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Inter-annual variability of the carbon dioxide oceanic sink south of Tasmania

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Abstract. We compiled a large data-set from 22 cruises spanning from 1991 to 2003, of the partial pressure of CO₂ (pCO₂) in surface waters over the continental shelf (CS) and adjacent open ocean (43° to 46° S; 145° to 150° E), south of Tasmania. Climatological seasonal cycles of pCO₂ in the CS, the subtropical zone (STZ) and the subAntarctic zone (SAZ) are described and used to determine monthly pCO₂ anomalies. These are used in combination with monthly anomalies of sea surface temperature (SST) to investigate inter-annual variations of SST and pCO₂. Monthly anomalies of SST (as intense as 2°C) are apparent in the CS, STZ and SAZ, and are indicative of strong inter-annual variability that seems to be related to large-scale coupled atmosphere-ocean oscillations. Anomalies of pCO₂ normalized to a constant temperature are negatively related to SST anomalies. A reduced winter-time vertical input of dissolved inorganic carbon (DIC) during phases of positive SST anomalies, related to a poleward shift of westerly winds, and a concomitant local decrease in wind stress is the likely cause of the negative relationship between pCO₂ and SST anomalies. The observed pattern is an increase of the sink for atmospheric CO₂ associated with positive SST anomalies, although strongly modulated by inter-annual variability of wind speed. Assuming that phases of positive SST anomalies are indicative of the future evolution of regional ocean biogeochemistry under global warming, we show using a purely observational based approach that some provinces of the Southern Ocean could provide a potential negative feedback on increasing atmospheric CO₂.

1 Introduction

The ocean is a major and dynamic sink for anthropogenic CO₂ (e.g. Sabine et al., 2004) playing an important role in the mitigation of climate change. The inclusion in climate models of potential feedbacks of air-sea CO₂ fluxes on the increase of atmospheric CO₂ is required to improve the reliability of the predictions of the future evolution of the global carbon cycle and climate change. The collection of partial pressure of CO₂ (pCO₂) data for surface waters during the last 30 years has allowed the investigation and characterisation of changes in air-sea CO₂ fluxes on decadal scales in some regions of the open ocean (e.g., North Atlantic Ocean (Lefèvre et al., 2004; Corbière et al., 2007; Bates, 2007), Pacific Ocean (Feely et al., 2006; Midorikawa et al., 2006; Takahashi et al., 2006), and Southern Ocean (Inoue and Ishii, 2005)). These studies have described how pCO₂ in surface waters and the associated air-sea CO₂ fluxes are evolving, and provide data to constrain and understand feedbacks in the carbon cycle related to changes in oceanic physical and biogeochemical processes.

The investigation of long term trends in surface water CO₂ requires the description of the inter-annual variability of pCO₂. The main drivers of the inter-annual variability of surface pCO₂ described to date are large-scale atmosphere-ocean coupled climate oscillations including the El Niño Southern Oscillation (ENSO) for the equatorial and subtropical Pacific Ocean (Feely et al., 2002; Dore et al., 2003; Brix et al., 2005), and the North Atlantic Oscillation for the North Atlantic Ocean (Gruber et al., 2002). In the Southern Ocean, atmosphere-ocean coupled climate oscillations such as the Southern Annular Mode (SAM) and ENSO have been identified using biogeochemical ocean general circulation models as major drivers of inter-annual variability of pCO₂ (Le



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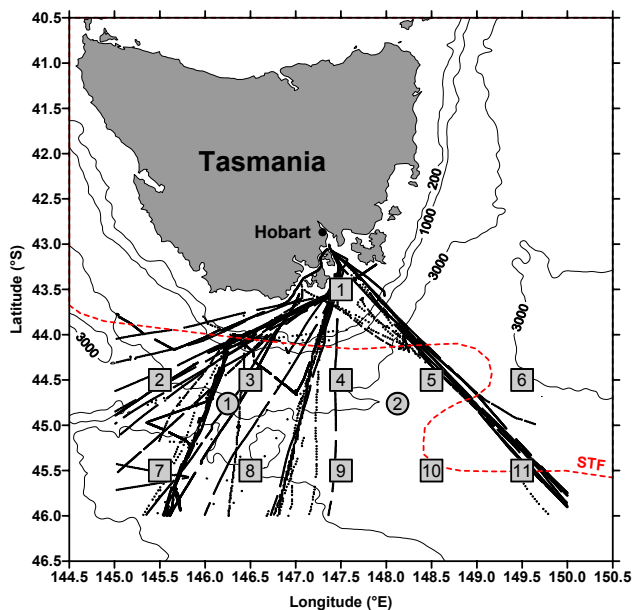


Fig. 1. Map showing ship tracks, bathymetry based on the Smith and Sandwell (1997) global seafloor topography, the climatological position of sub-tropical front (STF) based on Belkin and Gordon (1996) and Hamilton (2006), grid nodes from the Reynolds et al. (2002) sea surface temperature monthly climatology (squares), and the grid nodes of the Kalnay et al. (1996) National Centers for Environmental Prediction daily wind speeds (circles).

Quéré et al., 2007; Lenton and Matear, 2007; Lovenduski et al., 2007; Verdy et al., 2007).

Due to the relative scarcity of pCO₂ field data in the Southern Ocean, the inter-annual variability of air-sea CO₂ fluxes has seldom been investigated from a purely observational based approach. The impact of warm anomalies on air-sea CO₂ fluxes in the Southern Ocean has to some extent been investigated using cruise-to-cruise comparisons (Jabaud-Jan et al., 2004; Brévière et al., 2006). These two studies show that warm anomalies lead to significant, but opposing, effects on air-sea CO₂ fluxes in different regions of the high latitude Southern Ocean. This could be due to the spatially heterogeneous control of export production through either light or nutrient limitation in the Southern Ocean (Le Quéré et al., 2000, 2002, 2003), with some modulation by thermodynamic effects of sea surface temperature (SST) change on pCO₂.

The aim of the present work is to investigate at different time scales the variations of pCO₂ in the surface waters of the continental shelf (CS) and adjacent open ocean (43° to 46° S; 145° to 150° E) south of Tasmania, based on a compilation and synthesis of 40 transects obtained during 22 cruises (Fig. 1, Table 1). We first compose and describe the climatological seasonal cycles of pCO₂ in the CS, the subtropical zone (STZ) and the subAntarctic zone (SAZ). We then proceed to investigate the inter-annual variability of SST and pCO₂, using monthly anomalies to deseasonalize the data-

sets. We also examine the relationship between anomalies in air-sea CO₂ fluxes and SST on inter-annual time scales to provide insights into the biogeochemical responses of mid-latitudes of the Southern Ocean to global warming.

2 Methods

Measurements of pCO₂ were obtained with the equilibration technique as described by Frankignoulle et al. (2001) for the Université of Liège (ULg), by Lenton et al. (2006) for Commonwealth Scientific and Industrial Research Organisation (CSIRO), and by Poisson et al. (1993) for Laboratoire d'Océanographie et du Climat: Expérimentations et Approches Numériques/Institut Paul Simon Laplace (LOCEAN/IPSL). CSIRO and LOCEAN/IPSL systems were inter-calibrated during the *R.V. Meteor* international at-sea intercomparison (06/06–19/06/1996) in the North Atlantic (Körtzinger et al., 2000), and the results showed that the pCO₂ data were consistent within $\pm 1 \mu\text{atm}$. ULg and LOCEAN/IPSL systems were inter-calibrated during the OISO3 cruise (21/12–28/12/1998) in the central Indian sector of the Southern Ocean, and the results showed that the pCO₂ data were consistent within $\pm 4 \mu\text{atm}$. ULg and CSIRO systems were inter-calibrated during the AA0301 cruise (11/10–27/10/2003) in eastern Indian sector of the Southern Ocean, and the results showed that the pCO₂ data were consistent within $\pm 5 \mu\text{atm}$. Since 82% of the data were obtained by one single group (CSIRO; Table 1), we assume the uncertainty of the whole data-set to be better than $\pm 3 \mu\text{atm}$. All data were converted to pCO₂ in wet air at 1 atm. During the time-span of the data-set (from 1991 to 2003), atmospheric pCO₂ increased by $20 \mu\text{atm}$ ($1.7 \mu\text{atm yr}^{-1}$). Based on four cruises carried out between 1969 and 2002, Inoue and Ishii (2005) showed that pCO₂ in surface waters in the STZ and SAZ near Tasmania increases at a rate very close to the increase of atmospheric CO₂. Hence, data were referenced to 1997, the middle of the time-series, according to:

