



HAL
open science

Total ozone decrease in the Arctic after REP events

V. C. Roldugin, M. I. Beloglazov, G. F. Remenets

► **To cite this version:**

V. C. Roldugin, M. I. Beloglazov, G. F. Remenets. Total ozone decrease in the Arctic after REP events. *Annales Geophysicae*, 2000, 18 (3), pp.332-336. hal-00316599

HAL Id: hal-00316599

<https://hal.science/hal-00316599>

Submitted on 18 Jun 2008

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Total ozone decrease in the Arctic after REP events

V. C. Roldugin¹, M. I. Beloglazov¹, G. F. Remenets²

¹Polar Geophysical Institute, Apatity, Murmansk reg., Russia
E-mail: roldugin@pgi.kolasc.net.ru

²Institute of Radiophysics, University of St.-Petersburg, Russia

Received: 17 December 1997 / Revised: 23 April 1999 / Accepted: 10 May 1999

Abstract. Eight periods of relativistic electron precipitation (REP) with electron energies of more than 300 keV are identified from VLF data (10–14 kHz) monitored along the Aldra (Norway) – Apatity (Kola peninsula) radio trace. In these cases, anomalous ionization below 55–50 km occurred without disturbing the higher layers of the ionosphere. The daily total ozone values in Murmansk for six days before and six days after the REP events are compared. In seven of eight events a decrease in the total ozone of about 20 DU is observed. In one event of 25 March, 1986, the mean total ozone value for six days before the REP is bigger than that for six days after, but this a case of an extremely high ozone increase (144 DU during the six days). However, on days 3 and 4 there was a minimum of about 47 DU with regard to REP days, so this case also confirms the concept of the ozone decrease after REP. The difference between mean ozone values for periods six days before and six days after the REPs was found also for 23 points in Arctic on TOMS data. The difference was negative only in Murmansk longitudinal sector. Along the meridian of the trace it was negative at high latitudes in both hemispheres and was near zero at low latitudes.

Key words: Atmospheric composition and structure (middle atmosphere – composition and chemistry) – Meteorology and atmospheric dynamics (polar meteorology)

1 Introduction

The effects of high energy particle precipitation on the ozone layer are noted in many works. Stephenson and Scourfield (1992) have observed from satellite data,

ozone depletion over the polar caps caused by solar protons. Shumilov *et al.* (1992) described the same effect using ground data. Shumilov *et al.* (1997) also reported an ozone increase of 10% after Forbush decreases of galactic cosmic rays.

Hence, we might expect similar or even stronger effects after atmospheric penetration of relativistic electrons with energy of more than 300 keV. Baker *et al.* (1990a) discussed the ozone decrease in the upper stratosphere speculatively, but without any concrete investigations. Callis *et al.* (1991a, b) connect the ozone decrease in the 1980s with enhancement of relativistic electron fluxes which produce odd nitrogen molecules. Their model suggests that by 1985 the relativistic electrons are responsible for 2.6% decay of the global ozone.

2 Ground-based data

As a rule, relativistic electron precipitation (REP) events are registered by measurements on satellites. We do not know of published records or year books with these data. To identify REP events we have used data of amplitude and phase measurements on VLF traces (10.2, 12.1 and 13.6 kHz) at Aldra (Norway, 66.3°N, 13.2°E) and Apatity (Kola peninsula, 67.5°N, 33.3°E); see map in Fig. 1. Remenets and Beloglazov (1992) have shown that, on this trace during sunlit conditions, specific variations of VLF signal occur in REP events: amplitudes and phases of all three frequencies diminish simultaneously in the absence of magnetic disturbances and anomalous riometer absorption.

To model numerically the changes of the electron profiles in the low ionosphere during REP events, Beloglazov and Remenets (1997) have selected 12 such cases for the period 1982–1989. The principle of selection was the high quality of the experimental data in order to solve the VLF inversion problem. The results of the calculations have shown significant changes in electron profiles below 50 km, i.e., in the stratosphere.

