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# Diffuse solar radiation and associated meteorological parameters in India

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**Abstract.** Solar diffuse radiation data including global radiation, shortwave and longwave balances, net radiation and sunshine hours have been extensively analyzed to study the variation of diffuse radiation with turbidity and cloud discharges appearing in the form of atmospheric over the tropics. Results of surface radiation measurements at Calcutta, Poona, Delhi and Madras are presented together with some meteorological parameters. The monthly values of diffuse radiation and the monthly ratios of diffuse to global solar radiation have been examined, with a special emphasis in relation to the noise level of atmospheric at Calcutta in the very low frequency band. The results exhibit some definite seasonal changes which appear to be in close agreement with one another.

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## 1 Introduction

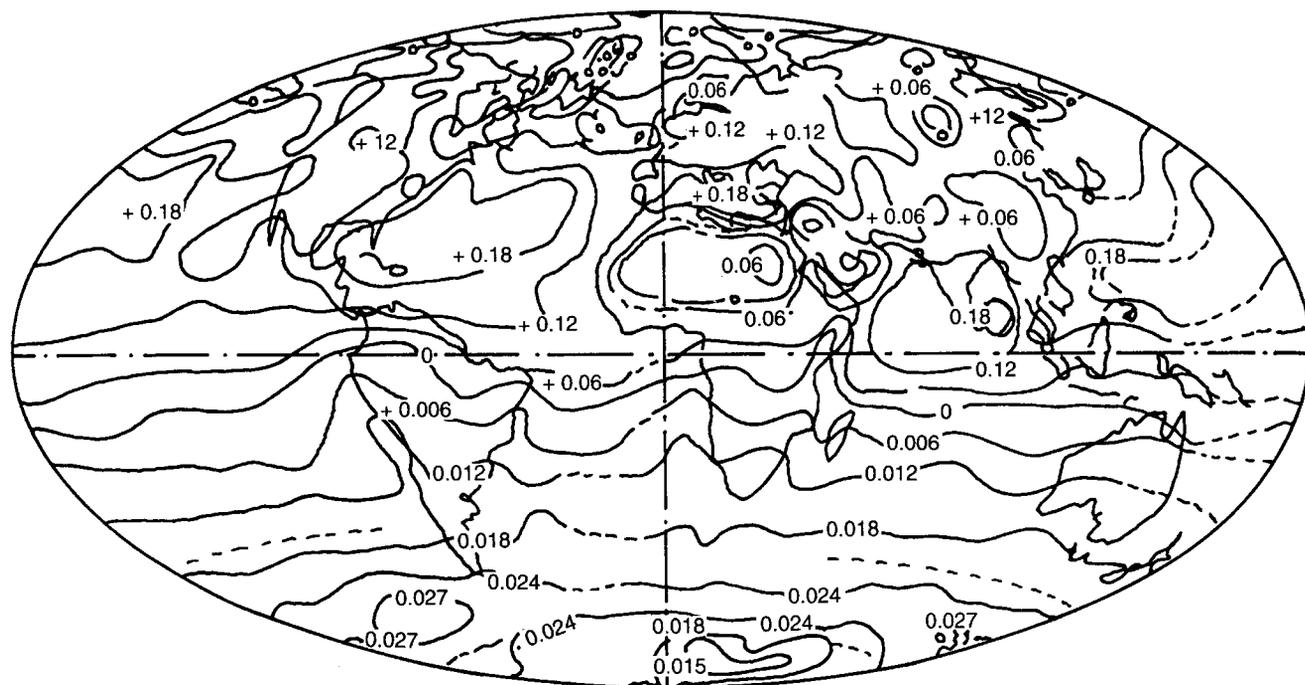
For a proper understanding of the radiation climatology of a region, apart from information on direct and global solar radiation, a knowledge of the diffuse or sky radiation is required. Diffuse solar radiation represents the shortwave energy of solar origin scattered downwards by gas molecules, aerosols (or suspended particulate matter), water vapour and clouds in the atmosphere. The variability in the amount and type of cloud has a dominant role on the value of diffuse radiation. An assessment of the radiation received from the sky assumes great importance in problems concerning the utilization of solar energy, in light studies and in architectural research, particularly in the tropics where the diffuse radiation reaches very high levels. This work presents the results of measurements of diffuse solar radiation in India as recorded at four cities situated in the NW, SW, NE and SE of the country. The relationships between solar activity and meteorological phenomena made an interesting study during the last few

decades, but the data obtained so far are not enough (Ideriah, 1992). Clouds are visible symbols of atmospheric activity and these are the seats of atmospheric electricity (Bhattacharya, 1994). No detailed analysis has yet been made between the solar diffuse radiation, turbidity and cloud discharges. We were thus prompted to study the variation of diffuse radiation with (1) turbidity and (2) cloud discharges appearing in the form of atmospheric over the tropics.

Solar radiation data for a period of 10 years from October 1982 to September 1992, including diffuse radiation, global radiation, shortwave and longwave balances, net radiation and sunshine hours, recorded by meteorological observatories in India have been collected and properly processed. These solar data have been extensively analyzed separately in conjunction with both turbidity and atmospheric data in the very low frequency (VLF) band.

