



Asian dust outflow in the PBL and free atmosphere retrieved by NASA CALIPSO and an assimilated dust transport model

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**Three-dimensional
transport structure of
Asian dust**

Y. Hara et al.

Asian dust outflow in the PBL and free atmosphere retrieved by NASA CALIPSO and an assimilated dust transport model

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Abstract

Three-dimensional structures of Asian dust transport in the planetary boundary layer (PBL) and free atmosphere occurring successively during the end of May 2007 were clarified using results of space-borne backscatter lidar, Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP), and results simulated using a data-assimilated version of a dust transport model (RC4) based on a ground-based NIES lidar network. Assimilated results mitigated overestimation of dust concentration by reducing 17.4% of dust emissions and improved the root mean square difference (RMSD) of dust AOT between the model and NIES lidar by 31.2–66.9%. The dust layer depths, vertical and horizontal structure simulated by RC4 agreed with CALIOP from the dust source region to a long-range downwind region for which 3-D distribution of dust clouds had not been clarified previously. Based on CALIOP and RC4, two significant transport mechanisms of Asian dust in the PBL and free atmosphere were clarified: a low level dust outbreak within the dry slot region of a well developed low-pressure system, and formation of an elevated dust layer within the warm sector of a low-pressure system. Finally, the aging of pure dust particles was investigated using the particle depolarization ratio (PDR) at 532 nm and the color ratio (CR) at 1064 nm and 532 nm for the low-altitude dust transport case. Aerosols with high PDR were observed uniformly over the dust source region. As the dust cloud was transported to the eastern downwind regions, aerosols with low PDR and high CR were found in the layer of less than 1 km height, suggesting changes from the external to internal mixing state of spherical aerosols and dust in the surface layer.

1 Introduction

Asian dust has important influences not only on air quality in the planetary boundary layer (PBL) but also on the regional/global radiative balance. Asian dust emitted from desert regions in northern China and Mongolia is usually transported within the PBL,

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but sometimes it is elevated up to the free atmosphere (FA), occasionally reaching thousands of kilometers to North America (Takemura et al., 2002; Uno et al., 2001). Numerical modeling is an effective method to clarify the detailed 3-D structure of wide-range dust transport. However, dust modeling engenders several uncertainties, especially in dust emission processes. Therefore, detailed 3-D dust observation is important for improving dust model performance (e.g., dust concentration level over source region, horizontal scale, and vertical thickness of the dust layer).

Ground-based lidar networks are powerful measurement systems that provide important vertical structures of aerosol and cloud layers as well as aerosol optical properties. They can provide the vertical profile of spherical (pollution), non-spherical (dust) aerosol particles, and particle size information from two-wavelength lidar signals (e.g., Sugimoto et al., 2006), which is very important for understanding of dust phenomena.

Space-based top-down active measurement is another powerful system for dust distribution. Actually, a NASA space-borne backscatter lidar, Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) onboard CALIPSO, has provided high-resolution vertical profiles of clouds and aerosols since June 2006 (Winker et al., 2007). It can detect dust particles quantitatively from the source to downwind regions, but it presents some difficulties in providing aerosol information under optically thick aerosol and cloudy conditions. For more detailed understanding of dust emission and transport, integration between the dust transport model and CALIOP observations is necessary (e.g., Uno et al., 2008).

Large-scale dust storms over eastern Asia occurred continually during 21–31 May 2007. In this paper, we will present detailed integrated analysis of 3-D dust structure based on the NIES Lidar network, CALIOP observations, and results obtained using a four-dimensional variational (4DVAR) data assimilation version of a dust transport model (Yumimoto et al., 2007a, b). We will also show the dust transport within the PBL and the dust lift-up within the warm sector of low pressure system, and underscore the importance of warm sector lift up for long-range transport of Asian dust.

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2 Dust emission/transport model and observation data

We used the RAMS/CFORS-4DVAR dust transport model (RC4; Yumimoto et al., 2007a,b). For this study, we applied RC4 by assimilating the dust extinction coefficient derived from the NIES Lidar network. The RC4 domain is centered at 37.5° N, 115° E on a rotated polar stereographic system (see Fig. 1). The horizontal grids comprise 180×100 grids with resolution of 40 km. The vertical grids comprise 40 grid points extending from the surface to 23 km with 40 stretching grid layers (140 m at the surface to 650 m at the top) in terrain-following coordinates. Meteorological boundary conditions to RAMS are taken from NCEP/NCAR reanalysis data with 2.5°×2.5° resolution of 6-h intervals. The simulation was performed from 21–31 May 2007 with zero initial dust concentration.

Table 1 shows the locations of NIES lidar sites. We selected nine lidar sites (Beijing, Seoul, Hedo, Nagasaki, Matsue, Toyama, Niigata, Tsukuba, and Sendai; see Fig. 1) from the NIES lidar network (Sugimoto et al., 2006), which measures vertical profiles of dust with high spatial and temporal resolution. The extinction coefficient was derived based on the backward Fernald's method (Fernald, 1984) by setting a boundary condition at 6 km and $S_1=50$ sr (Liu et al., 2002). The contribution of mineral dust to the total extinction coefficient is determined using the particle depolarization ratio (Shimizu et al., 2004). Total dust emission during the simulation period is also presented in Table 1; 17% of dust emission reduction was achieved using the RC4 model to give optimal agreement with the NIES lidar measurement. Details of the data assimilation of the RC4 dust model were already described in Yumimoto et al. (2007a, b)

The CALIOP measurement is used for comparison in this paper. Level 1B (version 1.20) CALIOP data (horizontal resolution of 333 m) provide profiles of the total attenuated backscatter coefficient at 532/1064 nm and the volume depolarization ratio (δ) at 532 nm. The CALIOP Level 2 (horizontal resolution of 5 km) processing extracts features (cloud, aerosol, surface, etc.) included in a lidar profile. Version 1.20 of Level 2 products are layer-averaged attenuated optical properties along with cloud-aerosol

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discrimination (CAD) results (Liu et al., 2004). The CAD is an indicator that enables discrimination between cloud (positive value) and aerosol (negative value) data. The dust extinction coefficients were derived using Fernald's inversion (Fernald, 1984) by setting the lidar ratio $S_1=30$ sr, as described by Shimizu et al. (2004). Inversion started from $z=14$ km down to the ground surface. In this study, Level 1B data are averaged by 20 profiles and interpolated to Level 2's horizontal resolution (5 km). Detailed descriptions can be found on the CALIPSO mission webpage (<http://www-calipso.larc.nasa.gov/>) and references therein.

