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► **To cite this version:**

Stuart Crampin, Yuan Gao. A review of a quarter century of International Workshops on Seismic Anisotropy in the crust (0IWSA–12IWSA). *Journal of Seismology*, 2008, 13 (2), pp.181-208. 10.1007/s10950-008-9121-3 . hal-00478436

HAL Id: hal-00478436

<https://hal.science/hal-00478436v1>

Submitted on 30 Apr 2010

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A review of a quarter century of International Workshops on Seismic Anisotropy in the crust (0IWSA–12IWSA)

Stuart Crampin · Yuan Gao

Received: 9 June 2007 / Accepted: 10 January 2008 / Published online: 10 September 2008
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Abstract In 25 years, the presence of azimuthally varying seismic anisotropy throughout the Earth's crust has progressed from general denial to universal acceptance, so that many international geophysical meetings now have sessions on seismic anisotropy. Over this period, the proceedings of the biennial series of International Workshops in Seismic Anisotropy (IWSAs) have captured many of the notable advances in the theory, calculation, observation and interpretation of particularly shear-wave splitting (seismic birefringence) in the Earth's crust. Shear-wave splitting is the almost-infallible indicator of seismic anisotropy

along the ray path. This paper reviews 13 IWSA meetings (0IWSA–12IWSA) as a catalogue of 25 years of progress in seismic anisotropy. The evidence now suggests that shear-wave splitting monitors the low-level pre-fracturing deformation of the stress-aligned fluid-saturated microcracks pervading almost all in situ rocks in the crust. Shear-wave splitting indicates that microcracks are so closely spaced they are critical systems with all the universality, calculability, predictability, “butterfly wing's” sensitivity, and deterministic chaos that that implies. This leads to a New Geophysics, where low-level deformation can be monitored with shear-wave splitting, future behaviour calculated–predicted with the anisotropic poro-elastic model of rock evolution, and in some circumstances even potentially controlled by feedback. We anticipate the New Geophysics will greatly invigorate IWSA.

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Keywords Anisotropy · International Workshop on Seismic Anisotropy (IWSA) · Shear-wave splitting · New Geophysics

1 Introduction

The first International Workshop on Seismic Anisotropy (1IWSA) was held in Suzdal, USSR, in 1982. The next, 2IWSA in 1986 in Moscow,

was the first of 20 years of biennial meetings leading to 11IWSA 2004 in St Johns, Newfoundland and 12IWSA 2006 in Beijing. There was also a precursory meeting in 1975 in Paris that can justifiably be labelled 0IWSA. The proceedings of this quarter of a century of IWSA meetings have captured most of the significant markers spanning our understanding of seismic azimuthal anisotropy and shear-wave splitting in the upper half of the crust. Nowadays, stress-aligned seismic anisotropy is widely recognised in almost all rocks in the upper crust, so that international meetings, in both earthquake and particularly exploration seismology, frequently have sessions on seismic anisotropy. Here, we review the past quarter of a century of research into seismic azimuthal anisotropy and shear-wave splitting as marked by IWSA meetings. This review is timely, as it is now argued that the crack-critical crust, as revealed by shear-wave splitting, leads to a fundamental revision of solid earth geoscience.

The continuing IWSA workshops have been in various countries, organised by various people and sponsored by various organisations. Details are listed in Table 1, and the “Appendix” records how IWSA workshops began.

The first 11 IWSAs (including 0IWSA) had proceedings (11IWSA had abstracts, and proceedings of 12IWSA are currently in press, including this issue) published in a variety of international research journals (Table 1). These proceedings contain many papers crucial to the development of the theory, observation and interpretation of seismic anisotropy and outline the overall development of particularly azimuthal anisotropy in both earthquake and exploration seismology. Initially, IWSA meetings were almost entirely devoted to earthquake seismology. After 3IWSA, however, oil company sponsorship led to the almost complete dominance of exploration seismology. The reduction in earthquake studies is probably an overall loss to seismic anisotropy (and to IWSAs). We need to gain information about inaccessible subsurface rocks from as many sources as possible. The difference in pore fluids between hydrocarbons and water-based salt solutions probably makes little change to the structure or behaviour of fluid-saturated cracks, except during hydrocarbon production procedures.

Here, we review the series of IWSA meetings and identify papers, both at IWSA meetings and elsewhere, that have had a significant effect on the development of the understanding of seismic azimuthal anisotropy in the crust. Many of the important papers in anisotropy have appeared in IWSA proceedings. There are two principal exceptions: Thomsen (1986) defines the *Thomsen parameters* convenient for describing reflection surveys in the transverse isotropy of sedimentary sequences in seismic exploration and is the most highly cited paper in *Geophysics*; Crampin (1981) is an early review of wave propagation in cracked and anisotropic rock and is another highly cited paper.

2 International workshops on seismic anisotropy

2.1 0IWSA, 1975, Paris, France: proceedings in *Geophys. J. R. Astron. Soc.*, 49, 1977

In 1975, the International Association of Seismology and Physics of the Earth’s Interior Commission for Controlled Source Seismology held a meeting at the Institut de Physique du Globe, Université Paris, organised by David Bamford, on Seismic Anisotropy and its Implications. This meeting, 0IWSA (Table 1), was a direct precursor to 1IWSA. 0IWSA was the first meeting at which there had been a session, let alone a whole meeting (albeit only 1 day), devoted to seismic azimuthally varying anisotropy. At that time, azimuthal anisotropy in the Earth’s crust had not been confirmed. In 1975, observations of seismic anisotropy in the Earth were confined to a comparatively thin layer at the top of the upper mantle, identified by two phenomena. Velocity variations of *Pn* waves propagating in oceanic basins displayed azimuthal velocity anisotropy, where the velocity perpendicular to the spreading centre are typically 5% greater than in directions parallel to the centre (Hess 1964; Raitt et al. 1969; and others). The second phenomenon was higher-mode surface-wave particle motions showing anisotropic coupling between higher-mode Love and Rayleigh waves which are controlled by the anisotropy in the uppermost few kilometres

of the mantle (Crampin 1966; Crampin and King 1977). There were several earlier tentative reports of possible shear-wave splitting in the crust, but these were isolated unconfirmed examples—see the introduction to *0IWSA: Seismic anisotropy—the state of the art* (Bamford and Crampin 1977).

Theory Three papers by Keith and Crampin (1977a, b, c) calculated synthetic seismograms of (plane wave) body-wave propagation in multi-layered azimuthally anisotropic media. Today, shear-wave splitting is the most diagnostic and widespread observation of azimuthal anisotropy in both the crust and upper mantle but, at that time (1975), anisotropy-induced shear-wave splitting was not a well-recognised phenomenon. The term *shear-wave splitting* for the differencing of shear-wave velocities in anisotropic media had been used (specifically *S-wave splitting*) by Crampin (1978) who calculated synthetic seismograms and polarisation diagrams (hodograms) of propagation through aligned cracks.

Observation and interpretation Bamford (1977) reported observations in the continental upper mantle of *Pn*-wave azimuthal velocity anisotropy from explosion refraction studies. These had similar characteristics to earlier examples of *Pn*-wave velocity anisotropy in oceanic basins from *Pn* refraction studies Hess (1964), which indicate a possibly thin layer of azimuthal anisotropy at the top of the upper mantle under oceanic basins.

Crampin and King (1977), using the techniques of Crampin and Taylor (1971), calculated the dispersion of surface waves in multi-layered anisotropic media, modelling previous observations of coupling between higher-mode Love and Rayleigh surface waves (Crampin 1966). Crampin and King showed that the coupling could be the result of as little as a 4-km thickness of 4% shear-wave velocity anisotropy at the top of the upper mantle analogous to the *Pn*-wave velocity anisotropy observed by Hess (1964) in oceanic basins and Bamford (1977) in the continental upper mantle. Fuchs (1977) discussed the possible causes of *P*-wave azimuthal anisotropy in the upper mantle in terms of dynamical processes.

Schlue and Knopoff (1977) inferred anisotropy in the upper mantle from the inability of isotropic

inversion to yield consistent structures from both Love and Rayleigh surface-wave phase velocity dispersion across the Pacific Basin. This phenomenon, known as *Rayleigh–Love polarisation anisotropy*, was suggested by McEvelly (1964) and was later used extensively to imply upper-mantle anisotropy.

2.2 1IWSA, Suzdal, USSR, 1982: proceedings in *Geophys. J. R. Astron. Soc.*, 76, 1984

The first International Workshop on Seismic Anisotropy, 1IWSA, was held in Suzdal, USSR (Table 1). Again a *state-of-the-art* paper summarised the current position (Crampin et al. 1984a). Byerly (1938) had suggested that double refraction (shear-wave splitting) might be caused by varying shear-wave travel times, and Nur and Simmons (1969) demonstrated anisotropy in laboratory stress cells. Positive identification of seismic azimuthal anisotropy in the Earth were made by observations of shear-wave splitting in the crust and upper mantle by Crampin et al. (1980) and Ando et al. (1980, 1983), respectively.

Theory Crampin (1984a) surveyed wave propagation in cracked and anisotropic media in a brief summary of Crampin (1981)—a comprehensive review—which had been published elsewhere. A significant advance in calculating anisotropy was rotating elastic tensors so that the horizontal projection of wave propagation was in the *x*-direction, with *z* vertical. This meant that analytical expressions and computer programs in multi-layered models could be written in concise general forms by making use of the summation convention for repeated suffices. This procedure is well adapted for computer manipulation and was one of the key features permitting the numerical developments in Crampin (1981) and many other theoretical and numerical calculations.

