



HAL
open science

On the Stability and Accuracy of the Harmonic and Biharmonic Isonutral Mixing Operators in Ocean Models

Florian Lemarié, Laurent Debreu, Alexander Shchepetkin, Jim McWilliams

► **To cite this version:**

Florian Lemarié, Laurent Debreu, Alexander Shchepetkin, Jim McWilliams. On the Stability and Accuracy of the Harmonic and Biharmonic Isonutral Mixing Operators in Ocean Models. *Ocean Modelling*, 2012, 52-53, pp.9-35. 10.1016/j.ocemod.2012.04.007 . hal-00665826v2

HAL Id: hal-00665826

<https://hal.science/hal-00665826v2>

Submitted on 26 Apr 2012

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

On the Stability and Accuracy of the Harmonic and Biharmonic Isonutral Mixing Operators in Ocean Models

F. Lemarié^{a,*}, L. Debreu^b A.F. Shchepetkin^a J.C. McWilliams^a

^a*Institute of Geophysics and Planetary Physics, University of California, 405 Hilgard Avenue, Los Angeles, CA 90095-1567, United States*

^b*INRIA, MOISE Project-Team and Laboratoire Jean Kuntzmann, 51 rue des Mathématiques, 38041 Grenoble Cedex 9, France*

Abstract

Ocean models usually rely on a tracer mixing operator which diffuses along isoneutral directions. This requirement is imposed by the highly adiabatic nature of the oceanic interior, and a numerical simulation needs to respect these small levels of dianeutral mixing to maintain physically realistic results. For non-isopycnic models this is however non-trivial due to the non-alignment of the vertical coordinate isosurfaces with local isoneutral directions, rotated mixing operators must therefore be used. This paper considers the numerical solution of initial boundary value problems for the harmonic (Laplacian) and biharmonic rotated diffusion operators. We provide stability criteria associated with the conventional space-time discretizations of the isoneutral Laplacian operator currently in use in general circulation models. Furthermore, we propose and study possible alternatives to those schemes. A new way to handle the temporal discretization of the rotated biharmonic operator is also introduced. This scheme requires only the resolution of a simple one-dimensional tridiagonal system in the vertical direction to provide the same stability limit of the non-rotated operator. The performance of the various schemes in terms of stability and accuracy is illustrated by idealized numerical experiments of the diffusion of a passive tracer along isoneutral directions.

Key words: Isonutral Mixing, Diffusion Equations, Time Schemes, Von-Neumann stability analysis, Initial-Boundary Value Problems

* Corresponding author.

Phone: +1-310-825-5402

Fax: +1-310-206-3051

E-mail address: florian@atmos.ucla.edu.

1 Introduction

Most ocean numerical models employ isoneutral¹ mixing operators either to parameterize the effect of unresolved mesoscale eddies (Gent and McWilliams, 1990; Smith and Gent, 2004), or more basically to control dispersive errors (Lemarié et al., 2012). It is thus very common for non-isopycnic models to implement a rotation of the diffusion tensor in a direction non-aligned with the computational grid. The benefits of a rotated mixing operator in simulating large scale flows are undeniable (e.g., Danabasoglu et al. (1994), Lengaigne et al. (2003)). Much of the improvements brought by the Gent and McWilliams (1990) parameterization of mesoscale eddies in coarse resolution models are also generally attributed to the orientation of lateral diffusive transport to be along isoneutral directions (Gent, 2011).

Redi (1982) provided the continuous form of the rotation tensor; however additional efforts were required to proceed to the actual implementation at the discrete level. Several works (Cox, 1987; Danabasoglu and McWilliams, 1995; Griffies et al., 1998; Mathieu et al., 1999; Beckers et al., 1998, 2000) tackled this problem that turned out to be more tedious than expected. The discretization in space raises difficulties to properly conserve the monotonicity (Mathieu and Deleersnijder, 1998; Beckers et al., 2000) and global tracer variance dissipation (Griffies et al., 1998) properties of the continuous operator once the problem is discretized. Moreover, due to the small vertical, relative to horizontal, grid distance typically used in numerical models the vertical and cross terms of the tensor can impose a severe restriction on the time step when explicit-in-time methods are used to advance the rotated operator. This stability problem is alleviated by the use of a standard backward Euler scheme for the vertical component of the tensor (Cox, 1987), at the expense of splitting errors² and associated errors in the balance between the active tracer isoneutral diffusive fluxes (Griffies, 2004, Chap. 16). This approach is used in all the state-of-the-art ocean climate models. The existing work on the isoneutral diffusion has been essentially carried out on the second-order (Laplacian) operator and under the small slope approximation (Cox, 1987; Gent and McWilliams, 1990).

Although very few studied so far, a rotated biharmonic operator may be of interest for high resolution simulations due to its known property of scale selectivity. A biharmonic operator non-aligned with the direction of the computational grid is used in Marchesiello et al. (2009) and Lemarié et al. (2012), and discussed in Griffies (2004) (Chap. 14). Because global climate models are now targeting increasingly higher horizontal resolution, the question of the viability of such an operator is not only relevant for the regional modeling community but also for the ocean climate community. As an illustration, Hecht (2010) shows that for a 0.1° resolution global model the use of a Lax-Wendroff scheme with an intrinsic numerically-induced diffusion aligned with the horizontal direction leads to too much of a spurious dia-neutral mixing in the Equatorial Pacific. This result suggests that even for eddy-resolving simulations an isoneutral mixing operator could be required. This motivates the design of a scale-selective (high-order) rotated operator. This is however not straightforward to maintain the stability of such an operator which can produce undesirable effects like overshooting/undershooting (Delhez and Deleersnijder, 2007) and spurious cabelling processes (Griffies, 2004, Chap. 14). To our knowledge, the current implementations of

¹ Throughout the paper we use the terminology "isoneutral direction" (sometimes referred as to "epineutral direction" or simply "neutral direction" in the literature) which is tangent to the locally-referenced potential density surface and whose definition is purely local (McDougall, 1987), rather than "isopycnal direction" which more generally refers to the direction tangent to a potential density surface referenced to an arbitrary fixed pressure.

² In the context of this paper, splitting errors are associated with the splitting of the isoneutral Laplacian operator into a time-explicit part (the horizontal components and cross-derivative terms) and a time-implicit part (the vertical component).

rotated biharmonic operators are based on an explicit Euler scheme in time with ad-hoc tapering or clipping of the neutral slopes to maintain good stability properties (Marchesiello et al., 2009). This approach has however the undesirable effect to allow spurious dianeutral mixing even at places where the slopes are modest and satisfy the small slope approximation.

The aim of this paper is to study a set of space-time discretizations of the rotated harmonic and biharmonic mixing, and to assess them in terms of accuracy, stability and monotonicity violations. One additional constraint we impose to ourselves is accuracy relative to large grid slope ratios (defined as the ratio between the neutral slope and the aspect ratio of the computational grid) in order to make the scheme adequate for use in a terrain-following σ -coordinate model. Indeed, problems with σ models are generally more pernicious than with z -level models because it is not unusual that the slope between the computational grid and the isoneutral direction steepens to be greater than the grid aspect ratio. The paper is organized as follows. In Sec. 2 we introduce the formulation of the isoneutral mixing problem as well as three different ways to discretize the problem in space. Then Sec. 3 and 4 are respectively dedicated to the temporal discretization of the rotated Laplacian and biharmonic operators. Sec. 5 provides the useful details to proceed to the actual implementation of the different schemes in ocean models. Finally, numerical experiments are designed to illustrate the properties of various space-time discretizations in Sec. 6. For clarity, the important notations used throughout the paper are given in Tab. 1.

An alternative approach to the use of isoneutral mixing operators is the design of a vertical coordinate system following the isopycnals (e.g., Hallberg and Adcroft, 2009; Hofmeister et al., 2010; Leclair and Madec, 2011). The present paper is a complementary effort in exploring the merits of different approaches to representing the nearly adiabatic flow in the oceanic interior.

2 Isoneutral Mixing Problem Formulation

2.1 Continuous Formulation

This section briefly introduces the continuous form of the problem under investigation throughout the paper. The three spatial directions are labeled x_1, x_2 for the horizontal coordinates, and x_3 for the vertical coordinate. We note q the tracer of interest, ∇ the three-dimensional gradient operator and $\partial_m = \partial_{x_m}$ ($m = 1, 3$). Following Redi (1982) a mixing operator \mathcal{D}_2 leading to diffusion along isoneutral directions with a diffusivity κ can be defined as

$$\mathcal{D}_2 = \nabla \cdot (\mathbf{R} \nabla q) = \sum_{m=1}^3 \partial_m \left(\sum_{n=1}^3 r_{mn} \partial_n q \right), \quad \mathbf{R} = [r_{mn}]_{1 \leq m, n \leq 3}, \quad (2.1)$$

with \mathbf{R} the following rotation tensor

$$r_{mn} = \kappa \left(\delta_{mn} - \frac{(\partial_m \rho)(\partial_n \rho)}{\|\nabla \rho\|^2} \right), \quad (2.2)$$

where δ_{mn} is the conventional Kronecker delta and ρ is the locally referenced potential density. This form of the tensor shows that the matrix \mathbf{R} is symmetric positive semi-definite ($r_{mm} \geq 0$ and $|r_{mn}| \leq \sqrt{r_{mm} r_{nn}}$) and thus defines a diffusion tensor. Moreover, by construction, the diffusive flux which takes the form $\mathbf{F} = (F^{(1)}, F^{(2)}, F^{(3)})$ with $F^{(m)} = - \sum_{n=1}^3 r_{mn} \partial_n q$ satisfies the orthogonality condition $\mathbf{F} \cdot \boldsymbol{\rho}_\perp = 0$, with

State variables	
q	Three dimensional tracer (can be temperature or salinity)
ρ	Three dimensional density field
Coordinates and spatial operators	
x_1, x_2	Horizontal coordinates
x_3	Vertical coordinate, pointing upward.
Δx_m	Measure of the grid-box interface in the x_m direction
Δt	Time-step for the temporal discretization
$\partial_m = \partial_{x_m}$	Partial derivative in the x_m direction
\mathcal{D}_2	Isonutral Laplacian operator, defined in (2.1)
\mathcal{D}_4	Isonutral biharmonic operator, defined in (2.8)
$\mathbf{F} = (F^{(1)}, F^{(2)}, F^{(3)})$	Diffusive flux, defined in (2.7) at a continuous level and in (2.13), (2.14), and (2.19) at a discrete level
$\mathcal{J}_m(q, \rho)$	Jacobian determinant, defined as $\partial_m q \partial_3 \rho - \partial_m \rho \partial_3 q$
$\delta_m q$	Discrete differentiation in the x_m direction, defined in (2.12) (for $\delta_m \rho$ a particular instance of differentiation to allow computation of isoneutral directions is presented in Sec. 5.5)
Parameters	
κ_m	Diffusivity in the x_m direction
B_m	Hyperdiffusivity in the x_m direction
$\boldsymbol{\alpha} = (\alpha_1, \alpha_2, 0)$	Neutral slope vector, defined in (2.5)
s_m	grid slope ratio, defined in (2.22)
β_m	Parameter controlling the stencil of the spatial discretization of isoneutral mixing operators, defined in (2.21)
w^\pm	Switches to select the computational stencil depending on the orientation of neutral slopes, defined in (2.15)
θ	Stabilizing parameter for the Method of Stabilizing Corrections, defined in (3.3)
$\tilde{\kappa}$	Stabilizing diffusivity for the Method of Stabilizing Corrections, defined in (4.1)
σ_m	Parabolic Courant number in the x_m direction, defined in (3.4)
$\sigma_m^{(4)}$	Square root of the biharmonic Courant number in the x_m direction, defined in (4.2)
z_{mn}	Discrete Fourier modes multiplied by Δt , defined in (3.14)
ϕ_m	Normalized Fourier frequency ($ \phi_m \leq \pi$) in the x_m direction
λ	Exact amplification factor of the isoneutral Laplacian (Sec. 3) and biharmonic (Sec. 4) operators
$\tilde{\lambda}$	Approximate amplification factor obtained after space-time discretization of the isoneutral Laplacian (Sec. 3) and biharmonic (Sec. 4) operators
μ_2	Ratio between the maximum time-steps allowed for stability of the horizontal and the isoneutral Laplacian operators discretized using a forward Euler scheme, defined in (3.33)
μ_4	Same as μ_2 for the biharmonic operator, defined in (4.14)

Table 1

Important notations for the three-dimensional analysis of the isoneutral mixing operators, where $m = (1, 2, 3)$ denotes the three spatial direction.

$\boldsymbol{\rho}_\perp$ the unit vector in the dianeutral direction. In the limit that horizontal density gradients are much more smaller than vertical gradients (i.e., $\sqrt{(\partial_1 \rho)^2 + (\partial_2 \rho)^2} / \partial_3 \rho \ll 1$), a simpler form of the tensor, preserving the isoneutral form of mixing, can be devised (Cox, 1987; Gent and McWilliams, 1990). This assumption corresponds to the so-called small slope approximation, and reduces the tensor (2.2) to

$$\begin{cases} r_{mn} = \kappa \delta_{mn} & (1 \leq m, n \leq 2) \\ r_{m3} = r_{3m} = -\kappa \partial_m \rho / \partial_3 \rho & (1 \leq m \leq 2) \\ r_{33} = \kappa ((\partial_1 \rho)^2 + (\partial_2 \rho)^2) / (\partial_3 \rho)^2 \end{cases} \quad (2.3)$$

This formulation of the rotated diffusion under the small slope approximation had also been derived early by Solomon (1971) in a two-dimensional (x_1, x_3) case. Throughout this paper we consider the small slope approximation and we allow an anisotropy in the diffusivities (i.e. $\kappa_1 \neq \kappa_2$), the corresponding matrix form of the tensor is

$$\mathbf{R} = \begin{pmatrix} \kappa_1 & 0 & \kappa_1 \alpha_1 \\ 0 & \kappa_2 & \kappa_2 \alpha_2 \\ \kappa_1 \alpha_1 & \kappa_2 \alpha_2 & \kappa_1 \alpha_1^2 + \kappa_2 \alpha_2^2 \end{pmatrix}, \quad (2.4)$$

with

$$\boldsymbol{\alpha} = (\alpha_1, \alpha_2, 0) = -\frac{(\partial_1 \rho, \partial_2 \rho, 0)}{\partial_3 \rho}, \quad (2.5)$$

the neutral slope vector. It is straightforward to check that this form of the tensor preserves the symmetry as well as the semi-positive definiteness of the full tensor (2.2). For a laplacian diffusion of a tracer field q in an unbounded domain $\Omega = \mathbb{R}^3$, we can cast the corresponding evolution problem over a time interval $[0, T]$ under a conservative form

$$\begin{cases} \partial_t q = \mathcal{D}_2(q) = \nabla \cdot (\mathbf{R} \nabla q) = -\nabla \cdot \mathbf{F} & \text{in } \Omega \times [0, T], \\ q|_{t=0} = q_0(x_1, x_2, x_3) & \text{in } \Omega, \end{cases} \quad (2.6)$$

with q_0 a given initial condition, and

$$-\mathbf{F} = \begin{pmatrix} -F^{(1)} \\ -F^{(2)} \\ -F^{(3)} \end{pmatrix} = \begin{pmatrix} \kappa_1 (\partial_1 q + \alpha_1 \partial_3 q) \\ \kappa_2 (\partial_2 q + \alpha_2 \partial_3 q) \\ -\alpha_1 F^{(1)} - \alpha_2 F^{(2)} \end{pmatrix}. \quad (2.7)$$

This problem is here temporarily defined on an unbounded spatial domain in order to facilitate the theoretical study in sections 3 and 4. For bounded domains, we provide the necessary boundary conditions in Sec. 5. Using the notations previously defined, the problem for the rotated biharmonic operator reads

$$\begin{cases} \partial_t q = \mathcal{D}_4(q) = -\mathcal{D}_2(\Psi), & \text{with } \Psi = \mathcal{D}_2(q), & \text{in } \Omega \times [0, T], \\ q|_{t=0} = q_0(x_1, x_2, x_3) & & \text{in } \Omega, \end{cases} \quad (2.8)$$

For the biharmonic case, the diffusivities κ_1 and κ_2 in (2.4) must be replaced by $\sqrt{B_1}$ and $\sqrt{B_2}$ with B_1 and B_2 the hyper-diffusivities. Following Griffies (2004) (Chap. 14), we formulate the biharmonic operator as the composition of two rotated Laplacian operators with coefficients $\sqrt{B_m}$ ($m = 1, 2$) because this form ensures that, at the continuous level, the variance of q is strictly dissipated. In the remainder of the paper we focus on the initial value problems (2.6) and (2.8) with the tensor of rotation defined in (2.4).

Unless explicitly said differently, we will use throughout the paper the subscript m to denote the horizontal directions, $m = (1, 2)$.

2.2 Spatial Discretization of the Isonutral Diffusion Operator

2.2.1 Semi-Discrete Considerations

The spatial discretization of the isoneutral diffusion operator is a difficult problem which have been thoroughly tackled by Cox (1987); Beckers et al. (1998, 2000); Griffies et al. (1998). We, first, briefly introduce the delicacies associated with the implementation of the operator \mathcal{D}_2 . Because the x_1 and x_2 directions are independent when adopting the small slope approximation, we consider only the problem defined in the (x_1, x_3) plane. As derived in Beckers et al. (2000), the continuous formulation of the rotated Laplacian operator can be formulated as

$$\mathcal{D}_2(q) = \partial_1 \left(\kappa_1 \frac{\mathcal{J}_1(q, \rho)}{\partial_3 \rho} \right) - \partial_3 \left(\kappa_1 \frac{\partial_1 \rho}{\partial_3 \rho} \frac{\mathcal{J}_1(q, \rho)}{\partial_3 \rho} \right)$$

where $\mathcal{J}_1(q, \rho) = \partial_1 q \partial_3 \rho - \partial_1 \rho \partial_3 q$ is a Jacobian determinant. The main difficulty resides in the evaluation of the $\mathcal{J}_1(q, \rho)/\partial_3 \rho$ term on a staggered grid. This term has to be evaluated at the cell interfaces referred as to u -points and w -points in Fig. 1. Note that, usually, the neutral slope $\partial_1 \rho / \partial_3 \rho$ used in practice is not the result of a discretization scheme only, but additional ad-hoc constraints taking the form of a smoothing, tapering, or clipping procedure are applied; more details concerning this point are given in Sec. 5. Defining the arithmetic average operators $\overline{\cdot}^{(1)}$ in the horizontal and $\overline{\cdot}^{(3)}$ in the vertical, and noting that the natural position of the $\partial_1 q$ and $\partial_1 \rho$ (resp. $\partial_3 q$ and $\partial_3 \rho$) terms is at u -points (resp. w -points), several linear discretizations of the Jacobian determinant can be proposed using a semi-discrete view of the problem:

- A simple discrete analog of the Jacobian \mathcal{J}_1 has been implemented in the early versions of the MOM-GFDL³ model (Cox, 1987; Danabasoglu and McWilliams, 1995), at u -points and w -points the discretization reads

$$\begin{cases} \mathcal{J}_1^u(q, \rho) = \partial_3 \rho \left(\partial_1 q - \partial_1 \rho \frac{\partial_3 \overline{q}^{(3,1)}}{\partial_3 \overline{\rho}^{(3,1)}} \right) \\ \mathcal{J}_1^w(q, \rho) = \partial_3 \rho \left(\partial_1 \overline{q}^{(3,1)} - \partial_1 \overline{\rho}^{(3,1)} \frac{\partial_3 q}{\partial_3 \rho} \right) \end{cases} \quad (2.9)$$

where q and ρ are subject to a double averaging in the x_3 and x_1 directions before differencing, which makes this approach prone to a computational mode (Griffies et al., 1998). Note that the $\partial_3 \rho$ in front of the parenthesis has a passive role here because this is formally $\mathcal{J}_1(q, \rho)/\partial_3 \rho$ that we aim at discretizing.

- An other way to compute the Jacobian is by differencing before averaging, as proposed in Griffies et al. (1998). In this case we have

$$\begin{cases} \mathcal{J}_1^u(q, \rho) = \partial_3 \rho \left(\partial_1 q - \partial_1 \rho \overline{\left[\frac{\partial_3 q}{\partial_3 \rho} \right]}^{(1,3)} \right) \\ \mathcal{J}_1^w(q, \rho) = \partial_3 \rho \left(\overline{\partial_1 q}^{(1,3)} - \overline{\partial_1 \rho}^{(1,3)} \overline{\left[\frac{\partial_3 q}{\partial_3 \rho} \right]} \right), \end{cases} \quad (2.10)$$

where the term in brackets could be simply replaced by $\partial q / \partial \rho$, however we keep $\partial_3 \rho$ explicitly to clearly identify the neutral slopes.

- To obtain more symmetry between the u and w points we can define the following quantity at the corners of the grid cells (sometimes referred as to Ψ -points)

³ Geophysical Fluid Dynamics Laboratory

$$\mathcal{J}_1^\Psi(q, \rho) = \partial_3 \rho \left(\overline{\partial_1 q}^{(3)} - \overline{\partial_1 \rho}^{(3)} \left[\overline{\frac{\partial_3 q}{\partial_3 \rho}} \right]^{(1)} \right)$$

which gives

$$\mathcal{J}_1^u(q, \rho) = \overline{\mathcal{J}_1^\Psi(q, \rho)}^{(3)}, \quad \text{and} \quad \mathcal{J}_1^w(q, \rho) = \overline{\mathcal{J}_1^\Psi(q, \rho)}^{(1)}.$$

This discretization has the interesting property to define the neutral slopes at Ψ -points only, which makes it more convenient to handle the tapering, clipping, or smoothing procedure. Indeed, for the schemes (2.9) and (2.10) this procedure has to be done twice, at u and w points. However this discretization has the major drawback not to reduce to the classical $(1, -2, 1)$ stencil when the neutral slopes vanish, which disqualifies it.

- We can suggest a last approach which would consist in providing more flexibility to the discretization by introducing two sets of weights ν_n^u and ν_n^w in the problem. We define the operator $\overline{\cdot}^{(\nu^u)}$ as the weighted average of the four w -points surrounding a u -point, and $\overline{\cdot}^{(\nu^w)}$ as the weighted average of the four u -points surrounding a w -point. A generalization of scheme (2.10) is

$$\begin{cases} \mathcal{J}_1^u(q, \rho) = \partial_3 \rho \left(\partial_1 q - \partial_1 \rho \left\{ \overline{\left[\frac{\partial_3 q}{\partial_3 \rho} \right]}^{(\nu^u)} \right\} \right) \\ \mathcal{J}_1^w(q, \rho) = \partial_3 \rho \left(\overline{\partial_1 q}^{(\nu^w)} - \overline{\partial_1 \rho}^{(\nu^w)} \left[\frac{\partial_3 q}{\partial_3 \rho} \right] \right), \end{cases} \quad (2.11)$$

The parameters ν_n^u and ν_n^w are set by requiring additional properties of the discretization scheme. An example is the LINEAR1 scheme of Beckers et al. (2000) which sets the ν_n^u and ν_n^w coefficients to get the most compact stencil in the dianeutral direction to reduce the amount of spurious dianeutral mixing. This scheme reduces to a $(1, -2, 1)$ stencil in the diagonal when the angle between the computational grid and the isoneutral direction is ± 45 degrees; this property is not satisfied by (2.10).

To our knowledge, three different schemes are currently in use in ocean models. The Nucleus for European Modeling of the Ocean (NEMO, Madec (2008)) and Coupled Large-scale Ice Ocean (CLIO, Goosse et al. (2008); Mathieu et al. (1999)) models use the Cox (1987) discretization (2.9). In NEMO, an horizontal two-dimensional Laplacian operator acting to smooth the neutral slopes provides the extra diffusion needed to stabilize the scheme (Cox, 1987; Mathieu and Deleersnijder, 1998; Griffies et al., 1998). In the Modular Ocean Model (MOM, Griffies (2010)) and Parallel Ocean Program (POP, Smith et al. (2010)) the discretization (2.10) based on the tracer/density triad formalism (Fig. 1) is used. A triad is defined as an elementary computational stencil of the jacobian $\mathcal{J}_1(q, \rho)$, and therefore of the rotated operator (Griffies et al., 1998). In MOM, an extra vertical smoothing of the $\partial_3 \rho$ term is used (Griffies, 2010, Chap. 16). Finally, even if not explicitly documented, the scheme (2.11) is implemented in the σ -coordinate Regional Oceanic Modeling System (ROMS, Shchepetkin and McWilliams (2005)) and is routinely used so far to rotate along geopotentials the explicit diffusion in the sponge layers near the open boundaries. We describe below the procedure to compute the weighted averages, and we show that this scheme can also be expressed in a tracer triad formalism. In the present study, the MOM/POP discretization is referred as to TRIADS, the NEMO/CLIO discretization as to COX, and the ROMS discretization as to SW-TRIADS (SW stands for switching).