## 2 Instrumentation

The equipment used for the continuous registration of diffuse solar radiation is the same as that for global solar radiation i.e. a Moll-Gorczynski pyranometer and a Cambridge recording millivoltmeter with the addition of a Schuepp shading ring to shade the thermopile element with its two hemispherical glass domes from the direct sun. The shading ring arrangement, the installation of the instrumental equipment, the correction to be applied for the sky inevitably obstructed by the ring, the standardization of instruments and the evaluation records have been carefully done throughout the work. Regular daily observations with Angstrom pyrgeometers have been taken at the four stations with some limitations. It can only be used at night when there is no precipitation and the winds are light. Precautions have to be taken to avoid the effects of temperature differences. The instrument has to be calibrated regularly and despite various practical difficulties regular nightly measurements have been taken at all the



**Fig. 1.** Radiation balance of the Earth-atmosphere system for July over the entire path (after Raschke *et al.*, 1967). The figure indicates

the net radiation flux ( $\times 4.186 \times 10^4$ ) at the top of the atmosphere in  $\text{J} \cdot \text{m}^{-2} \cdot \text{min}^{-1}$

stations. Three tuned radio frequency (TRF) receivers at 10, 20, and 30 kHz were used to record the atmospheric round-the-clock. The receivers used were designed to ensure a large dynamic range with a view to handling a fairly wide range of field intensities.

### 3 Factors controlling net radiation

The radiation balance of the whole Earth-atmosphere system in July is shown in Fig. 1. There is an energy surplus at this time of absorbed solar radiation over emitted longwave radiation everywhere north of  $10^\circ\text{S}$  (Raschke *et al.*, 1967). Maxima of energy occur over the relatively cloudless northern subtropical ocean and a deficit over North African and Arabian deserts. The deficit is due to high values of the albedo (30–40%) and high surface temperatures. Over the entire Southern Hemisphere the radiation balance is found to be negative.

The global average values of albedo, absorbed solar, outgoing longwave and net radiation are of the order as shown below:

Global albedo	29–30%
Absorbed radiation	$1.423 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{min}^{-1}$
Outgoing radiation	$1.444 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{min}^{-1}$
Net radiation for July	$-0.012 \times 10^4$ to $0.029$ $\times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{min}^{-1}$ .

The global distribution of the albedo, the outgoing longwave radiation and the radiation balances agree fairly well with the ground-based and airborne measurements of the same parameters (Rossow and Zhang, 1995). Some preliminary results of the Earth Radiation Budget

Experiment were reported by Barkstrom *et al.* (1990), while calculation of surface and top of atmosphere radiative fluxes was made by Rossow and Zhang (1995). Seasonal variation of the surface radiation budget was reported by Darnell *et al.* (1992). On the other hand, global climatologies of solar radiation budgets at the surface and in the atmosphere were analyzed by Li and Leighton (1993), while re-evaluation of the global energy balance was examined by Ohmura and Gilgen (1993). However, more extensive aircraft and balloon-borne measurements are required, particularly over tropical countries like India.

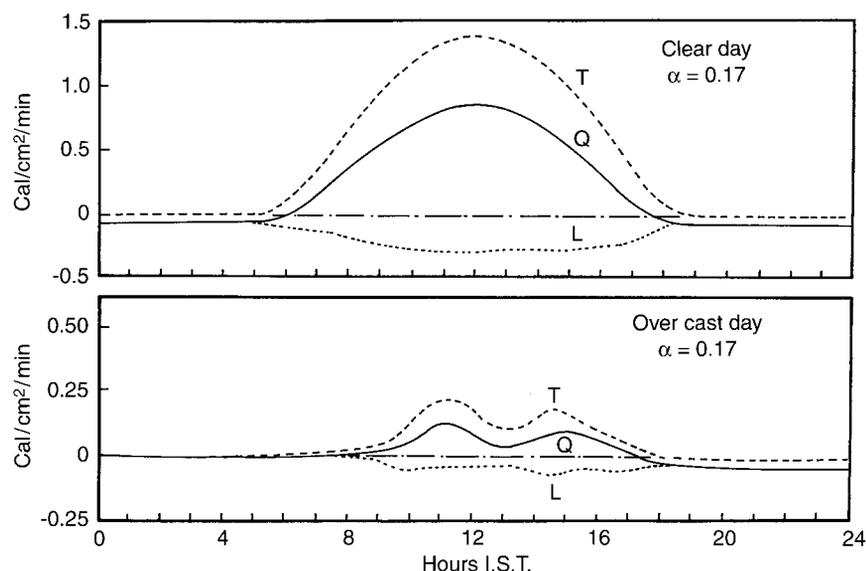
Net radiation  $Q$  is the difference between the total incoming and outgoing radiation fluxes, both shortwave and longwave and is expressed as,