Červený and Firbas (1984) and Petrashen and Kashtan (1984) presented the theory for body-wave propagation in inhomogeneous anisotropic media but did not calculate synthetic seismograms. Martynov and Mikhailenko (1984) presented the theory for wave propagation in inhomogeneous anisotropic media and used a finite-difference technique to calculate synthetic

Table 1 The International Workshops on Seismic Anisotropy

ID	Year	Place	Organisers	Sponsors	Proceedings	No. papers	No. pages	Editors
0IWSA	1975	Paris	Bamford, D.	IFP, IASPEI	1977 Geophys. J. R. Astron. Soc., 14 49: 1–243	14	243	Bamford, D.
1IWSA	1982	Suzdal, USSR	Magnitsky, V.A., Chesnokov, E.M. and Crampin, S.	Acad Sci USSR, IASPEI and UNESCO	1984 Geophys. J. R. Astron. Soc., 28 76: 1–272	28	272	Crampin, S., Hipkin, R.G. and Chesnokov, E.M.
2IWSA	1986	Moscow, USSR	Magnitsky, V.A., Chesnokov, E.M. and Crampin, S.	Acad Sci USSR, Amoco, IASPEI and UNESCO	1987 Geophys. J. R. Astron. Soc., 30 78: 261–554	30	293	Booth, D.C., Crampin, S. and Chesnokov, E.M.
3IWSA	1988	Berkeley, CA, USA	Leary, P.C., Crampin, S. and McEvilly, T.V.	AGU, Amoco, Arco, Chapman Conference, Exxon, SEG, Shell, Unocal and Western Atlas	1990 J. Geophys. Res., 95: 11,105–11,358	30	253	Leary, P.C., Crampin, S. and McEvilly, T.V.
4IWSA	1990	Edinburgh, UK	Crampin, S. and Lovell, J.H.	Amoco, Arco, BGS, CGG, Conoco, EAP, Exxon, PSTI, Schlumberger, Western Atlas	1991 Geophys. J. Int., 107: 385–714	30	329	Crampin, S. and Lovell, J.H.
5IWSA	1992	Banff, Canada	Brown, R.J. and Lawton, D.C.	Amoco, Arco, Chevron, Conoco, CREWES, CSPG, Gulf, NSERC, Schlumberger, Shell, Western Atlas	1993 Can. J. Expl. Geophys., 29: 1–390	35	390	Brown, R.J. and Lawton, D.C.
6IWSA	1994	Trondheim, Norway	Fjær, E., Holt, R.M. and Rathore, J.S.	Amoco, Elf, Norsk Hydro, Saga, Statoil,	1996 Seismic Anisotropy, Soc. Expl. Geophys., 1–763	22	763	Fjær, E., Holt, R.M. and Rathore, J.S.

7IWSA	1996	Miami, FL, USA	Berge, P., Cheng, A.H., Hood, J.A., Lynn, H.B. and Schoenberg, M.	Amoco, Chevron, Conoco, Schlumberger, Western Atlas	2001 Advances in Anisotropy. Soc. Expl. Geophys., Open File Publ., 5:1–322	18	322	Hood, J.A.
8IWSA	1998	Boussens, France	Arnaud, J., Helbig, K., Rasolofosaon, P. and Thomsen, L.	Amoco, Arco, CGG, Chevron, Elf, Geco-Prakla, IFP	1998 Spec. Issue., Rev. Inst. Franc. Pet., 53: 539–763	26	224	Rasolofosaon, P.
9IWSA	2000	Camp Allen, Houston, TX, USA	Arnaud, J., Dellinger, J., Ikelle, L., Lynn, H., MacBeth, C., Thomsen, L. and Tsvankin, I.	BP, Schlumberger, Texas A & M University, Total-Elf	2000 Anisotropy 2000: Fractures, Converted Waves, and Case Studies, Soc. Expl. Geophys., Open File Publ., 6: 1–425	25	425	Ikelle, L.T. and Gangi, A.
10IWSA	2002	Tutzing, Germany	Gajewski, D., Vannelle, C. and Pšeněk, I.	BP, EAGE, IASPEI, PGS,RWE-DEA, Statoil, TotalFinaElf, Veritas, WIT	2003 Spec. Issue., J. Appl. Geophys., 54: 161–454	23	293	Gajewski, D., Vannelle, C. and Pšeněk, I.
11IWSA	2004	St John's, Newfoundland, Canada	Bona, A., Brown, R. J., Harris, J.J., Helbig, K., Kendall, R., Pšeněk, I., Slawinski, M. and Tsvankin, I.	BP, Input/Output, Veritas	2006 Abstracts in Geophysics, 71: 13JF–29JF	Abst 43	17	Grechka, V., Helbig, K. and Pšeněk, I.
12WSA	2006	Beijing, China	Ding, Z. L., Gao, Y., Liu, E., Zhang, Z.-J. and 16 others	– ^(a)	2006 Volume of Expanded Abstracts 2007 J. Seism. and J. Seism. Explor., in press	Exp.Abst 233 112	– ^(b)	

^aChinese Acad. Sci.; National Natural Sci. Foundation, China; China Earthquake Admin.; China Univ. Mining and Tech.; Jilin Univ., State Key Lab. Lithospheric Evolution; Inst. Earthquake Sci., China Earthquake Admin.; Edinburgh Anisotropy Project, Brit. Geol. Survey; IASPEI; SEG

^bJ. Seism. Explor., Badal, J., Booth, D.C., Savage, M. and Zhang, Z. J. Seism., Liu, E.

seismograms in a transversely isotropic half-space with a vertical axis of symmetry.

Evans (1984) recognised that the restriction of Nuttli (1961) to surface observations of teleseismic shear-wave arrivals also applied to higher-frequency shear-wave arrivals from crustal earthquakes. The restriction limits useful shear-wave arrivals at a free surface to angles of incidence less than $\sin^{-1} V_s/V_p$ ($\approx 35^\circ$, for a Poisson's ratio of 0.25; Booth and Crampin 1985). This solid angle where true shear-wave arrivals are visible was later named the *shear-wave window*. Shear waves arriving outside the window are severely distorted by *S*-to-*P* conversions at the free surface. Note that refraction through near-surface low-velocity layers frequently means that the effective shear-wave window above small earthquakes can often be extended to an effective straight-line shear-wave window of 45° or 50° (Booth and Crampin 1985). The area of the shear-wave window may also be severely distorted by the typical irregular surface topography above most earthquakes.

Crampin (1984b) used the equations of (Hudson 1980, 1981) to derive anisotropic elastic moduli for both liquid- and gas-filled (dry) stress-aligned cracks—a problem first suggested to Hudson (1981) by Crampin. The effectiveness of these developments for the calculation of seismic propagation through cracked structures has been confirmed in many publications and Hudson (1981) is now the standard much-cited reference. Valid for crack densities up to ~ 0.05 , Hudson (1981) is often ineffectually challenged by attempts to extend the formulation to values beyond fracture criticality limit (~ 0.055) where no theory can correctly model elastic propagation in unstable material (Grechka 2006; Zheng et al. 2006). Crampin (1978) calculated synthetic seismograms through cracked media using anisotropic elastic moduli derived from the analyses of Garbin and Knopoff (1973, 1975a, b), which led to results compatible with those of Hudson (1981).

Using cracks to specify seismic anisotropy was an important advance for studies of anisotropy for at least two reasons. Distributions of stress-aligned cracks are now the accepted source of the shear-wave splitting observed in exploration surveys by the oil industry (see “Sections 2.4”,

4IWSA, and “Section 2.5”, 5IWSA, below) and above small earthquakes (Crampin 1994). The anisotropy of cracks also meant that anisotropic elastic moduli could be easily modified in meaningful ways in terms of crack density, crack aspect ratio and crack orientation, so that shear-wave splitting could be directly related to stress, porosity and other physical parameters. Since the elastic moduli of crystals are typically fixed and invariable, specifying cracks meant that shear-wave splitting became a flexible investigative tool.

Observation—field The proceedings contained 11 further papers on various observations of *P*-wave velocity anisotropy and shear-wave splitting in the crust and upper mantle and Rayleigh–Love polarisation anisotropy in the upper mantle.

Observation—laboratory Several authors discussed velocity anisotropy in laboratory tests of crystalline rocks. Christensen (1984) made comprehensive analyses of crystalline anisotropy in ultra-mafic upper-mantle rocks and suggested that shear-wave splitting would provide information about the mineral orientation and composition of the upper mantle, as indeed is now claimed. Babuška (1984) examined *P*-wave anisotropy of other crystalline rocks, and Babuška and Proz (1984) examined the closure of microcracks at increasing lithostatic pressures and concluded that *P*-wave anisotropy is only observed near the surface at low lithostatic pressures. Note however that these experiments were in drained specimens where fluids could escape. Cracks in undrained specimens do not close completely and shear-wave splitting believed to be due to stress-aligned *fluid-saturated* cracks is now observed throughout the crust.

Interpretation—aligned cracks Crampin et al. (1984b) suggested that observations of shear-wave splitting indicated that rocks in stressed fault zones are pervaded by fluid-saturated microcracks (named as *extensive-dilatancy* anisotropy or EDA). Stress-aligned shear-wave splitting is now seen in almost all rocks in the crust; see particularly Crampin (1993a, 1994). It is now generally, although not universally, accepted by both exploration and earthquake seismologists that the

shear-wave splitting observed in almost all rocks is caused by stress-aligned fluid-saturated cracks and microcracks. Clearly, EDA cracks would be highly compliant and Crampin et al. (1984b) stimulated the search for temporal variations in shear-wave splitting before earthquakes. This later led to observations of temporal variations before earthquakes by Peacock et al. (1988), Crampin et al. (1990, 1991), Booth et al. (1990) and others and a successfully stress-forecast earthquake (Crampin et al. 1999), as well as variations in hydrocarbon reservoirs caused by fluid injection modelled by Angerer et al. (2002).