$$p\text{CO}_{2\text{ sea }1997} = p\text{CO}_{2\text{ sea year}} + (p\text{CO}_{2\text{ air }1997} - p\text{CO}_{2\text{ air year}}) \quad (1)$$

where pCO_{2 sea 1997} is the pCO₂ in seawater referenced to 1997, pCO_{2 sea year} is the pCO₂ in seawater from a given year, pCO_{2 air 1997} is the atmospheric pCO₂ in 1997, and pCO_{2 air year} is the atmospheric pCO₂ for the same given year.

Atmospheric pCO₂ data from the Cape Grim station (40.7° S, 144.7° E; Tasmania) were obtained from the Cooperative Air Sampling Network of the National Oceanic and Atmospheric Administration/Earth System Research Laboratory/Global Monitoring Division (<http://www.cmdl.noaa.gov/>). Hereafter, pCO₂ refers to pCO_{2 sea 1997}.

Air-sea CO₂ fluxes were computed according to:

$$F = k\alpha\Delta p\text{CO}_2 \quad (2)$$

where F is the air-sea CO₂ flux, k is the gas transfer velocity, α is the CO₂ solubility coefficient, and $\Delta p\text{CO}_2$ is the

Table 1. Cruises, ships, institutes, dates (dd/mm/yyyy) of transects in the continental shelf (CS), the subtropical zone (STZ) and the subAntarctic zone (SAZ) south of Tasmania.

Cruise	ship	institutes	date	CS	STZ	SAZ
V191	<i>R.S.V. Aurora Australis</i>	CSIRO	05/10/1991	+	+	+
			25/10/1991	+	+	+
V192	<i>R.S.V. Aurora Australis</i>	CSIRO	21/11/1992	+		
V792	<i>R.S.V. Aurora Australis</i>	CSIRO	05/01/1993	+	+	
			07/03/1993	+	+	+
V993	<i>R.S.V. Aurora Australis</i>	CSIRO	12/03/1993	+	+	+
			08/05/1993	+	+	+
V193	<i>R.S.V. Aurora Australis</i>	CSIRO	07/08/1993	+	+	+
			08/10/1993	+	+	+
V493	<i>R.S.V. Aurora Australis</i>	CSIRO	19/11/1993	+		+
			27/12/1993	+		+
AA9407	<i>R.S.V. Aurora Australis</i>	CSIRO	02/01/1994	+	+	+
			28/02/1994	+	+	+
AA9401	<i>R.S.V. Aurora Australis</i>	CSIRO	31/08/1994	+	+	+
			13/12/1994	+	+	+
AA9404	<i>R.S.V. Aurora Australis</i>	CSIRO	31/01/1995	+	+	+
			18/07/1995	+	+	+
AA9501	<i>R.S.V. Aurora Australis</i>	CSIRO	07/09/1995	+		+
SS9511	<i>R.V. Southern Surveyor</i>	CSIRO	23/11/1995	+		+
			19/01/1996	+		+
AA9604	<i>R.S.V. Aurora Australis</i>	CSIRO	30/03/1996	+	+	+
			22/08/1996	+	+	+
AA9601	<i>R.S.V. Aurora Australis</i>	CSIRO	20/09/1996	+	+	+
			21/10/1996	+	+	+
MINERVE38	<i>S.V. Astrolabe</i>	LOCEAN/IPSL	23/11/1996	+	+	+
			02/02/1997	+	+	+
MINERVE39	<i>S.V. Astrolabe</i>	LOCEAN/IPSL	17/02/1997		+	
			14/11/1997		+	+
AA9703	<i>R.S.V. Aurora Australis</i>	CSIRO	26/11/1997	+	+	+
			28/02/1998	+	+	
AA9706	<i>R.S.V. Aurora Australis</i>	CSIRO/ULg	31/03/1998	+	+	
			05/02/1999	+		
SS9902	<i>R.V. Southern Surveyor</i>	CSIRO	16/02/1999	+	+	+
			16/07/1999	+	+	+
AA9901	<i>R.S.V. Aurora Australis</i>	CSIRO	05/09/1999	+	+	+
99R0	<i>S.V. Astrolabe</i>	ULg	22/10/1999	+	+	
99R1	<i>S.V. Astrolabe</i>	ULg	26/12/1999	+	+	+
OISO10	<i>R.V. Marion Dufresne</i>	LOCEAN/IPSL	29/01/2003		+	
			11/09/2003	+	+	
AA0301	<i>R.S.V. Aurora Australis</i>	CSIRO/ULg	29/10/2003	+	+	+

air-sea pCO₂ gradient. We used the *k*-wind parameterization of Wanninkhof (1992), and the Weiss (1974) formulation of α as a function of sea surface salinity (SSS) and SST. Atmospheric pCO₂ data from 1997 were expressed in wet air using the water vapour pressure formulation of Weiss and Price (1980) as a function of SSS and SST.

Data were collocated with bathymetry based on the Smith and Sandwell (1997) global seafloor topography (<http://topex.ucsd.edu/>). Data over the CS were gathered and averaged by sorting for depths <300 m, and data at depths >1000 m were considered as open ocean. The sub-Tropical

Front (STF) was identified from gradients of SSS and SST (e.g. Belkin and Gordon, 1996). The STF separates warmer and saltier STZ waters from cooler and fresher SAZ waters.

3 Results

3.1 Climatological SST and pCO₂ seasonal cycles

Climatological seasonal cycles of SST, pCO₂ and pCO₂ normalized to a temperature of 14°C (pCO₂@14°C, using the algorithms of Copin-Montégut, 1988, 1989) were obtained

by fitting monthly averages with a wave function in the form of:

$$y = a + b \sin\left(\frac{x}{c} + d\right) \quad (3)$$

where y is either SST, pCO₂ or pCO₂@14°C, x is time (julian days) and a , b , c , and d are fitted constants.

In the CS, STZ and SAZ, under-saturation of CO₂ is observed throughout the year (Figs. 2, 3, 4), showing these regions are perennial sinks for atmospheric CO₂. In the 3 regions, similar climatological pCO₂ and pCO₂@14°C seasonal patterns are observed in timing and amplitude: values decrease from late September to late February (austral spring-summer) as net biological uptake removes dissolved inorganic carbon (DIC) from surface waters. From March to September (austral fall-winter), pCO₂ and pCO₂@14°C values increase in relation to destratification and mixing of surface waters with DIC rich deeper waters (Goyet et al., 1991; Poisson et al., 1993; Metzl et al., 1991, 1995, 1998, 1999). The amplitude of the seasonal cycle of pCO₂ is lower than the one of pCO₂@14°C because warming of surface waters during spring and summer leads to a thermodynamic increase of pCO₂, that opposes a decrease due to net biological carbon uptake. The thermodynamic effect of SST change on the seasonal amplitude of pCO₂ is similar in the CS (51 μatm), the STZ (50 μatm), and the SAZ (47 μatm), because SST amplitude is similar in the 3 regions (3.7°C, 3.6°C and 3.3°C, respectively).