These events are compared with variations of the total ozone in Murmansk which is situated not far from

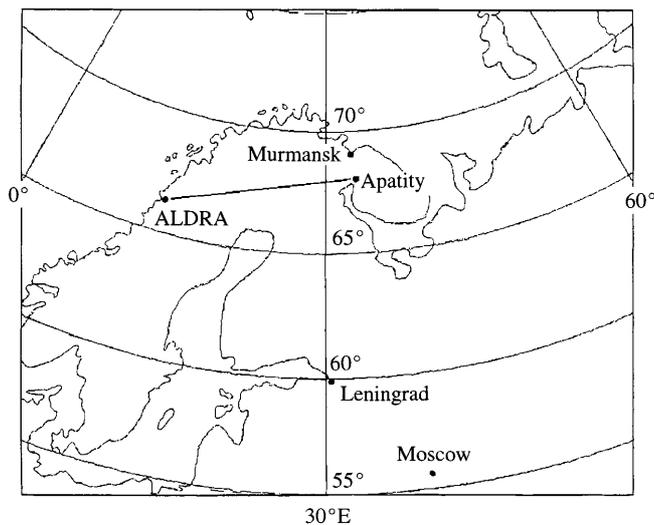


Fig. 1. Map of the VLF trace

the radio trace (see Fig. 1). The event number is reduced to eight because one case occurred in November when the ozone measurements are not carried out, and three others took place from one to three days after the preceding precipitation and were combined as single events. In Table 1, dates and times of the start and end of the chosen events are shown. They occurred during the daylight portion of the day and had durations from 5 min to 10 h. Magnetic disturbances and riometer absorption in Kola peninsula were absent during these events.

The cosmic ray intensity is examined for these events using data from the neutron monitor in Apatity. Six events occurred in quiet conditions or in small Forbush decreases. The case of 29 September, 1989, occurred during the strong solar proton event with the onset at 1150 UT. Table 1 shows that this REP event ended 2 h before the onset. Certainly, in this event both relativistic electrons and solar protons may influence on the ozone layer. The case of 23 October, 1989, was in the Forbush decrease just after SPE at 22 October.

All eight events have occurred when even the troposphere was illuminated by the Sun. It is in accordance with the result of Beloglazova and Beloglazov (1982), that the diurnal variation of such

phenomena in Apatity has maximum at 11–19 LT, or nearly 9–17 UT.

Baker *et al.* (1990b) found a strong 13- and 27-day periodicity of the fluxes of high-energy electrons. Despite the small number of our cases, this periodicity is seen well: (1) 15 and 29 September 1982, (2) 25–27 March and 22–23 April, 1986, (3) 22–23 April and 5 May, 1986, and (4) 29 September and 23 October, 1989. It confirms that the phenomena picked in the VLF propagation are REP events.

According to data from Baker *et al.* (1990a), the REP events occur on the geomagnetic latitude $62^\circ \pm 6^\circ$. The geomagnetic latitude of Murmansk is about 64° , the main part of the radio trace is situated also in the REP zone. Murmansk is to the east of the trace, and the main winds in this region are westerly, which carry ozone from the trace to the observatory. Thus, the disposition of the stations can be regarded as favourable for this investigation. The total ozone values in Murmansk are taken from the Soviet year-books edited by Gushchin (1984, etc).

3 Ozone changes in Murmansk

The mean daily ozone data were picked out for the intervals between 6 days before to 6 days after the REP events. They are shown in Fig. 2. To smooth over seasonal dependence, the ozone variations are shown here in relative units, normalized to the values for the day of the REP, which is marked in Fig. 2 by the vertical line. This relative scale is shown. For three cases of twinned REP occurrences with 1–3 days of lag combined into single events, the zero day is taken as the day of the first REP event. Annoying gaps in the ozone data can be seen, however, in Murmansk gaps in ozone data usually occur more frequently than in these cases. The bottom curve is the result of averaging of all eight normalized curves. The dotted horizontal lines show the mean ozone values before and after the REP events.