2.3 2IWSA, 1986, Moscow, USSR: proceedings in Geophys. J. R. Astron. Soc., 91, 1987

The second International Workshop was in Moscow in 1986 (Table 1). As was anticipated in the *state-of-the-art* paper for 1IWSA (Crampin et al. 1984a), increasing interest in anisotropy meant that the whole field could no longer be reviewed in a few pages, and state-of-the-art papers were discontinued. As an example of the rapidly increasing interest, there was one report of shear-wave splitting in the crust in 1980, three in 1985 and about 30 in 1986 (Booth et al. 1987).

Theory Several papers discussed various techniques for calculating full-wave synthetic seismograms in anisotropic structures including double-contour integration (Taylor 1987) and a ray algorithm for 3-D laterally inhomogeneous anisotropic layers (Gajewski and Pšenčík 1987), both of which were part of continuing developments.

Observation The Turkish Dilatancy Projects (TDP) were three deployments of radio-linked seismic networks spanning the North Anatolian Fault east of the Marmara Sea (TDP1 in 1979, TDP2 in 1980 and TDP3 in 1984). These experiments were designed to search for shear-wave splitting above a persistent swarm of small earthquakes some 10 km SE of Izmit. (Coincidentally, the TDP networks happened to span the epicenter, some 20 years later, of the devastating 1999, $M7.5$, Izmit earthquake. Unfortunately,

the TDP sites have not been re-occupied). Shear-wave splitting was indeed confirmed in the crust by Crampin et al. (1980), and five papers, following Crampin et al. (1985), reporting TDP1 and TDP2, set many of the parameters for observations of shear-wave splitting above small earthquakes. Evans et al. (1987) reporting TDP3 confirmed and enhanced previous results of TDP1 and TDP2.

TDP3 also deployed magneto-telluric (MT) stations (Evans et al. 1987). These MT sites showed regional contributions to the geoelectric-strike azimuths approximately parallel to the shear-wave polarisations. Local contributions to the geoelectric-strike azimuths appeared to be *exactly* along the maxima of shear-wave polarisations, except at one station where there was severe topographic irregularity as it was on the edge of a ~ 200 -m-deep gorge. (The effect of the shear-wave window makes observations of shear-wave splitting extremely sensitive to near-station topography, and this was not recognised at the time of deployment). Anomalies in electromagnetic signals associated with earthquakes were directly correlated with stress-aligned microcracks.

Chen et al. (1987), aware of the in-house results of Peacock et al. (1988), also searched for temporal variations in the TDP records. In the 5-month deployment of TDP3, Chen et al. found an overall increase in time delays between split shear waves in time delays in Band-1 directions within the shear-wave window normalised to millisecond per kilometre. Band-1 is the double-leafed solid angle 15° to 45° to the plane of the inferred parallel cracks (Crampin 1999). Increases in average time delays in this solid angle indicate increasing stress as in the accumulation of stress before earthquakes. The Band-1 time delays showed an increase at five out of seven seismic stations, suggesting stress accumulation before an impending earthquake. This is likely to be associated with stress accumulation before an $M4.8$ earthquake, some 45 km SW of the centre of the network, on 28th October, 1984, 4 days after the monitoring had ended. At that time, possible associations had not been identified, as the great sensitivity of fluid-rock deformation to comparatively small stress changes at substantial distances had not yet been recognised.

Observation Leary et al. (1987) and Li et al. (1987) presented observations and modelling of seismic propagation in the vicinity of the seismically active Oroville Fault in California and provided information about the aligned fractures surrounding the fault. Three papers reported various observations of anisotropic effects in USSR, including the very detailed multi-azimuth reflection surveys and vertical seismic profiles (VSPs) in transverse isotropy by Galperina and Galperin (1987).

Interpretation Crampin (1987) speculated on how changes in shear-wave splitting could forecast earthquakes (first suggested by Crampin 1978 and Crampin et al. 1984b). These ideas eventually led to observations of temporal changes (with hindsight) in Band-1 time delays before many earthquakes, including one successful stress forecast of an $M5$ earthquake in Iceland (Crampin et al. 1999). Note that Gupta (1973a) had earlier claimed to have observed temporal variations in shear-wave splitting before earthquakes. Unfortunately, Gupta (1973a) had misinterpreted laboratory experiments (Gupta 1973b) and identified crustal observations of shear-wave splitting so far outside the effective shear-wave window that they were completely uninterpretable in terms of our current understanding of shear-wave splitting (Crampin et al. 1981; Crampin 1994).

Note that Thomsen (1986) was an important paper, published elsewhere, defining appropriate parameters for exploration processing of weakly anisotropic transversely isotropic media with a vertical axis of symmetry (transverse isotropy (*TIV*) anisotropy). The *Thomsen parameters* are one of the major factors that opened up TIV anisotropy to exploration seismologists as is evident in 4IWSA and thereafter elsewhere.

2.4 3IWSA, 1988, Berkeley CA, USA: proceedings in J. Geophys. Res., 95, 1990

The Third International Workshop was in the Berkeley Campus of the University of California, USA (Table 1). The reports at this meeting were almost exclusively on laboratory and field observations of fracture and particularly microcrack anisotropy (Leary et al. 1990).

Theory Several papers calculated synthetic seismograms including: Gajewski and Pšenčík (1990) extending previous work in dynamic ray tracing (Gajewski and Pšenčík 1987) to calculate synthetic seismograms in multi-layered anisotropic substrates; Tsvankin and Chesnokov (1990) used double-contour integrals to calculate synthetic seismograms in orthorhombic media.

Liu and Crampin (1990) showed that, just as there is a shear-wave window limiting observations of shear waves at the free surface, there are also internal shear-wave windows limiting observations at each interface due to the different behaviour of *SV* and *SH* waves at horizontal interfaces. In some cases, the behaviour can suggest anisotropy-induced shear-wave splitting at isotropic-to-isotropic internal interfaces.

Observation Shih and Meyer (1990) used an automatic cross-correlation technique (Shih et al. 1989) to measure shear-wave splitting in Long Valley Caldera, CA, USA. They found anomalies associated with faults and larger time delays above the resurgent dome. Larger time delays are now recognised as being associated with rocks with high heat flow (Volti and Crampin 2003a, b). Savage et al. (1990) examined shear-wave splitting in Long Valley Caldera by visual techniques and obtained similar conclusions but attributed the scatter in time delays to near-receiver anisotropy. The large scatter above small earthquakes is now attributed to 90° flips in shear-wave polarisations caused by the high pore-fluid pressures associated with all seismically active faults rearranging microcrack orientations (Crampin et al. 2004). (90° flips are where the faster and slower split shear waves abruptly exchange polarisations, Angerer et al. 2002). Li et al. (1990) modelled observations of VSPs in the Mohave Desert, CA, USA, with synthetic seismograms and found shear-wave splitting consistent with interpretations in terms of sub-parallel sub-vertical cracks in other studies (for example, Evans et al. 1987). Li and Leary (1990) observed and modelled fault-zone-trapped waves in a thin anisotropic wave guide.

Kaneshima (1990) reviewed shear-wave splitting at some 40 seismic stations in Japan. Shear-wave splitting had been observed over a wide range of source zones and azimuths and

Kaneshima concluded that shear-wave splitting may be limited to the uppermost 15–25 km of the crust.

Interpretation Shepherd (1990) gave a comprehensive geological analysis of microcracks, fluid inclusions and stress which generally supported the interpretation of shear-wave splitting in terms of propagation through sub-parallel sub-vertical fluid-saturated microcracks.

In an integrated study, Queen and Rizer (1990) found that the orientations of pressure-cracked core-sample discs matched the orientations of shear-wave splitting at the Conoco Borehole Test Facility in Oklahoma. This was a direct correlation of shear-wave splitting with microcrack lithology and suggested that the shear-wave splitting was directly related to overall fracture orientations and densities, where shear-wave polarisations indicate directions of preferred fluid flow. This has now been widely established in production reservoirs and elsewhere.

Crampin et al. (1990) showed that, when more data became available, the statistically significant increase of Band-1 time delays identified by Peacock et al. (1988) at Station KNW of the Anza Seismic Network, Southern California, was immediately followed by the M 6 North Palm Springs of 8th July, 1986, ~30 km distant from KNW. (The aftershocks of the earthquake had previously overloaded the location programme in California, which was why Peacock et al. could not recognise the association). This was the first time that temporal changes in Band-1 time delays monitoring the accumulation of stress before earthquakes had been observed (with hindsight) before earthquakes as suggested by Crampin (1978, 1987) and Crampin et al. (1984b). Similar increases were also seen (with hindsight) in Band-1 time delays before an M 3.8 swarm earthquake in Enola, AR, USA by Booth et al. (1990).

Note that the interpretation of Peacock et al. (1988) and Crampin et al. (1990) was challenged by Aster et al. (1990), who used an automatic measurement technique to infer that there were no temporal changes. Crampin et al. (1991; reply, Aster et al. 1991) showed that visual inspection of rotated seismograms indicated that the comments of Aster et al. (1990) could be discounted as some

of their automatic measurements were in error by up to 200%. Note that shear-wave splitting above small earthquakes is difficult to measure automatically on typical seismograms in the crust (Crampin and Gao 2006a). No wholly successful fully automatic technique has yet been established, except when there is rigorous selection of seismograms and rejection of, typically, well over 50% of the records. Except for vigorous aftershocks sequences, there are seldom enough earthquakes within the shear-wave window to make this useful, and aftershock sequences do not appear to have consistent patterns of stress variations. Note that an effective *semi-automatic* technique has now been developed (Gao et al. 2006).