From our experience, besides the instability identified by Griffies et al. (1998), the TRIADS and COX schemes provide very similar results. The differences are greater between the TRIADS and the SW-TRIADS schemes, we thus focus our study on those two schemes only.

2.2.2 Discrete Fluxes

We define

$$\delta_1 q_{i+\frac{1}{2},j,k} = q_{i+1,j,k} - q_{i,k} \quad \text{and} \quad \delta_3 q_{i,j,k+\frac{1}{2}} = q_{i,j,k+1} - q_{i,j,k}. \quad (2.12)$$

The metric terms $(\Delta x_2)_{i+\frac{1}{2},j}$ and $(\Delta x_3)_{i+\frac{1}{2},j,k}$ are the horizontal and vertical measures of the corresponding grid-box interfaces and $(\Delta x_1)_{i+\frac{1}{2},j}$ is the distance between $q_{i+1,j,k}$ and $q_{i,j,k}$. The vertical index k varies from $k = 1$ for the first grid cell next to the ocean floor to $k = N$ at the surface, N corresponds to the number of vertical levels of the discretization. The methodology to compute the $\delta_m \rho$ terms to obtain the isoneutral directions is given explicitly later in Sec. 5.5 and differs from the formula (2.12).

The interfacial $F^{(1)}$ flux discretized using the TRIADS scheme reads

$$-F_{i+\frac{1}{2},j,k}^{(1,\text{triads})} = \kappa_1 (\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k} \left[\frac{\delta_1 q_{i+\frac{1}{2},j,k}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{\delta_1 \rho_{i+\frac{1}{2},j,k}}{(\Delta x_1)_{i+\frac{1}{2},j}} \cdot \frac{1}{4} \left\{ \frac{\delta_3 q_{i,j,k-\frac{1}{2}}}{\delta_3 \rho_{i,j,k-\frac{1}{2}}} + \frac{\delta_3 q_{i+1,j,k+\frac{1}{2}}}{\delta_3 \rho_{i+1,j,k+\frac{1}{2}}} + \frac{\delta_3 q_{i+1,j,k-\frac{1}{2}}}{\delta_3 \rho_{i+1,j,k-\frac{1}{2}}} + \frac{\delta_3 q_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} \right\} \right]. \quad (2.13)$$

Using the notations introduced in Fig. 1, we see that the TRIADS scheme uses the four triads labelled 2,4,5, and 6 with the same weight $w = 1/4$. The derivatives in the x_1 -direction are computed along the horizontal segment of a triad while the derivatives in the x_3 -direction are computed along the vertical segment. Alternatively, the spirit of the LINEAR1 scheme introduced in Beckers et al. (2000) is to weight those triads depending on the orientation of the slope by keeping only two of them. The LINEAR1 scheme was originally derived for a constant slope, we extend this scheme to the case with spatially variable neutral slopes to obtain the more general SW-TRIADS scheme. If we assume a stable stratification (i.e. $\delta_3 \rho_{i,j,k+\frac{1}{2}} < 0, \forall k$), this scheme reads

$$-F_{i+\frac{1}{2},j,k}^{(1,\text{sw-triads})} = \kappa_1 (\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k} \left[\frac{\delta_1 q_{i+\frac{1}{2},j,k}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{1}{2} \left\{ \frac{w_{i+\frac{1}{2},j,k}^+}{(\Delta x_1)_{i+\frac{1}{2},j}} \left(\frac{\delta_3 q_{i,j,k-\frac{1}{2}}}{\delta_3 \rho_{i,j,k-\frac{1}{2}}} + \frac{\delta_3 q_{i+1,j,k+\frac{1}{2}}}{\delta_3 \rho_{i+1,j,k+\frac{1}{2}}} \right) + \frac{w_{i+\frac{1}{2},j,k}^-}{(\Delta x_1)_{i+\frac{1}{2},j}} \left(\frac{\delta_3 q_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} + \frac{\delta_3 q_{i+1,j,k-\frac{1}{2}}}{\delta_3 \rho_{i+1,j,k-\frac{1}{2}}} \right) \right\} \right]. \quad (2.14)$$

with

$$w_{i+\frac{1}{2},j,k}^+ = \max(\delta_1 \rho_{i+\frac{1}{2},j,k}, 0), \quad w_{i+\frac{1}{2},j,k}^- = \min(\delta_1 \rho_{i+\frac{1}{2},j,k}, 0). \quad (2.15)$$

Depending on the orientation of the slope, either triads 4 and 5 are selected, or triads 2 and 6. This results in a compact four point stencil for the discretization of the $F_{i+\frac{1}{2},j,k}^{(1,\text{sw-triads})}$ interfacial flux. The cancellation of the contribution of cross-isoneutral points with the SW-TRIADS scheme tends to reduce the amount of spurious dianeutral mixing associated with discretization errors. This point will be exemplified in Sec. 6. The discretization of the vertical flux $F^{(3)}$ does not raise any additional difficulty for the TRIADS scheme and can be found in Griffies et al. (1998). We show in the next subsection a way to construct the vertical flux $F_{i,j,k+\frac{1}{2}}^{(3,\text{sw-triads})}$ so that the SW-TRIADS scheme satisfies a globally diminishing tracer variance.

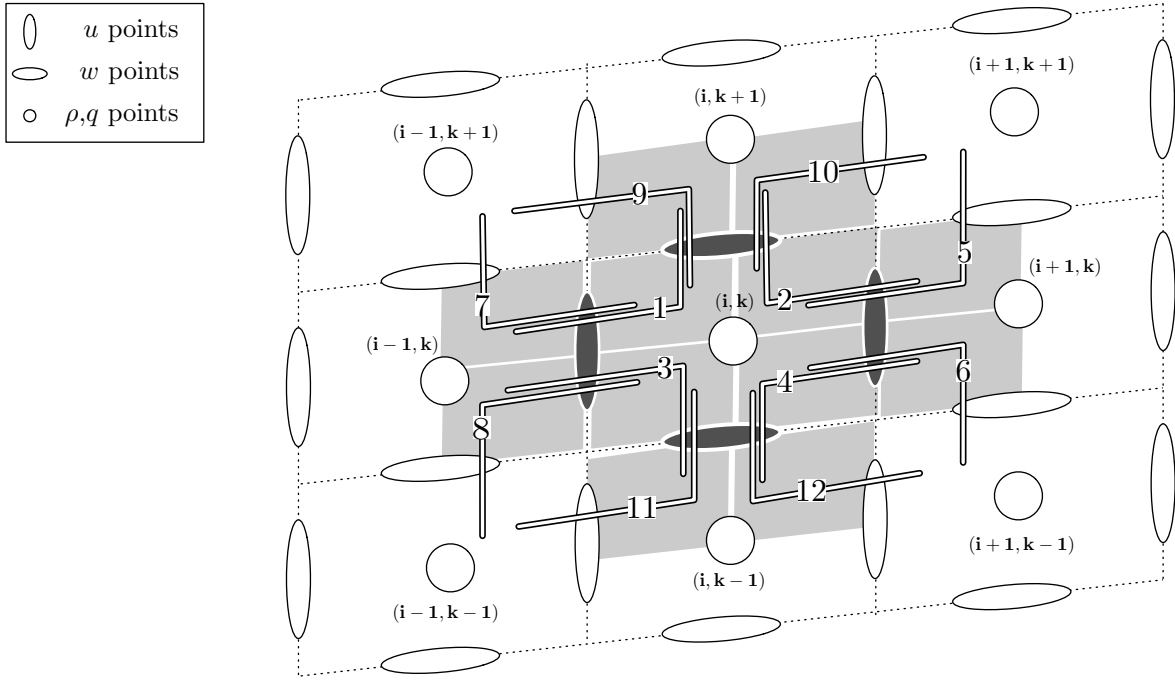


Fig. 1. Grid stencil, constructed upon 12 triads, involved in the computation of the rotated Laplacian operator in the (x_1, x_3) plane. Each triad has an associated quarter cell (gray shaded areas). The zonal $F^{(1)}$ and vertical $F^{(3)}$ components of the isoneutral diffusive flux are computed at the cell interfaces surrounding the (i, k) location (dark grey shaded points). The numbering of the triads is meant to be consistent with the one used in Griffies (2004) (Chap. 16).

2.3 Global Tracer Variance Dissipation

As shown in Griffies et al. (1998), the TRIADS scheme has been constructed to ensure that the discretized operator globally satisfies the tracer variance dissipation property of the continuous operator. This scheme is designed on the basis of the variational principle. A weak form of the problem under investigation can be defined through

$$\mathcal{G}[q] = -\frac{1}{2} \int_{\Omega} \nabla q \cdot (\mathbf{R} \nabla q) \, d\Omega \quad (2.16)$$

where $\mathcal{G}[q]$ is a functional whose Fréchet derivative $\delta \mathcal{G}[q]/\delta q$ gives the diffusion operator. Due to the symmetric positive semi-definite property of the diffusion tensor \mathbf{R} , \mathcal{G} is negative semi-definite. This property implies that the corresponding operator acts to decrease the total tracer variance, indeed Griffies et al. (1998) showed that $\partial_t \int_{\Omega} q^2 \, d\Omega = 4\mathcal{G}[q]$. Unlike the COX scheme, the TRIADS scheme provides a negative semi-definite functional at the discrete level, thus ensuring that the corresponding discretized operator is globally strictly dissipative (Griffies et al., 1998; Smith and Gent, 2004; Griffies, 2004). The reader is referred to Griffies (2004) (Chap. 16) for more details about the foundations for the dissipation functional and its discretization.

Consistent with the notations introduced in Griffies (2004) (Chap. 16) the discretization of the functional \mathcal{G} , in the (x_1, x_3) plane, is given by

$$\mathcal{G}^{(x_1-x_3)}[q] = -\frac{1}{2} \sum_{i,k} \sum_{n=1}^{12} A(n)V(n) \left[\frac{\delta_1 q(n)}{\Delta x_1(n)} + \alpha_1(n) \frac{\delta_3 q(n)}{\Delta x_3(n)} \right]^2 \equiv \sum_{i,k} \left(\sum_{n=1}^{12} L_{i,k}^{(n)} \right), \quad (2.17)$$

where the subscripts (i, k) run over all the cells of the computational domain. In (2.17), $A(n)$ is the diffusivity associated with triad n , $V(n)$ is the volume of quarter-cell n , $\Delta x_1(n)$ and $\Delta x_3(n)$ are respectively the length of the horizontal and vertical segments of triad n . For a given cell (i, k) , the contribution to the discrete functional is given by $\sum_{n=1}^{12} L_{i,k}^{(n)}$ which corresponds to 12 nonpositive components associated with 12 quarter cells, built in such a way that there is a unique volume $V(n)$ and diffusivity $A(n)$ for each quarter-cell (represented as light gray shaded areas in Fig. 1). This procedure ensures that $\mathcal{G}^{(x_1-x_3)}[q]$ is negative semi-definite. For example, the contribution of triad 1 to the discretized diffusion operator is given by

$$\frac{\delta L_{i,k}^{(1)}}{\delta q_{i,k}} = -A(1)V(1) \left(\frac{\delta_1 q(1)}{\Delta x_1(1)} + \alpha_1(1) \frac{\delta_3 q(1)}{\Delta x_3(1)} \right) \left(\underbrace{\frac{1}{\Delta x_1(1)}}_{\text{contribution to } F^{(1)}} - \underbrace{\frac{\alpha_1(1)}{\Delta x_3(1)}}_{\text{contribution to } F^{(3)}} \right) \quad (2.18)$$

which shows that when a triad is used to compute $F^{(1)}$, the same triad is used to compute $F^{(3)}$. The only difference between the TRIADS and the SW-TRIADS schemes is in the weighting of the 12 triads, the SW-TRIADS scheme simply cancels the contribution of certain triads depending on the orientation of the neutral slope α_1 , which is equivalent to set $A(n)$ to zero for those triads. This means that once a triad n is rejected during the computation of $F^{(1, \text{sw-triads})}$ it can not be used again to compute $F^{(3, \text{sw-triads})}$ because the diffusivity $A(n)$ associated with this triad is zero. If we follow this simple rule, we find that there is a unique way to define the vertical interfacial flux $F_{i,j,k+\frac{1}{2}}^{(3, \text{sw-triads})}$:

$$\begin{aligned} -F_{i,j,k+\frac{1}{2}}^{(3, \text{sw-triads})} &= \kappa_1 \frac{(\Delta x_3)_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} \left[\frac{w_{i-\frac{1}{2},j,k}^+}{(\Delta x_1)_{i-\frac{1}{2},j}} \left(\frac{\delta_1 \rho_{i-\frac{1}{2},j,k}}{(\Delta x_1)_{i-\frac{1}{2},j}} \cdot \frac{\delta_3 q_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} - \frac{\delta_1 q_{i-\frac{1}{2},j,k}}{(\Delta x_1)_{i-\frac{1}{2},j}} \right) \right. \\ &\quad + \frac{w_{i+\frac{1}{2},j,k+1}^+}{(\Delta x_1)_{i+\frac{1}{2},j}} \left(\frac{\delta_1 \rho_{i+\frac{1}{2},j,k+1}}{(\Delta x_1)_{i+\frac{1}{2},j}} \cdot \frac{\delta_3 q_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} - \frac{\delta_1 q_{i+\frac{1}{2},j,k+1}}{(\Delta x_1)_{i+\frac{1}{2},j}} \right) \\ &\quad + \frac{w_{i-\frac{1}{2},j,k+1}^-}{(\Delta x_1)_{i-\frac{1}{2},j}} \left(\frac{\delta_1 \rho_{i-\frac{1}{2},j,k+1}}{(\Delta x_1)_{i-\frac{1}{2},j}} \cdot \frac{\delta_3 q_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} - \frac{\delta_1 q_{i-\frac{1}{2},j,k+1}}{(\Delta x_1)_{i-\frac{1}{2},j}} \right) \\ &\quad \left. + \frac{w_{i+\frac{1}{2},j,k}^-}{(\Delta x_1)_{i+\frac{1}{2},j}} \left(\frac{\delta_1 \rho_{i+\frac{1}{2},j,k}}{(\Delta x_1)_{i+\frac{1}{2},j}} \cdot \frac{\delta_3 q_{i,j,k+\frac{1}{2}}}{\delta_3 \rho_{i,j,k+\frac{1}{2}}} - \frac{\delta_1 q_{i+\frac{1}{2},j,k}}{(\Delta x_1)_{i+\frac{1}{2},j}} \right) \right] \times \frac{1}{W} \end{aligned} \quad (2.19)$$

where W corresponds to the number of selected triads, i.e. the number of nonzero w^\pm terms. For $W = 2$, it is easy to check that the rotated operator discretized with the SW-TRIADS scheme is based on exactly six triads and is consistent with the functional discretization (2.17) where the contribution of the six other triads is rejected by setting their associated diffusivities to zero. As illustrated in Fig. 2, we have $W = 2$ for the common situations encountered in ocean models. There are however some degenerated cases (Fig. 2, e) and f) for which $W \neq 2$. In those very specific cases, the scheme does not satisfy the functional

discretization because some triads are used to compute $F^{(3)}$ with a different weight than it is used to compute $F^{(1)}$. We can show easily that when the stratification has a $2\Delta x_1$ mode, as studied in Griffies et al. (1998), the SW-TRIADS schemes ensure a global tracer variance reduction.

The SW-TRIADS scheme is a special instance of the TRIADS scheme, it thus makes the implementation of this scheme as straightforward. When spatially variable diffusion coefficients are used, each triad must be weighted by the corresponding coefficient. Usually, those coefficients are computed at the horizontal interfacial u points. In this case, in (2.19) the constant diffusivity κ_1 should be dropped and each coefficient w^\pm must be multiplied by the corresponding diffusivity. Because each triad crosses exactly one u point (Fig. 1) there is a unique diffusivity associated with a given triad.

It is worth mentioning that the tracer variance diminishing property should not be confused with the total variation diminishing (TVD) property which is monotonicity preserving. As we show in Sec. 6, the rotated operators discretized with the TRIADS or the SW-TRIADS scheme do not preserve monotonicity (Mathieu and Deleersnijder, 1998; Beckers et al., 2000).

Because the rotated biharmonic operator corresponds to two successive rotated Laplacian operators, it does not raise any additional difficulties as long as the spatial discretization is concerned. In the two following sections we consider the time integration of the rotated operators, and we assume that the slope vector α is spatially constant and that the grid is uniform to make the stability analysis tractable. For a constant slope, the discrete approximation $\widetilde{\mathcal{D}}_2$ of the rotated operator at the position (i, j, k) is given by

$$\begin{aligned} \left(\widetilde{\mathcal{D}}_2\right)_{i,j,k} &= - \left[\frac{F_{i+\frac{1}{2},j,k}^{(1)} - F_{i-\frac{1}{2},j,k}^{(1)}}{\Delta x_1 \Delta x_2 \Delta x_3} + \frac{F_{i,j+\frac{1}{2},k}^{(2)} - F_{i,j-\frac{1}{2},k}^{(2)}}{\Delta x_1 \Delta x_2 \Delta x_3} + \frac{F_{i,j,k+\frac{1}{2}}^{(3)} - F_{i,j,k-\frac{1}{2}}^{(3)}}{\Delta x_3} \right] \\ &= \sum_{p=-1}^1 \sum_{l=-1}^1 \mu_{p,l}^{(1)} q_{i+p,j,k+l} + \sum_{p=-1}^1 \sum_{l=-1}^1 \mu_{p,l}^{(2)} q_{i,j+p,k+l}, \end{aligned} \quad (2.20)$$

with

$$\left[\mu_{p,l}^{(m)}\right]_{-1 \leq p,l \leq 1} = \frac{\kappa_m}{\Delta x_m^2} \begin{pmatrix} (\beta_m - 1) \frac{s_m}{2} & s_m(s_m - \beta_m) & (1 + \beta_m) \frac{s_m}{2} \\ 1 - s_m \beta_m & -2(1 + s_m^2) + 2s_m \beta_m & 1 - s_m \beta_m \\ (1 + \beta_m) \frac{s_m}{2} & s_m(s_m - \beta_m) & (\beta_m - 1) \frac{s_m}{2} \end{pmatrix} \quad (2.21)$$

where

$$s_m = \frac{\Delta x_m}{\Delta x_3} \alpha_m \quad (2.22)$$

is the grid slope ratio, $\beta_m \equiv 0$ gives the TRIADS scheme and $\beta_m = 1$ for $s_m \geq 0$ (resp. $\beta_m = -1$ for $s_m < 0$) the SW-TRIADS scheme. The rotated Laplacian operator in the x_m -direction is thus discretized on a centered 9-point stencil if $\beta_m = 0$ and on a more compact 7-point stencil if $|\beta_m| = 1$. As far as the rotated biharmonic is concerned, the discretization is based on a 25-point stencil with the TRIADS and a 19-point stencil with the SW-TRIADS. We do not investigate this possibility in this study, but the parameter β_m could be used as a degree of freedom to derive alternative properties of the discretization.

The grid slope ratio s_m , introduced in (2.21), is a key dimensionless parameter that will be used throughout this paper. This parameter corresponds to the ratio between the neutral slope α_m and the aspect ratio $\Delta x_3/\Delta x_m$. We thus have $s_m = \pm 1$ for a ± 45 degrees slope. Values of $|s_m|$ greater than 1 generally lead to a degradation of the accuracy of the rotated diffusion due to the need for extrapolation to compute the isoneutral direction. For typical applications with a coarse resolution global climate model, $|s_m|$ can

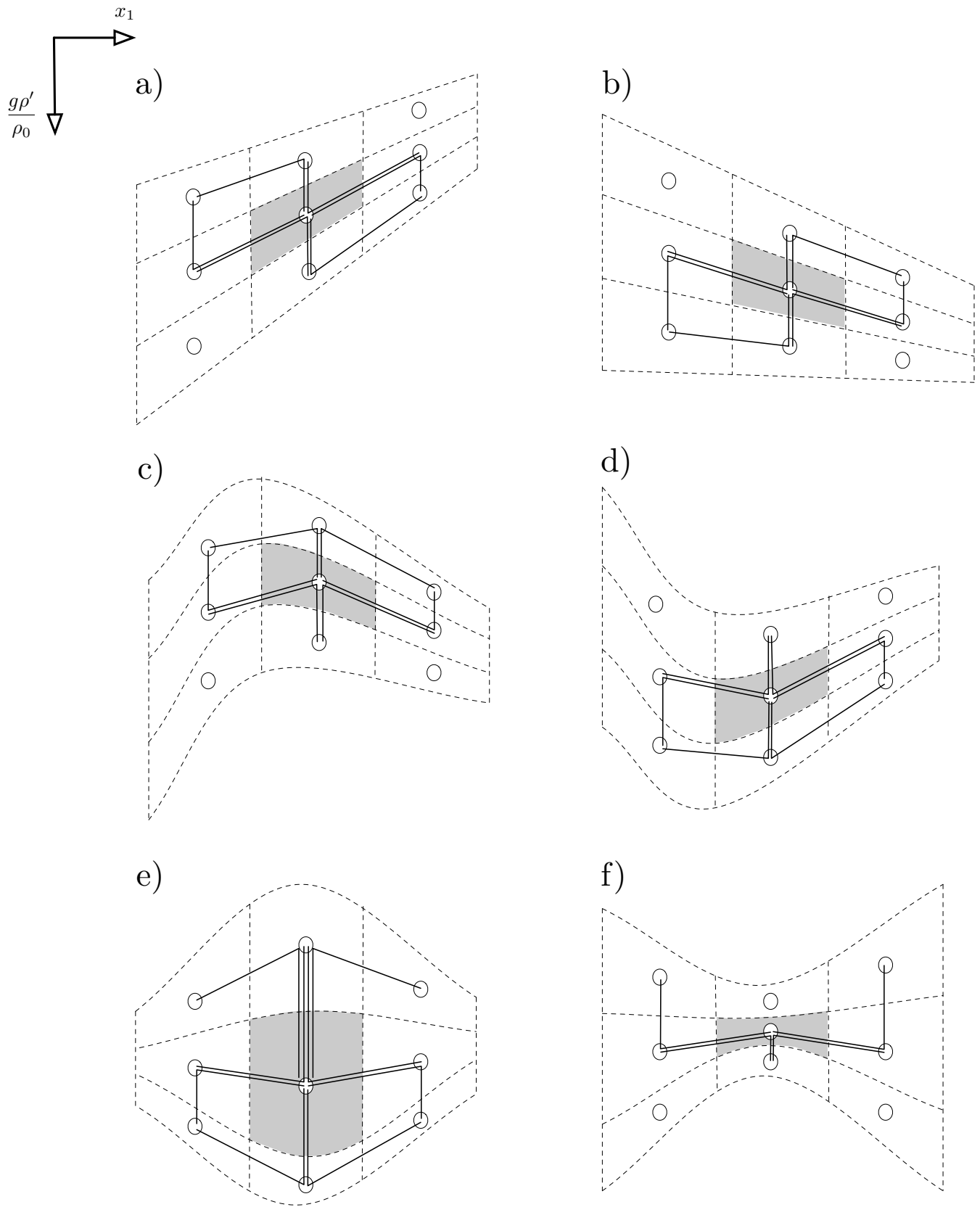


Fig. 2. Computational stencil, represented as triads (black lines), involved in the computation of the rotated Laplacian operator with the SW-TRIADS scheme for different orientations between the computational grid (dashed lines) and the isoneutral direction which corresponds to horizontal lines in the $(x_1, g\rho'/\rho_0)$ frame. In cases a) to d), 6 triads are selected while 8 triads are selected in e) and 4 in f).

reach $\mathcal{O}(10)$ values (e.g.; $s_m = 10$ for $\alpha_m = 5 \times 10^{-3}$, $\Delta x_1 = 100$ km, and $\Delta x_3 = 50$ m). Note that the use of slope clipping or tapering sets an upper bound on s_m which can not be arbitrarily large in this case. Large grid slope ratios should not be confused with large neutral slopes, meaning that large values of s_m can exist even under the small slope approximation. In the context of a σ to geopotentials rotation, as in Marchesiello et al. (2009), s_m corresponds to the so-called hydrostatic inconsistency number.

