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**Droplet growth on
organic aerosols
from field
measurements**

N. C. Shantz et al.

The effect of organic compounds on the growth rate of cloud droplets in marine and forest settings

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Organic matter represents an important fraction of the fine particle aerosol, yet our knowledge of the roles of organics in the activation of aerosol particles into cloud droplets is poor. A cloud condensation nucleus (CCN) counter is used to examine the relative growth rates of cloud droplets for case studies from field measurements on the North Pacific Ocean and in a coniferous forest. A model of the condensational growth of water droplets, on particles dissolving according to their solubility in water, is used to simulate the initial scattering of the droplets as they grow in the CCN counter. Simulations of the growth rates of fine particles sampled in the marine boundary layer of the North Pacific Ocean indicate that the main influence of the marine organic material on the water uptake rate is from its effect on the size distribution of the sulphate. Simulations of the observations of water uptake on biogenic organic aerosol particles sampled in a coniferous forest indicate an impact of the organic on the water uptake rates, but one that is still smaller than that of pure sulphate. The solubility of the organic becomes an important factor in determining the water uptake as the organic mass increases relative to sulphate. The values of the organic component of the hygroscopicity parameter κ that describes the CCN activity were found to be negligible for the marine particles and 0.02–0.05 for the forest particles.

1 Introduction

One of the largest uncertainties in determining the human impact on the global climate is the effect of aerosol particles on clouds or the indirect effect (IPCC, 2007). Atmospheric aerosol particles have varying abilities to nucleate cloud droplets depending on the size and composition of the particles. Historically, it was believed that primarily highly soluble inorganic species, such as sulphate and sea salt, contributed to cloud droplet growth. More recently, it has become apparent that organics are a significant component of the atmospheric aerosol, but we lack sufficient knowledge about how

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organic compounds act as cloud condensation nuclei (CCN), and this contributes to the large uncertainty in the indirect effect (IPCC, 2007). There may be thousands of individual species comprising the organic aerosol fraction and characterization of the organic constituents in the atmosphere has only scratched the surface (e.g., Alves et al., 2001; Carvalho et al., 2003; Gao et al., 2003; Graham et al., 2003; Yu et al., 2005). Because the known organic compounds range from soluble to insoluble (Saxena and Hildemann, 1996), the question of how the particulate organic contributes to cloud droplet growth is more difficult.

The complexity of the organic aerosol has led some, when examining aerosol-CCN and aerosol cloud droplet closure, to assume that all of the aerosol mass was sulphate. In some cases, this assumption has led to reasonable agreement between the observed number of cloud droplets or CCN (e.g. Conant et al., 2004; VanReken et al., 2003). In other cases, the agreement was relatively poor and a lack of information about organics is one of the explanations (e.g. Kaku et al., 2006; Snider et al., 2003). Some found that the inorganic mass present during the field measurements accounted for the majority of the water uptake or CCN activity and the organic mass did not contribute substantially to the water uptake; thus the organic could be assumed to be insoluble (Broekhuizen et al., 2006; Cantrell et al., 2001; Fountoukis et al., 2007; Lowenthal et al., 2003; Medina et al., 2007; Roberts et al., 2002). Chang et al. (2007) found specific times of high organic mass during the field study when the insoluble organic assumption led to an underestimation in the number of CCN, and it was necessary to assume that water uptake by the organic was significant. Organics have been shown to contribute to water uptake at relative humidities less than 100% (e.g. Dick, 2000; Gysel et al., 2004), and Kreidenweis et al. (2006) suggest that treating the slightly soluble component of the aerosol as either fully soluble or fully insoluble may be the reason why closure has not always been achieved for polluted air in field measurements.

Overall, organic particles tend to be less soluble than the major inorganic compounds in the atmospheric aerosol, and this increases the diameter at which an organic particle will activate to a cloud droplet (critical diameter) (Shulman et al., 1996). Organic com-

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pounds may also inhibit the condensation of water (e.g. Cantrell et al., 2001; Feingold and Chuang, 2002). On the other hand, organics may decrease the surface tension of a droplet, tending to increase the critical diameter and increasing its ability to take up water (Anttila and Kerminen, 2002; Ervens et al., 2004; Facchini et al., 1999; Mircea et al., 2002; Shulman et al., 1996). This surface tension affect could potentially counteract the delayed activation imposed by reduced particle solubility (Shulman et al., 1996). However, the effect of surface active organics on water uptake is still an unknown (Abbatt et al., 2005; Bilde and Svenningsson, 2004).

Laboratory studies of the effect of organics on cloud droplet activation have shown that water-soluble organics contribute to CCN activity, insoluble organics do not and slightly soluble organics have a range of cloud-forming capabilities (Abbatt et al., 2005; Broekhuizen et al., 2004; Corrigan and Novakov, 1999; Cruz and Pandis, 1997; Giebl et al., 2002; Hartz et al., 2006; Henning et al., 2005; Hori et al., 2003; Kumar et al., 2003; Prenni et al., 2001; Raymond and Pandis, 2002; Shantz et al., 2003). The addition of even a few percent of inorganic material (by mass) to a mostly insoluble organic particle can have a large impact on the CCN activation (Abbatt et al., 2005; Bilde and Svenningsson, 2004; Broekhuizen et al., 2004; Lohmann et al., 2004; Raymond and Pandis, 2003).

Results from modelling studies of the effects of organics solubility on the formation of cloud droplets have varied. Abdul-Razzak and Ghan (2005) found that in mixtures of slightly soluble organics with inorganics, the organics can be considered fully soluble yet Ervens et al. (2004) found that including low solubility organics decreased the number of CCN relative to a fully soluble case. Mircea et al. (2005) showed that with a situation with a high organic fraction, knowledge of the chemical composition (and thus the solubility) of the water soluble component of the aerosol is needed.

The delay in the growth of cloud droplets formed on organic particles due to reduced solubility observed in a laboratory and modelling study can have a significant effect on the ability of a particle to activate in clouds in the atmosphere (Shantz et al., 2003). The work presented here extends the approach of Shantz et al. to field observations. The

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initial growth of aerosols as they activate into cloud droplets is examined using measurements from two field projects, one sampling from a ship on the North Pacific Ocean as part of the Canadian SOLAS study (C-SOLAS) and one sampling in Golden Ears Provincial Park (GE) on the west coast of Canada. These sites provide strong contrasts, a marine environment with the aerosol mass dominated by sulphate (Phinney et al., 2006) and a forest site with a small urban influence and dominated by organic aerosol (Shantz et al., 2004).

A model that treats changes in droplet growth with respect to time and the Mie scattering off of these droplets is used to simulate the CCN measurements and to study the differences of the two contrasting aerosol regimes on the nucleation of cloud droplets. As it can be difficult to characterize the water activity of organic aerosols due to their complexity and insufficient information about their chemical composition, a single hygroscopicity parameter that considers both hygroscopic and CCN activity (κ : Petters and Kreidenweis, 2007) was also determined. Using this parameter can simplify the physical and chemical information needed for describing activation in models.

2 Instrumentation

2.1 Particle size distributions

Size distributions of particles from 0.01–0.4 μm were measured with a Scanning Mobility Particle Sizer (SMPS) comprised of a TSI 3071 Electrostatic Classifier and a TSI 3025 Ultrafine Condensation Particle Counter. Larger particle size distributions were measured using an optical particle counter, a PMS Passive Cavity Aerosol Spectrometer Probe (PCASP). Both the SMPS and PCASP are used in the two field projects, the C-SOLAS study and the Golden Ears (GE) study (see Phinney et al., 2006; Shantz et al., 2004 respectively for further details).

2.2 Particle chemical constituents

During the C-SOLAS study, a Quadrupole Aerodyne Aerosol Mass Spectrometer (Q-AMS, see Jayne et al., 2000) provided the size distribution of sulphate, nitrate, ammonium, methanesulphonic acid (MSA) and total organics in the aerosol. In the Q-AMS, the particles are focused into a narrow beam and directed into an oven, which vapourises the particles. The vapourised components of the particles are ionized by electron impact. The resulting ion fragments are detected with a quadrupole mass spectrometer. A comprehensive table of fragmentation patterns is used to combine the fragments over 300 amu into mass concentrations of the above major compounds. The MSA fragmentation patterns were first implemented in this study and were determined in laboratory experiments (Phinney et al., 2006). In addition to the total sampled particle masses, the composition is also determined as a function of particle size. The particles are sized based on the time it takes for a particle to travel from the inlet to the detection point. This size is measured as a vacuum aerodynamic diameter (D_{va}). A comprehensive description of the design of the Q-AMS, its operation, quantification methods and calibration procedures are given elsewhere (Alfarra et al., 2004; Allan et al., 2003a; Allan et al., 2003b; Jayne et al., 2000; Jimenez et al., 2003).

