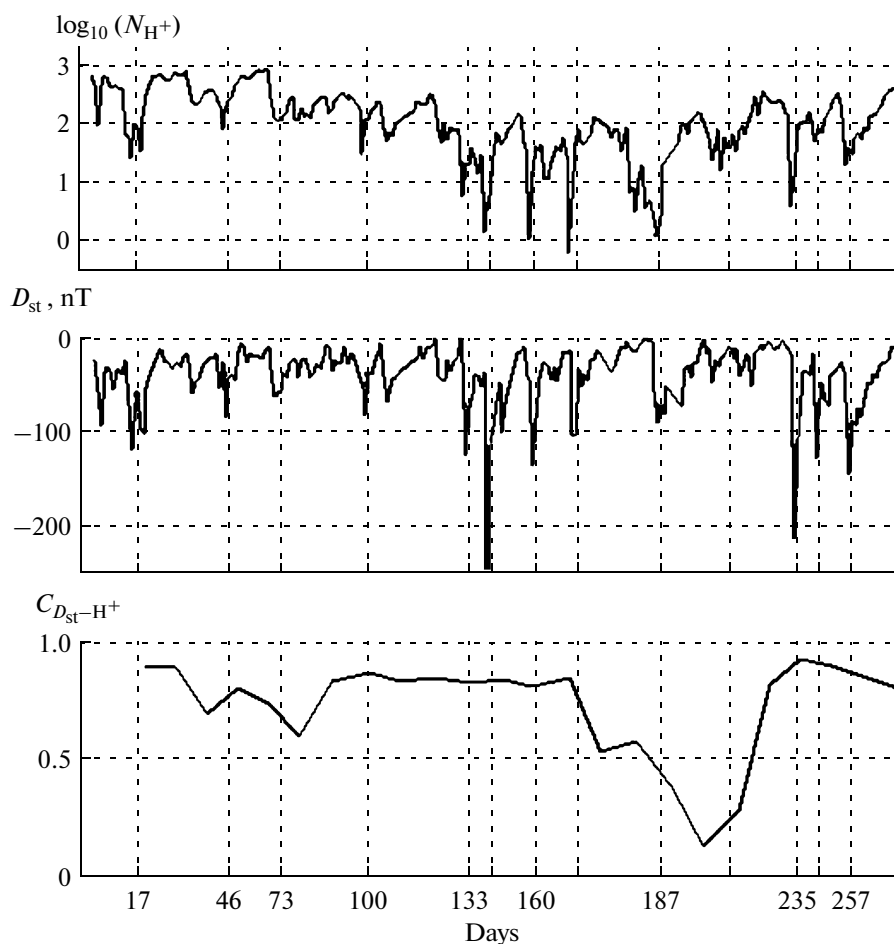


Fig. 9. Logarithm of the H^+ concentration within the latitudinal band $50^\circ\text{--}60^\circ\text{N}$, the D_{st} index, and running 30-point correlation coefficient $C_{D_{st}-H^+}$ between the D_{st} index and logarithm of the concentration.

related to the solar rotation due to different dependence of concentrations of O^+ and light ions on solar and geomagnetic activity.

Variations of the O^+ concentrations at a height of 700 km are governed mainly by solar radiation. It is confirmed by the positive correlation with solar activity in the 11-year cycle and at variations on time scales from several days to several months. The seasonal behavior with a maximum on late local summer, almost similar shape of the curve of the seasonal behavior at all latitudes, and the low-latitude global maximum in the concentration are results of the determining influence of solar radiation on the O^+ concentration.

Variations of the concentrations of H^+ and He^+ in the 11-year cycle of solar activity show the negative correlation with variations in the O^+ concentration and solar activity known from publications [7, 21]. In addition to the dependence on season itself, in the seasonal behavior of the light ions concentration appears (and even dominates at the nightside at middle latitudes) a dependence on global factors: the October

maximum of the H^+ concentration is observed simultaneously in both hemispheres. On the time scales from several days to several weeks, the correlation coefficient between the logarithm of the H^+ concentration and D_{st} index of geomagnetic activity, $C_{D_{st}-H^+}$, is higher than with the solar activity index, the values of $C_{D_{st}-H^+}$ being higher at middle and subauroral latitudes than at low latitudes. The revealed features of the time variations and latitudinal distributions of the light ion concentrations confirm the statement that variations in the light ion concentrations both in the topside ionosphere and in the plasmasphere [23] are governed first of all by geomagnetic activity.

