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L. P. Borovkov, B. V. Kozelov, L. S. Yevlashin, S. A. Chernouss. Variations of auroral hydrogen emission near substorm onset. *Annales Geophysicae*, 2005, 23 (5), pp.1623-1635. hal-00317815

**HAL Id: hal-00317815**

**<https://hal.science/hal-00317815>**

Submitted on 18 Jun 2008

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# Variations of auroral hydrogen emission near substorm onset

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Received: 17 February 2005 – Revised: 9 May 2005 – Accepted: 19 May 2005 – Published: 27 July 2005

Part of Special Issue “Atmospheric studies by optical methods”

**Abstract.** The results of coordinated optical ground-based observations of the auroral substorm on 26 March 2004 in the Kola Peninsula are described. Imaging spectrograph data with high spectral and temporal resolution recorded the Doppler profile of the  $H\alpha$  hydrogen emission; this allows us to estimate the average energy of precipitating protons and the emission intensity of the hydrogen Balmer line. Two different populations of precipitating protons were observed during an auroral substorm. The first of these is associated with a diffuse hydrogen emission that is usually observed in the evening sector of the auroral oval and located equatorward of the discrete electron arcs associated with substorm onset. The average energy of the protons during this precipitation was  $\sim 20\text{--}35$  keV, and the energy flux was  $\sim 3 \times 10^{-4}$  Joule/m<sup>2</sup>s. The second proton population was observed 1–2 min after the breakup during 4–5 min of the expansion phase of substorm into the zone of bright, discrete auroral structures (N-S arcs). The average energy of the protons in this population was  $\sim 60$  keV, and the energy flux was  $\sim 2.2 \times 10^{-3}$  Joule/m<sup>2</sup>s. The observed spatial structure of hydrogen emission is additional evidence of the higher energy of precipitated protons in the second population, relative to the protons in the diffuse aurora. We believe that the most probable mechanism of precipitation of the second population protons was pitch-angle scattering of particles due to non-adiabatic motion in the region of local dipolarization near the equatorial plane.

**Keywords.** Auroral ionosphere; Particle precipitation; Storms and substorms

## 1 Introduction

The hydrogen Balmer lines of  $H\alpha$  (656.3 nm) and  $H\beta$  (486.1 nm) were first identified in aurora spectra in Norway (Vegard, 1939). Doppler broadening of the line profiles indicated their extraterrestrial origin. More extensive studies

of the hydrogen emissions started during the International Geophysical Year (IGY). Almost simultaneously, observations of hydrogen emissions drew the interest of many researchers over the world: the USA, Canada, Sweden, Norway, Japan and some other countries (see review in Eather (1967)). In the Soviet Union, the pioneering studies of hydrogen emissions were published by Galperin (1959) and Yevlashin (1961). During next four decades, numerous studies featured the relationship of hydrogen emissions with various forms and types of auroras, latitudinal and longitudinal expansion of the hydrogen emission zone, dependence of hydrogen emissions on magnetic activity, and the relationship with ionospheric ionization, riometric absorption and ULF-emission, radio reflections from auroras. There the altitude of hydrogen emissions was estimated using sounding rockets and the observation of the absolute values of  $H\alpha$  and  $H\beta$  intensity was performed, as well as measurements of the Balmer decrement, emission fluctuations and, finally, the relationship of hydrogen emissions to auroral breakup and substorms in general (see a review by Yevlashin (2000)).

The main patterns of substorm development in proton auroras were described already in the 1970s in a few papers (Montbriand, 1971; Fukunishi, 1975; Yevlashin and Yevlashina, 1980). If one carefully compares results of these and others papers, one can conclude that they all coincide qualitatively in their basic aspects. So, at the growth phase of substorm in the evening, the proton aurora is found to be at lower latitudes than the discrete arcs of the electron aurora, whereas the situation is opposite in the recovery stage in the morning hours. These conclusions have been supported by statistics (Hardy et al., 1991).

As far as the expansion phase is concerned, some important quantitative differences can be found. In the pattern by Fukunishi, the enhancement of proton luminosity brightness is mentioned; by contrast, the paper by Yevlashin and Yevlashina reports the general decrease of proton aurora intensity, while the paper by Montbriand draws attention to a flare of proton aurora at the initial moment of the expansion phase with its subsequent decrease down to the values less than ones at the growth phase. The possible reasons

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of the differences were, first, the different techniques used for registration of hydrogen emissions (and therefore, for example, different exposure time) and secondly, difficulties in the interpretation of experimental data, due to strong intensification of electron-excited emissions ( $N_2$  1PG and Vegard-Kaplan bands), that blend the hydrogen spectral lines during the expansion phase of substorm (Gattinger and Vallance Jones, 1974).

Extensive studies of substorm events during the last two decades have revealed the important role of ions fluxes in the development of instabilities associated with the expansion phase (Delcourt et al., 1994, 1996). The significant variations of proton fluxes found by satellites in the zone of substorm expansion, both near magnetic equator (Sauvaud and Winckler, 1980; Kozelova et al., 2004) and at low altitudes (Sergeev and Kubyshkina, 1996), also attract one's attention. Despite the great achievements of satellite techniques, the direct observation of proton fluxes in various magnetospheric domains is usually lost in the space-time ambiguity associated with satellite data. Therefore, the development of optical observations, both ground and satellite, is continuing (see review in Lanchester et al. (2003)). However, the dynamics of precipitated proton fluxes and the resulting hydrogen emissions during the late growth and early expansion phases of the substorm are not completely understood.

Using global observations of electron and proton produced auroras observed by IMAGE satellite FUV instruments, Mende et al. (2001) concluded that the peak intensity of the proton aurora did not change substantially during substorms. They found that the initial brightening of the substorm was embedded in proton precipitation, which coincides with the results of Samson et al. (1992). Gérard et al. (2004) supports the results by large statistics. As the electron surge expands poleward it leaves the protons behind. Deehr and Lummerzheim (2001), while using a combination of a high sensitivity all-sky imager and meridian scanning photometers, studied emission parameters in 33 auroral substorms in 1992 with proton auroras at Alaska stations. They do not support the argument that the substorm onset begins from within the proton precipitation, but the discovery of the soft electron pulse during "auroral fading" does support theories requiring a pulse of Alfvén speed electrons. Takahashi and Fukunishi (2001) studied auroral breakup events at Syowa station, Antarctica for 35 events of auroral substorms during the period of 1990–1993. The authors believe that at the poleward boundary of auroral bulge the Doppler shift of hydrogen line is quite large, while the maximum intensity of hydrogen emission is observed somewhat equatorward of the leading edge. These facts imply that precipitating protons with higher energies and highly field-aligned pitch angles are produced at the poleward boundary through a plasma injection process in the magnetosphere.

The goal of this paper is a detailed study of the proton precipitation during late growth and early expansion phases of an auroral substorm on 26 March 2004, using the data of coordinated optical ground-based observations in the Kola Peninsula. In addition to the usual observatory equipment

(all-sky TV camera and scanning photometer), a special imaging spectrograph with high spectral and temporal resolution was used. The data from this device allows us to estimate variations of the energy of precipitating proton flux and the hydrogen emission intensity, using the Doppler profile of the auroral  $H\alpha$  emission during the event. An interpretation of optical observations, as well as any indirect method of diagnostics, is usually based on a theoretical model.

