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# High spatial and temporal resolution observations of an impulse-driven field line resonance in radar backscatter artificially generated with the Tromsø heater

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**Abstract.** The CUTLASS Finland HF radar has been operated in conjunction with the EISCAT Tromsø RF ionospheric heater facility to examine a ULF wave characteristic of the development of a field line resonance (FLR) driven by a cavity mode caused by a magnetospheric impulse. When the heater is on, striating the ionosphere with field-aligned ionospheric electron density irregularities, a large enough radar target is generated to allow post-integration over only 1 second. When combined with 15 km range gates, this gives radar measurements of a naturally occurring ULF wave at a far better temporal and spatial resolution than has been achieved previously. The time-dependent signature of the ULF wave has been examined as it evolves from a large-scale cavity resonance, through a transient where the wave period was latitude-dependent and the oscillation had the characteristics of freely ringing field lines, and finally to a very narrow, small-scale local field line resonance. The resonance width of the FLR is only 60 km and this is compared with previous observations and theory. The FLR wave signature is strongly attenuated in the ground magnetometer data. The characterisation of the impulse driven FLR was only achieved very crudely with the ground magnetometer data and, in fact, an accurate determination of the properties of the cavity and field line resonant systems challenges the currently available limitations of ionospheric radar techniques. The combination of the latest ionospheric radars and facilities such as the Tromsø ionospheric heater can result in a powerful new tool for geophysical research.

high latitudes (e.g., Nopper *et al.*, 1982; Kivelson *et al.*, 1984; Baumjohann *et al.*, 1984; Allan *et al.*, 1985; Potemra *et al.*, 1989). Impulsive excitation of the magnetosphere has also recently been the subject of a considerable number of theoretical and modelling studies, principally concerned with the coupling of the field line resonance activity with the compressional mode cavity resonance activity (Allan *et al.*, 1986a, b, 1987; Kivelson and Southwood, 1985, 1986; Inhester, 1987; Krauss-Varban and Patel, 1988; Zhu and Kivelson, 1988, 1989; Lee and Lysak, 1991; Lysak and Lee, 1992; Samson *et al.*, 1992). In discussing these models the phrase “cavity mode” should in some cases be replaced by “wave guide mode”, as an azimuthally travelling wave is assumed (Samson *et al.*, 1992).

Up to the present some of the strongest experimental evidence for the existence of magnetospheric cavity/wave guide modes has come from the inner magnetosphere, the plasmasphere. For example, Sutcliffe and Yumoto (1989; 1991), Yeoman and Orr (1989), Yumoto (1990), Yumoto *et al.* (1990), Lin *et al.* (1991), Yeoman *et al.* (1991) have all proposed that low- and mid-latitude Pi2 pulsations are driven by compressional cavity resonances.

At high latitudes, Crowley *et al.* (1987) studied the damping of a ULF wave observed in ground magnetometer data and by the EISCAT radar, and concluded that energy was being fed into the field line resonance by a global cavity mode. Ruohoniemi *et al.* (1991) and Samson *et al.* (1992) have demonstrated that, on occasions, nighttime and early morning data from the high latitude Goose Bay HF radar, Canada, showed structured spectra with distinct spectral peaks at a discrete set of frequencies. These were interpreted as field line resonances (FLRs) driven by magnetospheric cavity/wave guide modes. Similar structured spectra have recently been detected at sub-auroral latitudes by Provan and Yeoman (1996).

One highly characteristic aspect of the wave signatures predicted from a coupled cavity/wave guide – FLR system is the initial transient wave form. This transient

## 1 Introduction

The transient magnetohydrodynamic response to the impulsive excitation of the magnetospheric cavity has been the subject of a number of observational studies at

occurs as coupling between the magnetosonic wave mode of the resonant cavity and the Alfvénic wave mode of the resonant field line leads to the irreversible transfer of energy from the cavity resonance to the field line resonance. McDiarmid and Allan (1990) produced a numerical simulation of the signature expected from such a coupled magnetospheric cavity/wave guide mode in auroral radar data, such as the Scandinavian Twin Auroral Radar Experiment (STARE, Greenwald *et al.*, 1978) or the Sweden And Britain Radar-auroral Experiment (SABRE, Nielsen *et al.*, 1983). McDiarmid and Allan (1990) noted that a latitude-dependent period occurred early in the wave packet, as the oscillation evolved from a cavity resonance, where phase is constant with latitude, to a FLR which has a characteristic phase change of  $180^\circ$  in latitude. McDiarmid and Allan (1990) also noted that such wave characteristics were very difficult to derive from the simulated radar data.

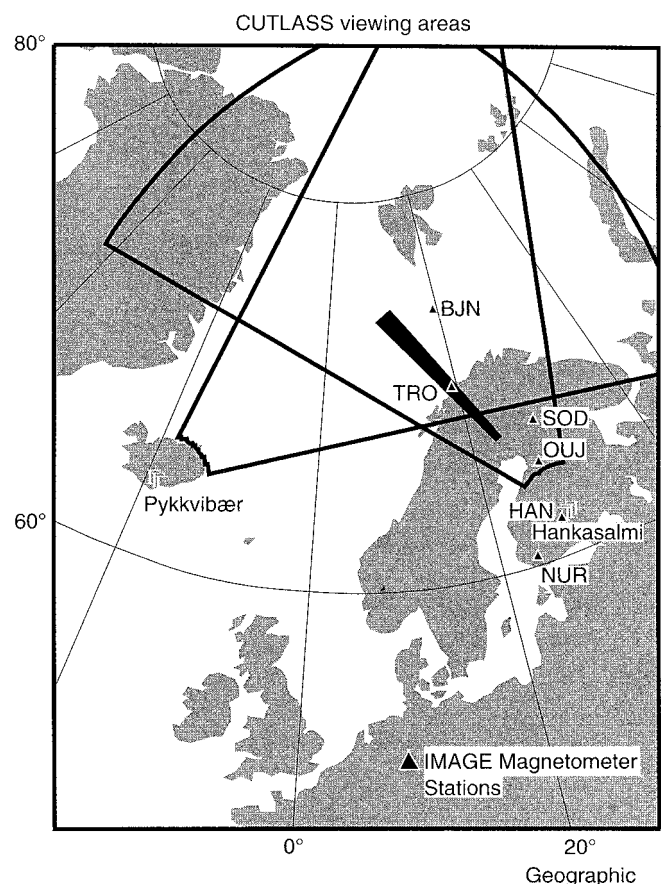
A transient wave signature, reminiscent of the ULF waves simulated by McDiarmid and Allan (1990), was noted in the SABRE Wick radar data by McDiarmid *et al.* (1994). This oscillation was demonstrated to start as a compressional mode wave in spacecraft magnetometer data, and then to evolve into field line resonances in both the morning and afternoon sectors, although the wave structure was quite different in the pre- and post-noon regions.