It is evident from our consideration that the correlation coefficients vary depending on the governing parameter level. For example, the correlation coefficient between logarithm of the H^+ concentration and D_{st} index decreases in the periods of low geomagnetic activity. A similar regularity is seen for the relation between solar activity and the O^+ concentration. Thus, the dependence of concentrations on solar and

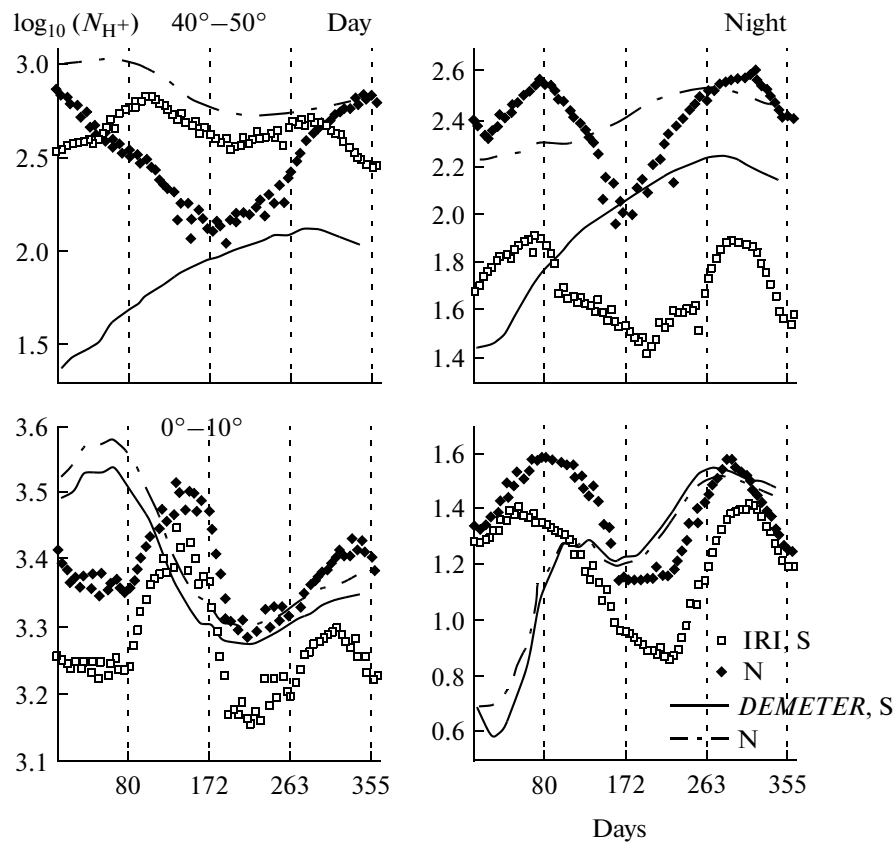


Fig. 10. Logarithm of the H^+ concentration within the latitudinal bands $40^\circ\text{--}50^\circ$ and $0^\circ\text{--}10^\circ$. The downward shift of the calculated curves is 0.5 for the dayside and 2 and 3 for the nightside at middle and low latitudes, respectively. The vertical dashed lines show the positions of equinoxes and solstices.

geomagnetic activity is not reduced to a simple linear dependence, as it has been assumed in [3].

The analysis of variations of ion concentrations in the topside ionosphere performed in this paper can be continued in two directions: elaborating a quantitative description of mean values of ion concentrations and studying irregular short-term variations of ion concentrations. Irregular variations on the time scales from tens of minutes to a day till now were not taken into account in empirical models. However, at least under strong disturbances, ion concentrations in the topside ionosphere vary several-fold during tens of minutes [22], i.e., amplitudes of irregular variations are comparable to the amplitudes of the regular variations considered above. Thus, inclusion into a model of statistical parameters of short-term irregular variations can be important for description of a disturbed state of the ionosphere.

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