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Group and phase velocities in the free-surface visco-potential flow: new kind of boundary layer induced instability

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Abstract

The present work is entirely devoted to thorough analysis of the dispersion relation of the novel visco-potential formulation. Namely, in this study we relax all assumptions of the weak dependence of the wave frequency on time. As a result, we have to deal with complex integro-differential equations that describe transient behaviour of the phase and group velocities. Using numerical computations, we show several snapshots of these important quantities at different times as functions of the wave number. Good qualitative agreement with previous study [Dut08] is obtained. Thus, we validate in some sense approximations made anteriorly. There is an unexpected conclusion of this study. According to our computations, the bottom boundary layer creates disintegrating modes in the group velocity. In the same time, the imaginary part of the phase velocity remains negative for all times. This result can be interpreted as a new kind of instability which is induced by the bottom boundary layer effect.

Key words: free-surface flows, viscosity, dissipation, visco-potential flow, bottom boundary layer

1 Introduction

The classical potential free-surface flow theory is known to be a good and relatively inexpensive model of water waves (especially in comparison with free-surface Navier-Stokes equations formulation [HW65, SZ99, WY07]). However, there are some physical situations where viscous effects cannot be neglected.

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The necessity of including some dissipation into various water waves models was pointed out explicitly in a number of experimental studies [ZG71, Wu81, BPS81]. For example, in the “Résumé” section of [BPS81] one finds:

...it was found that the inclusion of a dissipative term was much more important than the inclusion of the nonlinear term, although the inclusion of the nonlinear term was undoubtedly beneficial in describing the observations...

One can find many other evidences in the literature which point out the importance of viscous effects.

Historically, the researchers tried first to include the dissipative effects into various long wave models such as Burgers, Korteweg-de Vries and Boussinesq equations. There is a vast literature on this subject [Keu48, OS70, KM75, Mil76, Mat76, Kha87, Sug91, Kha97, DD07a, Dut07].

In order to include some dissipation into the framework of free-surface potential flows, we developed the so-called visco-potential formulation [LO04, DDZ08, DD07b, Dut07, Dut08]. The kinematic viscosity (or eddy viscosity more precisely) appears through local dissipative terms in kinematic and dynamic free-surface boundary conditions. The main peculiarity consists in modifying the bottom kinematic condition due to the presence of the boundary layer. Mathematically, this correction procedure leads to a nonlocal in time term. Note, that from fractional calculus point of view this nonlocal term is also a half-order integral. The physical relevance of the visco-potential formulation was shown in [LSVO06]. They compared model predictions with experiments on the damping and shoaling of solitary waves. It was shown that the viscous damping due to the bottom boundary layer is well represented by this theory.

To complete our literature review, recall that there is also an alternative approach to potential flows of viscous fluids developed by Daniel Joseph and his collaborators [JL94, FJ02, JW04, Jos06].

The main goal of the present article is twofold. On one hand, we refine our previous dispersion relation analysis. In particular, we do not make any additional assumption of the weak dependence of the wave frequency on time. Consequently, we have to deal with complex integro-differential equations. On the other hand, we show that bottom boundary layer can induce a disintegrating instability of the wave packets. This result is new to author’s knowledge.

The present article is organized as follows. In Section 2 we describe the governing equations and perform classical dispersion relation analysis. Then, we derive an equation for the group velocity in Section 2.1. Section 3 contains several numerical results and their discussion. Finally, this article is ended by

outlining main conclusions of the study in Section 4.

2 Mathematical formulation and dispersion relation analysis

Consider the 3D fluid domain bounded above by the free-surface $z = \eta(\vec{x}, t)$, $\vec{x} = (x, y)$ and below by the rigid boundary $z = -h(\vec{x})$. A Cartesian coordinate system with the z -axis pointing vertically upwards and the xOy -plane coinciding with the still-water level. The flow is assumed incompressible and the fluid viscous with the kinematic viscosity ν . The governing equations of the free-surface visco-potential flow have the following form [DD07b, Dut07, Dut08]:

$$\Delta\phi = 0, \quad (\vec{x}, z) \in \Omega = \mathbb{R}^2 \times [-h, \eta], \quad (1)$$

$$\eta_t + \nabla\eta \cdot \nabla\phi = \phi_z + 2\nu\Delta\eta, \quad z = \eta, \quad (2)$$

$$\phi_t + \frac{1}{2} |\nabla\phi|^2 + g\eta + 2\nu\phi_{zz} = 0, \quad z = \eta, \quad (3)$$

$$\phi_z = -\sqrt{\frac{\nu}{\pi}} \int_0^t \frac{\phi_{zz}(\vec{x}, z = -h, \tau)}{\sqrt{t - \tau}} d\tau, \quad z = -h. \quad (4)$$