$$Q = R_{S\downarrow} - R_{S\uparrow} + R_{L\downarrow} - R_{L\uparrow},$$

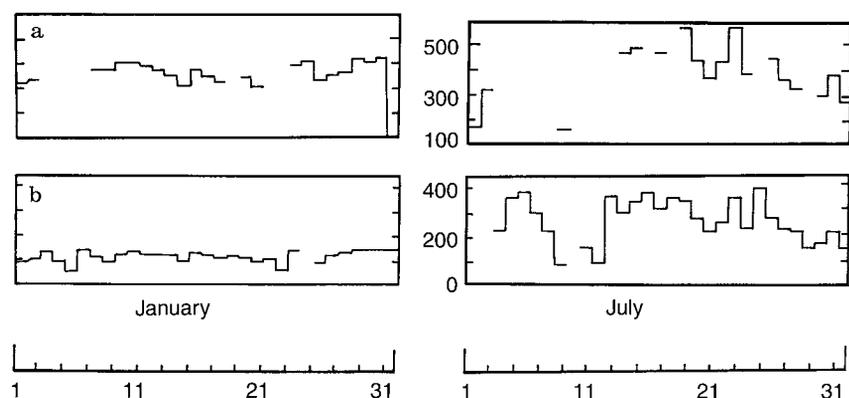
where  $R_S$  and  $R_L$  are the shortwave and longwave radiation fluxes. All fluxes towards the surface are considered to be positive and all fluxes away from the surface as negative.

$R_{S\downarrow}$  is the downward flux of solar radiation at the surface or global solar radiation  $T$ . This is affected mainly by cloud amount and also partly by atmospheric aerosols and atmospheric absorption. The nature of the surface itself has no direct influence on  $T$ , although it has a profound influence on the reflected solar radiation  $R_{S\uparrow}$  and hence on shortwave radiation balance ( $R_{S\downarrow} - R_{S\uparrow}$ ).

The downward terrestrial radiation  $R_{L\downarrow}$  depends on air temperature, on cloud amount and temperature and on the water content in the atmosphere, particularly when the layers are relatively low. The upward terrestrial radiation  $R_{L\uparrow}$  depends on both temperature and emissivity of



**Fig. 2.** Global radiation ( $T$ ), net radiation ( $Q$ ) and net terrestrial radiation ( $L$ ) at Calcutta in Indian Standard Time (IST). The data plotted along the Y-axis is to be multiplied by  $4.186 \times 10^4$  to represent the values in  $J \cdot m^{-2} \cdot min^{-1}$



**Fig. 3.** **a** Global radiation and **b** net radiation ( $\times 4.186 \times 10^4$ ) in January and July at Calcutta in  $J \cdot m^{-2} \cdot day^{-1}$

the Earth’s surface. Hence the net radiation, the difference between the amount of shortwave radiation absorbed by the Earth’s surface ( $R_S\downarrow - R_S\uparrow$ ) and the amount of effective outgoing longwave radiation ( $R_L\downarrow - R_L\uparrow$ ) is rather complex as concerns interacting variables, with no clear separation of “cause” and “effect”. For a given regime of  $T$ , the net radiation and its variations during the day will depend on the nature of the surface.

In India, the surface albedo  $\alpha$  varies with elevation, cloud amount, and the degree of moisture in the soil. Albedo is low when the soil is moist during the monsoon months and high when the soil is dry. Increase in cloudiness reduces the degree of dependence of albedo on the height of the Sun, since increase in the cloud amount decreases direct solar radiation and augments diffuse sky radiation (Chacko *et al.*, 1968; Desikan *et al.*, 1969; Mani and Chacko, 1973).

The diurnal variation of net radiation  $Q$ , on cloudless and overcast days, at Calcutta is shown in Fig. 2. The albedo for the station is estimated from the nature of the underlying surface as 0.17 values of global solar radiation  $T$  and the estimated net terrestrial radiation  $L = (R_L\downarrow - R_L\uparrow)$  are also plotted.  $Q$  is calculated by using the relation  $T(1 - \alpha) + L$ .

On clear days, the curves for all three fluxes,  $T$ ,  $Q$  and  $L$  are smooth and regular and attain a maximum at about noon. Net radiation is negative from sunset to sunrise and shows a constant value throughout the night. On a cloudy day, both global solar and net radiation curves are again phase, but both values are very low, being about  $0.627 \times 10^4$  and  $0.334 \times 10^4 J \cdot m^{-2} \cdot min^{-1}$  respectively. During the second half of the night the net radiation is almost zero. The effect of increasing surface temperatures during the day is noticeable in the net radiation profile, even though the sky is covered with cloud.

Figure 3 shows the daily totals of  $T$  and  $Q$  for the months January and July for Calcutta. These 2 months have been chosen on purpose to examine the characteristic variation during winter and the rainy season. It is seen that, when  $T$  is high, cloudless days are associated with high positive  $Q$ . An increase of cloud reduces both  $T$  and  $Q$  considerably.