2.5 4IWSA, 1990, Edinburgh, UK: proceedings in Geophys. J. Int., 107, 1991

The Fourth International Workshop on Seismic Anisotropy, 4IWSA, was held in Edinburgh (Table 1). Previous IWSAs had principally discussed earthquake seismology. However, shear-wave splitting had now been observed in the oil reservoirs. The 1986 Annual International Society of Exploration Geophysicists (SEG) Meeting had 15 abstracts on exploration-based crack-induced shear-wave splitting—see for example, Alford (1986) and Lynn and Thomsen (1986). Recognising this advance, Crampin founded in 1988 the Edinburgh Anisotropy Project (EAP) at the British Geological Survey. Sponsored by ~18 oil companies, EAP still continues, directed by Xiang-Yang Li. As a result of this change of emphasis, 4IWSA in Edinburgh had well over 50% of the proceedings related to exploration seismology, and all succeeding IWSA have been increasingly dominated by exploration seismology as oil companies provided increasing sponsorship (Table 1).

The first paper, by Crampin and Lovell (1991), reviewed the first decade of observations of shear-wave splitting in the Earth. Crampin and Lovell listed seven then unexplained features of shear-wave splitting ranging from the coincidence of the degree of stress-aligned shear-wave velocity anisotropy in virtually all rock types to the huge ($\pm 80\%$) scatter in shear-wave time delays and polarisations above small earthquakes. It is a mark of

the progress in understanding shear-wave splitting that these seven unexplained features in 1991 have all been largely resolved, principally as a result of the recognition that the fluid-saturated microcracks are so closely spaced that they are critical systems with all the new behaviour that this *New Geophysics* implies (Crampin 1998, 2003a, 2006; Crampin and Chastin 2001; Crampin and Peacock 2005; Crampin et al. 2003, 2004).

Theory Hudson (1991) reviewed the background and overall progress and conditions for modelling cracks in heterogeneous media begun 13 years earlier by Crampin (1978) and Hudson (1980, 1981).

Li and Crampin (1991a, b) developed the linear-transform technique, LTT, to analyse vector-polarised shear-wave data by writing the two horizontal components as complex numbers and successfully applied LTT to four very different case studies. This was the first time that a comprehensive analysis technique had been developed to analyse shear-wave splitting and was important as demonstrating that sophisticated fundamentally new techniques for analysing shear-wave splitting could be developed.

Lou and Crampin (1991) calculated synthetic seismograms and dispersion of interface waves propagating along thin anisotropic wave guides. Such guided waves in cross-hole seismics have been used to test continuity of wave guides in hydrocarbon reservoirs (Leary 2006).

Theory and observation—shear-wave singularities Wild and Crampin (1991) demonstrated the theoretical effects of solids with combinations of transverse isotropy (*periodic thin layer (PTL) anisotropy*) and azimuthal *anisotropy* (known as extensive-dilatancy anisotropy or *EDA anisotropy*) believed to be common in hydrocarbon reservoirs. PTL anisotropy is the transverse isotropy of periodic thin layers and the lithology of shales and mudstones with vertical axes of symmetry. EDA anisotropy is the azimuthal anisotropy of parallel vertical microcracks with horizontal axes of symmetry (Crampin et al. 1984b). One of the principal effects of a combination of PTL and EDA anisotropy is to yield orthorhombic anisotropic symmetry with

numerous (theoretically up to 27) shear-wave point singularities (Crampin and Yedlin 1981).

Point singularities, where the faster and slower shear-wave sheets touch or intersect, can cause severe disturbances to shear-wave ray paths in directions passing within ten or more degrees of the singularity (depending on the relative strengths of PTL and EDA anisotropy). Crampin (1991) used ANISEIS (Taylor 1987, 1990) to calculate exact full-wave synthetic seismograms demonstrating the pronounced disturbances to shear-wave polarisations (including 90° flips, Angerer et al. 2002) and time delays caused by point singularities, including abrupt reversal of signs (positive to negative, say) without passing through zero. Note that both radial and azimuthal integration were used and the full-wave seismograms are exact without approximations. Bush and Crampin (1991) confirmed the effects of point singularities in observations of a multi-offset VSP in the Paris Basin.

Observation The Proceedings of 4IWSA had some ten papers on observations of *P*-wave velocity anisotropy and shear-wave splitting in hydrocarbon reservoirs. There were also several examples of inverting both field and laboratory data for elastic parameters (Artemieva and Chesnokov 1991; Brodov et al. 1991; Brown et al. 1991; de Parscau 1991; MacBeth 1991a, b).

Interpretation In a classic paper, Mueller (1991) showed that the lateral variability of observed shear-wave splitting could be used to guide horizontal drilling into aligned fissures for enhanced oil recovery at depth in the Austin Chalk, TX, USA. This is one of the first published demonstrations of the direct value of azimuthally varying shear-wave splitting to the hydrocarbon industry.

Leary (1991) showed for the first time that both sonic and electrical resistivity well logs display fractal distributions for over three orders of magnitude of crack dimensions. Referred to as *1/f* noise, such effects have now been established as routinely present in virtually all well logs worldwide and are an important demonstration of the critical nature of cracks in the crust (Crampin 1998, 2003a, 2006; Crampin and Chastin 2001; Crampin et al. 2003).

Liu et al. (1991), continuing the investigations of Queen and Rizer (1990), showed the effects of aligned fractures on shear-wave splitting in shallow VSPs at the Conoco Borehole Test Facility.

2.6 5IWSA, 1992, Banff, Canada: proceedings in Can. J. Expl. Soc., 29, 1993

The Fifth International Workshop on Seismic Anisotropy was held in 1992 in Banff, Alberta, Canada (Table 1), with again the majority of papers in exploration seismics, this time with a significant number of case studies. The small number of earthquake studies included the first papers from mainland China. The proceedings of 5IWSA contain ten papers calculating synthetic seismograms (four of which directly model observed shear-wave splitting), and it is difficult to identify which of these papers is most significant for future developments. Note that, in the first 10 years of IWSAs, many of the original features of seismic anisotropy have now been discovered. This means that as time progresses there are smaller numbers of significant papers and it is more difficult to identify those papers which will be important for the future.

Theory Crampin (1993a) presented the arguments for distributions of microcracks (EDA cracks) rather than large cracks or fractures as the source of the nearly ubiquitous stress-aligned shear-wave splitting seen in almost all rocks. The evidence for microcracks has now recently been reviewed by Crampin and Peacock (2005). There are three key features that directly indicate microcracks: (1) the only anisotropy symmetry system with parallel polarisations is transverse isotropy with a horizontal symmetry axis (TIH anisotropy) or a minor perturbation thereof; (2) the only geological phenomenon with this symmetry is stress-aligned microcracks; and (3) fluid-saturated cracks are the source of anisotropy that can display the observed rapid temporal changes. The overwhelming evidence for microcracks is that very similar phenomena are seen at all crustal depths in almost all rocks whose only common feature is the presence of stress-aligned fluid-saturated grain-boundary cracks in crystalline rocks and thin pores and pore throats in sedimentary rocks.

Although larger cracks clearly exist, propagation through distributions of larger cracks is likely to severely attenuate second split shear waves, as demonstrated by Mueller (1991), and cause reflections and refractions but not the ubiquitous shear-wave splitting. The overall conclusion is that most shear-wave splitting is caused by microcracks rather than macrocracks.

Crampin (1993b) gave a theoretical review of the effects of crack parameters on wave propagation through cracked solids. Note that there is no simple relationship between the degree of velocity anisotropy and crack density. Almost counter-intuitively, the degree of velocity anisotropy and behaviour of both P waves and shear waves is *heavily dependent* on crack aspect ratios, pore-fluid velocities and V_s -to- V_p ratios (Poisson's ratio), as well as crack density. This means that unique inversions for crack parameters from observations of shear-wave splitting may be difficult if not impossible without very extensive datasets, which are seldom, if ever, available. There were several other papers about various anisotropic phenomena, including a number of papers about preferred techniques for observing or processing appropriate record sections.

Theory—synthetic seismograms Several 5IWSA papers calculated synthetic seismograms in cracked media with TIV anisotropy. Guest and Kendall (1993) used Maslov ray theory to calculate synthetic seismograms in anisotropic and inhomogeneous structures. Igel et al. (1993) presented a finite-difference grid technique which avoids problems with reflectivity (restriction to plane-layered models), ray techniques (limited heterogeneity) and finite difference (errors in phase and group velocities). The techniques of Igel et al. showed satisfactorily small relative errors but do not present synthetic seismograms. Leary et al. (1993) used finite-difference calculations of synthetic trapped-wave propagation in fractured low-velocity layers.

Wild et al. (1993) used ANISEIS to display instantaneous amplitude and polarisation attributes of synthetic VSPs in layered anisotropic structures (see note on ANISEIS in “Section 2.4”). Yao and Xiong (1993a, b) used the reflectivity technique of Booth and Crampin (1983) to model with

synthetic seismograms radiation from an anisotropic point source and shear-wave splitting above local earthquakes. Zhang et al. (1993) used finite-difference modelling for synthetic seismograms in TIV.

Synthetic seismograms were also used to model observed shear-wave splitting. Gledhill (1993) used the reflectivity technique of Booth and Crampin (1983) through cracked structures (Hudson 1980, 1981) to match synthetic to observed shear-wave splitting in Wellington Peninsula, New Zealand. Liu et al. (1993a) used the ANISEIS software (Taylor 1990) to successfully match synthetic seismograms in a multi-azimuth reverse VSP in the fractured Conoco Borehole Test Facility. Niitsuma et al. (1993) used wavelet transforms to model observed shear-wave splitting in the Kakkonda geothermal area in Japan.

Observation and interpretation There were several unique observations confirming various anisotropic phenomena. Holmes et al. (1993) monitored shear-wave splitting in a sparse but nearly complete solid angle of directions in a granite batholith at an underground research laboratory. (Continuing analysis showed that excavation damage probably extended to at least three diameters from a 3-m-diameter tunnel, Holmes et al. 2000).