The SSS values for all three water masses do not show a distinct seasonal signal like SST, which is strongly influenced by seasonal heating and cooling of surface waters. The water mass on the Tasmanian CS is a mixture of STZ and SAZ waters (Harris et al., 1987, 1991). This is apparent in our data-set, as the average SSS value in the CS (35.02±0.17) lies between those of the STZ (35.19±0.15) and the SAZ (34.73±0.15). The SSS values in the CS show more scatter than in the other two regions (Figs. 2, 3, 4), suggesting a variable degree of mixing between the STZ and SAZ water masses. For a two end-member mixing model, the water mass on the CS is on average composed of 64% STZ water, consistent with the annual average of SST in the CS (13.8±1.3°C) being similar to the STZ (13.4±1.3°C) and distinctly different from the SAZ (11.5±1.2°C) annual averages.

3.2 Inter-annual SST and pCO₂ variations

In order to deseasonalize the data-sets and investigate inter-annual variability, monthly anomalies of SST, pCO₂, and pCO₂@14°C were computed as the difference between observations and averaged monthly values for all the cruise data. SST anomalies as intense as 2°C are observed in the CS, STZ and SAZ (Figs. 2, 3, 4). SST values that are below the climatology tend to coincide with less saline waters suggesting a greater contribution of SAZ waters. Warm anomalies are typically associated with saltier waters implying a

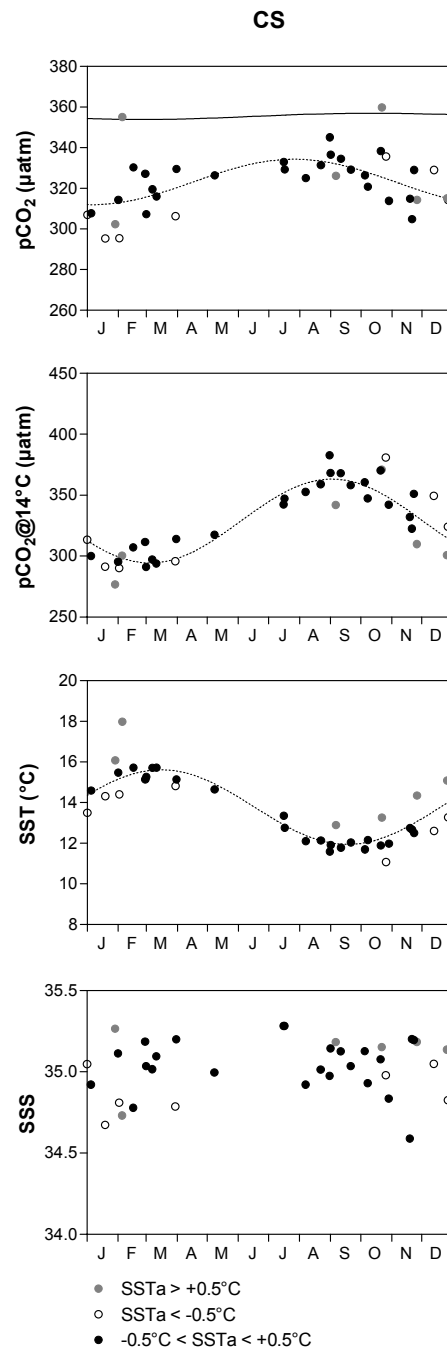


Fig. 2. Seasonal cycles of the partial pressure of CO₂ (pCO₂), the pCO₂ normalized to a temperature of 14°C (pCO₂@14°C), sea surface temperature (SST) and sea surface salinity (SSS) over the continental shelf (CS) south of Tasmania. Solid line shows atmospheric pCO₂ for 1997. Dotted lines show the climatological cycles based on a wave function curve fitted to the monthly averages. Data were sorted for significant positive SST anomalies (SSTA>0.5°C, grey circles), significant negative SST anomalies (SSTA<-0.5°C, open circles), and no significant SST anomalies (-0.5°C<SSTA<0.5°C, black circles).

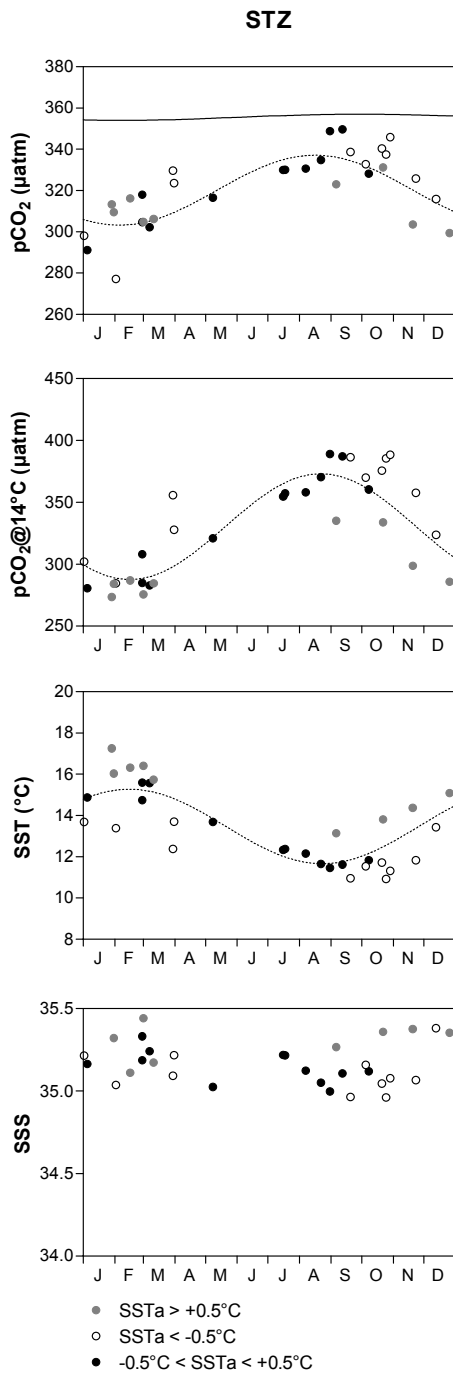


Fig. 3. Seasonal cycles of the partial pressure of CO₂ (pCO₂), the pCO₂ normalized to a temperature of 14°C (pCO₂@14°C), sea surface temperature (SST) and sea surface salinity (SSS) in the subtropical zone (STZ) south of Tasmania. Solid line shows atmospheric pCO₂ for 1997. Dotted lines show the climatological cycles based on a wave function curve fitted to the monthly averages. Data were sorted for significant positive SST anomalies (SSTA>0.5°C, grey circles), significant negative SST anomalies (SSTA<-0.5°C, open circles), and no significant SST anomalies (-0.5°C<SSTA<0.5°C, black circles).

