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J. M. Sánchez-Pérez, C. Bouey, Stéphane Sauvage, S. Teissier, I. Antiguiedad, et al.. A standardised method for measuring in situ denitrification in shallow aquifers: numerical validation and measurements in riparian wetlands. *Hydrology and Earth System Sciences Discussions*, 2003, 7 (1), pp.87-96. hal-00330844

HAL Id: hal-00330844

<https://hal.science/hal-00330844>

Submitted on 18 Jun 2008

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A standardised method for measuring *in situ* denitrification in shallow aquifers: numerical validation and measurements in riparian wetlands

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Abstract

A tracer test to examine *in situ* denitrification in shallow groundwater by a piezometer with a packer system used bromide as a tracer of dilution and acetylene (10%) to block the denitrification process at the nitrous oxide stage. During the test, dissolved oxygen, nitrate (NO₃⁻), bromide (Br⁻), nitrous oxide (N₂O) and dissolved organic carbon (DOC) were measured. To calibrate the experimental method, comparison with numerical simulations of the groundwater transfer were carried out, taking into account the environmental characteristics. The method was tested by measurements undertaken in different environmental conditions (geology, land use and hydrology) in two riparian wetlands. Denitrification rates measured by this method ranged from 5.7 10⁻⁶ g N-NO₃⁻ L⁻¹ h⁻¹ to 1.97 10⁻³ g N-NO₃⁻ L⁻¹ h⁻¹. The method is applicable in shallow aquifers with a permeability from 10⁻² to 10⁻⁴ m s⁻¹.

Keywords : denitrification, shallow aquifer, groundwater modelling, wetlands, nitrate-nitrogen, packer system

Introduction

Riparian zones can play an important role in reducing nitrate input from groundwater to river water and, thus, in regulating the flux of nitrate of agricultural origin. The capacity of these interface zones to retain and remove diffuse contamination of agricultural origin in the transversal axis (river - aquifer) has been the subject of numerous studies (Peterjohn and Correll, 1984; Pinay and Décamps, 1988; Gregory *et al.*, 1991; Correll *et al.*, 1992; Sánchez-Pérez *et al.*, 1991a, b; Sánchez-Pérez and Trémolières, 2003, 1997; Cooper, 1990; Jordan *et al.*, 1993; Cey *et al.*, 1999).

One of the major processes associated with the removal of nitrate was denitrification (Jacobs and Gilliam, 1985; Cooper, 1990; Lowrance *et al.*, 1995; Burt *et al.*, 1999). The denitrification rate has been quantified widely in different types of media or aquifer lithology and climate (Cooper, 1990; Pinay *et al.*, 1995).

While denitrification has been studied and quantified within soils, few studies have measured this process within an alluvial aquifer. The role of riparian vegetation in decreasing the nitrate concentrations from aquifers has been demonstrated predominantly by a decreasing nitrate concentration gradient from groundwater to rivers (Hill and

Shackleton, 1989; Haycock and Pinay, 1993). The role of the denitrification process within this decreasing nitrate gradient has been measured by running *in situ* denitrification rate analyses in the surface layer of the riparian soils (Pinay *et al.*, 1993). However, if the objective is to address the reduction of nitrate coming from groundwater by riparian zones, one must consider that the greatest flux of nitrate (i.e., that passing through the saturated zone of the aquifer) is often disconnected from the surface soil layer. Therefore, it is essential to be able to measure denitrification within the saturated zone in order to compare and quantify the role of microbial denitrification and root uptake processes quantitatively.

Denitrification refers to the dissimilatory reduction of one or both of the ionic nitrogen oxides (nitrates NO₃⁻, nitrites NO₂⁻) to the gaseous oxides (nitric oxide NO, nitrous oxide N₂O), which may themselves be further reduced to dinitrogen N₂ by nitrous oxide reductase. Denitrification occurs in anaerobic media, when nitrate and organic matter can be found by bacteria (Payne, 1973; Knowles, 1982).

The four reductases involved in denitrification are susceptible to inhibition by a variety of compounds. The N₂O reductase, located in the bacterial membrane is inhibited

by acetylene (C_2H_2) which is a non-competitive inhibition (inhibition independent of substrate concentration). The presence of acetylene in the medium causes the stoichiometric accumulation of N_2O as the terminal product of the reduction of other oxides, instead of an accumulation of dinitrogen (Knowles, 1982); the latter is difficult to measure as the reaction is blocked by the acetylene at the N_2O phase. The levels of N_2O can then be measured accurately.

The acetylene blocking method has been heavily criticised (Knowles, 1990), mainly because of a lack of N_2O reductases inhibition (Seitzinger *et al.*, 1993). However, this method provides a valid basis for measurements of denitrification (Balderston *et al.*, 1976 ; Yoshinari and Knowles, 1976 ; Yeomans and Beauchamp, 1978; Knowles, 1981). Therefore, it is deemed appropriate for *in situ* denitrification measurements in groundwater.

Denitrification was measured within the saturated zone of a porous aquifer by acetylene blocking method. The saturated zone was sampled by a piezometer, which became both the injecting point for acetylene and the sampling site for measuring the rate of denitrification. The objective was to standardise a method to measure denitrification *in situ* within shallow alluvial riparian groundwaters.