A smaller value of $S_1=30$ is used in the CALIOP extinction retrieval to obtain the overall semi-quantitative aerosol profile. As a result, the extinction coefficient retrieved from the CALIOP measurement might be smaller than that from the NIES lidar network measurements by a factor of ca. 0.6 ($=30/50$) for optically thinner cases. Fixed S_1 might engender excessive decay within the cloud layer (S_1 is around 20) and then give over-estimation of aerosol under that cloud. However, this is not critical because we are mainly targeting semi-quantitative analysis of dust transport (i.e., a comparison of dust vertical structure and horizontal scale) using the CALIOP measurement.

3 Results and discussion

3.1 Daily variation of meteorological conditions, dust AOT and CALIPSO observation paths

Continuous dust storms occurring in the Gobi Desert and Mongolia swept across vast areas of eastern Asia during 21–31 May 2007. The large dust event was also observed in Japan as latest spring record within the Japan Meteorological Agency observation history.

Figure 2 shows the day-by-day dust transport combined with meso-scale low-pressure systems and cloud image between 21 May and 28 May 2007. It includes a MODIS cloud image at 00:00 UTC, the RC4 aerosol optical thickness (AOT; red con-

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tour) at 09:00 UTC, simulated surface wind vector (yellow), and the CALIPSO observation path (light blue line). The path-time information of CALIPSO measurement is also written at the bottom of the figure; L in the figure shows the location of the low-pressure system. Symbols of L_1 –1–10, L_2 –1–5, upper troposphere dust layer UTD-1–4 (>6 km) in the figure denote locations of the vertical profile analysis. The selected CALIPSO path and location do not always coincide with the center of the simulated dust area because CALIPSO observations are limited to catching the center of the dust storm; moreover, the cold front with cloud layer quite frequently accompanies the strong dust storm area, even if CALIPSO has passed over the center of the dust storm area.

Table 2 shows the dust layer depth, layer averaged statistics (potential temperature, RH %, attenuated backscatter coefficient and dust extinction coefficient) taken from CALIOP and the RC4 dust model for the region of L_1 , L_2 , and UTD.

The first low-pressure system L_1 was located in northern Mongolia on 21 May. In fact, L_1 deeply developed as it moved eastward; it followed very strong surface winds which triggered dust mobilization in the Gobi Desert. On 23 May, a band of clouds over the eastern Mongolia region is apparent in the MODIS cloud image (Fig. 2c). The maximum surface wind speed was higher than 12.5 m s^{-1} over the Gobi Desert on 06:00 UTC, as simulated by RC4. A dense dust cloud is located within the dry slot area, and moved to the northeast China on 24 May; then arrived in Korea on 25 May. Actually, L_1 continues developing as it moves beyond the Sea of Japan, with a central sea level pressure down to 986 hPa. From late 25 to 26 May, the dust cloud behind the cold front had come to reach a wide coverage of the Japanese archipelago. At the northern part of Mongolia, the subsequently developed low-pressure system L_2 appears on 25–26 May. Strong winds on the western side of the low-pressure system continued to mobilize dust; then dust clouds flowed out continuously to southeast Korea and southern Japan.

These two subsequent low-pressure systems (identified as L_1 and L_2 in Fig. 2) triggered the continuous outbreak of dust clouds over eastern Asia. The region between L_1 and L_2 has two phase regions: a cold sector and warm sector. The NIES ground-

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based lidar network detected a heavy dust layer mainly within the planetary boundary layer (PBL), but it is also found in elevated dust layer of the 3–5 km altitudes at downwind northern Japanese sites, as depicted in Fig. 3.

We applied the RC4 dust model assimilated by NIES lidar observation for these events to clarify meteorological mechanisms for generating a vertical structure of dust transport. In the following sub-section, the RC4 dust model is used for detailed analysis of the 3-D dust transport structure from dust source regions to long-range downwind regions both with the NIES lidar network and with NASA/CALIOP measurements.

3.2 Time-height analysis with NIES Lidar network

Figure 3 shows (a) a time-series of dust AOT ($z < 6$ km), (b) a time-height cross-section of observed dust extinction coefficient by NIES lidar, and (c) a time-height cross section of the potential temperature and dust extinction coefficient by RC4 model for four NIES lidar sites. In Fig. 3a, the blue line represents the result without assimilation; the red line shows that with assimilation. The dust AOT from CALIOP was also calculated as shown by red symbols in Fig. 3a when CALIPSO passed within 200 km of the NIES lidar site. The solid circle denotes CALIOP data taken under an almost cloud-free condition, whereas the open-cross circle denotes that under cloudy conditions, which might cause large error in inversion results.

The observed vertical structures of the dust layer were simulated quite well by RC4. The first mid-latitude low-pressure system L_1 mobilizes heavy dust storms from 21 to 23 May in the Gobi Desert. The time delay (transport) of the first dust peak from Beijing and Seoul to Japanese sites is apparently from 23 May to 26 May. At Beijing, a dense deep dust layer from the surface to 3–4 km is apparent on 24 May; this dust layer was so thick that the lidar signal was unable to penetrate into the free atmosphere. The dust cloud accompanied by the cold front approached Seoul on 25 May, then Nagasaki and Toyama on 25–26 May. The RC4 model reproduced both correct timing of the arrival and dust layer height correctly.

The subsequent L_2 passed Beijing on 26–27 May, showing an important meteorolog-

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ical condition. The vertical gradient of potential temperature takes an isothermal profile from the surface to 5 km altitude on 26–27 May; then the dust layer depth deepened both in observation and RC4 results. The MODIS cloud images (Fig. 2f) show that the warm sector region of L_2 locates in the vicinity of Beijing on 26 May; the transported dust was within this sector. In the warm sector of low-pressure system, the ascending air motion plays an important role in upward transport of dust from PBL to the free atmosphere. At Seoul, elevated dust layer is apparent on 27–28 May, reaching 4 km altitude; an elevated dust layer was also detected at Toyama on 28–29 May in the vertical range of 3 km to 5 km elevation. While in Nagasaki, dust continues to appear only within the PBL, which indicates that the structure of the second dust cloud is split into two parts, i.e., the northern part of the Japan islands are mainly affected by the warm sector dust liftup and formation of the elevated dust layer, although the southern part of the Japan islands are controlled by the cold front line. The dust layer remains mainly within the PBL. We will explain the observed and simulated warm sector dust liftup in Sect. 3.4.