Figure 2 shows that for seven events out of the eight studied the mean ozone values after the REPs are smaller than before. The differences in Dobson Units are given in Table 1, where the mean ozone values before and after the events together with the mean square

Table 1. List of REP events and changes of total ozone in Murmansk

No	Day	Time (UT)		Ozone		
		Beginning	End	Before REP	After REP	Difference
1	15 September 1982	12 h 55 m	13 h 50 m	343 ± 2	319 ± 8	25
2	29 September 1982	13 h 20 m	13 h 25 m	333 ± 4	314 ± 6	20
3	20 October 1985	10 h 50 m	12 h 10 m	292 ± 3	272 ± 18	20
4	23 October 1985	8 h 48 m	11 h 30 m	403 ± 21	439 ± 6	-36 (47)
	25 March 1986	9 h 30 m	17 h 00 m			
5	27 March 1986	12 h 35 m	19 h 30 m	472 ± 8	435 ± 8	37
	22 April 1986	12 h 15 m	22 h 00 m			
6	23 April 1986	18 h 20 m	20 h 20 m	417 ± 6	410 ± 6	7
	5 May 1986	7 h 30 m	9 h 40 m			
7	29 September 1989	4 h 00 m	10 h 00 m	273 ± 11	252 ± 6	21
8	23 October 1989	8 h 08 m	9 h 15 m	226 ± 6	206 ± 7	20