3 Time Discretization of the Isonutral Laplacian Operator

3.1 Proposed Schemes

The overall objective of this section is to derive a time-integration scheme whose stability limit is imposed by the horizontal components of the tensor. In this case the constraint on the time step Δt of the temporal discretization would be equivalent between the rotated and the non-rotated operators. Moreover, because we are considering a diffusive process, we do not feel necessary to strive to design a high-order scheme in time, the aim is to keep the study as simple as possible. The rotated Laplacian operator \mathcal{D}_2 , with neutral slopes α_m and diffusivities κ_m , can be linearly decomposed into a sum of three functions $\mathcal{D}_2(q) = \sum_{m=1}^3 G_m(q)$, where

$$\begin{cases} G_m(q) = \partial_m (\kappa_m [\partial_m q + \alpha_m \partial_3 q]) + \partial_3 (\alpha_m \kappa_m \partial_m q) & (m = 1, 2), \\ G_3(q) = \partial_3 \left([\kappa_1 \alpha_1^2 + \kappa_2 \alpha_2^2] \partial_3 q \right). \end{cases} \quad (3.1)$$

We consider that the piece $G_0(q) = G_1(q) + G_2(q)$ is always treated explicitly in time-integration schemes, whereas G_3 represent a stiff and unidirectional contribution that can be treated implicitly, if needed. The integration of the G_0 term in an implicit manner would require the solution of a complicated implicit system in the horizontal direction which would be laborious to implement in parallel and would significantly affect the performances of the numerical model. The common practice in climate models is to use a standard backward Euler scheme to advance the vertical G_3 component of the tensor (Cox, 1987). Using the time step $\Delta t > 0$, we note q^n the approximation $q^n \approx q(t_n)$, with $t_n = n\Delta t$. The semi-discretized version of the scheme introduced by Cox (1987) is

$$q^{n+1} = q^n + \Delta t \{G_0(q^n) + G_3(q^{n+1})\}. \quad (3.2)$$

In the following this scheme will be referred as to (IMP) scheme. Moreover, the fully explicit version, i.e. with $G_3(q^n)$ instead of $G_3(q^{n+1})$ in (3.2), will be denoted by (EXP).

Following the work of Douglas (1962), Andreev (1967) or Craig and Sneyd (1988) (see also in 't Hout and Welfert (2009) for a review), split schemes such as alternating direction implicit (ADI) have proved valuable in the approximation of the solutions of multi-dimensional parabolic problems with mixed derivatives. This type of scheme is usually implemented with unidirectional implicit corrector steps in each spatial direction to pursue an unconditional stability. Alternatives to this scheme allowing one or more spatial directions to be treated explicitly can be found in Douglas and Gunn (1964); van der Houwen and Verwer (1979) or Hundsdorfer (2002). Those authors propose a multi-stage method : at the first stage, a consistent (explicit) approximation of the operator is evaluated, while all succeeding stages serve to improve the stability. This scheme has been called Method of Stabilizing Corrections (referred as to (MSC) hereafter) in van der Houwen and Verwer (1979) and seems particularly well suited for our problem

because it provides a consistent, efficient, and easy-to-implement scheme with degrees of freedom to ensure good stability properties. Moreover, we see in Sec. 4 that this approach can be extended to the time-integration of the rotated biharmonic operator. For the diffusion problem (2.6), the (MSC) scheme reads

$$\begin{cases} q^* = q^n + \Delta t \{G_0(q^n) + G_3(q^n)\} \\ q^{n+1} = q^* + \theta \Delta t \{G_3(q^{n+1}) - G_3(q^n)\} \end{cases} \quad (3.3)$$

where $\theta \geq 0$ is a real parameter. Note that for our purpose we only allow a stabilizing correction in the vertical direction. In (3.3), $\theta = 0$ gives the (EXP) scheme and $\theta = 1$ the (IMP) scheme. For $\theta = 1/2$, we retrieve the Crank-Nicolson scheme, however we will see later that this scheme is not particularly well-suited in the context of the isoneutral Laplacian operator because it does not stabilize the cross-derivative terms as effectively as the (IMP) scheme. In the next section we study the stability range of the (EXP), (IMP) and (MSC) methods subject to the TRIADS or SW-TRIADS discretizations in space.

3.2 Important Results

We give here the important results of our study on the Laplacian operator, the associated proof is provided in Sec. 3.3. Those results are given in terms of the parabolic Courant number

$$\sigma_m = \kappa_m \frac{\Delta t}{\Delta x_m^2}. \quad (3.4)$$

Horizontal Laplacian Operator

When using an Euler forward scheme, the stability limit of the two-dimensional horizontal Laplacian operator, discretized on a five-point stencil⁴, is

$$\sigma_1 + \sigma_2 \leq \frac{1}{2}. \quad (3.5)$$

⁴ A nine-point discrete Laplacian operator would provide a stability criterion which is somewhat less restrictive than for the usual five-point one.

• **TRIADS discretization**

- The (EXP) scheme is stable for

$$\sigma_1(1 + s_1^2) + \sigma_2(1 + s_2^2) \leq \frac{1}{2} \quad (3.6)$$

which is always more restrictive than (3.5).

- The stability limit of the (IMP) scheme is given by (3.5).
- The same stability constraint (3.5) is obtained for the (MSC) scheme when θ is chosen such that

$$(s_1^2\sigma_1 + s_2^2\sigma_2)\theta = \frac{1}{2} \max \left\{ -1 + 2 \left[\sigma_1(1 + s_1^2) + \sigma_2(1 + s_2^2) \right], 0 \right\}. \quad (3.7)$$

The value of θ ranges from $\theta = 0$ when the (EXP) scheme is stable to $\theta = 1$ when $\sigma_1 + \sigma_2 = 1/2$. The amount of implicit diffusion is thus always smaller with the (MSC) scheme compared with the (IMP) scheme.

• **SW-TRIADS discretization**

- The (EXP) scheme is stable for

$$\sigma_1 \max \{ s_1^2, 1 \} + \sigma_2 \max \{ s_2^2, 1 \} \leq \frac{1}{2} \quad (3.8)$$

which shows that for $s_1^2 \leq 1$ and $s_2^2 \leq 1$ the stability constraint is the same as the non-rotated operator.

- As for the TRIADS case, the (IMP) scheme is stable if condition (3.5) is satisfied.
- The stability condition (3.5) applies to the (MSC) scheme for

$$\theta = \max \left\{ \frac{|s_1| - 1}{|s_1|}, \frac{|s_2| - 1}{|s_2|}, 0 \right\} \quad (3.9)$$

Before demonstrating those results, we draw a few remarks :

- As shown in Sec. 3.3, the values of θ given in (3.7) and (3.9) are sufficient conditions for stability. Some conservative choices have been made during the analysis to simplify the results.
- The value of θ in (3.7) is equal to 1 when $\sigma_1 + \sigma_2 = 1/2$. Consequently, if the diffusivities κ_m are constant everywhere and such that $\sigma_1 + \sigma_2 = 1/2$, the (MSC) scheme is equivalent to the (IMP) scheme. The (MSC) scheme can, however, be particularly interesting when flow-dependent diffusion coefficients are used because it minimizes the amount of implicit vertical diffusion required for stability.
- The Crank-Nicolson scheme (i.e. $\theta = 1/2$) does not provide a sufficient condition ensuring that the isoneutral Laplacian operator can be advanced with the same time step as the horizontal diffusion operator. Indeed, we show in (3.7) that values of θ larger than $1/2$ are required, especially when $\sigma_1 + \sigma_2 = 1/2$.

We now provide the methodology to derive the stability conditions.

3.3 Proof Through Linear Stability Analysis

We first assume that the tracer q can be Fourier decomposed as

$$q(x_1, x_2, x_3, t) = \sum_{\mathbf{k}} A_{\mathbf{k}}(t) \exp\left(i \sum_{m=1}^3 k_m x_m\right), \quad i = \sqrt{-1} \quad (3.10)$$

with $\mathbf{k} = (k_1, k_2, k_3)$ the three-dimensional wave-vector. Substitution in (2.6) leads to the linear damping equation

$$\frac{dq}{dt} = -\eta q, \quad \text{with } \eta = \sum_{m=1}^2 \kappa_m (k_m + \alpha_m k_3)^2, \quad (3.11)$$

and subsequently

$$q(t_{n+1} = t_n + \Delta t) = \exp(-\eta \Delta t) q(t_n) = \lambda q(t_n). \quad (3.12)$$

$\lambda = e^{-\eta \Delta t}$ provides the exact damping obtained with a ‘‘perfect’’ discretization. We can derive the approximate damping $\tilde{\lambda}$ provided by the space-time discretization

$$q(t_{n+1}) = \tilde{\lambda}(z_{11}, z_{22}, z_{33}, z_{31}, z_{13}, z_{32}, z_{23}) q(t_n) \quad (3.13)$$

where the z_{mn} terms are all real and obtained by substitution of discrete Fourier modes in (2.20) and multiplication by Δt :

$$\begin{aligned} z_{mm} &= -2\sigma_m(1 - \cos \phi_m) & m = 1, 2 \\ z_{33} &= -2(s_1^2 \sigma_1 + s_2^2 \sigma_2)(1 - \cos \phi_3) \\ z_{m3} &= z_{3m} = -s_m \sigma_m [\sin \phi_m \sin \phi_3 - \beta_m(1 - \cos \phi_m)(1 - \cos \phi_3)] & m = 1, 2 \end{aligned} \quad (3.14)$$

with $\phi_m = k_m \Delta x_m$ ($|\phi_m| \leq \pi$) the normalized Fourier frequencies. Using those notations, we obtain $\eta \Delta t = \sum_{m=1}^2 \sigma_m (\phi_m + s_m \phi_3)^2$ in (3.12).

The response function of the (MSC) time scheme (3.3) subject to the spatial discretization (2.20-2.21) is

$$\tilde{\lambda} = 1 + \frac{z_0 + z_{33}}{1 - \theta z_{33}}, \quad z_0 = \sum_{m=1}^2 (z_{mm} + z_{m3} + z_{3m}). \quad (3.15)$$

Stability of the scheme under investigation is obtained for $|\tilde{\lambda}| \leq 1$.

In the subsequent paragraphs we first show that the condition $\tilde{\lambda} \leq 1$ is satisfied whatever the value of θ , and whatever the spatial discretization. Then, we show that the TRIADS scheme always provides a more restrictive stability range than the SW-TRIADS scheme. Eventually, we derive the requirements to satisfy the constraint $\tilde{\lambda} \geq -1$ which ensures that the corresponding scheme is stable.

Upper Bound on $\tilde{\lambda}$

To emphasize the fact that the z_{mm} terms are negative, we define $y_m = \sqrt{2\sigma_m(1 - \cos \phi_m)}$, and $y_3 = \sqrt{y_{31}^2 + y_{32}^2}$ (with $y_{3m} = \sqrt{2s_m^2\sigma_m(1 - \cos \phi_3)}$) so that $z_{mm} = -y_m^2$. Moreover, we introduce the vector $\mathbf{v}_m = (\sin \phi_m, \varsigma_m(1 - \cos \phi_m))$, with ς_m such that $\beta_m = -\varsigma_m\varsigma_3$ and $|\varsigma_m| \leq 1$. We can easily show that

$$z_{m3} = z_{3m} = -s_m\sigma_m(\mathbf{v}_m \cdot \mathbf{v}_3) \quad (3.16)$$

and for $m = 1, 2$ (thanks to the condition $|\varsigma_m| \leq 1$)

$$\sigma_m\|\mathbf{v}_m\|^2 \leq \sigma_m\{\sin^2 \phi_m + (1 - \cos \phi_m)^2\} = y_m^2, \quad (3.17)$$

where $\|\bullet\|$ defines the ℓ_2 -norm. (3.16) and (3.17) imply that

$$\begin{aligned} -z_{33} - \sum_{m=1}^2 (z_{mm} + z_{m3} + z_{3m}) &= y_1^2 + y_2^2 + y_3^2 + 2s_1\sigma_1(\mathbf{v}_1 \cdot \mathbf{v}_3) + 2s_2\sigma_2(\mathbf{v}_2 \cdot \mathbf{v}_3) \\ &\geq \sum_{m=1}^2 \sigma_m\|\mathbf{v}_m\|^2 + s_m^2\sigma_m\|\mathbf{v}_3\|^2 + 2s_m\sigma_m(\mathbf{v}_m \cdot \mathbf{v}_3) \\ &= \sum_{m=1}^2 \sigma_m\|\mathbf{v}_m + s_m\mathbf{v}_3\|^2 \geq 0. \end{aligned} \quad (3.18)$$

This result is sufficient to show that $\tilde{\lambda} \leq 1$, indeed thanks to (3.15) we have

$$\tilde{\lambda} = 1 - \left\{ \frac{-z_{33} - \sum_{m=1}^2 (z_{mm} + z_{m3} + z_{3m})}{1 + \theta y_3^2} \right\} \leq 1, \quad (3.19)$$

because we showed that the term in curly brackets is positive. Note that this result is valid whatever $|\beta_m| \leq 1$ and hence whatever the spatial discretization. This result means that stability of the time discretization is obtained for $\tilde{\lambda} \geq -1$. In the following we first show that if this condition is satisfied by the TRIADS scheme the same applies to the SW-TRIADS scheme.

Effect of the Cross-Terms

As shown in (3.14), the difference between the two spatial discretizations under consideration appears only in the z_{3m} terms. The definition of the slope dependent parameter β_m for the SW-TRIADS scheme is such that for all s_m , $\beta_m s_m = |s_m|$. Substitution in (3.14) leads to

$$-z_{3m} = \sigma_m [s_m \sin \phi_m \sin \phi_3 - |s_m|(1 - \cos \phi_m)(1 - \cos \phi_3)] \leq \sigma_m s_m \sin \phi_m \sin \phi_3. \quad (3.20)$$

This inequality shows that the cross-terms discretized using the SW-TRIADS scheme always provide a less restrictive stability constraint than when using the TRIADS scheme (which corresponds to $\beta_m = 0$ in (3.14)). It means that the condition $\tilde{\lambda} \geq -1$ is more difficult to satisfy with the TRIADS discretization. We thus consider in the remainder of this section the most restrictive case $\beta_m = 0$. We provide in App. B the results obtained for a non-zero value of β_m with the SW-TRIADS discretization.

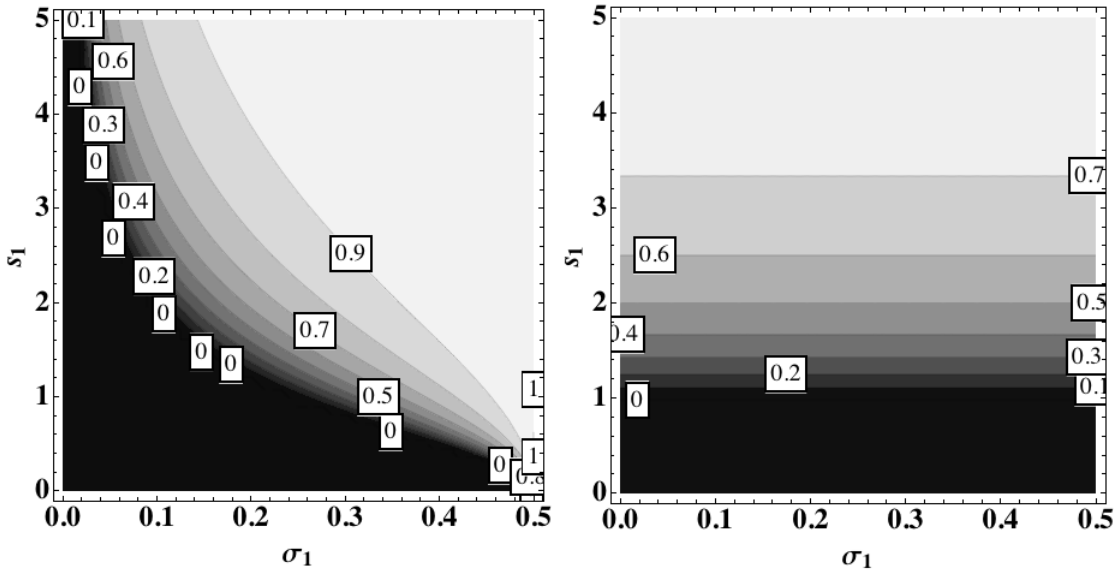


Fig. 3. Optimized values θ^* (for $\sigma_2 = 0$) of the stabilizing parameter θ obtained with the TRIADS scheme (left) and the SW-TRIADS scheme (right) with respect to σ_1 and s_1 .

Stability Conditions

Starting from (3.15), stability is obtained for

$$\tilde{\lambda} \geq -1 \Leftrightarrow \sum_{m=1}^2 \varphi_m - 2(1 + \theta y_3^2) \leq 0, \quad \varphi_m = y_m^2 + y_{3m}^2 - z_{3m} - z_{m3}. \quad (3.21)$$

Taking $s_m = 0$ (i.e. $y_{3m}=z_{3m}=0$) and $\theta = 0$ in (3.21) we find the stability constraint $y_1^2 + y_2^2 \leq 2$ which corresponds to the usual constraint (3.5) obtained when the horizontal Laplacian operator is integrated using an Euler explicit scheme. Ideally, this is the stability constraint we are targeting for the general case $s_m \neq 0$ so that the rotated Laplacian operator does not involve the use of a smaller time step than the non-rotated Laplacian operator. The stability condition (3.21) can be more conveniently written as

$$\varphi_1^\theta + \varphi_2^\theta \leq 2, \quad \text{with } \varphi_m^\theta = \varphi_m - 2\theta y_{m3}^2. \quad (3.22)$$

The stability analysis therefore requires the derivation of a proper upper bound for the φ_m^θ functions. For clarity, we describe the computation of such an upper bound in App. A, and we recall here only the final result :

$$\varphi_1^\theta + \varphi_2^\theta \leq 2\sigma_1 \mathcal{M}(\theta, s_1) + 2\sigma_2 \mathcal{M}(\theta, s_2) \quad (3.23)$$

where

$$\mathcal{M}(\theta, s_m) = 1 + s_m^2(1 - 2\theta) + \sqrt{(1 + s_m^2)(1 + s_m^2(1 - 2\theta)^2)}. \quad (3.24)$$

Combining (3.22) and (3.23) we get the following criteria for stability,

$$2\sigma_1 \mathcal{M}(\theta, s_1) + 2\sigma_2 \mathcal{M}(\theta, s_2) \leq 2. \quad (3.25)$$

In the fully-explicit (EXP) case (i.e. $\theta = 0$) we get $\mathcal{M}(0, s_m) = 2(1 + s_m^2)$, the associated stability criteria reads

$$\sigma_1(1 + s_1^2) + \sigma_2(1 + s_2^2) \leq \frac{1}{2}. \quad (3.26)$$

This result leads to the stability condition (3.6) and is consistent with the result found in Mathieu et al. (1999) (eqn. 31). Furthermore, for $\theta = 1$ (i.e.; the (IMP) scheme), we get $\mathcal{M}(1, s_m) = 2$ and thus $\sigma_1 + \sigma_2 \leq 1/2$ which corresponds to the stability limit of the non-rotated operator. This result shows that by treating implicitly the vertical component of the tensor there is no additional constraint arising from the cross-terms. Finally, for the (MSC) scheme the optimal value of θ providing a minimum of implicit diffusion is solution of the equation $\sigma_1 \mathcal{M}(\theta, s_1) + \sigma_2 \mathcal{M}(\theta, s_2) = 1$. However, the analytical solution of this equation is extremely complex and would be impractical anyway for real applications. We, therefore, proceed in a more conservative way by noting that for $\theta \leq 1$

$$\mathcal{M}(\theta, s_m) \leq \mathcal{M}^\ddagger(\theta, s_m) = 2(1 + s_m^2) - 2\theta s_m^2. \quad (3.27)$$

It is now straightforward to determine the analytical solution θ^* of the equation $\sigma_1 \mathcal{M}^\ddagger(\theta, s_1) + \sigma_2 \mathcal{M}^\ddagger(\theta, s_2) = 1$:

$$(s_1^2 \sigma_1 + s_2^2 \sigma_2) \theta^* = \frac{1}{2} \max \left\{ -1 + 2(1 + s_1^2) \sigma_1 + 2(1 + s_2^2) \sigma_2, 0 \right\}. \quad (3.28)$$

We can check that $\theta^* = 0$ if the (EXP) scheme is stable and $\theta^* = 1$ for $\sigma_1 + \sigma_2 = 1/2$. In Fig. 3 we show the values of the stabilizing parameter θ^* associated with the TRIADS scheme as well as the values of the stabilizing parameter obtained in App. B with the SW-TRIADS. We, first, want to emphasize the fact that those values of θ are sufficient conditions for stability but are not the optimal values, conservative choices have been made during our analysis. This remark explains why in Fig. 3 larger values of θ are shown for the SW-TRIADS, compared to the TRIADS (particularly for small values of σ_m), although we argue earlier that the TRIADS scheme always provide a more restrictive stability condition than the SW-TRIADS scheme. It can be shown that the value of θ given in (3.28) for the TRIADS scheme is also a sufficient condition for stability with the SW-TRIADS scheme.

In general, we see that the value of θ required to maintain stability of the scheme can be smaller than 1 in numerous cases and hence that splitting errors (as defined in footnote 2) associated with the (IMP) scheme can be further reduced. It, however, still needs to be checked that this reduction of splitting errors has a clear and meaningful impact on the physical solution of numerical models. This question is left for a future study.

3.4 Non-oscillatory Scheme

Throughout the previous subsection we have considered the requirement $|\tilde{\lambda}| \leq 1$ which is a necessary condition for stability. However a more severe condition $0 \leq \tilde{\lambda} \leq 1$ may be required to ensure a non-oscillatory behavior of the temporal integration (Mathieu et al., 1999). This corresponds to the so-called "no flip-flop" condition, as opposed to the condition $|\tilde{\lambda}| \leq 1$ which allows a flip-flop behavior in time unlike the exact solution of the problem under investigation (Cushman-Roisin and Beckers, 2011, Chap. 5). Indeed, we see in (3.13) that negative values of $\tilde{\lambda}$ would allow the solution to oscillate at each time step because $\tilde{\lambda}, \tilde{\lambda}^2, \tilde{\lambda}^3, \dots$ would change sign. It is straightforward to extend the results described in Sec. 3.3 to the "no flip-flop" case. In the non-rotated case, (3.5) reduces to

$$\sigma_1 + \sigma_2 \leq 1/4, \quad (3.29)$$

which means that the time step must be divided by two. In the rotated case, the (EXP) scheme satisfies the condition $\tilde{\lambda} \geq 0$ for

$$\sigma_1(1 + s_1^2) + \sigma_2(1 + s_2^2) \leq \frac{1}{4}. \quad (3.30)$$

For the (IMP) scheme, the absence of flip-flop requires

$$\sigma_1 \left(1 + \sqrt{1 + s_1^2}\right) + \sigma_2 \left(1 + \sqrt{1 + s_2^2}\right) \leq \frac{1}{2} \quad (3.31)$$

which is the condition used in the CLIO model to set Δt (Mathieu et al., 1999). The inequality (3.31) shows that the no flip-flop condition is satisfied under a stability constraint more restrictive than (3.29). In the case of the (MSC) scheme, the no flip-flop condition is achieved if (3.29) is satisfied and if θ is chosen such that

$$(s_1^2\sigma_1 + s_2^2\sigma_2)\theta^* = \frac{1}{2} \max \left\{ -1 + 4(1 + s_1^2)\sigma_1 + 4(1 + s_2^2)\sigma_2, 0 \right\}. \quad (3.32)$$

For this value of θ the scheme satisfies the "no flip-flop" condition if the horizontal terms do, which reduces to (3.29). In (3.32), the value of θ^* varies from $\theta^* = 0$ when (3.30) is satisfied to $\theta^* = 2$ when $\sigma_1 + \sigma_2 = 1/4$. The aim here is not to claim that flip-flops should not be allowed. Indeed, the widely used horizontal Laplacian operator advanced with a forward Euler scheme allows flip-flops for $1/4 \leq \sigma_1 + \sigma_2 \leq 1/2$ and it did not turn out to be problematic in practice. This paragraph was meant to illustrate that, unlike the (IMP) scheme which is restricted by the cross-terms, the (MSC) scheme is general enough to ensure additional properties of the time discretization like the absence of flip-flops.

3.5 Comments

The literature on multi-dimensional parabolic problems with mixed derivatives has been focused on demonstrating the stability of the method of stabilizing corrections (and/or the ADI method). To our knowledge, there are no systematic studies on the impact of the stabilizing step on the accuracy of the solution. This point is discussed in this section and turns out to be very helpful to anticipate the behavior of the scheme in practical situations.