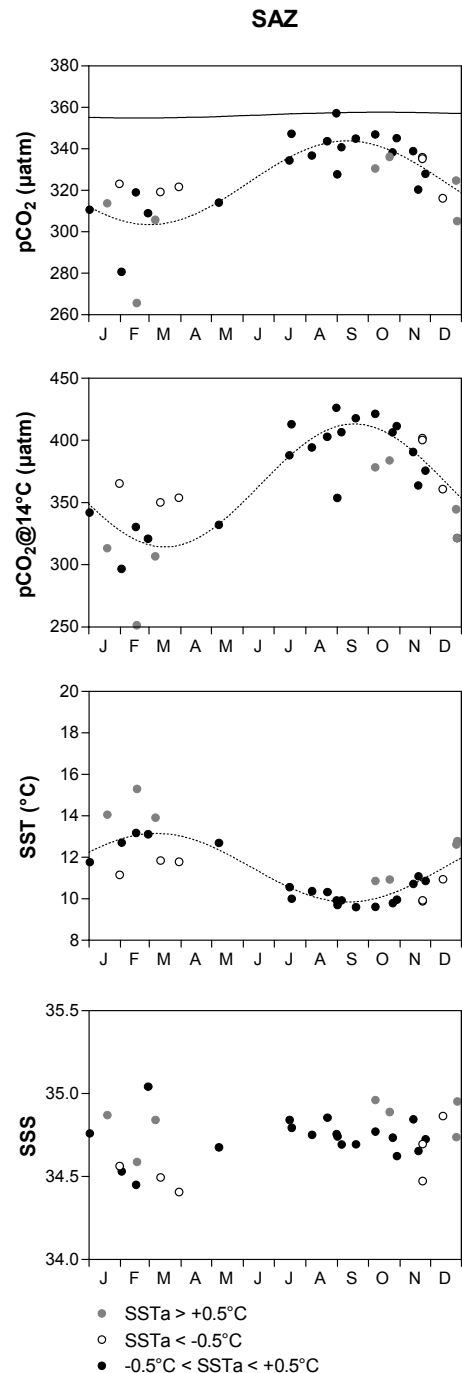


Fig. 4. Seasonal cycles of the partial pressure of CO₂ (pCO₂), the pCO₂ normalized to a temperature of 14°C (pCO₂@14°C), sea surface temperature (SST) and sea surface salinity (SSS) in the subAntarctic zone (SAZ) south of Tasmania. Solid line shows atmospheric pCO₂ for 1997. Dotted lines show the climatological cycles based on a wave function curve fitted to the monthly averages. Data were sorted for significant positive SST anomalies (SSTA>0.5°C, grey circles), significant negative SST anomalies (SSTA<-0.5°C, open circles), and no significant SST anomalies (-0.5°C<SSTA<0.5°C, black circles).

greater contribution of subtropical waters and consistent with observations of increases in warming and salinity for the region (Rintoul and Sokolov, 2001; Rintoul and Trull, 2001; Morrow et al., 2007).

A comparison of the monthly SST anomalies with anomalies for pCO₂ and pCO₂@14°C show consistent patterns in the three regions (Figs. 2, 3, 4). Values of pCO₂@14°C associated with positive SST anomalies (>0.5°C) generally lie below the climatological pCO₂@14°C cycle. Values associated with significant negative SST anomalies (<-0.5°C) generally lie above the climatological pCO₂@14°C cycle. The patterns for pCO₂ anomalies are not as clear. In the STZ, positive pCO₂ and pCO₂@14°C anomalies correspond to negative SST anomalies during spring (October–December). However, during the summer (January–March) the pCO₂ anomalies in the STZ show the opposite behavior of pCO₂@14°C. For summer, the pCO₂ values associated with positive SST anomalies generally occur above the climatological pCO₂ cycle, and values associated with significant negative SST anomalies tend to lie below the climatological pCO₂ cycle.

In the 3 regions, pCO₂@14°C anomalies are well correlated with SST anomalies, with different slopes for the spring-summer and fall-winter periods (Fig. 5). The pCO₂ anomalies in the 3 regions show more scatter than pCO₂@14°C anomalies versus SST anomalies (Fig. 5). This is due to the thermodynamic effect of temperature change that leads to a decrease of the positive pCO₂ anomalies associated with negative SST anomalies, and conversely to an increase of negative pCO₂ anomalies associated with positive SST anomalies. In some extreme cases there is a reversal of the direction of the pCO₂ anomalies, particularly in the STZ (Fig. 5).

The SST anomalies in the STZ, SAZ and CS are roughly consistent in timing and amplitude (Fig. 6), suggesting a similar driver of these anomalies in the three water masses. Large-scale coupled atmosphere-ocean oscillations likely to drive large SST anomalies in the mid-latitude band of the Southern Ocean are ENSO through atmospheric bridges (Li, 2000; Verdy et al., 2006; Morrow et al., 2007), the subtropical dipole pattern (Behera and Yamagata, 2001), or SAM (Hall and Visbeck, 2002; Lovenduski and Gruber, 2005; Morrow et al., 2007).

The patterns of inter-annual variability of SST and wind speed (u_{10}) were investigated south of Tasmania using time-series including years 1982 to 2005 from the Reynolds et al. (2002) monthly SST climatology (Reyn.SmithOIV2, <http://iridl.ldeo.columbia.edu/>) and the Kalnay et al. (1996) National Centers for Environmental Prediction (NCEP) Reanalysis-1 daily u_{10} (<http://www.cdc.noaa.gov/>). Deseasonalized SST and u_{10} reveal strong inter-annual variability for the 1982–2005 period, with low wind speed associated to higher SST (Fig. 7). Such patterns have been previously reported at mid-latitudes of the Southern Ocean in conjunction with changes in SAM (Hall and Visbeck, 2002; Lovenduski

and Gruber, 2005; Morrow et al., 2007). Figure 7 shows that low wind speed years are in general associated with positive phases of SAM, although not as clear during the 1993–1998 period, possibly related to ENSO and SAM interactions (Verdy et al., 2006, 2007).

3.3 Inter-annual air-sea CO₂ flux variations

In order to examine the relationship of SST anomalies to inter-annual variations of air-sea CO₂ fluxes, we lengthen the oceanic pCO₂ time-series to include the years 1982 to 2005 using SST time-series and our previously established relationships between observed SST and pCO₂@14°C from 1991–2003. Wind speed data from the two NCEP grid nodes in the study region (Fig. 1) were averaged and assumed representative of u_{10} over the CS, STZ and SAZ. The Reyn.SmithOIV2 SST grid node n°1 (Fig. 1) was used to compute monthly SST anomalies for the CS. The remaining SST grid nodes (n°2 to n°11, Fig. 1) were averaged, and used to compute monthly SST anomalies for the STZ and SAZ. This assumes that SST anomalies in the STZ and SAZ are synchronous and similar in amplitude and direction as shown in Fig. 6. Also, the position of the STF is highly variable in time (Sokolov and Rintoul, 2002) and it is not possible to arbitrarily associate the SST grid nodes to either STZ or SAZ water masses.

The pCO₂@14°C anomalies were computed from the Reyn.SmithOIV2 monthly SST anomalies using the linear relationships shown in Fig. 5. These pCO₂@14°C anomalies were added to the climatological pCO₂@14°C cycles shown in Figs. 2, 3 and 4. The Reyn.SmithOIV2 monthly SST anomalies were added to the climatological SST cycles shown in Figs. 2, 3 and 4. The pCO₂ values were then computed from pCO₂@14°C and SST (both including the respective anomalies), using the algorithms of Copin-Montégut (1988, 1989). The daily air-sea CO₂ fluxes were calculated using Eq. (2), and the NCEP u_{10} data.

For the whole 1982–2005 period, the CS, STZ and SAZ act annually as sinks for atmospheric CO₂ at the rate of, respectively, -6.4 ± 0.7 , -6.8 ± 0.5 , and -5.7 ± 0.5 mmol C m⁻² d⁻¹. Significant potential inter-annual variability of annual F is apparent in the 3 regions (Figs. 7, 8). Consistent with pCO₂, negative annual F anomalies (stronger sink for atmospheric CO₂) are associated with positive annual SST anomalies, and conversely positive annual F anomalies (weaker sink for atmospheric CO₂) are associated with negative annual SST anomalies for the 3 regions. For positive annual SST anomalies, the strongest F anomalies (Table 2) lead to an increase of the 1982–2005 average annual sink for atmospheric CO₂ of 26, 21, and 59%, in the CS, STZ and SAZ, respectively. For negative annual SST anomalies, the strongest F anomalies (Table 2) lead to a decrease of the 1982–2005 average annual sink for atmospheric CO₂ of 15, 11, and 51%, in the CS, STZ and SAZ, respectively. For positive annual SST anomalies,