The RC4 dust model assimilated with the NIES lidar data shows improved agreement with lidar data. The magnitude and daily variation of both lidar and RC4 AOT show good agreement at most sites. Table 1 shows the root-mean square differences (RMSDs) for results achieved with and without data assimilation. The RMSD of dust AOT between model and NIES lidar was improved by 31.2–66.9% after introduction of data assimilation. In particular, at Beijing and Seoul, assimilated results mitigate overestimation of the dust extinction coefficient during 26–27 May. However, in Nagasaki and Toyama, the first dust cloud remains underestimated because the lidar at Beijing was unable to retrieve the whole dust layer above 1.5 km because of the dust layer, which was too thick for optical observations; these missing observation data are important for inverse correction of this underestimation.

Apparently, the variation of CALIOP AOT in Toyama was explained well by ground-based lidar and RC4. However, we also found that the CALIOP AOT under clear-sky conditions is smaller than both NIES lidar and RC4. This underestimate might result

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from difference of S_1 values between NIES lidar and the CALIOP inversion process. Available observation timing between CALIOP and NIES lidar remains limited. For that reason, we must accumulate more observations to produce sufficient quantitative verification.

5 3.3 Cross-section analysis along the cold front movement with CALIOP and RC4 model

This section presents tracking analyses of the vertical cross-section of dust following the movement of the L_1 cold front.

Figure 4 shows the vertical cross-section analysis of the particle depolarization ratio (PDR) at 532 nm, dust extinction coefficient along CALIPSO paths, and averaged vertical profiles of the dust extinction coefficient for L_1-2 , L_1-4 , L_1-6 and L_1-8 across the first dust storm. The CALIOP data were screened with conditions of $0.02 < \text{PDR} < 1$, attenuated backscatter coefficient $> 8 \times 10^{-4} \text{ (km}^{-1}\text{)}$ to avoid observation noise and $\text{CAD} < 90$ to screen out dense clouds.

An aerosol layer with high PDR, thought to comprise mainly dust, is shown in Fig. 4a. Examination over the source region (L_1-2) clarified that the model reproduced the dust mixing height correctly during the dust outbreak. According to Fig. 4b and c, both modeled and observed cross sections have very similar structures both vertically and horizontally. Cloud bands created by the cold front exist along the region for which the modeled relative humidity is considerably high, and dust aerosols are visible in the PBL behind the cloud band. As the low-pressure system L_1 develops, the potential temperature drops gradually; it reaches a minimum value of 296.9 K on L_1-6 at 25 May, 17:00 UTC. Based on this cold front tracking analysis, we found that the first dust cloud behind a cold front maintains the PBL height of around 2–3 km from the continent to the Pacific Ocean within the latitude range of 30–44° N; it then covers the wide region of eastern Asia with high-level dust concentrations. Another interesting point is the change of PDR from the dust source region to the downwind region, i.e., PDR decreases gradually in direct relation to transport distances, especially in a part

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of surface layer. Detailed analyses will be described in Sect. 3.6.

3.4 Analysis of warm sector dust uplift

In this section, we specifically address the vertical structure of dust between L_1 and L_2 , and present cross-section analyses of warm sector uplift and cold sector transport within PBL.

Figure 5 is the same as Fig. 4, but for the vertical structure L_2-1 , L_2-2 , L_2-4 , and L_2-5 across the second low L_2 . Please note that figures for L_2-2 and L_2-4 are unclear because these are taken during the CALIPSO daytime path with a small signal-to-noise ratio. Therefore, to outline the dust layer, we added a red broken line from the RC4 extinction coefficient value of $0.07 \text{ (km}^{-1}\text{)}$. From a snapshot of the L_2-4 path on 04:00 UTC 27 May, a marked dust structure is visible, with an elevated dust layer at northern latitudes ($38-48^\circ \text{ N}$) between 2 km and 6 km altitude and a low-altitude dust layer in southern latitudes ($30-38^\circ \text{ N}$).

Figure 5c is important for presenting clear evidence of the warm sector dust uplift from PBL to the free atmosphere near the source region. A MODIS image (Fig. 2g) shows the L_2-4 path just across the warm sector of low-pressure system L_2 . This fact suggests that the driving force producing the elevated dust layer is a dynamic warm sector updraft associated with low-pressure system activity. Trajectory analysis of the elevated dust layer using the Hybrid Particle Transport Model (HYPART) (Walko et al., 2001) based on RAMS output shows that the updraft dominates from the surface to upper troposphere in the vicinity of Beijing on 26 May; then the dust cloud is elevated to about 4 km height with a typical ascending speed of 2 km/12 h. This trajectory analysis will be described in Sect. 3.5.

The vertical structures retrieved by CALIOP shown in Fig. 5 are consistent with those by NIES lidar measurements in Sect. 3.2. As portrayed in Figs. 2 and 5, the dust layer is located under a thick cloud. However, RC4 results reveal that the boundary dust layer of thickness of 2 km passed over the East China Sea for the latitude range of $30-36^\circ \text{ N}$, as shown in Fig. 5c and d.

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3.5 3-D Structure of dust cloud transport by low-pressure system

We conduct Lagrangian forward trajectory analysis using HYPACT to investigate different transport structures of dust accompanying L_1 and L_2 . Table 3 shows HYPACT trajectory information.

5 The 3-D transport structure of the first dust cloud, which occurred on 23 May from the Gobi Desert, is presented in Fig. 6 with T-1 forward trajectory analysis. Figure 6a–d show dust extinction coefficients along CALIPSO paths for L_1 –2, L_1 –4, L_1 –6, and L_1 –8. The vertical cross section of the dust extinction coefficient and potential temperature from RC4 along the T-1 trajectory are shown in Fig. 6e. Figure 6f shows the dust AOT and surface wind speed interpolated temporally and spatially along the CALIPSO path time.