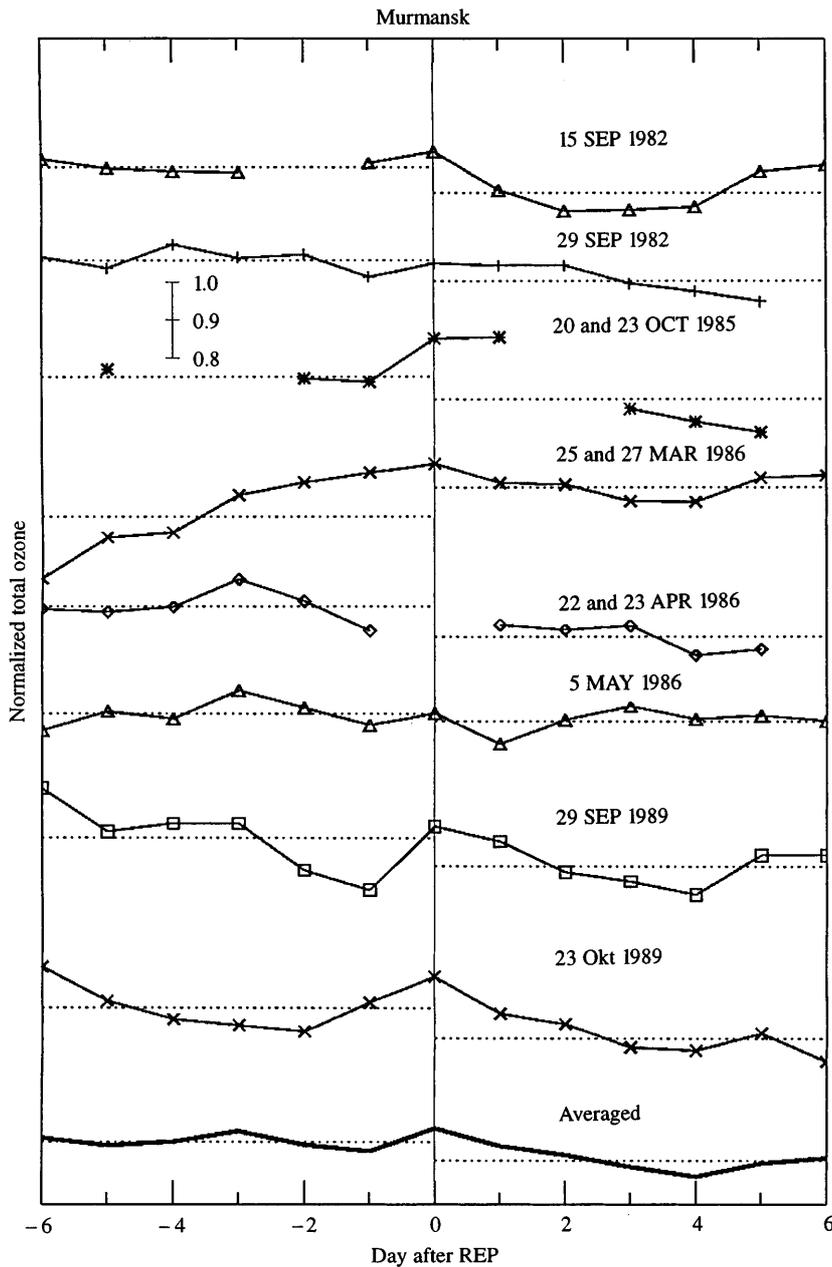


Fig. 2. The normalized ozone variations in Murmansk before and after REP events. The dotted lines are mean values. The bottom curve is an average of all eight events

errors are given also, in DU. One can see that the differences are less than the errors, and only for the 5 May, 1986, case is the ozone decrease comparable with the error. The typical value of the difference is about 20 DU.

In one example for 25–27 March, 1986, the mean ozone value before the REP is significantly larger than after it. However in that case the extremely large total ozone increase happened when the ozone rise of 144 DU occurred for these six days. Later however, an appreciable ozone decrease was seen the next day and also on the third fourth day, when the ozone value was smaller by 47 DU than the highest on the REP day. Thus, an ozone decrease took place during this event as well.

At the bottom of Fig. 2, the averaged data for all eight normalized ozone variations are shown. The averaged-out normalized values in the space of six days

are equal to 0.964 before the REP events and 0.915 after, the difference is 0.049, which is significantly less than errors of both mean values 0.007 and 0.011, correspondingly.

4 Satellite ozone data

For those eight events the ozone changes were examined also for 23 stations situated in high latitudes by using TOMS data. We averaged the total ozone values for 6 days before and 6 days after the events without normalizing to zero day value, and determined mean differences. The number of stations with negative difference, corresponding to an ozone decrease, was twice as big as positive one for all events with the exception of 25 March, 1986.

Then, for every event and every station, we excluded the seasonal difference “6 days after – 6 days before”. To do this we calculated the daily averaged variation using the TOMS 1979–1993 data, approximated them by cubic polynomials and then found the required difference. Figure 3 shows the mean values of eight event differences, excluding seasonal changes, at all 23 stations by circles: solid circles show a negative difference, open circles mean a positive difference. There are 14 stations with decreased total ozone and nine stations with increased ozone. It is clear the stations with ozone decreases are grouped in the longitudinal sector between 60°W–90°E, around the Murmansk meridian. In Murmansk the mean ozone decrease is 14 DU (21 DU without the event of 25 March 1986).

On the base of the TOMS data the difference “6 days after – 6 days before” was determined also along the geographic meridian from the pole to the equator, passing through the middle point between Aldra and Apatity (67°N, 23°E), and then along the geographic meridian in the Southern Hemisphere, passing through the geomagnetic conjugated point (56°S, 63°E). The result is shown in Fig. 4. The vertical bars at both points present the statistical errors. Here the negative deflections exceed the error twice. Figure 4 testifies to the ozone decrease at high latitudes after the REP events.