### 3.1 Terrestrial radiation fluxes at the ground

The Angstrom pygeometers measure the longwave effective outgoing radiation,  $L$ , as the difference between

**Table 1.** Location of the four stations

Stations	Coordinates		
	Height (m)	Latitude ( $^{\circ}$ N)	Longitude ( $^{\circ}$ E)
Calcutta	06	22 $^{\circ}$ 32'	88 $^{\circ}$ 20'
Poona	559	18 $^{\circ}$ 32'	73 $^{\circ}$ 51'
Delhi	216	28 $^{\circ}$ 35'	77 $^{\circ}$ 12'
Madras	16	13 $^{\circ}$ 00'	80 $^{\circ}$ 11'

longwave radiation flux downward from the atmosphere  $R_{L\downarrow}$  and that upward from the instrument. Measurements of the effective outgoing longwave radiation  $L$  at the ground were made at 2030, 2330, 0230 and 0530 IST (Indian Standard Time) on all nights when there was no rain at the time of observation. The values of  $R_{L\downarrow}$  and  $L$  are considered here for four Indian stations. Of the four stations, Calcutta and Madras are coastal stations where the diurnal variations of temperature are small and water vapour content high throughout the year. Madras is situated in the southern half of the Indian peninsula, while Poona lies further north and closer to the western coast. Calcutta is typical of the northeast region with its hot humid summer followed by the monsoon and mild winter. New Delhi on the other hand, is typical of the semi-arid zones to the north and northwest with its extreme summer and winter conditions. Madras on the east coast has climatic conditions quite different from that of the rest of the country, the northeast monsoon occurring during October and November being a special feature of this area. The coordinates of the four stations with extreme climatic variations are noted in Table 1.

The mean monthly values for  $R_{L\downarrow}$  and  $L$  at 2030 IST for the four stations are shown in Table 2.

#### 4 Diffuse solar radiation

The mean daily values of diffuse solar radiation  $D$  over a period of 10 years for the four stations, i.e. Calcutta,

**Table 3.** Mean daily values of diffuse solar radiation ( $\times 10^4$ ) (in  $\text{J}\cdot\text{m}^{-2}\cdot\text{day}^{-1}$ )

	Calcutta	Poona	Delhi	Madras
January	506.50	397.67	414.41	699.06
February	560.92	401.85	493.94	606.97
March	761.85	565.11	653.01	669.76
April	862.31	703.24	833.01	724.17
May	1029.75	803.71	958.59	908.36
June	1084.17	1163.70	1213.94	1025.57
July	1071.61	1180.45	1059.05	1096.73
August	987.89	1205.56	950.22	1067.43
September	874.87	1025.57	749.29	883.24
October	586.04	590.22	481.39	791.15
November	485.57	464.64	372.55	715.80
December	456.27	414.41	389.29	715.80
Annual	774.41	740.92	715.80	824.64

**Table 4.** Values of mean daily diffuse solar radiation (in  $\text{J}\cdot\text{m}^{-2}\cdot\text{day}^{-1}$ ) during the premonsoon (March–May), monsoon (June–September) and postmonsoon (October–February) seasons with their annual percentages

	March–May		June–September		October–February	
	$D \times 10^4$	(%)	$D \times 10^4$	(%)	$D \times 10^4$	(%)
Calcutta	883.24	36.7	1004.64	41.7	519.06	21.6
Poona	690.69	30.2	1142.77	50.0	452.08	19.8
Delhi	812.08	36.4	992.08	44.5	426.97	19.1
Madras	766.03	30.8	1017.19	40.9	703.24	28.3

Poona, Delhi and Madras, are given in Table 3. Table 4 shows the values of mean daily diffuse radiation during the premonsoon (March–May), monsoon (June–September) and postmonsoon (October–February) seasons and their annual percentages.

It can be seen that, in general, for all the stations, diffuse radiation is highest during the summer and monsoon months and least during the clear winter months. The higher values in summer and monsoon are due to

**Table 2.** Mean monthly values (in  $\text{J}\cdot\text{m}^{-2}\cdot\text{min}^{-1}$ ) of downward longwave radiation flux ( $R_{L\downarrow}$ ) and net outgoing longwave radiation ( $L$ ) on clear nights

	Calcutta		Poona		Delhi		Madras	
	$R_{L\downarrow} \times 10^4$	$L \times 10^4$						
January	2.034	0.464	2.072	0.544	1.871	0.439	2.281	0.406
February	2.180	0.380	2.130	0.611	2.021	0.435	2.310	0.422
March	2.331	0.426	2.394	0.439	2.017	0.523	2.415	0.037
April	2.494	0.385	2.432	0.460	2.327	0.539	2.494	0.385
May	2.582	0.334	2.377	0.502	2.524	0.510	2.545	0.339
June	2.561	0.305	2.264	0.506	2.716	0.372	2.519	0.351
July	2.641	0.221	2.440	0.334	2.746	0.226	2.582	0.297
August	2.595	0.263	2.281	0.318	2.662	0.267	2.528	0.359
September	2.599	0.234	2.335	0.347	2.679	0.322	2.582	0.309
October	2.419	0.318	2.243	0.456	2.197	0.460	2.524	0.313
November	2.130	0.426	2.093	0.498	1.963	0.481	2.272	0.385
December	1.942	0.468	2.076	0.506	1.850	0.447	2.298	0.376
Mean	2.369	0.351	2.260	0.460	2.298	0.439	2.444	0.364