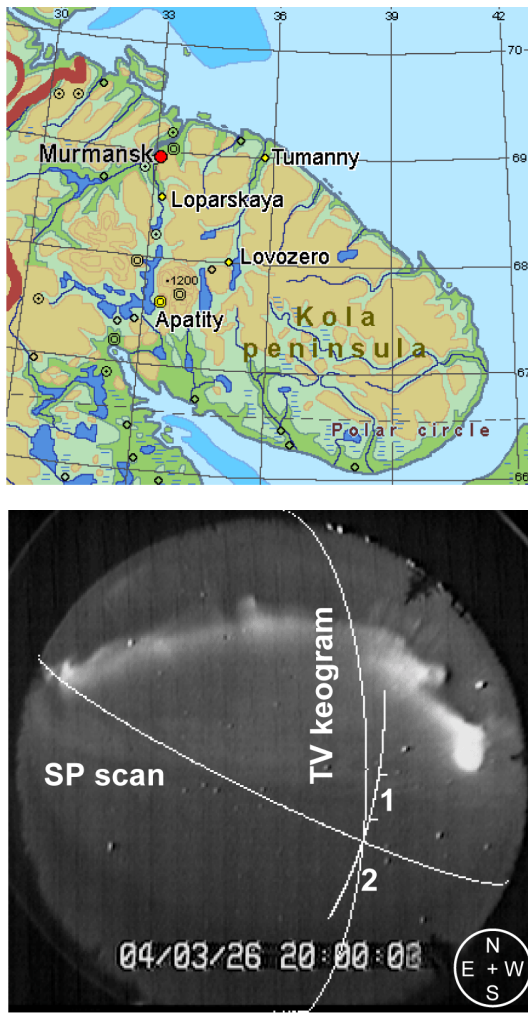
Our estimations of the proton-hydrogen flux parameters are based on the results of simulations of the Doppler-shifted hydrogen line profiles using the Monte-Carlo method (Kozelov, 1993, 1994). The simplified model from Kozelov et al. (1995), based on results of detailed Monte-Carlo transport modeling, is used to calculate the spatial structure of the  $H\alpha$  luminosity resulting from such precipitation, for different values of precipitated proton flux parameters. The observed spatial structure of the  $H\alpha$  luminosity supports our conclusion about the characteristic energy of proton precipitation before and after substorm onset. Relations between the dynamics of proton flux observed during the event of interest and the protons' motion in the region of local dipolarization near the equatorial plane are discussed.

## 2 Equipment

This study is based on observations provided by the Polar Geophysical Institute at two points: Lovozero observatory (67.97 N, 35.02 E) and Apatity station (67.58 N, 33.31 E). Positions of the observation points are shown on the map in Fig. 1a. During the evening of 26 March 2004, two optical devices were operating at Lovozero observatory: TV All-Sky Camera (TVASC) and Scanning Photometer (SP). The TVASC observed the aurora in "white" light with a broad maximum at blue-green wavelengths. The TV data were recorded by a VHS recorder in a PAL video system, at 25 images per second. The TV data were also digitized using an Acorn TV-Capture plate with an output resolution of 5 images per second,  $320 \times 288$  pixels, 8 bits per pixel.

The Scanning Photometer (SP) was scanning approximately in the magnetic east-west direction and monitoring auroral emissions at the wavelengths 427.8 nm ( $N_2^+$  1NG), 557.7 nm (OI) and 630.0 nm (OI). The SP was calibrated by calibration lamp TRS-28-50, with an estimated discrepancy of emission intensity in the 630.0-nm channel of 25%. The trajectory of SP scans in the field of view of the TVASC is shown in Fig. 1b. The magnetometer and riometer at Lovozero observatory was also available to monitor the geophysical environment.

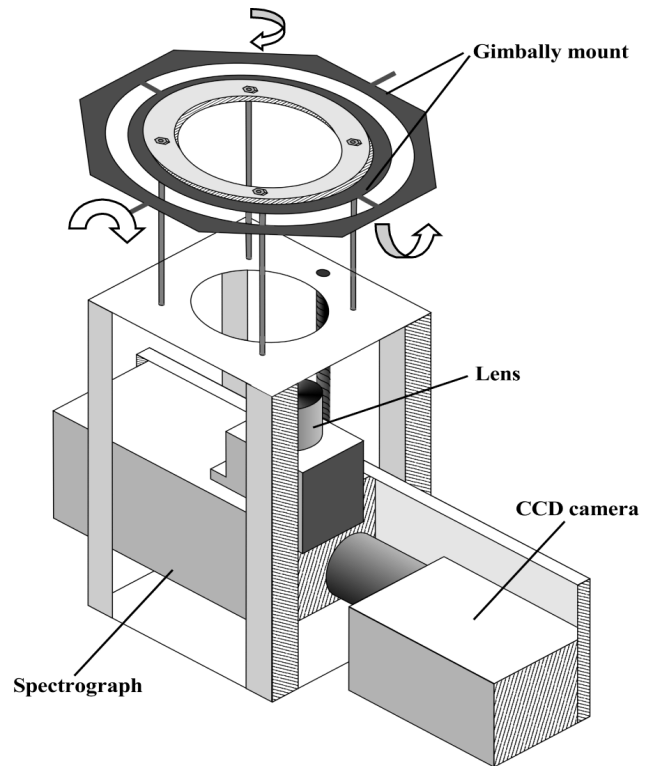
During the evening of 26 March 2004, at the Apatity station, spectrograph observations of auroras were provided by a special device consisting of an imaging spectrograph SpectraPro 306, made by Acton Research Corp. and an intensified CCD camera IpenaMax, made by Princeton Instruments. The specifications of the spectrograph and CCD camera are the following:



**Fig. 1.** Geometry of observations: (a) observation points; (b) fields of view for optical devices.

#### Spectrograph SpectraPro 306

- Optical design: Imaging Czerny-Turner with polished aspherical mirrors
- Focal length: 300 mm
- Aperture Ratio: f/4
- Grating with profiled grooves, 1200 gr/mm, 68×68 mm
- Resolution: 0.1 nm at 435.8 nm with 100 μm entrance slit
- Dispersion: 2.7 nm/mm
- Focal plane size: 27 mm wide×14 mm high
- Free spectral interval ~25 nm with IPentamax camera.

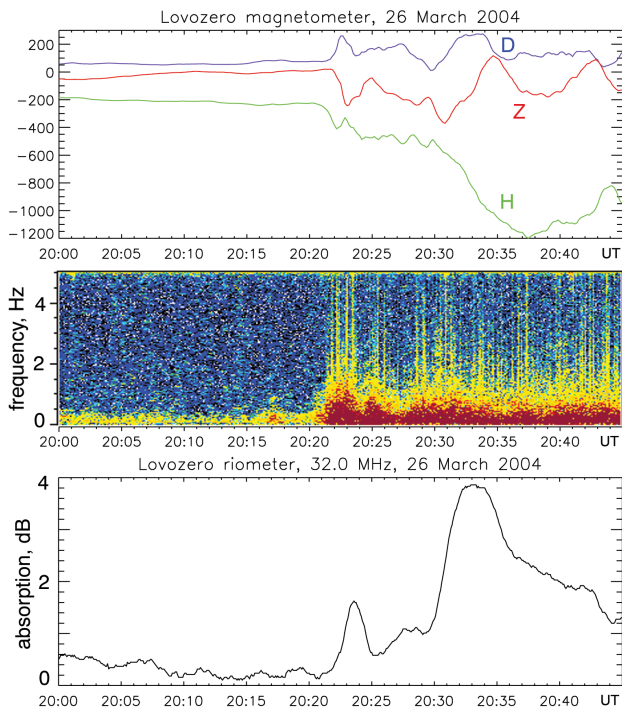


**Fig. 2.** Schematic of the spectrograph SpectraPro 306 with CCD camera IPentaMax.

#### CCD camera IPentaMax

- Thermoelectrically cooled CCD matrix 512×1024, 15×15 μm pixels with frame transfer
- Microchannel image intensifier fiber-optically coupled to CCD
- 12 bits 5 MHz ADC converter
- Dark current ~20 electrons  $\text{pix}^{-1}\text{s}^{-1}$  at  $-20^\circ\text{C}$
- Exposure range 50 μs to 23 h
- Computer controls camera and acquires data via high speed serial connection.