We can derive corresponding long wave models from the visco-potential formulation (1) – (4). The derivation can be found in [LO04, DD07b, Dut07].

In order to perform the dispersion relation analysis, we have to linearize equations (1) – (4) over the flat bottom $z = -h$. This procedure is classical and we do not detail it here. Then, we look for the following periodic plane wave solutions:

$$\phi(\vec{x}, z, t) = \varphi(z)e^{i(\vec{k}\cdot\vec{x} - \omega(t; \vec{k})t)}, \quad \eta(\vec{x}, t) = \eta_0 e^{i(\vec{k}\cdot\vec{x} - \omega(t; \vec{k})t)}, \quad (5)$$

where \vec{k} is the wavenumber and $\omega(t; \vec{k})$ is the wave frequency.

Remark 1 *It is important to assume from the beginning of the derivation that the wave frequency ω explicitly depends on the time t . Consequently, we get some additional important terms of the form $it\omega'(t; \vec{k})$ (see equation (6)) which were neglected in our previous study [Dut08].*

We plug the special form of solutions (5) into the linearized version of the governing equations (1) – (4). After performing some simple computations (details can be found in [Dut07, Dut08]), we come to the following necessary condition of periodic solution (5) existence:

$$\Omega^2 + gk \tanh(kh) - kF(\omega, t)(\Omega^2 \tanh(kh) + gk) = 0, \quad (6)$$

where we introduced several notations $k := |\vec{k}|$, $\Omega := it\partial_t\omega + i\omega - 2\nu k^2$ and $F(\omega)$ is inherited from the nonlocal term:

$$F(\omega, t) := \sqrt{\frac{\nu}{\pi}} \int_0^t \frac{e^{i\omega(\tau)(t-\tau)}}{\sqrt{t-\tau}} d\tau.$$

Once the wave frequency ω is computed from equation (6), the phase speed can be immediately deduced by its definition:

$$c_p(t; k) := \frac{\omega(t; k)}{k}.$$

Remark 2 *The classical dispersion relation for water waves $\omega^2 = gk \tanh(kh)$ can be immediately recovered from (6) if we replace Ω by $i\omega$ and $F(\omega)$ by 0.*

2.1 Group velocity

The group velocity is defined as follows:

$$c_g(t; k) := \frac{\partial\omega(t; k)}{\partial k}.$$

There are several physical interpretations of this quantity. One can see it as the wave energy propagation speed. Another interpretation consists in viewing it as the wavetrain amplitude variation speed. Anyhow, the group speed c_g plays an important rôle in the description of the wavetrains and wavepackets. The evolution equation of this quantity can be obtained by simple differentiation of (6) with respect to the wavenumber modulus k :

$$\begin{aligned} 2\Omega\Gamma + g \tanh(kh) + gkh(1 - \tanh^2(kh)) - (F + k\partial_k F)(\Omega^2 \tanh(kh) + gk) \\ - kF(2\Omega\Gamma \tanh(kh) + \Omega^2(1 - \tanh^2(kh))h + g) = 0, \end{aligned} \quad (7)$$

where Ω was defined above and $\Gamma := it\partial_t c_g + ic_g - 4\nu k$.

Integro-differential equations (6) and (7) are too complex for any mathematical analysis. Therefore, we use numerical methods in order to explore some properties of the solutions.

3 Numerical results and discussion

In order to have an insight into the transient behaviour of the phase and group velocities, we have to solve numerically equations (6) and (7) correspondingly.

<i>parameter</i>	<i>definition</i>	<i>value</i>
ν	eddy viscosity	$10^{-3} \frac{m^2}{s}$
g	gravity acceleration	$9.8 \frac{m}{s^2}$
h	water depth	1 m
Δt	time step	0.05 s

Table 1

Values of the parameters used in the phase and group velocities numerical computations.