3.5.1 Spatial Discretization

In Fig. 4 we show the amplification factor obtained when the (EXP) scheme is stable. Because in this case there are no splitting errors, this figure is indicative of the differences between the two spatial discretizations under consideration. To make the interpretation of the results easier, we consider the no flip-flop case. In the figure, we can not formally identify the direction of the computational grid because it represents a wavenumber space, but it is instructive to look at the amplification factor in several directions. Using (3.10), we find that ∇q and (k_1, k_3) are collinear. Moreover it is straightforward to see that the dianeutral direction and $\rho_{\perp} = (-\alpha_1, 1)$ are collinear too, as well as the isoneutral direction and $\rho_{\parallel} = (1, \alpha_1)$. We specifically look at two directions in the (ϕ_1, ϕ_3) plane corresponding to two different angles between ∇q and the isoneutral/dianeutral direction :

- $D_1 = \phi_1 + s_1\phi_3 = 0$ (i.e.; $\nabla q \cdot \rho_{\parallel} = 0$) corresponds to the direction along which ∇q is perpendicular to the isoneutral direction. In this case, q is constant in the isoneutral direction (because the iso- q lines are perpendicular to ∇q). Because an isoneutral mixing operator should not affect a tracer constant in the

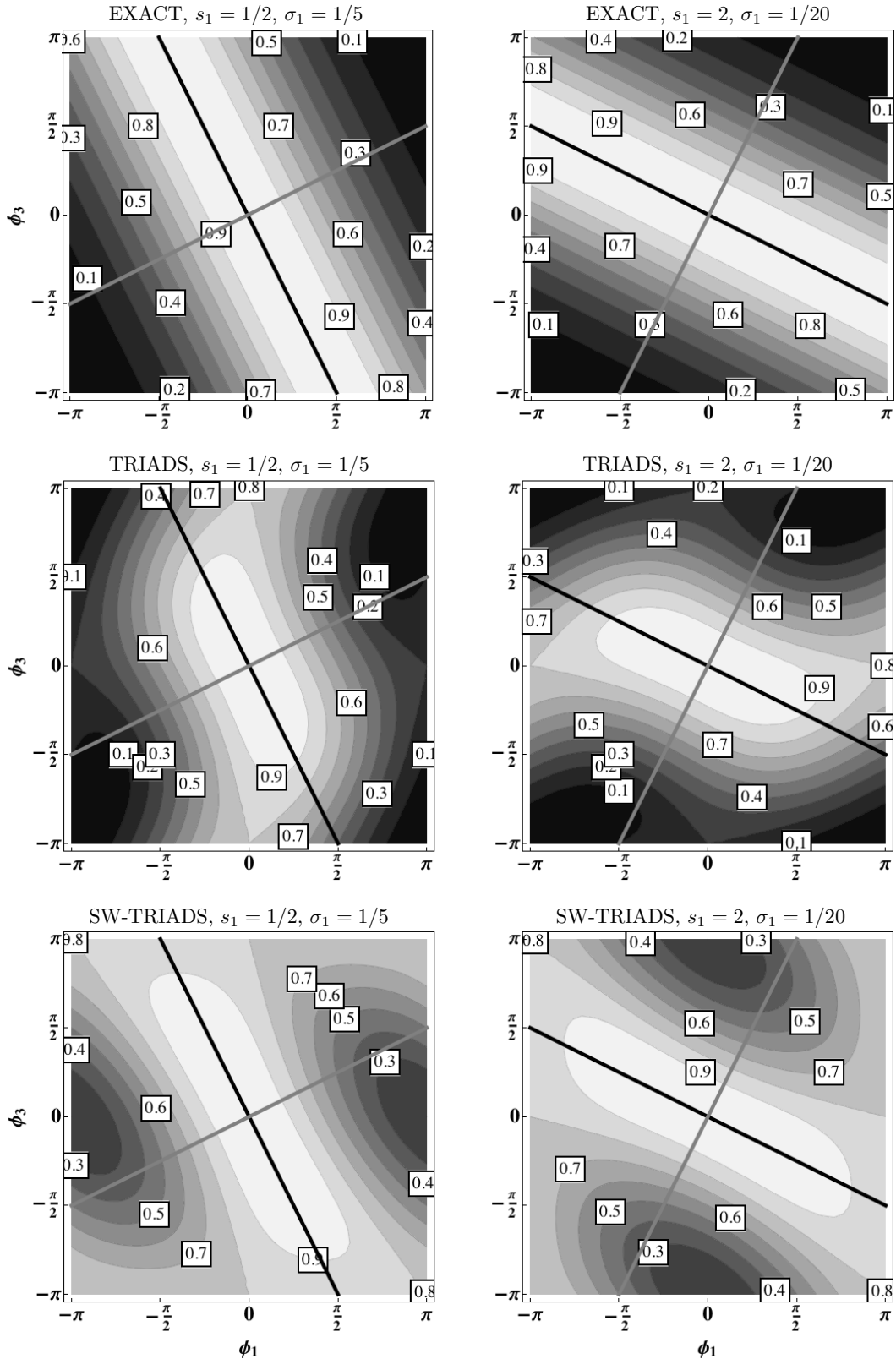


Fig. 4. Exact amplification factor in the (ϕ_1, ϕ_3) plane for $\sigma_1 = 1/4(1 + s_1^2)$ (i.e.; when the (EXP) scheme is stable) with $s_1 = 1/2$ (top, left) and $s_1 = 2$ (top, right) in the two-dimensional (x_1, x_3) case (i.e. $\kappa_2 = 0$). Amplification factor for the (EXP) scheme with the TRIADS discretization (middle, left for $s_1 = 1/2$ and middle, right for $s_1 = 2$) and with the SW-TRIADS discretization (bottom, left for $s_1 = 1/2$ and bottom, right for $s_1 = 2$) for the same values of σ_1 . The thick black line is $D_1 = 0$, the thick gray line is $D_2 = 0$.

isoneutral direction, we expect the amplification factor λ to be equal to one along this line, otherwise it would indicate that the slope is computed in an inaccurate way and that dianeutral mixing occurs.

- A perpendicular to D_1 is $D_2 = \phi_3 - s_1\phi_1 = 0$ (i.e.; $\nabla q \cdot \boldsymbol{\rho}_\perp = 0$) which corresponds to the direction along which ∇q is aligned with the isoneutral direction. In this case, we see from Fig. 4 (top panels) that the exact amplification factor is equivalent to the one of a one-dimensional Laplacian operator, indeed we would have $\eta\Delta t = \sigma_1(1 + s_1^2)^2\phi_1^2$.

Along the line $D_1 = 0$, the SW-TRIADS always provide a damping in better agreement with the exact one, compared to the TRIADS scheme, thus indicating a more accurate computation of the direction of diffusion for $s_1 = 1/2$ as well as $s_1 = 2$. We see that, for $\phi_3 = 0$, all the discretizations have the same behavior and that the mode $|\phi_1| = \pi$ is effectively damped. This was expected because the difference between the TRIADS and SW-TRIADS schemes is in the cross-terms which vanish for $\phi_3 = 0$. Unlike the SW-TRIADS, the TRIADS scheme has also the property to efficiently damp the checkerboard mode $(|\phi_1|, |\phi_3|) = (\pi, \pi)$, whatever the grid slope ratio (Fig. 4). The lack of damping associated with the SW-TRIADS is not problematic to solve the initial value problem (2.6) because we assume some regularity of the initial condition, and the checkerboard mode is not present in this case. However, when advective terms are considered they can allow the creation and accumulation of dispersive errors which are expected to be controlled by diffusive processes. We usually do not rely on the rotated operator to damp small scale noise in the vertical direction because the vertical mixing parameterization does it efficiently (otherwise grid-scale noise would arise when an horizontal Laplacian operator is used). In the horizontal direction, those operators are generally the only source of numerical filtering, and we thus expect them to control small-scale noise. For example, for $s_1 = 1/2$, the TRIADS scheme does it efficiently while the SW-TRIADS scheme damps the $2\Delta x_1$ mode ($|\phi_1| = \pi$) only for well-resolved scales in the vertical ($|\phi_3| \leq \pi/2$). Note that it is, however, expected that the $2\Delta x_3$ mode ($|\phi_3| = \pi$) is already significantly damped by the vertical mixing scheme. The weak damping with the SW-TRIADS scheme can be seen also along the line $D_2 = 0$. In this direction, the exact amplification factor and the TRIADS scheme ensure a monotonic damping, indeed λ monotonically decreases when we go from well-resolved to poorly-resolved scales (i.e. when $|\phi_1|$ or $|\phi_3|$ increases). This property is not satisfied by the operator discretized with the SW-TRIADS. We thus expect the TRIADS to provide a better control of numerical noise and the SW-TRIADS to provide a more accurate computation of the direction of diffusion. Moreover, it is worth mentioning that numerical noise in the tracer fields can project into irreversible dianeutral mixing errors, especially when local Richardson number-dependent vertical mixing schemes are used in the oceanic interior.

3.5.2 Space-Time Discretization

We show in Fig. 5 the amplification factor for $s_1 = 1/2$ and $s_1 = 2$ when the (MSC) scheme is used with values of θ chosen such that there are no flip-flops in time. If we note Δt_ρ the time step of the (EXP) scheme and Δt_0 the time step of the (MSC) scheme, we get the ratio μ_2 defined as

$$\mu_2 = \frac{\Delta t_0}{\Delta t_\rho} = 1 + s_1^2. \quad (3.33)$$

For $s_1 = 1/2$, we have $\mu_2 = 5/4$, and $\mu_2 = 5$ for $s_1 = 2$. When μ_2 is close to one, as for $s_1 = 1/2$, splitting errors are very small. We see that the left panels in Fig. 5 are very similar to the left panels in Fig. 4. For $\mu_2 = 5$, the damping along the line $D_1 = 0$ increases which indicates that the computation of the direction of diffusion become less and less accurate (Fig. 5, right panels). The space-time discretization is also relatively inaccurate along the line $D_2 = 0$ and the SW-TRIADS scheme suffers from non-monotonic damping. However, the schemes still perform well as long as the vertical scales are well-resolved (i.e.

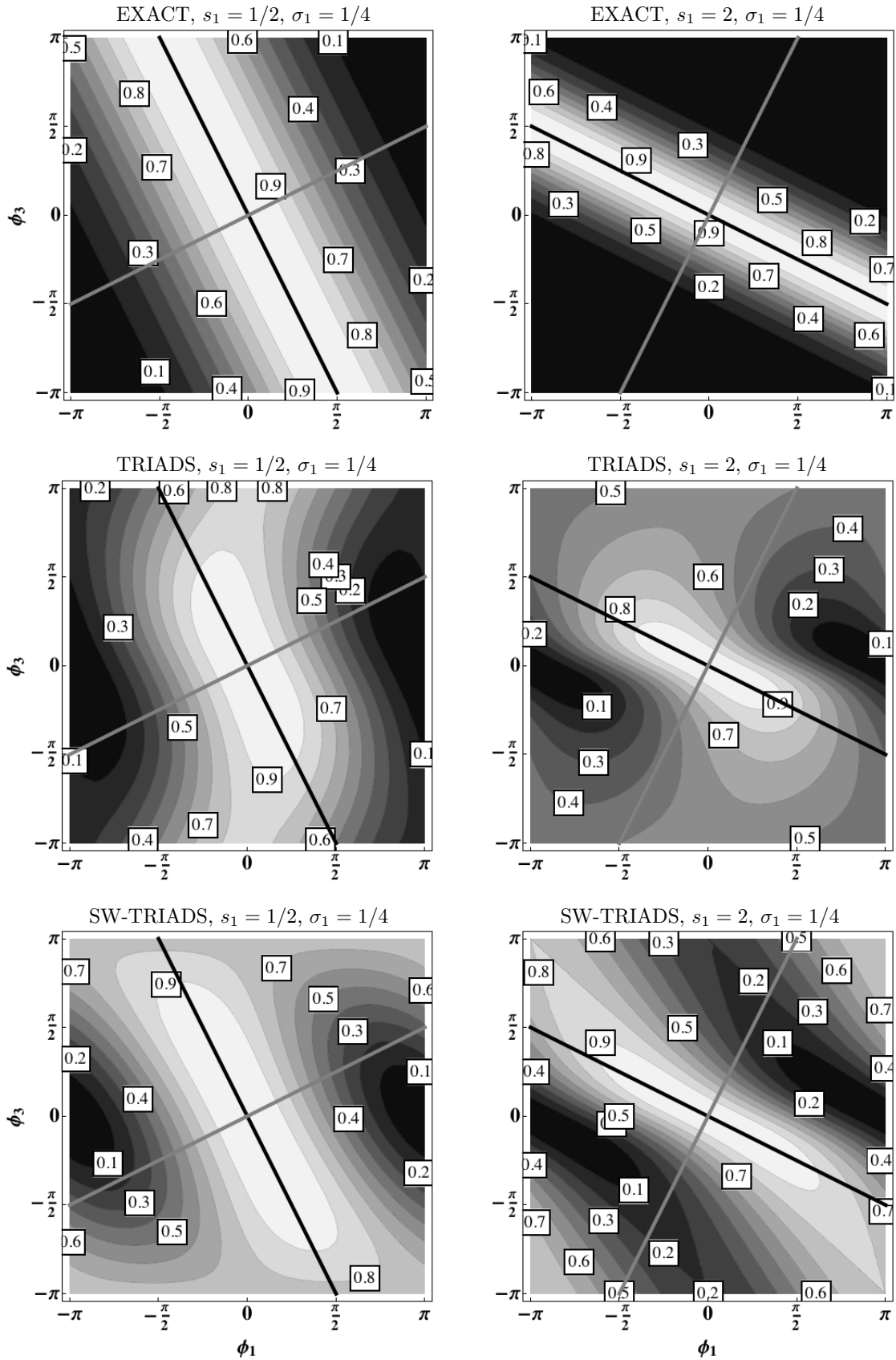


Fig. 5. Same as Fig. 4 for the (MSC) scheme with $\sigma_1 = 1/4$ (i.e.; when the (EXP) scheme does not satisfy the no flip-flop condition).

$$|\phi_3| < \pi/2).$$

4 Time Discretization of the Isonutral Biharmonic Operator

4.1 Proposed Scheme

As mentioned in the paper's introduction, the isoneutral biharmonic operator has never been thoroughly studied in the literature so far. One possible explanation is that people working on rotated operators were mainly interested in large-scale flows for which the scale-selectivity property is not a priority. Moreover, a very drastic reduction of the time step is required to maintain stability of such an operator. The aim of the present section is to tackle this last point by deriving a time integration scheme enabling an efficient and easy implementation of a rotated biharmonic operator. To our knowledge, no alternative to the usual explicit Euler scheme has been proposed. It is of course feasible to mimic the (IMP) scheme used for the rotated Laplacian operator but it would require the solution of a penta-diagonal system, and additional constraints would arise from the cross terms anyway. Indeed, an implicit vertical biharmonic operator can not stabilize the second order terms in x_3 . Even if it is relatively counter-intuitive we show in the remainder of this section that it is possible to stabilize a biharmonic operator by means of a simple Laplacian operator. An alternative approach could be to use a combination of an implicit Laplacian operator and an implicit biharmonic operator in the vertical to stabilize separately the second-order and fourth-order terms, but it would increase substantially the complexity of the scheme.

In the same spirit as the method of stabilizing corrections introduced in Sec. 3.1, we thus propose to integrate the rotated biharmonic operator \mathcal{D}_4 in the following way :

$$\begin{cases} q^* = q^n + \Delta t \mathcal{D}_4(q^n) \\ q^{n+1} = q^* + \Delta t \partial_3 \{ \tilde{\kappa} \partial_3 q^{n+1} - \tilde{\kappa} \partial_3 q^n \}, \end{cases} \quad (4.1)$$

In (4.1), an explicit and consistent approximation of the operator is first computed, followed by a second stage aiming at improving the stability. For the aforementioned reasons, we decided to use a second-order diffusive operator for this second stage. Instead of the stabilizing parameter θ used in the previous section, we use here a stabilizing vertical diffusivity $\tilde{\kappa}$.

4.2 Important Results

We first provide the results and prove them in Sec. 4.3. We keep the notations from the previous section but now with σ_m replaced by

$$\sigma_m^{(4)} = \frac{\sqrt{\Delta t B_m}}{\Delta x_m^2}, \quad (4.2)$$

where B_m is the hyperdiffusivity in the x_m direction.

Horizontal Biharmonic Operator

When using an Euler forward scheme, the stability limit of the two-dimensional horizontal biharmonic operator is

$$\left(\sigma_1^{(4)} + \sigma_2^{(4)}\right)^2 \leq \frac{1}{8}. \quad (4.3)$$

This result is also given in Griffies (2004) (Chap. 18).

Isonutral Biharmonic Operator

• TRIADS discretization

- The (EXP) scheme is stable for

$$\left(\sigma_1^{(4)}(1 + s_1^2) + \sigma_2^{(4)}(1 + s_2^2)\right)^2 \leq \frac{1}{8}. \quad (4.4)$$

- The stability constraint of the (MSC) scheme is the same as the stability limit (4.3) for

$$\tilde{\kappa} = 8 \left(\frac{\Delta x_3^2}{\Delta t}\right) \left(\sigma_1^{(4)} s_1^2 + \sigma_2^{(4)} s_2^2\right) \left((1 + s_1^2)\sigma_1^{(4)} + (1 + s_2^2)\sigma_2^{(4)}\right) \quad (4.5)$$

• SW-TRIADS discretization

- The (EXP) scheme is stable for

$$\left(\sigma_1^{(4)} \max\{s_1^2, 1\} + \sigma_2^{(4)} \max\{s_2^2, 1\}\right)^2 \leq \frac{1}{8}. \quad (4.6)$$

- The stability constraint of the (MSC) scheme is given by (4.3) for

$$\tilde{\kappa} = 8 \left(\frac{\Delta x_3^2}{\Delta t}\right) \left(S_1 \sigma_1^{(4)} + S_2 \sigma_2^{(4)}\right) \left((1 + S_1)\sigma_1^{(4)} + (1 + S_2)\sigma_2^{(4)}\right), \quad (4.7)$$

with $S_m = \max\{s_m^2 - |s_m|, 0\}$.

Note that, if we take $S_m = s_m^2$ in (4.7), we retrieve (4.5). Using the notations introduced in Sec. 2.2, we could write in a generic way $S_m = s_m(s_m - \beta_m)$, thus showing that S_m corresponds to the weight of the vertical points $(i, j, k + 1)$ and $(i, j, k - 1)$ in the stencil (2.21). We now give the proof of the stability results.

4.3 Linear Stability Analysis

We, first, define a vertical Courant number $\tilde{\sigma} = \tilde{\kappa} \Delta t / \Delta x_3^2$ associated with the stabilizing stage in (4.1), and we introduce $\tilde{y}_3 = \sqrt{2\tilde{\sigma}(1 - \cos \phi_3)}$. The aim of this stability analysis is to determine a value of $\tilde{\sigma}$ (and hence $\tilde{\kappa}$) ensuring that the stability constraint of the rotated biharmonic operator is imposed

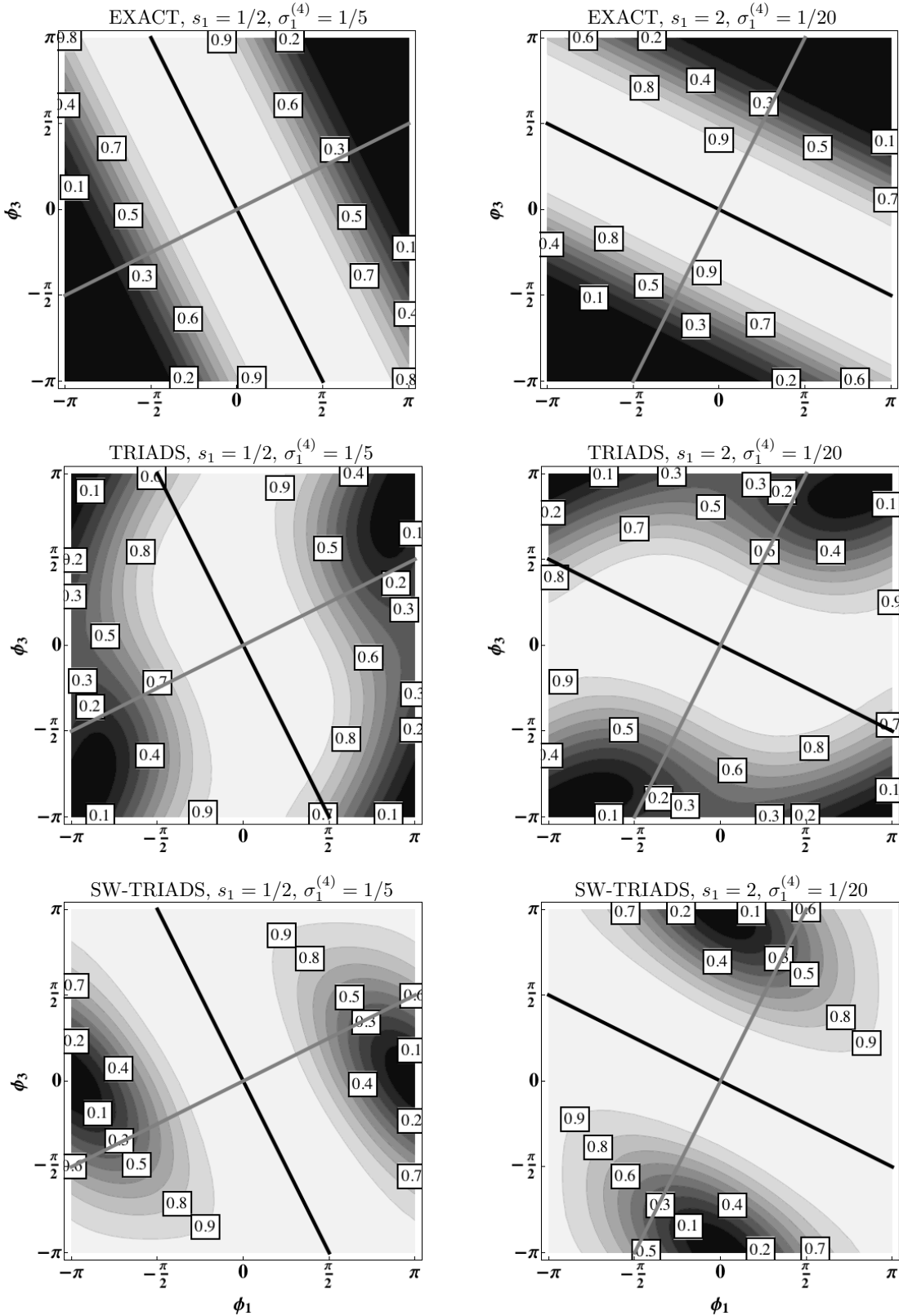


Fig. 6. Exact amplification factor in the (ϕ_1, ϕ_3) plane for $\sigma_1^{(4)} = \sqrt{1/16(1 + s_1^2)^2}$ (i.e.; when the (EXP) scheme is stable) with $s_1 = 1/2$ (top, left) and $s_1 = 2$ (top, right) in the two-dimensional (x_1, x_3) case (i.e. $\kappa_2 = 0$). Amplification factor for the (EXP) scheme with the TRIADS discretization (middle, left for $s_1 = 1/2$ and middle, right for $s_1 = 2$) and with the SW-TRIADS discretization (bottom, left for $s_1 = 1/2$ and bottom, right for $s_1 = 2$) for the same values of σ_1 . The thick black line is $D_1 = 0$, the thick gray line is $D_2 = 0$.

by the horizontal components of the operator. Note that the exact amplification factor is now given by $\lambda = \exp(-\eta\Delta t)$ with $\eta = \sum_{m=1}^2 B_m(k_m + \alpha_m k_3)^4$, while the amplification factor after discretization is

$$\tilde{\lambda} = 1 - \frac{(y_1^2 + y_2^2 + y_3^2 - z_{13} - z_{31} - z_{23} - z_{32})^2}{1 + \tilde{y}_3^2} = 1 - \frac{(\varphi_1 + \varphi_2)^2}{1 + \tilde{y}_3^2}, \quad (4.8)$$

where φ_m is defined in (3.21). It is straightforward to check that $\tilde{\lambda} \leq 1$, stability is thus obtained for $\tilde{\lambda} \geq -1$; this condition is equivalent to

$$(\varphi_1 + \varphi_2)^2 - 2\tilde{y}_3^2 \leq 2. \quad (4.9)$$

We can first easily check that the stability limit of the horizontal biharmonic operator (i.e., with $s_m = 0$) is given by $(y_1^2 + y_2^2)^2 \leq 2$, leading to $(\sigma_1^{(4)} + \sigma_2^{(4)})^2 \leq 1/8$. Furthermore, using the upper bound on φ_m found in App. A, we can derive the stability constraint for the forward Euler scheme (i.e., $\tilde{\kappa} = 0$),

$$(\varphi_1 + \varphi_2)^2 \leq 4 \left(\sigma_1^{(4)} \mathcal{M}(0, s_1) + \sigma_2^{(4)} \mathcal{M}(0, s_2) \right)^2 = 16 \left(\sigma_1^{(4)}(1 + s_1^2) + \sigma_2^{(4)}(1 + s_2^2) \right)^2 \leq 2, \quad (4.10)$$

where \mathcal{M} is defined in (3.24); (4.10) leads to the expected result (4.4). As an illustration, this stability constraint implies that for $s_1 = s_2 = 2$, the time step used with the horizontal biharmonic operator should be reduced by a factor of 25 when using the rotated biharmonic operator.