The CCD sensor was cooled up to  $-45^\circ\text{C}$  to reduce the noise level. The resulting mean dark charge for the 10-s exposition was about 4 units of the device ADC (Analog to Digit Converter). The device is mounted by gimbal to the ceiling and obtains emissions through a plexiglass dome. The schematic view of the device is presented in Fig. 2. We can tilt the device's optical axis up to 45 deg relative to the vertical direction in 2 orthogonal planes. In addition, it is possible to turn the device (and the entrance slit) relative to the optical axis in any direction. With a Nikkor fisheye lens, the full aperture of the device is  $\sim 80 \times 1$  deg. We control the spectrograph and camera with our own original software, using the



**Fig. 3.** Substorm event of 26 March 2004: (a) magnetogram at Lovozero observatory; (b) Lovozero magnetometer spectrogram (Y component); (c) riometric absorption observed by Lovozero riometer.

low level programming possibilities of the software obtained with both instruments.

The field of view of the spectrograph is shown in Fig. 1b. The SP scan crosses the spectrograph's field of view (see point 2 in Fig. 1b), and this allows us to calibrate the spectrograph sensitivity relative to a 630.0-nm channel of the SP using the 636.4-nm emission line. According to Chamberlain (1961) the ratio of these emissions is approximately equal to 3. For calibration we used the diffuse auroral emissions of 630.0 nm in the range 0.3–2.5 kR.

### 3 Observations

For a detailed study we selected the interval 20:00–20:45 UT, on 26 March 2004. The observations are presented in Figs. 3–5. Figure 3 shows three components of the geomagnetic field, magnetic pulsations and riometric absorption obtained at Lovozero observatory. Figure 4 illustrates the results of the optical observation by the SP and the TVASC at the same points.

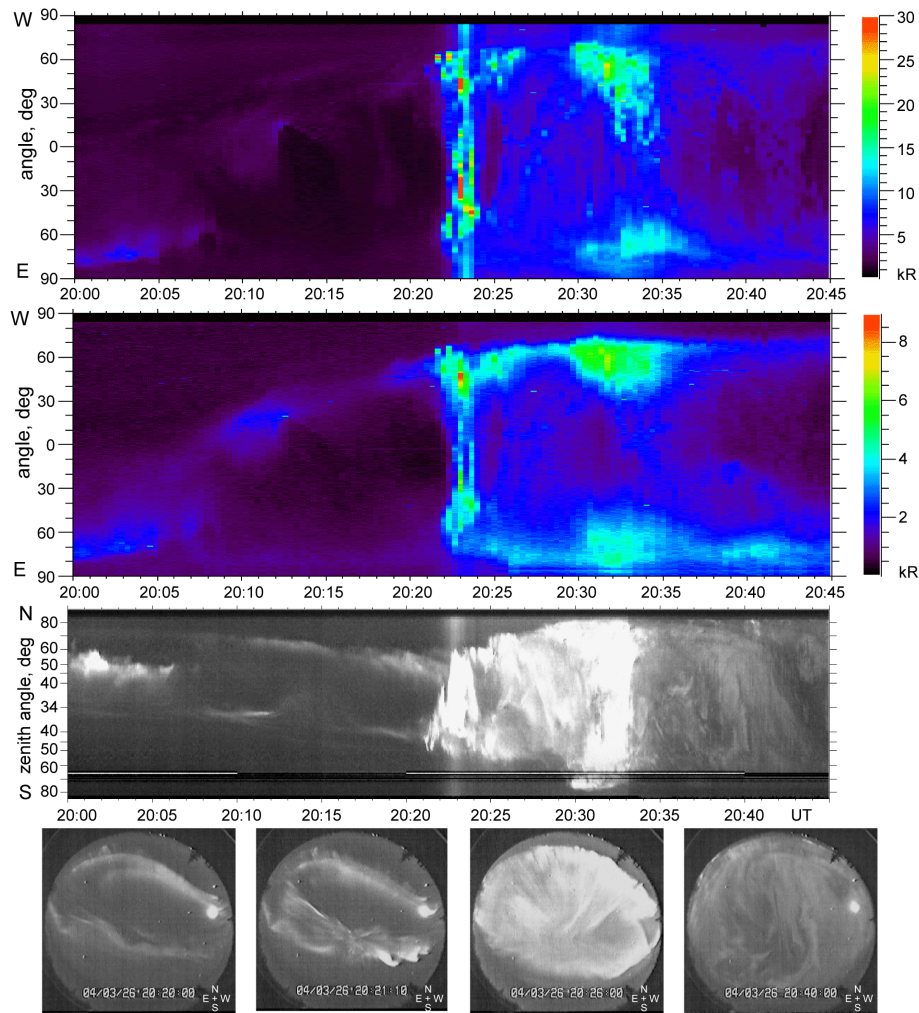
Figure 5 contains the spectrograph data from Apatity station. There is a gap in the data for  $\sim 30$  s from 20:23:35 UT to 20:24:05 UT, due to file saving. The data has both spectrometric and spatial information about the aurora luminosity. The time exposure of the spectrograph data was  $\sim 10$  s. Under quite, pre-breakup conditions, the mean signal level at  $H\alpha$  profile maximum was  $\sim 4$  units of ADC (Analog to

Digit Converter), that means the signal-to-noise ratio at raw data was  $\sim 1$ . The top plate of Fig. 5 presents an evolution of the auroral spectrum in the spectral region from 635.5 to 661.0 nm. By summing 3 frames of spectral images and taking an average of 100 rows of image, we improve the signal-to-noise ratio by  $\sqrt{3 \times 100} \approx 17.3$  times. So, the statistical error in the data is less than 6%. The averaged rows in the spectral image corresponds to the range of view angles from zenith to  $\sim 18$  northward degrees, marked in Fig. 1b by the number “1”.

The shape of the spectrum at several important moments is shown in the second plate. Two spectral regions are marked at the spectrum with “A” and “B”. The spectral region of 655.7–656.2 nm is mainly filled with the Doppler-shifted  $H\alpha$  hydrogen line, however, some contribution from emissions of atmospheric species is also possible, especially for strong electron precipitation (Gattinger and Vallance Jones, 1974). The second spectral region is formed only by atmospheric emissions, mostly by a line from  $N_2$ IPG system. Spatial evolution of the auroral luminosity in these spectral regions is illustrated in two plates with keograms in the north-south direction.