Principles of the method

Since the same piezometer is used to inject acetylene and to collect water sampled for denitrification analysis, it is necessary to check how the injected products are dispersed and diluted around the piezometer. A numerical model supports these experiments to compute groundwater motion and, consequently, to check if what has been injected is really sampled. This method was tested in the Garonne floodplain (France) and in the Vitoria-Gasteiz aquifer (Spain). In the Garonne floodplain, six piezometers were sampled; two were in a riparian wetland, two in a poplar tree plantation and two in agricultural land. In the Vitoria-Gasteiz Aquifer, two piezometers were sampled at two depths, in a marny substrate and in a silty soil.

Material and methods

WATER SAMPLING

Nineteen litres of groundwater were sampled with a peristaltic pump in a piezometer using a packer system (UVITEC, Austria) after removal of ten times the volume of water from the piezometer. This system enables water to be pumped from the aquifer rather than just from the piezometer at the same depth during the experiment.

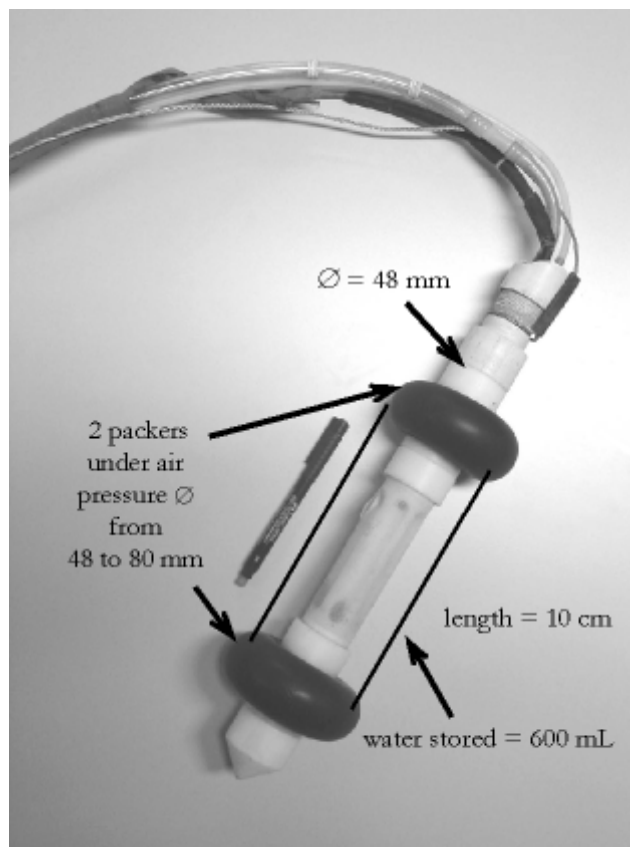


Fig. 1. The packer system.

The packer system (Fig. 1), consists of a pump within a 48 mm in diameter tube, 10 cm long, which can store 600 ml of water; above and below this tube are two packers that can be inflated from 48 to 80 mm. Water is pumped at a pre-determined depth, and stored in a 25 litre container. One litre of a solution containing bromide (as NaBr, final concentration 13.4 mg l^{-1} of Br^-) and acetylene (10%) is added to the pumped water sample. Bromide is used as a dilution tracer and acetylene is used to block the denitrification process at the N_2O step. Bromide and acetylene were added to the sampled water in a sampling bag (container). The C_2H_2 was obtained by an exothermic reaction between water (20 ml) and calcium carbide, using speleological equipment; then it is injected by bubbling into the water until the C_2H_2 has displaced the air. The C_2H_2 injection in the sampling bag is considered complete when the C_2H_2 goes out of the water sample bottle through a thin tube. It is considered that N_2O is diluted in the aquifer in the same way as bromide. The injections are done with all the precautions necessary to prevent oxygen from entering the container.

After mixing, the 20 litre solution is re-injected with the peristaltic pump into the piezometer, at the same depth at

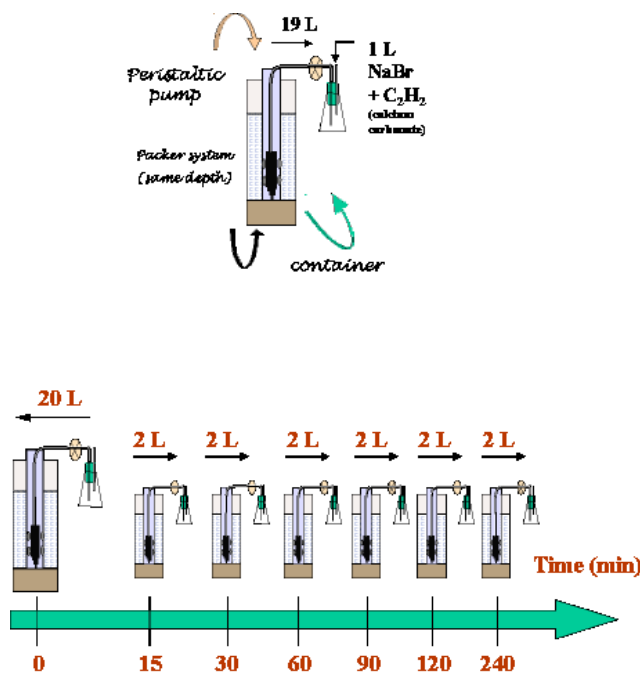


Fig. 2. Application of the methodology in the field.

which the water had previously been sampled. t_0 is the time at which the solution has been injected entirely. Two litres of water, assumed to be mixed with acetylene and bromide, were sampled successively in the same piezometer 0, 15, 30, 60, 90, 120 and 240 minutes after injection (Fig. 2).