In the general case $\tilde{\kappa} \neq 0$, we again use the upper bound on φ_m to get

$$(\varphi_1 + \varphi_2)^2 - 2\tilde{y}_3^2 \leq \mathcal{B}(\theta) = 4 \left(\sigma_1^{(4)} \mathcal{M}(\theta, s_1) + \theta y_{13}^2 + \sigma_2^{(4)} \mathcal{M}(\theta, s_2) + \theta y_{23}^2 \right)^2 - 2\tilde{y}_3^2, \quad (4.11)$$

stability is obtained for $\mathcal{B}(\theta) \leq 2$. For simplicity, we take $\theta = 1$ (i.e. $\mathcal{M} = 2$)

$$\begin{aligned} \mathcal{B}(\theta = 1) &= 16 \left(\sigma_1^{(4)} + \sigma_2^{(4)} + (\sigma_1^{(4)} s_1^2 + \sigma_2^{(4)} s_2^2)(1 - \cos \phi_3) \right)^2 - 4\tilde{\sigma}(1 - \cos \phi_3) \\ &= 16 \left(\sigma_1^{(4)} + \sigma_2^{(4)} \right)^2 \\ &\quad + (1 - \cos \phi_3) \left\{ 16(\sigma_1^{(4)} s_1^2 + \sigma_2^{(4)} s_2^2) \left[2(\sigma_1^{(4)} + \sigma_2^{(4)}) + (\sigma_1^{(4)} s_1^2 + \sigma_2^{(4)} s_2^2)(1 - \cos \phi_3) \right] - 4\tilde{\sigma} \right\}. \end{aligned} \quad (4.12)$$

The stability constraint $\mathcal{B}(\theta = 1) \leq 2$ thus reduces to $(\sigma_1^{(4)} + \sigma_2^{(4)})^2 \leq 1/8$ if the term in curly brackets in (4.12) is negative or zero. This requirement is satisfied for

$$\tilde{\sigma} \geq \sigma^* = 8(\sigma_1^{(4)} s_1^2 + \sigma_2^{(4)} s_2^2) \left((1 + s_1^2)\sigma_1^{(4)} + (1 + s_2^2)\sigma_2^{(4)} \right), \quad (4.13)$$

which demonstrates (4.5). Less conservative values of $\tilde{\sigma}$ could be derived (like the one used in Lemarié et al. (2012)), however the algebra becomes very quickly complicated and tedious. The optimal value (4.7) of the stabilizing parameter $\tilde{\sigma}$ obtained with the SW-TRIADS scheme is provided in App. B.

Non-Oscillatory Scheme

As for the Laplacian operator, we can derive the necessary constraints to satisfy $\tilde{\lambda} \geq 0$ instead of $\tilde{\lambda} \geq -1$. In this case, $1/8$ should be replaced by $1/16$ in (4.3), (4.4), and (4.6). Furthermore, if the

stabilizing diffusivity (4.5) is multiplied by two, the (MSC) scheme satisfies the "no flip-flop" condition for $(\sigma_1^{(4)} + \sigma_2^{(4)})^2 \leq 1/16$.

4.4 Comments

We show in Fig. 6 and 7 the amplification factor obtained with a rotated biharmonic operator for $s_1 = 1/2$ and $s_1 = 2$ with the (EXP) and the (MSC) schemes in the no flip-flop case. We can draw the same remarks as for the Laplacian operator: the SW-TRIADS computes more accurately the direction of diffusion, however it suffers from a lack of damping of the checkerboard mode. The ratio μ_4 between the time step of the (MSC) scheme and the one of the (EXP) scheme is given by

$$\mu_4 = (1 + s_1^2)^2. \quad (4.14)$$

For $s_1 = 1/2$, we get $\mu_4 \approx 1.56$. In this case, the results are very similar between the (MSC) and the (EXP) schemes. For $s_1 = 2$, the ratio becomes stiffer with $\mu_4 = 25$. However, the scheme behaves in a similar manner (Fig. 7, right panels) than the (EXP) scheme (Fig. 6, right panels).

4.5 Partial Conclusion and Limitations of Our Approach

We have shown so far that the Laplacian and biharmonic rotated operators can be made stable by combining an explicit-in-time evaluation with a semi-implicit stabilizing step. This approach has the advantage to be extremely easy to implement because it requires only the inversion of a simple tridiagonal system in the vertical. For quite large grid slope ratios ($s_1 = 2$) the scheme performs relatively well but as the slope steepens to $s_1 = 10$ (Fig. 8) it can suffer from a lack of damping of the smallest resolved scales, especially for the rotated biharmonic operator. This behavior is reminiscent of the Crank-Nicolson scheme for which the small scale structures are not damped in the limit of large time-steps (Manfredi and Ottaviani, 1999). This problem thus arises when the time step used with the (MSC) scheme is significantly larger than the stability limit of the (EXP) scheme, i.e.; when μ_4 or μ_2 are large (for $s_1 = 10$ we have $\mu_4 \approx 10000$). This lack of damping affects only the vertical direction. The horizontal terms of the tensor, which are advanced explicitly, are still "active". Because the TRIADS scheme requires a larger value of $\tilde{\kappa}$ to stabilize the time-integration, this lack of damping is generally more pronounced with this scheme. This issue is also more obvious when we try to suppress the flip-flops in time by increasing the value of $\tilde{\kappa}$ (Fig. 8).

The fact that we restrict ourselves to implicit integrations only in the vertical direction reduces significantly the range of possible methods to try to mitigate this problem. If steep grid slope ratios are expected and accuracy matters, we could combine the (MSC) with a time-splitting approach, not to penalize the whole model. When using an Euler explicit scheme, if s_1 reaches $\mathcal{O}(10)$ values we should decrease the time step by a factor 10^4 compared to the time step of the horizontal biharmonic operator (indeed, $\mu_4 \approx 10000$ for $s_1 = 10$). Several sub-steps of the (MSC) scheme to advance the rotated operator from time n to $n + 1$ would still be significantly more computationally efficient than using an explicit Euler scheme. The use of a small number of sub-step clearly improves the damping properties of the rotated biharmonic operator which smoothes more efficiently the poorly resolved scales (Fig. 9). Based on the results shown in Fig. 9 we can recommend the use of 4 sub-steps to maintain a good accuracy of the rotated operator for steep grid slope ratios. We checked that for very large values of s_1 ($s_1 = 10^6$), 4 sub-steps are still sufficient to damp the smallest scales (not shown). Note that an even number of sub-steps ensures the absence of flip-flops.

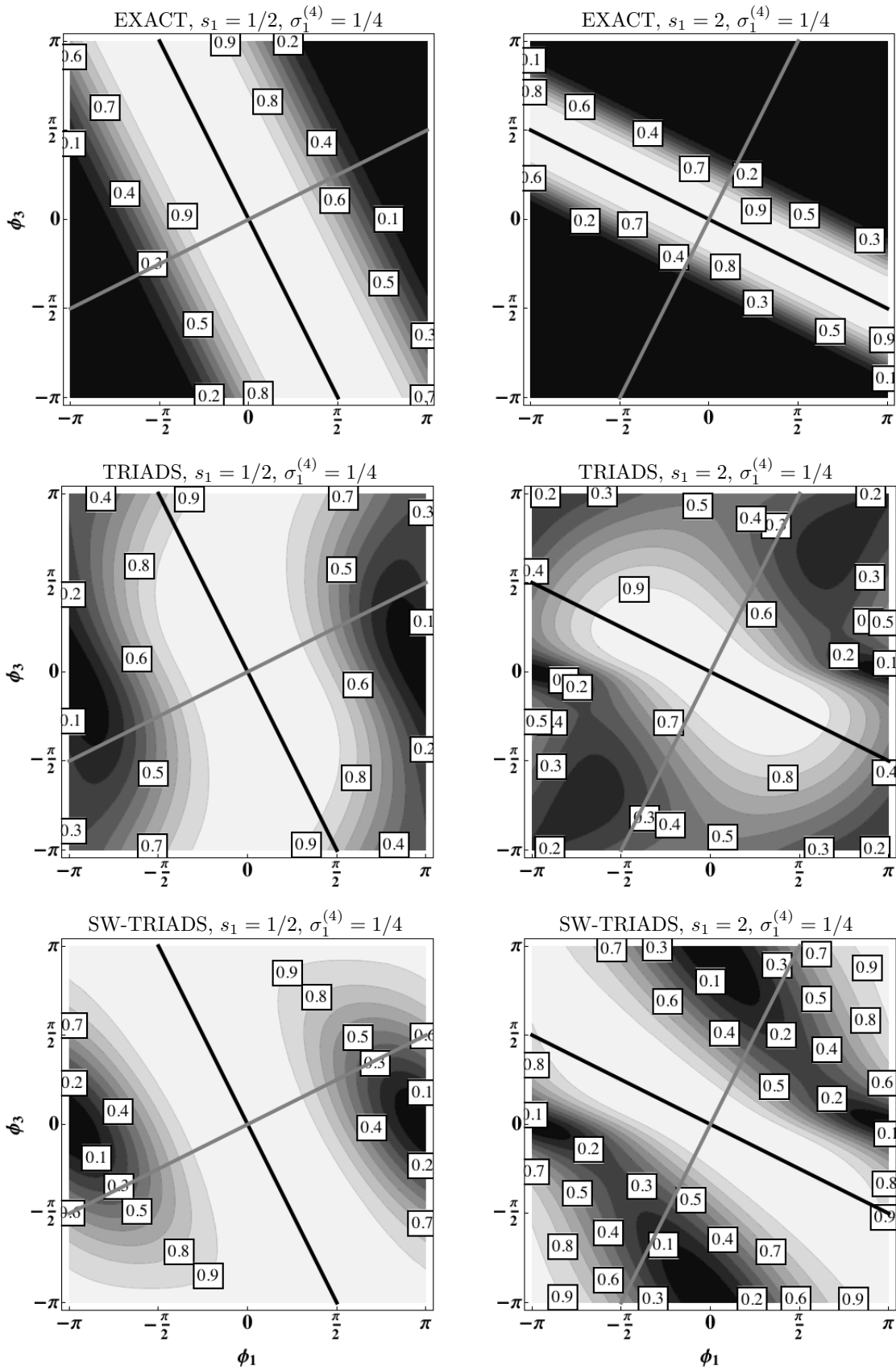


Fig. 7. Same as Fig. 6 for the (MSC) scheme with $\sigma_1^{(4)} = 1/4$ (i.e.; when the (EXP) scheme does not satisfy the no flip-flop condition).

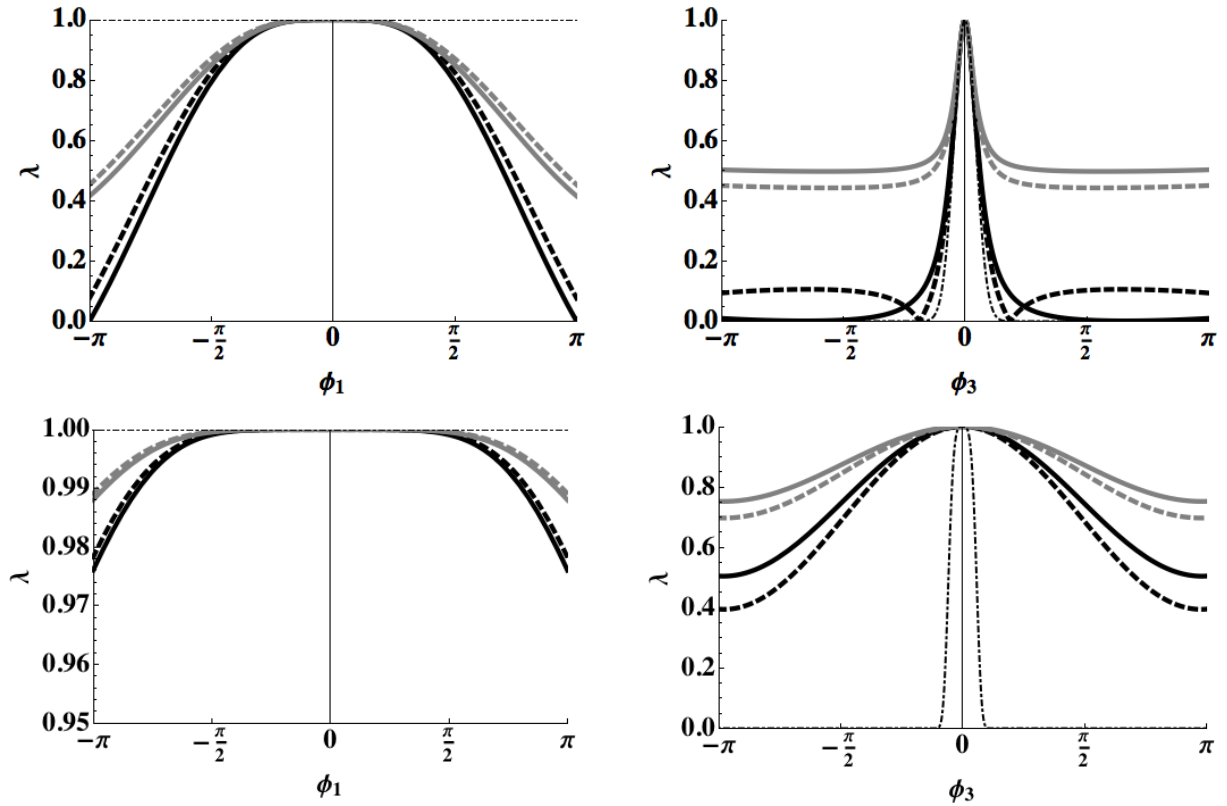


Fig. 8. Amplification factor of the rotated Laplacian operator along the line $D_1 = 0$ (top, left) and $D_2 = 0$ (top, right), and of the rotated biharmonic operator along $D_1 = 0$ (bottom, left) and $D_2 = 0$ (bottom, right) for $s_1 = 10$, $\sigma_1 = 1/4$ and $\sigma_1^{(4)} = 1/4$. The gray lines (resp. black lines) are obtained with a stabilizing diffusivity preventing flip-flops (resp. allowing flip-flops) with the (MSC) scheme. The results obtained with the TRIADS (resp. SW-TRIADS) discretization are represented with thick lines (resp. thick dashed lines). The thin black lines correspond to the exact amplification.

5 Implementation of the Proposed Schemes in Existing Models

In this section, we provide additional information required for the implementation of the various schemes introduced in the previous sections.

5.1 Time-Scheme for Diffusive Processes

As mentioned earlier, the majority of the global ocean climate models makes use of a rotated Laplacian operators advanced with the (IMP) scheme. For those models, the implementation of the (MSC) scheme is straightforward and does not require any additional global array for storage. In numerical models like ROMS or MitGCM⁵ using a predictor-corrector scheme where provisional values $q^{n+\frac{1}{2}}$ are given by a predictor step (e.g. a Leapfrog-Adams-Moulton interpolation in ROMS or an Adams-Bashforth extrapolation in MitGCM), the (MSC) scheme and the implicit vertical diffusion can be combined with the corrector step in the following way :

⁵ Massachusetts Institute of Technology General Circulation Model, mitgcm.org

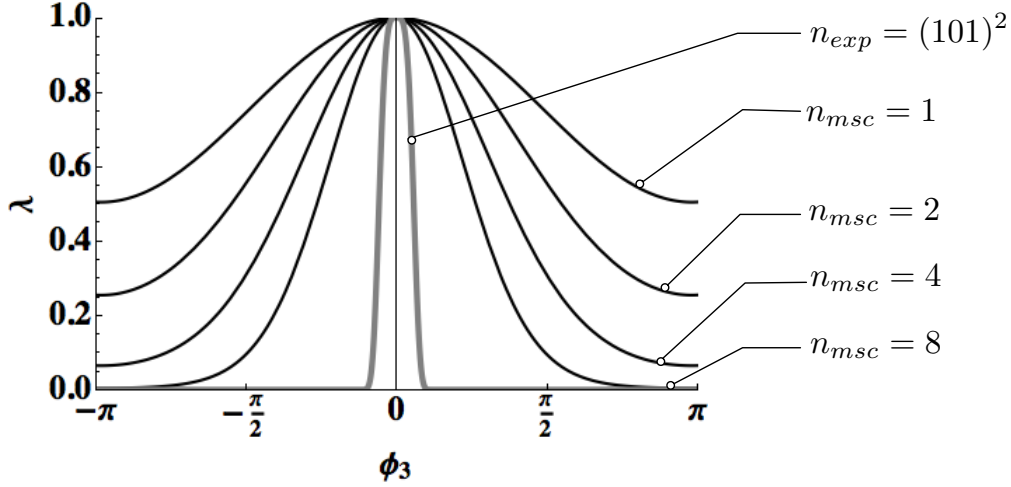


Fig. 9. Amplification factor of the rotated biharmonic operator along the line $D_2 = 0$, for $s_1 = 10$. The results are shown for the TRIADS scheme when a sub-stepping of the operator is used with $n_{msc} = 1, 2, 4$ or 8 sub-steps of the (MSC) scheme, and with $n_{exp} = (1 + s_1^2)^2$ sub-steps of the (EXP) scheme. The case $n_{msc} = 1$ corresponds to $\sigma_1^{(4)} = 1/4$.

$$\begin{cases} q^{n+1,*} = q^n + \Delta t \text{ RHS}^{n+\frac{1}{2}} + \Delta t \mathcal{D}_4(q^n) - \Delta t \partial_3 [\tilde{\kappa} \partial_3 q^n] \\ q^{n+1} = q^{n+1,*} + \Delta t \partial_3 [(K_3 + \tilde{\kappa}) \partial_3 q^{n+1}], \end{cases} \quad (5.1)$$

where K_3 is the vertical eddy-diffusivity given by an appropriate parameterization of the sub-grid scale vertical mixing, and RHS contains all the terms other than diffusive terms. More generally, the explicit part of the stabilizing correction must be applied at the initial time of the tracer time step (i.e. at time n for ROMS/MitGCM, $n - 1/2$ for MOM4p1 (Griffies, 2010, Eqn (8.8)), or $n - 1$ for NEMO/POP), and the implicit part at the final time (i.e. at time $n+1$ for ROMS/MitGCM/NEMO/POP, or $n+1/2$ for MOM4p1). For models using a Leapfrog scheme, like NEMO or POP, Δt must be replaced by $2\Delta t$ (Madec, 2008; Smith et al., 2010).

5.2 Boundary Conditions

We have considered so far a model problem defined on an unbounded domain. For realistic applications, the specification of boundary conditions is required and is based upon two requirements : the mixing operator conserves the tracer content and reduces the global tracer variance. Conservation of the tracer being diffused imposes $F^{(3)} = 0$ at $x_3 = \zeta$ (with ζ the free-surface) and $x_3 = -H$ (with H the depth of the water column). To ensure that the boundary conditions are consistent with the global tracer variance diminishing property it is convenient to use the tracer/density triads formalism defined in Sec. 2.3. Using the notations introduced in Fig. 1, we consider that the grid cell (i, k) is located at the ocean floor. The no-flux boundary condition requires $F_{i,k-\frac{1}{2}}^{(3)} = 0$ which involves that the triads 3,4,11 and 12 cancel. More generally, each time a given triad crosses the boundary the diffusivity $A(n)$ associated with this triad in the discrete functional (2.17) should be zero, and this triad can not participate in the computation of the $F^{(1)}$ interfacial fluxes in the grid-cell next to a boundary. At the bottom (i.e. for $k = 1$) we have

$$-F_{i+\frac{1}{2},j,1}^{(1,\text{triads})} = \frac{\kappa_1}{2} (\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k} \cdot \frac{1}{2} \left(\frac{\delta_1 q_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{\delta_1 \rho_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} \frac{\delta_3 q_{i+1,j,\frac{3}{2}}}{\delta_3 \rho_{i+1,j,\frac{3}{2}}} \right. \\ \left. + \frac{\delta_1 q_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{\delta_1 \rho_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} \frac{\delta_3 q_{i,j,\frac{3}{2}}}{\delta_3 \rho_{i,j,\frac{3}{2}}} \right), \quad (5.2)$$

and

$$-F_{i+\frac{1}{2},j,1}^{(1,\text{sw-triads})} = \frac{\kappa_1}{2} (\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k} \left(\frac{\delta_1 q_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{w_{i+\frac{1}{2},j,1}^+}{(\Delta x_1)_{i+\frac{1}{2},j}} \frac{\delta_3 q_{i+1,j,\frac{3}{2}}}{\delta_3 \rho_{i+1,j,\frac{3}{2}}} - \frac{w_{i+\frac{1}{2},j,1}^-}{(\Delta x_1)_{i+\frac{1}{2},j}} \frac{\delta_3 q_{i,j,\frac{3}{2}}}{\delta_3 \rho_{i,j,\frac{3}{2}}} \right). \quad (5.3)$$

This approach can be seen as a diffusivity tapering because κ_1 is divided by two in the bottom most grid box. An other way to specify a boundary condition is to consider that $F_{i,j,\frac{1}{2}}^{(3)} = 0$ because $(\delta_3 \rho_{i,j,\frac{1}{2}})^{-1} = 0$ at the bottom. In this case we get

$$-F_{i+\frac{1}{2},j,1}^{(1,\text{triads})} = \kappa_1 (\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k} \left[\frac{\delta_1 q_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{\delta_1 \rho_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} \cdot \frac{1}{4} \left\{ \frac{\delta_3 q_{i+1,j,\frac{3}{2}}}{\delta_3 \rho_{i+1,j,\frac{3}{2}}} + \frac{\delta_3 q_{i,j,\frac{3}{2}}}{\delta_3 \rho_{i,j,\frac{3}{2}}} \right\} \right] \quad (5.4)$$

and

$$-F_{i+\frac{1}{2},j,1}^{(1,\text{sw-triads})} = \kappa_1 (\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k} \left(\frac{\delta_1 q_{i+\frac{1}{2},j,1}}{(\Delta x_1)_{i+\frac{1}{2},j}} - \frac{1}{2} \left\{ \frac{w_{i+\frac{1}{2},j,1}^+}{(\Delta x_1)_{i+\frac{1}{2},j}} \frac{\delta_3 q_{i+1,j,\frac{3}{2}}}{\delta_3 \rho_{i+1,j,\frac{3}{2}}} - \frac{w_{i+\frac{1}{2},j,1}^-}{(\Delta x_1)_{i+\frac{1}{2},j}} \frac{\delta_3 q_{i,j,\frac{3}{2}}}{\delta_3 \rho_{i,j,\frac{3}{2}}} \right\} \right). \quad (5.5)$$

This second approach is recommended only if a specific procedure is used to ensure that $(\partial_3 \rho)^{-1}$ goes smoothly to zero as a boundary is approached. The same remarks apply at the surface. For realistic applications, the direction of diffusion is generally progressively aligned with the horizontal direction when approaching the surface. This ‘‘boundary rotation’’ is generally done to avoid the rotated operator to interact with the parameterization of sub-grid scale vertical mixing and also to stay consistent with the small slope limit in the well mixed surface layer.

In the horizontal, homogeneous Neumann boundary conditions (i.e., $\partial_1 q = \partial_1 \rho = 0$) are imposed through masking at the coast. Again, because the biharmonic operator corresponds to two successive Laplacian operators, the specification of the boundary conditions does not imply additional difficulties, and the second boundary condition at the coast required for this operator is simply $\partial_1^3 q = \partial_1^3 \rho = 0$. It is straightforward to extend the results of this paragraph to the x_2 -direction. The boundary condition for the stabilizing step in the (MSC) scheme is an homogeneous Neumann condition to ensure conservation of the tracer content.

A major difficulty when rotating the diffusion in the isoneutral direction is to maintain the numerical integrity of the scheme in poorly stratified regions, including boundary layers and convective regions. Here, the notion of numerical integrity encompasses several aspects : stay consistent with the small slope approximation, smoothly satisfy the boundary condition when conditions (5.4) and (5.5) are used, avoid any infinite slopes and associated infinite fluxes which would introduce numerical instabilities. In the next two subsections, we discuss more specifically those delicacies.

5.3 Boundary Rotation

A specific procedure is required to safely consider that the neutral slope varies smoothly in the boundary layer and cancels at the surface. Several ways to handle the transition from isoneutral to horizontal mixing have been proposed in the literature (e.g. Treguier et al., 1997; Fox-Kemper et al., 2008; Ferrari et al., 2008, 2010, see also Griffies (2004) (Chap. 15) for a review on this topic through 2003). Those methodologies are generally specifically designed in the framework of the Gent and McWilliams (1990) parameterization to avoid any spurious recirculation in the surface layer when an eddy-induced velocity is used. Indeed, the eddy-induced velocity, parameterizing the effect of mesoscale eddies, is usually defined with respect to the vertical derivative of the neutral slopes, the type of boundary rotation adopted in the surface layer therefore interacts with the parameterization. For this reason, a linear clipping, ensuring a constant eddy-induced velocity in the surface layer, is usually applied to the neutral slopes between the base of the surface boundary layer and the surface (Gnanadesikan et al., 2007; Madec, 2008). However, recent advances in the understanding of the role of eddy fluxes in the mixed layer have led to a redesign of the function responsible for the boundary rotation (Ferrari et al., 2008; Fox-Kemper et al., 2008; Colas et al., 2012). A somehow different way to proceed is presented in Ferrari et al. (2010) where the eddy-induced velocity is computed from a vertical mode decomposition under the assumption that the parameterized eddy-transport is dominated by low baroclinic modes. In this case, the computation of the eddy-induced velocity does not require any slope limiting or boundary rotation procedure, and those delicacies are transferred to the isoneutral diffusion operator only.