Hence, in this study we discretize all local terms with an implicit scheme, while the integral term $F(\omega, t)$ is computed in explicit way for the sake of simplicity. Resulting algebraic equations are solved analytically. Initial conditions were chosen according to [DDZ08]:

$$c_p|_{t=0} = \sqrt{\frac{g}{k} \tanh(kh)} - 2i\nu k, \quad c_g|_{t=0} = \left(\frac{1}{2} + \frac{kh}{\sinh(2kh)} \right) \text{Re } c_p|_{t=0} - 4i\nu k.$$

The values of all parameters used in numerical computations can be found in Table 1.

In this section upper and lower images always refer to the real and imaginary parts respectively unless special indications are given. Presented here results are normalized by \sqrt{gh} . Thus, all real parts at the infinitely long wave limit $kh \rightarrow 0$ take the unitary value.

On Figures 1 and 2 we show two snapshots of the phase velocity at different times. The comparison with our weakly dynamic approach [Dut08] gives an excellent agreement for the real part of the phase velocity $c_p(t; k)$. There is a little discrepancy in the behaviour of the imaginary part in the region of long waves, while the dissipation of short waves is perfectly described. Clearly, our previous analysis overestimates the dissipation rate for long and, perhaps, some intermediate wavelenghtes. However, the simplified approach gives a reasonable accuracy for applications.

This comparison validates the simplifications made in our previous study [Dut08]. We refer to [Dut08] for the discussion and physical interpretation of the phase velocity behaviour.

Group velocity snapshots are presented on Figures 3 – 7. The real part of the group velocity c_g is shown only on Figure 7 since its evolution is slow as in the phase velocity case. The whole animation of the group velocity can be downloaded at [GRP].

The imaginary part of the group velocity remains negative until about 1 s (see Figure 4). Then, it was very unexpected for the author to observe the forma-

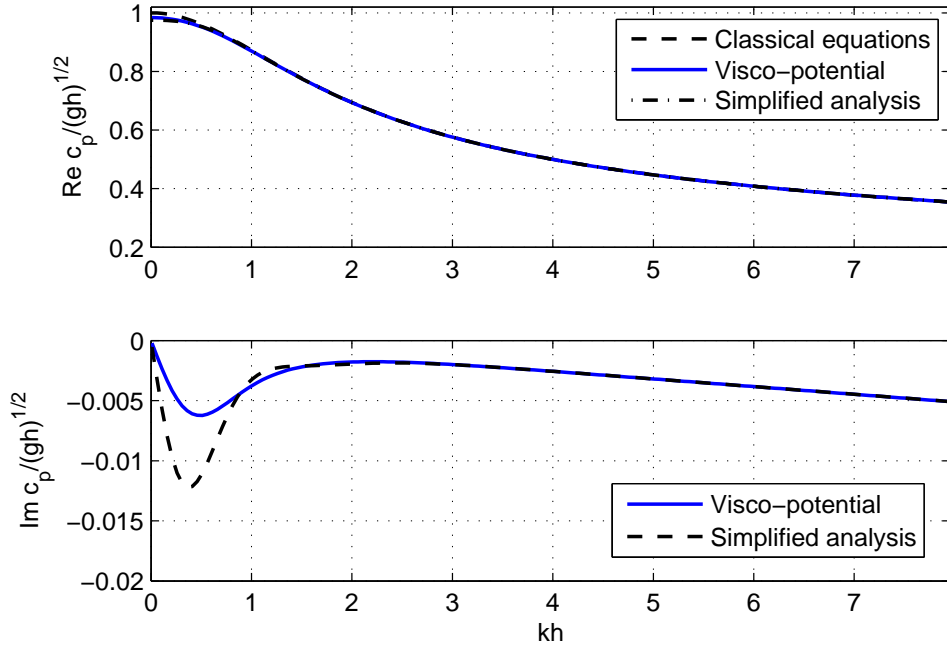


Figure 1. Real and imaginary parts of the phase velocity at $t = 2$ s.