Li et al. (1993), continuing the analysis of Mueller (1991), showed that hydrocarbon production rates across the Austin Chalk, TX, USA, approximately correlated with degree of shear-wave splitting in reservoirs at two producing and one non-producing reservoir. This was a direct confirmation that varying degrees of shear-wave splitting are directly correlated with variations in production.

Liu et al. (1993a), see above, successfully matched synthetic seismograms in a multi-azimuth reverse VSP in fractured rock, confirming the association of shear-wave splitting to both micro- and macro-fracturing.

Liu et al. (1993b) monitored shear-wave splitting showing temporal variations, modelling stress accumulation, before a M 4 earthquake at Parkfield on the San Andreas Fault similar to variations reported previously (Peacock et al. 1988; Crampin et al. 1990; Booth et al. 1990).

This was the fourth example of a now frequently observed phenomenon. The third example was by Gao et al. (1998) on Hainan Island, China. Note that other papers at 5IWSA also searched for temporal changes in shear-wave splitting above small earthquakes.

Graham and Crampin (1993) examined regional earthquakes recorded by the TDP experiments in Turkey (Evans et al. 1987) and showed that shear-wave splitting in S_n -waves, refracted along the Moho, indicate substantial crack-induced anisotropy in the lower crust with similar orientation as the upper crust but with larger time delays of up to 1 s.

Observation—anisotropic cusps Slater et al. (1993) successfully modelled synthetic seismograms using ANISEIS software (Taylor 1990) in walkaway VSPs in an oil field in the Caucasus Basin, matching confirmed field observations of anisotropic cusps.

2.7 6IWSA, 1994, Trondheim, Norway:
proceedings on seismic anisotropy, Soc. Expl. Geophys., 1996

The Sixth International Workshop on Seismic Anisotropy, 6IWSA, was in Trondheim, Norway (Table 1). The first paper was *A geophysicist's view on seismic anisotropy* by Lynn (1996). The original earthquake stimulation for IWSA has now almost completely disappeared and Lynn made no mention of earthquake seismology. The next paper was by Helbig (1996) who, at several of the following IWSAs, presents papers outlining the historical background of seismic anisotropy. On this occasion, it was the contribution of William Thomson, later Lord Kelvin (1824–1907), to strain and stress tensors.

There were 20 other papers. Many of these were theoretical modifications of well-known phenomena, whose overall significance is difficult to assess. They included four papers on the theoretical modelling of anisotropic matrices, seven papers modelling wave propagation through material containing aligned cracks, several papers on various record-section processing techniques and several on rock mechanics experiments.

Theory Rasolofosaon and Yin (1996) set up the equations for non-linearity (NL) in anisotropic elastic media and present rock mechanics results that suggest that the sensitivity of the NL response to changing parameters in anisotropy may be far greater than that of the linear parameters such as wave speeds and elastic moduli. This sensitivity was later to be confirmed in several papers (Angerer et al. 2002; Crampin et al. 2003; Volti and Crampin 2003b).

Observation and interpretation There were a small number of case studies. Horne et al. (1996) devised a successful genetic algorithm (GA) technique for inverting for elastic moduli for a 620-m-deep VSP at the Conoco Borehole Test Facility, whose shallow structure was modelled by Liu et al. (1991, 1993a). Kebaili et al. (1996) presented a new slant stack processing technique and successfully applied it to a shallow VSP in Alberta, and Stawicki and Lynn (1996) investigated the effects of complex tectonic strain on multi-azimuth VSPs in the Lower Indus Basin, Pakistan. Mjelde (1996) interpreted shear-wave velocities in the lower crust from two air-gun reflection surveys recorded on three-component ocean bottom seismographs at Lofoten, Norway, in terms of 14% shear-wave velocity anisotropy. This is one of the few examples, together with Graham and Crampin (1993) and Kaneshima (1990), of realistic estimations of lower-crustal shear-wave velocity anisotropy. Both Mjelde and Graham and Crampin agree that the azimuthal shear-wave velocity anisotropy in the lower crust may be over 10% and substantially higher than in the upper half of the crust.

2.8 IWSA, 1996, Miami, FL, USA: proceedings in advances in anisotropy, Soc. Expl. Geophys., 2001

The Seventh International Workshop on Seismic Anisotropy was held in Miami, FL, USA (Table 1). The meeting was now typical in that there were significant contributions on a wide range of topics that are too many to mention individually.

Theory—frequency dependent anisotropy Hornby (2001) found systematic differences between

‘intrinsic’ anisotropy of shales at ultrasonic frequency in cores with seismic-scale anisotropy in walkaway VSPs. The conclusion was that the effects of fine-layer anisotropy on the seismic scale are underestimated in sonic logs from the Dipole Shear Sonic Indicator borehole sonic tool.

Observation—laboratory Skjærstein and Fjær (2001) measured attenuation in the model of open fluid-saturated cracks of Rathore et al. (1995) which had been used to compare various theoretical models of cracks. They confirmed theoretical results that shear waves with polarisations perpendicular to parallel cracks are more attenuated than shear waves with polarisations parallel to the crack face (Crampin 1984b; Hudson 1981).

Theory—temporal changes Bokelmann (2001) presented a method for resolving small temporal variations between clusters of doublets (earthquakes with similar seismograms and similar source). The results showed that small temporal variations are common but Bokelmann offered no explanation. We can now suggest that the sensitivity of critical systems of fluid-saturated microcracks is a probable explanation (see papers in the next three IWSAs).

Observation and interpretation Two papers examined the effects of dipping TIV structures. Kühnel and Li (2001) developed an algorithm for separating anisotropy and structure from: a first-order isotropic dip term; two second-order residual terms, isotropic dip, dip-independent anisotropy; and one third-order dip- and anisotropy-dependent residual. The algorithm is usually effective except for particular orientations of the anisotropy of the matrix material, when the technique is no longer applicable. Leslie and Lawton (2001) propagated *P*-wave signals through an anisotropic laboratory model with structure simulating a dipping sequence of shales. Comparatively severe distortion can be interpreted in terms of anisotropic and dip parameters as long as the degree of anisotropy is not too great.

2.9 8IWSA, 1998, Boussens, France: proceedings in spec. issue, *Rev. Inst. Franc. Pet.*, 53, 1998

The Eighth International Workshop on Seismic Anisotropy was held in Boussens, France (Table 1). Approximately half the papers were on various aspects of theory and half on processing techniques but with very few case studies.

Theory Two papers calculated synthetic seismograms. Caddick et al. (1998) compared asymptotic- with Maslov-ray theory, which give different behaviour near shear-wave singularities. (Note that Crampin 1991 gave exact full-wave behaviour near singularities, using synthetic seismograms calculated by ANISEIS, Taylor 1987, 1990.) Chichinina and Oblentseva (1998) compared waveforms in anisotropic and gyrotropic media. Gyrotropy was not specifically defined but seems to refer to media with a continuous rotation of preferred shear-wave polarisations. As far as we know, such properties have not yet been demonstrated in elastic media nor in in situ rocks.

There were a number of papers presenting variations on standard procedures whose overall significance is difficult to judge. Helbig (1998) presented a formalism for the non-linearity suggested by Rasolofosaon and Yin (1996), and Chesnokov et al. (1998) developed a diagrammatic technique for the calculation of the dynamic properties, including the frequency response, of propagation through randomly distributed cracks and pores.

Observation Berthet et al. (1998) showed that imaging dipping structures was improved by anisotropic rather than isotropic post-stack processing. They demonstrated the improvement in a record section from offshore Africa.

Interpretation Raymer and Kendall (1998) showed synthetic seismograms through models containing salt, with anisotropic properties determined from laboratory texture analysis. They demonstrated that the anisotropy of salt produced substantial travel-time variations and significant shear-wave splitting that could be diagnostic of in situ salt.

Li (1998), assuming fracture-induced anisotropy, showed that near-vertically propagating

converted *PS* waves can indicate the orientation and intensity of in situ fractures. These effects were demonstrated with synthetic (ANISEIS) seismograms showing phase changes and variations in time delays.

Interpretation—critical systems Crampin (1998) suggested that the match of anisotropic poroelasticity (APE) modelling (Zatsepin and Crampin 1997; Crampin and Zatsepin 1997) to observed shear-wave splitting implies that the fluid-saturated stress-aligned microcracks in most rocks in the crust are so closely spaced that they are critical systems. This is now supported by a large variety of observations relating to cracks, stress and shear-wave splitting.

2.10 9IWSA, 2000, Camp Allen TX, USA: proceedings in anisotropy 2000, *Soc. Expl. Geophys.*, 2001

The Ninth International Workshop on Seismic Anisotropy was held in Camp Allen, Houston, TX, USA (Table 1).

Theory Rommel and Tsvankin (2001) developed a simple analytical approach for group- and phase-velocity relationships for rays of *P* waves in TIV and orthorhombic media, based on the notation of Tsvankin (1996) for the Thomsen (1986) parameters applied to orthorhombic media. Helbig and Rasolofosaon (2001), continuing the work of Rasolofosaon and Yin (1996), developed a formalism for incorporating non-linearity and hysteresis into a “unifying theory”. Garmany (2001) showed analytically that phase shifts (90° flips) in shear-wave polarisations occur whenever a ray touches a caustic surface (or passes near a shear-wave singularity as demonstrated with synthetic full-wave seismograms by Crampin 1991).