The T-1 trajectory corresponds with CALIOP observation in the time and trajectory location. The trajectory T-1 maintains altitude of 1 km, and travels nearly $1500 \text{ km} \cdot \text{d}^{-1}$. The 3-D transport structure in which dust is trapped within the potential temperature between 290 and 300 K is readily apparent; the dust travels within PBL for a long distance. Figure 6f shows that the dust AOT outbreak follows the westerly wind: a clear change in wind direction is visible at the center of the low-pressure system. Figure 6 presents the very typical 3-D structure of dust transport accompanied by the cold front and low-pressure system.

20 Figure 7 is the same as Fig. 6, except for the second dust cloud associated with low-pressure system L_2 . Therein, T-2 shows the trajectory analysis of a dust cloud taken in the cold sector; T-3 shows analysis of the elevated dust cloud. Figure 7e is for the trajectory T-3. The masked area in Fig. 7e indicates an updraft region simulated by the RC4 dust model.

25 The T-2 trajectory time of passage coincides with the CALIOP cross-section measurement, although the T-3 trajectory does not completely match up with the CALIOP cross section because of slower transportation of T-3 trajectory. Therefore, the elevated dust layer is not clearly visible in Fig. 7c and d. Nevertheless, this figure clearly

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shows two significant transport structures, associated with cold and warm air advection around the low-pressure system L_2 . The T-2 trajectory was departing from a northern point, and traveled southward at a horizontal speed of $1500 \text{ km}\cdot\text{d}^{-1}$, maintaining 2 km altitude, whereas the T-3 trajectory was departing within a warm sector, and ascended from ground to 4 km at about 116° E . Subsequently, the T-3 trajectory traveled eastward slowly with speed of $900 \text{ km}\cdot\text{d}^{-1}$; it reached Toyama on 00:00 UTC 29 May. As depicted in Fig. 7e, the dynamical updraft in the warm sector of the low-pressure system is an important meteorological mechanism producing an elevated dust layer that plays an important role in long-range transport. The formation of the elevated dust layer tends to occur under a meteorological situation of sequential outbreaks of two low-pressure systems; a very typical example is apparent in the so-called ACE-Asia “Perfect Dust Storm” (Uno et al., 2004). This mechanism is important for supplying nutrition through long-range transport to the distant open ocean.

Transport within the cold sector is also an important process because human society is considerably affected by dust-polluted PBL and dust cloud passage through the industrial areas of eastern Asia. Therefore, complicated processes of internal or external mixing occur between dust and anthropogenic aerosols, in which the mixing status plays an important role in human health and radiative balance. In the next section, we examine the changing of optical properties of dust cloud during outflow using wavelength ratio and aerosol depolarization ratio retrieved by CALIOP.

An L_1 event is characterized as a severe dust storm with the low average potential temperature of 301.7 K, whereas an L_2 dust event is understood to be a more moderate dust event than L_1 with the average potential temperature of 305.5 K at each altitude of the dust layer. The maximum dust extinction coefficient derived from CALIOP is 0.349 km^{-1} on the L_1 –2 path. The dust layer depth (see Table 2) behind the cold front of L_1 is 1320–3120 m (1847–4149 m) from CALIOP (from RC4). For the L_2 event, the dust depth is 2040–4200 m (2721–4666 m) for observation (RC4). Actually, RC4 tends to overestimate the dust layer depth because of coarser vertical resolution than CALIOP measurement. However, we can obtain a good correlation coefficient of 0.743

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between the simulated and observed layer depth. The upper troposphere dust (UTD as presented in Table 2) layer also can be found on 22–24 May; the dust layer heights are 5520–11 880 km (5350–11 201 km) based on CALIOP (RC4), and peaks of the dust layer are higher than 6 km. The origin of this UTD can be considered as regions other than the Gobi Desert because of the considerably high altitude of the dust layer (e.g., Taklimakan Deser region; Liu et al., 2008).

3.6 Correlation of color ratio and particle depolarization ratio

The CALIOP device provides aerosol information for two wavelengths of 532 and 1064 nm. The wavelength ratio of aerosol backscattering coefficients (defined as the color ratio; CR) indicates the aerosol size information. The relationship between CR and the particle depolarization ratio (PRD) plays an important role in characterization of aerosols (Sugimoto et al., 2002). Therefore, the scatter diagram derived from CALIOP is useful to clarify aging of pure dust particles during long-range dust transport. The region of analysis was selected based on trajectory analysis T-1 to track the same dust mass. We applied the screening method described in Sect. 3.3 to exclude the cloud layer.

Figure 8 shows a scatter diagram of CR and PDR from L_1 –2, 4, 6, and 8 data for three altitude levels. CALIOP data were averaged within 39.3 km horizontally and 1 km vertically. The solid triangles show arbitrary values for (A) the pure sulfate, (B) pure dust, (C) sea salt accumulation mode and (D) sea salt coarse mode, respectively. The solid line shows the theoretical curve for an external mixture of A and B.

A clear difference of aerosol characteristic is apparent among paths, particularly in Fig. 8a: in the dust source region L_1 –2, aerosol particles uniformly have high PDR greater than 0.3, and CR of 0.7–0.86. The L_1 –2 data are considered to show the information of pure dust particles which have just been emitted from the Gobi Desert region. As the dust cloud is transported down to the east, PDR decreases gradually; however, the CR does not change much. The L_1 –4 data points across northeast China fall on a theoretical curve of the external mixture of sulfate and dust, and L_1 –6 data

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points across Japan fall under the theoretical curve. Finally, regarding L_1 –8 data over the Pacific Ocean, some dots are distributed in high CR and low PDR. In addition to the internal mixing of dust and sulfate, there is a possibility that dust particles were mixed with sea salt since data was obtained for PBL over ocean.