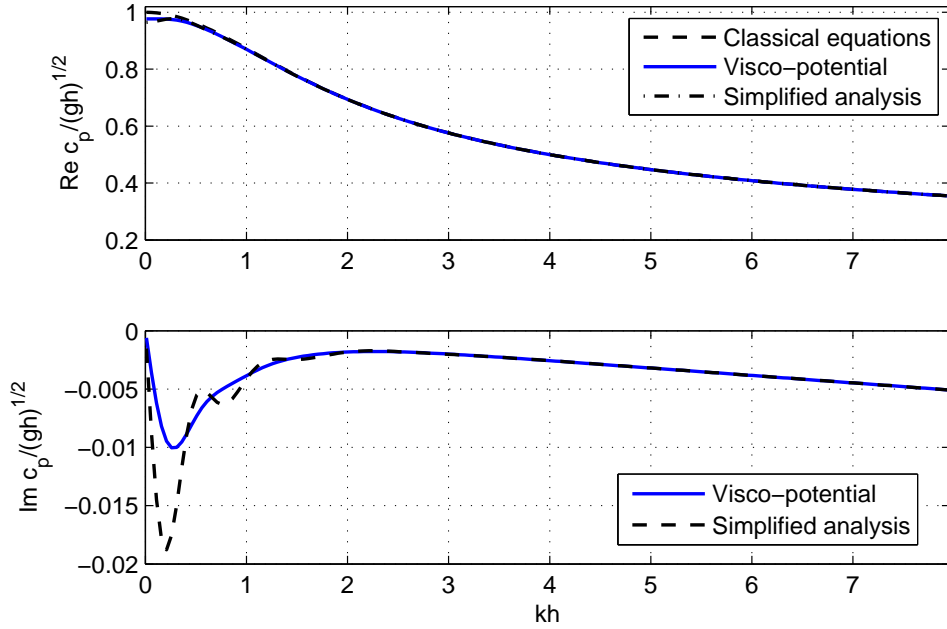


Figure 2. Real and imaginary parts of the phase velocity at $t = 4$ s.

tion of modes with positive imaginary part (see Figure 5). These destabilizing modes may be responsible of wavepackets disintegration. We refer to the next section for the discussion of this interesting result.

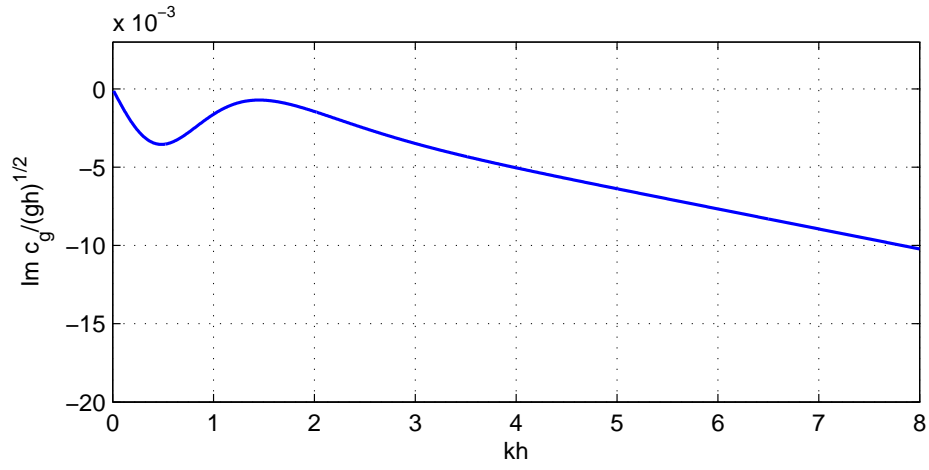


Figure 3. Imaginary part of the group velocity at $t = 1$ s.