The interval under consideration started from the late growth phase of the substorm. According to TVASC data, in the beginning of the interval, an active bright arc was located on the north side of the all-sky image ( $\sim 130$  km from Lovozero), and a slow diffuse arc was observed near the zenith; see Fig. 1b. During the next 20 min there was a slow evolution of the arc system: motion of the first arc to the zenith of TVASC, occurrence of a new arc at the previous position of the first arc, intensification of the zenith arc ( $\sim 20:07$  UT), and then dissipation at  $\sim 20:12$  UT and the next intensification at  $\sim 20:17$  UT. All this activity was accompanied by large-scale motion of the aurora in the southward direction; this is a normal feature for the growth phase of the substorm (Akasofu, 1968). However, remarkable variations in the magnetic field and riometer absorption were not observed; see Fig. 3. During this interval, the  $H\alpha$  hydrogen line with Doppler shift to the blue side was observed at all view angles of spectrograph (see green spectra in Fig. 5b), that indicates a wide zone of the proton precipitation. Figure 5c shows that in the northern part of the view angles the intensity of the hydrogen line was gradually decreasing, but in the southern part, it was increasing. Therefore, we can deduce that the proton precipitation zone was moving towards the south, too. The estimated width of the diffuse zone of the hydrogen emission was  $\sim 200$  km. The slow electron arc mentioned above was also the spectrograph's field of view. The evolution of the discrete arc intensity is shown by  $N_2$ IPG emission (659.3–659.8 nm emission band, marked in Fig. 5b as “B”); see Fig. 5d. However, the arc had no influence on the spectral region of blue-shifted  $H\alpha$  emission; see Fig. 5c.

An intensification of the near-zenith arc, which started from 20:21 UT simultaneously with the increase of the magnetic pulsations (Fig. 3b), developed into the breakup observed at 20:22 UT. The expansion phase of the substorm



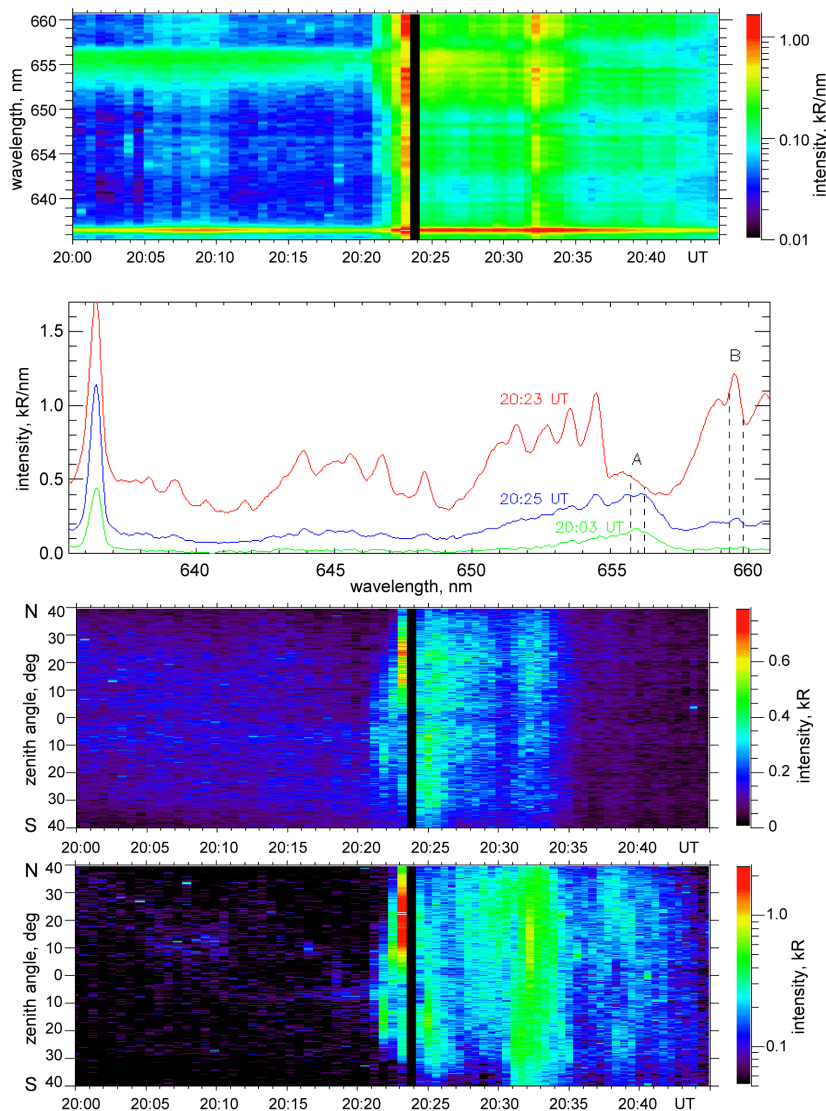
**Fig. 4.** Optical observations at Lovozero: **(a)** data of east-west Scanning Photometer (SP) in 427.8 nm emission; **(b)** SP data in 630.0 nm emission; **(c)** TV keogram in north-south direction (see Fig. 1); **(d)** TV all-sky images at several moments.

proceeded in two steps shown by a drop in the H-component of the magnetic field (Fig. 3a), indicating the development of the magnetosphere-ionosphere current system. Each drop corresponds to an increase in riometer absorption (Fig. 3c) and an intensification of optical aurora. These two steps corresponded to the aurora expansion in the north-south direction and are obvious in the TVASC keogram, Fig. 4. The longitude region of the auroral event was located between the positions of two LANL satellites, 01A and 02A. The satellite data (not presented here) show the appearance of two injections of high-energy particles at  $\sim 20:21$  UT and  $\sim 20:28$  UT.

During the expansion phase of the substorm, the H $\alpha$  hydrogen line was blended with very strong emissions of the excited atmospheric species (see red spectra in Fig. 5b), mainly N<sub>2</sub> 1PG (Gattinger and Vallance Jones, 1974). Therefore, we can deduce that strong electron precipitation occurred. However, after the first intensification of the electron precipitation during several minutes (20:24–20:29 UT) one can also see an increased hydrogen emission (see blue spec-

tra in Fig. 5b). This intensification of the hydrogen emission corresponds to several patches at northward view angles in spatial distribution in Fig. 5c. This structure is absent in Fig. 5d, therefore, this is not an intensification of the blended emissions excited by electron precipitation. Simultaneously with this intensification of the hydrogen emission, N-S oriented arcs appeared on the north side of TVASC images; see Fig. 4, image at 20:26:00 UT. After 20:35 UT the hydrogen emission is absent, all other aurora emissions are decreasing, TVASC images were filled with pulsating aurora structures, which is a normal picture for the recovery phase of substorm.

During the event of interest we observed two “populations” of precipitated protons. One of them produced a diffuse hydrogen zone during the preliminary phase of the substorm. Another one was observed during the expansion phase; it had a more sharp and active precipitation region. What difference exists between these populations and the mechanisms of their precipitation? To answer the question in the next sections we will try to separate the hydrogen



**Fig. 5.** Spectrograph data at Apatity: **(a)** evolution of spectrum averaged in  $0^{\circ}$ – $18^{\circ}$  North zenith angles (the region marked by “1” in Fig. 1); **(b)** spectra at several moments; **(c)** north-south keogram in “A” emission band (655.7–656.2 nm, mainly Doppler-shifted  $H\alpha$ ); **(d)** north-south keogram in “B” emission band (659.3–659.8 nm,  $N_2$  1PG).

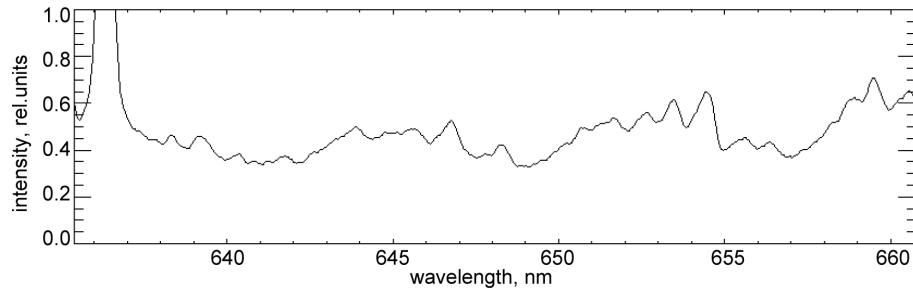
emission from blended spectra and to estimate the energy of precipitated protons.