Also during C-SOLAS, a Micro-Orifice, Uniform-Deposit Impactor (MOUDI) was used to collect particles for major ion analysis using ion chromatography (IC). This was done using Teflon filters on each of the 12 stages of the MOUDI. Despite its name, the MOUDI was operated in a static mode as opposed to rotating mode that gives rise to the uniform-deposit. The static operation was sufficient as the filters were completely extracted in de-ionized water prior to analysis by IC. The MOUDI data in this work was used to determine the contribution from sea salt. The agreement between the Q-AMS measurements and the MOUDI results for chemical species measured using both methods was relatively close (Phinney et al., 2006).

During the GE study (Shantz et al., 2004), particles were collected on single Teflon filters analyzed by IC for information on the total mass concentrations of inorganic

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ion constituents (not separated by size) including Na^+ , Cl^- , NH_4^+ , SO_4^{2-} , NO_3^- and K^+ . Particles for analysis of total mass concentrations of particulate organic carbon (OC), black carbon (BC) and water-soluble organic carbon (WSOC) were collected on single Quartz filters. The BC and OC concentrations were determined by a modified NIOSH method using a Sunset Labs Thermal Optical Transmission Instrument (Lee et al., 2003; Sharma et al., 2002). Water-soluble organic carbon (WSOC) was also measured in samples from the quartz filters as described by Shantz et al. (2004).

2.3 Water uptake: cloud condensation nucleus counter

A University of Wyoming Model MA100 static thermal gradient cloud condensation nucleus counter (CCNc) was used for the CCN measurements during both studies (Shantz et al., 2003; Snider et al., 2006). Light scattered off the growing droplets within the supersaturated environment of the chamber is detected as a voltage from a photodiode. This voltage is proportional to both the number concentrations and sizes of the growing droplets. For a given supersaturation, the size of the droplets is dependent on their composition, and thus it is difficult to distinguish the effects of number and size from the voltage measurement unless one is relatively constant. For a given particle composition, the difference between the maximum voltage (which is determined when the droplets are at their largest size prior to falling out of the detection region) and the baseline voltage (the signal prior to the exposure of the particles to a water supersaturation in the chamber) is proportional to the number concentration of the droplets. Hereafter, this is referred to as “deltaV”.

CCN observations are commonly expressed as the number concentration of particles that activated or CCN number concentration (e.g., Bilde and Svenningsson, 2004; Cruz and Pandis, 1997; Hegg et al., 2001; Leaitch et al., 1999; Snider et al., 2003; Stroud et al., 2007). One could determine the CCN number concentration from a calibration constant applied to deltaV (e.g. Leaitch et al., 1999). Due to the limitations discussed above, it is unreasonable to expect to determine accurate CCN number concentrations using this approach to CCN measurements. Thus, the focus of this

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work takes the approach of Shantz et al. (2003) and examines the variation in the light scattered by the particles and growing droplets. The rate of change in the scattered light, represented by the time-resolved voltage, is proportional to the growth rate into cloud droplets.

5 CCNc supersaturations should be experimentally determined and not assumed to be the theoretical values (Frank et al., 2007). The effective supersaturations (hereafter: S_{eff}) within the CCNc were estimated to be 0.19–0.50% using nearly monodisperse ammonium sulphate particles generated in the laboratory, as discussed in Shantz et al. (2003). All supersaturations discussed here are based on these calibrations. The
10 evaluation of S_{eff} results in the largest uncertainty in these measurements. The errors in S_{eff} are estimated as approximately $\pm 15\%$. It is difficult to evaluate how this error affects the growth rates from the observations, and so this uncertainty is applied in a model and reflected as the maximum and minimum limits of S_{eff} .

3 Model description

15 3.1 Simulations of cloud droplet growth rates in the CCNc

An adiabatic cloud parcel model (hereafter called the “ACP model”) described by Leitch et al. (1986) was modified to include partially soluble material (Lohmann et al., 2004; Shantz et al., 2003) and to simulate the growth rates in the CCNc (hereafter referred to as the “CCNc model”) (Shantz et al., 2003). The difference between the
20 ACP model and the CCNc model is how the supersaturation (S) is calculated. In the ACP model, the rate of cooling of the air parcel is balanced by the rate of water uptake to determine S . In the CCNc model, S is specified by a transient rise (that occurs in the CCNc after the flush of particles into the CCNc chamber) followed by a constant S , in an attempt to mimic the variation of S in the chamber. The transient S is de-
25 termined from comparisons of simulated growth rates with those measured for pure, nearly monodisperse ammonium sulphate (Shantz et al., 2003).

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The mass accommodation coefficient (α_c) can have a strong effect on the simulated condensation rate of water, and this is an outstanding uncertainty in quantifying the indirect effect of aerosols on climate forcing. There are a wide range of α_c values used in the literature spanning two orders of magnitude, from 0.01 to 1 (e.g. Davidovits et al., 2004; Laaksonen et al., 2005; Marek and Straub, 2001; Mozurkewich, 1986) and the references therein). For cases where an organic film may form on the surface of the cloud droplet, it has been suggested that the value of α_c could be as low as 10^{-5} (Chuang, 2003; Feingold and Chuang, 2002). Ruehl et al. (2008) determined α_c for field measurements (normalized to α_c for laboratory ammonium sulphate) and found that 60% of ambient CCN grew similarly to ammonium sulphate but a number of cases had a low α_c compared to ammonium sulphate. We did not have a measurement of the droplet size from this CCNc as was the case for Ruehl et al. (2008) so we were unable to determine α_c using their method. Also, α_c can not be determined from the growth rates here, as attempted by Stroud et al. (2007), because of the uncertainties in these CCNc's such as turbulence and edge effects. Consequently, α_c is assumed to be unity in all of the present simulations as suggested in many of the above references. What this means for this study is that although the effect of the organic on water uptake is represented primarily by the dissolution of the organic, any real impact of the organic could alternatively or in some combination be through effects on α_c or the surface tension.

For all model runs outlined here, the thermal accommodation coefficient (α_T) are assumed to be 1. Based on other results (e.g. Chuang, 2006), it is assumed that changing the value for α_T will not affect the growth rate significantly. All dry particles are assumed to be spherical. It is also assumed that the compounds present in the particles and droplets do not chemically react with each other.

The output of the model includes information about the droplet size, number concentration and the scattering efficiency of the growing droplets. For direct comparison of the model output to the first five seconds of light scattering off of the droplets from the CCNc measurements (a detector voltage), the scattering cross-sectional area of the

growing droplets is calculated, using Mie theory.

3.2 Simulations of organics

All the organics in the model are assumed to resemble adipic acid in terms of the physical and chemical properties, except for solubility. The latter is varied. This assumption is used because the CCN and water uptake properties of adipic acid (Table 1) have been widely studied (e.g., Bilde and Svenningsson, 2004; Broekhuizen et al., 2004; Corrigan and Novakov, 1999; Cruz and Pandis, 1997; Ervens et al., 2004; Hori et al., 2003; Joutsensaari et al., 2001; Prenni et al., 2001; Raymond and Pandis, 2002; Shulman et al., 1996) and adipic acid has been observed in the atmosphere (Grosjean et al., 1978; Kawamura and Yasui, 2005; Wang et al., 2006). The determination of the hygroscopicity parameter κ (Petters and Kreidenweis, 2007) (Sect. 5) is also a means of evaluating this assumption.

Sensitivity to solubility is examined because the previous laboratory work and simulations suggest that solubility is an important factor when dealing with organics (Shantz et al., 2003) and it is better constrained than other properties such as the surface tension and the accommodation coefficient. Thus, the dissolution of the organic in the model depends on its solubility and is based on its equilibrium with the amount of water present on the particles and droplets at each time step, as described in Shantz et al. (2003). The dissolution kinetics discussed in Asa-Awuku and Nenes (2007), which is more important at low temperatures, was not accounted for in this work. The surface tension is defined according to Ervens et al. (2004) and is shown in Table 1. This does not take into account the possibility of a lower solubility film-forming organic compound (Feingold and Chuang, 2002), the effects of which are uncertain (Kanakidou et al., 2005).

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4 Observations and simulations of cloud droplet growth rates in the CCNc

4.1 Pacific Ocean- C-SOLAS field study

The Sub-arctic Ecosystem Response to Iron Enrichment Study (SERIES), an experiment of the Canadian Surface Ocean and Lower Atmosphere Study (C-SOLAS), took place in the Pacific Ocean in July 2002. The research platform for the atmospheric aerosol measurements was the Mexican oceanographic vessel El Puma. The majority of the measurements were made near Station Papa, an ocean station located in the sub-arctic northeast Pacific, nearly 1400 km off shore. All times used in the work for C-SOLAS are in Pacific Daylight Savings Time (i.e. not local time in the ocean, which was 1.5 h behind PDT). Table 2 outlines the measurements made during this project. Details of the inlets and measurements can be found in Marshall et al. (2005), Phinney et al. (2006) and Phinney et al., (2008¹).