In the present study, a high-latitude observation of a wave signature characteristic of a coupled cavity/wave guide – FLR is presented. This wave has been detected at very high spatial and temporal resolution by the Cooperative UK Twin Located Auroral Sounding System (CUTLASS) radar, running a SuperDARN discretionary mode experiment (Greenwald *et al.*, 1995). The very high spatial and temporal resolution has been enabled through the generation of field-aligned electron density irregularities with the European Incoherent Scatter (EISCAT) high power RF ionospheric heating facility at Tromsø, in northern Norway (see e.g. Stubbe, 1996). These irregularities then act as an artificially produced target for the CUTLASS coherent scatter radar system.

## 2 Instrumentation

The ionospheric convection velocities in this study are provided by the CUTLASS Finland radar. CUTLASS is a bistatic HF coherent radar, with stations in Finland and Iceland, and forms part of the international SuperDARN chain of HF radars (Greenwald *et al.*, 1995). Each radar of the system is a frequency agile (8–20 MHz) radar, routinely measuring the line-of-sight (l-o-s) Doppler velocity and spectral width of, and the backscattered power from, ionospheric plasma irregularities. The radars each form 16 beams of azimuthal separation  $3.24^\circ$ . Each beam is gated into 75 range bins, each of length 45 km in standard operations, when the dwell time for each beam is 7 s, giving a full 16 beam scan, covering  $52^\circ$  in azimuth and over 3000 km in range (an area of over  $4 \times 10^6$  km<sup>2</sup>), every 2 min. Common-

volume data from the two stations can be combined to provide convection velocities perpendicular to the magnetic field, although no common volume backscatter was detected during the interval under study here. During this interval the Finland radar was operating a non-standard scan mode. In this mode, rather than the usual anticlockwise sweep through beams 15, 14, 13, ..., 0 the Finland radar was confined to a single look direction, beam five. This beam overlies both the EISCAT radar facility and the ionospheric heater at Tromsø. The radar operations at this time also employed a range gating of 15 km, the first gate being at a range of 480 km, and an integration time of 1 s. This arrangement was designed to explore the generation of field-aligned electron density irregularities at very high temporal and spatial resolution. This dataset is supplemented here by data from ground magnetometers, provided by the International Monitor for Auroral Geomagnetic Effects, IMAGE) (Lühr, 1994). These instruments are fluxgate magnetometers, with a sampling interval of 10 s for the interval under consideration here. The IMAGE stations employed in this study, and the location of the high resolution CUTLASS Finland Beam 5 are illustrated in Fig. 1.



**Fig. 1.** Locations of the IMAGE magnetometer stations used in this study and the fields-of-view of the CUTLASS radars. The location of the CUTLASS Finland 15 km resolution single beam experiment is marked as a solid black region

### 3 Observations

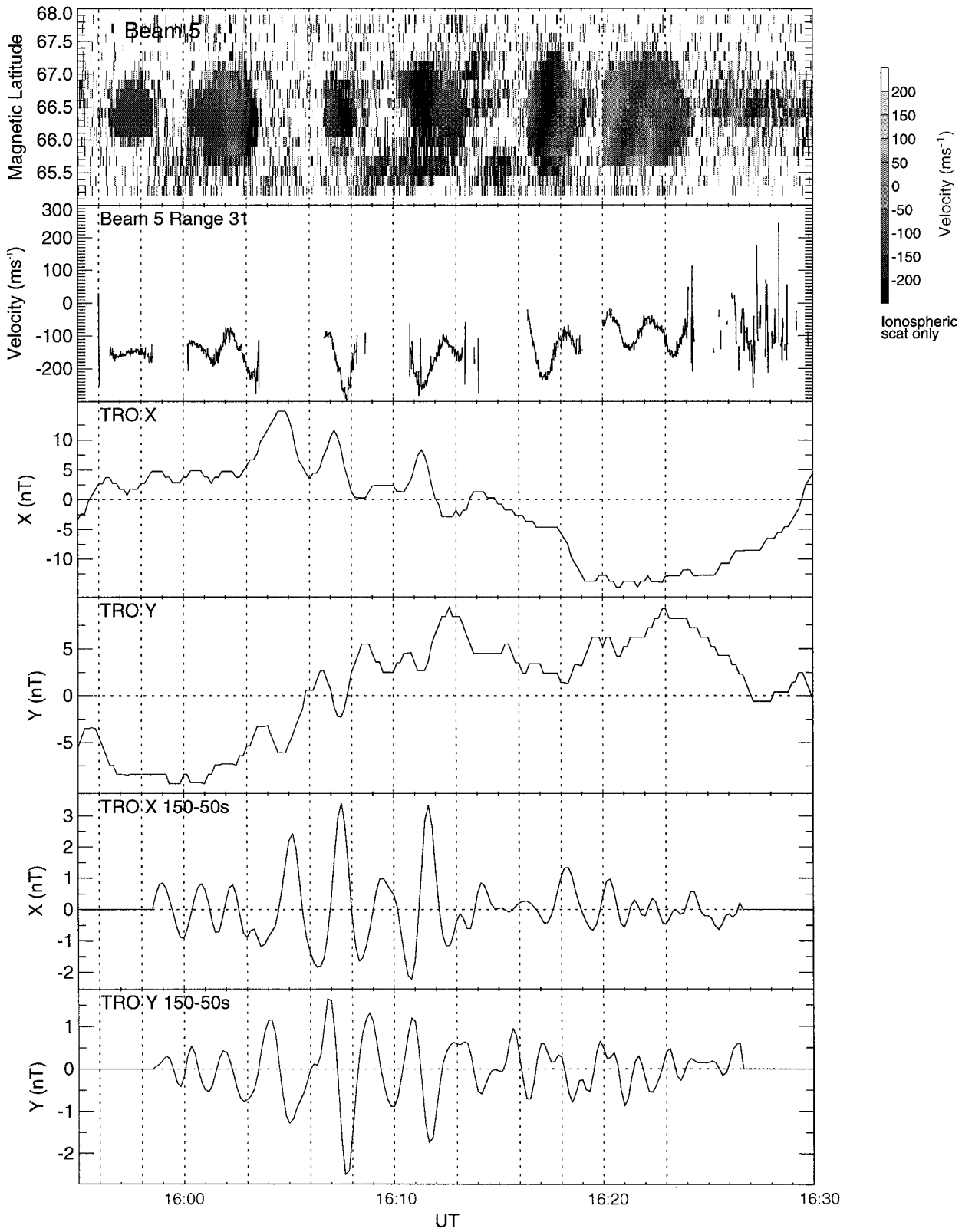
During the interval from 18 to 26 April 1996, a series of experiments were conducted with the EISCAT Incoherent Scatter radar, the Tromsø ionospheric heating facility, the CUTLASS HF radar and a number of low power HF diagnostics. One well-known result of artificial modification of the ionosphere with a high power RF facility such as the Tromsø heater is the generation of field-aligned ionospheric electron density irregularities (see Stubbe, 1996). One of the principal scientific aims of this experimental campaign was the delineation of the spatial and temporal development of these irregularities with the CUTLASS coherent scatter radar, which will backscatter from irregularities of wavelength of order 10–20 m, depending on the operating frequency. The results of these experiments will be presented elsewhere (T. R. Robinson, G. E. Bond, personal communication, 1996). Coherent backscatter from these artificially induced targets is recorded over a region where the heater power is sufficient to generate irregularities. The power backscattered is proportional to the mean square deviation of the electron density associated with the artificial irregularities. Backscatter signal levels obtained from artificial irregularities appear to be significantly higher than those from natural irregularities and this allows a reduction in the integration period of the radar. Integration periods of only one second have been possible, giving observations of regions of backscatter of  $\sim 150$  km extent in range with signal strengths in excess of 30 dB.