5 Several processes change and age dust and air pollution particles. Size-dependent gravitational settling and dry/wet deposition of dust particles and internal or external mixing of dust and air pollutant particles are main processes of decreasing PDR. In eastern Asia, emission of anthropogenic pollutants has increased with recent rapid Asian economic growth. Therefore, it is usually difficult to separate effects of pollutants
10 from those of dust deposition processes for decreasing PDR. Application of a regional chemical transport model, CMAQ, shows that high concentrations of anthropogenic pollutants such as O_3 and sulfate aerosols were also transported within the dry slot behind the cold front line of L_1 . The CMAQ results support the strong possibility of mixing of dust and sulfate during the low-altitude dust transport. Additionally, RC4 predicted
15 high concentration field of sea salt in boundary layer over L_1 –8, therefore aerosols in the layer were suggested to be mixture of dust, sulfate and sea salt. Figure 8a can be understood as a typical aging process of pure dust with air pollutants and sea salt. In the upper altitude cases (b) and (c), such a clear aging step is not apparent, which suggests that the mixing of air pollutants occurred mainly at low altitude. As presented
20 in Fig. 8a, both high and low PDR values distributed below 1 km in latitudes of 32° – 38° N, which suggests that, in low-altitude dust transport cases, changing (or aging) from external to internal mixing phases of dust and pollutants occurs frequently. This scatter diagram using data obtained by tracking the same dust cloud enables analyses of the aging of pure dust particles.

25 4 Conclusions

Detailed analyses by integrating the NASA space-borne lidar CALIOP, RC4 dust transport model, and the ground-based NIES lidar network clarified 3-D structures of Asian

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dust transport triggered by two subsequent low-pressure systems L_1 and L_2 during 21–28 May 2007. The NIES lidar observation data were used for dust model assimilation. Assimilated results mitigated overestimation of dust concentration by reducing 17.4% of dust emissions and improved the RMSD of dust AOT between the model and NIES lidar data by 31.2–66.9%. Assimilated dust profiles by RC4 were verified using the CALIOP dust cross-section from the dust source region to the long-range downwind region down to the Pacific Ocean. From comparison, we found that the vertical and horizontal structure of dust cloud coincided well between RC4 results and CALIOP measurements.

We clarified two important transport mechanisms of Asian dust in the PBL and the free atmosphere; the first is a low-altitude dust outbreak within the dry slot region of the well-developed low-pressure system; the other is the formation of an elevated dust layer within the warm sector of the low-pressure system. The warm sector liftup dust layer was presented clearly by CALIOP and RC4. Results showed that the dust cloud located in the cold (warm) sector is transported within the boundary layer (up to the free atmosphere) through the low-pressure system that has a descending (ascending) motion.

The aging of pure dust particles was investigated using the particle depolarization ratio (PDR) at 532 nm and the color ratio (CR) at 1064 nm and 532 nm for the low-altitude dust transport case by tracking nearly the same dust air mass. A clear difference of aerosol characteristic among paths was found the layer below 1 km altitude: aerosols with uniformly high PDR were observed over the dust source region, whereas the mixing states between air pollution aerosol and dust were found in the downwind region over Japan and the Pacific Ocean. This difference suggests a change from the external to internal mixing state of spherical aerosols and dust in the surface layer.

Results presented in this paper showed that the integrated analysis of 3-D dust observation and numerical dust transport model plays an important role in elucidating and characterizing dust phenomena. For this study, we used different S_1 values for CALIOP and NIES lidar inversion. It is important to revisit and reanalyze the same dust event

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using the official CALIOP extinction coefficient based on the variable S_1 setting from NASA, which is scheduled for release in 2008.

Acknowledgements. This work was partly supported by the Global Environment Research Fund of the Ministry of the Environment, Japan (C-061), and a Grant-in-Aid for Scientific Research under grant 17360259 from the Ministry of Education, Culture, Sports, Science and Technology, Japan. Lidar observations at Beijing were conducted with the cooperation of Z. Wang of the Institute of Atmospheric Physics, China and NIES. The CALIPSO data were obtained from the NASA Langley Research Center Atmospheric Sciences Data Center. The MODIS cloud image used in Fig. 2 was obtained from the NRL/Monterey Aerosol Page (<http://www.nrlmry.navy.mil/aerosol>).

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**Table 1.** Locations of NIES lidar Sites, RMSD, and Dust emission statistics.

Station	Longitude °E	Latitude °N	CTM/RMSD AOT	RC4/RMSD AOT	Reduced ratio ^a %
Beijing	116.46	39.92	0.487	0.223	−54.6
Seoul	126.90	37.50	0.329	0.109	−66.9
Hedo/Okinawa	128.25	26.87	0.070	0.076	+8.6
Nagasaki	129.86	32.78	0.166	0.171	+3.0
Matsue	133.01	35.21	0.252	0.173	−31.3
Toyama	137.10	36.70	0.217	0.134	−38.2
Niigata	138.94	37.84	0.189	0.130	−31.2
Tsukuba	140.12	36.05	0.189	0.111	−41.3
Sendai	140.84	38.25	0.205	0.134	−34.6
			CTM/ Emission (Tg)	RC4/ Emission (Tg)	Difference ^b Emission (Tg)
Dust Emission	–	–	74.9	61.9	−17.4

^a Improved ratio by data assimilation for AOT^b Difference in dust emission.

Table 2. Characteristics of dust layers associated with low-pressure systems L_1 and L_2 , and the downwind upper troposphere dust (UTD) layer.

