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Electron temperatures during rapid subauroral ion drift events

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Abstract. Examples of data from DE-2 satellite instruments are presented. These illustrate the behaviour of plasma parameters in the F-region and adjacent topside ionosphere during rapid sub-auroral ion drift (SAID) events. In particular, a variety of behaviours of the electron temperature (T_e) is demonstrated, both within and equatorward of the SAID region. The Sheffield University plasmasphere-ionosphere model (SUPIM) is used to perform calculations in which a model SAID is applied to a plasma flux tube. The model results indicate that strongly elevated ion temperature (a recognised signature of SAID events) is on occasion sufficient to raise T_e to observed values by ion-electron heat transfer. On other occasions, an additional heat source is required. It is suggested that such a source for the electron gas may be due to interaction between the ring current and the plasmasphere at high altitudes. The magnitude of the downward heat flux is consistent with that necessary to produce sub-auroral red arcs. The resulting strongly heated electron gas causes vibrational excitation of molecular nitrogen in the thermosphere.

Key words. Electron temperature · Subauroral ion drift · Plasmasphere-ionosphere SUPIM

1 Introduction

Occurrences of a narrow band of rapid westward ion drift (of order 1 km s^{-1}) in the evening sector of the subauroral ionosphere were detected by instruments on board satellites. References to an early observation (Galperin *et al.*, 1973) and to subsequent clarifications of the properties and signatures of these drifts (known as

SAID) may be found in the work of Anderson *et al.*, (1993), who investigated the temporal relationship between SAID and auroral substorms. A phenomenological model of SAID production, based on the ideas of Southwood and Wolf (1978) and Harel *et al.* (1981), has been proposed by Anderson *et al.* (1993) in order to account for the observed ionospheric signatures and their temporal relationship with substorms.

Using the observed ion drifts as inputs, mathematical models of the ionosphere have been used to study the causes of the observed signatures of SAID (Sellek *et al.*, 1991, 1992; Moffett *et al.*, 1991, 1992a,b, 1993; Körösmezey *et al.*, 1992; Heelis *et al.*, 1993). Elevated and anisotropic ion temperatures, depletion of electron density, changes in ion composition, enhanced field-aligned plasma flows and the behaviour of H⁺ and He⁺ in the topside ionosphere have been modelled for various SAID conditions.

The behaviour of the electron temperature (T_e) has been neglected, apart from a passing mention by Körösmezey *et al.* (1992). The chief reason for the neglect is the sparseness of published data of T_e during SAID events. The present paper shows examples of observed T_e (Sect. 2). Also, the Sheffield University plasmasphere-ionosphere model (SUPIM), with inclusion of ion-temperature anisotropy and vibrational excitation of molecular nitrogen (Sect. 3), is used to interpret the T_e data (Sect. 4). It is found that ionelectron heat transfer due to the elevated ion temperature (T_i) is sometimes sufficient to generate the T_e values, but on other occasions an additional heat source for the electron gas is required, since T_e may become comparable with T_i .

2 Observations of SAID

The Dynamics Explorer-2 satellite was launched into a 90° inclination orbit in August, 1981, with apogee near 1000 km altitude and perigee near 300 km altitude. The DE-2 instruments giving data of interest in this study are the ion drift meter (Heelis *et al.*, 1981), the retarding

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potential analyser (RPA) (Hanson *et al.*, 1981) and the Langmuir probe (Krehbiel *et al.*, 1981). The drift meter provided data on the bulk ion drift velocity in the plane perpendicular to the satellite track, nominally east-west and up-down. The RPA provided the ion concentration and composition and the ion temperature and bulk velocity in the direction of satellite motion. The Langmuir probe provided the electron concentration and temperature.

Figure 1 shows DE-2 data from a SAID encountered at about 420 km altitude on day 280 (7 October) in 1982. The top panel illustrates the characteristic spike in the east-west ion drift, with a maximum magnitude of about 2 km s⁻¹ at 08:34:40 UT. The peak in T_i , trough in N_i and peak in T_e are collocated with the maximum drift. The N_e measured by the Langmuir probe (not shown) agrees with the total ion concentration in the bottom panel.