During the experimental campaign, between 1600 and 1630 UT on 23 April, quite by chance, an opportunity arose to study a naturally occurring magnetospheric ULF wave within the region of artificially generated radar backscatter. The Tromsø heater was operating at frequencies between 3.90 and 4.01 MHz, and its interaction height in the ionosphere is estimated at  $\sim 200$  km. Unfortunately, the EISCAT radar had ceased operation at 1600 UT. Although modification of the D and E region ionosphere at ULF frequencies can be used to artificially generate a magnetic response of up to 10 nT (see Stubbe, 1996 and references therein), in the current experiment of F region heating the ULF wave is clearly of natural origin. Some evidence for observations of naturally occurring travelling ionospheric disturbances with HF diagnostics scattered by the Sura ionospheric heater has been presented by Blagoveshchenskaya and Troshichev (1996). They also observed wave activity in the ULF band, but attributed these to artificially generated waves. Although the experimental arrangement in the present study was optimised for plasma physics studies, the data collected offer a unique opportunity to study an important naturally occurring geophysical phenomenon with a very high spatial and temporal resolution. Figure 2 presents an overview of the radar and magnetometer data during the interval. A greyscale representation of the Finland line-of-sight velocity data is given in the top panel, with flow toward the radar (equatorward flow, positive Doppler velocity) plotted in light shading and flow away from the radar

(poleward flow, negative Doppler velocity) plotted in dark shading. Patches of radar backscatter are observed, covering a spatial region of  $\sim 1.5^\circ$  of geomagnetic latitude (throughout this study altitude adjusted corrected geomagnetic (AACGM) coordinates, based on Baker and Wing, 1989 are employed). The backscatter patches occur during periods when the Tromsø heater is operational (marked with vertical dashed lines). During this interval the heater operated the following sequence. A 2-min “tuning” period, during which the heater power was gradually increased, was followed by a 2-min “off” period. This was then followed by a 3-min full power heater “on” period, and then a 3 min “off”. This sequence was repeated three times. Note that the artificial backscatter does not start immediately after the onset of a tuning cycle, as irregularities are only generated when the heater power has reached a threshold value. In addition, the artificial backscatter persists after the heater has been switched off, as the irregularities generated have a characteristic decay time. The latest results suggest that the heater-induced irregularities detected by CUTLASS decay to a level of 3 dB in  $\sim 100$  s (G. E. Bond, personal communication). The second panel presents a time series for a single range gate, gate 31. Sections of the ULF wave forms can be clearly distinguished in this plot, with the background ionospheric drift velocity of  $\sim 150$  m s $^{-1}$  being modulated by the ULF wave to values between 100 m s $^{-1}$  and 300 m s $^{-1}$ . The ULF wave signature persists with a peak amplitude of between 100 and 200 m s $^{-1}$  for the interval 1600–1624 UT. The lower four panels of Fig. 1 present the  $X$  and  $Y$  component magnetic field measurements from the Tromsø (TRO) magnetometer which lies underneath beam 5, gate 31 cell of the CUTLASS Finland radar in the discretionary radar programme running at this time. A ULF wave can be seen both in the filtered and unfiltered data from TRO at a similar frequency to that observed in the radar data. The wave signature in the magnetometer data is strongest in the interval 1603–1613 UT. Fourier analysis of the wave signature at TRO reveals a dominant wave in the  $X$  component with a period of 143 s (7 mHz).