Label	Date May	Path time hhmmss UTC	Lat. Range (° N)	Surface Elev. (m)	Z_0 (Layer depth)* (m)	Θ (K)	RH (%)	Atn. Bsc. $\text{km}^{-1} \text{sr}^{-1}$	Dust Ext. Coef. * km^{-1}
L_{1-1}	21	190843	40–42	1021	3120 4149	308.7	57.3	0.00240	0.057 0.108
L_{1-2}	23	185624	36–39	1331	1320 2208	300.7	29.8	0.00815	0.349 0.255
L_{1-3}	24	053616	35–37	915	2040 2458	308.0	14.8	0.00422	0.150 0.163
L_{1-4}	24	180052	36–37	0	2640 2193	300.4	26.0	0.00387	0.181 0.377
L_{1-5}	25	044044	33–35	0	2040 1847	301.0	20.1	0.00403	0.137 0.295
L_{1-6}	25	170515	34–36	0	2400 1994	296.9	48.6	0.00330	0.086 0.165
L_{1-7}	25	184410	33–37	0	2400 1994	298.7	49.9	0.00451	0.130 0.162
L_{1-8}	26	160943	34–35	0	2040 1847	299.2	38.5	0.00213	0.082 0.109
L_{1-9}	27	165301	31–33	0	2520 3961	303.4	35.0	0.00279	0.079 0.091
L_{1-10}	28	155724	30–32	0	1320 3292	300.7	47.9	0.01012	0.160 0.072
L_{2-1}	25	184410	40–41.5	1141	3240 3533	309.5	21.5	0.00156	0.020 0.081
L_{2-2n}	26	052402	42–44	965	3000 3769	306.4	50.3	0.00343	0.110 0.129
L_{2-2s}	26	052402	30–31.5	0	2040 2721	310.1	23.5	0.00206	0.022 0.171
L_{2-3}	26	174833	38.5–40	0	3960 4519	303.3	33.1	0.00158	0.013 0.085
L_{2-4}	27	042825	37.5–39	0	4200 4519	304.9	31.5	0.00189	0.019 0.107
L_{2-5}	28	173620	34.5–35	0	3360 4666	302.2	56.8	0.00200	0.052 0.046
UTD1	22	181306	39–40	0	8260–10800 8601–9901	334.5	17.0	0.00126	0.021 0.014
UTD2	23	171734	36–37	0	7080–9120 6001–8601	327.5	49.2	0.00116	0.015 0.018
UTD3	23	031407	37.5–38	0	6720–11880 6001–11201	324.0	19.1	0.00080	0.004 0.003
UTD4	24	162157	34–35	0	5520–7800 5350–8601	325.6	27.3	0.00093	0.004 0.015

* upper=CALIPSO observation; lower=RC4 dust model. For UTD, bottom-top height.

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Table 3. Starting locations of forward trajectories.

Label	Date May	UTC	Starting point of forward trajectory	Altitude from sea level (m)	Surface Elevation (m)
T-1	23	19:00	42.0° N, 108.3° E	2022	1273
T-2	25	19:00	42.0° N, 111.4° E	4638	1430
T-3	25	19:00	40.0° N, 110.8° E	1460	1084

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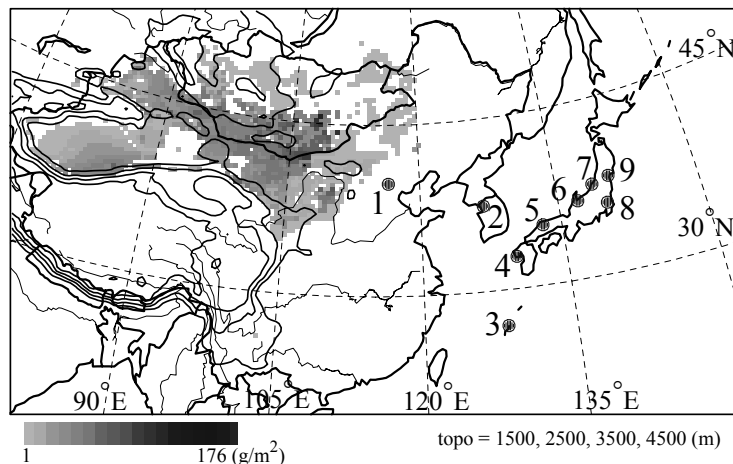


Fig. 1. Model domain and distribution of emission intensity (gray tone level). Dots represent the locations of NIES lidar network. Contour lines are the topographical elevations. The NIES lidar sites are: 1, Beijing; 2, Seoul; 3, Hedo; 4, Nagasaki; 5, Matsue; 6, Toyama; 7, Niigata; 8, Tsukuba; and 9, Sendai.

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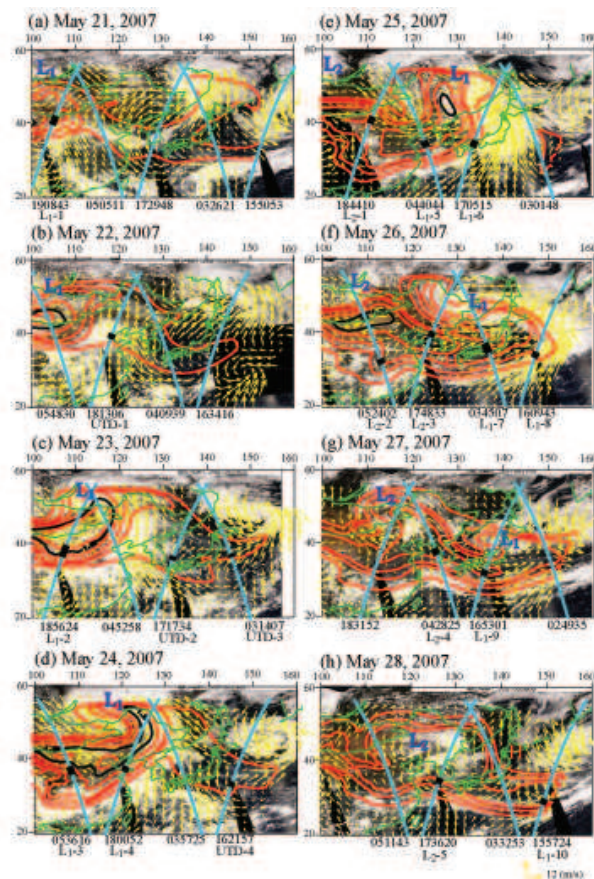


Fig. 2. MODIS cloud images, CALIPSO orbit paths (light blue line) and RC4 dust AOT (thick red contour). Vectors show the wind field at 140 m a.g.l. Six-digit numbers at the bottom of each figure are the CALIPSO path times (HHMMSS).