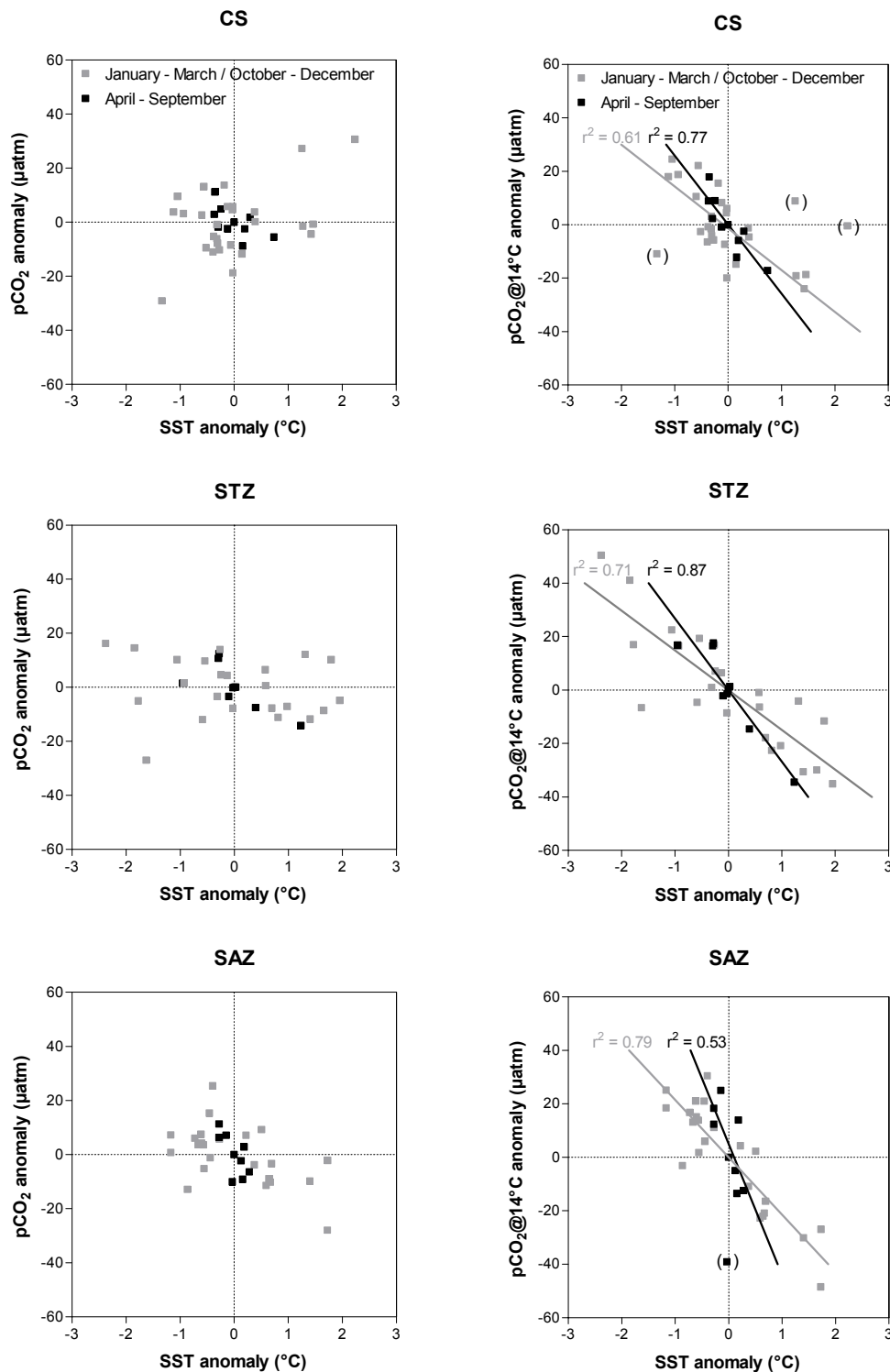


Fig. 5. Monthly anomalies of the partial pressure of CO₂ (pCO₂) and of the pCO₂ normalized to a temperature of 14°C (pCO₂@14°C) plotted against the sea surface temperature (SST) monthly anomalies in the continental shelf (CS), the subtropical zone (STZ) and sub-Antarctic zone (SAZ) south of Tasmania, for the spring-summer period (grey squares) and the fall-winter period (black squares). Solid lines correspond to linear regression functions, and r^2 to the corresponding coefficient of determination (winter-fall in black, spring-summer in grey). Symbols in brackets were excluded from the linear regressions.

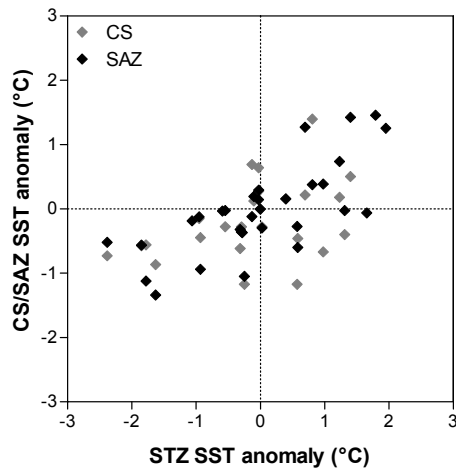


Fig. 6. Sea surface temperature (SST) monthly anomalies in the continental shelf (CS) and subAntarctic zone (SAZ) plotted against the SST monthly anomalies in the subtropical zone (STZ), south of Tasmania.

the average F anomalies (Table 2) cause an increase of the 1982–2005 average annual sink for atmospheric CO₂ of 5, 3, and 21%, in the CS, STZ and SAZ, respectively. For negative annual SST anomalies, the average F anomalies (Table 2) lead to a decrease of the 1982–2005 average annual sink for atmospheric CO₂ of 4, 3, and 18%, for the CS, STZ and SAZ, respectively.

The annual F anomalies computed from the 1982–2005 mean of daily u_{10} data show a consistent negative relationship with the annual SST anomalies for all three regions (Fig. 8). This suggests the scatter in the annual F anomalies computed from daily u_{10} data versus the annual SST anomalies is to a large extent related to inter-annual variability in u_{10} . In particular, the annual F anomalies in 1985, 1989, 1999 and 2001 are smaller compared to other years for the 3 regions, leading to positive annual F anomalies in the CS and the STZ during 1985, 1989 and 1999, and to virtually neutral annual F anomalies in the SAZ during 1985 and 1989.

The F anomalies computed for the fall-winter and spring-summer periods using the 1982–2005 mean of daily u_{10} (Fig. 9) show that the annual F anomalies (Fig. 8) are mainly driven by the fluxes during the fall-winter period. This is related to the thermodynamic effect of temperature change that tends to reduce the magnitude of pCO₂ anomalies, and due to SST anomalies being more marked during the periods of stratification (spring-summer, Figs. 2, 3, 4). Higher wind speeds during the fall-winter period (not shown) are also responsible for the larger contribution of F anomalies during this period. These patterns are reflected in the F anomalies computed for the fall-winter and spring-summer periods using the daily u_{10} , but with more scatter due to inter-annual variability of u_{10} (Fig. 9). The small annual F anomalies in 1985, 1989, 1999 and 2001 (Fig. 8) are due to the spring-

Table 2. The strongest and average annual anomalies of air-sea CO₂ fluxes (F) for positive and negative sea surface temperature (SST) annual anomalies in the continental shelf (CS), the subtropical zone (STZ) and subAntarctic zone (SAZ) south of Tasmania, from 1982 to 2005. Annual anomalies were computed as the difference between the annual mean value and the average of annual mean values of the whole data-set.

	CS	STZ	SAZ
Strongest annual F anomaly (mmol C m ⁻² d ⁻¹)			
Positive SST anomalies	-1.7	-1.4	-3.4
Negative SST anomalies	1.0	0.8	2.9
Average annual F anomaly (mmol C m ⁻² d ⁻¹)			
Positive SST anomalies	-0.3	-0.2	-1.2
Negative SST anomalies	0.3	0.2	1.0

summer F anomalies (Fig. 9). Figure 7 shows that u_{10} in 1985, 1989, 1999 and 2001 was lower than other years. This is most likely related to a poleward shift of westerly winds related to changes in the SAM which causes a local decrease in wind stress and positive SST anomalies (Hartmann and Lo, 1998) and reduced air-sea fluxes of CO₂ (Fig. 7).