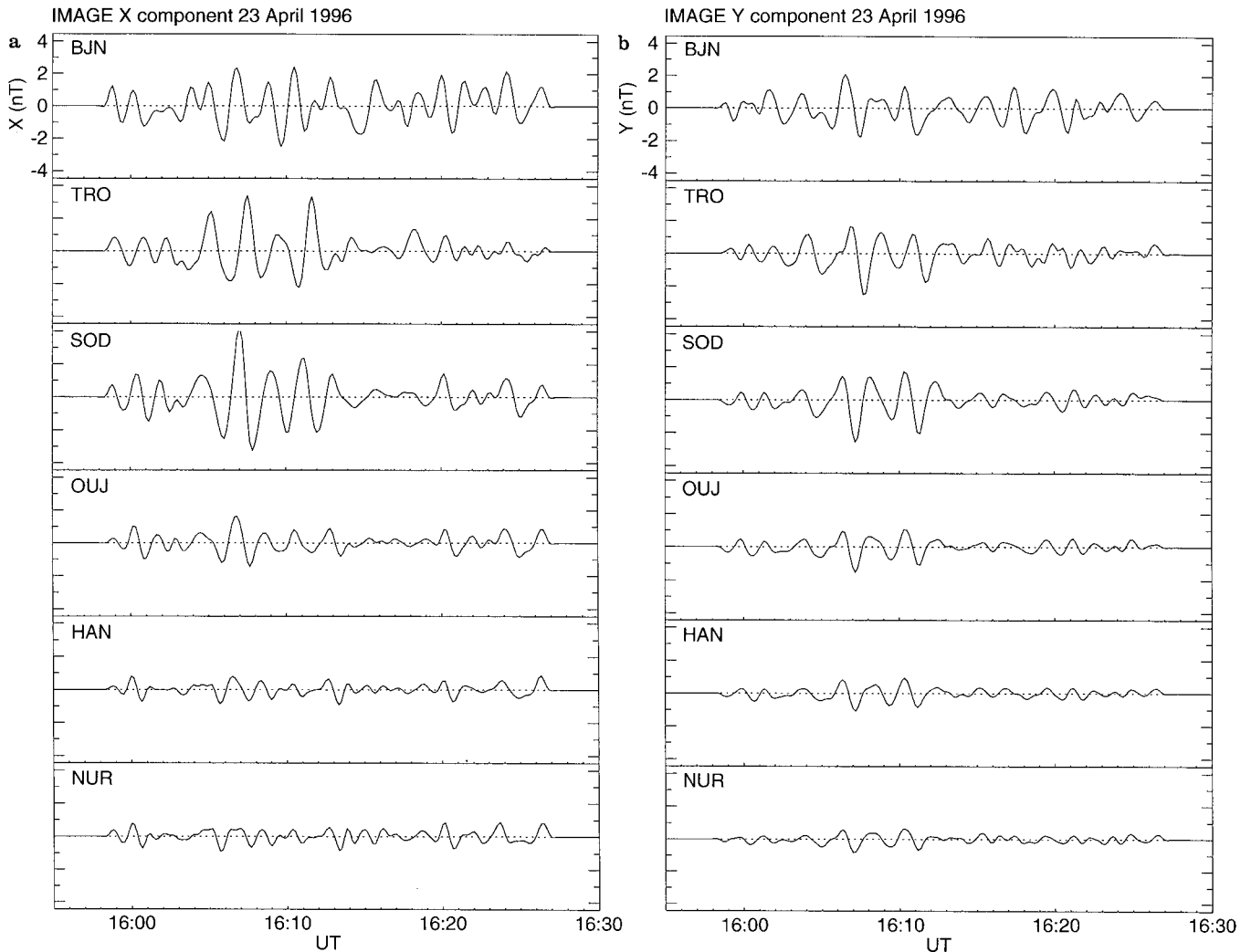
The ULF wave in ground-based magnetic field measurements is examined in more detail in Fig. 3, which displays  $X$  and  $Y$  component magnetic field data from six IMAGE magnetometer stations, covering magnetic latitudes from  $56.8^\circ$ – $71.4^\circ$  geomagnetic latitude, bandpass filtered between 150 and 50 s (6.67–20.0 mHz). A coherent wave packet can be seen across the magnetometer chain, with maximum amplitude of  $\sim 10$  nT (peak-to-peak) observed in the  $X$  component magnetic field between the latitudes of TRO and SOD. Fourier power spectral analysis of the magnetometer data is presented in Fig. 4. This confirms that the dominant frequency is generally 7–8 mHz, and that the largest power at this frequency is found in the  $X$  component at SOD and TRO. Examining the relative phase of the wave packets in Fig. 3 near the time of their maximum in amplitude, at 16:07 UT, the  $Y$  components can be seen to be approximately in phase across the stations. In contrast, a phase change in excess of  $180^\circ$

SUPERDARN / IMAGE 23 April 1996  
 CUTLASS Finland and Tromsø



**Fig. 2.** An overview of the radar and magnetometer data during the interval. A greyscale representation of the Finland line-of-sight velocity data is given in the top panel (positive velocities are towards the radar). Patches of radar backscatter are observed during periods when the Tromsø ionospheric heater is operational (marked with vertical dashed

lines). The second panel plots a time series for a single range gate, gate 31, which overlies the Tromsø magnetometer. Filtered (150 – 50 s, 6.67 – 20 mHz) and unfiltered *X* and *Y* component magnetic field measurements are presented in the lower four panels. A correlated ULF wave signature can be seen in data from both instruments