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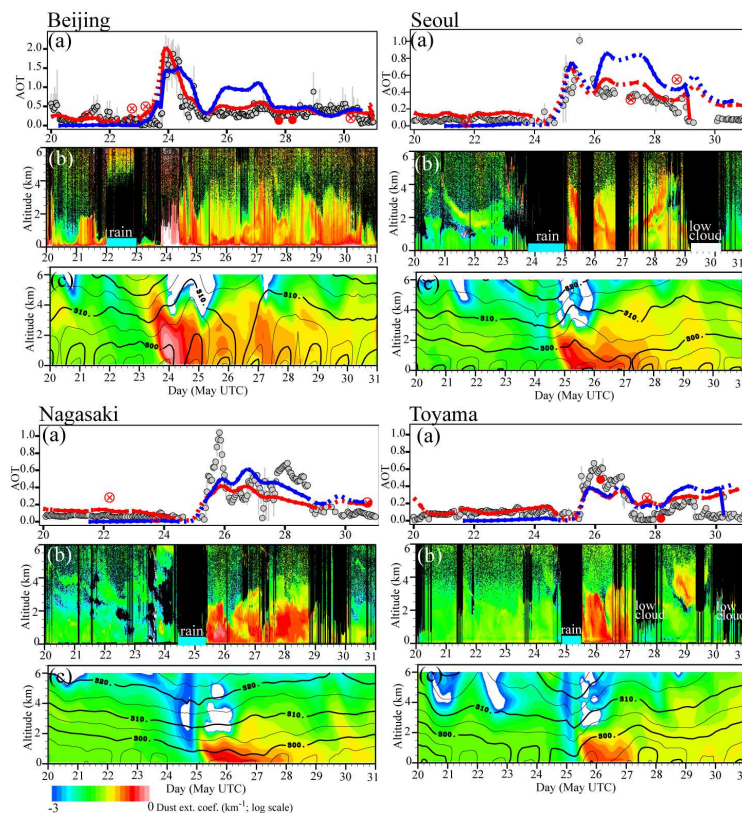


Fig. 3. (a) Comparisons of NIES lidar (gray circles) and RC4 dust AOT (lines). AOT is calculated to a height of less than 6 km for the Beijing, Seoul, Nagasaki, and Toyama locations. The blue line shows the result without assimilation; the red line shows that with assimilation. Symbols (solid and open-cross circle) are AOT derived from CALIOP, as described in the text. The time-height cross section of (b) dust extinction coefficient observed by lidar (km^{-1}) and (c) the simulated dust extinction coefficient (km^{-1}).

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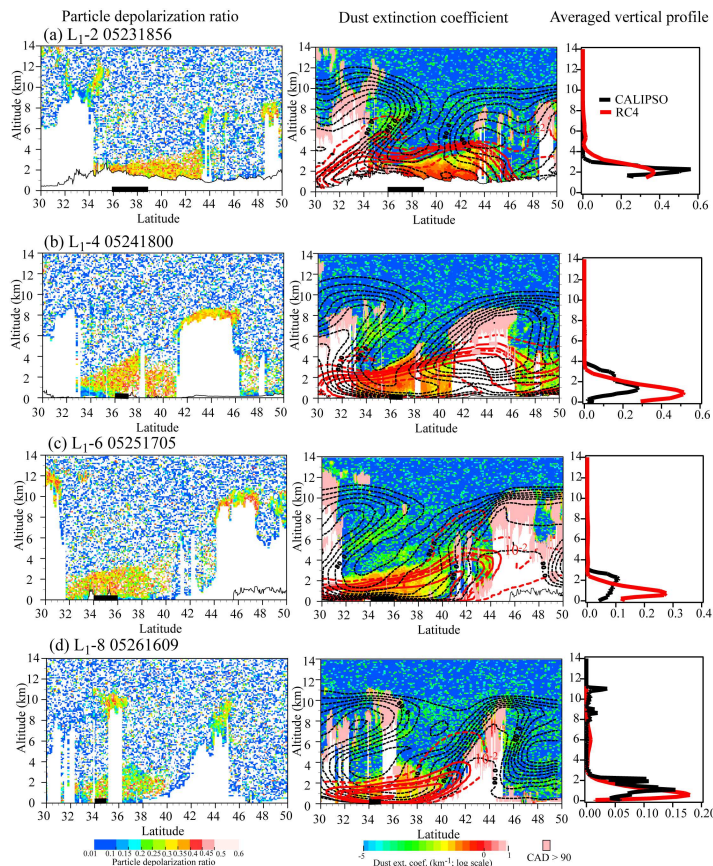


Fig. 4. (Left) Vertical cross-section of the CALIOP particle depolarization ratio, (Middle) CALIOP dust extinction coefficient (color), RC4 dust extinction coefficient (thick red contour), RAMS relative humidity (black dashed contour) and region of CAD > 90 (by light pink area) for paths L₁-2, 4, 6, and 8. (Right) Averaged vertical profile of dust extinction coefficient for CALIOP and RC4 for L₁-2, 4, 6, and 8.

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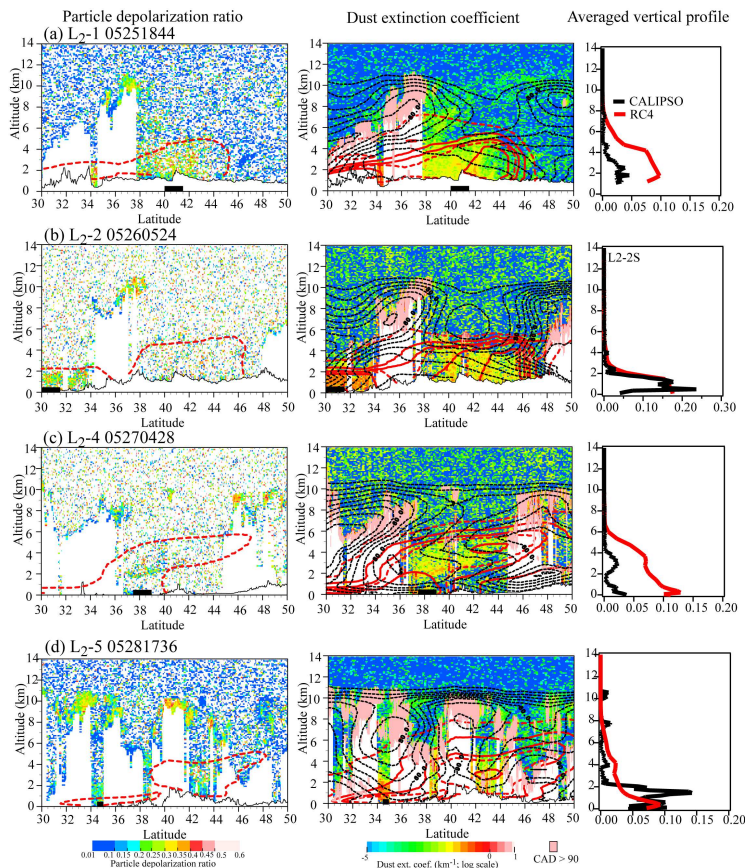


Fig. 5. Same as Fig. 4, but for $L_2-1, 2, 4$, and 5.