4 Discussion

4.1 Causes of the observed inter-annual variability of pCO₂ and SST

SAM, which is characterized by changes in atmospheric mass between 20° S and 90° S is the principal mode of atmospheric forcing and climate variability in the Southern Ocean. SAM induces changes in the strength of westerly winds which induce significant changes in ocean circulation with potential to significantly impact on biogeochemical carbon cycling and air-sea CO₂ fluxes (Le Quéré et al., 2007; Lenton and Matear, 2007; Lovenduski et al., 2007; Verdy et al., 2007). A change in the SAM index leads to synchronous SST anomalies in the STZ and SAZ (Lovenduski and Gruber, 2005; Lovenduski et al., 2007), as observed in our dataset (Figs. 6, 7). An increase in the SAM index is associated with a poleward shift of westerly winds with a local decrease in wind stress (Fig. 7; Hartmann and Lo, 1998; Thompson and Wallace, 2000). This causes an increase in Ekman convergence in the SAZ, with deepening isopycnals, surface warming in the mid-latitudes, and a southward expansion of subtropical waters (Cai et al., 2005; Roemmich et al., 2007). Near Tasmania, surface warming and an increased outflow of warm and salty subtropical waters from the South Tasman Sea have been observed and linked to a polewards shift in the wind stress over the Southern Ocean (Rintoul and Sokolov,

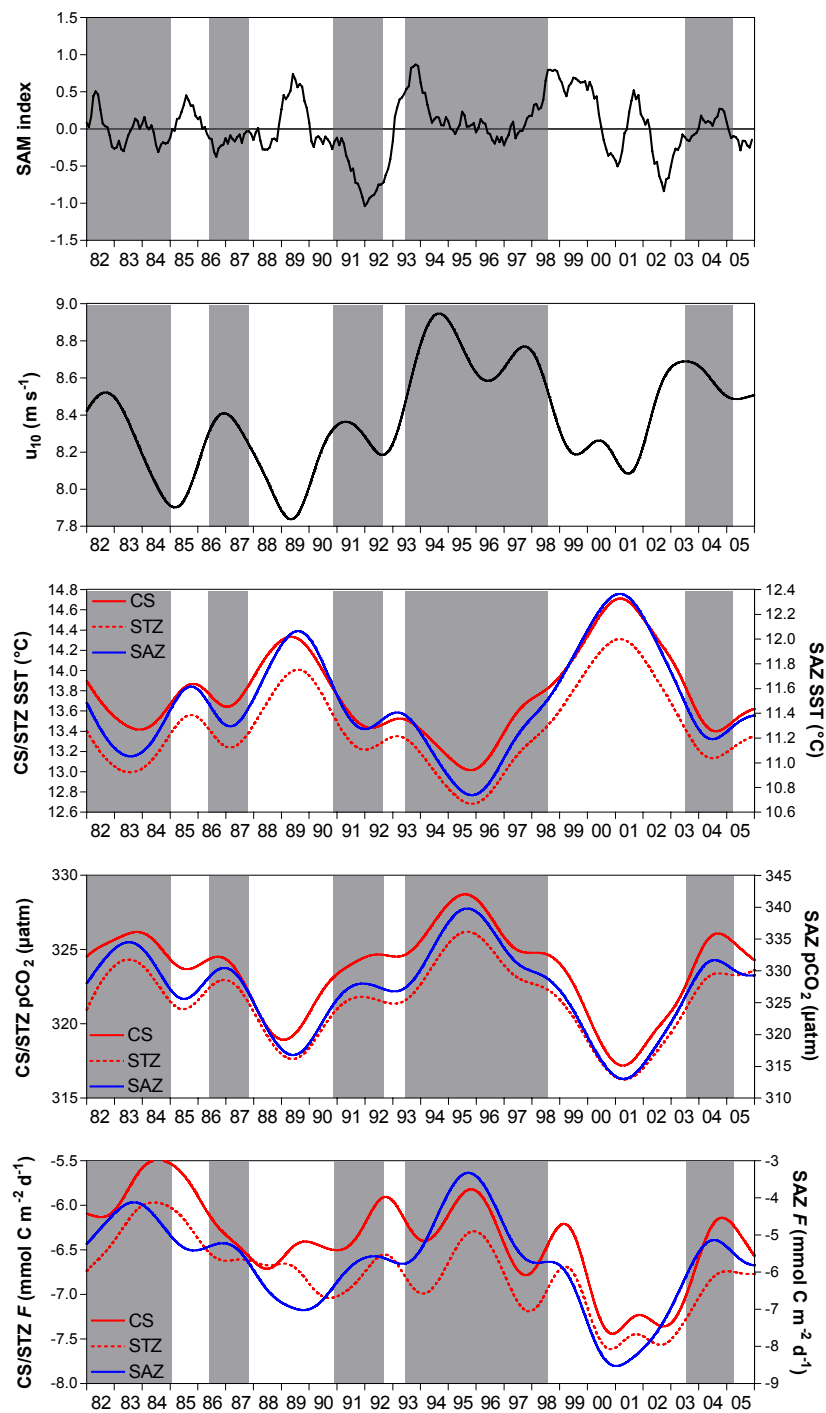


Fig. 7. Time-series from 1982 to 2005 of southern annular mode (SAM) index and of deseasonalized wind speed (u_{10}), sea surface temperature (SST), partial pressure of CO₂ (pCO₂) and air-sea CO₂ fluxes (F) in the continental shelf (CS), the subtropical zone (STZ) and subAntarctic zone (SAZ) south of Tasmania. Grey areas correspond to periods of sustained negative SST anomalies. Signals were deseasonalized using a Gaussian one year low pass filter. SAM index was obtained from the National Oceanic and Atmospheric Administration Climate Prediction Center (<http://www.cpc.ncep.noaa.gov/>) and smoothed with a one year running mean. Details on the computations of SST, pCO₂ and F for the 3 regions are given in Sect. 3.3.