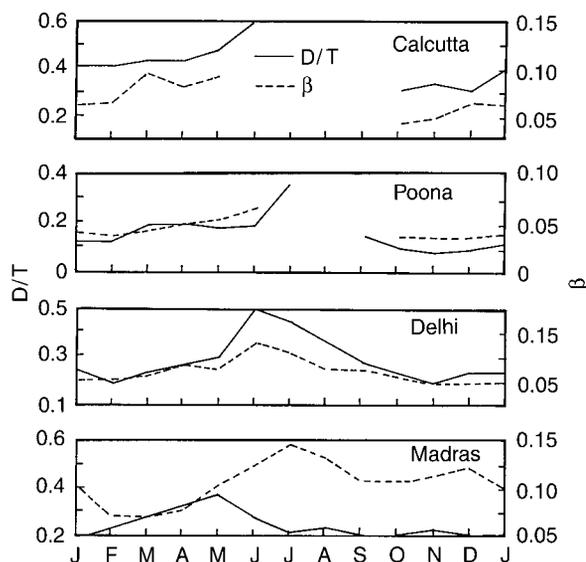
**Table 5.** Ratios of mean daily values of diffuse to global solar radiation for all days and for clear days only. The dashes in the table indicate that a sufficient number of clear days was not available during those months for taking a mean

	Calcutta		Poona		Delhi		Madras	
	All days	Clear days						
January	0.34	0.24	0.21	0.17	0.29	0.20	0.39	0.20
February	0.31	0.25	0.18	0.15	0.28	0.20	0.27	0.18
March	0.39	0.37	0.24	0.18	0.30	0.22	0.28	0.20
April	0.38	0.32	0.27	0.20	0.32	0.26	0.30	0.17
May	0.45	0.36	0.32	0.22	0.36	0.25	0.39	0.29
June	0.63	–	0.55	0.27	0.51	0.36	0.49	0.21
July	0.65	–	0.76	–	0.59	0.32	0.58	–
August	0.64	–	0.74	–	0.56	0.25	0.54	–
September	0.57	–	0.57	–	0.38	0.25	0.44	–
October	0.35	0.17	0.31	0.16	0.25	0.22	0.45	0.17
November	0.32	0.19	0.25	0.15	0.24	0.19	0.45	0.17
December	0.33	0.25	0.25	0.15	0.29	0.20	0.47	0.19
Annual	0.45	0.28	0.37	0.18	0.37	0.23	0.42	0.20

increased turbidity and cloudiness during this period; cloudiness having the greatest influence on diffuse radiation. However, Delhi receives the maximum diffuse radiation in June, the hot summer month, when the atmosphere is very turbid and there are frequent duststorms. The highest values of nearly  $1255.8 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{day}^{-1}$  are recorded during June at Delhi and during August at Poona. At the other two stations, diffuse radiation is of the order of  $1046.5 \times 10^4$ – $1172.1 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{day}^{-1}$  during the monsoon months.

The lowest diffuse radiation values recorded are about  $376.7 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{day}^{-1}$  during the clear winter month, December at Delhi. At other stations the minimum is of the order of  $418.6 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{day}^{-1}$ , except at Madras where cloud-free skies are very rare and the minimum is as high as  $627.9 \times 10^4 \text{ J} \cdot \text{m}^{-2} \cdot \text{day}^{-1}$ . In fact, Madras receives relatively uniform diffuse radiation throughout the year because it is more or less uniformly cloudy throughout the year. The annual values of diffuse radiation are therefore highest at Madras with  $301.4 \times 10^7 \text{ J} \cdot \text{m}^{-2} \cdot \text{year}^{-1}$ . It is seen from Table 4 that about 40–50% of the diffuse radiation is received during the cloudy monsoon months and 30–35% during the turbid summer months. At Madras about 30% of the diffuse radiation is received during the winter season due to cloud formation during the northeast monsoon.

Ratios of mean daily values of diffuse to global solar radiation ( $D/T$ ) for all days and the same ratios for clear days only are given in Table 5. Taking all days, clear as well as cloudy, into consideration, the values of  $D/T$  are seen to vary between 15 and 80%, the maximum values occurring in June–August and the minimum in November–February. This indicates that about 80% of the short-wave radiation received is scattered sky radiation during monsoon months. The variation of this ratio is maximum at Poona, where it varies from 25% in December to 76% in August, and least in Madras, where the range is only from 27% in February to 58% in July. The lower variation at Madras is caused by the general cloudiness throughout the year. Stations which are characterized by clear skies and lower turbidity in winter and higher



**Fig. 4.** The monthly variations of diffuse to global solar radiation ( $D/T$ ) and Angstrom turbidity coefficient ( $\beta$ ) for the four Indian stations

turbidity and severe cloud in summer and monsoon. For example, Poona reveals a very sharp increase in  $D/T$  during the summer and monsoon seasons. Calcutta, with the industrial smoke and turbid and foggy atmosphere presents comparatively higher  $D/T$  values during the winter months.