The elevation of T_i above its background value in the plasmasphere is caused by ion-neutral frictional heating (Schunk et al., 1976; Anderson et al., 1991; Sellek et al., 1991). St.-Maurice and Schunk (1977) predicted theoretically that the ion thermal velocity distribution becomes anisotropic and is distorted towards a toroidal form when the ion drift exceeds the neutral wind speed by more than the neutral thermal speed. Toroidal distributions had been observed by the RPA on the AE-C satellite (St.-Maurice et al., 1976). More recently, incoherent-scatter radar observations have given evidence for $T_i^{\perp} > T_i^{\parallel}$, where T_i^{\perp} and T_i^{\parallel} are the perpendicular and parallel ion velocity distribution temperatures, respectively (Lockwood et al., 1993, and references therein). In the present context, the RPA essentially sees the perpendicular velocity distribution and thus measures T_i^{\perp} , which is expected to be significantly greater than T_i^{\parallel} (Moffett *et al.*, 1993, and references therein). It is to be noted that the DE-2 satellite passes almost radially through the N_e trough region where the ion drifts are aligned almost perpendicular to the meridian, and thus gradients in the ram ion drift are not expected to be large. Along the satellite track, the maximum density gradient is about one order of magnitude in about 12 s. This amounts to less than a 10% change in the density over the nominal 0.5 s it takes to apply the retarding potential. The effects of this gradient on the derived temperature are small and less than the 8% error in the overall least-square analysis procedure.

The increase in the effective temperature in the loss rates for O⁺ with N₂ and O₂ (St.-Maurice and Torr, 1978) leads to increased abundance of NO⁺ and O₂⁺ relative to O⁺ (Anderson *et al.*, 1991; Moffett *et al.*, 1992a) but a net decrease in N_e in the F-region. The depth of the trough is emphasised by being embedded in the evening sector mid-latitude trough (Anderson *et al.*, 1991), which may be substantially formed by plasma stagnation (Spiro *et al.*, 1978).

The T_e value reaches 4500 K, comparable with T_i , during the SAID event. Since heat transfer from the ion gas to the electron gas is roughly proportional to temperature difference, this is *prima facie* evidence that

DE-B convection parameters Day = 82280 Orbit = 6469 Northern hemisphere



Fig. 1. Observations by DE-2 of a SAID event seen after 0834 UT on day 180, 1982. *Top panel*: zonal component of horizontal ion drift; *second panel*: ion temperature; *third panel*: electron temperature; *bottom panel*: total ion concentration

the electrons have an additional heat source. The latitudinal profile of T_e also shows a "shoulder" that is located equatorward of the SAID T_e spike. This shoulder lies within the outer plasmasphere. There is experimental evidence that T_e typically increases with increasing latitude in the outer plasmasphere on the nightside (Brace *et al.*, 1982, 1988). The nightside heating could arise from photoelectron energy that is stored in the plasmasphere during the daytime (Brace *et al.*, 1988) but may also necessitate another and stronger heat source.

In the second example of SAID data (Fig. 2), the ion drift velocity at about 320 km altitude reaches 4 km s⁻ westward at about 0943 UT, with the ion temperature as measured by the RPA about 5000 K. The T_e profile shows four features worthy of note: (1) a shoulder from about 0941 to 0942UT: (2) an elevation to about 3000 K, between 0942 and 0943UT, followed by (3) a minimum; and (4) a spike reaching 2300 K, where the ion drift speed is at its greatest value. Remarks on feature (1) are similar to those already made on the T_e shoulder in Fig. 1. In this case, in addition, N_e is decreasing with latitude and thus less heat (from whatever source) will be necessary to cause the shoulder. Features (2) and (3) indicate that additional heating of the electron gas may take place equatorward of the SAID or that the observed plasma may have been heated previously in a SAID region. On the other hand, feature (4) (the spike in T_e in the SAID event) may be produced by ion-electron heat transfer alone, given the relatively large T_i observed.

The final example (Fig. 3) shows another form of T_e behaviour. The value of T_e increases steadily with latitude in the outer plasmasphere but is still less than 2000 K before the SAID event is encountered (just after 0912 UT). A complication is the reducing solar zenith angle (<100°), leading to some direct solar heating of the electron gas. In this case, within the SAID region, T_e rises towards T_i , as in Fig. 1.