Theory—AVOs and fractures There were eight papers dealing with various aspects of processing Amplitude Variations with Offset (AVOs) of models with anisotropic-layered structure containing fractures. All seem useful, but space restrictions limit comments to two papers. Ikelle and Amundsen (2001) presented preliminary

techniques for interpreting P - P , P - SV and P - SH data with the ambiguity of TIV anisotropy and heterogeneity. Since anisotropy and heterogeneity may always be present in AVOs, these techniques could be important discriminants. Liu et al. (2001) presented techniques for extracting information about crack compliance and pore-fluid parameters from AVO analysis. Using synthetic seismograms, Liu et al. showed that the combined use of AVOs of P - P and P - SV waves offers advantages in determining fractures and fluid saturations.

Observation—case studies There were two case studies. Granger et al. (2001) used the shear-wave splitting of P -to- S -converted waves (C waves) recorded on 3D–4C ocean bottom geophones to make a preliminary evaluation of azimuthal anisotropy of the North Sea Valhall Field. Results showed that principal directions coincide with fault systems at two levels from this feasibility demonstration, but the authors do not attempt any interpretive conclusions. Berthet et al. (2001) applied anisotropic pre-stack depth migration to offshore data in West Africa. The technique showed sufficient sensitivity that can be used as a lithology discriminate to detect isotropic sands of a few wavelengths thickness.

Interpretation—processing TIV media Five papers described various techniques for processing TIV media. We shall only refer to Williamson and Maocec (2001) who estimated local anisotropy using polarisations and travel times from the Oseberg 3D VSP. They achieved better determination of anisotropy by inverting polarisations than by standard slowness inversion.

Interpretation—critical systems Crampin and Chastin (2001) claimed that three recent results support the arguments of Crampin (1998) that the stress-aligned fluid-saturated microcracks in almost all rocks are so closely spaced that they are critical systems. The papers reported: modelling the response of a reservoir to a high-pressure fluid injection (published later by Angerer et al. 2002); modelling frequency dispersion in laboratory experiments (published later in Chapman et al. 2002); and a successfully

stress-forecast earthquake (Crampin et al. 1999). It is suggested that critical systems of stress-aligned fluid-saturated microcracks are the physical reality underlying the well-known self-organised criticality of the Gutenberg and Richter relationship. They also have a profound effect on a wide range of solid earth geoscience including carbon production and earthquake science (Crampin 2006).

2.11 10IWSA, 2002, Tutzing, Germany: proceedings in spec. issue, J. Appl. Geophys., 54, 2003

The Tenth International Workshop on Seismic Anisotropy, 10IWSA, 2002, was in Tutzing, Germany (Table 1). There were too many useful and interesting papers at this meeting to refer to each individually.

Theory—shear-wave triplication Thomsen and Dellinger (2003) discussed analytically the occurrence of cusps and shear-wave triplications and suggest that P -wave kinematics indicate that cusps in the SV -wave sheet in TIV anisotropy are common. Although such triplications have been observed in a wide-angle VSP (Slater et al. 1993), the angles of incidence at which cusps would be seen are further from the vertical than typical P -to- S -converted waves (C waves) as would usually be sampled in surface seismic surveys.

Theory and observation—frequency-dependent anisotropy Gurevich (2003) developed the mathematics for the elastic properties of saturated porous rock with aligned cracks. Wave-induced fluid flow between pores and cracks in rock with parallel cracks was shown to alter the elastic behaviour for low-frequency waves. Chapman et al. (2003) showed that the movement of fluids between cracks and pore space strongly alters the anisotropic behaviour, where a simplified model matched laboratory data. In particular, the analysis showed the effects are frequency dependent. Li and Yuan (2003) derived improved equations for calculating the moveout and conversion points of the non-hyperbolic C waves. Comparison of analysis with real data shows that the technique is sufficiently accurate to yield reliable structures.

Liu et al. (2003) observed and interpreted frequency-dependent anisotropy in terms of variations of time delays with depth in different frequency pass bands. The anisotropy diminished from 28% at 10 Hz to 2% at 25 Hz.

Observation and interpretation Gao and Crampin (2003) reported temporal variations of shear-wave splitting in field and laboratory studies in China, which show changes of shear-wave splitting with increasing stress and pressure. The laboratory experiments show an abrupt decrease in time delays immediately before fracturing occurs. Such stress relaxation is also observed before earthquakes (Gao and Crampin 2004), where it is now interpreted as the effects of crack coalescence as the impending fracture is identified in the fluid-stress evolution.

Observation and interpretation—critical system sensitivity Crampin et al. (2003; see also Crampin 2004) reported preliminary results from a prototype borehole Stress-Monitoring Site (SMS) adjacent to the Húsavík–Flatey Transform Fault in Northern Iceland. Well-recorded anomalies (accurate to ± 0.02 ms) in P , SH , SV travel times and shear-wave splitting (SV – SH travel times) correlated with low-level seismicity (equivalent energy to one M 3.5 earthquake) at 70-km distance. This exceptional sensitivity (in layered basalts), at hundreds of times the conventional source dimensions, confirms the criticality of the fluid-saturated microcracks in the crust as advocated by Crampin (1998) and Crampin and Chastin (2001). Crampin et al. (2003) also reported 90° flips in shear-wave polarisations caused by critically high pore-fluid pressures on all seismically active faults (Crampin et al. 2002). These 90° flips are the cause of the large ($\pm 80\%$) scatter in shear-wave splitting time delays (Crampin et al. 2004) invariably observed above small earthquakes (Crampin et al. 1990; Booth et al. 1990).

2.12 11IWSA, 2004, St. Johns, Newfoundland, Canada: no published proceedings, but abstracts in *Geophysics*, 71, 13JF–29JF, 2006

The 11th International Workshop on Seismic Anisotropy, 11IWSA, was in St John's, New-

foundland, Canada (Table 1). This meeting had the largest attendance of any IWSA. The majority of the abstracts at 11IWSA was theoretical and appears to be modifications of well-known phenomena. At this time, it is difficult to judge their future relevance.

Theory—frequency-dependent anisotropy Vikhorev et al. (2006) suggest that different scales of inhomogeneities can cause differences between the values of anisotropy in ‘sonic waves’ (~ 20 kHz) and ‘seismic waves’ (5–10 Hz).

Theory—wave propagation in viscoelastic media Červený and Pšenčík (2006) investigate the properties of plane waves in viscoelastic anisotropic media by specifying the attenuation by complex values of the slowness vector. Note that the use of complex elastic moduli for modelling attenuation was suggested by Crampin (1981), and ANISEIS routinely uses complex moduli to model full-wave synthetic seismograms in multi-layered attenuating media (Taylor 1990). In some ways, complex moduli may be a more natural way to specify attenuating media, as one complex specification of the media avoids the complications of the complex slowness of Červený and Pšenčík (2005, 2006).

Theory—wave propagation in penny-shaped cracks—a necessary correction Grechka (2006) uses finite-element modelling to determine Thomsen parameters for media with parallel cracks and shows results that diverge from those of Hudson (1980), particularly for crack densities greater than 0.05. Grechka attributes the divergence to the first-order theory of Hudson (1980, 1981) not treating crack-to-crack interactions. This is not surprising. Hudson (1980, 1981) is a first-order theory. The second-order theory of Hudson (1986) correctly models crack-to-crack interactions. Grechka also claims that “conventional wisdom” suggests Hudson’s “first-order results are correct up to crack densities ≈ 0.1 ”. Conventional wisdom is incorrect. Microcrack densities in stable unfractured rock are limited to *less than* the fracture criticality of ~ 0.05 (Crampin 1994, 1999; the exact values may vary slightly with different matrix Poisson’s ratios

and different fluid properties Crampin 1993b). This is not a question of formulaic derivation or mathematical modelling. The physical reality of crack densities of aligned cracks higher than ~ 0.05 is that the rock is so heavily fractured that it is without shear strength and is physically unstable if the external constraints are relaxed in any way. Hudson's (1980, 1981, 1986) formula are valid for parallel cracks with crack-to-crack interactions in stable rock for crack densities less than the physical limitation of ~ 0.05 .

Observation—CO₂ injections Jenner (2006) investigates the effects of injecting high-pressure CO₂ at 10–14 MPa on the anisotropic behaviour of *P* waves in a reservoir at the Weyburn Field, Saskatchewan, Canada and finds that there is no “seismically discernible change in anisotropy”. Note that Angerer et al. (2002), who found substantial changes in shear-wave anisotropy for similar high-pressure CO₂ injections, also found negligible changes in *P*-wave signals. Theory and observations suggest that the principal effect of changing pressures is to modify crack aspect ratios (Crampin 1999). Nearly-vertically propagating *P* waves are not very sensitive to aspect ratios of nearly-vertically aligned cracks. Jenner (2006) confirms that shear-wave splitting is the most sensitive phenomenon for time-lapse monitoring of small changes in conditions.

Interpretation—GEMS, the opportunity for a global network of borehole stress-monitoring sites to forecast all damaging earthquakes worldwide Crampin et al. (2005) suggest that the sensitivity of the prototype borehole stress-monitor site (Crampin et al. 2003) indicates that a global network of (both onshore and offshore) borehole SMSs, on a 400-km grid in seismic areas and a 1,000-km grid elsewhere, would be able to forecast all damaging earthquakes ($M \geq 5$) worldwide.

2.13 12IWSA, 2006, Beijing, China: proceedings in press in *Journal of Seismology* and *Journal of Seismic Exploration*. We refer to 12IWSA expanded abstracts, 2006

The 12th International Workshop on Seismic Anisotropy, 12IWSA, was held in Beijing, China

(Table 1). This was the largest IWSA meeting, with ~ 60 national and 60 international participants, 59 oral presentations and 53 posters, with significant advances in theory, observation and interpretation. 12IWSA in Beijing stimulated a significant input from China. There were 15 oral presentations and 37 posters where the lead author was from China and many more from Chinese nationals working abroad.

There were three entertaining and informative invited Keynote Addresses: Leon Thomsen discussed the various levels of anisotropy and symmetry; Heloise B. Lynn discussed the present and future of anisotropy in exploration seismics; and Savage (2006) discussed anisotropy from the inner core to the crust with particular reference to New Zealand.