#### 4 Separation of the $H\alpha$ profile

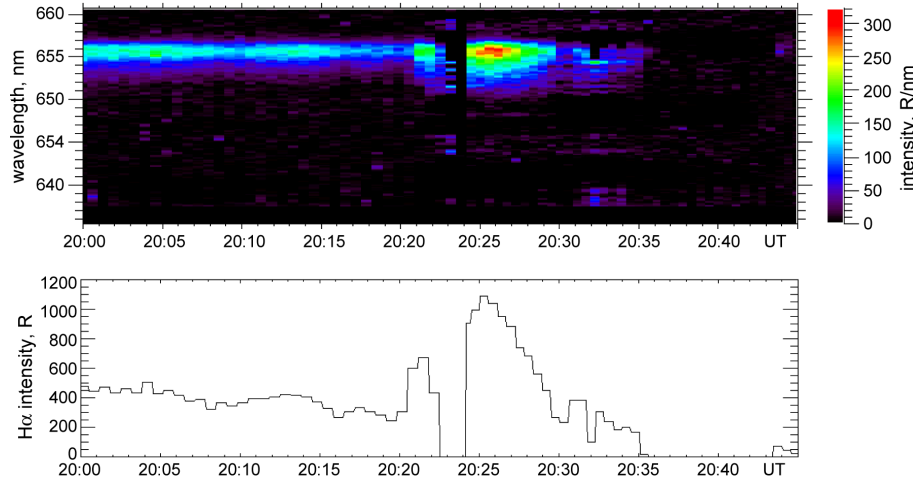
The main problem for detection of the  $H\alpha$  Doppler-shifted line is blending by numerous other emissions (Gattinger and Vallance Jones, 1974). Theoretical models of the auroral emission spectrum are not available with the quality which would allow one to separate the shape of the  $H\alpha$  line from the experimental spectrum with sufficient precision. Therefore, here we use a semiempirical method based on experimental data itself.

We will assume that the spectrum of all emissions in the range of 640–660 nm, excluding  $H\alpha$ , has the same shape,

and that relations between intensities of different lines in the range do not depend on characteristics of incoming particle flux. This assumption has a theoretical substantiation. The auroral spectrum near  $H\alpha$  during nighttime is formed mainly by lines of a 1-PG band emitted from the  $B^3\Pi_g$  state of molecular nitrogen. The efficiency (so-called “energetic cost”) of the state excitation by electrons does not depend on the electron energy of the auroral range and it does not depend on altitude (from 100 to 200 km); see Sergienko and Ivanov (1993). The excitation rate of the  $B^3\Pi_g$  state of  $N_2$  in a volume unit is directly proportional to the energy deposition rate by electron flux in this volume unit. Integration along the line of sight does not change this proportion, independent of the geometry (shape) of the electron precipitation and the direction of the line of sight. The lifetime of the  $B^3\Pi_g$  state is too short to have a remarkable quenching at



**Fig. 6.** Average spectrum of electron excited lines in the spectral range of 636.0–661.0 nm.



**Fig. 7.** Evolution of separated  $H\alpha$  hydrogen emission: (a) spectra after removing the background emissions; (b) integral intensity in the range of 651–657 nm.

the auroral altitudes (Kirillov, 2004). So, all these excitations of the state lead to emission of 1-PG band lines. The remaining problem is relations between intensities of different rotational lines. Ordinarily, these relations must depend on the local atmospheric temperature (Gattinger and Vallance Jones, 1974). The temperature changes dramatically in changing precipitation energy. However, for weak incoming fluxes of electrons and protons, we make the assumption that the temperature dependence of the spectrum shape is possible weak.

To obtain the shape of blended emission spectrum, we calculate an average normalized spectrum during the time interval 20:38–20:45 UT, when the shifted  $H\alpha$  emission was absent. The integral intensity in 642–647 nm has been used for normalization. The average normalized spectrum is shown in Fig. 6. Then, using the normalized spectrum and integral intensity in 642–647 nm for each spectrum in the considered event in order to obtain the magnitude, we then remove the blended emissions from spectrograph data. The results are presented in Fig. 7a. One can clearly see the  $H\alpha$  emission line in spectra from 20:00 to 20:21 UT, where blended emissions were fully removed in the spectra.

However, the residual emissions are clearly seen during an interval of strong electron precipitation, accompanied by

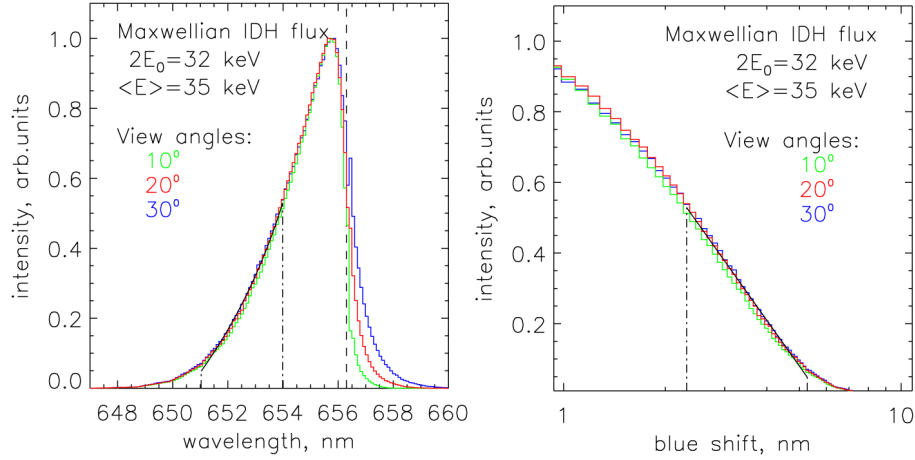
increased riometer absorption. Obviously, this is a result of dependencies not taken into account in our separation procedure. It is interesting that after removal of the blending emissions, the  $H\alpha$  emission is practically absent in the spectrum at 20:23 UT, which corresponds to the strongest aurora emission during the breakup of the aurora arc.

The total intensity of the  $H\alpha$  emission has been estimated by the integration of the spectra from Fig. 7a. Figure 7b illustrates the temporal evolution of the integral intensity of the separated  $H\alpha$  spectra in the range of 651–657 nm.

## 5 Estimation of proton energy

Estimations of the proton-hydrogen flux parameters are based on the results of simulations of the Doppler-shifted hydrogen line profiles using the Monte-Carlo method. The algorithm was successfully validated in Kozelov (1994) by comparison with particle data from a rocket campaign by Søråas (1974). The algorithm takes into account the 3-component atmosphere ( $N_2$ , O,  $O_2$ ) (Hedin, 1987), the influence of the Earth's dipolar magnetic field (Kozelov, 1993) and the collisional angle scattering based on differential cross sections evaluated from laboratory measurements (Fleischmann et al., 1967, 1974; Newman et al., 1986;





**Fig. 8.** Simulated  $H\alpha$  profiles for view angle relative to the magnetic field line direction: green –  $10^\circ$ ; red –  $20^\circ$ ; blue –  $30^\circ$ . Black solid line is the least-squares fit of Eq. (1), dash-dotted line marks the fit range, dashed line is a position of non-shifted  $H\alpha$  line. (a) full profiles; (b) blue-shifted wing in log-linear coordinates.