WATER ANALYSIS

Water samples were collected in glass bottles and filtered through glass fibre filters (Whatman GFF). Nitrate (NO_3^- -N) and bromide (Br) were analysed by ionic chromatography using a DIONEX system. Water samples collected for dissolved organic carbon (DOC) were filtered using precombusted GFF filters (450°C for 4 h) and analysed using a platinum catalyst at 650°C (Shimadzu, Model TOC 5000). A multi-parameter probe (YSI 6920) provided physico-chemical data at each time (temperature, dissolved oxygen, conductivity, pH, redox potential).

A small sample (3 ml) was also injected with a plastic syringe in a vacuum tube (Venoject, Terumo; evacuated blood collection tubes, diameter = 13 mm, length = 75 mm, vacuum volume 5 ml), with formaldehyde (0.5%) to prevent bacteria from continuing denitrification. N_2O was further analysed by a gas chromatograph equipped with a ^{63}Ni electron capture detector. The carrier gas was a mixture of argon (95%) and methane (5%) under 1 bar. The inert gases were separated on a Poparak Q column at 80°C, the injector temperature was 130°C and the detector temperature 280°C.

The calculation takes into account the N_2O mass in the sampled water (using N_2O solubility, Weiss and Price, 1980) plus the N_2O mass in the gaseous phase minus the N_2O mass introduced during equilibration.

CALCULATIONS OF THE DENITRIFICATION RATE

The N_2O concentrations measured in the experiment were corrected for dilution by dispersion and advection processes, using a dilution factor determined from the bromide concentrations. This dilution factor is defined as the ratio of the injected bromide in the piezometer $[\text{Br}^-]_{t_0}$ to the bromide concentration at time of sampling t_i ($[\text{Br}^-]_{t_0} / [\text{Br}^-]_{t_i}$). Corrected N_2O concentrations were calculated as the product of the measured N_2O concentration at time t and the dilution factor at the same time.

After the injection, the N_2O production time rose regularly to a plateau. The maximum slope of the tangent to this curve in the ascending part represents the N_2O production rate ($= \Delta\text{N}_2\text{O corrected} / \Delta t$). The calculated N_2O production rate is finally transformed to give the N-NO_3^- denitrification rate.

GROUNDWATER MODELLING

To validate the method and determine the limits of the validation, the experiment was modelled. Transient flow and conservative solute transport during the experiment were simulated. Each transient step unit corresponds to each step of the experiment (injecting period, pumping periods, latent periods between pumps).

The model used, Processing Modflow (PM5), offers a totally integrated simulation system for modelling groundwater flow and transport processes (Chiang *et al.*, 1998). The program MODFLOW-96 (McDonald and Harbaugh, 1998) simulated three-dimensional groundwater flow during the whole experiment and revealed the impact of the injection and the different pumping operations in the piezometer. MT3D is a modular 3D solute transport model for MODFLOW for simulation of advection, dispersion and chemical reaction. The physical characteristics of the aquifer used in the model correspond to those of the study site as presented in Table 1.

STUDY SITES

This method was tested in the Garonne alluvial floodplain (south-western France) and in the quaternary aquifer of Vitoria-Gasteiz (North Spain). The French site is located on the Monbequi study site (40 km north of Toulouse). The Spanish site is located on the Arkaute site near the city of Vitoria (northern Spain). At the Monbequi site, the denitrification studies were made in six piezometers (30

Table 1. Characteristics of the piezometers tested with the method.

| <i>Location</i> | <i>Piezometer number</i> | <i>Soil use</i> | <i>Diameter of drilling (mm)</i> | <i>Diameter of piezometer (mm)</i> | <i>Total depth (m)</i> | <i>Nature of deposit</i> | <i>Depth of denitrification measurement</i> | <i>Nature of the deposit at the depth of the denitrification measurement</i> |
|-----------------|--------------------------|-------------------|----------------------------------|------------------------------------|------------------------|---------------------------------------|---|--|
| Monbequi | P8 | Poplar plantation | 76 mm | Int. : 52 mm Ext. : 60 mm | 4.9 m | Sand | 3.1 m | Sand |
| Monbequi | P9 | Riparian zone | 76 mm | Int. : 52 mm Ext. : 60 mm | 6 m | Gravel and sand | 4.1 m | Gravel and sand |
| Monbequi | P10 | Poplar plantation | 76 mm | Int. : 52 mm Ext. : 60 mm | 6 m | Gravel and sand | 4.2 m | Gravel and sand |
| Monbequi | P13 | Riparian zone | 76 mm | Int. : 52 mm Ext. : 60 mm | 3 m | Sand and clay | 1.7 m | Sand and clay |
| Monbequi | P26 | Agricultural land | 76 mm | Int. : 52 mm Ext. : 60 mm | 5.5 m | Gravel and sand | 4 m | Gravel and sand |
| Monbequi | P29 | Agricultural land | 76 mm | Int. : 52 mm Ext. : 60 mm | 7.5 m | Gravel and sand | 6 m | Gravel and sand |
| Arkaute | P5 | Meadows | 125 mm | Int. : 80 mm Ext. : 90 mm | 12 m | Clay and sand. Marl at 4 m | 3.5 m 8 m | Clay Marls |
| Arkaute | P11 | Meadows | 125 mm | Int. : 80 mm Ext. : 90 mm | 8 m | Sand, clay and gravel. Marl at 4 m | 6 m | Marls |