3 Plasmasphere-ionosphere model

Bailey and Balan (1996) have described the standard version of SUPIM that is based on the model of Bailey and Sellek (1990) and now includes an eccentric dipole geomagnetic field. This version is used for low- and moderate-latitude applications. Briefly, values of the concentrations, field-aligned fluxes and temperatures of the $O^+, H^+, He^+, N_2^+, O_2^+$ and NO^+ ions, and the electrons, are obtained from time-dependent mathematical equations describing the chemical and physical processes controlling the behaviour of the thermal plasma. These equations, for the continuity, momentum and energy balance of the plasma confined within closed magnetic flux tubes, are solved along the axis of an eccentric dipole flux tube connecting the conjugate hemispheres.

For the present application, two major recent amendments to SUPIM have been utilised: (1) allowance for temperature anisotropy in O^+ and NO^+ , with



Fig. 2. SAID event observed after 0942 UT on day 94, 1982

these ions also having separately calculated temperatures (Moffett *et al.*, 1993; Jenkins *et al.*, 1997); (2) inclusion of vibrational excitation of molecular nitrogen R. J. Moffett et al.: Electron temperatures during rapid subauroral ion drift events



Fig. 3. SAID event observed after 0911 UT on day 105, 1982

(Ennis *et al.*, 1995). In amendment (1), the ion-neutral frictional heating, F_{in} , is partitioned into the parallel and perpendicular ion distributions according to ion fric-

tional *heating* partition coefficients, B_{\parallel} and B_{\perp} . Thus $B_{\parallel}F_{in}$ is input to the parallel ion energy balance equation and $B_{\perp}F_{in}$ to the perpendicular ion energy balance equation. The values of B_{\parallel} and B_{\perp} are based on the Monte Carlo cross-sections of Winkler et al. (1992), whose results were expressed in terms of temperature coefficients, β_{\parallel} and β_{\perp} . The coefficients B_{\parallel} and B_{\perp} depend on the composition of the ion and neutral gases. A more detailed explanation is given by Moffett et al. (1993), in which it is pointed out that expressing T_{\parallel} and T_{\perp} in terms of β_{\parallel} and β_{\perp} is valid only in the region where the frictional heating is intense and the ion density relatively high (whereas the $B_{\parallel}F_{in}$ and $B_{\perp}F_{in}$ terms can be used at all altitudes). Typical values for B_{\parallel} and B_{\perp} have been presented by Jenkins et al. (1997). For amendment (2), the inclusion of vibrationally excited N_2 (denoted by N_2^*), the continuity and diffusion equations for the first five vibrational states of N₂ have been solved. The main source of N_2^* is excitation by the high-energy tail of the energy distribution of the thermal electron gas. The main loss process of N₂^{*}, at the altitudes of interest, is through collisions with atomic oxygen.

4 Model runs and results

The model has been used to investigate the processes that may have to be invoked to explain the data. Model results shown here are for 1982 equinox (day 82), an *L* value of 4, $F10.7 = 150 \times 10^{-22}$ W m⁻² Hz⁻¹ and $A_p = 4$. The MSIS-86 neutral model atmosphere (Hedin, 1987) gives the neutral concentrations and temperatures; the neutral air wind velocities are taken from the HWM 90 thermospheric wind model (Hedin *et al.*, 1991); and the solar EUV fluxes from the EUV 91 solar EUV flux model (Tobiska, 1991).

The model was run for 32 model hours, with a corotating flux tube starting at local noon and finishing at 2000 LT. The neutral atmosphere conditions were then held at those relevant to 2000 LT. In order to see the undisturbed evolution of the model ionosphere, a 90min continuation run was carried out with only the corotation velocity and no applied SAID. Starting from the same initial conditions, the undisturbed evolution was followed for 15 min, after which time a sudden westward $\mathbf{E} \times \mathbf{B}$ drift of 2 km s⁻¹ was applied for a period of 30 min, so that comparison can be made with the data of Fig. 1. The drift was then turned off and the subsequent behaviour followed. The time after a continuation run is started is called *elapsed time*. A further set of calculations was made with the use of an additional high-altitude heat source for the electron gas. The heat source was applied linearly about the equator, switched on at the same time as the applied SAID (i.e. at 15 min elapsed time) and lasted for 20 min. In all cases, the model calculations were repeated with the inclusion of vibrationally excited nitrogen.

