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1 **Estimation of the Free Core Nutation parameters from SG data: sensitivity study and**
2 **comparative analysis using linearized Least-Squares and Bayesian methods**

3

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13

14 **Abstract**

15 The Free Core Nutation (FCN) is investigated with the help of its resonance effect on the tidal
16 amplitudes in Superconducting Gravimeter (SG) records of the GGP network. The FCN
17 resonance parameters are combined in a resonance equation involving the Earth's interior
18 parameters. The sensitivity of the FCN parameters to the diurnal tidal waves demonstrates
19 that the quality factor of the FCN is strongly dependent on the accuracy of the imaginary part
20 estimates of the gravimetric factors close to the resonance. The weak amplitude of Ψ_1 tidal
21 wave on the Earth, which is the closest in frequency to the FCN, in addition to errors in ocean
22 loading correction, explains the poor determination of the quality factor Q from surface
23 gravimetric data. The inversion of tidal gravimetric factors leads to estimates of the period, Q
24 and resonance strength of the FCN. We show that, by inverting $\log(Q)$ instead of Q , the
25 results using the least-squares method optimized using the Levenberg-Marquardt algorithm
26 are in agreement with the Bayesian probabilistic results and agree with the results obtained

1 from VLBI nutation data. Finally, a combined inversion of 7 GGP European SG data is
2 performed giving $T = 428 \pm 3$ days and $7762 < Q < 31989$ (90% C.I.). An experimental
3 estimate of the internal pressure Love number is also proposed.

4
5 **Keywords:** Free Core Nutation, Superconducting Gravimeter, Bayesian analysis

6 7 **1. Introduction**

8 Because of the fluidity of the core, the Earth has a rotational mode, called the free core
9 nutation (FCN) with a period almost diurnal in Earth-fixed coordinates. The FCN parameters
10 (period, damping) strongly depend on the coupling mechanism at the core-mantle boundary
11 (flattening, topography, electro-magnetic coupling...). The FCN can be detected by its effect
12 on the Earth's rotation, using the VLBI network analyses, or by studying its effects on the
13 gravity field. As the tidal potential contains some diurnal components, a resonance occurs in
14 the diurnal frequency band. This resonance effect can be observed in time-varying gravity
15 data continuously recorded on the Earth's surface by Superconducting Gravimeters (SGs) of
16 the Global Geodynamics Project (GGP) network (Crossley et al., 1999). The FCN resonance
17 in gravity data is commonly represented by a damped harmonic oscillator model that we
18 invert in order to determine the FCN frequency, quality factor Q and the transfer function of
19 the mantle (or the resonance strength). The usual approach to solve this non-linear inverse
20 problem is to use a linearized least-squares method optimized based on the Levenberg-
21 Marquardt algorithm (Marquardt, 1963 – Numerical Recipes Fortran Chapter 15.5- see for
22 instance Defraigne et al. 1994 and 1995; Sato et al., 2004; Ducarme et al., 2007). However
23 Florsch and Hinderer (2000) have demonstrated the inadequacy of using such a least-squares
24 method, because the statistical distribution of Q is definitely not Gaussian. They have

1 proposed instead the use of a Bayesian approach to invert the FCN parameters, since the
2 Bayesian method better propagates the information to the parameters.

3 Neuberg et al. (1987) first proposed an inversion of stacked gravity tide measurements
4 in central Europe to retrieve the FCN parameters using the Marquardt optimized linearized
5 least-squares. Then Defraigne et al. (1994) extended the gravity stack to the nutation
6 observations. In those past studies, the obtained Q -value was abnormally small and sometimes
7 even negative. Sato et al. (2004) used $1/Q$ instead of Q as a parameter to be inverted using a
8 modified Marquardt least-squares method since $1/Q$ seems to be Gaussian (Sato et al., 1994).
9 However they obtained a Q -value still smaller than the one retrieved from the VLBI nutation
10 analysis (Table 1). The first application of the Bayesian method was proposed by Florsch and
11 Hinderer (2000), who introduced $\log(Q)$ as a parameter instead of Q , in order to preserve the
12 positivity of Q , and obtained a Q -value greater than 20,000. More recently, Ducarme et al
13 (2009) inverted $\log(Q)$ using both a Bayesian and a least-squares approach but applied on
14 averaged gravimetric factors from European sites. They obtained a value for Q consistent with
15 the VLBI result. The other studies based on the least-squares method are summarized in Table
16 1. Note that Koot et al. (2008) performed an estimation of the FCN resonance parameters
17 from VLBI nutation series using a Bayesian statistical approach in the time domain.

18 Here we propose a comparison of the results given by the linearized least squares
19 method optimized by the Levenberg-Marquardt algorithm with the Bayesian inversion applied
20 on SG gravity records. We show that the Q -value obtained from SG data is now in agreement
21 with the value inverted from VLBI nutation series whatever the method used, least-squares or
22 Bayesian inversion. Besides, we demonstrate that the poor constraint on the Q -value obtained
23 by Florsch and Hinderer (2000) was due to the large uncertainty on the phase of the diurnal
24 tidal waves close to the resonance.

1 In the first part we describe the FCN resonance model. Then, we review the theory of
 2 the Bayesian method and the Levenberg-Marquardt optimization algorithm applied to
 3 linearized least-squares. A qualitative study is then performed to check the sensitivity of the
 4 gravity factors to the FCN parameters. Finally, we invert the FCN resonance parameters using
 5 a combination of 7 European SG time-series.