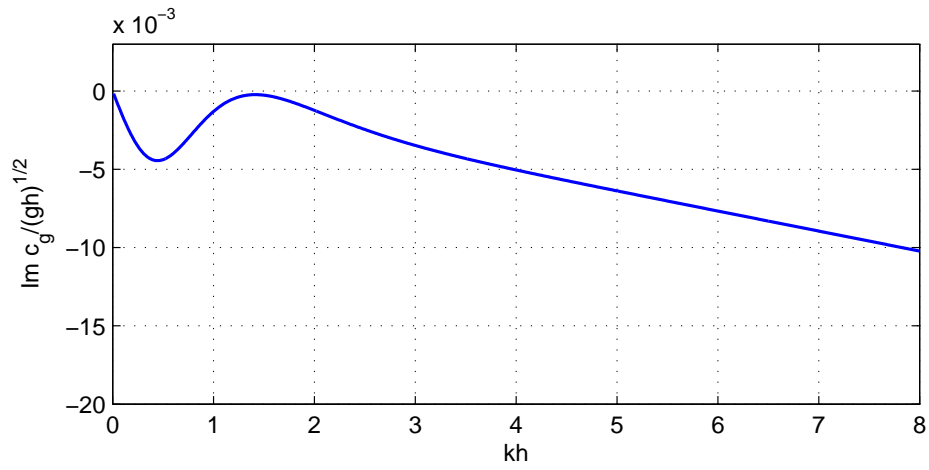


Figure 4. Imaginary part of the group velocity at $t = 1$ s.

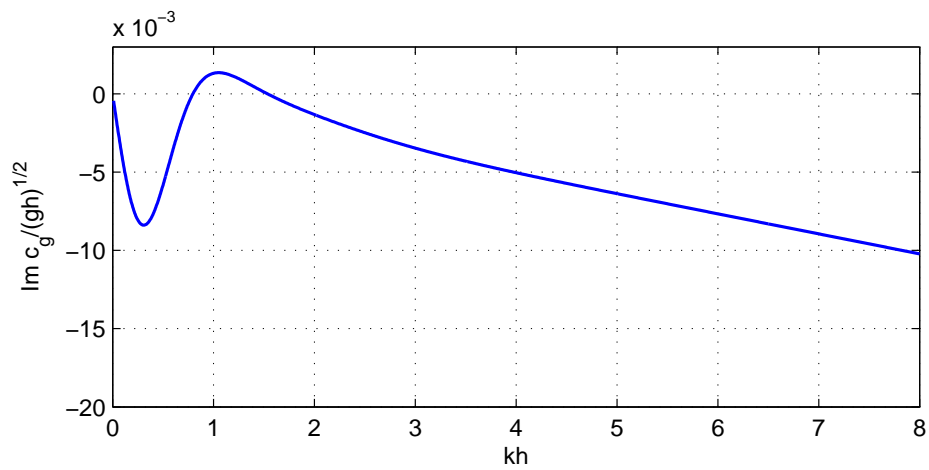


Figure 5. Imaginary part of the group velocity at $t = 2$ s.

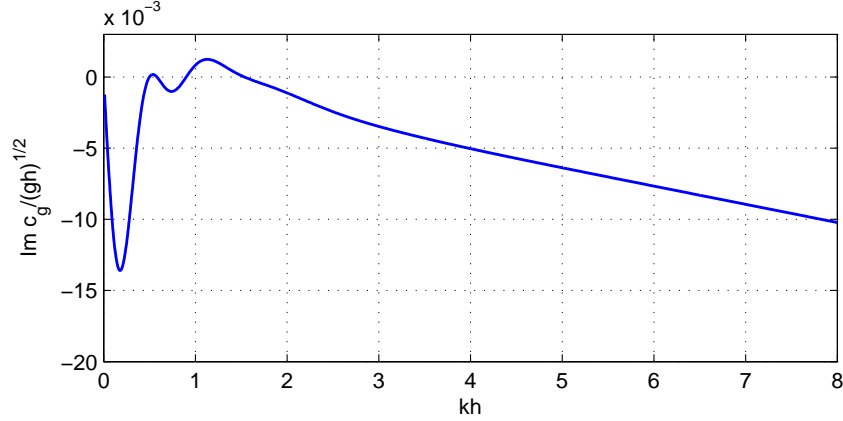


Figure 6. Imaginary part of the group velocity at $t = 4$ s.