Van Zyl et al., 1978; Gao et al., 1990). The formalism of the calculation of the Doppler profiles from transport algorithms is well-known; see, for example, Lorentzen et al. (1998). Here we only note that our interest is the high-energy ( $>10$  keV) proton precipitation; therefore, we assume that the proton fluxes are isotropic in the downward hemisphere (IDH) in pitch-angle distribution and have a Maxwellian energy distribution. Figure 8 presents an example of the simulated profiles of the hydrogen line for precipitated proton flux having a Maxwellian energy distribution with a characteristic energy of  $E_0=16$  keV, so the average energy of particles in precipitated flux is  $\langle E \rangle = 2E_0 = 32$  keV. The isotropic pitch-angle distribution results in the independence of the blue wing of the profile on the view angle relative to the magnetic field line direction in the range from  $0^\circ$  to  $30^\circ$ .

As we have noted above, the Doppler-shift of the maximum of the hydrogen line profile is not a good characteristic of the energy of precipitated proton flux. The blue wing of the profile is the most informative region to characterize the flux energy. For precipitation of the monoenergetic proton flux, the most shifted region of the profile directly corresponds to the maximal velocity of excited particles in proton-hydrogen atom flux and, therefore, to the energy of precipitated particles. For energy-distributed precipitated proton flux, the situation is more complicated due to the possible existence of a long tail from a small flux of high-energetic particles. However, to obtain the energetic characteristic of the precipitated particles, we will use the similar approach: the least-squares fit of the blue wing of the profile in the range of 651–654 nm ( $E_{||}=5.8$ –30.6 keV) to determine a crossing point of the fitted line with intensity axis (background level).

To fit the experimental profile we have used the approximation:

$$I(\Delta\lambda) = a \ln \Delta\lambda + b, \quad (1)$$

where  $\Delta\lambda = \lambda_{H\alpha} - \lambda$  is the shift relative to the  $\lambda_{H\alpha} = 656.3$  nm;  $I(\Delta\lambda)$  is emission intensity at the  $\Delta\lambda$ ;  $a$  and  $b$  are fitting

parameters. Then the crossing point of the fitted line with the intensity axis corresponds to

$$\Delta\lambda_{max} = \exp(-b/a). \quad (2)$$

From the relation  $\Delta v/c = \Delta\lambda_{max}/\lambda_{H\alpha}$  we can obtain:

$$\Delta v = c \Delta\lambda_{max}/\lambda_{H\alpha} \quad (3)$$

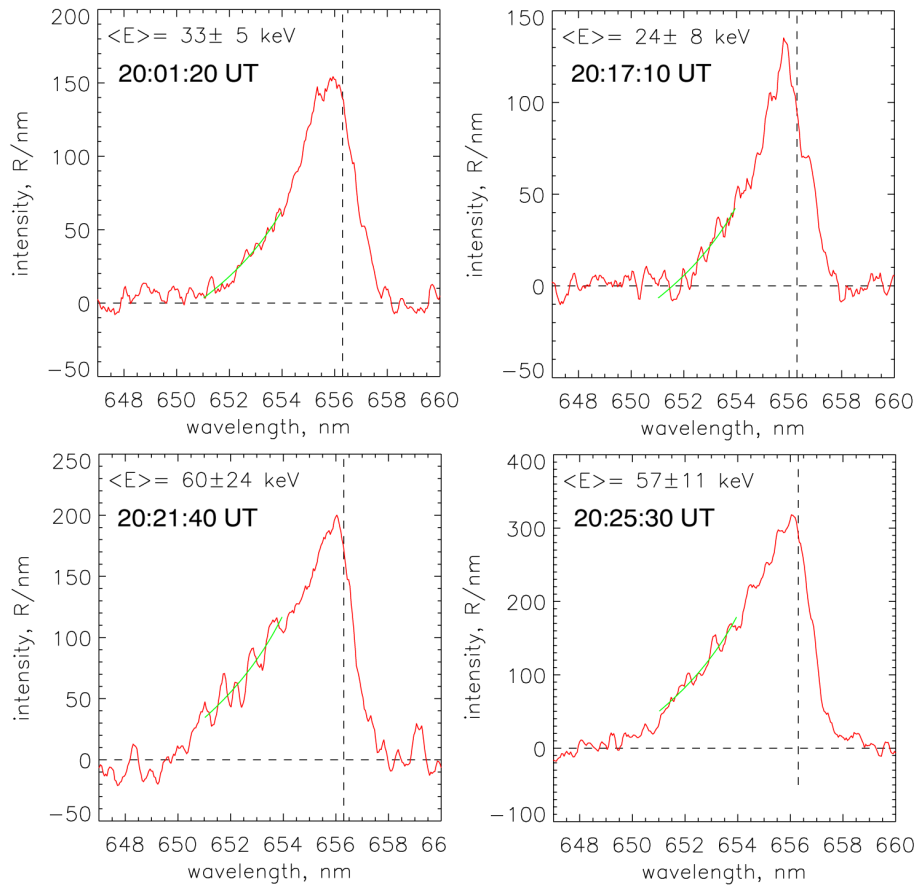
$$\langle E \rangle = m_p \Delta v^2 / 2 \approx (\Delta\lambda_{max}/0.9575)^2. \quad (4)$$

Here  $c$  is the velocity of light,  $m_p$  is the proton mass;  $\langle E \rangle$  is an effective (average) energy of particles in the flux.

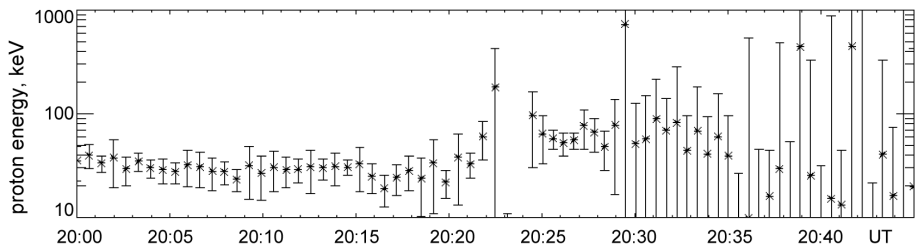
The application of the approach to simulated profiles shows that for a Maxwellian energy distribution of the initial proton flux, the approach allows us to estimate the average energy of precipitated protons in the range from 16 to 100 keV with a discrepancy of less than 15%. Figure 8 illustrates the estimation for flux with characteristic Maxwellian energy of  $E_0=16$  keV, so the average energy is  $2E_0=32$  keV. An estimated value is 35 keV. So, we deduce that this approach to energy estimation is working sufficiently well.

The experimental profiles, even after removal of the blended emissions, contains significant noise fluctuations. This is in addition to the error in the estimated energy. Figure 9 presents several examples of the experimental profiles after removal of the blended emissions. The least-square fitted lines by Eq. (1) are marked with smooth curves.

The evolution of the estimated average energy during the event of interest is presented in Fig. 10. One can see that during the growth phase (20:00–20:20 UT) the energy of protons decreased gradually from  $\sim 35$  keV to  $\sim 20$  keV. The protons of the second population during the expansion phase (20:25–20:28 UT) have the higher energy,  $\sim 60$  keV. We can note that during intervals of small  $H\alpha$  intensity (see Fig. 7b) the large discrepancy of estimated average energy mimics the relatively large noise fluctuation in the wavelength region used for the fitting procedure.