In a more general view, when considering the isoneutral diffusion independently of the eddy-induced velocity, the important term to keep under control is $(\partial_3\rho)^{-1}$ and no longer the neutral slopes.

5.4 Diffusion Slope Limit

We provide here a simple example of a procedure to stay consistent with the small slope limit and to satisfy the boundary conditions, although many other procedures might be perfectly valid. We first define the quantity

$$(\partial_1\rho)_{i,j,k+\frac{1}{2}}^{\max} = \max \left\{ \frac{|\delta_1\rho_{i-\frac{1}{2},j,k}|}{(\Delta x_1)_{i-\frac{1}{2},j}}, \frac{|\delta_1\rho_{i+\frac{1}{2},j,k}|}{(\Delta x_1)_{i+\frac{1}{2},j}}, \frac{|\delta_1\rho_{i-\frac{1}{2},j,k+1}|}{(\Delta x_1)_{i-\frac{1}{2},j}}, \frac{|\delta_1\rho_{i+\frac{1}{2},j,k+1}|}{(\Delta x_1)_{i+\frac{1}{2},j}} \right\}$$

and the maximum slope α_{\max} allowed. The value of $(\delta_3\rho)^{-1}$, defined at a cell vertical interface, that can be used in practice to compute the neutral slopes is given by

$$\left(\delta_3\rho_{i,j,k+\frac{1}{2}}\right)^{-1} = \max \left\{ -\frac{F(x_3) \cdot \alpha_{\max}}{(\Delta x_3)_{i,j,k+\frac{1}{2}} (\partial_1\rho)_{i,j,k+\frac{1}{2}}^{\max}}, \min \left\{ -\varepsilon, \left(\delta_3\rho_{i,j,k+\frac{1}{2}}\right)^{-1} \right\} \right\} \quad (5.6)$$

In the case of an unstable stratification, the enforcement of a minimum stratification ε maintains the diffusion in the horizontal direction. In the case of a stable stratification, but with large slopes, the slope is bounded by a maximum value α_{\max} . The procedure (5.6) ensures that the limiting of $(\delta_3\rho)^{-1}$ is "felt" the same way when the slopes are computed at u and w points. The analytical function F in (5.6) enables the implementation of a boundary rotation. The value of this function is typically 1 in the oceanic interior and decreases to 0 when approaching the boundaries.

In Sec. 7, we provide a critical discussion on the choice of the actual slope α_m (as in (5.6)) rather than

on the grid slope ratio s_m to maintain the numerical integrity of the isoneutral mixing operators.

5.5 Isoneutral Directions

The first implementations of the rotated operators were using the *in-situ* density to compute the $\delta_m \rho$ terms in the rotation tensor (Cox, 1987). As mentioned earlier in Sec. 2.2.1, an extra horizontal diffusion was generally required to control an instability in the corresponding scheme. The source of this instability was found later in Griffies et al. (1998) (see also Griffies (2004), Chap. 14). They showed that using the *in-situ* density to compute the $\delta_m \rho$ terms leads to a non-linear instability due to an imbalance between the active tracer isoneutral diffusive fluxes. The balance between those fluxes is achieved when the diffusive flux for potential temperature and salinity combine to give zero locally referenced potential density flux. In Griffies et al. (1998), a procedure to compute the neutral slopes is presented to ensure this constraint. For completeness, we briefly recall the basis of this procedure and express it with our notations. We first note that

$$\delta_m \rho(\Theta, S, x_3) = \rho^{(\Theta)} \delta_m \Theta + \rho^{(S)} \delta_m S, \quad (5.7)$$

where

$$\rho^{(\Theta)} = \left[\frac{\partial \rho}{\partial \Theta} \right]_{S, x_3 = \text{Const}}, \quad \rho^{(S)} = \left[\frac{\partial \rho}{\partial S} \right]_{\Theta, x_3 = \text{Const}}$$

with Θ the potential temperature and S the salinity. At a discrete level, the natural placement for the $\rho^{(\Theta)}$ and $\rho^{(S)}$ terms is the same as Θ and S , and each triad has a unique associated value of $\rho^{(\Theta)}$ and $\rho^{(S)}$ (corresponding to the central ρ -point of a triad, Fig. 1). For the potential temperature Θ , the flux (2.13) in the x_1 -direction should thus be rewritten as

$$\begin{aligned} -F_{i+\frac{1}{2},j,k}^{(1,\text{triads})} = \kappa_1 \frac{(\Delta x_2)_{i+\frac{1}{2},j} (\Delta x_3)_{i+\frac{1}{2},j,k}}{(\Delta x_1)_{i+\frac{1}{2},j}} & \left[\delta_1 \Theta_{i+\frac{1}{2},j,k} - \frac{1}{4} \left\{ \delta_3 \Theta_{i,j,k-\frac{1}{2}} \frac{\rho_{i,j,k}^{(\Theta)} \delta_1 \Theta_{i+\frac{1}{2},j,k} + \rho_{i,j,k}^{(S)} \delta_1 S_{i+\frac{1}{2},j,k}}{\rho_{i,j,k}^{(\Theta)} \delta_3 \Theta_{i,j,k-\frac{1}{2}} + \rho_{i,j,k}^{(S)} \delta_3 S_{i,j,k-\frac{1}{2}}} \right. \right. \\ & + \delta_3 \Theta_{i+1,j,k-\frac{1}{2}} \frac{\rho_{i+1,j,k}^{(\Theta)} \delta_1 \Theta_{i+\frac{1}{2},j,k} + \rho_{i+1,j,k}^{(S)} \delta_1 S_{i+\frac{1}{2},j,k}}{\rho_{i+1,j,k}^{(\Theta)} \delta_3 \Theta_{i+1,j,k-\frac{1}{2}} + \rho_{i+1,j,k}^{(S)} \delta_3 S_{i+1,j,k-\frac{1}{2}}} \\ & + \delta_3 \Theta_{i+1,j,k+\frac{1}{2}} \frac{\rho_{i+1,j,k}^{(\Theta)} \delta_1 \Theta_{i+\frac{1}{2},j,k} + \rho_{i+1,j,k}^{(S)} \delta_1 S_{i+\frac{1}{2},j,k}}{\rho_{i+1,j,k}^{(\Theta)} \delta_3 \Theta_{i+1,j,k+\frac{1}{2}} + \rho_{i+1,j,k}^{(S)} \delta_3 S_{i+1,j,k+\frac{1}{2}}} \\ & \left. \left. + \delta_3 \Theta_{i,j,k+\frac{1}{2}} \frac{\rho_{i,j,k}^{(\Theta)} \delta_1 \Theta_{i+\frac{1}{2},j,k} + \rho_{i,j,k}^{(S)} \delta_1 S_{i+\frac{1}{2},j,k}}{\rho_{i,j,k}^{(\Theta)} \delta_3 \Theta_{i,j,k+\frac{1}{2}} + \rho_{i,j,k}^{(S)} \delta_3 S_{i,j,k+\frac{1}{2}}} \right\} \right], \quad (5.8) \end{aligned}$$

same rules apply to the salinity flux and to the vertical flux $F^{(3)}$. As a consequence of this methodology, the isoneutral diffusive flux of the locally referenced potential density discretely vanishes triad-by-triad, as long as the (EXP) scheme is used for the time integration. As discussed earlier, the splitting errors associated with the (IMP) time discretization of the isoneutral Laplacian operator reintroduces an imbalance in the active tracer isoneutral diffusive fluxes. The use of the (MSC) scheme is a way to minimize this imbalance because it introduces only the minimum amount of implicit diffusion necessary to stabilize the temporal integration.

In Lemarié et al. (2012), the neutral slopes are computed using the "elementary adiabatic differences" as defined in Shchepetkin and McWilliams (2011) (Eq. 4.8). In this case, the exact balance between the active tracer isoneutral diffusive fluxes was not satisfied, however the absence of such a balance did not

turn out to be problematic for a realistic multi-decadal basin-scale Pacific simulation, although this issue may be more acute elsewhere and for longer timescales.

5.6 Diagnosing the Stability Condition of the Rotated Biharmonic Operator in a Basin-scale Model Simulation

Based on the prescriptions to handle poorly stratified regions, introduced in (5.6), and on the ROMS Pacific configuration described in Lemarié et al. (2012), we show in Fig. 10 an offline diagnostics of the maximum value of the quantity $(\sigma_1^{(4)}(1 + s_1^2) + \sigma_2^{(4)}(1 + s_2^2))^2 / (1/8)$ for each water column. This quantity, derived from the stability condition (4.4) for the rotated biharmonic operator advanced with the (EXP) scheme, provides an estimate of the stiffness of the stability condition for the isoneutral biharmonic operator. To generate the figure, we considered $\Delta t = 4200$ s (i.e. the ROMS baroclinic time-step used in Lemarié et al. (2012) to integrate the Pacific configuration in time), $B_1 = |u|\Delta x_1^3/12$, $B_2 = |v|\Delta x_2^3/12$, $\alpha_{\max} = 1/10$, and $F(x_3) = (x_3/h_{bl})^2 [3 - 2(x_3/h_{bl})]$ for $x_3 \leq h_{bl}$ ($F(x_3) = 1$ otherwise) where h_{bl} is the depth of the surface boundary layer diagnosed by the parameterization of the vertical mixing. For this configuration, the value of Δx_m varies from 50 km at the equator to 35 km at high latitudes. The hyperdiffusivities B_m are those given in Marchesiello et al. (2009) and Lemarié et al. (2012).

As mentioned earlier, it is not unusual to get very stiff stability conditions especially in shallow areas, including coastal regions, and in the Kuroshio extension. The maximum stability constraint is generally found either at the depth of the thermocline (green shaded area in Fig. 10, bottom panel) or at the bottom (blue shaded area in Fig. 10, bottom panel). Fig. 10 is meant to illustrate that even if the time-step of numerical models is generally set by hyperbolic terms, the use of a forward Euler scheme to integrate the rotated biharmonic operator is impractical and would impose a severe restriction on the time-step of the model, as could be expected from our study.

6 Numerical Results

In this section, we present a set of idealized experiments to investigate the behavior of the various space-time discretization schemes we introduced earlier in the paper.

6.1 Experimental Setup

We have implemented a two-dimensional (x_1, x_3) testcase defined on the computational domain $\Omega = [0, 1] \times [0, 1]$ with a time-independent stratification

$$\rho_\xi(x_1, x_3) = -\tanh\left(5 \left\{x_3 - \frac{1}{4} - \xi 8\pi^3 x_1^3 \left[\sin(\pi x_1) - \frac{1}{2} \sin(2\pi x_1)\right]^2\right\}\right). \quad (6.1)$$

The slope associated with this stratification is

$$\alpha_1 = -\frac{\partial_1 \rho_\xi}{\partial_3 \rho_\xi} = \xi \cdot R(x_1), \quad (6.2)$$

where $R(x) = 64\pi^3 x^2 \cos(\pi x/2) \sin(\pi x/2)^5 (2\pi x + 4\pi x \cos(\pi x) + 3 \sin(\pi x))$. In (6.1), ξ is a stiffness parameter used to define the amplitude of the angle between the computational grid and the local isoneutral

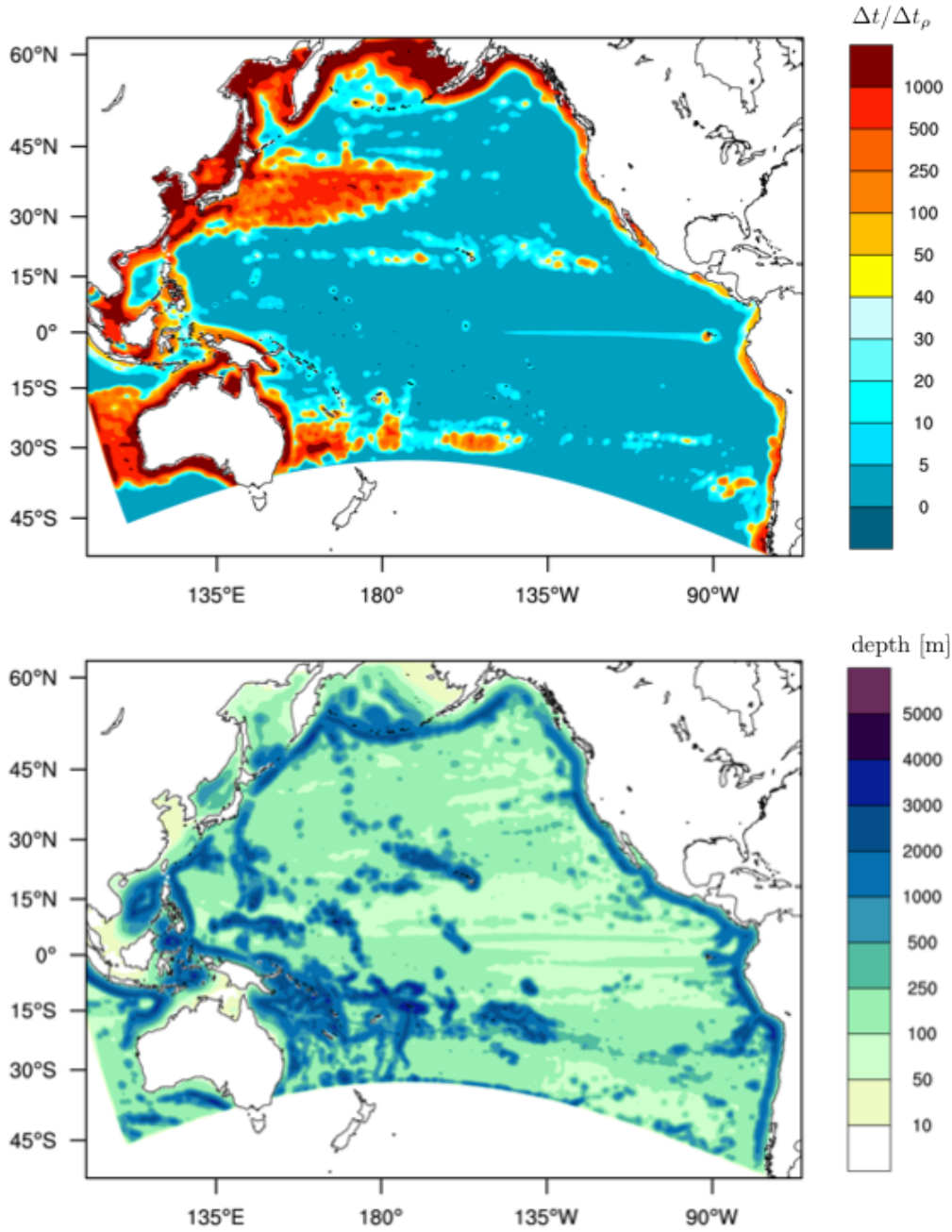


Fig. 10. Annual mean of the maximum value of the ratio $\Delta t / \Delta t_\rho$ for each water column (top), with Δt the time-step of the ROMS model using the (MSC) scheme to advance the rotated biharmonic operator and Δt_ρ the time-step which would be needed to advance the same operator but with the (EXP) scheme. Depth of the maximum (bottom). $\Delta t / \Delta t_\rho$ is first computed using seasonal averages and is then averaged to get the annual mean.

direction. Choosing two values of ξ , we build two set of experiments which are described in Tab. 2 : a first one with a small grid slope ratio between the computational grid and the isoneutral direction (hereafter SMALL) and a second with a large grid slope ratio (hereafter LARGE). The largest value we consider is $s \approx 3$ which, for example, would correspond to $\alpha_1 = 7.5 \times 10^{-4}$, $\Delta x_1 = 100$ km and $\Delta x_3 = 25$ m.

The stratification (6.1) is shown in Fig. 11 along with the initial condition

Resolution	Scheme	Time Step	Maximum slope parameter (s_1)		$\Delta t_0/\Delta t_\rho$			
			SMALL	LARGE	TRIADS		SW-TRIADS	
					\mathcal{D}_2	\mathcal{D}_4	\mathcal{D}_2	\mathcal{D}_4
1024x96	(EXP)	Δt_ρ	0.08	0.35	1	1	1	1
256x48	(MSC)	Δt_0	0.16	0.70	1.5	2.25	1	1
128x48	(MSC)	Δt_0	0.32	1.4	2.96	9	1.96	3.84
64x24	(MSC)	Δt_0	0.32	1.4	2.96	9	1.96	3.84
32x24	(MSC)	Δt_0	0.64	2.8	8.8	77	7.84	61.5

Table 2

Description of the numerical experiments discussed in Sec. 6. Those experiments are carried out using the rotated Laplacian operator (with $\kappa_1 = 5 \text{ m}^2 \text{ s}^{-1}$) and the rotated biharmonic operator (with $B_1 = 5 \times 10^{-1} \text{ m}^4 \text{ s}^{-1}$) with either the SW-TRIADS or the TRIADS spatial discretizations. Δt_ρ is given by the stability limit (3.6) for the Laplacian operator and (4.4) for the biharmonic operator. Δt_0 corresponds to the time step restriction arising from the horizontal component of the rotated operators; i.e. $\Delta t_0 = \Delta x_1^2/2\kappa_1$ for the diffusion, resp. $\Delta t_0 = \Delta x_1^4/8B_1$ for the hyperdiffusion. The ratio $\Delta t_0/\Delta t_\rho$ is given for the LARGE experiments with the rotated Laplacian operator \mathcal{D}_2 and for the rotated biharmonic operator \mathcal{D}_4 .

$$q_0(x_1, x_3) = \frac{1}{4} \left[\cos \left((20 x_3 - 5) \frac{\pi}{3} \right) + 1 \right] \left[\cos \left((20 x_1 - 5) \frac{\pi}{3} \right) + 1 \right] \mathbf{I}_{[1/10, 4/10]}(x_1) \mathbf{I}_{[1/10, 4/10]}(x_3), \quad (6.3)$$

where $\mathbf{I}_\Omega(x)$ is the indicator function⁶. All the simulations described in Tab. 2 are run with the same diffusivity $\kappa_1 = 5 \text{ m}^2 \text{ s}^{-1}$ and hyperdiffusivity $B_1 = 5 \times 10^{-1} \text{ m}^4 \text{ s}^{-1}$.

6.2 Accuracy of the Rotated Operators

The accuracy of different space-time discretizations is checked by computing a high resolution solution on a 1024x96 grid using an explicit Euler scheme in time. This solution is taken as a solution of reference for a set of four coarser grids, 256x48, 128x48, 64x24, and 32x24. With the exception of the reference solution, all the simulations are done using the (MSC) scheme with the stabilizing parameters (3.7), (3.9), (4.5) and (4.7) with $\kappa_2 = 0$ (resp. $B_2 = 0$). To build this hierarchy of grids we have considered that when the horizontal resolution is refined by n^2 the vertical resolution is refined by n (Tab. 2). This rule is meant to be quite consistent with the usual practice in climate models. Due to the excessive computational cost of the explicit-in-time rotated biharmonic operator on the 1024x96 grid, we use the 256x48 grid as a solution of reference in this case.

The stratification (6.1) and the initial condition (6.3) are first defined on the reference 1024x96 grid and then a coarsening procedure is used to provide it to the entire hierarchy of grids. The same coarsening procedure is then applied to the solution of reference to compute the ℓ_2 -norm of the error for each grid of the hierarchy. This procedure ensures that the error is initially zero. The results obtained for the SMALL and the LARGE experiments are shown in Fig. 12. When using the TRIADS scheme, the error is monotonically decreasing when the grid is refined. For large grid slope ratios this discretization leads

⁶ $\mathbf{I}_\Omega(x) = 1$, if $x \in \Omega$; $\mathbf{I}_\Omega(x) = 0$ otherwise.

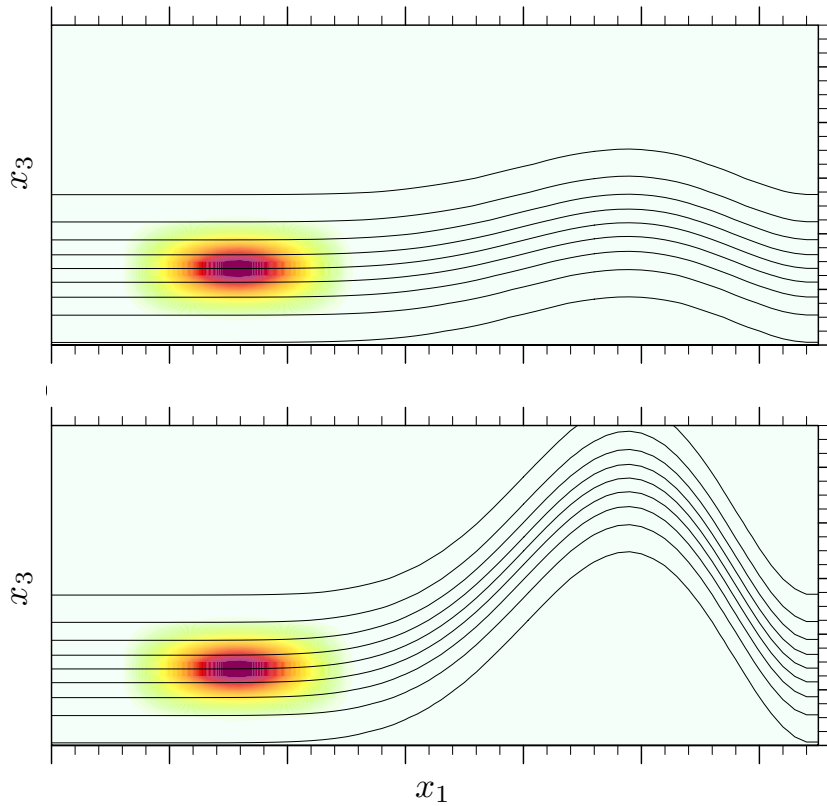


Fig. 11. Two-dimensional (x_1, x_3) initial conditions for different experiments of a passive tracer (shaded) diffused along a density field (contours) with a moderate slope between the computational grid and the isoneutral direction (top), and with a large slope (bottom).

to a significant amount of spurious diapycnal diffusion due to discretization errors (Fig. 13, g) and i). The SW-TRIADS scheme systematically leads to smaller errors whatever the grid resolution and the grid slope ratio (Fig. 12). However, the evolution of the errors with the resolution is relatively uncommon. For example, in the LARGE experiments, the error is smaller on the 128x48 grid than on the 256x48 grid. This is explained by the fact that the errors of discretization associated with the SW-TRIADS scheme do not monotonically increase when the grid slope ratio s_1 increases, as this is the case for the TRIADS scheme. This point will be further discussed in Sec. 6.4 but it can be seen from (2.21) that the discretization reduces to the classical 1,-2,1 stencil in the oblique direction for $s_1 = 1$, thus explaining why the errors are small for the experiments for which the grid slope ratio is close to one (e.g., for the 64x24 and 128x48 grids in the LARGE case). As far as the biharmonic operator is concerned, the results are very consistent with the one obtained with the rotated Laplacian operator (Fig. 14). The SW-TRIADS performs better than the TRIADS when the grid slope ratio exceeds one, which is an important asset when considering grid cells with a small aspect ratio $\Delta x_3/\Delta x_1$ as it can happen for a σ -coordinate models in shallow areas. For very large values of s_1 the direction of diffusion becomes less and less accurate (Fig. 13 j), Fig. 14 h)) and a larger vertical stencil should potentially be required to lessen the discretization errors.