Over the oceans, sulphate and methanesulphonic acid (MSA) are formed from the oxidation of dimethyl-sulphide (CH_3SCH_3) (Seinfeld and Pandis, 1998). Natural organic material, other than MSA, is also present in fine particles in marine environments (e.g. O'Dowd et al., 2004; Phinney et al., 2006), and ships contribute primary organic particles as well as SO_2 (e.g. Hobbs et al., 2000). In the absence of significant anthropogenic aerosols, either from long-range transport or from ship emissions, the marine aerosol consists primarily of sulphate, MSA and sea salt as well as low concentrations of organics. This was the scenario for most of the C-SOLAS measurements (Phinney et al., 2006).

¹ Phinney, L., Lohmann, U., Leaitch, W. R., and Shantz, N.: CCN observations over the summertime North Pacific Ocean: Contributions from DMS and from ship emissions, in preparation, 2008.

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4.1.1 C-SOLAS – time series

The Q-AMS time series of mass concentrations of total particulate sulphur (sulphate + MSA) and total organics for C-SOLAS are shown in Fig. 1. Periods of fumigation from the ship's exhaust, during times of mooring, have been removed as described by Phinney et al. (2006). With the exception of the 18–19 July period, sulphur dominates the fine particle aerosol. The ammonium mass concentrations (not shown) were low over the time series (on average, the ammonium to sulphate ratio was 0.2) and the sulphate was mostly acidic (Phinney et al., 2006).

For 18–19 July, there was a substantial increase in the organics. The wind speeds were relatively low during this time and the air trajectories looped over the ocean. The Q-AMS mass spectrum for this period resembles a mix of a diesel signature and the background organic (Phinney et al., 2006; Phinney et al., 2008¹). It is believed that in this case the aerosol was impacted by emissions of ship traffic in the region.

Three periods selected for the present study of CCN droplet growth analysis are highlighted in Fig. 1. On 16 July, the fine particle aerosol is made up of 90% sulphate and MSA and 10% organic material. On 27 July, the organic fraction was <10%. The third case, 18 July, is during the peak of the period of highest organic fraction, up to 30%.

4.1.2 Simulations of 27 July 2002, from C-SOLAS

The observed particle number distribution for 27 July is shown in Fig. 2a. It is approximated by three lognormal distributions centred at $0.045\ \mu\text{m}$ (mode 1), $0.16\ \mu\text{m}$ (mode 2), and $0.65\ \mu\text{m}$ (mode 3). The chemical mass size distributions from the Q-AMS are compared with the mass size distribution estimated from the SMPS and PCASP in Fig. 2b. The majority of the mass is in sizes $>0.1\ \mu\text{m}$, and there is reasonable agreement between the physical and Q-AMS measurements for mode 2.

The Q-AMS measurements lack sufficient sensitivity to resolve the mass in mode 1, so it is assumed to be 100% sulphuric acid. Both the Q-AMS and the measurements

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from the MOUDI indicate the mode 2 particles are composed mostly of acidic sulphate with a smaller amount of MSA. As the Q-AMS total organic is below detection limit in both the time-of-flight and mass measurement modes of the Q-AMS, the organic is assumed to be zero in this case. Assuming that the MSA behaves as sulphate, then mode 2 is modelled as 100% sulphuric acid. Mode 3 was dominated by sea salt (Phinney et al., 2006), and it is assumed to consist entirely of NaCl. Table 3 shows the assumptions of the chemical breakdown for all modes for all of the field case studies.

Figure 3 shows the scattering cross-sectional area from the simulations compared with the observed voltage normalized to the base voltage from the CCNc. The scale of the right hand axis has been adjusted to match the simulated growth curves with the observed values. As this is a case of nearly pure polydisperse sulphate, this scaling is used for the remaining cases and provides a reference for the other cases to nearly pure sulphuric acid. The error bars reflect the model runs performed at the $S_{\text{eff}} \pm 15\%$, giving the maximum and minimum range in S_{eff} (Sect. 2.3). The minimum activation diameters from the simulations are indicated with vertical lines for the different chamber S_{eff} (Fig. 2a). At the lowest S_{eff} , 0.19%, particles in modes 2 and 3 are activated. For $S_{\text{eff}}=0.34\%$, a small fraction of particles are activated in mode 1. At $S_{\text{eff}}=0.50\%$, just under 50% of the mode 1 particles are activated. Sensitivity tests with the model indicate that including up to 15% soluble or insoluble organic or changing the sulphuric acid in mode 1 to ammonium bisulphate slightly decreases the growth rate at $S_{\text{eff}}=0.34\%$ and 0.50%. These same sensitivity tests for mode 2 do not alter the growth rate at all. Overall, because the organic mass was low and these composition sensitivity tests do not modify the results substantially, we believe that this is a good case to use as a reference for the comparisons between the CCNc observed growth rates and CCNc model results. The next case study (16 July) will be used to verify this reference case.

4.1.3 Simulations of 16 July 2002, from C-SOLAS

C-SOLAS 16 July has a similar size distribution (Fig. 4a) and chemical composition (Fig. 4b) as 27 July. One difference with 27 July is that 16 July has a 65% lower

number concentration in mode 1. Note that on 16 July, mode 2 is split into two modes, 2a at $0.1\ \mu\text{m}$ and 2b at $0.22\ \mu\text{m}$. Both 2a and 2b are assumed to have the same chemical composition.

The Q-AMS measurements for mode 1 are below detection limit. Mode 2 is dominated by acidic sulphate, and organics comprise 10% of the mode 2 mass, exclusive of MSA. There were no MOUDI measurements on 16 July, so mode 3 is assumed to be NaCl, as consistent with the rest of the study.

The simulated and observed CCN growth rates for 16 July are shown in Fig. 5. With the same right-hand axis scaling as used in Fig. 3, the simulated and observed growth rates compare closely, lending confidence that this scaling properly reflects a sulphate reference for the remaining cases. It does not matter what solubility is assumed for the organic; soluble and insoluble organics gave the same results.

4.1.4 Simulations of 18 July 2002, from C-SOLAS

The measured and modelled size distributions for 18 July are shown in Fig. 6a, the case with the highest organic mass concentrations. As discussed in Sect. 4.1.1, the mass spectrum of the organic in this case indicates that there was a significant contribution to this organic from diffuse diesel emissions, and the source was possibly from ships crossing the region. The number distribution is represented in the simulations by four modes: mode 1a centred at $0.035\ \mu\text{m}$, mode 1b at $0.08\ \mu\text{m}$ (hereafter mode 1a and mode 1b will be referred to as “mode 1”), mode 2 at $0.19\ \mu\text{m}$ and mode 3 at $0.65\ \mu\text{m}$. The observations of Phinney et al. (2008¹) show that the geometric diameter of mode 1b is larger than the mode diameters on 27 and 16 July due to the sulphate condensing on primary organics from the diesel emissions.

Figure 6b shows the mass distributions. More information about the chemical composition of mode 1 is available from the Q-AMS for 18 July than 27 or 16 July because mode 1 is larger in size and mass concentration. From the Q-AMS measurements, mode 1 is assumed to be composed of 55% sulphuric acid and 45% organic, and mode 2 is 70% sulphuric acid and 30% organic. These mixtures are assumed to be

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internally mixed based on the observations described by Phinney et al. (2008¹). The MOUDI measurements show mode 3 to be a combination of sulphate, sodium nitrate and sodium chloride. Because all three are quite soluble, the choice is insignificant, and mode 3 is modelled here as pure NaCl.

Figure 7a shows the simulations and the CCNc observations of the droplet growth. The S_{eff} error bars are not shown in this case (for simplicity), but all results fall within that uncertainty. All organics in the distribution are assumed to have the same solubility. The shaded areas encompass the growth rates for assumptions ranging between relatively soluble (solubility=200 g L⁻¹) and insoluble organic (solubility=10⁻⁶ g L⁻¹). The soluble case dissolves readily and accumulates more water early on in the simulation; hence the slightly higher growth curves for the soluble case.

Figure 6a shows the minimum particle diameters activated at the different S_{eff} values (the vertical lines) for the insoluble organic case. The soluble organic case activates at the same or very slightly smaller sizes than for the insoluble case.