tests): two in a riparian zone (P9 and P13), two in a poplar plantation (P8 and P10) and two in agricultural land (P26 and P29). At the Arkaute site, measurements were made in two different piezometers (three tests) : P11 and P5 (at two different depths). The characteristics of the piezometers are given in Table 1.

Results

BACKGROUND CONDITIONS

Nitrate concentrations ranged from 2.76 to 16.45 mg N l⁻¹ except for one sample at a depth of 8 m in piezometer 11 in the Vitoria Gasteiz Aquifer; this had a nitrate-nitrogen content concentration of 40.90 mg N l⁻¹. Concentrations in each piezometer were rather stable over time (Table 2).

Dissolved oxygen ranged from 1.4 to 9.3 mg l⁻¹. Dissolved organic carbon (DOC) ranged from 0.96 to 1.99 mg l⁻¹, except for the 3.5 m sample in piezometer 5 in the Vitoria-

Gasteiz aquifer, which had a DOC concentration of 3.56 mg l⁻¹.

INJECTION TEST RESULTS

Figure 3 shows how bromide and N₂O behave with time in piezometer P8 at Monbequi. The kinetic of the denitrification is exponential. The calculations of the denitrification rate are represented in Table 3. The concentrations in nitrate and in DOC remained constant during the measurement. Pumping groundwater causes oxygenation but, during the injection of water into the piezometer, the dissolved oxygen concentration decreased to its initial value.

Figure 4 shows two other cases for which the method is not applicable. In the first case (P26 at Monbequi), dilution is too high and inhibition of denitrification is not guaranteed. It gives very high dilution rates and the N₂O correction for dilution consequently provides a very high and inaccurate N₂O production rate. There may have been a problem with

Table 2. Background concentration of N-NO₃⁻, DOC and Dissolved oxygen in the piezometers (mean ± SE). N=number of samples. SE = standard error.

| Localisation | N | Piezometer number | N-NO ₃ ⁻ (mg l ⁻¹) | Dissolved Oxygen range (mg l ⁻¹) | DOC (mg l ⁻¹) |
|--------------|---|-------------------|--|--|---------------------------|
| Monbequi | 6 | P8 | 13.43 ± 2.47 | 4.5 – 7.1 | 1.54 ± 0.11 |
| Monbequi | 5 | P9 | 8.19 ± 0.44 | 1.6 – 5.4 | 1.60 ± 0.17 |
| Monbequi | 5 | P10 | 13.08 ± 1.27 | 1.4 – 6.2 | 1.17 ± 0.13 |
| Monbequi | 5 | P13 | 7.05 ± 1.47 | 1.0 – 3.5 | 1.75 ± 0.16 |
| Monbequi | 5 | P26 | 11.42 ± 1.07 | 5.4 – 7.7 | 1.58 ± 0.18 |
| Monbequi | 5 | P29 | 13.44 ± 0.09 | 5.4 – 8.1 | 1.10 ± 0.05 |
| Arkaute | 2 | P5 | 1.21 ± 0.09 | 0.6 – 1.4 | 3.56 |
| Arkaute | 1 | P11 | 40.90 ± 0.75 | 5.1 | – |

the watertightness of the packer system and bromide may have leaked. This situation can be compared with the case of a very permeable medium.

In the second case (piezometer P5 at Arkaute), in a clayey soil, the injected water was blocked in the piezometer. The decrease in Br⁻ concentration between injection and sampling at time t₁₅ is due to the dilution by the water stored in the piezometer. These results show that the N₂O production rate is very low.

The average rate of nitrate consumption measured with this method ranged from 6.51 10⁻⁶ g N-NO₃⁻ L⁻¹ h⁻¹ to 2.06 10⁻³ g N-NO₃⁻ L⁻¹ h⁻¹ (Table 4). On the Monbequi site, the rate of N₂O production differs among the three compartments. The riparian zone showed a maximum production of 2.06 10⁻³ g N-NO₃⁻ L⁻¹ h⁻¹, the poplar plantation of 0.97 10⁻³ g N-NO₃⁻ L⁻¹ h⁻¹ while in the agricultural land the minimum denitrification measured was 0.10 10⁻³ g N-NO₃⁻ L⁻¹ h⁻¹.