7 **2. The FCN resonance model**

8 The basic equation used to describe the resonance of the FCN in the tidal gravity is
 9 usually written as (Hinderer et al., 1991a):

$$10 \quad \tilde{\delta}_j = \tilde{\delta}_{ref} + \frac{\tilde{a}}{\sigma_j - \tilde{\sigma}_{nd}}, \quad (1)$$

11 where $\tilde{\delta}_j$ corresponds to the complex gravimetric factor observed for every tidal wave of
 12 frequency σ_j , $\tilde{\sigma}_{nd} = \sigma_{nd}^R + i\sigma_{nd}^I$ is the complex eigenfrequency of the FCN, $\tilde{a} = a^R + ia^I$ refers
 13 to the resonance strength corresponding to the response of the whole Earth to the FCN. The
 14 quantity $\tilde{\delta}_{ref}$ is the value of the gravimetric factor without any resonance process (classical
 15 tidal gravimetric factor); it is also the asymptotic value of $\tilde{\delta}_j$ for frequencies far away from
 16 the resonance frequency. The eigenperiod T of the FCN expressed in sidereal days in the
 17 rotating frame is related to σ_{nd}^R by:

$$18 \quad T = \frac{2\pi}{\sigma_{nd}^R},$$

19 where σ_{nd}^R is expressed in radian per sidereal day. In the inertial reference frame, the period
 20 can be written:

$$21 \quad T' = \frac{1}{\sigma_{nd}^R C - 1}$$

1 where $C = \frac{86164}{15 \times 86400}$ and σ_{nd}^R is given in degrees/solar hour. The quality factor Q ,
 2 expressing the damping due to all the physical processes involved in the resonance, is defined
 3 as $Q = \frac{\sigma_{nd}^R}{2\sigma_{nd}^I}$. The quantities σ_{nd}^R and σ_{nd}^I are positive by definition, therefore they should
 4 follow a log-normal distribution law (Tarantola, 2005; Florsch and Hinderer, 2000) to avoid
 5 possible negative values. It is therefore recommended to include the *a priori* positivity of Q in
 6 the model by changing the variable $Q = 10^x$ and inverting for x , instead of Q .

7 Florsch and Hinderer (2000) performed the inversion by treating $\tilde{\delta}_{ref}^R$ as an unknown
 8 and showed that a correlation exists between the real parts δ_{ref}^R and σ_{nd}^R , and between a^R
 9 and δ_{ref}^R , but the correlation between a^R and T is much stronger. As $\tilde{\delta}_{ref}^R$ has a weak influence
 10 on the values of T and Q , we do not include this parameter in the inversion process. In
 11 previous studies (e.g. Defraigne et al. 1994 – 1995; Ducarme et al., 2007), the observed value
 12 for the tidal wave O_1 was used as the reference gravimetric factor. In our case we will use the
 13 mean value of the theoretical inelastic amplitude factors of the O_1 and OO_1 waves computed
 14 for the Wahr-Dehant model (Wahr, 1974; Dehant, 1987). By doing so, we suppose that the
 15 scale factors of the SG used here are accurate enough, which is usually true (better than 0.3%
 16 accuracy, e.g. Amalvict et al., 2001; Sato et al., 2004). We could also have normalized by the
 17 observed O_1 amplitude as done by Sato et al. (2004) but we suppose that the scaling error is
 18 negligible with respect to the ocean loading uncertainty.

19 The equations relative to the resonance model are written:

$$\left\{ \begin{array}{l} \delta_j^R = \delta_{ref} + \frac{a^R (\sigma_j - \sigma_{nd}^R) - a^I \frac{\sigma_{nd}^R 10^{-x}}{2}}{(\sigma_j - \sigma_{nd}^R)^2 + \left(\frac{\sigma_{nd}^R 10^{-x}}{2} \right)^2} \\ \delta_j^I = \frac{\frac{a^R \sigma_{nd}^R 10^{-x}}{2} + a^I (\sigma_j - \sigma_{nd}^R)}{(\sigma_j - \sigma_{nd}^R)^2 + \left(\frac{\sigma_{nd}^R 10^{-x}}{2} \right)^2} \end{array} \right. \text{with } \delta_{ref} = \frac{\delta_{O1}^R + \delta_{OO1}^R}{2}. \quad (2)$$

Equations (2) will be our formulation for the resonance model.

Hinderer (1986), Hinderer and Legros (1989), Defraigne (1992) and Legros et al. (1993) have written the analytical expression for the resonance strength as a function of the Earth's interior parameters:

$$\tilde{\alpha} = -\tilde{\delta}_2 \frac{A}{A_m} e \left(\sigma \left(1 - \frac{A}{A_f} \frac{\gamma}{e} \right) - \frac{A_m}{A} \tilde{\sigma}'_{nd} \right) - \tilde{\delta}_1 \frac{A}{A_m} e \Omega \left(1 - \frac{\gamma}{e} \right) \quad (3)$$

where γ is a compliance (Mathews et al., 2002) related to the geodetic parameter q_0 (ratio of centrifugal acceleration to gravitational acceleration) and the Love number h^f by $\gamma = \frac{q_0}{2} h^f$. $\tilde{\delta}_1$ is the pressure Love number that represents the gravity response to the pressure potential acting at the CMB (Hinderer et al. 1991a). e is the dynamic ellipticity of the Earth, Ω its rotation rate and A/A_m is the ratio of Earth to mantle moments of inertia.

$\tilde{\sigma}'_{nd} = -\Omega \frac{A}{A_m} (e^f - \beta)$ is the FCN frequency in the inertial reference frame ($\tilde{\sigma}'_{nd} = \tilde{\sigma}_{nd} + \Omega$),

e^f is the dynamical flattening of the CMB and β the compliance defined by $\beta = \frac{q_0}{2} h_1^f \cdot h_1^f$ is

the displacement Love number at the CMB associated with the fluid pressure on the mantle caused by the relative rotation.

1 Legros et al. (1993) also considered the effect of the inner core including the Free
 2 Inner Core Nutation (FICN) resonance effect. We neglect it here as the effect would be too
 3 small to be seen in SG records.

4 In the following, we will estimate the period T , the quality factor Q , as well as the real
 5 and imaginary parts of the resonance strength \tilde{a} using two methods: a probabilistic approach
 6 based on the Bayesian inversion and the more classical linearized least-squares optimized
 7 using the Levenberg-Marquardt algorithm.