At $S_{\text{eff}}=0.19\%$, all of mode 2 is activating, and there is little difference between the soluble and insoluble organic growth curves because there is only 30% organic in this mode and a small number concentration (the growth is less sensitive to changes in the chemical composition in such a situation). There is slightly more sensitivity to the choice of soluble and insoluble scenarios at $S_{\text{eff}}=0.34\%$ and 0.5% , because the activation is occurring in regions of the size distribution where the number concentrations are changing more rapidly with size and the organic fraction is higher. Still the relative differences are very small and the results from this case agree with those of Abdul-Razzak and Ghan (2005) where any solubility assumption is a close enough approximation as long as the organic is not assumed to behave as an inorganic. To illustrate the effect of assuming the organic is equivalent to an inorganic, and because that assumption has been used in some in previous closure studies (e.g., Conant et al., 2004), the results of Fig. 7a are reproduced for the case where the entire distribution is assumed to be sulphuric acid (Fig. 7b). In this case, the simulated CCN scattering cross sections are significantly higher than the corresponding observations, and most of the simulated

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points lie at the edge of or outside the uncertainty defined by the chamber supersaturation. This result indicates the significant error that arises from assuming the organic portion behaves similarly to the inorganic portion, and underscores the need to treat the components of the aerosol separately.

5 The results shown in Fig. 7 indicate that the organic reduces the effectiveness of the individual particle as a CCN compared to pure inorganics. However, compared with the cases from 27 and 16 July, considerably more of the mode 1 particles activate on 18 July at the higher S_{eff} values. This is why the measured growth rates are higher on 18 July than the other two cases. In this case the organic appears to have acted as a
10 substrate for the sulphate to condense upon (e.g. Virkkula et al., 1999; Phinney et al., 2008¹), redefining the size distribution of the sulphate. Thus, on 18 July the organic appears to be ineffective at water uptake but it contributes to the formation of a broader, more effective CCN distribution by redistributing the sulphate.

4.2 Golden Ears Provincial Park (GE) field study

15 The GE provincial park study was part of the Pacific 2001 field campaign (e.g. Li, 2004). The park (covered predominantly with tall coniferous trees) is on the north slope (203 m elevation) of the Lower Fraser Valley and about 45 km northeast of Vancouver, British Columbia. Aerosol and trace gas measurements were made in the southern part of the park from 6–11 August 2001. Table 4 summarizes the measurements made during
20 this project that have been discussed previously by Shantz et al. (2004). Note all times mentioned with respect to GE are in local time (i.e. Pacific Daylight Savings Time). The goal of the present study is to examine the cloud-forming properties from this forest aerosol that is also influenced by urban sources.

4.2.1 GE time series

25 The GE particle chemistry was obtained from open filter samples averaged over 4–8 h. A time series of sulphate (SO_4^-) and organic carbon (OC) mass concentrations is

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shown in Fig. 8. The aerosol was dominated by OC through the sampling period, and increasingly towards the end of the period. The organic aerosol is believed to be a mix of primary and secondary organics from both anthropogenic and biogenic sources; sulphate was likely an ammonium salt (Shantz et al., 2004). Figure 8 also shows water-soluble organic carbon (WSOC) concentration. Because a range of solubilities is tested, WSOC is simply provided as a reference. WSOC is 20–60% of the total OC, which is in the range of what was observed in other studies for polluted, rural and background sites (Carvalho et al., 2003; Decesari et al., 2000; Decesari et al., 2001; Saxena and Hildemann, 1996; Zappoli et al., 1999).

The particulate matter volume for particles $>0.1 \mu\text{m}$ determined from the size distribution measurements is shown in Fig. 8. If we assume a mass density such as $1.0\text{--}1.5 \text{g cm}^{-3}$, there is reasonable agreement between the particulate matter volume and the sum of the OC plus sulphate mass concentrations indicating that most of the OC and sulphate was in the fine particles. The sulphate concentrations at GE were similar or lower than measured over the Pacific Ocean.

The CCNc ΔV (proportional to CCN number concentration) at $S_{\text{eff}}=0.19\%$ switches from being above the particle volume estimate for the first half of the study to being below the particle volume from 9 August onwards (Fig. 8). Although sulphate is low compared to the organic throughout the study, the sulphate to organic ratio is significantly higher during the period prior to 9 August (17%) than after 9 August (4%). The change in the sulphate to organic ratio is a major reason for the change in the correspondence of the CCN with the particle volume, and it indicates a reduced effectiveness of the organic relative to the sulphate as CCN. To examine this further, two periods are selected representing before and after 9 August. These periods are highlighted in Fig. 8 (afternoon of 7 August and morning of 11 August).

4.2.2 Simulations of 7 August 2001, from the GE study

The particle number and volume distributions for the period selected from 7 August are shown in Fig. 9a. The fitted size distribution used in the CCNc simulations is matched

to these observations using 3 modes as shown in Fig. 9b. There is one dominant number mode at about $0.034 \mu\text{m}$ diameter with a very high number concentration, a shoulder mode near $0.07 \mu\text{m}$ diameter, and a large particle mode at $0.36 \mu\text{m}$.

For the purpose of modelling the CCN growth rates in this case, all of the inorganic ions (sulphate, ammonium, nitrate, sodium and chloride) are assumed to be ammonium sulphate. All 3 modes are assumed to have the same composition as each other based on the filter measurements from 7 August: 76% organic, 4% BC (an insoluble organic with molecular weight= 12.01 g mol^{-1} , index of refraction= $(1.82, 0.74)$ and density= 2 g cm^{-3}) and 20% ammonium sulphate. Particles are assumed to be internally mixed based on hygroscopic growth measurements at this site (Aklilu and Mozurkewich, 2004; Shantz et al., 2004).

Figure 10 shows the scattering cross-sectional area from the model output compared with the voltage from the CCNc observations. Absorption by the BC is included in the calculation, however its influence is negligible because it is present in such small quantities. As for the C-SOLAS 18 July case, the shaded areas encompass the growth rates between relatively soluble and insoluble organic. The vertical lines on Fig. 9b show the minimum particle diameters activated for the organics with a solubility of 10^{-6} g L^{-1} and at all 3 S_{eff} 's, these fall within the shoulder mode 2. The 200 g L^{-1} organic activates at slightly smaller sizes (not shown), but still only activates minimum diameters within mode 2. Mode 1 does not activate regardless of the organic solubility. However, the shoulder mode 2 contains a large number of particles (and mass) that may explain the larger growth rates here than during the C-SOLAS cases.

The differences between the modelled soluble (upper curves of shaded sections) and insoluble (lower curves) organic cases are much larger than for the C-SOLAS cases, indicating the importance of organic solubility with increasing organic fraction in the aerosol. For a highly soluble organic, the observations should match the upper curves. The insoluble organic simulation has only sulphate influencing the droplet growth and if the organic does not contribute at all, the observations will match the lower curves. Since the observed growth rates are much higher than the modelled insoluble curves

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and are also substantially higher than those observed for the C-SOLAS cases which had similar or higher sulphate concentrations, the organic aerosol makes a significant contribution to the water uptake.

5 Simulated growth rates for an organic with an effective solubility of about 10 g L^{-1} best match the observed growth rates for 7 August. Note that if we change our initial assumptions, this effective solubility may be altered, such as the sulphate and organic may not be distributed equally across all sizes or the organic may have affected the surface tension or mass accommodation coefficient in ways that are not properly referred in the model.

10 In these simulations, with ammonium sulphate internally mixed with organics, ammonium sulphate dissolves initially and dominates the droplet growth until more water collects on the droplets allowing the slightly soluble organics to start to dissolve. At that point the organic may also contribute to the droplet growth.

15 The WSOC/OC fraction is approximately 25% (Fig. 8). Simulations run with 20% ammonium sulphate, 20% organic with 200 g L^{-1} solubility and 60% insoluble organic to match the WSOC measurements show reasonable agreement between these simulations and the observations. The use of a lower solubility (i.e. $<200 \text{ g L}^{-1}$) worsens the agreement.

4.2.3 Simulations of 11 August 2001, from the GE study

20 Figure 11a shows the number and volume distributions for 11 August. The modelled size distribution, represented by 3 modes, is shown in Fig. 11b. Although the total number concentration is less than on 7 August, there are more particles in mode 2 on 11 August than on 7 August and the volume is substantially higher on 11 August. The increased mass is explained by the condensation of secondary organic aerosol as suggested by the gradual increase in particle mass from 9–11 August. Most of the additional material on 11 August is suspected to be due to the oxidation of monoterpenes as measured concentrations of a known product of terpene oxidation, cis-pinonic acid, also increased during this period (Cheng et al., 2004; Shantz et al., 2004). Each mode

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is assumed to have the same composition, an internal mixture of 85% organic, 2% BC and 13% ammonium sulphate (note that the sum of all inorganic ions is assumed to be ammonium sulphate).

Figure 12 shows the scattering cross-sectional area from the simulations compared with the observations of CCNc droplet growth rates. The vertical lines in Fig. 11b show the minimum activated diameters at the 3 S_{eff} 's (none of the mode 1 particles activate, even when the organic is assumed to be soluble). The growth rates and the number of particles activated are much greater for 11 August than for any of the previously discussed cases.