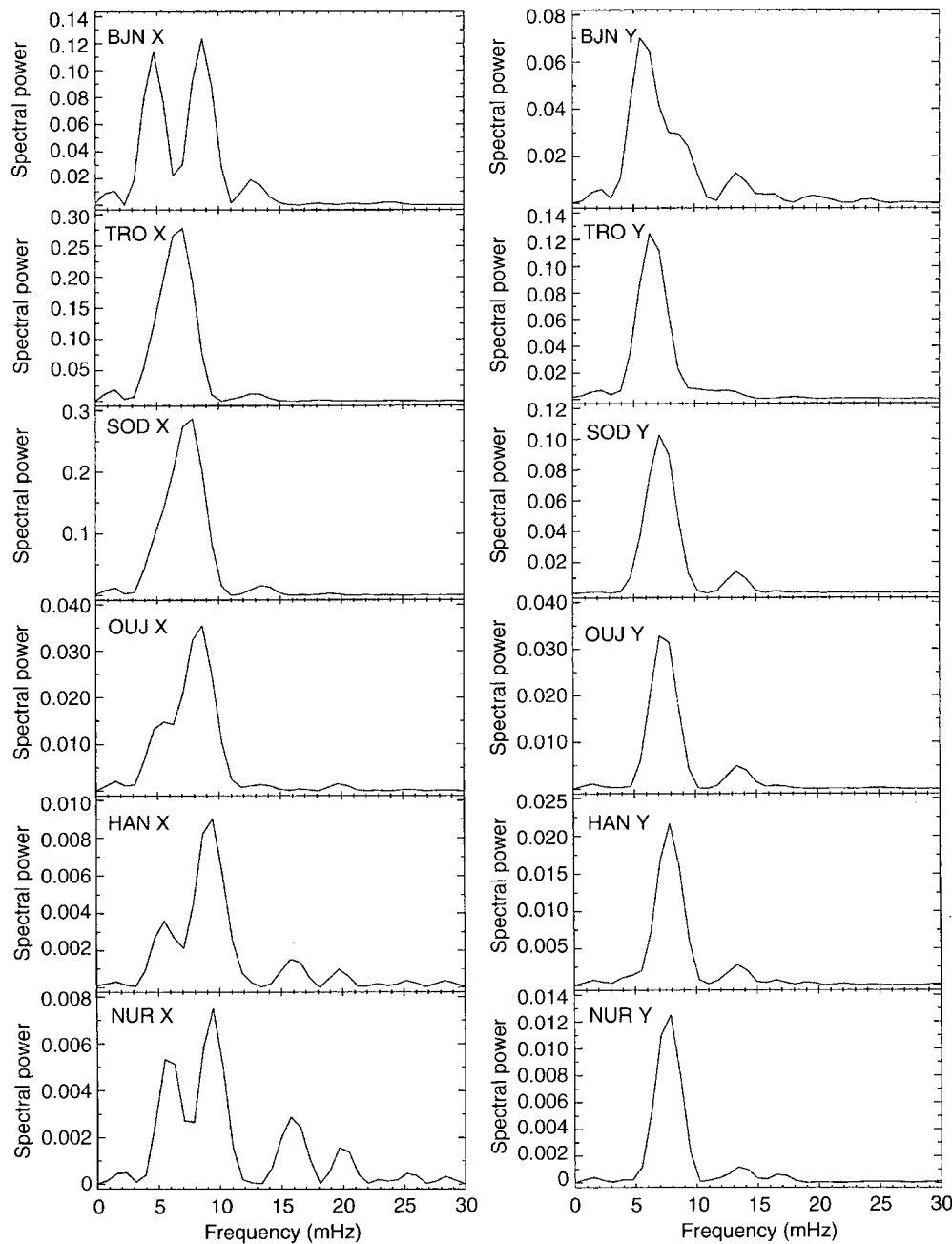
**Fig. 3 a,b.** Magnetic field data from six IMAGE magnetometer stations, bandpass filtered between 150 and 50 s (6.67–20.0 mHz), for **a** the  $X$  component and **b** the  $Y$  component

can be seen in the  $X$  component field measurements, with the lower latitude stations leading in phase.

The latitude variation of the line-of-sight CUTLASS data is displayed in Fig. 5. Stacked time series of Finland line-of-sight velocity data for nine 15 km range gates (gates 28–36, ranges 900–1035 km) are included. Although the analysis of the ULF wave, which has a period of  $\sim 2$  min, is limited by the heater cycle, which restricts the available data to 2- and 3- min sections, the very high temporal and spatial resolution of the data enable a detailed picture of the wave parameters to be obtained for a number of intervals during the wave packet. The ULF wave signature is well defined over the 135 km of latitude illuminated by the Tromsø heater. The wave is of a fairly constant amplitude over the nine range gates early in the wave packet, but has a peak amplitude in the equatorward range gates by 1620 UT. The early wave cycles are also in phase over the nine range gates, whereas by the later parts of the wave packet a phase change of  $\sim 180^\circ$  can be seen over the 135 km separation, with the lower latitude range gates leading in phase. The characteristics of these wave

packets in the radar data are presented in detail in Fig. 7.

The spatial and temporal variation of the wave parameters in the ground magnetometer and ionospheric drift velocity measurements will now be examined in more detail. In Fig. 6 wave parameters derived from the magnetometer data presented in Fig. 3 are examined. Panels a and b present the amplitude and phase of the dominant Fourier component of the  $X$  and  $Y$  component magnetometer data. The magnetic field data for the entire ground-observed wave packet, between 1557 and 1618 UT, bandpass filtered between 150 and 100 s (6.67–10.0 mHz) has been included in the Fourier analysis, and the average frequency of peak spectral power over the array is 7.9 mHz. The spectral amplitude and phase at this frequency are presented. Analysis of this spectral component reveals a pattern typical of a classical FLR signature. There is a clear peak in spectral power in latitude, in this case at SOD, with the  $X$  component the dominant component in the ground magnetic field. The spectral phase is fairly constant in latitude for the  $Y$  component magnetic field, but has a

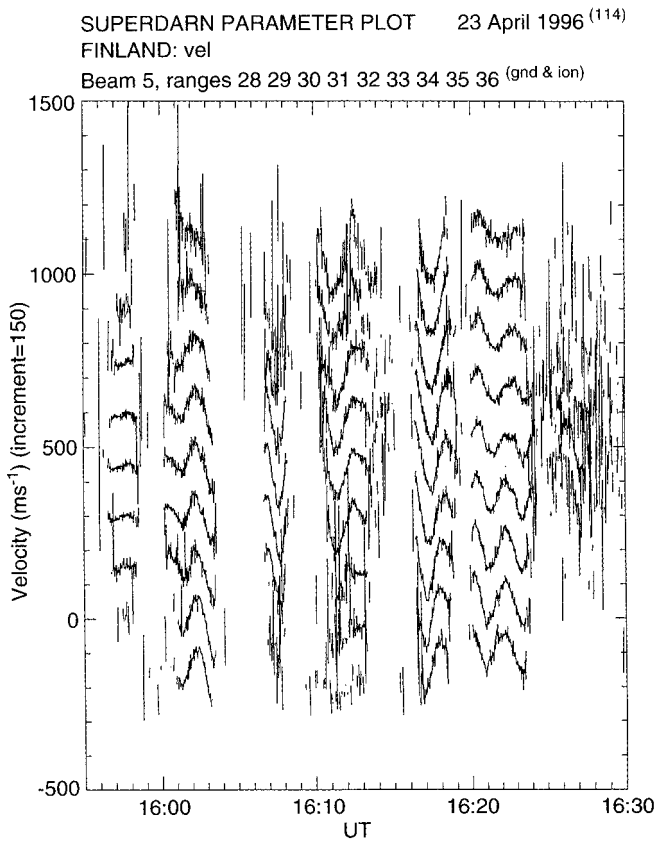


**Fig. 4.** Fourier power spectral analysis (linear arbitrary units) for the magnetometer data presented in Fig. 3, in this case detrended with a 300 s (3.3 mHz) high pass filter