Monotonicity violations are obvious from Fig. 13 and 14 where the blue-shaded areas correspond to negative values whereas the initial condition has positive values only. When the discretization errors are increasing, thus leading to a spurious diapycnal diffusion, the false extrema are significantly moderated. The biharmonic operator generates large negative values (Fig. 14), however those min-max violations are not necessarily associated with the rotation of the tensor because the horizontal biharmonic operator also

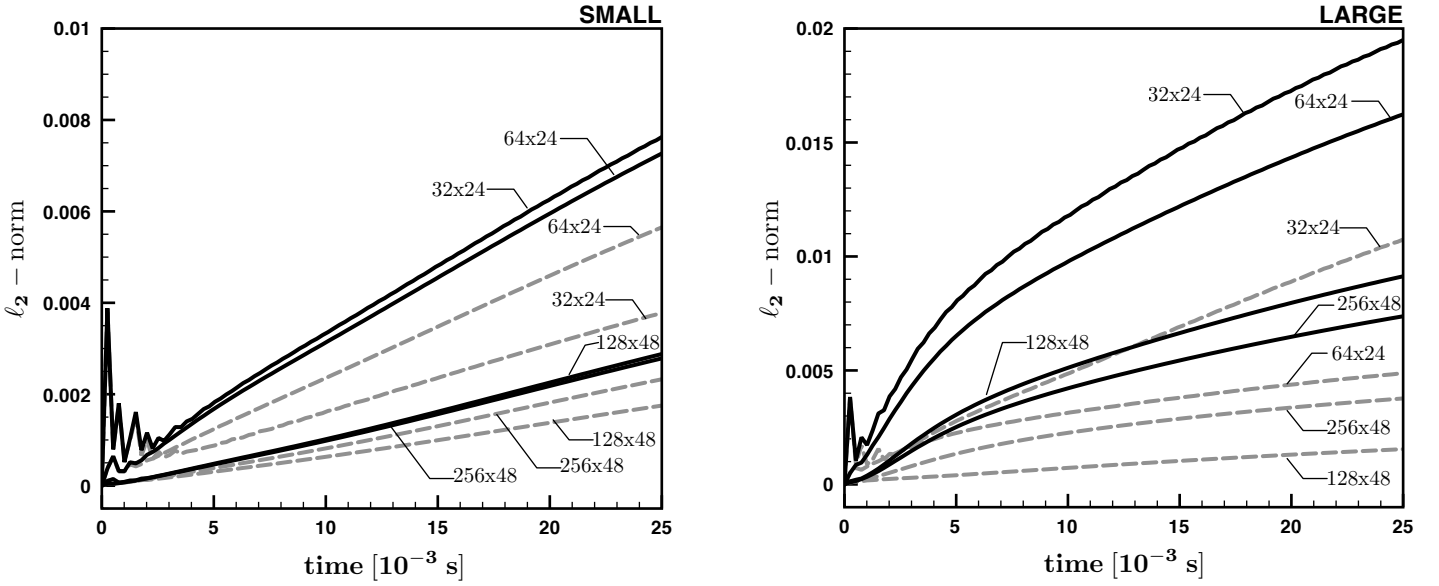


Fig. 12. Time history of the error ℓ_2 -norm for different grid resolutions, in the SMALL experiments (left) and in the LARGE experiments (right). The black lines correspond to the results obtained with the TRIADS scheme, and the dashed gray lines to the results obtained with the SW-TRIADS. Note the different vertical axis between the two panels.

produces overshoots with a similar order of magnitude (not shown). It is worth mentioning that the part of the computational domain with a flat stratification coincides with the location of the initial condition (Fig. 11). Most of the overshoots with the biharmonic operator are generated during the initialization, this is the reason why relatively large monotonicity violations are located in this part of the computational domain.

6.3 Overshootings

Mathieu and Deleersnijder (1998) emphasized the non-monotonic behavior of the discretized rotated diffusion operator. This can be seen simply by looking at the diffusion of a dirac signal along a direction non-aligned with the computational grid (Fig. 15). Negative bands are generated on both sides of the isoneutral direction along which the Dirac signal is propagating. The rotated biharmonic operator presents two bands of negative values separated by positive values. We can also see from those figures that the cancellation of the contribution of the dianeutral points in the discretization stencil of the SW-TRIADS scheme leads to a thin signal elongated in the isoneutral direction while the TRIADS scheme produces a thicker signal in both the $s_1 = 1/2$ and $s_1 = 2$ cases. Note that the case $s_1 = 1/2$ corresponds to a maximum of discretization errors for the SW-TRIADS scheme in the limit that $s_1 \leq 1$; see discussion in Sec. 6.4.

Monotonicity is verified at a position (i,k) only if all the coefficients in the discretization stencil of the operator under consideration are positive for points others than (i,k) . It can be seen from (2.21) that neither the TRIADS, nor the SW-TRIADS scheme verifies this property. The continuous form of the rotated Laplacian operator has however the property not to amplify existing extrema (Mathieu and Deleersnijder, 1998), this property is lost at the discrete level due to the discretization of the cross-derivatives. We quantify the monotonicity violation associated with the SW-TRIADS and TRIADS schemes by computing

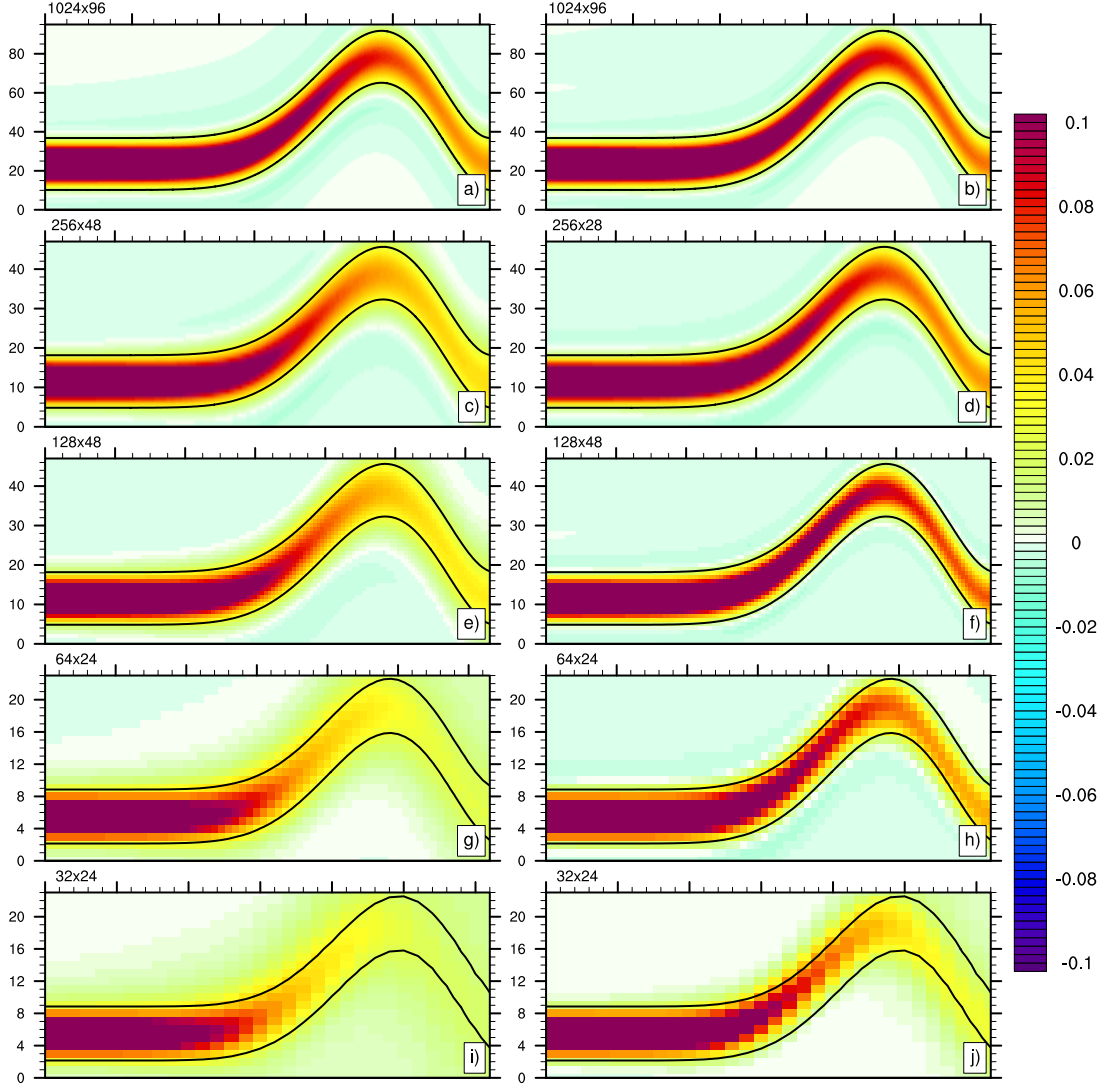


Fig. 13. Snapshots of a passive tracer diffused along isoneutral directions after 25×10^{-3} s of integration in the LARGE case. Results obtained with the TRIADS scheme (resp. SW-TRIADS scheme) are shown in the left panels a), c), e), g) and i) (resp. the right panels b), d), f), h) and j)) for different grid resolutions.

$$\varepsilon_{i,k} = \max(q_{i,k}^{n+1} - q_{i,k}^{\max}, 0) - \min(q_{i,k}^{n+1} - q_{i,k}^{\min}, 0) \quad (6.4)$$

where

$$q_{i,k}^{(\min)} = \min(q_{i\pm c, k\pm c}^n), \quad q_{i,k}^{(\max)} = \max(q_{i\pm c, k\pm c}^n), \quad (6.5)$$

with $c = 1$ for the Laplacian operator and $c = 2$ for the biharmonic operator. The evolution of the maximum value of $\varepsilon_{i,k}$ is shown in Fig. 16 for the diffusion of a Dirac signal. This experiment is done on the 256x48 grid with a diffusivity κ_1 and an hyperdiffusivity B_1 chosen so that it provides the same dissipative time scale at the scale of the grid (i.e., $B_1 = \kappa_1 \Delta x_1^2$).

Fig. 16 shows that the magnitude of the min-max violations is of the same order for the rotated harmonic and biharmonic operators. Even if Fig. 15 suggests that larger negative values are produced by the

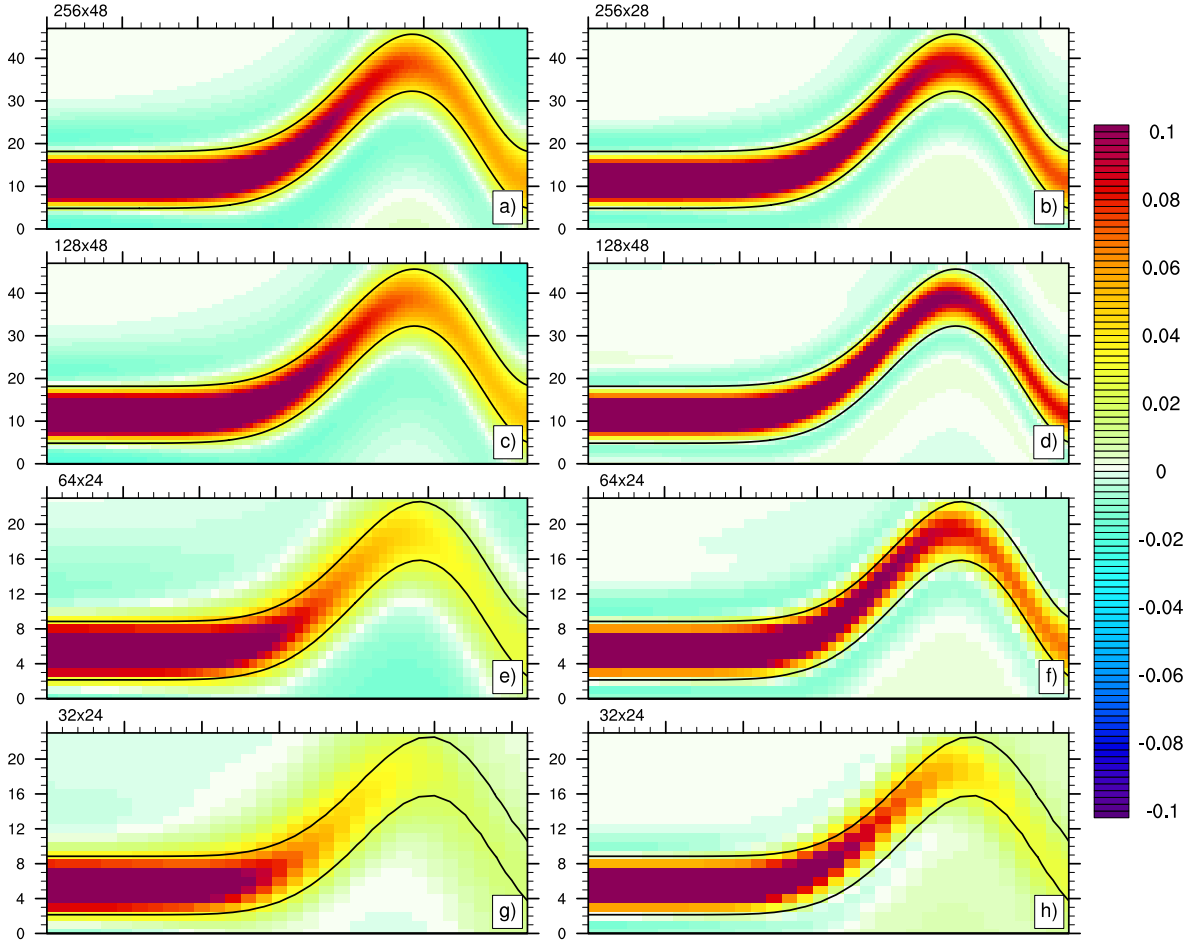


Fig. 14. Snapshots of a passive tracer diffused with a biharmonic operator along isoneutral directions after 25×10^{-3} s of integration in the LARGE case. Results obtained with the TRIADS scheme (resp. SW-TRIADS scheme) are shown in the left panels a), c), e), and g) (resp. the right panels b), d), f) and h)) for different grid resolutions.

biharmonic operator, this does not mean that the min-max violations are larger in this case, as we see from Fig. 16. This is explained by the fact that the amount of diapycnal diffusion, associated with discretization errors, is smaller with the rotated biharmonic operator than with the rotated Laplacian operator and is insufficient to moderate the false extrema.

6.4 SW-TRIADS-COMBI Scheme and Scaling of the Discretization Errors

In this paragraph, we borrow the concept of the so-called COMBI scheme, introduced in Beckers et al. (2000), to assess the evolution of the discretization errors of the SW-TRIADS scheme with respect to s_m . To minimize overshootings, Beckers et al. (2000) propose an inconsistent COMBI scheme which consists in adding an additional horizontal or vertical diffusion to cancel the negative weights in the stencil (2.21) and thus to maintain monotonicity. With the TRIADS discretization this is not straightforward because the negative weights are systematically in the diagonal direction. For the SW-TRIADS scheme, it is easier to apply the concept of the COMBI scheme; in (2.21) we see that, for s_m positive and $s_m < 1$, it is sufficient to add a vertical diffusion with a coefficient proportional to $s_m(1 - s_m)$ and, for $s_m > 1$, it is sufficient to add an horizontal diffusion with a coefficient proportional to $s_m - 1$. In Fig. 17, we show the diffusion

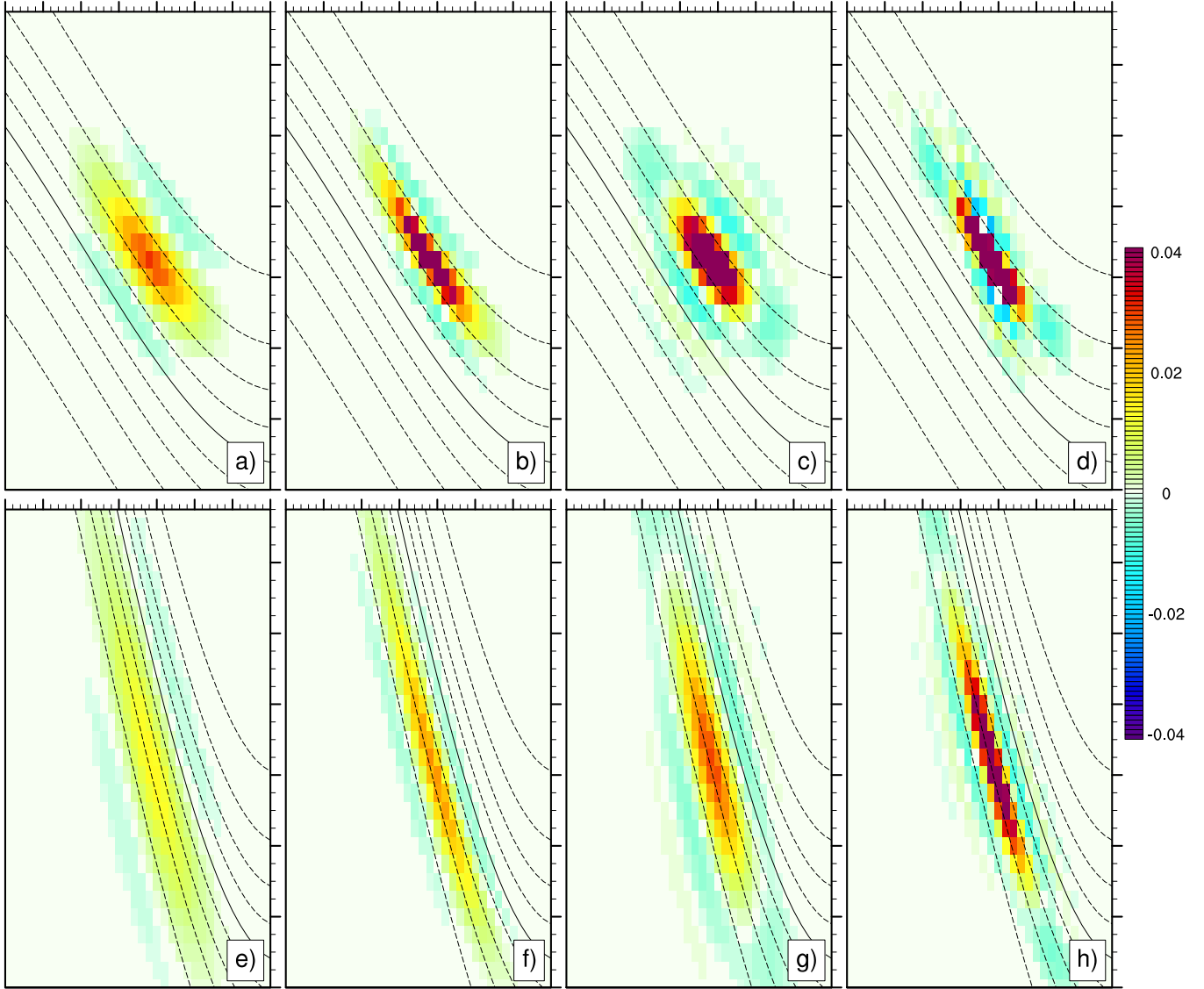


Fig. 15. Diffusion of a Dirac signal along isoneutral directions (contours) after $t = 2.5 \times 10^{-4}$ s : a) and e) with a rotated Laplacian operator using the TRIADS discretization; b) and f) with a rotated Laplacian operator using the SW-TRIADS discretization; c) and g) with a rotated biharmonic operator using the TRIADS discretization; d) and h) with a rotated biharmonic operator using the SW-TRIADS discretization. The tracer field is deliberately not smoothed, cell-averaged values are shown. Note that blue-shaded areas indicate negative values. The grid slope ratio is $s_1 \approx 1/2$ for panels a),b),c) and d), and $s_1 \approx 2$ for e),f),g) and h).

of a Dirac signal using the SW-TRIADS-COMBI scheme (defined as the SW-TRIADS scheme combined with the COMBI approach). The scheme seems to behave relatively well for $s_1 \approx 2$, however the amount of spurious dianeutral diffusion is larger compared with the SW-TRIADS scheme. For $s_1 \approx 1/2$, the orientation of the diffusion is less accurate and the dirac signal is propagated too vertically relative to the isoneutral direction. We can show that the extra diffusivity κ_v used for $s_m < 1$ in the SW-TRIADS-COMBI scheme is

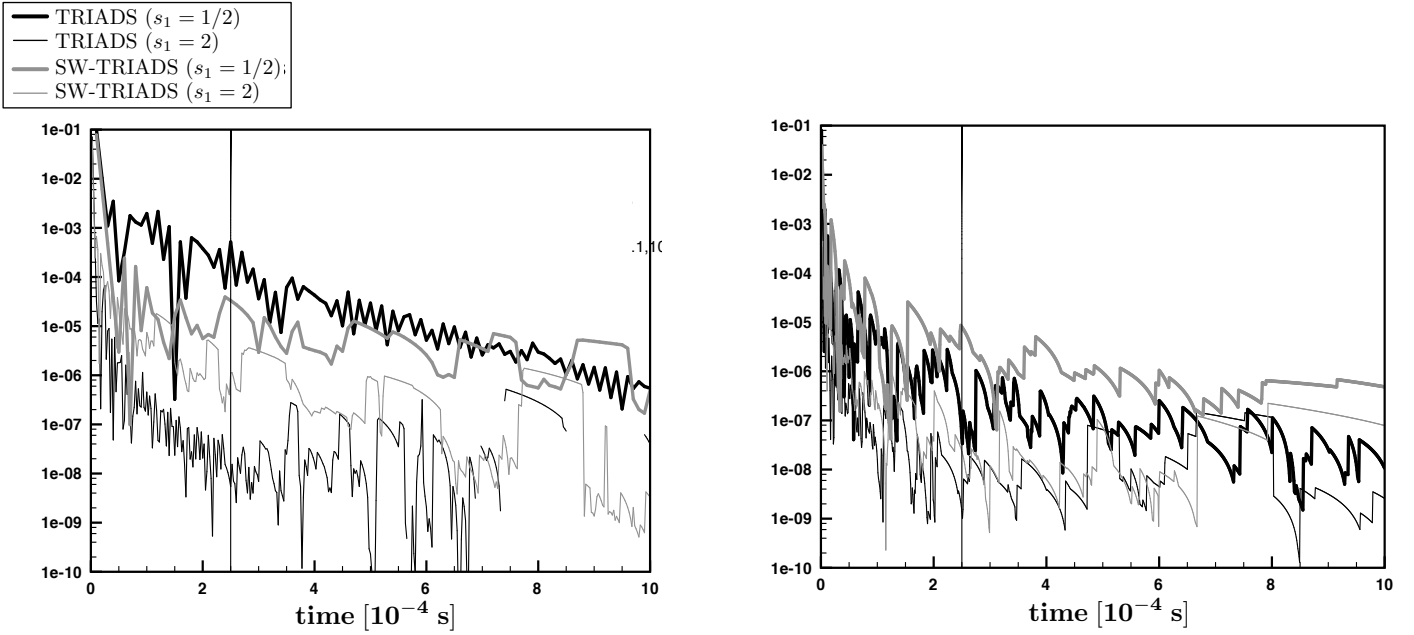


Fig. 16. Magnitude of min-max violations ε for the rotated Laplacian operator (left) and for the rotated biharmonic operator (right) discretized either with the TRIADS scheme (black lines) or the SW-TRIADS scheme (gray lines). The vertical line at $t = 2.5 \times 10^{-4}$ s represents the time corresponding to the plots in Fig. 15.

$$\kappa_v = \kappa_m \left(\frac{\Delta x_3}{\Delta x_m} |\alpha_m| - \alpha_m^2 \right).$$

The expression of κ_v is also useful to quantify the discretization errors associated with the SW-TRIADS scheme. The minimum of the error is obtained for $\alpha_m = 0$, or $\alpha_m = \Delta x_3 / \Delta x_m$ (i.e., $s_m = 1$) because κ_v cancels in those cases, hence the scheme reduces to the classical 1,-2,1 stencil and is monotonic. The maximum of the discretization error is for $\alpha_m = \Delta x_3 / 2 \Delta x_m$ (i.e., $s_m = 1/2$), which leads to $\kappa_v^{\max} = \frac{\kappa_m}{4} \left(\frac{\Delta x_3}{\Delta x_m} \right)^2$. For $s_m > 1$, the discretization error scales with $s_m - 1$ (which corresponds to the background horizontal diffusivity of the SW-TRIADS-COMBI scheme) and thus keeps increasing with increasing values of s_m .

The value of κ_v^{\max} can be compared with the background vertical diffusivity associated with the vertical mixing parameterizations currently in use in ocean models, which is of the order of $10^{-5} \text{m}^2 \text{s}^{-1}$ below the mixed layer. We see that it is not unusual, regarding the typical horizontal and vertical resolutions used by climate models, that the background vertical diffusion in the oceanic interior is sufficient to significantly moderate the overshootings associated with the use of a rotated mixing operator discretized with the SW-TRIADS scheme, as long as the grid slope ratio is smaller than one.