At $S_{\text{eff}}=0.50\%$ the observed growth curves are close to the insoluble curve, suggesting a smaller contribution from the organics to the water uptake on 11 August than 7 August at this supersaturation. The observed growth rates are higher than the modelled insoluble cases at $S_{\text{eff}}=0.34\%$. The difference between the modelled results at these two S_{eff} values is similar to those from 7 August. For $S_{\text{eff}}=0.5\%$, there may have been competition of water vapour between the particles in the CCNc at the highest supersaturation. Sensitivity tests suggest that an aerosol composed of soluble inorganics may deplete the supersaturation substantially but slightly soluble organics do not as they don't take up water to the same extent. If, however, the organics are soluble, competition for water vapour may drive the supersaturation to a lower value especially at $S_{\text{eff}}=0.5\%$, thus taking up less water, scattering less light and showing a lower voltage. This may indicate a limitation to the growth rate measurement. Overall, the modelled growth rates best agree with the observations for organics with an effective solubility of 5 g L^{-1} , slightly less soluble than 7 August. Keep in mind this is an effective solubility and by changing the initial assumptions, we could find a different value for solubility.

Despite the fact that we found a lower overall effective solubility from the simulations for 11 August compared to 7 August, the WSOC/OC fraction is higher on 11 August compared to 7 August (over 50% – Fig. 8). The WSOC measurements do not give the solubility of the organic, only that it is or is not soluble. Simulations (not shown) are run with 13% ammonium sulphate, 47% organic with 5 g L^{-1} solubility and 40%

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insoluble organic based on the WSOC measurements. Reasonable agreement with the observations is achieved for this situation. However, there are uncertainties in these measurements such as the long filter sampling time, variations in the chemical composition with respect to size and assumptions about the organic. These simulations are fairly sensitive to the amount of sulphate present and if we decrease the amount of ammonium sulphate from 13% to 7%, the WSOC solubility would be higher.

Overall, these results indicate that the organic is less effective as CCN compared to sulphate, but in these situations with relatively low sulphate, the organic particles do act as CCN. Even with lower sulphate masses on 11 August, there is an increase in the number of CCN as there is an increase in the particle number concentration, and similar to C-SOLAS 18 July, it is suspected that the influence of the organic on the sulphate contributes to the increased growth rates. This argument depends on whether the organic coats the sulphate or whether the sulphate is readily available to the water vapour as this model assumes.

5 Implementing a single κ parameter representing water uptake (κ)

Recently Petters and Kreidenweis (2007) discussed the hygroscopicity parameter κ that encompasses water uptake capabilities (both above and below supersaturation) and simplifies the physical and chemical properties needed for aerosols in cloud microphysical models. Because the majority of organic constituents are unknown and thus properties such as molecular weight and solubility are also unknown, a κ value is useful in describing the water activity of aerosols from different sources. A version of the present CCNc model using κ was produced (hereafter “ κ -CCN model”) in order to compare estimated the total κ values (κ_{tot}) for the five case studies discussed above and to determine the organic κ values (κ_{org}) from the field measurements.

For the nearly pure sulphate case from C-SOLAS (27 July), the best agreement between the CCNc model and the κ -CCN model is for $\kappa_{\text{tot}}=0.7$ (see Table 5 for a summary of all κ values determined) and a slightly lower value of $\kappa_{\text{tot}}=0.65$ for 16 July due to the

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small amount of organic. These values are slightly below that suggested by Petters and Kreidenweis (2007) for H_2SO_4 ($\kappa=0.9$), which may be due to the differences in the way the water activity is calculated in the CCNc model compared to Petters and Kreidenweis (2007).

In Fig. 13a, the results for modelling C-SOLAS 18 July using κ are shown. This case, with a marine aerosol mixed with organic, shows the best agreement with $\kappa_{\text{tot}}=0.35$. The κ_{tot} can be constrained in this case to 0.2–0.6.

Petters and Kreidenweis (2007) show that for aerosols that consist of a few chemical species, the individual chemical κ values can be weighted by the volume fraction to determine the κ_{tot} (using $\kappa_{\text{tot}}=\sum \varepsilon_i \kappa_i$ where ε_i is the volume fraction and i represents the individual species). Using the measured aerosol composition from the Q-AMS for 18 July (assuming sulphate has $\kappa=0.7$ based on 27 July results) and $\kappa_{\text{tot}}=0.35$, then κ_{org} is found to be zero. In other words, for 18 July over the North Pacific Ocean the organic does not contribute to the water uptake.

GE forest 7 August show the best agreement between the observed and modelled growth rates using $\kappa_{\text{tot}}=0.16$ (Fig. 13b). Assuming the ammonium sulphate has a $\kappa=0.61$ (Petters and Kreidenweis, 2007), κ_{org} is estimated as 0.05. For GE 11 August, the estimated κ_{tot} is 0.08, which falls in the range of κ values provided by Petters and Kreidenweis (2007) for the oxidation products of monoterpenes. The 11 August κ_{org} is estimated to be 0.02, lower than the κ_{org} for 7 August. The measurements of water uptake at relative humidities <100% that Aklilu and Mozurkewich (2004) made during the GE study gave values of κ_{tot} (called “ b ” in their paper) in the range of 0.04 to 0.10, with most values between 0.05 and 0.07. These fall within the uncertainty range of the values determined here. The uncertainties in our measurements do not allow the absolute determination of κ , but based on the time series, it seems that κ is decreasing over time during the GE study based on these 2 case studies (consistent with Aklilu and Mozurkewich, 2004). This could reflect a change in the properties of the organic aerosol over time, or it might be the result of a change in the nature of the mixing of the organic and inorganic components.

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6 Summary and conclusions

A cloud condensation nucleus counter (CCNc) is used to investigate the role of organics as CCN by comparing the measured growth rates of cloud droplets to those simulated using a kinetic model (CCNc model). Five cases of ambient measurements of CCN, particle size distribution and particle chemistry are examined, three from the North Pacific Ocean and two from a forest in British Columbia, Canada. One of the marine cases, for which the aerosol was dominated by sulphate (27 July 2002) is used as a reference for the CCNc model. This reference case is verified using a second marine case also dominated by sulphate (16 July 2002). In these cases with the fine particle mass $\geq 90\%$ inorganic, the organic is of no consequence.

The organic fraction is higher in the third marine case (30% of the total mass – 18 July 2002). Evidence suggests that this aerosol was a mix of somewhat aged ship emissions and marine aerosol. The comparison of the observations and simulations indicates that the sulphate in these aerosol particles was responsible for most of the water uptake. However, the organics influenced the CCN growth rates by enhancing the number concentration of particles containing sulphate and redistributing the sulphate to larger and more CCN active sizes.

Two forest cases (from Golden Ears Provincial Park) had higher number concentrations, much higher organic mass fractions (80–90% of the total mass), and similar or smaller sulphate mass concentrations than the marine cases. The observed CCN growth rates from these two cases are much higher than for the marine cases. This is due to an increase in the number concentrations and some influence of the organic material on the water uptake. The best fits of the simulations to the observations were obtained for organics with solubilities of 5–10 g L⁻¹. This result assumes an accommodation coefficient of unity and a specified surface tension. Differences in those two quantities from the modelled values might also contribute to the water uptake.

Values of the parameter κ (Petters and Kreidenweis, 2007) are determined for the five cases using the κ -CCN model. For the combined aerosol, the Pacific marine values

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are 0.35–0.7 and the organic components are found to be negligible. For the forest aerosol, the total κ value is found to be 0.08–0.16 with an organic κ of 0.02–0.05. The observations suggest that the κ values decreased with time during the GE study as the aerosol accumulated a larger organic mass.

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Table 1. Summary of properties for adipic acid, a dicarboxylic acid.

Alternate name ^a	Hexanedioic acid
Number of carbons	6
Molecular formula ^a	HO ₂ C(CH ₂) ₄ CO ₂ H
Molecular weight (g mol ⁻¹) ^a	146.14
Index of refraction ^b	Real: 1.45, imaginary: 0
Density (g cm ⁻³) ^a	1.36
Solubility (g L ⁻¹) ^c	25
Ionic dissociation factor ^d	1
Osmotic coefficient	1
Surface tension equation in model ^e	$\sigma_s = \sigma_w \times \left(1 - b \times \log_{10} \left(\frac{c_{\text{org}}}{a} + 1 \right) \right)$ <p> σ_s=surface tension of solution at the air-water interface σ_w=surface tension of water at the air-water interface a and b=coefficients, indicated in next row c_{org}=molarity of organic acids in the water </p>
Surface tension coefficients ^f	$a=0.019$, $b=0.105$
Growth factor at 95% ^g	1.01 ^h

^a CRC (1974).

^b Unavailable for adipic acid, used value for succinic acid (CRC, 1974).

^c Saxena and Hildemann (1996).

^d Bilde and Svenningsson (2004); Broekhuizen et al. (2004); Corrigan and Novakov (1999); Cruz and Pandis (1997).

^e Equation from Ervens et al. (2004).