There were a large range of abstracts on theory, observation and interpretation, with far too many for detailed description. Only overall trends and a few key contributions can be noted. For example, there were 11 abstracts describing various techniques for removing the effects of VTI anisotropy, transverse isotropy with a vertical axis of symmetry. (Note that VTI now appears to be the agreed mnemonic for what is also called TIV anisotropy.) There were also seven abstracts on processing and analysing “Tilted Transverse Isotropy” or TTI a comparatively frequent occurrence in reservoirs where fine layering is not horizontal. One general problem with many of these processing techniques is that the effects are a better resolved “anisotropy” but there is little interpretation of what the anisotropy means, other than the rather vague supposition that the anisotropy refers to unresolved ‘fractures’.

Theory The most important theoretical advance at 12IWSA was probably the virtual source method (VSM) suggested by Bakulin and Calvert (2006) where a virtual shear-wave source can be constructed at downhole geophones by time reversal and deconvolution with the signals from surface sources. This effectively places sources at downhole geophone positions and eliminates the highly heterogeneous near-surface effects. VSM is believed to have great value for constructing shear-wave signals propagating through deeper layers.

Observation The Common Offset–Common Azimuthal (COCA) Cube of Gray (2006) is an effective way of displaying pre-stack seismic data that highlight azimuthal anisotropy. Just as the Common Offset Stack serves as a quality control and visual confirmation of AVO anomalies, the COCA Cube serves as a visual quality control of azimuthal anisotropy attributes.

Interpretation—attenuation Several interpretations gave direct information about the subsurface. Chapman (2006) showed that *P*-wave attenuation anisotropy in a fractured layer is a better indicator of fractures than *P*-wave travel-time anisotropy and can potentially provide new information about fracture dimensions and fluid content. Chapman hopes to develop interpretation of shear-wave anisotropy in the future.

Interpretation—kinematic and dynamic anisotropy As an example, Liu et al. (2006) showed theoretically that kinematic anisotropy (travel-time- and velocity-based measurements) and dynamic anisotropy (amplitude and attenuation measurements) may have different symmetries when there are two crack sets with different fluid contents (dry and brine-filled, say). These different symmetries have measurable effects on seismograms. This means that differences between different field measurements potentially contain valuable information about complex stress-fluid fracture systems.

Interpretation—the New Geophysics Crampin and Gao (2006b) argue that worldwide observations of shear wave indicate that the stress-aligned fluid-saturated microcracks causing the splitting are so closely spaced that they verge on fracture criticality and fracturing and are critical systems with all that that implies (Crampin 1998, 2003a, 2004, 2006; Crampin and Chastin 2001). Below criticality, critical systems of complex interactive heterogeneous phenomena act locally and have complicated reactions to local heterogeneities. However, at criticality, the whole system abruptly acts in unison. It is said to be “one of the universal miracles of nature” that at critical points whole assemblages of elements “are capable of organising

themselves into patterns of cooperative activity” (Davis 1989).

Critical systems have several crucial differences from conventional geophysics:

1. Behaviour near criticality is calculable so that seemingly complicated overall phenomena may be calculated by simplified effects neglecting detailed behaviour, hence the “unreasonable” success of APE modelling. This may be the reason why the Gassmann Theory is apparently applicable outside the expected range (Rasolofosaon 2006).
2. There is extreme (“butterfly wing’s”) sensitivity to seeming negligible changes (see Crampin et al. 2003, Rasolofosaon and Yin 1996). These two effects mean that low-level deformation can be monitored with shear-wave splitting and calculated (predicted) with APE.
3. Since the behaviour is calculable, as demonstrated by Angerer et al. (2002), then any desired level of cracking, as imaged by shear-wave splitting, can be potentially controlled. If the desired effects are not achieved, they can be adjusted by feedback (Crampin 2006). This ‘New Geophysics’ is fundamentally different from conventional sub-critical geophysics and has such potentially important applications, where shear-wave splitting is the key observable that the future of IWSA meetings seems assured.
4. One of the consequences of the New Geophysics of crack-critical systems is that universality (Davis 1989; Bruce and Wallace 1989) tends to apply over the full range of appropriate phenomena. Since a very similar range of shear-wave splitting parameters is found both in the crust and upper mantle and hydrated melt is expected, in at least some circumstances, in the upper mantle, the universal character of critical systems suggests that fluid-saturated cracks, where the in-fill is hydrated melt, may contribute to shear-wave anisotropy in the upper mantle (Crampin 2003b). Crack anisotropy in the upper mantle was originally suggested by Ando et al. (1980, 1983).

3 Discussion and conclusions

There have been some 250 papers published in the proceedings of 0IWSA to 11IWSA (part of proceedings of 12IWSA is published in this issue). This review cites papers from IWSA proceedings, which are believed to be significant for the continuing development of shear-wave splitting and seismic anisotropy, usually because they were first observations, first analyses or important leads to the continuing developments. With ten or more years' hindsight, one can recognise the significant papers. Recent papers tend to be smaller advances over broader fronts, whose significance is more difficult to assess. Consequently, this review is necessarily a rather subjective selection, and there are many excellent papers in the recent proceedings that are too numerous to cite individually, and we beg forgiveness of those authors whose papers have been omitted.

What are believed to be the most significant advances are summarised in Table 2. Each IWSA tends to be unique with the principal subject matter changing from one workshop to the next. Almost all significant developments in crustal shear-wave splitting have appeared in IWSA proceedings, and most of the few papers published elsewhere have been cited in this review. One of the principal benefits of IWSA workshops has been that all the major research workers in anisotropy have known (and trusted) each other. Consequently, workshops have been comparatively open, so that most impending developments have been immediately known to the whole research community of 'anisotropists'. This has meant that frequently, without prior consultation, several, sometimes many, people work on similar topics, so that IWSA meetings often have a dominating topic of research. This has no doubt contributed to the comparatively rapid advance, for a fundamentally new concept, which has taken anisotropy, particularly shear-wave splitting, from the first occasional observations in the crust, to observations worldwide, and an understanding of fluid-rock deformation in a critical system of closely spaced fluid-saturated microcracks in almost all rocks in the Earth's crust.

Note that the Society of Exploration Geophysicists (SEG) has promoted much of the research

in seismic anisotropy and shear-wave splitting. Its own annual meetings have had many sessions on anisotropy. It has sponsored several IWSAs and has published proceedings or abstracts of several IWSAs (Table 1). SEG awarded three Virgil Kauffman Gold Medals for research in anisotropy (Stuart Crampin in 1988; Richard M. Alford in 1990; Ilya Tsvankin in 1996). Leon Thomsen (2002) presented a Distinguished Short Course on anisotropy, and Lynn (2004) was Distinguished Lecturer on anisotropy. Leon Thomsen is currently President of SEG—an honour for Thomsen and also perhaps for anisotropy and IWSA.

Note also that Joe Dellinger of BP maintains an Internet email address where information can be posted on an Anisotropists Digest list. This greatly aids communication between 'anisotropists', who register at <anisotropists-request@freeusporng>.

IWSAs have reported observations of, particularly, shear-wave splitting, with very similar characteristics over an enormous range of frequencies from the upper mantle (0.05–0.5 Hz) in surface-wave studies, above crustal earthquakes (2–20 Hz), reflection seismology (10–80 Hz), VSPs (10–200 Hz), sonic logs (kHz) and laboratory measurements (MHz). Since such universality and calculability are characteristic of critical systems (Davis 1989; Bruce and Wallace 1989), these are comparatively direct evidence for critical systems of fluid-saturated cracks as extensive all-embracing phenomena (Crampin and Chastin 2001; Crampin et al. 2003; Crampin and Gao 2006b; Crampin 2006).

As Vikhorev et al. (2006) suggested, anisotropy covers an enormous range of lineament dimensions, and indeed, as Heffer and Bevan (1990) observed, the frequency and lengths of microcracks, cracks, fractures, faults and lineaments are self-similar over some ten orders of magnitude. This is related to the other outstanding self-similar distribution in geophysics, the Gutenberg–Richter relationship between frequency and magnitude of earthquakes, over some eight orders of magnitude. Consequently, we recognise that the deformation of fluid-saturated cracks is the physical reality underlying the Gutenberg–Richter relationship. This had been suspected in the past (Turcotte 1992), but this is now a direct correlation with other phenomena.