4.1 Relation between turbidity and diffuse radiation

The mean daily values of the ratios of diffuse to global solar radiation ( $D/T$ ) for clear days and the mean daily values of the Angstrom turbidity coefficient ( $\beta$ ) for the four stations are plotted in Fig. 4. The coefficient  $\beta$  is defined as the extinction due to scattering and absorption of aerosols. On a clear day one may expect  $D/T$  to depend mainly on atmospheric turbidity. It is seen from Fig. 4

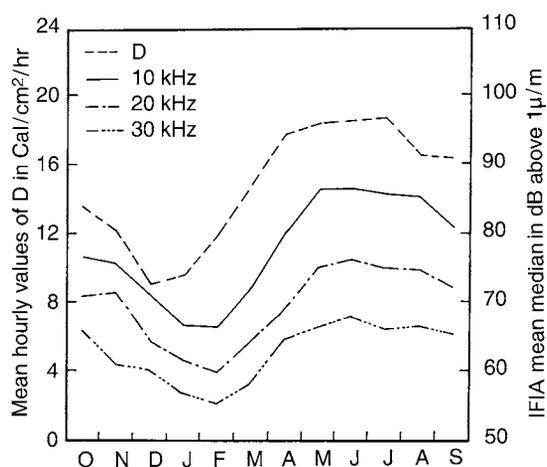


Fig. 5. Mean hourly values of diffuse radiation and median IFIA plotted month by month

that the general trend in the curves of  $\beta$  and  $D/T$  are similar. The trends are, however, a little different for Madras during the period June–August. This might be due to the fact that at Madras pyrheliometer measurements are taken when the Sun's disc and the sky immediately around it are free from clouds while the rest of the sky is cloudy. As a result  $\beta$  values do not correspond exactly to the conditions under which the values of  $D/T$  are determined. There is a close correlation between turbidity and the ratio of diffuse to global solar radiation on cloudless days. Aerosols, contributing as much as 30% of diffuse radiation over the country, play a vital role in the zonal radiation climatology.

#### 4.2 Relation between cloud discharges and diffuse radiation

The mean monthly values of diffuse radiation ( $D$ ) in  $J \cdot m^{-2} \cdot h^{-1}$  over Calcutta are plotted in Fig. 5. The median noise level of the integrated field intensity of atmospheric (IFIA) at the same station corresponding to each observed hour of solar radiation was first noted, and the mean value for each month has been calculated from this. These mean values are also shown in Fig. 5. Figure 5 reveals that the variation of diffuse radiation closely resembles the variation of atmospheric at all the three frequencies, 10, 20, and 30 KHz. The intensities of diffuse radiation and of atmospheric are predominant during the hot summer and monsoon months, May–August, while a minimum value is noticed during the clear winter

Table 6. Correlation coefficients between diffuse radiation and IFIA

Events considered	Correlation coefficient
$D$ and $(IFIA)_{10 \text{ KHz}}$	0.87
$D$ and $(IFIA)_{20 \text{ KHz}}$	0.82
$D$ and $(IFIA)_{30 \text{ KHz}}$	0.79

months, November–February. An analysis of the correlation coefficients between the plots of diffuse radiation and IFIA at each of the frequencies has been made. The results obtained are presented in Table 6.

Table 6 shows that there is a positive correlation coefficient of 0.8 in all the cases.