^f Ervens et al. (2004) published these coefficients, that had been fitted to the data from Shulman et al. (1996) using the equation in the previous row of this table.

^g The growth factor is expressed as a ratio of the wet radius of the droplet over the dry radius of the original particle or CCN.

^h Joutsensaari et al. (2001); Prenni et al. (2001) state that the growth factor at 90% relative humidity for adipic acid is 1. Thus, in the model it was assumed to be 1.01.

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Table 2. Summary of the relevant instrumentation from the C-SOLAS field study.

Measurement	Instrument	Size range	Frequency of measurements
Particle Size Distributions	SMPS	0.008–0.3 μm diameter in 26 size bins	Every 5 min
	PCASP	0.15–3 μm diameter in 15 size bins	Every 5 min
Particle chemical constituents	MOUDI	0.03–18.0 μm diameter on 12 stages	4–10 h per sample
	Q-AMS	0.06–0.6 μm vacuum aerodynamic diameter (D_{va})	Every 15 min
Water Uptake	CCNc	All particle sizes that would activate in the supersaturation range 0.19–0.5%	Every \approx 35 s

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Table 3. Assumptions made in the simulations about the size distribution based on the SMPS and PCASP measurements and chemical composition based on Q-AMS and MOUDI measurements (C-SOLAS) and filter measurements (from Golden Ears –GE). Details are shown for each of the 5 case studies and for each mode of the distributions.

	Number Median Diameter (μm)	Number Conc. (cm^{-3})	Composition Assumption
C-SOLAS 27 July 2002: 16:00–17:30 PDT			
Mode 1	0.045	641	H_2SO_4
Mode 2	0.160	147	H_2SO_4
Mode 3	0.650	1.3	NaCl
Total		789	
C-SOLAS 16 July 2002: 18:45–20:00 PDT			
Mode 1	0.040	435	H_2SO_4
Mode 2a	0.100	36.5	10% organic and 90% H_2SO_4
Mode 2b	0.220	112	10% organic and 90% H_2SO_4
Mode 3	0.650	1.5	NaCl
Total		585	
C-SOLAS 18 July 2002: 12:15–15:15 PDT			
Mode 1a	0.035	146	45% organic and 55% H_2SO_4
Mode 1b	0.080	754	45% organic and 55% H_2SO_4
Mode 2	0.190	133	30% organic and 70% H_2SO_4
Mode 3	0.650	1.8	NaCl
Total		1035	
GE 7 August 2001: 17:00–19:00 PDT			
Mode 1	0.034	11221	
Mode 2	0.070	2196	76% organic, 4% black carbon & 20% ammonium sulphate
Mode 3	0.360	141	
Total		13558	
GE 11 August 2001: 00:00–01:00 PDT			
Mode 1	0.050	3471	
Mode 2	0.090	3943	85% organic, 2% black carbon and 13% ammonium sulphate
Mode 3	0.360	111	
Total		7525	

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Table 4. Summary of the relevant instrumentation from the Golden Ears field study.

Measurement	Instrument	Size range	Frequency of measurements
Particle Size Distributions	SMPS	0.01–0.2 μm diameter in 25 size bins	Every 5 min
	PCASP	0.15–3 μm in 15 size bins	Every 5 min
Particle chemical constituents	Single Teflon filters	Total mass of all particles that impact	4–14 h/sample
	Single Quartz filters	Total mass of all particles that impact	4–14 h/sample
Water Uptake	CCNc	All particle sizes that would activate in the supersaturation range 0.19–0.5%	Every \approx 35 s

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Table 5. Hygroscopicity parameter (κ) values determined for each case study.

	C-SOLAS			Golden Ears	
	27 July	16 July	18 July	7 August	11 August
κ_{tot}	0.7	0.65	0.35	0.16	0.08
Range of $\kappa_{\text{tot}}^{\text{a}}$	0.4–1.1	0.4–1.0	0.2–0.6	0.09–0.27	0.04–0.17
$\kappa_{\text{org}}^{\text{b}}$	–	0	0	0.05	0.02
Range of $\kappa_{\text{org}}^{\text{c}}$	–	– ^d	0–0.5	0.001–0.2	0.001–0.11

^a Simulations are performed for κ_{tot} at $S_{\text{eff}} \pm 15\%$ to reflect the error in the calculated supersaturation. The ranges of κ values are determined by finding κ_{tot} simulations that fall within these error bars.

^b The κ_{org} are found by using the measured composition of the aerosol and κ_{tot} and applying the mixing rule $\kappa = \sum \varepsilon_j \kappa_j$.

^c The range of κ_{org} is determined as those that fall within the error bars of $S_{\text{eff}} \pm 15\%$.

^d Because there was only 10% organic in this case, it is difficult to determine the range of κ_{org} 's.

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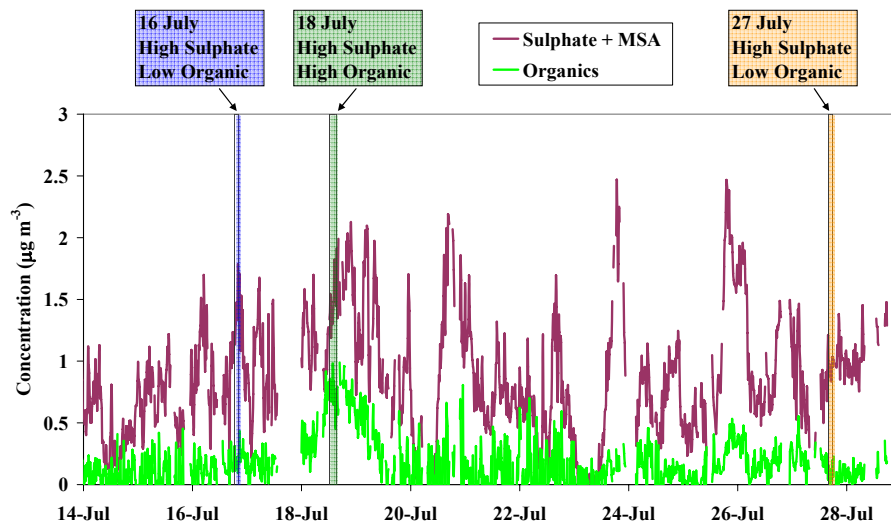


Fig. 1. C-SOLAS time series of Q-AMS mass concentration, with the diesel exhaust spikes removed. Shown are organics and total particulate sulphur (sulphate plus methanesulphonic acid (MSA)). Indicated with arrows are the periods of interest for this work. Each date marker on the x-axis indicates midnight in Pacific Daylight Savings Time (i.e. not local time in the ocean, which was 1.5 h behind PDT).

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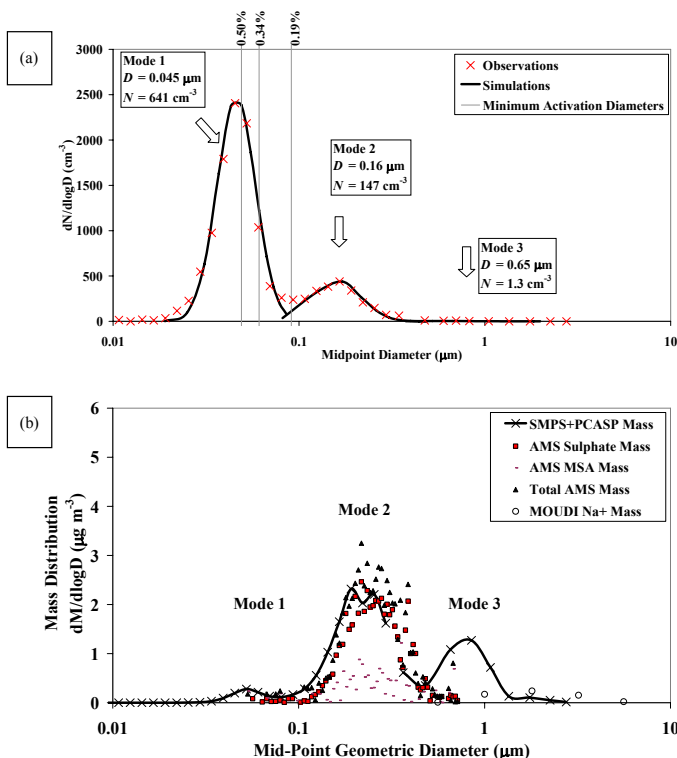


Fig. 2. (a) C-SOLAS 27 July 2002 observed size distribution (from the SMPS and PCASP measurements) and fitted size distribution used in the CCNc simulations. Indicated are the number median diameter, D , and the number concentration, N , for the mode. The total number concentration for this period is 789 cm^{-3} . The vertical lines correspond with the minimum activation diameter found in the simulations at the supersaturations indicated at the top of the graph (i.e. all particles at this diameter and larger are activated in the simulations). (b) Mass distributions from the MOUDI filter measurements and Q-AMS “time of flight” measurements and estimated from the SMPS and PCASP measurements. The SMPS and PCASP masses are estimated from the size distributions (Fig. 2a) using a density approximated from the Q-AMS chemical measurements. The Q-AMS mass size distributions are converted from vacuum aerodynamic diameter (D_{va}) to geometric diameter by dividing the D_{va} by this density.