8

9 **3. A review of the methods**

10 a. The Bayesian approach

11 The Bayesian inversion consists in propagating the information (or knowledge) provided by
 12 the measurements through an assumed physical model (perfectly or probabilistically known)
 13 to the parameters and to include the *a priori* knowledge of the model parameters. Both the
 14 data and the model parameters are described with probability distributions. The Bayesian
 15 approach preserves the full knowledge provided by the data combined with the physical law
 16 and the *a priori* information on the data and model parameters. Therefore, it is the most
 17 suitable method to perform the inversion of non-linear problems (Tarantola and Valette,
 18 1982a, b). For more details about the Bayesian method, please refer to Florsch and Hinderer
 19 (2000).

20 The Bayesian probability distribution of the parameter vector θ is given by:

$$21 \quad p(\boldsymbol{\theta}) = p(x, \sigma_{nd}^R, a^R, a^I) = k \exp \left\{ -\frac{1}{2} \sum_j \left[\left(\frac{\text{Re}(\delta_j^{th}) - \text{Re}(\delta_j)}{\Delta \text{Re}(\delta_j)} \right)^2 + \left(\frac{\text{Im}(\delta_j^{th}) - \text{Im}(\delta_j)}{\Delta \text{Im}(\delta_j)} \right)^2 \right] \right\}, \quad (4)$$

22 where k is a normalisation factor in order that the integral of this equation is unity, Re
 23 denotes the real part and Im the imaginary part, δ_j holds for the j^{th} measurement value of the

1 gravimetric factor, *th* refers to the theoretical value, and $\Delta\delta$ is the error on δ (standard
2 deviation).

3 The previous formula gives the general probability laws for the parameter vector θ . In
4 order to obtain the law for one or two parameters, we compute the marginal pdfs by
5 integration of the probability function over selected parameters. For instance, the joint pdf
6 integrated with respect to σ_{nd}^R is defined by:

$$7 \quad p_{x,a^R,a^I}(x,a^R,a^I) = \int p(x,\sigma_{nd}^R,a^R,a^I)d\sigma_{nd}^R.$$

8 Notice that when integrating over selected parameters, we limit *a priori* the space parameter
9 on a finite domain. Two further integrations of the pdf lead to the marginal probability law for
10 each of the parameters.

11

12 b. The linearized least-squares approach

13 As for the Bayesian method, the linearized least-squares approach is a subset of the
14 non-linear least squares generalized by Tarantola and Valette (1982b). The only difference is
15 how to treat *a priori* information, and what pdfs are involved. The least-squares method is
16 based on the determination of the best-fit parameters by minimizing the merit function:

$$17 \quad \chi^2 = \sum_j \frac{\left| \tilde{\delta}_j - \left(\delta_{ref} + \frac{\tilde{a}}{\sigma_j - \tilde{\sigma}_{nd}} \right) \right|^2}{(\Delta\tilde{\delta}_j)^2} \quad (5)$$

18 However, when the model is non-linear, the minimization must proceed iteratively given
19 initial values for the parameters. The procedure finishes when χ^2 stops decreasing. The χ^2
20 function is linearized to a quadratic form depending on the Hessian matrix for the inverse-
21 Hessian method or the gradient of χ^2 for the steepest descent method. The Levenberg-
22 Marquardt (also called Marquardt method; Marquardt, 1963) algorithm is in fact a smooth
23 variant between these two methods. The Marquardt method works very well in practice and

1 has become a standard of non-linear least-squares routines. It has been widely used in
 2 previous FCN retrieval studies like in Neuberg et al. (1987), Richter and Zürn (1986), Zürn
 3 and Rydelek (1991), and Defraigne et al. (1994, 1995). We refer to the Numerical Recipes
 4 (Press et al., 1992) Chapter 15.5 for the optimization subroutines.

5 The set of linear equations $\sum_j \alpha_{kj} \Delta \theta_j = \beta_k$ is solved for the increments $\Delta \theta_j$ that, added to the

6 current trial parameters, give the next approximation. The matrix α is called the curvature

7 matrix and is equal to one-half times the Hessian matrix: $\alpha_{kj} = \frac{1}{2} \frac{\partial^2 \chi^2}{\partial a_k \partial a_j}$ and $\beta_k = -\frac{1}{2} \frac{\partial \chi^2}{\partial a_k}$,

8 where a_k (or a_j) corresponds to one parameter. In the Marquardt optimization algorithm, the

9 increments are related to the curvature matrix through a “damping factor” λ as $\Delta \theta_j = \frac{1}{\lambda \alpha_{jj}} \beta_j$.

10 When λ is very large, the diagonal of the curvature matrix is dominant and the solution tends
 11 to the initial parameters (the increments tend toward zero). On the other hand, as λ approaches
 12 zero, the increments become very large and we may have a divergence problem. The damping
 13 factor λ is also known as the Marquardt factor. The damping factor effectively constrains the
 14 range of values that the increments $\Delta \theta_j$ can take.

15 The model used for the inversion is described by equation (2). In order to impose the
 16 positivity of the quality factor Q , we perform the change of variable $x = \log_{10}(Q)$, as for the
 17 Bayesian approach, then we minimize equation (5). We also estimate the damping factor λ ,
 18 which best minimizes χ^2 .

19

20 **4. Sensitivity and correlations between the FCN parameters**

21 In order to check the sensitivity of the diurnal tidal wave amplitudes to the FCN
 22 parameters, we perform some qualitative studies of the resonance transfer function presented
 23 in equation (2) and based on the previously estimated values. We have $\sigma_{nd}^R \cong 7.30910^{-5} \text{ rad} / \text{s}$,

1 $Q \approx 2 \cdot 10^4$ (based on VLBI results, e.g. Mathews et al., 2002), $a^R \approx 3.4 \cdot 10^{-9} \text{ rad/s} \approx 7 \cdot 10^{-4}$
 2 deg/h and $a^I \approx 2 \cdot 10^{-10} \text{ rad/s} \approx 4 \cdot 10^{-5} \text{ deg/h}$ (e.g. Hinderer et al., 1989; Florsch and Hinderer,
 3 2000; see next section).