clear lead at the lower latitude stations in the *X* component. The latitudinal phase change in the *X* component measured here is, however, almost  $360^\circ$ , rather than the  $180^\circ$  predicted by field line resonance theory (Southwood, 1974; Chen and Hasegawa, 1974). The horizontal shaded bar shows the latitudinal extent of the radar data presented in Figs. 5 and 7.

Figure 7 presents wave parameters derived from the CUTLASS radar line-of-sight velocity data. Following McDiarmid and Allan (1990), wave parameters are calculated directly in the time domain from the velocity measurements, with the time shifts between the wave peaks and troughs being used to derive wave period and phase. This technique was found to be the most useful by McDiarmid and Allan (1990) and, in addition, allows

the effect of the heater cycle to be deconvolved from the wave parameter analysis. In each panel, wave parameters have been derived for five short sections of the CUTLASS velocity timeseries during heater “on” intervals at 1601–1603, 1606–1608, 1611–1613, 1616–1618, and 1621–1622 UT. Figure 7a presents the calculated wave periods. There is, in general, no clear variation of wave period with latitude, but the profile measured at 1611–1613 UT shows the strongest latitude variation. The wave phase, relative to the lowest latitude range gate available for each time interval is shown in Fig. 7b. This analysis clearly shows the evolution of the wave packet from a latitudinal sequence of in-phase time series from 1601–1613 UT to a sequence displaying the characteristic latitudinal phase variation of a FLR (Southwood,



**Fig. 5.** Stacked time series of Finland line-of-sight velocity data for nine 15 km range gates (gates 28–36, ranges 900–1035 km). Range gates 28–35 are used to derive the wave parameters in Fig. 7

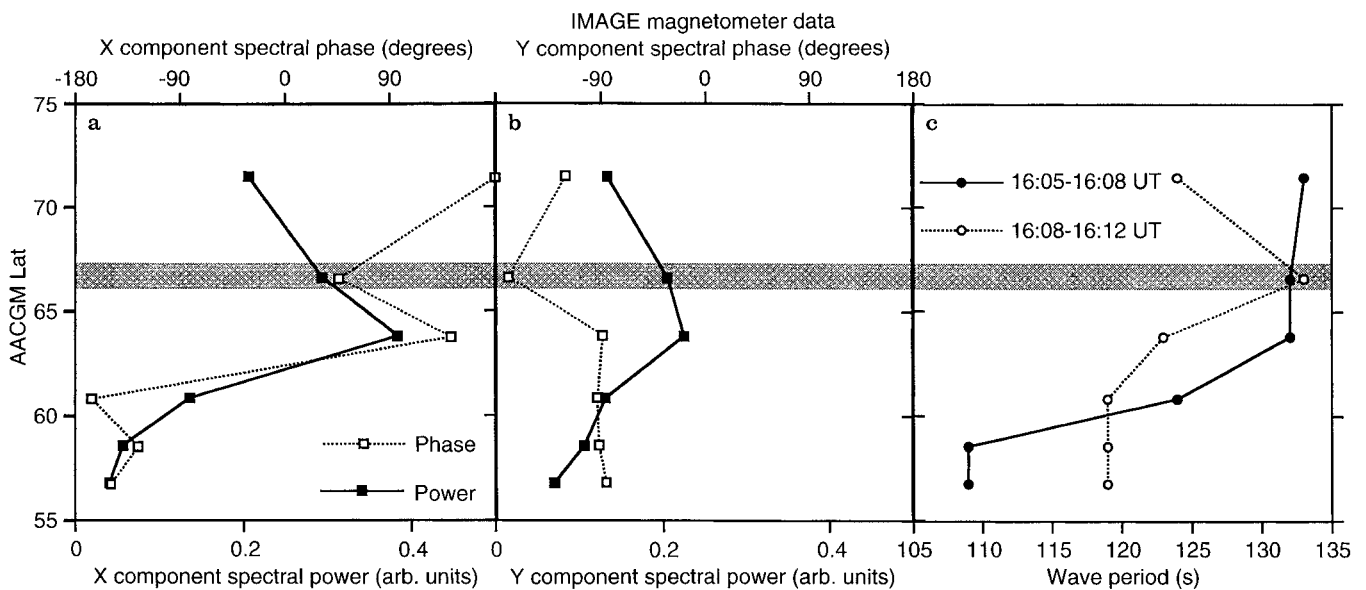
1974; Chen and Hasegawa, 1974). Figure 7c displays the peak-to-peak wave amplitude. Again a clear evolution can be seen during the wave packet, from a latitude-

independent profile early in the oscillation to a sharply peaked resonant structure at 66.25° geomagnetic latitude. This transition appears to be surprisingly sharp, occurring between 16:13 and 16:16 UT, although this may, in part, be an artefact of the observational constraints imposed by the heater cycle.

Wave period estimates have also been derived from the ground magnetometer data using the technique employed for Fig. 7. Figure 6c presents the dominant wave period in the horizontal ground magnetic field data for two intervals during the wave event, one early in the wave packet (1605–1608 UT) and one late in the ground-observed wave packet (1608–1612 UT). The wave period can be seen to have a distinct latitudinal variation, with the field oscillating at a longer period at higher latitudes in the early part of the ground magnetometer recordings. However, this trend becomes less apparent as the wave packet develops.