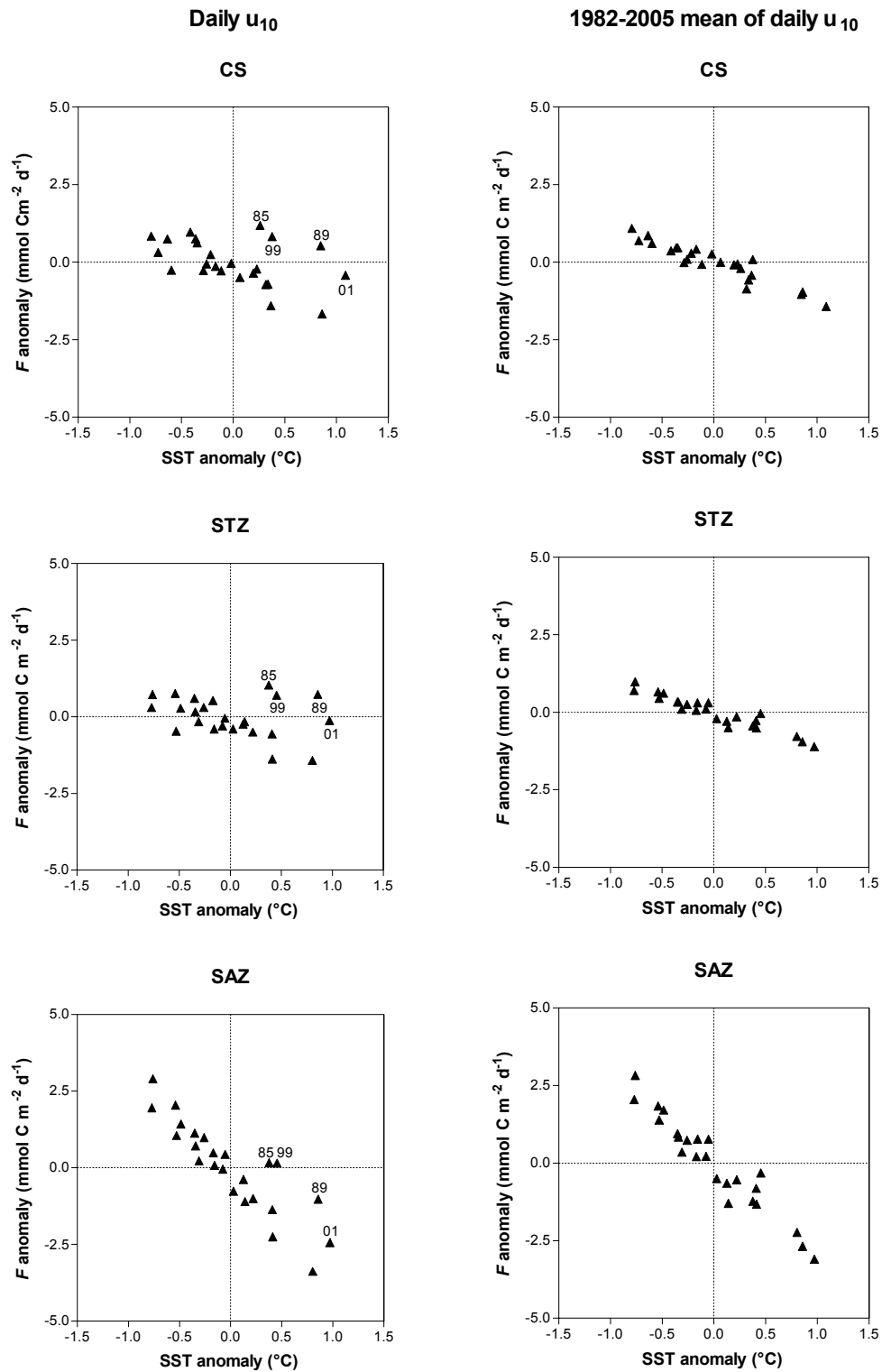


Fig. 8. Annual anomalies of air-sea CO₂ fluxes (F) computed with daily wind speeds (u_{10}) and 1982-2005 mean of daily u_{10} , plotted against the sea surface temperature (SST) annual anomalies in the continental shelf (CS), the subtropical zone (STZ) and subAntarctic zone (SAZ) south of Tasmania, from 1982 to 2005. Annual anomalies were computed as the difference between the annual mean value and the average of annual mean values of the whole data-set.

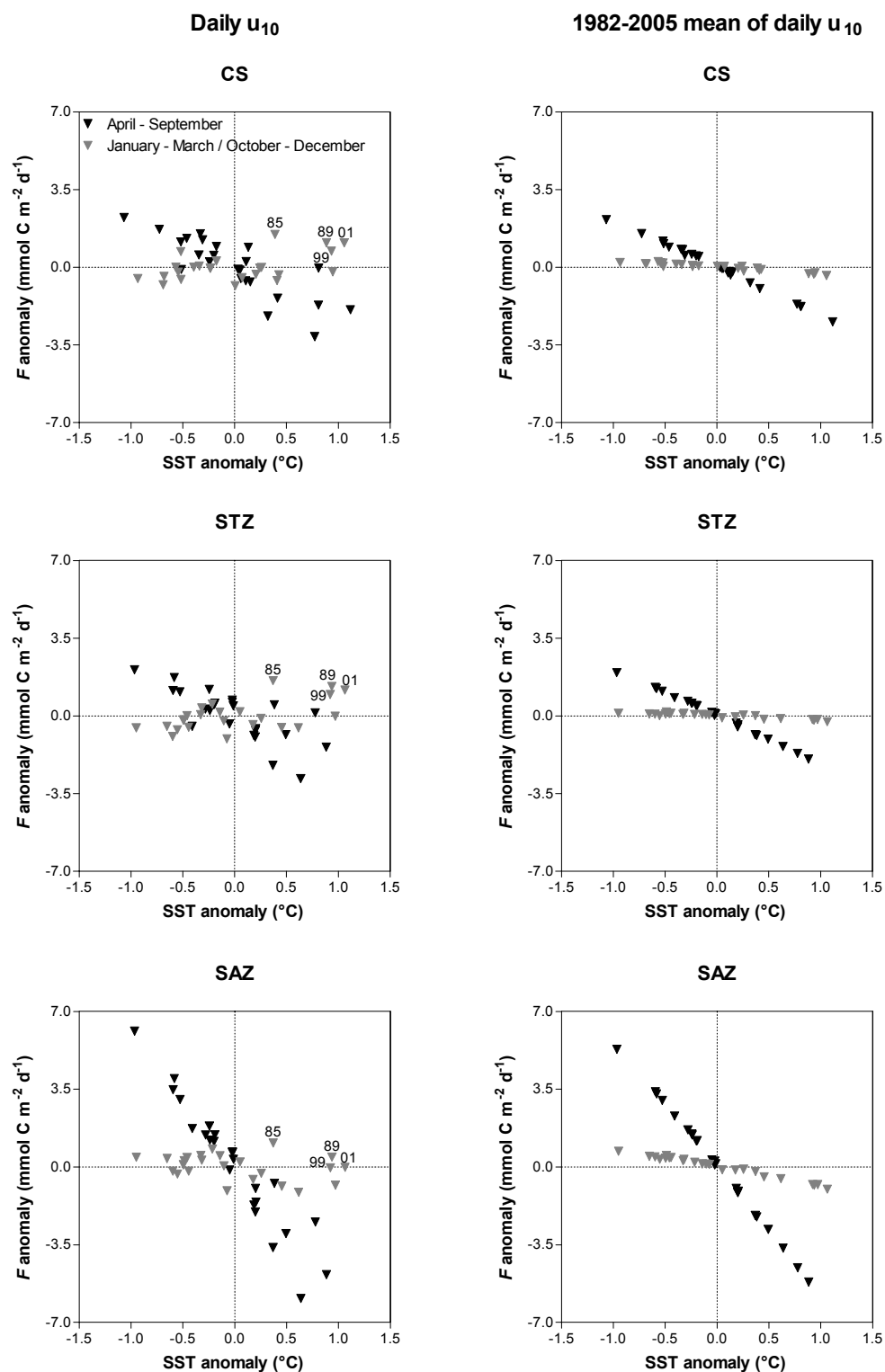


Fig. 9. Anomalies of air-sea CO₂ fluxes (F) computed with daily wind speeds (u_{10}) and 1982–2005 mean of daily u_{10} , plotted against the sea surface temperature (SST) anomalies in the continental shelf (CS), the subtropical zone (STZ) and subAntarctic zone (SAZ) south of Tasmania, for the spring-summer period (grey triangles) and the fall-winter period (black triangles), from 1982 to 2005. Spring-summer (fall-winter) period anomalies were computed as the difference between the spring-summer (fall-winter) period mean value and the average of spring-summer (fall-winter) period mean values of the whole data-set (winter-fall in black, spring-summer in grey).

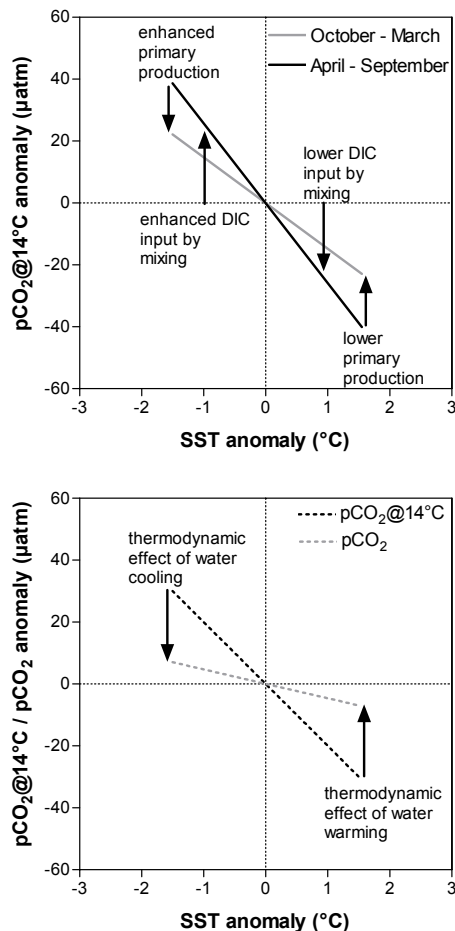


Fig. 10. Conceptual model relating the partial pressure of CO₂ (pCO₂) normalized to a temperature of 14°C (pCO₂@14°C), and pCO₂ anomalies to sea surface temperature (SST) anomalies south of Tasmania.