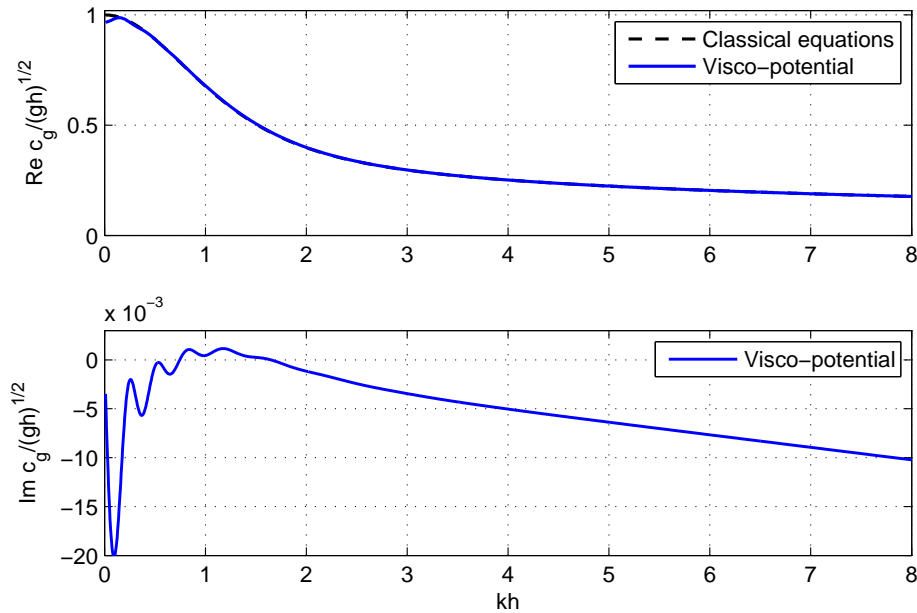


Figure 7. Real and imaginary parts of the group velocity at $t = 8$ s.

4 Conclusions

In this work we performed an analysis of linear dispersion relation for visco-potential flow formulation. Several numerical snapshots of the phase and group velocities were presented. The main particularity in this study is that we make no assumption of the weak dependence on time. Therefore, we come up with additional terms of the form $it\omega'(t; k)$ which complicate analysis and numerical computations. We can conclude that we obtain a good qualitative agreement with preceding results [Dut08]. Thus, we validate in some sense approximations made anteriorly. However, our computations reveal less oscillatory behaviour in the imaginary part of the phase velocity for intermediate

wavelengths.

Furthermore, we studied the evolution of the group velocity with time for the visco-potential formulation. Physically, this quantity represents the energy propagation speed. There exists also another interpretation. The group velocity $c_g(t; k)$ can be seen as the wavetrain amplitude variation speed. Numerical simulations show the appearance of modes with positive imaginary part for $\frac{1}{2} < kh < \frac{3}{2}$ in the group velocity. These modes may destabilize wave packets. Physically, this effect comes from the bottom boundary layer. Previous numerical study [WLY06] did not reveal this instability since the authors did not take into account for the boundary layer effect. Hence, they included only a local dissipative term which has a stabilizing effect.

It is well known that Stokes wavepacket can be disintegrated by the Benjamin-Feir instability [Ben67, BF67]. This instability comes from the resonant quartet wave interaction. The classical result of T. Brooke Benjamin [Ben67], refined later by R.S. Johnson [Joh77], states that no instability is present if $kh < \sigma_0$ and the modes inside a certain band-width grow without bound if $kh \geq \sigma_0$. The critical value of σ_0 is estimated about 1.363. There is a certain surprising interplay with our results that we would like to point out here. Our theory gives less sharper results. That is why we prefer to speak about the instability zone defined as $\mathbb{U}(t) = \{kh \in \mathbb{R}^+ : \text{Im } c_g(t; k) > 0\}$ in kh -space. According to our computations, the topological structure of the set $\mathbb{U}(t)$ is a union of finite number of intervals (see the animation [GRP]). It is curious that the critical value σ_0 lies inside this zone (see Figure 5).

Recently it was shown that the Benjamin-Feir instability can be enhanced by dissipation [BD07]. The peculiarity lies in the form of dissipative terms that we put into the model. Our computations show the stabilizing effect of local dissipative terms [SHC+05, WETG01, WLY06] and the destabilizing effect of our nonlocal dissipation.

The experiments should confirm our theoretical result, provided that a generated wavetrain have sufficiently big wavelength to interact with the bottom boundary layer.

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