4 From equation (2) and with these values, for any diurnal waves of angular frequency σ_j , we
 5 can see that the quantity $a^I \frac{\sigma_{nd}^R}{2Q}$ is negligible with respect to

6 $a^R(\sigma_j - \sigma_{nd}^R)$ and $\left(\frac{\sigma_{nd}^R}{2Q}\right)^2 \ll (\sigma_j - \sigma_{nd}^R)^2$, hence the amplitude gravimetric factor can be
 7 approximated by:

$$8 \quad \begin{cases} \delta_j^R \approx \delta_{ref} + \frac{a^R}{(\sigma_j - \sigma_{nd}^R)} \\ \delta_j^I \approx \frac{\frac{a^R \sigma_{nd}^R 10^{-x}}{2} + a^I (\sigma_j - \sigma_{nd}^R)}{(\sigma_j - \sigma_{nd}^R)^2} \end{cases} \quad (6)$$

9 For diurnal waves situated far from the resonance, i.e. for $(\sigma_j - \sigma_{nd}^R) \gg 3 \cdot 10^{-8} \text{ rad/s}$, we
 10 have $\delta_j^I \approx \frac{a^I}{(\sigma_j - \sigma_{nd}^R)}$. Therefore, the FCN quality factor is only constrained by the imaginary

11 part of the gravimetric factor of the diurnal waves close to the resonance (K_I , Ψ_I and Φ_I) and
 12 from equation (6) we can see that the imaginary part of the resonance strength is mainly
 13 constrained by the imaginary part of the gravimetric factors. However, the imaginary parts of
 14 the tidal gravity factors are poorly determined, especially for Ψ_I and Φ_I that have small
 15 amplitudes (particularly a small imaginary part) and are therefore very sensitive to the ocean
 16 loading correction error (Sato et al., 2004). We have tested the influence of the error of the
 17 imaginary part of the gravimetric factors on the resulting pdfs of the Bayesian inversion. For
 18 that, we have considered synthetic gravimetric factors computed from equation (2) with $x = 4$,

1 $T = 430$ days, $a^R = 6 \cdot 10^{-4}$ deg/h and $a^I = -5 \cdot 10^{-5}$ deg/h. Then we assume an error of 0.1% on
 2 the real part of the gravimetric factors for nine diurnal waves (from Q_1 to OO_1 waves). We
 3 increase the relative error on the imaginary part from 1% to 100%, and we also test the case
 4 when the error on $\delta^I(\Psi_l)$ and $\delta^I(\Phi_l)$ becomes larger than their values themselves (relative
 5 uncertainty larger than 100%). The corresponding pdfs are plotted in Fig. 1. When the relative
 6 error on the imaginary part of Φ_l and Ψ_l reaches 100%, the pdf for x begins to be non-
 7 Gaussian and tends to an asymmetric probability law (Fig. 1 (b)). If we still increase the
 8 relative error on $\delta^I(\Psi_l)$ and $\delta^I(\Phi_l)$ to 200% (Fig. 1 (c)) or increase the error for all the waves
 9 to 100% (Fig. 1 (e)), then the probability law for x is definitely not Gaussian any more and
 10 tends to the infinity towards larger values. For a relative error of 50%, the pdf for x can still
 11 be approximated by a Gaussian law (Fig. 1 (d)). In a previous paper, Florsch and Hinderer
 12 (2000) obtained a probability law for x that tends to infinity: that is because of the large errors
 13 they had on the imaginary parts of the smallest diurnal waves (Ψ_1 and Φ_1). As a consequence,
 14 if the estimations of the gravimetric factors were more accurate (i.e. with relative errors
 15 smaller than 50%), then the probability law for x should be Gaussian. Moreover, using VLBI
 16 nutation data, for which the amplitudes of the equivalent Φ_l and Ψ_l nutations are high and
 17 well-determined, Rosat and Lambert (2009) have indeed obtained a Gaussian-law for x .

18 We can also check the influence of the precision on the estimated Ψ_1 gravimetric
 19 factor on the FCN resonance parameters. We use the generalized non-linear least-squares
 20 formulation of Tarantola and Valette (1982b) and we make the Ψ_1 gravimetric factors varying
 21 around its observed value at Strasbourg ($\tilde{\delta}_{\psi_1} = 1.2684 \pm 0.004 + i 0.0063 \pm 0.004$) with a
 22 range defined by twice the observed standard deviation. Note that the $\text{std}(\tilde{\delta}_{\psi_1}) = 0.004$ value
 23 attributed as an error on both the real and imaginary parts of $\tilde{\delta}_{\psi_1}$ does not take into account the
 24 ocean loading correction error as it must reflect the minimal error that we have on the

1 determination of the gravimetric factors at the site. The results of the least-squares inversion
 2 lead to the conclusion that, with an error of 0.004 on the real part of $\tilde{\delta}_{\psi_1}$, the *a posteriori* error
 3 on T from the least-squares inversion is about 2.3 days. For $x = \log_{10}(Q)$, the correlation with
 4 $\delta_{\psi_1}^l$ is much stronger and x varies between 4.2 and 4.6, corresponding to Q -values between
 5 15849 and 39811, when $\delta_{\psi_1}^l$ ranges from 0.002 to 0.012. Therefore the uncertainty on the
 6 estimation of Q is large even if the ocean loading correction was perfect. The error on Q is
 7 therefore mostly due to the fitting error of Ψ_l because of its small amplitude. There is also a
 8 strong correlation between $\delta_{\psi_1}^l$ and a^l . Similar computations could be done for K_l and Φ_l .