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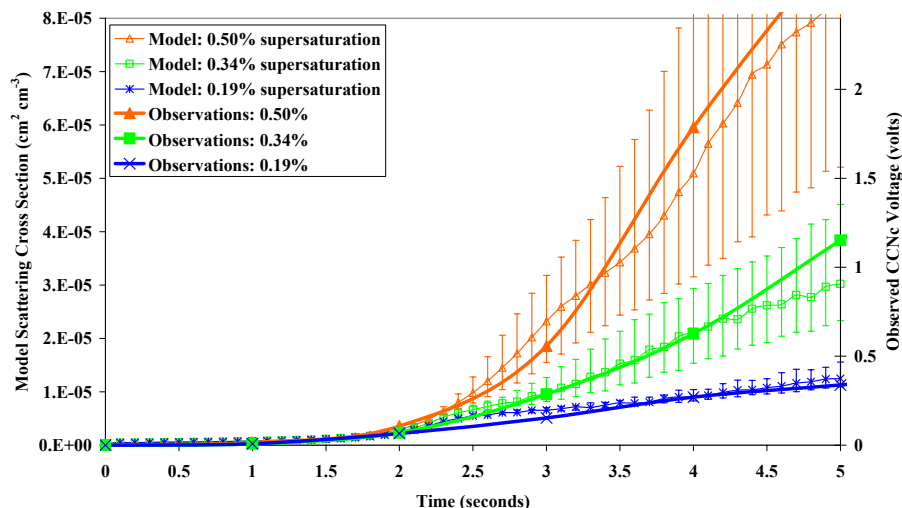


Fig. 3. C-SOLAS 27 July 2002 observations and simulations of CCNc droplet growth rates. The x-axis is the sample time in seconds as the particles activate and grow into cloud droplets. The observed growth rates at 3 supersaturations (shown as a voltage) from the CCNc are the thicker lines, corresponding to the y-axis on the right. The other curves with the many symbols are the modelled curves, on the left y-axis shown as a scattering cross section. These y-axes are scaled to match in this case. The error bars are simulations performed at the effective supersaturation $\pm 15\%$ to reflect the error in the calculated supersaturation.

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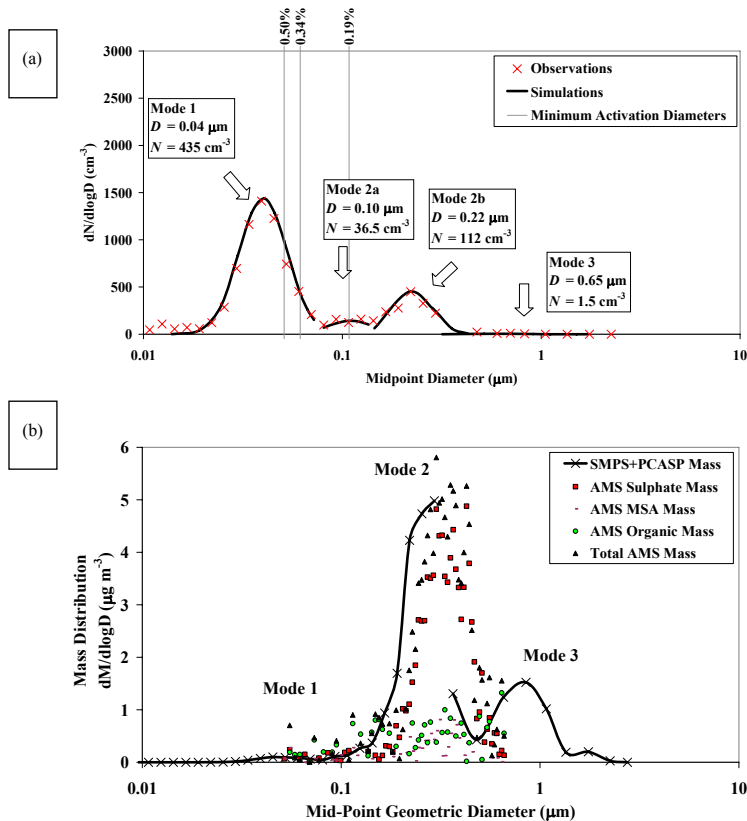


Fig. 4. (a) C-SOLAS 16 July 2002 observed size distribution (from the SMPS and PCASP measurements) and fitted size distribution used in the CCNc simulations. The total number concentration for this period is 585 cm^{-3} . Mode 2a and mode 2b are assumed to have the same chemical composition. (b) Mass distributions from the Q-AMS “time of flight” measurements and estimated from the SMPS and PCASP measurements. There were no MOUDI measurements on 16 July. See Fig. 2 caption for further details.

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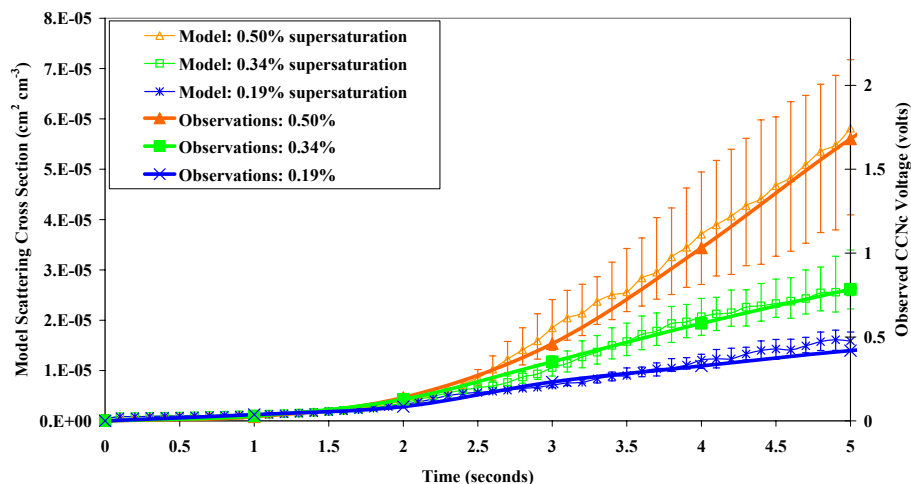


Fig. 5. C-SOLAS 16 July 2002 observations and simulations of CCNc droplet growth rates. The axis lineup for 27 July (Fig. 3) is tested in this case to give confidence in these comparisons between the observations and simulations. See Fig. 3 caption for details.

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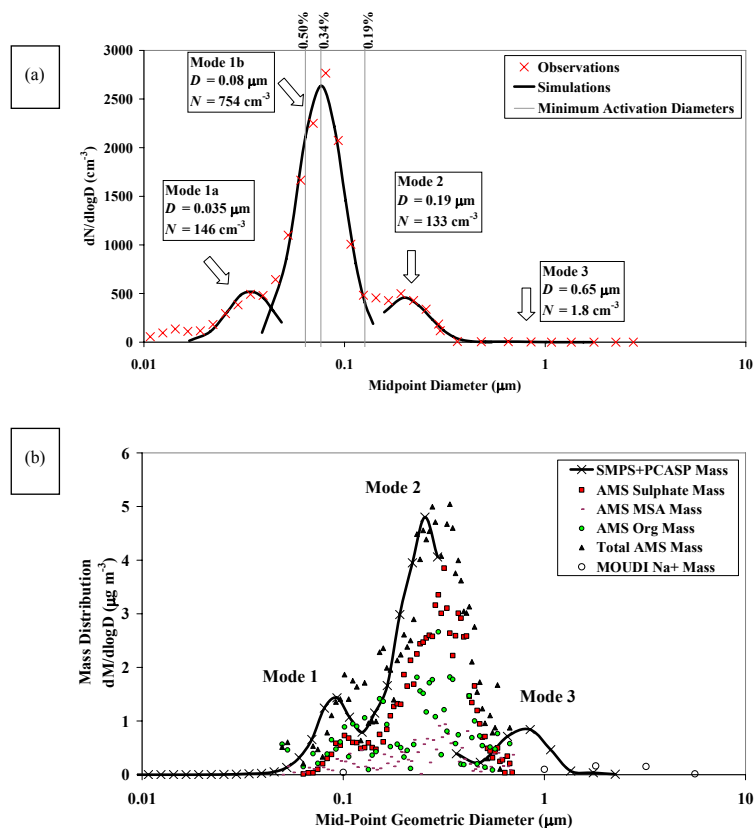


Fig. 6. (a) C-SOLAS 18 July 2002 observed size distribution and fitted size distribution used in the CCNc simulations. The total number concentration for this period is 1035 cm^{-3} . Mode 1a and mode 1b are assumed to have the same chemical composition. The vertical lines show activation for the insoluble case. (b) Mass distributions. See Fig. 2 caption for further details.