5 Discussion

The comparison of the total ozone variations in the ground-based observatory at Murmansk before and after eight REP events over Scandinavia reveals an

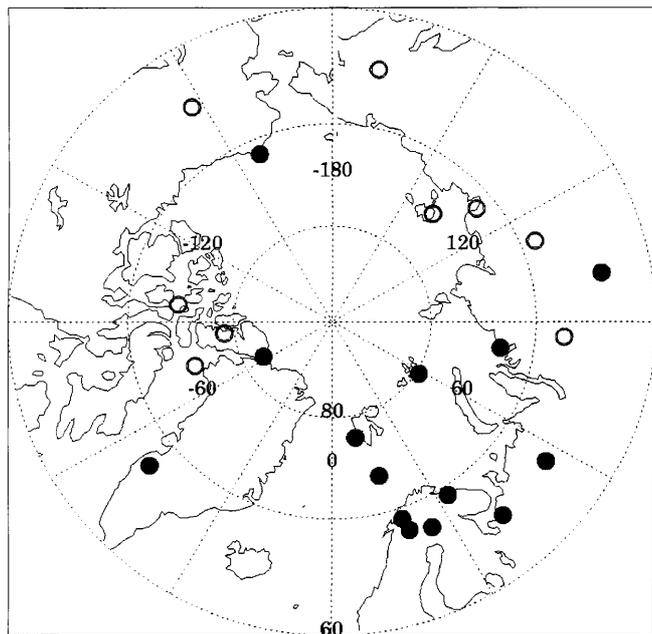


Fig. 3. Main geophysical observatories in the northern polar cap used to examine the difference of the total ozone after and before the REP events. The *solid circles* mean the ozone decrease, the *open circles* mean a positive difference

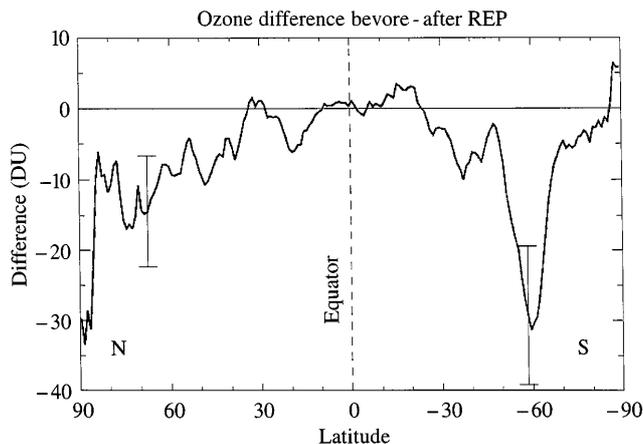


Fig. 4. The meridian profile of the ozone decrease after the REP events. The vertical bars are standard errors

ozone decrease of about 20 DU over several days. This value exceeds by several times, the mean season ozone diminution for the six day period by about 4 DU and seems to be statistically significant. Accordingly to TOMS data, the ozone decrease is observed also at the area near Murmansk in the Arctic and in the magnetic conjugated area in the Antarctic. The decrease is absent in middle latitudes (see Fig. 4), where the relativistic electrons do not precipitate.

However there are lots of stations in the polar cap where the ozone decrease after REP is not observed, e.g. over eastern Siberia or western Canada. Two different reasons for this phenomenon occur to us.

1. Because there is no decrease on the dark side, we can guess that the process of the ozone decay includes any photoreactions. In particular, it may be photolysis of N_2O that is involved, see Callis *et al.* (1991b)
2. The eight REPs chosen were local.

The second explanation seems to be in contradiction with apparent global character of REP. For example, Reeves *et al.* (1998) demonstrate synchronous temporal variations of relativistic electron fluxes of $E > 2$ MeV on 13 satellites in January 1997 over several days.

However, our events were not chosen on the basis of satellite data, but from data of VLF propagation over Scandinavia, so the physical distinction may exist between REPs determined from space and ground-based observations. Really, accordingly to Fig. 1 of Imhof *et al.* (1991), the precipitating fluxes of electrons $E > 1$ MeV at low altitudes from the S81-1 satellite during our REP cases 15 and 29 September 1982 (258 and 272 days), were very moderate. The authors give information of spikes of precipitating relativistic electrons exceeding by tenfold the daily averaged trapped fluxes at high altitudes.