It should be pointed out that many studies (for example, Spiro *et al.*, 1978; Quegan *et al.*, 1982; Anderson *et al.*, 1993) suggest that a depleted density

in the mid-latitude trough region can be explained by the presence of slow or stagnant flow that is moving poleward in darkness. Thus the background density level, upon which the SAID and/or the high-altitude heat source must act, may be lower than that used in these calculations; we simply have assumed that the background plasma density profile used to define the trough is taken to be constant and equal to the values in the auroral zone poleward of the event and in the corotating plasmasphere equatorward of the event.

The results displayed in Figs. 4–8 illustrate the response of the model ionosphere to the input drift magnitude and atmospheric conditions appropriate to the SAID event of Fig. 1. The ion temperature profiles of Fig. 4 show the elevated ion temperatures, compared with the unperturbed temperature, which are a characteristic signature of the SAID event. The O⁺ gas has a greater temperature anisotropy than the NO⁺ gas because the frictional heat due to the ion-neutral relative motion is partitioned more strongly into the perpendicular distribution for the O⁺ compared to the NO⁺. The curve labelled T'_i shows the composition-weighted average of the three-dimensional temperatures for the O⁺ and NO⁺, i.e.,

$$T'_{i} = \left\{ N(\mathbf{O}^{+}) \left(\frac{1}{3} T_{\parallel}(\mathbf{O}^{+}) + \frac{2}{3} T_{\perp}(\mathbf{O}^{+}) \right) + N(\mathbf{NO}^{+}) \left(\frac{1}{3} T_{\parallel}(\mathbf{NO}^{+}) + \frac{2}{3} T_{\perp}(\mathbf{NO}^{+}) \right) \right\} / \left\{ N(\mathbf{O}^{+}) + N(\mathbf{NO}^{+}) \right\} ,$$
(1)

where $N(O^+)$, $N(NO^+)$ are the O⁺ and NO⁺ concentrations, respectively, $T_{\parallel}(O^+)$ and $T_{\parallel}(NO^+)$ are the parallel temperatures of the ion gases and $T_{\perp}(O^+)$ and $T_{\perp}(NO^+)$ the perpendicular temperatures; here O₂⁺ has been neglected, since it appears to play a minor role in the molecular ion abundance during SAID events (Moffett *et al.*, 1992a). This average ion temperature enters into the electron-ion heat transfer term in the heat balance equations.

The following factors determine the behaviour of the ion temperatures. Frictional heating is the dominant heat input below about 350 km altitude and this is countered by collisional cooling to the neutral gas. In the topside ionosphere the initial temperature increase is dominated by field-aligned advective transport of hot plasma from lower altitudes until the plasma attains a new scale height (Heelis *et al.*, 1993). The time evolution of the ion temperature is determined by the relatively small contributions of heat conduction in the parallel and perpendicular ion distributions, and by collisional heat transfer between the ion gases and between the ions and electrons.

This time evolution is illustrated in Fig. 5 for 420 km altitude, corresponding to the observations of Fig. 1. The average ion temperature, T'_i , is shown because, as already noted, this is the ion temperature that enters into the electron-ion collisional heat transfer term in the heat balance equations. The other ion temperature, that of the O⁺ perpendicular distribution, dominates the temperature measured by the RPA. The satellite is moving almost perpendicular to the magnetic field and, at 420 km altitude, NO⁺ contributes only a few percent to the ion abundance. The agreement of the model $T_{\perp}(O^+)$ and the observed quantity (Fig. 1) is satisfactory.

Also presented in Fig. 5 are the T_e and N_e values during the model SAID event. The electron temperature rises quickly in response to heat transfer from the ion gas at elevated temperature. During the remainder of the event, T_e remains almost constant at around 200 K.



Fig. 4. Model ion temperature profiles at 30 min elapsed time for a drift of 2 km s⁻¹ for atmospheric conditions appropriate to Fig. 1. Shown are parallel and perpendicular O⁺ and NO⁺ temperatures; T'_i is the composition-weighted three-dimensional average [see Eq. (1)]



Fig. 5. Model ion temperatures, electron temperature and electron concentration, for an altitude of 420 km, as a function of elapsed time for the same inputs as those of Fig. 4 (without an additional heat source at high altitudes)

Examination of the heat balance terms for the electron gas (results not shown) reveals that this value of T_e and those of the F-region represent a balance between ionelectron and electron-neutral collisional heat transfers. The decay in the electron density is caused by increased loss rate of O⁺ due to increased effective temperature for O⁺ – N₂ and O⁺ – O₂ collisions.