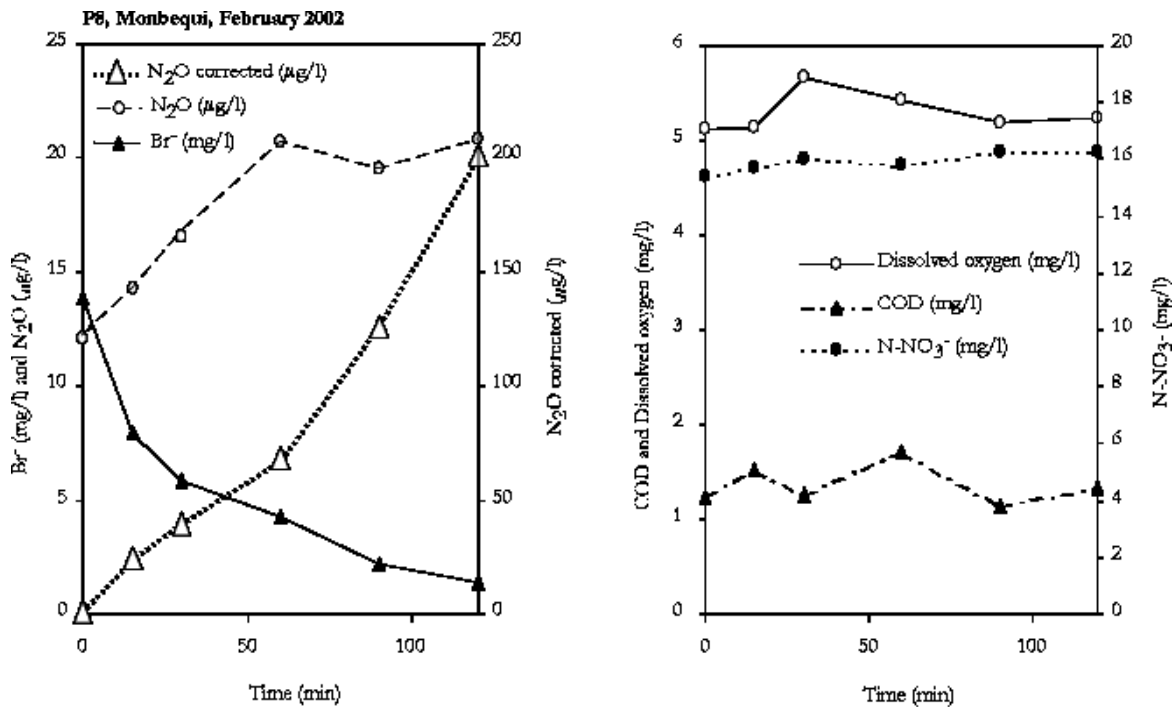


Fig. 3. Bromide, N₂O and corrected N₂O, DOC, Dissolved oxygen and nitrogen nitrate concentrations versus time in a piezometer during an injection test.

Table 3. Example of calculation of the denitrification rate (Piezometer P8, Monbequi, February 2002).

| Time (min) | N_2O produced ($mg\ l^{-1}$) | Br^- ($mg\ l^{-1}$) | Dilution rate [Br_{10}/Br_{it}] | N_2O corrected ($\mu g\ l^{-1}$) | N_2O produced ($\mu g\ l^{-1}\ min^{-1}$) | $N-NO_3^-$ denitrified ($g\ l^{-1}\ h^{-1}$) |
|------------|----------------------------------|-------------------------|-------------------------------------|--------------------------------------|---|--|
| 0 | 12.1 | 13.9 | 1.0 | 12.1 | | |
| 15 | 14.3 | 7.9 | 1.1 | 25.1 | | |
| 30 | 16.6 | 5.9 | 1.8 | 39.3 | | |
| 60 | 20.7 | 4.2 | 4.3 | 68.0 | 0.94 | $3.77\ 10^{-5}$ |
| 90 | 19.5 | 2.1 | 6.2 | 126.3 | | |
| 120 | 20.8 | 1.4 | 9.3 | 201.5 | | |