9 We have shown that the diurnal waves close to the resonance are very sensitive to the
 10 FCN parameters. Besides, some correlations exist between the FCN resonance parameters. As
 11 shown by the tilted shapes in Fig. 1, two strong correlations exist between a^R and T , on the
 12 one hand, and between a^l and x , on the other hand. Physically it comes from the response of
 13 the whole Earth to the FCN that depends also on the complex eigenfrequency of the core
 14 oscillation. Indeed, the damping of the nearly-diurnal free wobble introduced through the
 15 complex eigenfrequency could be due to the dissipation by viscomagnetic and topographic
 16 coupling at the CMB and to the anelasticity (Mathews et al. 2002) while the imaginary part
 17 introduced in the resonance strength, reflects the anelasticity of the Earth which is related to
 18 the imaginary part of the Love numbers. As a consequence, the Q -factor is physically
 19 correlated to the imaginary part of the resonance strength through the anelasticity of the
 20 interface.

21 Florsch and Hinderer (2000) have also shown that the reference gravimetric factor is
 22 correlated to a^R and hence to the real part of $\tilde{\delta}_1$. As we have an uncertainty (usually less than
 23 0.3%) on the SG scale factor, this error will be propagated to the estimation of $\tilde{\delta}_1$. As we will

1 see in the next part, this scaling error of 0.3% is smaller than the uncertainty that we obtain
 2 for $\tilde{\delta}_1$.

3 The exploration performed by Florsch and Hinderer (2000) has shown that the
 4 imaginary part of the strength a^l is close to zero with an error much larger than the value
 5 itself. They proposed to cancel this parameter in order to restrict the scanning to a 3-parameter
 6 space. However, a^l is strongly correlated to Q particularly for Q -values less than 1000: putting
 7 $a^l = 0$ forces Q to be around 2600 (cf. Fig. 2 (a)). Note that in this case (elastic model), the
 8 obtained T value is close to 465 days, which is the value inferred by Sasao et al. (1980) for an
 9 elastic Earth with a hydrostatic CMB flattening.

10 When imposing a^l equal to a value close to the estimation from the 4-D inversion, i.e. $a^l = -5$
 11 10^{-5} deg/h, then we obtain a most probable values for Q of 21241 +/- 6956 (cf. Fig. 2 (b))
 12 which is in agreement with predictions of the MHB model (Mathews et al., 2002) and with
 13 VLBI nutation observations (e.g. Lambert and Dehant, 2007; Koot et al., 2008). In the next
 14 part, we will finally estimate the FCN resonance parameters.

15 To conclude this part, we have seen that the FCN resonance damping is strongly
 16 correlated to the anelasticity of the mantle through the imaginary part of the internal Love
 17 number. The large errors on the determined phases of the diurnal waves close to the resonance
 18 make the parameter x follow a non-Gaussian probability law.

19

20 **5. Combined GGP inversion**

21 Here we propose to analyze, in the spectral domain, 7 tidal European SG observations
 22 to invert the FCN parameters. Contrary to Ducarme et al. (2009), we do not compute averaged
 23 gravimetric factors but combine them by using relation (4) and we suppose that the resonance
 24 strength \tilde{a} is the same for all the sites. Because \tilde{a} corresponds to the mantle response to the

1 FCN, supposing a homogeneous response at all SG sites enables to retrieve the global
 2 parameters (cf. paragraph 2) like the internal pressure Love number at the CMB. The weight
 3 in equation 4 is the combination of the standard deviation of the fitted gravimetric factor by
 4 ETERNA software and of the ocean loading error.

5 The SG sites considered here are Bad-Hambourg (H1, Germany), Moxa (M1,
 6 Germany), Membach (MB, Belgium), Medicina (MC, Italy), Strasbourg (ST, France), Vienna
 7 (VI, Austria) and Wettzell (W1, Germany). H1 (resp. M1, W1) refer to the lower sphere of the
 8 double-sphere SG installed at these German sites. The record length that has been used to
 9 retrieve the gravimetric factors from the ETERNA tidal analysis is larger than 5 years for
 10 every time-series. These sites have been chosen because they are relatively far from the ocean
 11 (except Membach) and their SGs are known to be well calibrated (better than 0.3% accuracy).

12 The raw data have been corrected for gaps, spikes, steps and other disturbances so that
 13 a tidal analysis with ETERNA software (Wenzel, 1996) is possible. Before the tidal analysis,
 14 the minute data are decimated to 1h (using a filter with a cut-off period of 3h). The ETERNA
 15 software then performs a least-squares fit to tides, local air pressure and instrumental drift to
 16 retrieve the complex gravimetric factors, the residual gravity, an adjusted barometric
 17 admittance, and a polynomial drift function.

18 The data to be inverted are the complex gravimetric factors corrected for the ocean
 19 tide loading effect according to FES2004 ocean model (Lyard et al. 2006, for ocean loading
 20 computation see for instance Llubes et al. 2008). We have attributed a nominal error of 0.02
 21 nm/s^2 on the ocean loading correction vector. This value has been roughly estimated based on
 22 the comparison between different ocean tide models. The inversion is carried out for the four
 23 parameters $(x, \sigma_{nd}^R, a^R, a^I)$ on a 101^4 points space. The explored parameter space is defined
 24 based on the results by Florsch and Hinderer (2000), namely x spans from 3 to 8, T varies
 25 between 380 and 520 sidereal days, a^R is ranging from 0.0004 to 0.001 deg/h and a^I varies

1 between -10^{-4} and 10^{-4} deg/h. Therefore we have defined *a priori* locally uniform laws for
 2 these parameters.

3 The pdfs obtained from the Bayesian estimation are plotted in Fig. 3. The mean values
 4 of the FCN parameters computed from the marginal probability laws are $T = 428 \pm 3$ days,
 5 $a^R = (0.667 \pm 0.005) 10^{-3}$ deg/h and $a^I = (-0.492 \pm 0.05) 10^{-4}$ deg/h. The law for Q is not
 6 symmetric (and not Gaussian) so we define confidence intervals: Q is between 7762 and
 7 31989 within 90% C.I. Note the good agreement with the Levenberg-Marquardt least-squares
 8 results represented by the vertical dotted lines.