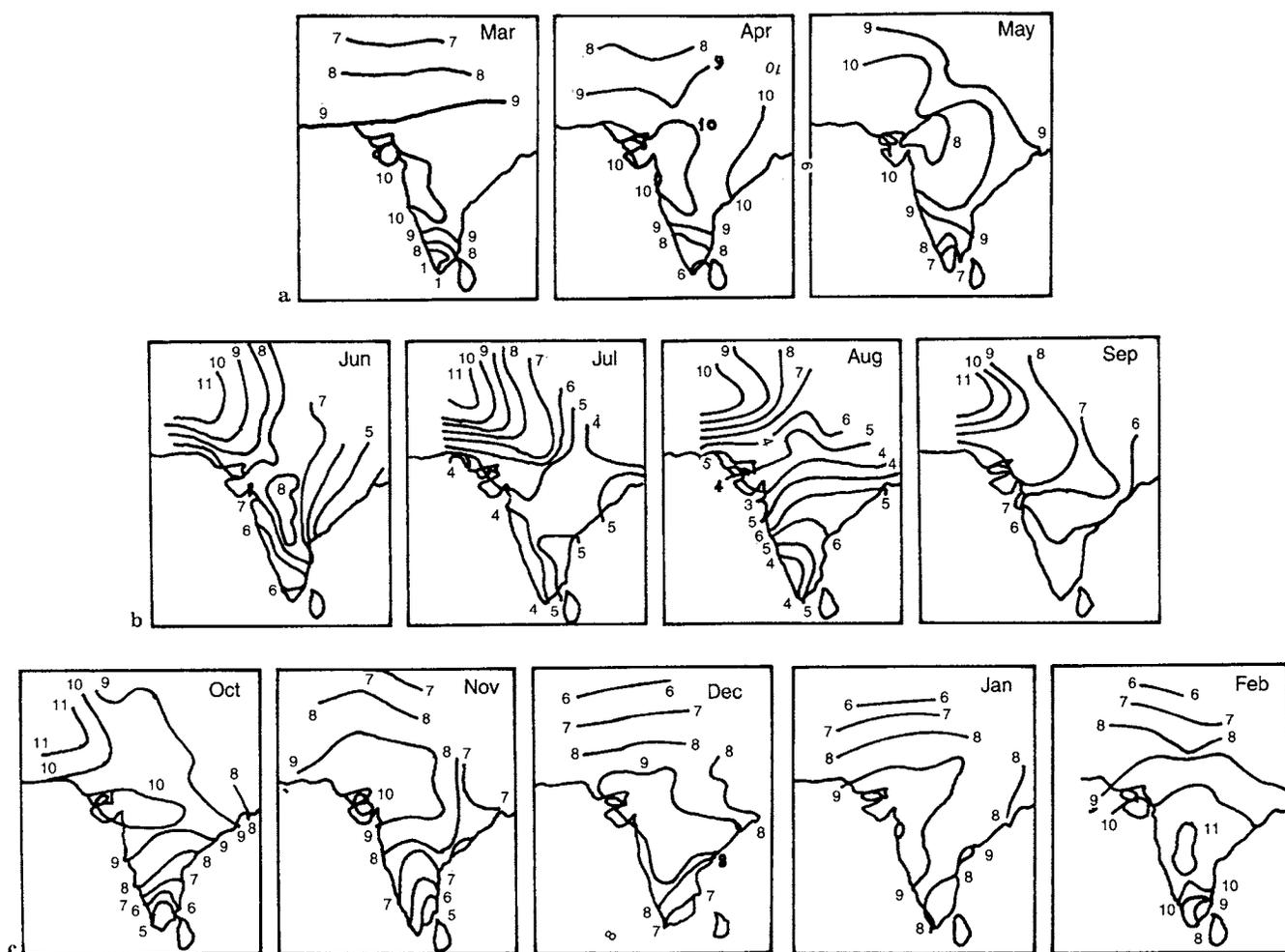
The seasonal values of diffuse solar radiation per hour for the three main seasons, postmonsoon, premonsoon, and monsoon, have been calculated. The results are shown in Table 7. It appears from Table 7 that the diffuse radiation during the premonsoon and monsoon seasons is much higher than during the postmonsoon season. For comparison, the mean hourly median noise level of IFIA at 10, 20 and 30 KHz for the three seasons are also shown in the same table. It is found from Table 7 that the values of IFIA during the postmonsoon is least, intermediate during the premonsoon and highest during the monsoon season at all the three frequencies. However, it should be noted that there is a difference in the amount of diffuse sky radiation for Calcutta compared with the values given in Table 4. These differences are due to the fact that the data we have considered here are only for locally clear days. This is required for a suitable comparison with the atmospheric data being considered in the analysis.

#### 4.3 Distribution of sunshine hours, normalized $D$ and IFIA

The distribution of the mean daily hours of bright sunshine for each month over the Indian subcontinent is shown in Fig. 6 using data from 40 stations. The duration of sunshine shows a maximum over central and northwest India, with an average value between 8 and 10 h per day. During the cold season, it is highest over central India, while the northern regions receives less sunshine due to the passage of winter disturbances from the west and the decrease in the length of the day with latitude and in the south due to cloudiness associated with disturbances in the Bay of Bengal and Arabian sea. As the year advances, till the onset of the monsoon in June, the region of clear skies and maximum sunshine shifts to the north and west

Table 7. Seasonal values of diffuse radiation and IFIA

Season	Mean diffuse radiation ( $J \cdot m^{-2} \cdot h^{-1}$ )	Relative values (%)	Mean hourly median IFIA (dB above $1 \mu v/m$ ) at		
			10 KHz	20 KHz	30 KHz
Postmonsoon	$46.88 \times 10^4$	24.5	71.2	65.4	59.6
Premonsoon	$70.74 \times 10^4$	37.1	78.6	69.4	63.0
Monsoon	$73.25 \times 10^4$	38.4	84.6	74.6	66.3



**Fig. 6.** Distribution of sunshine hours during **a** premonsoon; **b** monsoon; and **c** postmonsoon seasons over India

Pakistan. In the arid and semi-arid zones of central and northwest India, the skies are clear for generally more than 8–10 h a day.

The mean actual hours of sunshine over Calcutta are plotted in Fig. 7. A normalized curve for the diffuse radiation obtained from the mean values is also plotted monthwise in the figure. For this purpose, the mean monthly value is first calculated for different months, then normalized values of diffuse sky radiation are obtained by dividing the individual values by the maximum value obtained. A similar normalized curve for IFIA is further superimposed by considering the median noise level obtained from the means at 10, 20, and 30 KHz. Figure 7 reveals that the trends for the two normalized curves are similar, with a little more fluctuation for the *D* values. It is also interesting to note that the curve for sunshine hours is almost a mirror image of the other two curves, with a minimum during the cloudy monsoon season and a maximum during the remaining clear months.