6 Conclusion

Both ground-based ozone observations and TOMS data testify to an ozone decrease of about 20 DU during

relativistic electron precipitation events, which were determined from disturbance of VLF propagation.

Acknowledgements. Topical Editor F. Vial thanks D. N. Baker and E. Kyrö for their help in evaluating this paper.

References

- Baker, D. N., J. B. Blake, R. W. Klebesadel, D. D. Sentman, D. J. Gorney, and P. R. Higbie**, Relativistic magnetospheric electrons: lower ionospheric conductivity and long-term atmospheric variability, *Adv. Space Res.*, **10**, 229–233, 1990a.
- Baker, D. N., R. L. McPherron, T. E. Cayton, and R. W. Klebesadel**, Linear prediction filter analysis of relativistic electron properties at 6.6 RE, *J. Geophys. Res.*, **95**, 15 133–15 140, 1990b.
- Beloglazov, M. I., and G. F. Remenets**, Unique ground VLF monitoring of the relativistic electron precipitation, in *Physics of auroral phenomena. XX annual Apatitian seminar. 25–28 February 1997*, Abstracts, Preprint PGI-97-01-101, Apatity, 1997.
- Beloglazova, G. P., and M. I. Beloglazov**, On one type of phase anomaly of VLF-signal at auroral trace, *Geomagn. Aeron.*, **22**, 56–59, 1982.
- Callis, L. B., R. E. Boughner, M. Natarajan, J. D. Lambeth, D. N. Baker, and J. B. Blake**, Ozone depletion in the high latitude lower stratosphere: 1979–1990, *J. Geophys. Res.*, **96**, 2921–2937, 1991a.
- Callis, L. B., D. N. Baker, J. B. Blake, J. D. Lambeth, R. E. Boughner, M. Natarajan, R. W. Klebesadel, and D. J. Gorney**, Precipitating relativistic electrons: their long-term effect on stratospheric odd nitrogen levels, *J. Geophys. Res.*, **96**, 2939–2976, 1991b.
- Gushchin, G. P.**, (Ed.) Total content of the ozone and spectral transparency of the atmosphere // Leningrad, *Hydrometeoizdat*, 1984 and etc. (in Russian).
- Imhof, W. L., H. D. Voss, J. Mobilia, D. W. Datlow, J. P. McGlennon, and D. N. Baker**, Relativistic electron enhancements: simultaneous measurements from synchronous and low altitude satellites, *Geophys. Res. Lett.*, **18**, 397–400, 1991.
- Reeves, G. D., D. N. Baker, R. D. Belian, J. B. Blake, T. E. Cayton, J. F. Fennel, R. H. W. Friedel, M. M. Meier, R. S. Selesnick, and H. E. Spence**, The global response of relativistic radiation belt electrons to the January 1997 magnetic cloud, *Geophys. Res. Lett.*, **25**, 3265–3268, 1998.
- Remenets, G. F., and M. I. Beloglazov**, Dynamics of an auroral low ionospheric fringe of geophysical disturbances on 29.09.89, *Planet. Space. Sci.*, **40**, 1101–1108, 1992.
- Shumilov, O. I., K. Henriksen, O. M. Raspopov, and E. A. Kasatkina**, Arctic ozone abundance and solar proton events, *Geophys. Res. Lett.*, **19**, 1647, 1992.
- Shumilov, O. I., E. A. Kasatkina, O. M. Raspopov, and K. Henriksen**, Influence of Forbush decreases of galactic cosmic rays on the ozone layer, *Geomagn. Aeron.*, **37**, 24–31, 1997 (in Russian).
- Stephenson, J. A. E., and Scourfield, M. W. J.**, Ozone depletion over the polar caps caused by solar protons, *Geophys. Res. Lett.*, **12**, 2425, 1992.