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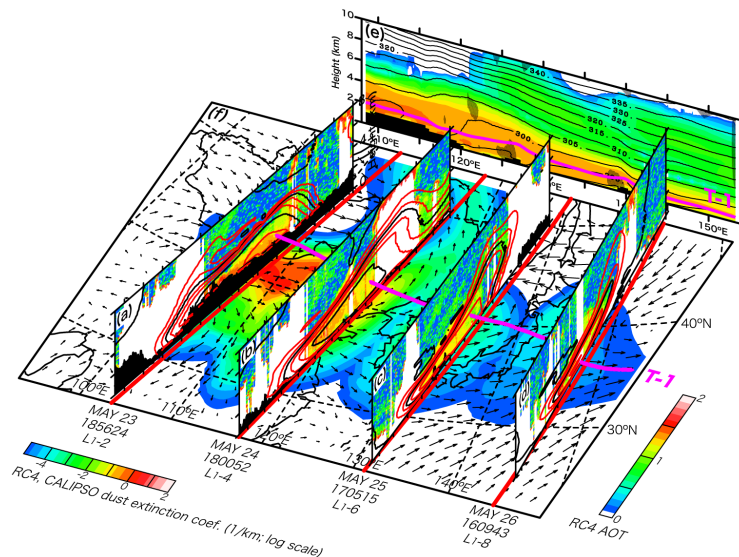


Fig. 6. 3-D analysis of CALIOP and RC4 cross-section. **(a)–(d)** Vertical cross-sections of CALIOP path (L_1 –2, 4, 6, and 8) by CALIOP dust extinction (color) and RC4 dust extinction coefficient (thick red line). Red lines (T-1) show forward trajectories starting from north of L_1 –2 (start 23 May, 19:00 UTC at 2022 m MSL). **(e)** The longitudinal cross-section along the T-1 from the RC4 dust extinction coefficient (color) and potential temperature (contour).

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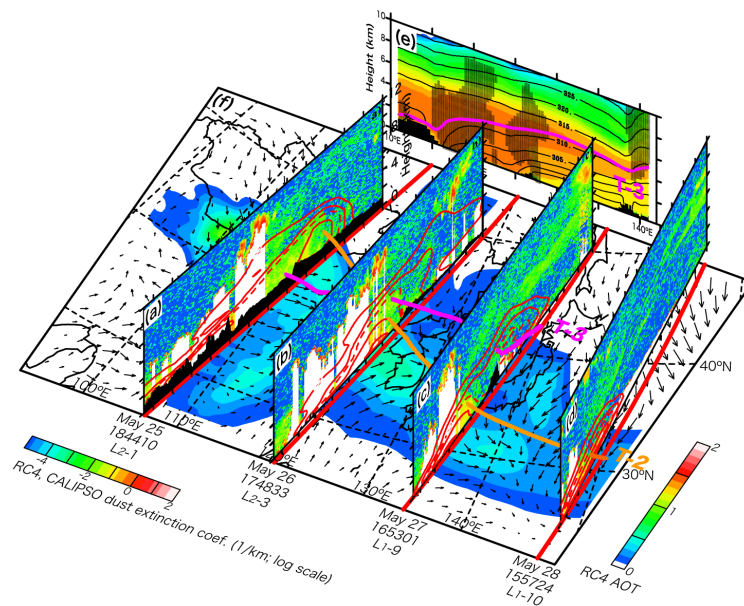


Fig. 7. 3-D analysis of CALIOP and RC4 cross-section. **(a)–(d)** Vertical cross-sections of CALIOP path (L_2-1 , 3, L_1-9 , and L_1-10) by CALIOP dust extinction (color) and RC4 dust extinction coefficient (thick red line). The orange line (T-2) and red line (T-3) respectively depict forward trajectories from the L_2-1 path starting at 25 May, 19:00 UTC at 4638 m and 1460 m MSL. **(e)** Longitudinal cross-section along the T-3 from RC4 dust extinction coefficient (color) and the potential temperature (contour).

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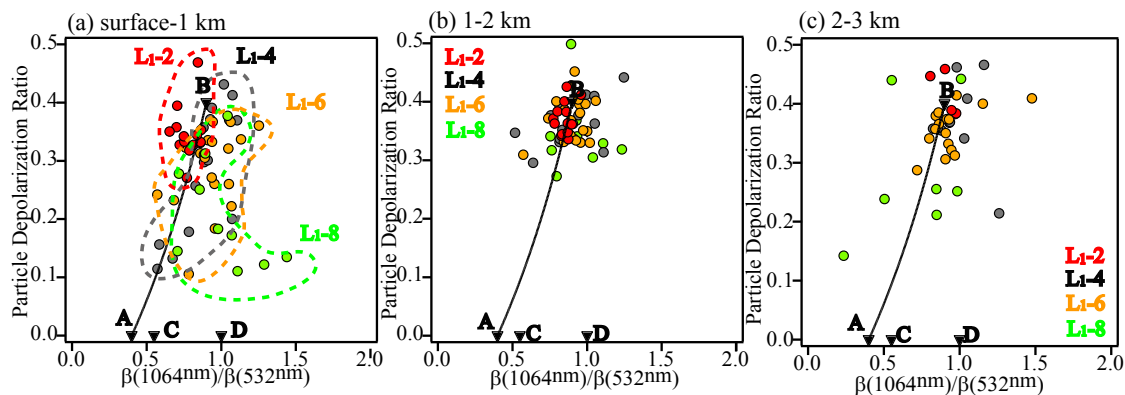


Fig. 8. Scatter diagram of the particle depolarization ratio (PRD) and the color ratio (CR) derived from L_1-2 , 4, 6, and 8 CALIOP data for three altitude levels. The solid line shows the theoretical curve for an external mixture of sulfate (A) and dust (B).

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