The results in Figs. 4 and 5 were obtained without allowance for vibrational excitation of N₂. When the runs were repeated with the inclusion of vibrationally excited N₂, the results (not shown here) differed very little. This is readily explained by noting that the T_e values of Fig. 5 are not high enough to give significant amounts of N₂^{*}; it was found that the abundance of N₂^{*} was less than 5% at 420 km altitude at 45 min elapsed time. Comparison of Figs. 1 and 5 shows clearly that, while the model results cannot be expected to mimic the observational results exactly, the disagreement in the maximum value of T_e is severe.

The effects on T_e of the uncertainty in the calculated values of N_e can be assessed both qualitatively and quantitatively. If the initial values of N_e are increased, the modelled trough values will be far from agreement with the observed values. If the initial values of N_e are decreased, ion-electron heat transfer will be affected more (since it is proportional to N_e^2) than electronneutral heat transfer or electron thermal conduction. Thus the need for an additional heat source would be strengthened. The magnitude of the proposed heat source depends of the local value of N_e and could be lower than that used in our calculations. These qualitative assessments were confirmed by computations that used differing initial N_e and so gave differing N_e results within the SAID event. Detailed results are not presented. To quote one example, however: for results corresponding to Fig. 5, when the N_e minimum is lowered to 4×10^4 cm⁻³, T_e reaches 2000 K within the SAID event.

The obvious candidate to improve the model results is an additional heat source for the electron gas due to ring current-plasmasphere interaction (Kozyra et al., 1987; Chandler et al., 1988; Fok et al., 1991). Heat produced at great altitudes is conducted downwards by the electron gas to the ionosphere, raising the electron temperature and thus sometimes causing SAR arcs. The heat flux to the ionosphere in the absence of the additional heat source is shown in Fig. 6. It is seen that at 1000 km altitude the downward flux is about $10^9 \text{ eV cm}^{-2} \text{ s}^{-1}$. It is found that about an order of magnitude greater flux is required to elevate T_e to the observed level. Such a flux is consistent with the studies by Rees and Roble (1975) and Kozyra et al. (1987; 1990) on SAR arcs. For the results presented in Figs. 6-8, the equatorial heat source was turned on at the time of onset of the model SAID (15 min elapsed time) and switched off at 36 min elapsed time. Figure 6 also shows the boosted downward heat flux.

The effects of this flux are demonstrated in Fig. 7 (without N_2^*) and Fig. 8 (with N_2^*). In both cases, the average ion temperature, T'_i , and the O⁺ perpendicular temperature show little change from their values in Fig. 5, for 420 km altitude. In the F-region and adjacent topside, frictional heating and ion-neutral heat transfer dominate the ion thermal balance. Larger changes are evident, however, in the topside ionosphere (above about 500 km; results not shown) since ion-electron heat transfer contributes to the ion heat balance there.

By contrast, T_e responds significantly to the increased electron heat flux, both directly and indirectly. In Fig. 7



Fig. 6. Model profiles of heat flux in the electron gas: without additional heat source (*dashed line*) with additional heat source (*solid line*)



Fig. 7. As for Fig. 5, but with inclusion of additional heat source for the electron gas

(without N_2^*) T_e peaks at 4000 K; the strong convergence of the heat flux as the T_e altitude gradient changes leads to deposit of heat in the electron gas. The further reduction of the N_e value in the trough (Fig. 7) relative to that in Fig. 5 is due to a change of shape of the N_e altitude profile (results not shown); the level of the F2peak is lowered by the increase in the plasma temperature ($T_e + T_i$). Vibrationally excited N₂ gives a greater loss rate with O⁺ than unexcited N₂, leading to reduced N_e, which decreases the heat capacity of the electron gas, tending to raise T_e again. Thus the effects of inclusion of N₂^{*} (Fig. 8) are to lower the N_e minimum to 5×10^4 cm⁻³ and to raise T_e further, to a maximum of



Fig. 8. As for Fig. 7, but with inclusion of vibrationally excited N_2

4200 K. For the results of Fig. 8, the abundance of N_2^* was about 10% (at 45 min elapsed time at 420 km).

Control results in Figs. 5, 7 and 8 (i.e. with no SAID) also reveal the effects of the additional heat source for the electrons. Figure 5 simply shows that, with neither the additional heat source nor the frictional heat source, T_e remains unperturbed. Figure 8 shows that the additional heat source, with the help of N_2^* , can alone raise T_e to about 3800 K.