2001; Morrow et al., 2007; Ridgway, 2007), in agreement with our inter-annual analysis of SST (Fig. 7).

A reduction in available nutrients and reduced DIC input through either the expansion of subtropical waters or less vertical mixing, could explain the observed relationship between pCO₂@14°C and SST anomalies for all 3 regions (Figs. 5, 10). Surface waters of the Tasman Sea have relatively low macronutrient concentrations (Condie and Dunn, 2006) and low pCO₂ values (Takahashi et al., 2002), compared to subAntarctic waters to the south of the STF. The surface warming and decreased wind stress associated with a positive SAM index will also produce a deepening of isopycnal surfaces (Roemmich et al., 2007) and result in more stratified surface mixed layers, thus reducing the potential to entrain subsurface waters with high nutrient and DIC concentrations into the surface mixed layer over winter. The net effect of altered vertical mixing and the southerly expansion of the subtropical waters should be to produce warm SST anomalies and low pCO₂@14°C values in the mid-latitude

waters. Conversely, during periods when the SAM index is reduced, there is likely to be a greater supply of nutrients and DIC to the surface layer through vertical mixing and a retreat of subtropical waters to the north. Under these conditions the surface waters are expected to contain a greater component of relatively high nutrient and DIC waters of the SAZ, producing cool SST anomalies and high pCO₂@14°C anomalies (Fig. 10).

Inter-annual variability in physical forcing related to SAM could also explain different slopes between pCO₂@14°C anomalies and SST anomalies observed during spring-summer and fall-winter periods (Figs. 5, 10). Years with negative SST anomalies would be associated with enhanced nutrient and DIC inputs. Primary production in the region may be partly limited by nutrient availability (Boyd et al., 2001). Years of greater nutrient availability associated with negative SST anomalies would lead to enhanced primary production during spring and summer, and a decrease of the slope (becomes less negative) between pCO₂@14°C and SST anomalies for the fall-winter to spring-summer periods (Fig. 10). The overall negative relationship between pCO₂@14°C anomalies and SST anomalies (Fig. 5) shows that the enhanced primary production associated with negative SST anomalies does not overcome the enhanced winter inputs of DIC.

4.2 Inter-annual variability indicative of future change and potential feedback on increasing atmospheric CO₂?

Examining the relationship between anomalies in F and SST on inter-annual time scales provides insights of biogeochemical responses to projected long-term trends in SST and stratification of the Southern Ocean. Based on correlations of remotely sensed chlorophyll-*a* and SST, Le Quééré et al. (2002) suggested that an increase of SST would lead to an overall increase in primary production in the Southern Ocean. This was attributed to stronger light limitation compared to nutrient limitation for most regions of the Southern Ocean and supported by the simulation of propagating warm anomalies using a biogeochemical ocean general circulation model (Le Quééré et al., 2003). Bopp et al. (2001) have also suggested that export production in the Southern Ocean will increase with projected warming of surface waters. The comparison of pCO₂@14°C and SST anomalies during fall-winter and spring-summer periods (Fig. 5) suggests that in our study region an increase of SST leads to a decrease of export production (Fig. 10). This is in agreement with evidence for nutrient rather than of light limitation of primary production for the sub-polar waters near Tasmania (Boyd et al., 2001).

Besides modifying export production, warming of surface waters will also change air-sea CO₂ fluxes due to changes of the input of CO₂ from deeper layers. Indeed, numerical models of CO₂ dynamics and air-sea CO₂ exchange in the Southern Ocean highlight the very significant role of vertical mixing at seasonal (Louanchi et al., 1996; Metzl et al.,

1999, 2006) and inter-annual (Louanchi and Hoppema, 2000; Verdy et al., 2007) time scales. Changes in vertical mixing in the Southern Ocean can have an impact on the intensity of the air-sea CO₂ fluxes and can potentially modulate the atmospheric CO₂ content by changing the natural biogeochemical CO₂ cycle (Matear and Hirst, 1999). Indeed, to explain the decrease of atmospheric CO₂ during the last glacial period, some hypotheses are based on a decrease in the ventilation of deep DIC rich water in relation to an increase of stratification (Toggweiler, 1999; Sigman and Boyle, 2000), due to a northward shift of the westerly winds (Toggweiler et al., 2006) or due to an increase of salinity in deeper waters (Watson and Naveira Garabato, 2006). The decrease of vertical inputs of DIC at high latitudes combined with constant or enhanced export production could have increased the biological pump (in regions of light limitation for primary production) and enhanced the atmospheric CO₂ sink (Toggweiler, 1999; Sigman and Boyle, 2000; Toggweiler et al., 2006; Watson and Naveira Garabato, 2006). This clearly illustrates that a modification of the natural biogeochemical cycle of CO₂ in the Southern Ocean can lead to strong feedbacks on the atmospheric CO₂ content.

Our study region is small but includes a portion of two important waters masses: the STZ and SAZ. Our data-based inter-annual variability analysis does indicate that a decrease of vertical input of DIC during winter due to positive SST anomalies could lead to an increase of the sink for atmospheric CO₂. This suggests that warming of surface waters of some regions of the Southern Ocean, may lead to a negative feedback on increasing atmospheric CO₂.

5 Conclusions

We compiled data obtained along 40 transects during 22 cruises carried out between 1991 and 2003, on the CS and adjacent open oceanic waters south of Tasmania. This allowed us to analyze inter-annual variations of pCO₂ and air-sea CO₂ exchange and examine how they vary in relation to warm and cool anomalies. Strong SST anomalies up to 2°C were observed in the STZ and the SAZ. As the waters on the CS are a mixture of STZ and SAZ, these SST anomalies also propagate onto the CS. The consistency in timing and amplitude of SST anomalies in the STZ and SAZ can only be attributed to a large scale coupled atmosphere-ocean oscillation.

Overall, positive SST anomalies are associated with negative pCO₂@14°C anomalies, and negative SST anomalies with positive pCO₂@14°C anomalies, in the CS, STZ and SAZ. This seems to be related to a reduced input of DIC during the fall-winter period, during the phases of positive SST anomalies, in relation to a poleward shift of the westerly winds, and a local decrease in wind stress.

The potential effect of SST anomalies on air-sea CO₂ exchange were investigated using a 23 years consistent time-

series of SST and wind speed. The general pattern is an increase in the sink for atmospheric CO₂ associated with positive SST anomalies, mainly during the fall-winter period. However, this general pattern is strongly modulated by inter-annual variations of wind speed that affects the gas transfer velocity and the intensity of the air-sea CO₂ flux.

Assuming that phases of positive SST anomalies are indicative of the future evolution of ocean biogeochemistry under global warming, we show based on a spatially restricted observational data-set, that some provinces of the Southern Ocean could provide a potential negative feedback on increasing atmospheric CO₂ and associated climate change. The observations from our region are in agreement with recent modelling studies that show during positive phases of SAM a decrease of the CO₂ sink in high latitude areas of the Southern Ocean due to enhanced upwelling and an increase of the CO₂ sink in the low latitude areas of the Southern Ocean (Le Quéré et al., 2007; Lenton and Matear, 2007; Lovenduski et al., 2007). It is clear that larger scale investigation in the Southern Ocean is required to quantify more rigorously potential feedbacks on the increase of atmospheric CO₂ due to SST changes in surface waters.

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