#### 4 Discussion

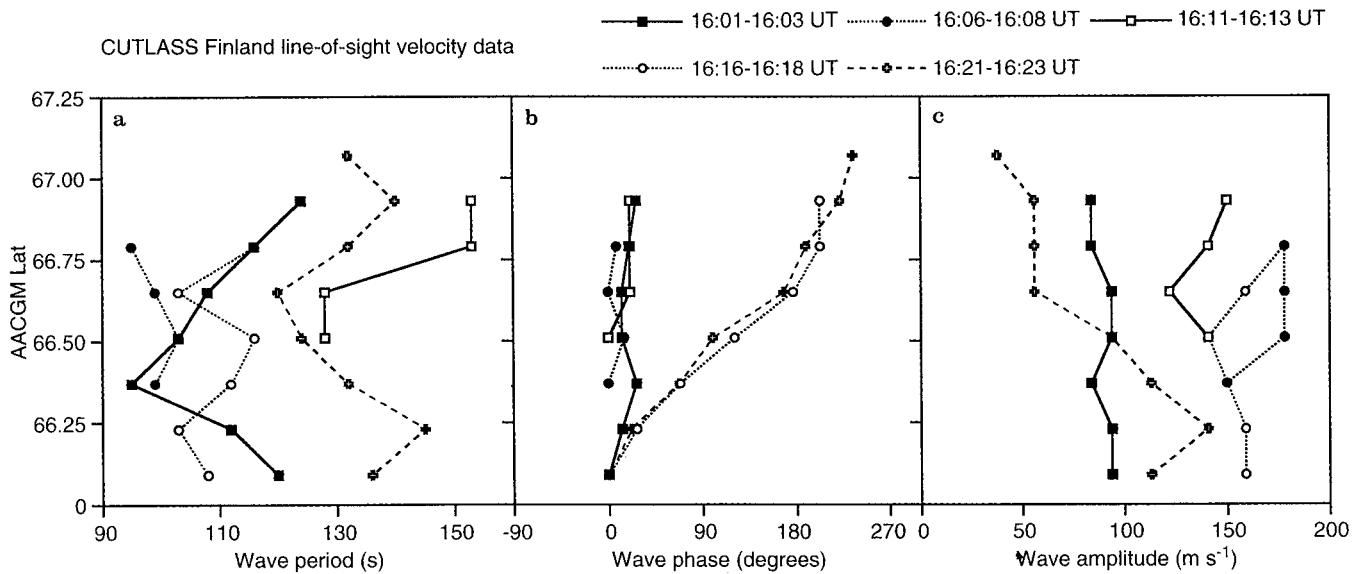
The combination of the Tromsø heater and the CUTLASS Finland HF coherent scatter radar has enabled a high temporal and spatial resolution study of a high-latitude ULF wave. The high resolution study possible with the radar data has been compared with an analysis of the lower resolution, but more spatially extensive, magnetometer data. The wave has the characteristics of the transient signature of a coupled cavity/waveguide – FLR driven by an impulse on the magnetopause. The period of the observed wave suggests that it is a harmonic of the local field line eigenperiod (Poulter *et al.*, 1984). A similar wave observed by McDiarmid *et al.* (1994) was demonstrated to be most probably a second harmonic. The observed wave signatures can



**Fig. 6 a–c.** Wave parameters derived from the magnetometer data presented in Fig. 3. Fourier spectral amplitude and phase at 7.9 mHz are presented for **a** the X component and **b** the Y component.

**c** Shows the dominant wave period for two intervals during the wave event (see text for details). The *horizontal shaded bar* shows the latitudinal extent of the radar data presented in Fig. 5 and 7





**Fig. 7 a–c.** Wave parameters derived from the CUTLASS radar line-of-sight velocity data presented in Fig. 5. **a** Wave periods, **b** wave phase relative to the lowest latitude range gate and **c** peak-to-peak

wave amplitude. Each panel presents parameters derived from five sub-intervals of the wave packet at 1601–1603, 1606–1608, 1611–1613, 1616–1618, and 1621–1623 UT (see text for details)

most readily be compared with the simulation of an impulse driven field line resonance in VHF coherent radar data presented by McDiarmid and Allan (1990). They used the output of the time-dependent numerical model of Allan *et al.* (1986b) to model the development of a FLR driven by a cavity mode caused by a magnetospheric impulse. The most characteristic signature of these simulations was the evolution of the response from the constant period-constant phase characteristic of a cavity resonance, through a transient, latitude-dependent frequency region, where each field line oscillated at its natural frequency, to a resonant signature with an amplitude peak accompanied by a  $180^\circ$  phase change in latitude, as the energy of the cavity/waveguide mode was channelled into a FLR. The simulation showed that the latitude-dependent pulsation period was observed most strongly in the early wave cycles of the oscillation, and had almost entirely disappeared by the latter half of the modelled wave. Another important result of this modelling was the prediction that even with the good spatial (20 km) and temporal (20 s) resolution of the STARE/SABRE type radars, the transient features of these resonances would be difficult to observe. The difficulty in making accurate observational studies of such wave phenomena is clearly illustrated in the contrast between Fig. 6 and 7. The magnetometer data suggest a latitudinally broad, resonant structure. However, this is a consequence of the temporal evolution of the signal during the wave packet and the spatial integration of the ground magnetometers (see Sect. 4.3). The radar data reveals that the resonance is actually confined to a very narrow region (Fig. 7; the latitude region of this figure is indicated in Fig. 6 by the horizontal shaded bar). The oscillation will be discussed in detail in three sections: (a) the initial cavity oscillation, (b) the coupled transient and (c) the driven field line resonance.

#### 4.1 The cavity resonance

The initial wave cycles of the transient pulsation observed here are most characteristic of the impulse-driven cavity/wave guide resonance. The amplitude and phase are relatively constant with latitude, as expected for a large-scale cavity resonance. This section of the wave packet appears strongly in the ground magnetometer data. The early wave packet does not have a simple damped envelope, however, as might be expected for an impulse-driven resonance. Rather it grows between 1601 and 1608 UT in both the ground magnetometer data and the ionospheric velocity data, before damping and coupling to a FLR reduce the amplitude.