**4.4 Significance test**

Using the normalized values of diffuse solar radiation and IFIA the significance test (*t*) has been calculated by the

usual relation (Croxtton and Cowden, 1964),

$$t = \frac{\bar{X}_d}{\sigma \bar{X}_d} = \frac{\bar{X}_d}{\hat{\sigma}_d / (n)^{1/2}}$$

Here,

$$\bar{X}_d = \frac{\Sigma d}{n}$$

and

$$\hat{\sigma}_d = \{[\Sigma d^2 / (n - 1)] - (\Sigma d)^2 / n(n - 1)\}^{1/2}$$

where *d* is the difference between the two mean values of a pair and *n* is the number of samples for which  $\Sigma d$  has been calculated. In our analysis, the three main Indian seasons of the year, i.e. winter, premonsoon and monsoon, have been considered for 5, 3, and 4 months, respectively, giving values of *n* for three different sets of *t*. The seasonal values of the significance ratio (*t*) and also its corresponding level of significance (*p*) are presented in Table 8. The level of significance (*p*) has been obtained from the knowledge of the significance ratio (*t*) and degrees of freedom (*f* = *n* - 1).

It appears from Table 8 that the level of significance is somewhat small during the winter months, but there is a marked rise in its value during the premonsoon and

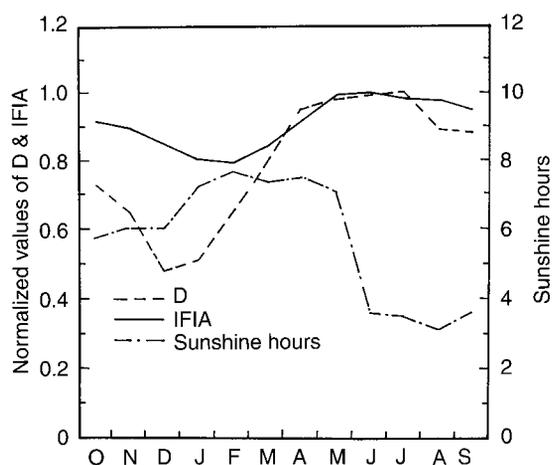


Fig. 7. Curves showing the sunshine hours, and normalized values of (1) diffuse radiation and (2) IFIA for Calcutta

Table 8. Significance ratio ( $t$ ) and level of significance ( $p$ )

Season	Degree of freedom ( $f = n - 1$ )	Significance ratio ( $t$ )	Level of significance ( $p$ )
Winter	4	6.5	$0.002 < p < 0.003$
Premonsoon	2	2.6	$0.15 < p < 0.20$
Monsoon	3	2.4	$0.15 < p < 0.20$

monsoon. The significance ratio during the latter two seasons is of the order of 2 only.

## 5 Discussion

The phenomena of scattered radiation in the atmosphere are very complex and varied. Numerous factors determining the quantity of scattered radiation change continuously in time and space and make computation difficult (Coppolino, 1992). The amount of diffuse radiation received for a particular geophysical location on the Earth's surface depends mainly on the solar elevation, turbidity in the atmosphere and cloudiness. On the other hand, the noise level in the IFIA at a place is solely controlled by lightning discharges around the observing station (Williams *et al.*, 1989). The seasonal variation of both the phenomena noted will be governed by the activity of their sources. In our results, close similarity between the variation of diffuse radiation with the variation of IFIA indicates that the common causes are more responsible for their variation. In the winter months when the local sky is almost free from clouds but dust particles are present diffuse radiation is least. Due to the poor activity of the sources, the noise level of IFIA also exhibits a minimum value during these months. With the advance of the year, there is a general increase of both diffuse radiation and IFIA due to an increase in cloudiness accompanied by nor'westers (local term) in the premonsoon months. With the establishment of the monsoon, both diffuse radiation and IFIA further increase. Both phenomena attain a

maximum value during the monsoon months as a result of the increased cloudiness over the region. With the withdrawal of the monsoon in October and onset of winter conditions, there is again a fall of  $D$  and IFIA.

The correlation coefficient between the values of diffuse solar radiation and the intensity of atmospheric is quite good. Besides the activity of clouds, some other parameters, e.g. pollution, also contribute to the diffuse radiation, either simultaneously or in succession, while the record of IFIA is influenced by a large number of natural sources over a very wide range. The positive sign associated with a high correlation coefficient between  $D$  and IFIA indicates that in producing those values the common causes are distinctly more effective than the contribution due to independent causes.

The lack of significance in the winter season is understandable, as the  $D$  values at such times are governed mainly by the characteristics of dust particles, gas molecules, etc. This is as a result of high local industrial pollution and partly by the activity of local clouds (Raja Rao, 1984). During the overcast cloudy months, on the other hand, there is a reduction in industrial pollution as a result of rainfall and wash-out in precipitation. The low level of significance during the winter season, thus, might be due to the combination of weak clouds and dust particles, while their high values during the premonsoon and monsoon months are perhaps caused by the vigorous activity of clouds at such times (Bhattacharya *et al.*, 1994).

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