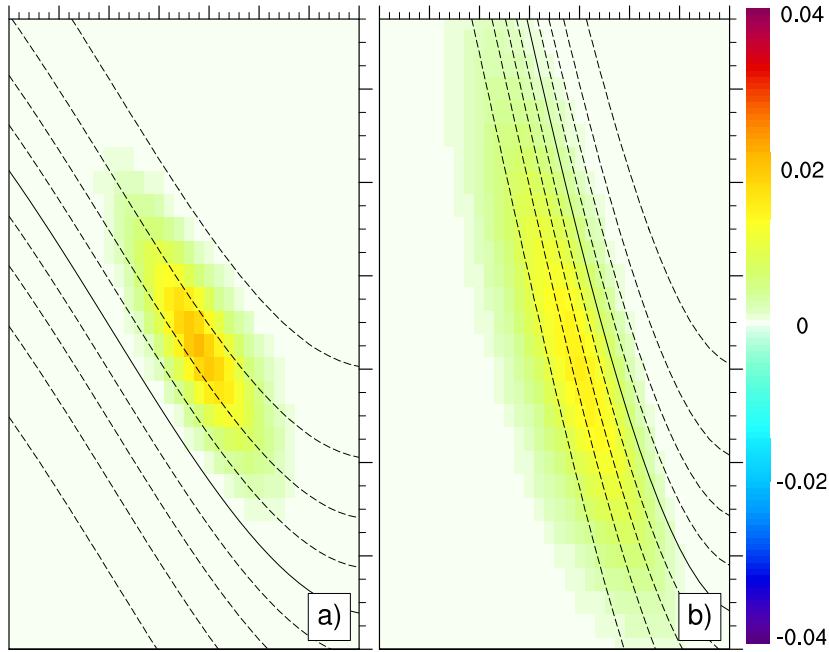


Fig. 17. Diffusion of a Dirac signal (shaded) along isoneutral directions (contours) using the SW-TRIADS-COMBI scheme : a) for $s_1 \approx 1/2$; b) for $s_1 \approx 2$. The results are shown after 2.5×10^{-4} s of integration. Panel a) should be compared with Fig. 15, b) and panel b) with with Fig. 15, f).

7 Conclusion

The use of a mixing operator non-aligned with the computational grid raises difficulties to maintain several properties of the non-rotated operators : numerical stability, strict global tracer variance dissipation, monotonicity (in the case of the Laplacian operator), and accuracy when the slopes steepen to greater than the grid aspect ratio. In this paper, we present a set of conservative space-time discretizations to investigate those delicacies. Since we do not want the numerical models to be penalized by diffusive processes, we consider only linear spatial discretizations and first-order accurate temporal integrations. We do not strive to design higher order schemes because the diffusive terms are used either for numerical or physical closure (or both) and are thus supposed to vanish anyway for sufficiently high resolution.

In the same spirit as the methods commonly used for multi-dimensional parabolic problems with mixed derivatives, we introduce a simple time-stepping scheme and we show its relevance to handle the time-integration of the rotated harmonic and biharmonic operators. This method, which is known as Method of Stabilizing Corrections (MSC), consists of a multi-stage approach : at the first stage, a consistent (explicit) approximation of the rotated operator is evaluated, while all succeeding stages serve only to improve the stability. For both the rotated harmonic and biharmonic operators, the stabilizing step is done using a vertical Laplacian operator whose diffusivity is chosen through linear stability analysis. The proposed scheme is made to enable the rotated operators to be advanced with the same time step as the non-rotated ones. For large grid slope ratios, the scheme can however suffer from a lack of damping of the smallest resolved scales in the vertical. To alleviate this problem a time-splitting of the diffusion could be used with a few (typically of the order of 4) small time steps using the (MSC) scheme within a larger baroclinic time step. This approach would still be significantly more computationally efficient than using an explicit Euler scheme.

As noted previously in Beckers et al. (2000), all the consistent linear spatial discretizations of the

rotated operators produce false dianeutral mixing and min-max violations. This issue is cumbersome because, in general, min-max violations are larger for schemes with small inherent dianeutral mixing, and conversely schemes with a larger dianeutral mixing associated with discretization errors tend to moderate the overshootings. The rotated biharmonic operator does not seem to generate significantly larger monotonicity violations than the rotated Laplacian operator and than the non-rotated biharmonic operator. Furthermore, we show that different spatial discretizations of the rotated operators can lead to a quite different behavior of the solution especially for large grid slope ratios. A compact discretization (referred as to SW-TRIADS) whose stencil adapts to the orientation of the slope provides more accuracy, hence has less spurious dianeutral mixing, for situations where the neutral slope is greater than the vertical to horizontal aspect ratio. However, this discretization does not damp very efficiently small-scale noise, and dispersive errors could thus go uncontrolled with this approach. As shown in Beckers et al. (2000), the SW-TRIADS scheme can easily be made monotonic by adding a vertical background diffusion (as long as the grid slope ratio is less than one) at the expense of a larger dianeutral contribution. However, under some circumstances, the order of magnitude of this background diffusivity is expected to be of the same order as the physically admissible dianeutral diffusivity.

When using the isoneutral mixing operators under realistic conditions, a specific procedure is required to maintain the consistency with the small slope approximation made at a continuous level to derive the rotation tensor, and to properly satisfy the boundary conditions. To do so, we consider a clipping or tapering based on a maximum value of the neutral slope α_m , as done traditionally in numerical climate models. However, our study shows that the stability and the accuracy of the rotated operators are much more dependent on the grid slope ratio s_m rather than on the actual slope α_m . This remark suggests that a clipping on the local value of s_m could be more appropriate. One drawback of this approach would be to allow the violation of the small slope approximation when large values of α_m are associated with small values of s_m . Considering that the small slope limit is valid for $\alpha_m^2 \ll 1$, a slope $\alpha_m = 1/10$ would imply a 1% error which suggests some flexibility in the choice of α_{\max} . It is therefore not clear whether violating the small-slope approximation (which may happen with a limiting acting on s_m) is much more damaging for the stratification than the loss of accuracy for large values of the grid slope ratio (which may happen with a limiting acting on α_m). A definitive answer to this question would require numerical experiments under fully-realistic conditions. This point is left for a future study, but this question certainly deserves more attention.

We show that rotated operators must be used with care if we rely on those operators to efficiently smooth grid-scale noise along the computational grid. Adding a background vertical, as it is commonly done through the vertical mixing parameterization, or horizontal diffusion is a way to erase the deficiencies in terms of smoothing but at the expense of additional numerically-induced spurious dianeutral mixing. The requirements in terms of noise control are tied to the numerical schemes used to handle the tracer advection. When low-order schemes are used, larger dispersive errors are expected and an operator providing an efficient numerical filtering is crucial. In this case, the TRIADS scheme would be recommended. An unexplored possibility could be to formulate the advection in an isoneutral framework hence the numerical noise would be prominently in the isoneutral direction and rotated operators would be more adequate to control this noise.

Besides the idealized configurations studied in the present paper, it is shown in Lemarié et al. (2012) that a rotated biharmonic operator advanced with the (MSC) scheme in time and using the SW-TRIADS discretization performs well for a realistic coarse-resolution basin-scale simulation with the Regional Oceanic Modelling System (ROMS, Shchepetkin and McWilliams, 2005) σ -coordinate model. Using this space-time discretization, we have demonstrated that the use of an isoneutral biharmonic operator sig-

nificantly improves the conservation of intermediate water properties and leads to a tightening of the thermocline compared to an iso- σ biharmonic operator. Because the grid slope ratio with a σ -coordinate is generally larger than with a z -level model we took advantage of the gain of accuracy provided by the SW-TRIADS scheme.

Acknowledgements

F. Lemarié, A. Shchepetkin and J.C. McWilliams acknowledge the support of the Office of Naval Research Grant N00014-08-1-0597. L. Debreu was funded by the ANR through contract ANR-11-MONU-005 (COMODO). The authors would also like to thank S. Griffies and three anonymous reviewers whose comments and suggestions during the review process helped to clarify the manuscript.

References

- Andreev, V.B., 1967. Splitting operator difference schemes for general second order p-dimensional parabolic equations with mixed derivatives. U.S.S.R. Comput. Math. Math. Phys 7, 92 – 104.
- Beckers, J.M., Burchard, H., Campin, J.M., Deleersnijder, E., Mathieu, P.P., 1998. Another reason why simple discretizations of rotated diffusion operators cause problems in ocean models: Comments on “Isoneutral diffusion in a z-coordinate ocean model”. J. Phys. Oceanogr. 28, 1552–1559.
- Beckers, J.M., Burchard, H., Deleersnijder, E., Mathieu, P.P., 2000. Numerical discretization of rotated diffusion operators in ocean models. Mon. Weather Rev. 128, 2711–2733.
- Colas, F., Capet, X., McWilliams, J.C., Li, Z., 2012. Mesoscale eddy buoyancy flux and eddy-induced circulation in eastern-boundary upwelling systems. J. Phys. Oceanogr. Under review.
- Cox, M.D., 1987. Isopycnal diffusion in a z-coordinate ocean model. Ocean Modelling, (unpublished manuscripts) 74, 1–5.
- Craig, I., Sneyd, A., 1988. An alternating-direction implicit scheme for parabolic equations with mixed derivatives. Comput. Math. Applic. 16, 341–350.
- Cushman-Roisin, B., Beckers, J.M., 2011. Introduction to Geophysical Fluid Dynamics: Physical and Numerical Aspects. Academic Press.
- Danabasoglu, G., McWilliams, J.C., 1995. Sensitivity of the global ocean circulation to parameterizations of mesoscale tracer transports. J. Climate 8, 2967–2987.
- Danabasoglu, G., McWilliams, J.C., Gent, P.R., 1994. The role of mesoscale tracer transports in the global ocean circulation. Science 264, 1123–1126.
- Delhez, E.J., Deleersnijder, E., 2007. Overshootings and spurious oscillations caused by biharmonic mixing. Ocean Modell. 17, 183 – 198.
- Douglas, J., 1962. Alternating direction methods for three space variables. Numer. Math. 4, 41–63.
- Douglas, J., Gunn, J.E., 1964. A general formulation of alternating direction methods. Numer. Math. 6, 428–453.
- Ferrari, R., Griffies, S.M., Nurser, A.J.G., Vallis, G.K., 2010. A boundary-value problem for the parameterized mesoscale eddy transport. Ocean Modell. 32, 143–156.
- Ferrari, R., McWilliams, J.C., Canuto, V., Dubovikov, M., 2008. Parameterization of eddy fluxes near oceanic boundaries. J. Climate 21, 2770–2789.
- Fox-Kemper, B., Ferrari, R., Hallberg, R., 2008. Parameterization of mixed layer eddies. part i: Theory and diagnosis. J. Phys. Oceanogr. 38, 1145–1165.

- Gent, P.R., 2011. The Gent-McWilliams parameterization: 20/20 hindsight. *Ocean Modell.* 39, 2 – 9.
- Gent, P.R., McWilliams, J.C., 1990. Isopycnal mixing in ocean circulation models. *J. Phys. Oceanogr.* 20, 150–155.
- Gnanadesikan, A., Griffies, S.M., Samuels, B.L., 2007. Effects in a climate model of slope tapering in neutral physics schemes. *Ocean Modell.* 16, 1–16.
- Goosse, H., Campin, J.M., Deleersnijder, E., Fichefet, T., Mathieu, P.P., Morales Maqueda, M.A., Tartinville, B., 2008. Description of the CLIO model version 3.0, 7 December 2008 <http://www.astr.ucl.ac.be/research/CLIO/cliio30.pdf>.
- Griffies, S.M., 2004. *Fundamentals of Ocean Climate Models*. Princeton University Press, Princeton, NJ, USA.
- Griffies, S.M., 2010. Elements of MOM4p1, in: GFDL Ocean Group Technical Report No. 6. NOAA/GFDL, February 19, 2010. 444 pp.
- Griffies, S.M., Gnanadesikan, A., Pacanowski, R.C., Larichev, V., Dukowicz, J.K., Smith, R.D., 1998. Isoneutral diffusion in a z-coordinate ocean model. *J. Phys. Oceanogr.* 28, 805–830.
- Hallberg, R., Adcroft, A., 2009. Reconciling estimates of the free surface height in lagrangian vertical coordinate ocean models with mode-split time stepping. *Ocean Modell.* 29, 15 – 26.
- Hecht, M.W., 2010. Cautionary tales of persistent accumulation of numerical error: Dispersive centered advection. *Ocean Modell.* 35, 270–276.
- Hofmeister, R., Burchard, H., Beckers, J.M., 2010. Non-uniform adaptive vertical grids for 3d numerical ocean models. *Ocean Modell.* 33, 70 – 86.
- in 't Hout, K., Welfert, B., 2009. Unconditional stability of second-order ADI schemes applied to multi-dimensional diffusion equations with mixed derivative terms. *Appl. Numer. Math.* 59, 677 – 692.
- van der Houwen, P., Verwer, J., 1979. One-step splitting methods for semi-discrete parabolic equations. *Computing* 22, 291–309.
- Hundsdoerfer, W., 2002. Accuracy and stability of splitting with stabilizing corrections. *Appl. Numer. Math.* 42, 213 – 233.
- Leclair, M., Madec, G., 2011. \tilde{z} -coordinate, an arbitrary lagrangian-eulerian coordinate separating high and low frequency motions. *Ocean Modell.* 37, 139 – 152.
- Lemarié, F., Kurian, J., Shchepetkin, A.F., Molemaker, M.J., Colas, F., McWilliams, J.C., 2012. Are There Inescapable Issues Prohibiting the use of Terrain-Following Coordinates in Climate Models ? *Ocean Modell.* 42, 57–79.
- Lengaigne, M., Madec, G., Menkes, C., 2003. Impact of isopycnal mixing on the tropical ocean circulation. *J. Geophys. Res.* 108.
- Madec, G., 2008. NEMO ocean engine, in: Note du Pole de modélisation No. 27. Institut Pierre-Simon Laplace (IPSL), France.
- Manfredi, G., Ottaviani, M., 1999. Finite-difference schemes for the diffusion equation, in: Leach, P., Bouquet, S., Rouet, J.L., Fijalkow, E. (Eds.), *Dynamical Systems, Plasmas and Gravitation*. Springer Berlin / Heidelberg. volume 518 of *Lecture Notes in Physics*, pp. 82–92.
- Marchesiello, P., Debreu, L., Couvelard, X., 2009. Spurious diapycnal mixing in terrain-following coordinate models: The problem and a solution. *Ocean Modell.* 26, 159–169.
- Mathieu, P.P., Deleersnijder, E., 1998. What is wrong with isopycnal diffusion in world ocean models? *Appl. Math. Modelling* 22, 367 – 378.
- Mathieu, P.P., Deleersnijder, E., Beckers, J.M., 1999. Accuracy and stability of the discretised isopycnal-mixing equation. *Appl. Math. Lett.* 12, 81–88.
- McDougall, T.J., 1987. Neutral surfaces. *J. Phys. Oceanogr.* 17, 1950–1964.

- Redi, M.H., 1982. Oceanic isopycnal mixing by coordinate rotation. *J. Phys. Oceanogr.* 12, 1154–1158.
- Shchepetkin, A., McWilliams, J.C., 2005. The Regional Oceanic Modeling System: A split-explicit, free-surface, topography-following-coordinate ocean model. *Ocean Modell.* 9, 347–404.
- Shchepetkin, A.F., McWilliams, J.C., 2011. Accurate Boussinesq oceanic modeling with a practical, “Stiffened” Equation of State. *Ocean Modell.* 38, 41–70.
- Smith, R.D., Gent, P.R., 2004. Anisotropic Gent-McWilliams Parameterization for Ocean Models. *J. Phys. Oceanogr.* 34, 2541–2564.
- Smith, R.D., Jones, P., Briegleb, B., Bryan, F., Danabasoglu, G., Dennis, J., Dukowicz, J., Eden, C., Fox-Kemper, B., Gent, P., Hecht, M., Jayne, S., Jochum, M., Large, W., Lindsay, K., Maltrud, M., Norton, N., Peacock, S., Vertenstein, M., Yeager, S., 2010. The Parallel Ocean Program (POP) Reference Manual, 23 March 2010, LAUR-10-01853.
- Solomon, H., 1971. On the representation of isentropic mixing in ocean circulation models. *J. Phys. Oceanogr.* 1, 233–234.
- Treguier, A.M., Held, I.M., Larichev, V.D., 1997. On the parameterization of quasi-geostrophic eddies in primitive equation ocean models. *J. Phys. Oceanogr.* , 567–580.

A Upper Bound on $\varphi_m^\theta = \varphi_m - 2\theta y_{m3}^2$

The aim of this appendix is to derive an upper bound for the function φ_m^θ introduced in Sec. 3. We keep the notations previously used in the paper, and for convenience we introduce the parameters $X_m = \tan(\phi_m/2)$, $X_m \in \mathbb{R}$ ($m = 1, 3$); so that the expression of φ_m^θ transforms to

$$\varphi_m^\theta = 2\sigma_m \left(\frac{[(X_m + s_m X_3)^2 + (1 + s_m^2)X_m^2 X_3^2]}{(1 + X_m^2)(1 + X_3^2)} - 2\theta s_m^2 \frac{X_3^2}{1 + X_3^2} \right). \quad (\text{A.1})$$

In the remaining of this section we successively study the behavior of φ_m^θ with respect to X_m and then to X_3 . We first get

$$\text{Sign} \left(\frac{\partial \varphi_m^\theta}{\partial X_m} \right) = \text{Sign} \left(s_m X_3 (1 - X_m^2) + X_m (1 + X_3^2) \right), \quad (\text{A.2})$$

where the roots of the second order polynomial in X_m are

$$X_m^{(\pm)} = \frac{1 + X_3^2 \pm \sqrt{(1 + X_3^2)^2 + (2s_m X_3)^2}}{2s_m X_3}. \quad (\text{A.3})$$

$\frac{\partial \varphi_m^\theta}{\partial X_m}$ is positive for $X_m^{(-)} \leq X_m \leq X_m^{(+)}$ which is sufficient to show that φ_m^θ reaches its maximum either for $X_m = X_m^{(+)}$ or $X_m \rightarrow -\infty$. In (A.1), we see that the asymptotic limit of φ_m^θ for X_m going to infinity either by positive or negative values is the same; the maximum of φ_m^θ is therefore obtained for $X_m = X_m^{(+)}$. Substituting X_m by $X_m^{(+)}$ in (A.1) leads to

$$\varphi_m^\theta \leq \varphi_m^\theta(X_m^{(+)}) = \bar{\varphi}_m^\theta = 2\sigma_m \left(1 + 2s_m^2(1 - 2\theta) \frac{X_3^2}{1 + X_3^2} + \sqrt{1 + 4s_m^2 \frac{X_3^2}{1 + X_3^2} \left(1 - \frac{X_3^2}{1 + X_3^2} \right)} \right). \quad (\text{A.4})$$

After some algebra, we find

$$\text{Sign} \left(\frac{\partial \bar{\varphi}_m^\theta}{\partial X_3} \right) = \text{Sign} \left((1 - 2\theta) \sqrt{1 + 4s_m^2 \tilde{X}_3 (1 - \tilde{X}_3)} + (1 - 2\tilde{X}_3) \right), \quad \tilde{X}_3 = \frac{X_3^2}{1 + X_3^2} \in [0, 1]. \quad (\text{A.5})$$

This derivative cancels for

$$\tilde{X}_3 = \tilde{X}_3^{(\pm)} = \frac{1}{2} \left[1 \pm \sqrt{\frac{(1 - 2\theta)^2 (1 + s_m^2)}{1 + (1 - 2\theta)^2 s_m^2}} \right]. \quad (\text{A.6})$$

The maximum of $\bar{\varphi}_m^\theta$ is reached for $\tilde{X}_3 = \tilde{X}_3^{(+)}$ if $\theta \leq 1/2$ and $\tilde{X}_3 = \tilde{X}_3^{(-)}$ otherwise. In both cases the maximum is the same and provides the following upper bound for φ_m^θ

$$\varphi_m^\theta \leq 2\sigma_m \cdot \mathcal{M}(\theta, s_m), \quad (\text{A.7})$$

where

$$\mathcal{M}(\theta, s_m) = 1 + s_m^2(1 - 2\theta) + \sqrt{(1 + s_m^2)(1 + s_m^2(1 - 2\theta)^2)}. \quad (\text{A.8})$$

B Stability Analysis With the SW-TRIADS Scheme

In this section, we extend the results found in Sec. 3 and 4 for the TRIADS scheme to the SW-TRIADS discretization. The function φ_m^θ (with $m = 1, 2$) defined in (3.22) for $\beta_m = 0$ (i.e., the TRIADS scheme) becomes

$$\varphi_m^{\theta, \text{SW}} = 2\sigma_m \left(\frac{[(X_m + s_m X_3)^2 + (s_m - 1)^2 X_m^2 X_3^2]}{(1 + X_m^2)(1 + X_3^2)} - 2\theta s_m^2 \frac{X_3^2}{1 + X_3^2} \right) \quad (\text{B.1})$$

for $\beta_m = 1$ (i.e., the SW-TRIADS scheme with $s_m \geq 0$). In this appendix, we consider only the case $s_m \geq 0$, it is straightforward to extend by symmetry the results to the case $s_m \leq 0$. As we did in App. A for φ_m^θ , we study here the behavior of $\varphi_m^{\theta, \text{SW}}$ with respect to X_m and X_3 . After some algebra, we get

$$\text{Sign} \left(\frac{\partial \varphi_m^{\theta, \text{SW}}}{\partial X_m} \right) = \text{Sign} \left(s_m X_3 (1 - X_m^2) + X_m (1 + X_3^2 (1 - 2s_m)) \right), \quad (\text{B.2})$$

where the roots of the second order polynomial in X_m are

$$X_m^{(\pm)} = \frac{1 + X_3^2 (1 - 2s_m) \pm \sqrt{(1 + X_3^2 (1 - 2s_m))^2 + (2s_m X_3)^2}}{2s_m X_3}. \quad (\text{B.3})$$

We can easily show that the maximum of $\varphi_m^{\theta, \text{SW}}$ is for $X_m = X_m^{(+)}$ which leads to

$$\varphi_m^{\theta, \text{SW}} \leq \varphi_m^{\theta, \text{SW}}(X_m^{(+)}) = \overline{\varphi}_m^{\theta, \text{SW}} = 2\sigma_m \left(1 + 2(s_m(s_m - 1) - 2\theta s_m^2) \frac{X_3^2}{1 + X_3^2} + \sqrt{1 + 4s_m(s_m - 1) \frac{X_3^2}{1 + X_3^2}} \right). \quad (\text{B.4})$$

The sign of the derivative of this function with respect to X_3 reads

$$\text{Sign} \left(\frac{\partial \overline{\varphi}_m^{\theta, \text{SW}}}{\partial X_3} \right) = \text{Sign} \left((s_m - 1) \left(1 + \sqrt{1 + 4s_m(s_m - 1) \widetilde{X}_3} \right) - 2s_m \theta \sqrt{1 + 4s_m(s_m - 1) \widetilde{X}_3} \right) \quad (\text{B.5})$$

where $\widetilde{X}_3 = X_3^2 / (1 + X_3^2)$.

Rotated Laplacian Operator

We recall that the space-time discretization under consideration is stable for $\varphi_1^{\theta, \text{SW}} + \varphi_2^{\theta, \text{SW}} \leq 2$. In (B.5), we see that for $s_m \leq 1$, $\partial \overline{\varphi}_m^{\theta, \text{SW}} / \partial X_3$ is negative, hence the maximum is for $\widetilde{X}_3 = 0$ which translates into $\varphi_m^{\theta, \text{SW}} \leq 4\sigma_m$ in (B.4). This shows that, in this case, the scheme is stable for $\sigma_1 + \sigma_2 \leq 1/2$ which corresponds to the stability condition of the non-rotated operator. Moreover, for $s_m > 1$, taking $\theta = (s_m - 1)/s_m$ still ensures that $\partial \overline{\varphi}_m^{\theta, \text{SW}} / \partial X_3$ is negative, and thus leads to the same stability constraint. This last remark is sufficient to conclude that the stability condition of the rotated Laplacian operator discretized using the SW-TRIADS scheme in space and the (MSC) scheme in time is equivalent to the stability condition of the non-rotated Laplacian operator for

$$\theta = \theta^{\text{SW}} = \max \left\{ \frac{|s_1| - 1}{|s_1|}, \frac{|s_2| - 1}{|s_2|}, 0 \right\}. \quad (\text{B.6})$$

If $|s_1|$ and $|s_2|$ are both smaller than 1 we get $\theta = 0$, and for $|s_m| \rightarrow \infty$, $\theta \rightarrow 1$ which also shows that the (IMP) scheme is stable as long as the horizontal terms of the tensor are stable.

Rotated Biharmonic Operator

Using (B.6), and following the methodology used in Sec. 4.3 to derive an expression for the stabilizing diffusivity $\tilde{\sigma}^{\text{SW}}$, we find that

$$\tilde{\sigma} = \tilde{\sigma}^{\text{SW}} = 8 (S_1 \sigma_1 + S_2 \sigma_2) ((1 + S_1) \sigma_1 + (1 + S_2) \sigma_2), \quad S_m = \max\{s_m^2 - |s_m|, 0\}. \quad (\text{B.7})$$

Note that $\tilde{\sigma}^{\text{SW}}$ vanishes when $|s_1|$ and $|s_2|$ are smaller than one.