9 Based on the MHB Earth's parameter values and using Equation (3), this resonance
 10 strength estimate leads to an observed determination of the $\tilde{\delta}_1$ internal pressure gravimetric
 11 factor of $\tilde{\delta}_1 = 0.0306 - i 0.00223$ with an error of $2 \cdot 10^{-4}$. Theoretical values of the rotational
 12 pressure gravimetric factor $\tilde{\delta}_1$ have been computed by Dehant et al. (1993). A first
 13 experimental determination of $\tilde{\delta}_1$ has been proposed by Hinderer et al. (1991b) by stacking
 14 gravity measurements. They have found $\tilde{\delta}_1 \approx 0.043$ and they have estimated that the
 15 imaginary part of $\tilde{\delta}_1$ is 1.6 % of the real part, i.e. around $7 \cdot 10^{-4}$. In our case $\tilde{\delta}_1^I$ is about 7 %
 16 of $\tilde{\delta}_1^R$. From the FCN frequency, the combination of the dynamic flattening of the fluid core
 17 and the compliance (or the displacement Love number h_1^f) can be evaluated
 18 from $\tilde{\sigma}_{nd}^i = -\Omega \frac{A}{A_m} (e^f - \beta)$ (cf. section 2). From our observed value, $(e^f - \beta) \approx 0.002$ which
 19 is in agreement with the MHB values of $e^f = 2.6456 \cdot 10^{-3}$ and $\beta = 6 \cdot 10^{-4}$ (Mathews et al.,
 20 2002).

21

22 **6. Conclusion**

1 The problem of negative Q -values encountered in previous studies has been avoided using
2 the logarithm of Q in the inversion scheme.

3 We have solved the FCN resonance damping equation for four parameters using on the
4 one hand, the more traditional linearized least-squares method optimized with the Levenberg-
5 Marquardt algorithm, and on the other hand, a statistical Bayesian approach. We have
6 demonstrated the good agreement between both results.

7 Because of the correlation existing between the quality factor Q and the imaginary part of
8 the resonance strength a' , and the necessity for the mantle to possess some anelasticity, this
9 latter parameter cannot be neglected and must be considered in the inversion scheme. Because
10 of the large error on the phase of the diurnal waves close to the resonance, the parameter $x =$
11 $\log_{10}(Q)$ does not follow a Gaussian law.

12 From stacking the SG tidal gravity data, we can improve the determination of the FCN
13 resonance parameters, which are in agreement with VLBI nutation determination. In order to
14 further improve these estimates, the ocean loading correction must be improved at these
15 diurnal frequencies. The nutation has the advantage to be less affected by the ocean loading
16 uncertainty as the contribution of the ocean to the Earth's wobble is much smaller than its
17 effect (mainly local) on the surface gravity.

18 Time-varying gravity data appear to provide complementary information to the VLBI
19 nutation data as the latter cannot determine the precession while gravity can determine the K_1
20 tidal wave. Besides, gravity observations enable to retrieve additional Earth's interior
21 parameters like the internal pressure Love numbers.

22

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1

2 **Table captions**

3 Table 1: A summary of various estimates of period and quality factor of the FCN. In addition
4 to theoretical results relative to an elastic Earth, to a slightly inelastic one and to MHB2000
5 model of Mathews et al. (2002), we have added experimental results from the IDA
6 (International Digital Accelerometers) network of spring gravimeters and from VLBI (Very
7 Long Baseline Interferometry). The other results are from superconducting gravimeter (SG)
8 datasets: B = Brussel (Belgium), BH = Bad Homburg (Germany), ST = Strasbourg (France),
9 CA = Cantley (Canada), J = 3 Japanese stations, ES = Esashi (Japan), MA = Matsushiro
10 (Japan), CB = Canberra (Australia), MB = Membach (Belgium).

11

1 **Figure captions**

2 Fig.1 Influence of the error of the imaginary part of the gravimetric factors on the resulting
 3 pdfs for a^R , a^I , x and T retrieved from the Bayesian analysis of synthetic data. The error is
 4 expressed in percent of the gravimetric factor (imaginary part) amplitude. The relative error is
 5 (a) 10% for the 9 diurnal waves Q_I , O_I , M_I , P_I , K_I , Ψ_I , Φ_I , J_I and OO_I ; (b) 10% for the 7
 6 waves Q_I , O_I , M_I , P_I , K_I , J_I , OO_I and 100% for Ψ_I and Φ_I ; (c) 10% for Q_I , O_I , M_I , P_I , K_I , J_I ,
 7 OO_I and 200% for Ψ_I and Φ_I ; (d) 50% for Q_I , O_I , M_I , P_I , K_I , Ψ_I , Φ_I , J_I and OO_I ; (e) 100%
 8 for Q_I , O_I , M_I , P_I , K_I , Ψ_I , Φ_I , J_I and OO_I . The vertical dot lines indicate the values of x , T , a^R
 9 and a^I used to compute the synthetic delta-factors: $x = 4$, $T = 430$ days, $a^R = 6 \cdot 10^{-4}$ deg/h and
 10 $a^I = -5 \cdot 10^{-5}$ deg/h.

11 Fig. 2 Joint and marginal pdfs for the FCN parameters (a^R , T and x) estimated from
 12 Strasbourg SG tidal gravity factors using the Bayesian method with (a) $a^I = 0$ deg/h; (b) $a^I = -$
 13 $5 \cdot 10^{-5}$ deg/h. The Marquardt-least squares inversion results are also indicated as vertical
 14 dotted lines.

15 Fig. 3 Joint and marginal pdfs for the FCN parameters (a^R , a^I , T and x) estimated from
 16 stacking 7 European SG (Bad-Homburg, Moxa, Membach, Medicina, Strasbourg, Vienna,
 17 Wetzell) records using the Bayesian method. The Marquardt-least squares inversion results
 18 are also indicated as vertical dotted lines.