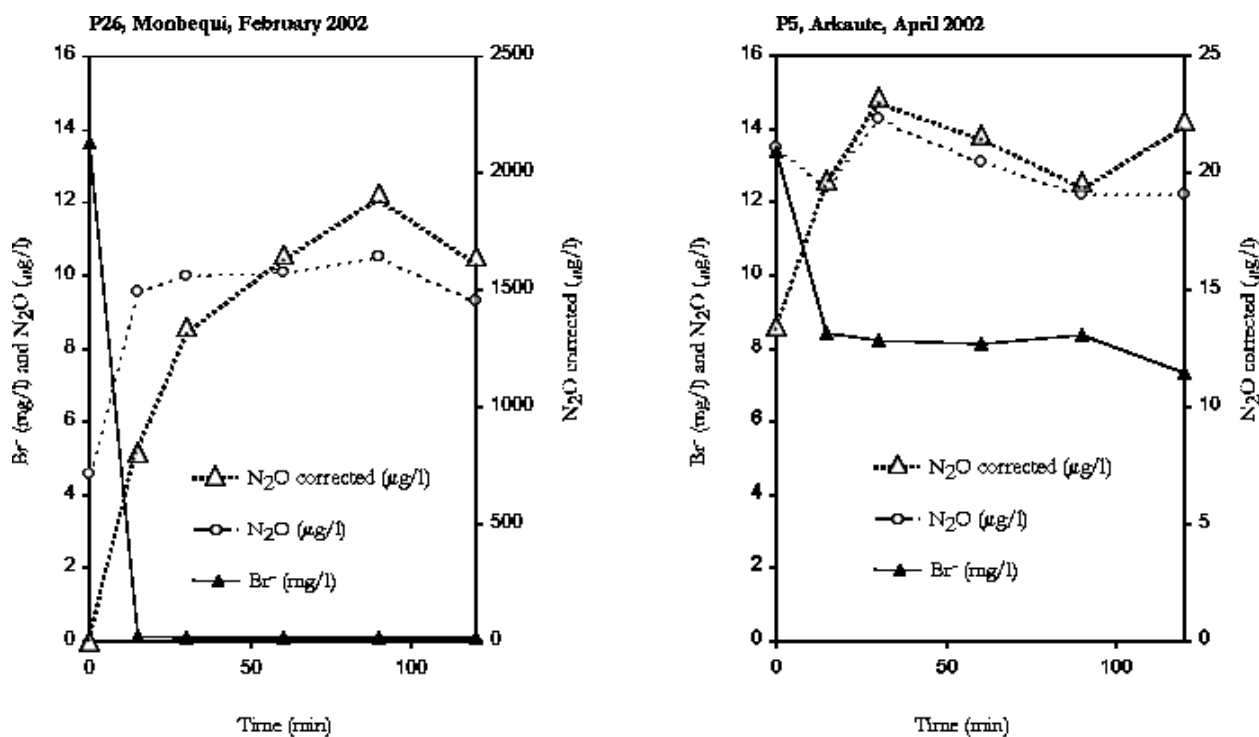


Fig. 4. Bromide, N_2O and corrected N_2O concentrations versus time with the dilution for two injection tests.

Table 4. Denitrification rate ($g\ N-NO_3^-\ L^{-1}\ h^{-1}$) measured by the tracer test (P10 and P13 were located in a riparian zone, P8 and P10 in a poplar plantation and P26 and P29 in an agricultural zone).

| | P8 | P9 | P10 | P13 | P26 | P29 |
|----------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|
| February | $3.77\ 10^{-5}$ | $1.77\ 10^{-5}$ | $6.50\ 10^{-6}$ | $4.06\ 10^{-5}$ | $2.14\ 10^{-5}$ | $2.34\ 10^{-5}$ |
| March | $1.20\ 10^{-5}$ | $8.24\ 10^{-6}$ | $1.08\ 10^{-4}$ | $8.30\ 10^{-5}$ | $1.20\ 10^{-5}$ | $1.51\ 10^{-5}$ |
| April | $2.14\ 10^{-4}$ | $8.24\ 10^{-6}$ | $1.08\ 10^{-4}$ | $2.06\ 10^{-3}$ | $5.23\ 10^{-5}$ | $1.57\ 10^{-5}$ |
| Mai | $6.54\ 10^{-4}$ | $2.10\ 10^{-5}$ | $9.73\ 10^{-4}$ | $2.04\ 10^{-3}$ | $1.02\ 10^{-4}$ | $1.53\ 10^{-5}$ |
| June | $3.03\ 10^{-4}$ | $2.02\ 10^{-4}$ | $1.10\ 10^{-3}$ | $6.33\ 10^{-4}$ | $2.43\ 10^{-5}$ | $1.87\ 10^{-5}$ |

HYDROGEOLOGICAL MODELLING

The model grid comprises three horizontal layers (Fig. 5), the thickness of the second layer corresponding to the thickness of the interval isolated by the packer system. The size dimension of the grid is 10 m × 10 m and the dimension of the cells is 20 cm × 20 cm. The depth of sampling section in the piezometer is around 3 m which corresponds to the mean underground depth at the study site during the experiments. Analysis of sensitivity to hydraulic conductivity was calculated over a hydraulic conductivity range of 10⁻¹ to 10⁻⁸ m.s⁻¹. The hydraulic gradient of the water table is 0.3 %.

For the transport simulations, the ratios, horizontal transverse dispersivity/longitudinal dispersivity and vertical transverse dispersivity/longitudinal dispersivity are equal to 0.1. The flow simulations by PM5 show that the pumping plan makes groundwater depth move by 0.16%. The transport simulation results by MT3D (Fig 5, with k = 10⁻⁴ m.s⁻¹) show that the pumping plan influenced only a 120 cm radius around the piezometer. Therefore, the sampled part of the aquifer was not disturbed significantly during the pumping operation. The top face (1) shows that the

tracer injected was spread homogeneously all around the piezometer; the front (2) and side views (3) show that the tracer stays at the depth of the packer during the experiment and, therefore, the sampling was done in water blocked by C₂H₂.

The results of the model for different hydraulic conductivities in the physical context of the experiments are represented in Fig. 6. Concentrations of Br⁻ were normalised according to the equation : $1 - [(C_{inj} - C_t) / (C_{inj} - C_{ini})]$, with C_{inj} the concentration of Br⁻ in the injection, C_{ini} the initial concentration of Br⁻ in the groundwater and C_t the concentration calculated by the model at each time step. Table 5 presents the initial and injected concentrations in the model. For a given hydraulic conductivity, normalisation gives the same results independently of the injected concentrations and of the initial conditions. The decrease of Br⁻ due to dilution by dispersion is linked to the hydraulic conductivity of the aquifer. To sample the water blocked by C₂H₂ during the whole experiment and to calculate the denitrification rate, the method was optimal for values of k from 10⁻² to 10⁻⁴ m s⁻¹ in the physical and hydrological context.