Table 2 Summary of significant advances in seismic anisotropy at IWSA meetings

ID	Year of IWSA	Year of proceedings	Significant advances
0IWSA–3IWSA: principally earthquake seismology			
0IWSA	1975	1977	<p>Synthetic seismograms in anisotropic layered models (Keith and Crampin 1977a, b, c)</p> <p>Anisotropy observed in upper mantle from <i>Pn</i>-wave velocity anisotropy beneath both oceans (Hess 1964) and continents (Bamford 1977)</p> <p>Anisotropy observed (Crampin 1966) and calculated in higher-mode surface-wave dispersion (Crampin and King 1977)</p> <p>Shear-wave splitting first named (as <i>S-wave splitting</i>) and displayed in synthetic seismograms (Keith and Crampin 1977c)</p>
1IWSA	1982	1984	<p>Observations of <i>P</i>-wave velocity anisotropy in field and laboratory (Artemieva and Chesnokov 1991; Brodov et al. 1991; Brown et al. 1991; de Parscau 1991; MacBeth 1991a, b)</p> <p>Elastic moduli calculated for aligned cracks (Crampin 1978, 1984b)</p> <p>Importance of shear-wave window above small earthquakes recognised (Evans 1984)</p> <p>Aligned microcracks (EDA cracks) suggested in most crustal rocks (Crampin et al. 1984b)</p> <p>General behaviour of wave propagation in cracked media reviewed (Crampin 1981)</p> <p>Shear-wave splitting displayed in synthetic seismograms through cracked media (Crampin 1978)</p> <p>Shear-wave splitting observed above earthquakes in both crust (Crampin et al. 1980) and upper mantle (Ando et al. 1980)</p>
2IWSA	1986	1987	<p>Observations of <i>P</i>-wave anisotropy in crust (Leary et al. 1987; Li et al. 1987; Galperina and Galperin 1987)</p> <p>Speculation on how monitoring shear-wave splitting could predict earthquakes (Chen et al. 1987)</p> <p>Determination of the Thomsen parameters for seismic exploration (Thomsen 1986)</p> <p>Double-contour integration for 3D full-wave propagation in anisotropic substrates (Taylor 1987)</p>
3IWSA	1988	1990	<p>Observations of anisotropy in laboratory studies in both field and laboratory (Kaneshima 1990; Li et al. 1990; Savage et al. 1990; Shih and Meyer 1990)</p> <p>First attempts at automatic measurement of shear-wave splitting (Shih and Meyer 1990)</p> <p>First correlation of microcracks with fractures and directions of preferred fluid flow (Queen and Rizer 1990)</p> <p>Shear-wave splitting time delays seen to increase before two earthquakes (Booth et al. 1990; Crampin et al. 1990)</p> <p>ANISEIS programme developed for applying double-contour integration for full-wave synthetic seismograms through anisotropic layered models (Taylor 1990)</p>
4IWSA–12IWSA: Principally exploration seismology			
4IWSA	1990	1991	<p>First 10 years of shear-wave splitting reviewed (Crampin and Lovell 1991)</p> <p>Review of theory of wave propagation in aligned cracks (Hudson 1991)</p> <p>Shear-wave singularities: identified in cracked VTI anisotropy (Wild and Crampin 1991)</p> <p>Synthetic seismograms of shear-wave singularities show 90° flips and amplitude anomalies (Crampin 1991)</p>

Table 2 (continued)

ID	Year of IWSA	Year of proceedings	Significant advances
5IWSA	1992	1992	<p>Effects of shear-wave singularities observed and modelled in VSP in Paris Basin (Bush and Crampin 1991)</p> <p>Many papers on observations of <i>P</i>-wave velocity anisotropy and shear-wave splitting</p> <p>Shear-wave splitting used to guide horizontal drilling into heavily fractured rock (Mueller 1991)</p> <p>Linear-transform technique developed for processing anisotropic record sections (Li and Crampin 1991a, b)</p> <p>Micro-cracks rather than macro-fractures suggested as source of shear-wave splitting (Crampin 1993a)</p> <p>Effects of aligned cracks reviewed for a range of parameters (Crampin 1993b)</p> <p>Many techniques for calculating synthetic seismograms presented (Guest and Kendall 1993; Igel et al. 1993; Leary et al. 1993; Wild et al. 1993; Yao and Xiong 1993a, b; Zhang et al. 1993)</p> <p>Correlation of degree of shear-wave splitting and hydrocarbon production (Li et al. 1993) and (Liu et al. 1993a)</p> <p>Anisotropic cusps observed in walkaway VSP and modelled with synthetic seismograms (Slater et al. 1993)</p>
6IWSA	1994	1996	<p>Historical background of Lord Kelvin (Helbig 1996)</p> <p>A number of case studies were presented (Horne et al. 1996; Kebaili et al. 1996; Stawicki and Lynn 1996; Mjelde 1996)</p> <p>Non-linear response of shear-wave splitting more sensitive to near-negligible changes than conventional wave speeds and moduli (Rasolofosaon and Yin 1996)</p>
7IWSA	1996	2001	<p>Frequency-dependent anisotropy: systematic differences between intrinsic anisotropy of shales in cores at ultrasonic frequencies and walkaway VSPs (Hornby 2001)</p>
8IWSA	1998	1998	<p>Various theoretical and experimental papers</p> <p>Papers on theoretical wave propagation (Helbig 1998; Chesnokov et al. 1998) and processing techniques (Raymer and Kendall 1998; Li 1998)</p> <p>Closely spaced cracks form critical systems first proposed (Crampin 1998)</p>
9IWSA	2000	2001	<p>Many papers on amplitude variation with offset (Ikelle and Amundsen 2001; Liu et al. 2001) and <i>P</i>-to-<i>S</i>-converted <i>C</i> waves (Granger et al. 2001)</p> <p>Many papers on processing TIV media</p> <p>Time, magnitude and fault of <i>M</i> 5 in SW Iceland stress forecast (Crampin et al. 1999)</p> <p>Supporting evidence for closely spaced cracks as critical systems (Crampin and Chastin 2001)</p>
10IWSA	2002	2003	<p>Theory and observation of frequency-dependent anisotropy (Chapman et al. 2003; Gurevich 2003; Li and Yuan 2003; Liu et al. 2003)</p> <p>Evidence for critical systems from sensitivity of prototype borehole Stress-Monitoring Site (Crampin et al. 2003)</p>
11IWSA	2004	2006	<p>The sensitivity of the prototype SMS suggests that GEMS, a global network of SMSs, presents the opportunity to forecast all damaging earthquakes worldwide</p>
12IWSA	2006	Abstr 2006	<p>Many papers on processing and analysing to remove the effects of VTI anisotropy</p> <p>Many papers on processing and analysing tilted VTI-anisotropy</p>

Table 2 (continued)

ID	Year of IWSA	Year of proceedings	Significant advances
			Virtual source method developed by Bakulin and Calvert (2006)
			Common offset–common azimuth cube for displaying azimuthal data (Gray 2006)
			<i>P</i> -wave attenuation anisotropy is a better indicator of fractures than <i>P</i> -wave travel-time anisotropy (Chapman 2006)
			In some circumstances, travel-time measurements may contain different anisotropic information from attenuation measurements (Liu et al. 2006)
			Crampin and Gao (2006b) suggest a New Geophysics where in most in situ rocks in the crust fluid-saturated microcracks are so closely spaced that they verge on fracture criticality and fracturing and are critical systems; evidence suggests the universality, calculability, predictability and great sensitivity expected of critical systems; New Geophysics would have massive implications for monitoring and measuring geophysics including hydrocarbon exploration and recovery

In the earth, there is a continuum of all scales of ordered heterogeneities, from 0.01 mm of microcracks to 100 s of kilometre plate boundaries in critical systems. The implications of critical systems are that low-level deformation before fracturing takes place: can be monitored with shear-wave splitting; future behaviour *calculated* by APE; even *predicted* if changing conditions are known, as in Angerer et al. (2002); and in some circumstances future behaviour potentially *controlled* by feedback (Crampin 2006). If correct, and the papers at this series of IWSAs reported in this review suggests overwhelming evidence that it is correct, this is such a fundamental advance in understanding fluid-rock (pre-fracturing) deformation in solid Earth geoscience that it is a New Geophysics.

A large proportion of papers by exploration seismologists presented at IWSAs have been about observing, processing and calculating the effects of anisotropy and shear-wave splitting, with much of the emphasis on how anisotropy effects conventional isotropic processing: how conventional processing for structure and stratigraphy is affected by anisotropy. The New Geophysics suggests that the implications of shear-wave anisotropy for fluid-rock interaction are likely to be crucially important for hydrocarbon production in the future, as well as almost all other solid earth geoscience, by their effects on fluid-rock deformation.

Acknowledgements We thank Klaus Helbig, Sheila Peacock, Ivan Pšenèk, Michael Slawinski and Leon Thomsen for information about IWSA workshops that neither of the authors attended. We particularly thank Heloise Lynn for her valuable comments. We also thank the Guest Editor Martha Savage and Yong-Gong Li, whose meticulous comments greatly improved the manuscript. YG was supported by NSFC Project: 40674021.

Appendix: how IWSAs began

Evgeni Chesnokov and Stuart Crampin, in a casual conversation over a vodka-fuelled dinner during a 3-month collaboration in Moscow in the winter of 1979–1980, speculated on the need for an international workshop on seismic anisotropy to discuss and publicise our ideas. In those days, anisotropy needed encouragement. Several eminent geophysicists (who shall be nameless) poured scorn on the very idea of widespread stress-aligned anisotropy in the Earth. Evgeni and I needed to gather support to show we were not wasting our time. One of us (EC) happened to mention this idle speculation to Academician V.A. Magnitsky, who immediately lent his authority to the idea, and the workshop was on its way. Organised by Chesnokov, Crampin and Magnitsky, the first IWSA was in 1982 in the beautiful ancient city of Suzdal, USSR.

The workshop happened to be in the right place at the right time. Shear-wave splitting had just

been reported both in the crust and upper mantle, and there was more general interest in anisotropy in USSR than elsewhere at that time.

1IWSA set the pattern for future IWSAs. IWSAs tend to be comparatively small (less than 100 participants) and single session, tend to be in interesting and memorable places and tend, after 2IWSA, to alternate across the Atlantic (and now the Pacific). IWSAs have been from Monday to Friday, with Wednesday set aside for the usually cultural excursions. These excursions are important since many of the most productive ideas, as with the original idea for IWSA, come from casual exchanges away from the seminar room.

The first IWSA, in 1982, was a resounding success and seemed worth continuing, so the next meeting was arranged 4 years hence in Moscow, 1986. This gap set another pattern. After 2IWA, biennial meetings were established. IWSAs are not repeated unless there is sufficient enthusiasm amongst participants at one meeting to organise the next in 2 years time. Thus, IWSA meetings have built-in quality assurance. When enthusiasm flags, IWSAs will immediately fold, which will be a right and proper finale to what has been a productive and rewarding series of meetings.

Note added in proof Papers from Proceedings of IWSA workshops are indicated by asterisks.

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