19

Author	Data	T	Q
Neuberg et al. (1987)	Stacked Gravity (B+BH)	431 +/- 6	2800 ± 500
Sasao et al. (1980)	Theory elastic	465	∞
Wahr & Bergen (1986)	Theory anelastic	474	78000
Herring et al. (1986)	VLBI	435 +/- 1	22000-10 ⁵
Cummins & Wahr (1993)	Stacked Gravity IDA	428 +/- 12	3300-37000
Sato et al. (1994)	Stacked Gravity J	437 +/- 15	3200- ∞
Defraigne et al. (1994)	Stacked Gravity VLBI Stacked Gravity + VLBI	424 +/- 14 432 +/- 4 433 +/- 3	2300-8300 Q>15000 Q>17000
Florsch et al. (1994)	Gravity ST	431 +/- 1	1700-2500
Merriam (1994)	Gravity CA	430 +/- 4	5500-10000
Hinderer et al. (1995)	Stacked Gravity (ST+CA)	429 +/- 8	7700-∞
Roosbeek et al. (1999)	VLBI	431-434	-
Florsch and Hinderer (2000)	Gravity ST (Bayes)	428	Q>20000
Hinderer et al. (2000)	Gravity + VLBI	431-434	15000-30000
Mathews et al. (2002)	MHB2000 model	430.20 +/- 0.28	20000
Sato et al. (2004)	Stacked gravity (ES+MA+CB+MB)	429.7 +/- 1.4	9350-10835
Vondrák & Ron (2006)	VLBI	430.32 +/- 0.07	20600 +/- 340
Ducarme et al. (2007)	Mean gravity	429.7 +/- 2.4	Not estimated
Lambert & Dehant (2007)	VLBI	430 +/- 0.4	17000 +/- 3000
Ducarme et al. (2009)	Mean gravity in Europe	430 +/- 2	15000 +/- 8000
Koot et al. (2008)	VLBI (Bayes)	430	13750 +/- 514
This paper	Stacked gravity of 7 European SGs (Bayes)	428 +/- 3	7762<Q<31989 (90% C.I.)

Table 1

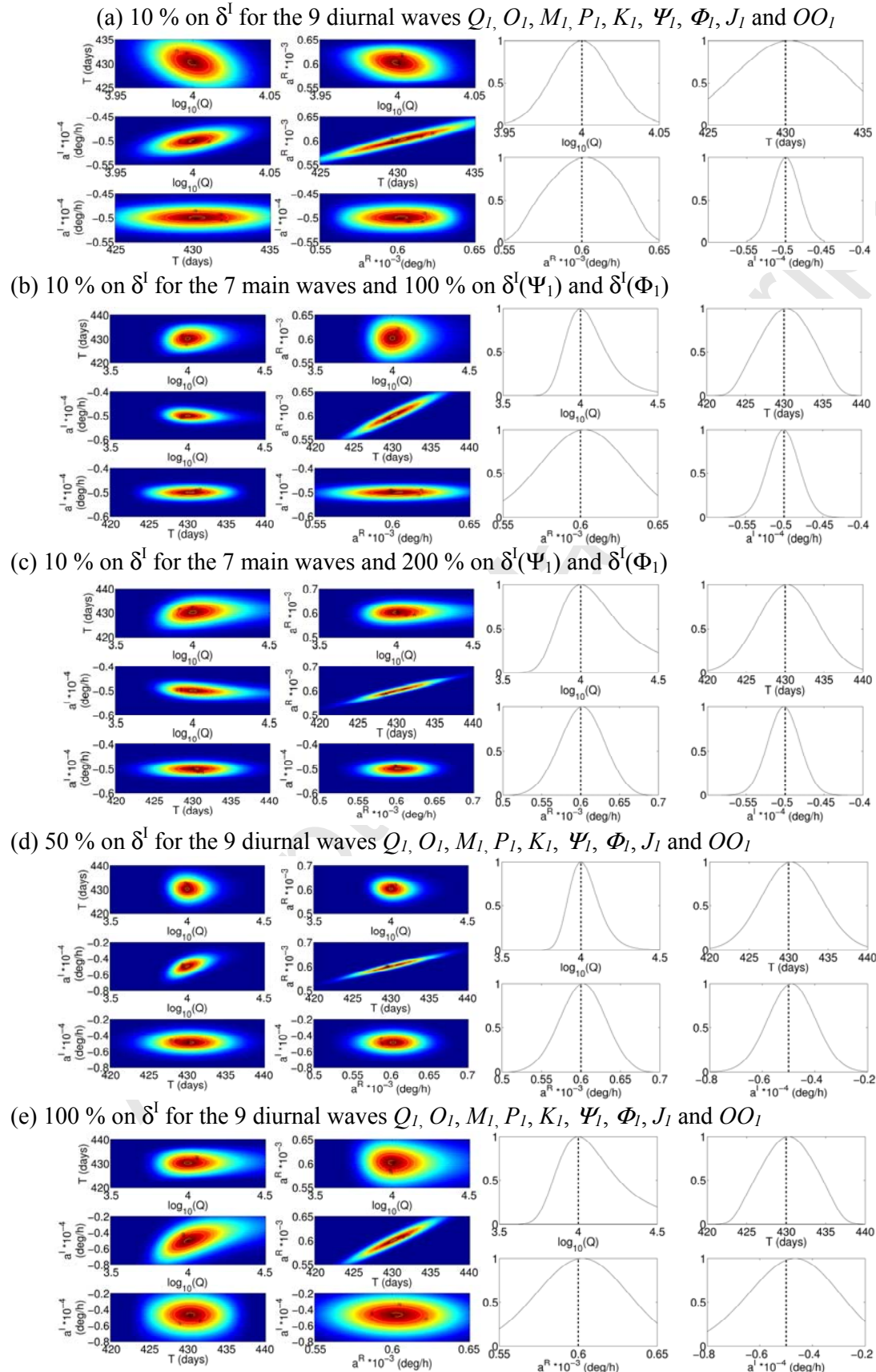
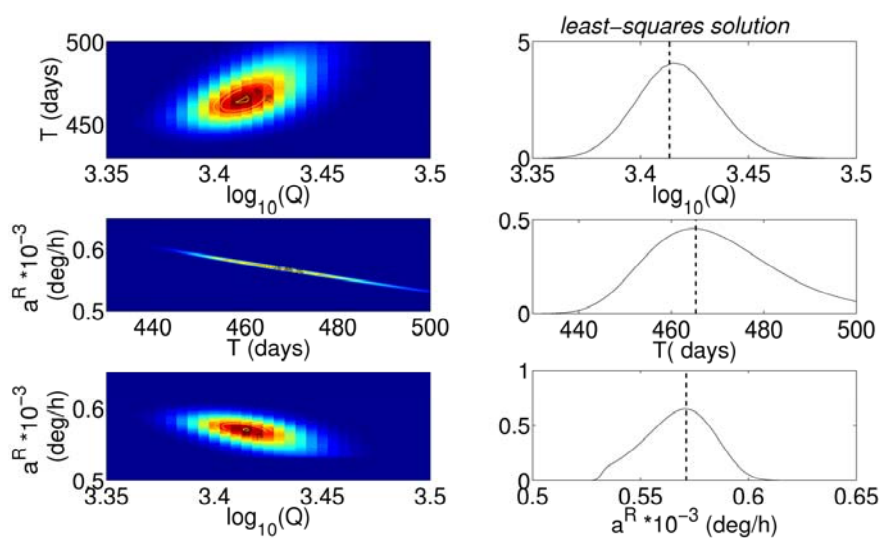