**Fig. 9.** Examples of observed H $\alpha$  profiles (red lines) and the least-squares fit of the blue wing by Eq. (1) (green lines).  $\langle E \rangle$  is an estimated average energy of precipitated protons. Dashed lines mark a position of non-shifted H $\alpha$  line and zero intensity level.



**Fig. 10.** Evolution of the estimated average energy of precipitated protons.

Due to specific features of the excitation cross sections, the efficiency of the hydrogen excitation decreased with increasing of particle energy. Therefore, the estimated energy flux of proton precipitation does not directly correspond to the average proton energy or intensity of the hydrogen emission. During the growth phase, the estimated energy flux of proton precipitation then decreased from  $\sim 6.5 \times 10^{-4}$  Joule/m $^2$  s to  $\sim 2.5 \times 10^{-4}$  Joule/m $^2$  s. During the intensification of 20:25–20:28 UT the proton energy flux was  $\sim 2.2 \times 10^{-3}$  Joule/m $^2$  s.

## 6 Discussion

### 6.1 Modelling of the degradation of the proton-H atom flux

Some attempts to estimate the parameters of precipitating protons from the hydrogen emission profiles were undertaken already in early papers; see review in Eather (1967). Such estimates are usually based on a comparison of observed profiles with the ones obtained by some theoretical calculations modelling the degradation of the proton-H atom flux in the atmosphere. However, the solution to this problem is complicated by several factors, which forced critical concerns in many previous papers. First, the cross sections

for collisional reactions of protons and H-atoms with atmospheric gases in the energy range needed for theoretical calculations were published only in Van Zyl et al. (1978, 1980). Many previous papers were based on erroneous assumptions about the energy dependencies of these cross sections; see review in McNeal and Birely, (1973).

Second, even after publication of these cross sections, many authors continued to discuss such a characteristic as a “Doppler shift of the profile maximum” (Takahashi and Fukunishi, 2001). However, the elementary consideration of the modern cross sections demonstrates that this shift is basically determined by the position of the maximum in the energy dependence of the effective cross section of hydrogen emission ( $\sim 1$  keV) and by the pitch-angle distribution in the precipitation flux at low energies ( $< 10$  keV). Higher-energetic particles with the large pitch angles can also give a contribution to the maximum of the hydrogen line profile, but this contribution is much weaker because of the decrease in excitation efficiency at the larger energies ( $> 20$  keV). Besides, in some papers, the dependence of the efficiency of hydrogen emission excitation on initial proton energy is ignored (for example, Mende et al. (2001)).

Third, the theoretical models of proton-H atom flux transport in the atmosphere have problems related to specific features of the various algorithms. The most powerful approach is the Monte-Carlo method, based on direct simulation of the particle trajectories in the framework of a “collision-by-collision” algorithm. This approach was used, for example, in the following papers (Davidson, 1965; Kozelov and Ivanov, 1992; Kozelov, 1993, 1994; Lorentzen et al., 1998; Synnes et al., 1998), where the authors obtained many transport characteristics of the proton-H atom fluxes in the Earth’s atmosphere. The Monte-Carlo algorithms naturally take into account the collisional scattering of particles and the influence of the inhomogeneous magnetic field. The calculation precision is limited only by the errors in the collision cross sections. The statistical discrepancy can be obtained in the algorithm itself during the model calculation.

A simpler method of transport calculations is based on the continuous slowing down approximation (CSDA) (Edgar et al., 1973, 1975). In reality, after several collisions with atoms and molecules of an absorber, the precipitated proton flux comes to a charge equilibrium state defined by the relation between cross sections of charge exchange and stripping processes. On average, the energy loss of a monodirectional proton-hydrogen atom flux in the charge equilibrium state is described by the equation  $dE/dx = -L(E)$ , where  $L(E)$  is the loss function (Edgar et al., 1973). The more refined version of the CSDA for non-equilibrium proton-hydrogen atom flux was considered in Decker et al. (1996). However, the methods based on CSDA cannot directly take into account such effects as the discrete nature of the energy loss, collisional angle scattering, and the influence of the magnetic field.

Several authors have developed different algorithms to solve directly the transport equation for particle fluxes (Jasperse and Basu, 1982; Basu et al., 1990; Galand et al.,

1997). Usually the algorithms need additional assumptions about transport processes to simplify the numerical calculations. However, the common problem of such algorithms is the number of interpolation procedures needed to recalculate the distribution of particle fluxes after each step of the spatial variable. Therefore, the algorithms tend to accumulate a systematic error which is hardly detectable in the framework of the method itself. The checking methods (for example, checking of the energy conservation rule, increasing of cell number in the phase space, etc.) cannot solve the problem. Usually the systematic error leads to artifacts near the boundaries of the spatial region occupied by the dissipated particles. In a paper by Decker et al. (1996) the above-mentioned three theoretical techniques were tested by precisely the same incoming parameters and input conditions. For the simplest conditions (without collisional angle scattering and magnetic field, and with a simplified set of cross sections) it was shown that these theoretical techniques gave practically the same distribution of the particle fluxes in the dissipation region.

## 6.2 Lateral spreading and structure of hydrogen emission region

It is well-known that the hydrogen emission region is diffuse because the precipitated protons undergo charge-exchange collisions and thereby spend a part of the time as neutral hydrogen atoms that cross magnetic field lines. Davidson (1965) modeled the lateral spread of hydrogen emission in the atmosphere due to this process using a thin sheet of initial precipitated protons. The characteristic scale of the hydrogen emission zone was found to be  $\sim 100$  km. Therefore, the spatial structure observed in our case for the second population of protons seems very strange in that it is less diffuse, or more structured than one would expect. Kozelov (1993) provided a more detailed study of the lateral spreading effect in the dipolar magnetic field and showed that the spreading effect decreases with the increasing of the initial proton energy. Thus, the observed structure is additional evidence of the higher energy of precipitated protons, relative to the protons in the diffuse aurora.

Let’s consider the ionospheric spatial structures produced by a set of three sheets of monoenergetic proton precipitation with a distance of  $\sim 20$  km between the centers of the sheets. The longitudinal orientation of the sheets is assumed, and the latitudinal section of simulated precipitation structure is shown in top panel of Fig. 11. A downward isotropic pitch-angle distribution in precipitation flux is assumed. The precipitation is characterized by values of initial energy ( $E_o$ ) and of energy flux ( $F_p$ ). The simplified model from Kozelov et al. (1995), based on the results of detailed Monte-Carlo transport modeling, was used to calculate the H $\alpha$  luminosity structure resulting from such precipitation for different values of  $E_o$  and  $F_p$ . The bottom panel of Fig. 11 presents the simulation results for two cases:  $E_o = 20$  keV,  $F_p = 2.5 \times 10^{-4}$  Joule/m<sup>2</sup> s and  $E_o = 60$  keV,  $F_p = 2.2 \times 10^{-3}$  Joule/m<sup>2</sup> s. The last case corresponds to

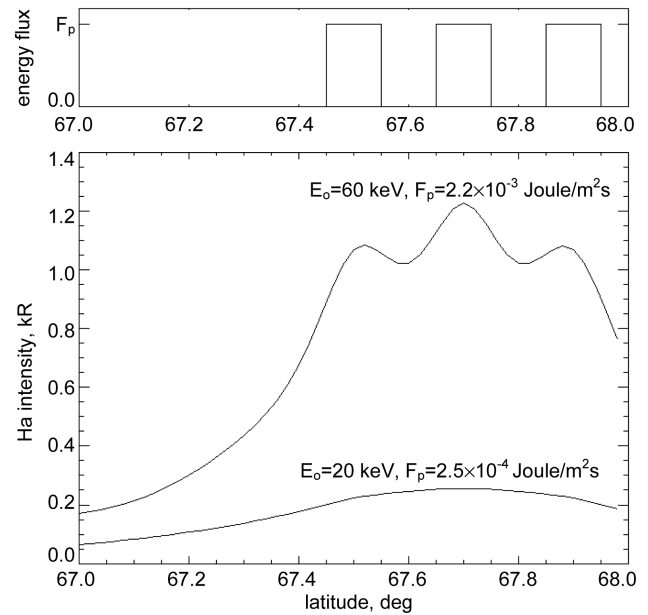
characteristics of a second population of the precipitated protons observed at 20:25–20:28 UT.