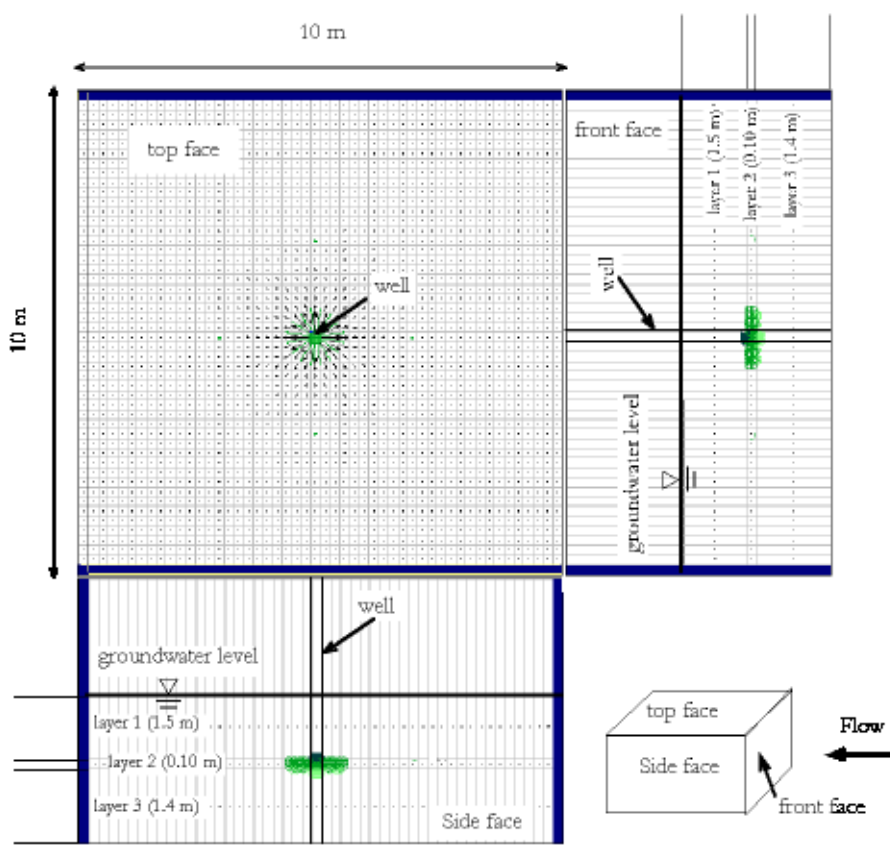


Fig. 5. Modelling injection in a piezometer with an aquifer hydraulic conductivity of 10⁻⁴ m s⁻¹.

Table 5. Concentrations used for modelling with different hydraulic conductivities.

| | Hydraulic conductivity from 10^{-1} to 10^{-8} $m\ s^{-1}$ | | | | Hydraulic conductivity 10^{-3} $m\ s^{-1}$ | | |
|--|---|-----|------|-----|---|----|----|
| Initial concentration ($mg\ L^{-1}$) | 0.1 | 0.1 | 0.1 | 0.1 | 0.5 | 1 | 5 |
| Injected concentration ($mg\ L^{-1}$) | 1 | 10 | 13.4 | 20 | 1 | 10 | 20 |

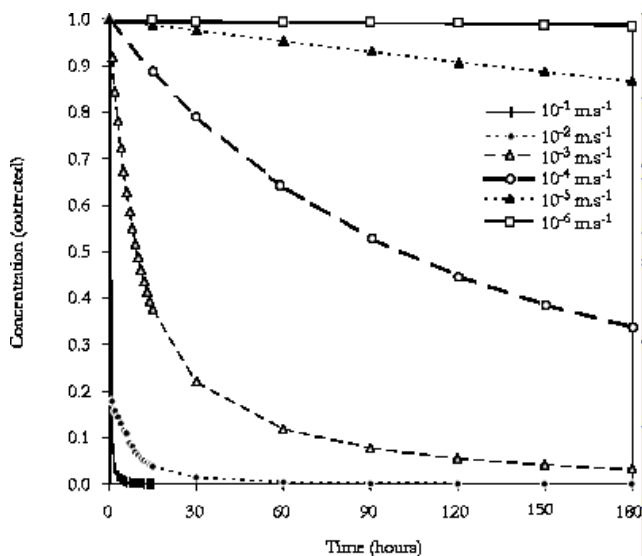


Fig. 6. Normalised concentration of bromide as a function of time for different hydraulic conductivities.

Discussion

The blockage of N_2O reduction by acetylene is indeed valid for measurement of denitrification in soils and sediments, as well in the laboratory and in studies *in situ*. Studies have been done on soils and sediments, using core samples (Smith and Duff, 1988 ; Pinay *et al.*, 1993; Bengtsson and Bergwall, 1995 ; Burt *et al.*, 1999) or *in situ* microcosms (Starr and Gillham, 1993).