Fig. 1

(a)

Strasbourg – 3 parameters – $a^I = 0$ deg/h

(b)

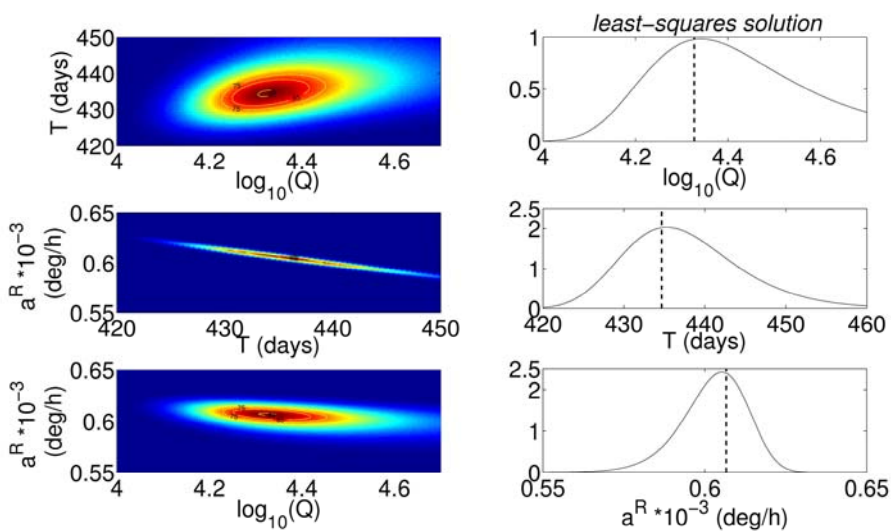
Strasbourg – 3 parameters – $a^I = -5 \cdot 10^{-5}$ deg/h

Fig. 2

Stack of 7 European GGP-time series

(H1, M1, MB, MC, ST, VI, W1)

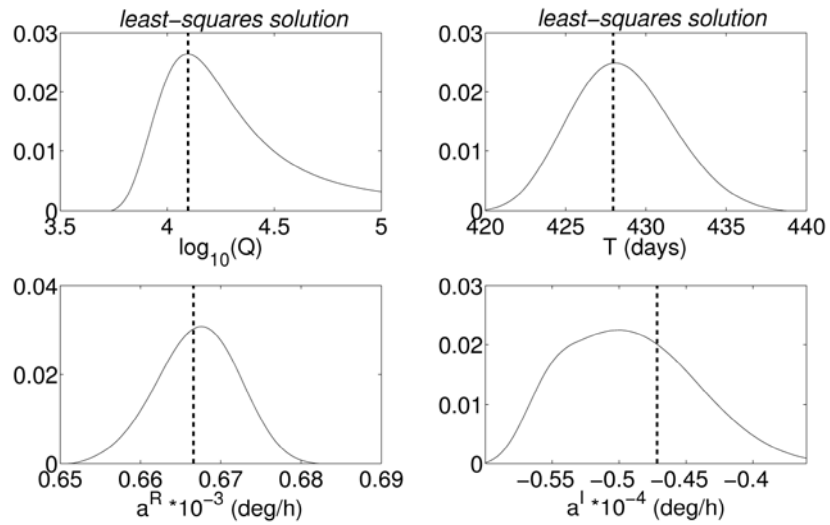
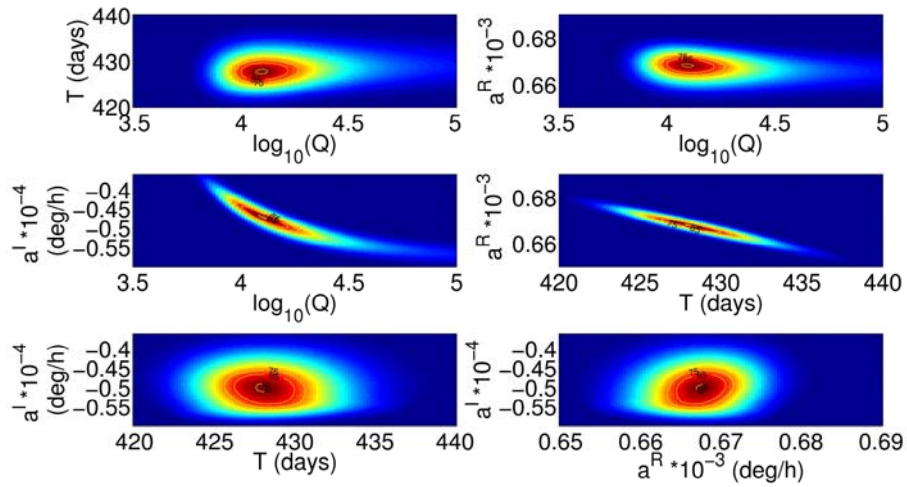


Fig. 3