The results illustrate that the structure of the precipitation for  $E_0=60$  keV is easily observed in the spatial luminosity distribution, while no structure is seen for  $E_0=20$  keV. The variations of 10–20% in the simulation luminosity for  $E_0=60$  keV correspond well to the observed spatial luminosity variations in 20:25–20:27 UT at  $0^\circ$ – $30^\circ$  zenith angles; see Fig. 5c. This structure is absent in Fig. 5d, therefore, we can deduce that this is the structure of hydrogen emission, but not of other emissions excited by electron precipitation. So, the simulation results of the spatial distribution of luminosity also supports the conclusion about the higher characteristic energy of protons precipitated in the second population.

### 6.3 Mechanisms of proton precipitation at different phases of the substorm

Before the substorm onset, and during the late growth phase of the event considered here the morphology of the hydrogen emission is comparable to previous observations (Yevlashin and Yevlashina, 1980). The wide ( $>100$  km) diffuse region of the equatorward hydrogen emission is slowly moving towards the equator. The integrated intensity of H $\alpha$  emission was 300–500 R, the average energy of the precipitated proton was 20–35 keV, and the estimated energy flux of proton precipitation was  $\sim 3 \times 10^{-4}$  Joule/m $^2$  s, which is typical for the region (Hardy et al., 1991). The proton drift in the large-scale electric and magnetic fields and subsequent interaction with ion-cyclotron waves lead to pitch-angle diffusion of the proton flux in the plasma sheet and, therefore, to the precipitation in the atmosphere. Inside the region of the hydrogen emission there are more mobile electron structures. The most equatorial discrete electron arc had several intensifications and the last of them ( $\sim 20:16$  UT) led to the auroral breakup ( $\sim 20:21$  UT). In the breakup, or onset, the arc was located at the poleward boundary of the hydrogen emission region.

The main peculiarity of this event is a clear manifestation of the dynamical structure of the hydrogen emission evolution during the substorm at the temporal scales of 2–3 min after onset, when both the intensity and the shape of the hydrogen emission line profile varied dramatically. During the first 2 min after onset, the hydrogen line was also increasing in intensity. However, when the electron arc expanded poleward with the development of the westward traveling surge (WTS), the hydrogen emission was absent in the leading edge of the WTS. The same features had been noted previously by Takahashi and Fukunishi (2001) and Deehr and Lummerzheim (2001). The next intensification of the hydrogen emission appeared 2–3 min later on the north side of the field of view inside the region of active aurora. Deehr and Lummerzheim (2001) found that “the H emission expands poleward after the formation of the N-S oriented arcs”. In our case, the N-S oriented electron arcs were also present during the H $\alpha$  intensification from 20:24 to 20:28 UT; see Fig. 4, all-sky TV image for 20:26 UT. The H $\alpha$  emission intensification region was located a bit northward from the poleward



**Fig. 11.** H $\alpha$  luminosity as a function of latitude, stimulated by three regions of precipitated protons. Top panel is a latitudinal section of precipitation structure. Bottom panel presents the spatial structures of the column-integrated luminosity for two cases of proton precipitation parameters.

boundary of the hydrogen emission at the end of the growth phase. However, as we deduce from the hydrogen line profile, the average energy of the precipitated protons during 20:24–20:28 UT was higher by a factor of 2–3 than before the breakup. An increase of the particle flux at energies up to  $>300$  keV was found by Sergeev and Kubyshkina (1996) 3–4 min after a substorm onset in data from low-altitude satellites.

During the expansion phase of substorm, the earthward injection of energetic particles occurs under the influence of large, short-lived electric fields. In the course of this injection, the particles are subjected to both impulsive parallel and perpendicular acceleration and nonadiabatic motion. The convection surge mechanism (Mauk, 1986) for the ion acceleration is efficient in producing energetic particles, especially when nonadiabatic effects are concerned (Delcourt et al, 1994). The dynamics of the non-adiabatic motion region of protons during the development of the expansion phase in the equatorial plane of the nightside magnetosphere at  $5$ – $8 R_E$  was analyzed in Kozelov and Kozelova (2003), using data from the CRRES satellite and the adapted time-dependent magnetospheric model. The spatial scale of the transition region (boundary) from adiabatic to non-adiabatic ion motion was found to be about  $1 R_E$ . This boundary is bent in a radial direction before the beginning of an expansion phase and then is moving in a tailward direction. This movement is accompanied by fast changes of pitch-angle distribution in proton flux. Therefore, we can assume that the second population of protons in the event discussed is probably produced by pitch-angle scattering of protons due to

the non-adiabatic motion in the region of local dipolarization near the equatorial plane.

Thus, the characteristics of the proton precipitation fluxes before the leading edge of WTS and inside the WTS region appear to be different from one another. The possible physical mechanisms that lead to the precipitation differ, too. Therefore, we have strong evidence for the existence of two different populations of protons.

## 7 Conclusions

The paper describes the results of coordinated, optical ground-based observations of the auroral substorm on 26 March 2004 in the Kola Peninsula. In addition to the usual observatory equipment, consisting of an all-sky TV camera and a scanning photometer, a special imaging spectrograph with high spectral and temporal resolution was used. The data from this device allowed us to use the Doppler profile of the  $H\alpha$  hydrogen emission to estimate the variation in the energy of the precipitating proton flux and the proton emission intensity.

Two different populations of precipitated proton flux were observed during the event. The first of them was a diffuse proton precipitation that is usually observed in the evening sector of the auroral oval and located equatorward of the discrete electron precipitation. The average energy of the protons during this precipitation was  $\sim 20\text{--}35$  keV, the energy flux was  $\sim 3 \times 10^{-4}$  Joule/m<sup>2</sup> s.

The second population of precipitated protons was separated from the first one by the breakup, or onset arc (boundary of WTS). The proton precipitation was observed 1–2 min after the breakup and 4–5 min into the expansion phase of the substorm into the zone of bright discrete auroral structures. The average energy of the protons during this precipitation was  $\sim 60$  keV, the energy flux was  $\sim 2.2 \times 10^{-3}$  Joule/m<sup>2</sup> s. We believe that the most probable mechanism of precipitation was pitch-angle scattering of protons due to non-adiabatic motion in the region of local dipolarization near the equatorial plane.

*Acknowledgements.* The work was partly supported by the basic research programs of Russian Academy of Sciences "Solar activity and physical processes in the Sun-Earth system" and the Division of Physical Sciences of RAS "Plasma processes in the solar system". Authors thank the Polar Geophysical Institute for the data from the Lovozero observatory and all referees for helpful comments.

Topical Editor U.-P. Hoppe thanks C. Deehr and two other referees for their help in evaluating this paper.

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