Some studies try to evaluate a denitrification rate in aquifers by bacteria located in the groundwater. C_2H_2 and a tracer, most often bromide, are injected in a piezometer and the evolution of bromide and N_2O transported by groundwater is followed by means of surrounding piezometers; however, these experiments often require several days (Starr and Gillham, 1993; McMahon and Böhlke, 1996) and do not reveal how the mixture of bromide and acetylene is transported and what is really sampled. The flow from an assumed upstream piezometer to an assumed downstream one is too uncertain; moreover the N_2O could

be transformed into N_2 as the C_2H_2 blockage may become inactive.

The results of the present model show that the concentration decrease in Br^- depends directly on the hydraulic conductivity of the aquifer. During the experiment (2 hours) the dilution processes are dominated by dispersion. For this method, the hydrogeological conditions are optimal for a hydraulic conductivity between 10^{-2} and 10^{-4} $m\ s^{-1}$. For a hydraulic conductivity higher than 10^{-2} $m\ s^{-1}$, dilution mainly by dispersion is too high and the effect of the inhibition with the acetylene is reduced with time, which in turn makes the measurements very difficult. In contrast, for a hydraulic conductivity lower than 10^{-4} $m\ s^{-1}$, the transfer into the aquifer is very weak, the tracer remains in the piezometer during the whole time of the experiment and C_2H_2 may decrease by dispersion or by exchange with air. As the concentration of C_2H_2 decreases during the experiment (sampling operation, dilution), total inhibition of N_2O reductases cannot be considered efficient after 180 min in the experimental conditions.

Modifications of the environment caused by this method (oxygen entry to the aquifer, dilution process of the water volume studied) did not seem to modify microbial action significantly.

The use of C_2H_2 to measure the denitrification rates does not allow measurement of the denitrification of NO_3^- possibly produced by nitrification : acetylene not only inhibits the N_2O reductases activity but also, even at lower concentrations, the ammonium mono-oxygenase (thus ammonium oxidizing bacteria and nitrification) (Hyman and Wood, 1985).

In the present study, because of the high nitrate concentrations, denitrification resulting from nitrification products was considered negligible compared to the denitrification of the circulating nitrates.

As neither nitrate, oxygen nor carbon concentrations change significantly during the injection experiment, it can be assumed that the linear part of the N_2O increase within the water sampled gives information on the effective *in situ* denitrification rate.

By this method, it is possible to measure the denitrification rates occurring within the alluvial aquifer and, therefore, the quantity of nitrogen that is exported to the atmosphere if the groundwater flux is known.

The denitrification rate measured here lies within the range of values reported in shallow groundwater of Ontario during a 10-day experiment by Trudell *et al.* (1986) : $7.8 \cdot 10^{-6}$ to $1.3 \cdot 10^{-3}$ g N-NO₃⁻ L⁻¹ h⁻¹ and Starr and Gillham (1993) : $2.4 \cdot 10^{-5}$ g N-NO₃⁻ L⁻¹ h⁻¹.

The acetylene method is not the only one to measure a denitrification rate; Addy *et al.* (2002) used the ¹⁵N method in a riparian aquifer at Rhode Island. However, the acetylene method is the most suitable for measuring denitrification in groundwater. The denitrification rates obtained by this method were in the range 17.6 to $22.4 \cdot 10^{-6}$ g N-NO₃⁻ L⁻¹ h⁻¹.

DENITRIFICATION WITHIN A POROUS MEDIUM IS PLAYED BY BIOFILMS

The results of this experiment also gave information on the sites where denitrification processes take place. Despite the presence of high concentrations of dissolved oxygen (from 0.6 up to 8.1 mg O₂ L⁻¹) during the experiments, denitrification detected in each sampled piezometer shows that microbial denitrification processes do not occur in interstitial water. Referring to previous studies on periphytic biofilms demonstrating that denitrification could occur within oxygenated water in superficial water (Nielsen *et al.*, 1990; Teissier *et al.*, 2002), it has been assumed that denitrification is localised within the biofilms growing around particles included in the sediments making up the porous aquifer.

Conclusion

A standardised method is proposed to measure *in situ* denitrification rates within the saturated zone of a porous aquifer. In addition, the hydrogeological model enables measurement of the hydraulic conductivity within the aquifer. This information, very important if the injection experiment can be undertaken in the given porous medium, also indicates the 'residence time' and dilution factor of interstitial water within the microenvironment of denitrification sites. This experiment shows that the denitrification process intensity is controlled, not only by the basic parameters, i.e. nitrate, carbon and dissolved oxygen, and by the permeability of the porous medium, but also by the biofilm metabolism and by the water exchanges between the biologically active sites (i.e. most probably biofilms) and the circulating interstitial water.

It is concluded that denitrification activity is negatively correlated with the hydraulic conductivity of the porous medium; the lower the permeability the higher the microbial community denitrification activity when carbon and nitrate are available. The conditions for this method are optimal for shallow aquifers with a hydraulic conductivity of 10^{-2} to 10^{-4} ms⁻¹.

Acknowledgements

These investigations were part of the ECOBAG program. The PNRZH (French National Programme of Research on Wetlands) and the Agence de l'Eau Adour-Garonne provided financial support for this project. The authors also wish to thank three anonymous referees for their valuable comments.

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