The behaviour of T_e around 20 min elapsed time in Figs. 7 and 8 may be tentatively suggested as relevant to the spatial shoulders in the observed T_e (Figs. 1 and 3) already mentioned. With N_e still relatively high, the effect of the downward heat flux shows a tendency to flatten out, whereas at later times, as the N_e trough deepens, the T_e value climbs to about 4000 K. Further investigation is needed to determine whether a further contribution to the increasing T_e within the shoulder is due to energy stored at high altitudes in the plasma-sphere.

Finally, for this set of model results related to Fig. 1, Fig. 9 illustrates the dependence of the T_e and N_e results on the temporal phasing of the additional heat input. If the additional heat input is moved forward by 15 min (now starting at 0 min elapsed time and ending at 21 min elapsed time, while the SAID still begins at 15 min elapsed time), T_e reaches a plateau at about 3750 K. It is noticeable in Figs. 7–9 that the time required for cooling of the electron gas after the SAID event is significantly greater than that required for heating the gas. About 45 min after the event, T_e has decreased to about 3000 K.

Computations have also been performed with a SAID V_{\perp} of 4 km s⁻¹. In this case the high ion-neutral relative velocity causes high ion temperatures (Fig. 10), leading to increased abundances of NO⁺. This is consistent with experimental data and model results



Fig. 9. As for Fig. 8, but with shift in timing of additional heat source (see text)



Fig. 10. Model ion temperatures, electron temperature and electron concentration, for an altitude of 320 km, as a function of elapsed time for input drift of 4 km s⁻¹ (without additional heat source and N_2^*)

presented previously (Anderson *et al.*, 1991; Moffett *et al.*, 1992a). In Fig. 10 is shown the compositionweighted average perpendicular ion temperature, denoted by T'_{\perp} , for 320 km altitude. This is the parameter most likely to be comparable with the ion temperature measured by the RPA instrument (Fig. 2). Also shown in Fig. 10 are T'_i and T_e . The former, defined by Eq. (1), is the composition-weighted average three-dimensional ion temperature. The difference $(T'_i - T_e)$ enters into the rate of collisional heat transfer between ions and electrons. It is striking that, although this difference is very much greater in the 4-km s⁻¹ case compared to the 2-km s⁻¹ case (Fig. 5), the resulting T_e is only slightly greater. It is readily shown that the reduction in heat transfer from ions to electrons is due to the severe reduction in N_e (Fig. 10); ion-electron heat transfer depends on N_e^2 , whereas electron-neutral heat transfer depends on N_e and the neutral concentration.

Results in Fig. 11 show the effects of including an additional heat input and vibrationally excited nitrogen. The heat input was chosen as the same at that used in the 2-km s⁻¹ V_{\perp} case, so that the downward fluxes of heat in the electron gas are similar to those of Fig. 6. As in Fig. 8, T_e exceeds 4000 K by the end of the event. The N_e value is only slightly affected.

5 Discussion

Examples of observed electron temperature behaviour in SAID events have been presented. These demonstrate responses of T_e to particular SAID conditions. Model simulations of SAID events have been performed in order to try to identify the main physical processes that lead to the observed T_e behaviour.