#### 4.2 The coupled transient

The transition between a cavity resonance-like and a FLR-like signature is characterised by a latitude-dependent period. Latitude-dependent periods, due to magnetospheric field lines freely ringing at their natural frequencies have been observed at high latitude by Poulter and Nielsen (1982) and Allan *et al.* (1985). In the present study this transition region can be seen in the early part of the ground-observed oscillation from 1605–1608 (Fig. 6c), where the wave period varies from 110–130 s. The later parts of the wave form give unclear results. The results from the analysis of the ionospheric velocity time series do not show a clear dependence of period with latitude, although a latitude variation of 130–150 s is observed from 1611–1613 UT.

#### 4.3 The field line resonance

The resonant field line section of the oscillation modelled by McDiarmid and Allan (1990) has a

resonance width of  $\sim 0.6^\circ$  latitude (67 km) and a phase slope of  $160\text{--}195^\circ/\text{latitude}$ . In the simulated radar data most appropriate for HF radars, with no velocity threshold effects, this appeared as a phase slope of  $110^\circ/\text{lat.}$ , with a  $0.8^\circ$  latitude (90 km) width. Probably the best observational candidates for such a wave previously reported were the oscillations studied by McDiarmid and Nielsen (1987) and McDiarmid *et al.* (1994). In the present study, using a radar experiment of spatial resolution 15 km, slightly superior but comparable to the VHF radars, but with much improved temporal resolution (1s rather than 20s), a peak phase slope of  $270^\circ/\text{Lat.}$  is observed and the full width at half maximum (FWHM) of the resonance can be estimated at  $0.54^\circ$  Latitude or 60 km. Previous observations of the widths of field line resonances using the STARE radar (20 km, 20 s resolution, Walker *et al.*, 1979) and the Goose Bay HF radar (45 km, 96 s resolution, Walker *et al.*, 1992) gave results of FWHM  $\sim 100$  km. McDiarmid *et al.* (1994) inferred a resonance width of  $2^\circ$  latitude (220 km) for their observation of an impulse-driven FLR.

The theoretical minimum FLR width in the magnetospheric equatorial plane was evaluated as  $0.005 LR_E$  by Newton *et al.* (1978), corresponding to  $\sim 6$  km in the ionosphere at the latitude of Tromsø, where a separation of 100 km in the ionosphere will map to  $0.59 R_E$  in the magnetospheric equatorial plane. This calculation assumed a light damping due to ionospheric Joule dissipation only, with the height integrated Pedersen conductivity,  $\sum_p = 10$  S. Any additional damping effects will act to widen the resonance.

In the present study, although the EISCAT radar ceased to operate before the occurrence of the ULF wave, at 16:00 UT, data from 15:00–16:00 UT show that the height-integrated Hall and Pedersen conductivities,  $\sum_H$  and  $\sum_p$ , both had a steady value of 3 S. It thus seems reasonable to take this value as appropriate for the wave from 16:00–16:30 UT. These values have been calculated from the EISCAT alternating code electron density profiles in the field-aligned pointing position, along with the IGRF magnetic field model and the MSIS86 model of the neutral atmosphere. The conductivities have been height-integrated between 90 and 200 km. A value of  $\sum_p = 3$  S implies a resonance width in the ionosphere of 30 km and 9 km respectively for a fundamental and second harmonic toroidal mode (based on the curves of Newton *et al.*, 1978). These narrow theoretical resonance widths suggest that damping mechanisms other than ionospheric Joule dissipation are important in the wave event under study. A similar conclusion was reached for an impulse-driven wave by McDiarmid *et al.* (1994). Recently, Mann *et al.* (1995) have suggested that the asymptotic width of a field line resonance coupled to a magnetospheric cavity mode is in fact determined by the damping of the cavity mode rather than the resonant field line, and predicted a resonance width of  $0.4 R_E$  in the magnetospheric equatorial plane at  $L = 7$  (corresponding to 68 km in the ionosphere above Tromsø). This value very closely matches the observations presented here.

It is difficult to confirm these calculations with a direct estimate of the wave damping from the observed time series. For a coupled cavity mode a damping rate of lower than that due to ionospheric Joule dissipation will be inferred, if the cavity mode is still driving the field line resonance. In addition the difficulty in estimating damping rates of transient pulsations from ground magnetic field data has been demonstrated in the modelling of Poulter and Allan (1985). In the example studied here, clearly the ground magnetometer data would give a very deceptive estimate of damping rate. Unfortunately in this case, the driven field line resonance is not detected for long enough for the ionospheric radar data to provide a good estimate of the damping either.

The field line resonance part of the oscillation, although clearly delineated in the radar data, is not apparent in the ground magnetometer data. For such a narrow resonance this is perhaps not surprising. Ground magnetometer signals are attenuated at small-scale lengths, due to the spatial integration of the E region current systems responsible for the magnetic signature over an area of the ionosphere of radius of the same order of the E region height. Hughes and Southwood (1976) demonstrated that a signal attenuation of  $\sim 1$  order of magnitude occurred if the azimuthal scale length of the ULF wave was 50 km. Attenuation increases of a factor of 5 were observed by Yeoman *et al.* (1992) in a comparison of pulsations in ionospheric radar and ground magnetometer data as the azimuthal wave number,  $m$ , increased from 3 to 18.

## 5 Summary

A ULF wave characteristic of the development of a FLR driven by a cavity mode caused by a magnetospheric impulse has been investigated by means of a novel experimental technique. The time-dependent signature of the waveform has been examined as it evolves from a large-scale cavity resonance, through a transient where the wave period was latitude-dependent and the oscillation had the characteristics of freely ringing field lines, and finally to a very narrow, small-scale local field line resonance. The FLR was only 60 km in width and was highly attenuated in the ground magnetometer data.

The characterisation of impulse driven FLRs can only be achieved very crudely by ground magnetometer arrays and, in fact, an accurate determination of the properties of the cavity and field line resonant systems challenges the currently available limitations of ionospheric radar techniques.

The results reported here demonstrate that the combination of the latest ionospheric radars and an ionospheric heater, such as that at Tromsø, provide a powerful new tool for active geophysical research.

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