However, when discussing the issues that arise from the model-observational comparisons, two general ca-



Fig. 11. As for Fig. 10, but with additional heat source and inclusion of N_2^{\ast}

veats must be borne in mind. First, the observations have been obtained along the single trajectories of the satellite. Each set of results represents a snapshot of a SAID event over a period of a few minutes. In contrast, the model calculations are performed for an $\mathbf{E} \times \mathbf{B}$ drifting tube of plasma. Care must therefore be exercised in attempting to separate spatial or temporal features. As a flux tube enters the SAID region, the temporal behaviour of the plasma is simulated by the model calculations; the spatial structure at a certain time is made up of the total set of flux tubes entrained in the SAID event at that time. Thus the behaviour of the plasma within a single flux tube is indicative of the behaviour that is likely to be observed, but an exact correspondence is not possible without detailed information on the time evolution and spatial extent of the SAID. Secondly, the model N_e values within the SAID event depend on the values of N_e previous to the event. These initial values are not well known. The period of entrainment of the plasma in the convection stagnation region, prior to the SAID event, is also not known. Thus computations have been performed with differing initial values of N_e ; these results confirm the discussion and conclusions that follow.

The modelled elevation and anisotropy of T_i due to the SAID is in accord with the present and previous observational results and with previous model results. The resulting difference in T_i and T_e leads to heat transfer from the ions to the electrons. The results show that this can raise T_e (from less than 1500 K) to between 2000 and 3000 K for both the $V_{\perp} = 2$ and $V_{\perp} = 4$ km s⁻¹ cases; in the latter case, N_e is likely to be less than in the former case and so compensates (in the rate of heat transfer) for the greater $T_i - T_e$ difference. This modest increase in T_e is (1) clearly compatible with the results within the SAID region of Fig. 2, and (2) perhaps compatible with the SAID results of Fig. 3. It is evident also that the greater elevation of T_e in the SAID region in Fig. 1 requires an additional heat source. The model results demonstrate that the downwards flux of heat usually invoked to explain the occurrence of SAR-arcs (Kozyra *et al.*, 1987) is of the correct order of magnitude to explain the elevated T_e . Furthermore, Caton *et al.* (1996) found that downward heat flow from the magnetosphere was needed as an input in order to obtain agreement with EISCAT topside measurements.

Liu *et al.* (1995) and Caton *et al.* (1996) have also invoked the effects of soft electron precipitation in modelling electron temperature and plasma upflows observed by DE-2 and by EISCAT, respectively. By contrast, the SAID events lie equatorward of the region of electron precipitation that has a source in the plasma sheet. Mechanisms for producing the electric field in the trough may require the SAID to be embedded in the region of weak ion precipitation associated with the inner edge of the ring current. Thus any heat source required to account for elevated electron temperatures in SAID events is identified as arising from a source in the ring current.

The persistence of elevated T_e after the SAID event in the model results bears on consideration of the elevated T_e feature seen equatorward of the SAID in Fig. 2. This feature may arise directly from heating of the electron gas that is occurring at that location and that presumably arises from the ring current. However, the plasma with elevated T_e may have recently emerged from a SAID region, and so the heating may have ceased. With the T_e value about 3000 K in the feature, it is likely that the origin of the heated plasma still requires additional heating either in, or equatorward of, a SAID region.

Allowance for the vibrational excitation of molecular nitrogen by thermal electrons has negligible effect when the increase in T_e is modest. This is to be expected, since the high-energy tail of the electron distribution is insufficiently intense. On the other hand, when the additional heat source is applied and T_e is further elevated, vibrationally excited N₂ plays a noticeable role. The excitation that occurs is sufficient to increase the loss rate of O⁺ with N₂ and so deepen the density trough within the SAID event. There is some positive feedback since lower N_e leads to higher T_e .

6 Conclusions

From comparison of the observational and model results, it is concluded that:

- 1. The ion temperatures observed during SAID events are caused by ion-neutral frictional heating which is preferentially partitioned into the ion perpendicular distribution.
- 2. Heat transfer from the ion gas to the thermal electron gas can raise the electron temperature to between 2000 and 3000 K.
- 3. To obtain greater elevations of electron temperature, an additional heat source for the electron gas is

required. It is suggested that the likely candidate is ring current-plasmasphere interaction; the downward flux heat required in the present study is consistent with that used in modelling SAR arcs and in modelling topside measurements by EISCAT. The persistence of increased electron temperatures after a SAID event raises the possibility that plasma with elevated T_e observed equatorward of a SAID may recently have emerged from such an event.

4. Vibrational excitation of molecular nitrogen by higher-energy thermal electrons can deepen the plasma density trough, associated with the SAID, by increasing the O⁺ loss rate. This in turn allows the electron temperature to increase further.

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