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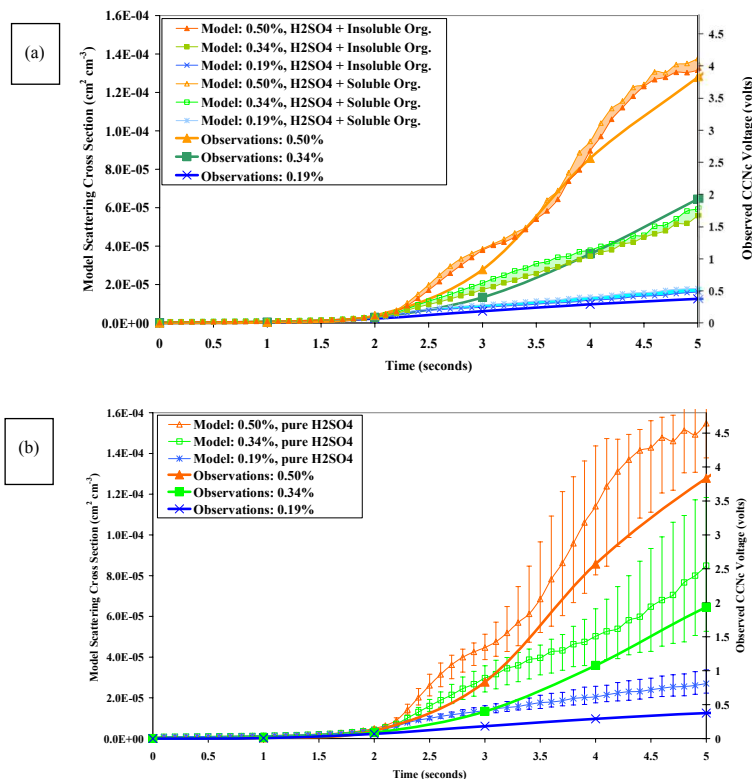


Fig. 7. C-SOLAS 18 July 2002 observations and simulations of CCNc droplet growth rates for **(a)** organics included, with two solubilities. The shaded section shows the range of solubilities of the organic, from very soluble (top of the shaded sections with solubility= 200 g L^{-1}) to insoluble (bottom of the shaded sections with solubility= 10^{-6} g L^{-1}) and **(b)** the assumption that the entire distribution was pure sulphuric acid. See Fig. 3 caption for more details.

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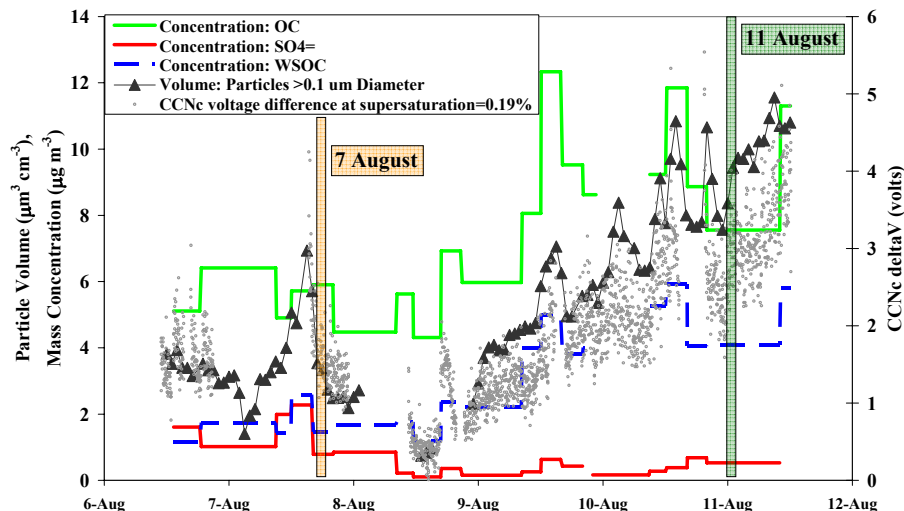


Fig. 8. Golden Ears (GE) time series of filter mass concentrations of organic carbon (OC) and water-soluble organic carbon (WSOC) from the quartz filter samples and sulphate ($\text{SO}_4=$) from the Teflon filter samples. Also included is the particulate volume from the SMPS and PCASP and deltaV (see Section 2.3 text) from the CCNc at 0.19% effective supersaturation. Indicated are the periods of interest for this work. Each date marker on the x-axis indicates midnight in Pacific Daylight Savings Time (i.e. local time).

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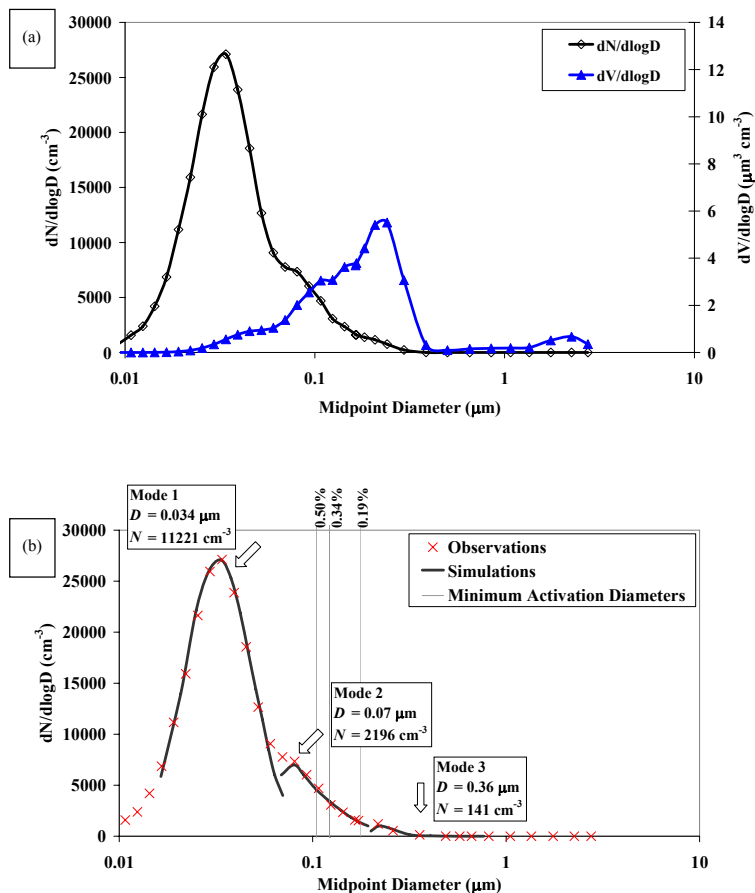


Fig. 9. (a) Golden Ears 7 August 2001 number and volume distributions from SMPS and PCASP. (b) The observed size distribution and fitted size distribution used in the CCNc simulations. The total number concentration for this period was $13\,558 \text{ cm}^{-3}$. The vertical lines show activation for the insoluble case. See Fig. 2a for further details.

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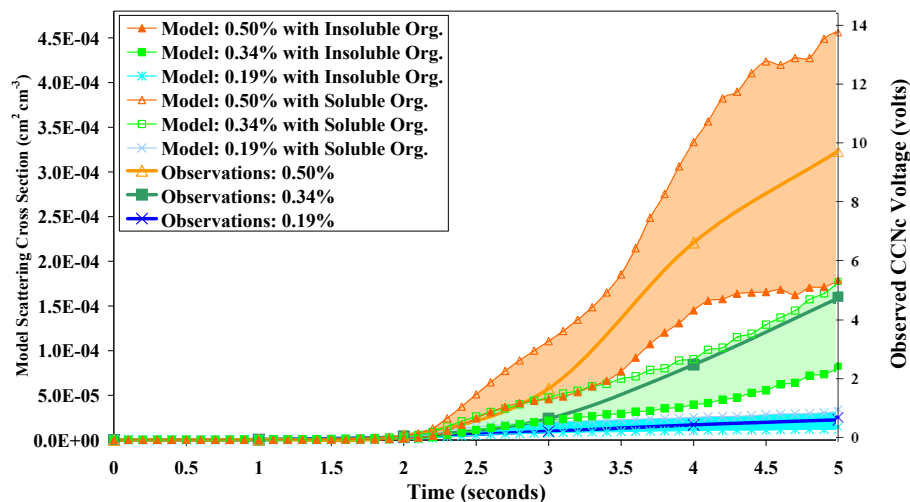


Fig. 10. Golden Ears 7 August 2001 observations and simulations of CCNc droplet growth rates, showing soluble and insoluble scenarios. See Figs. 3 and 7b figure captions for details.

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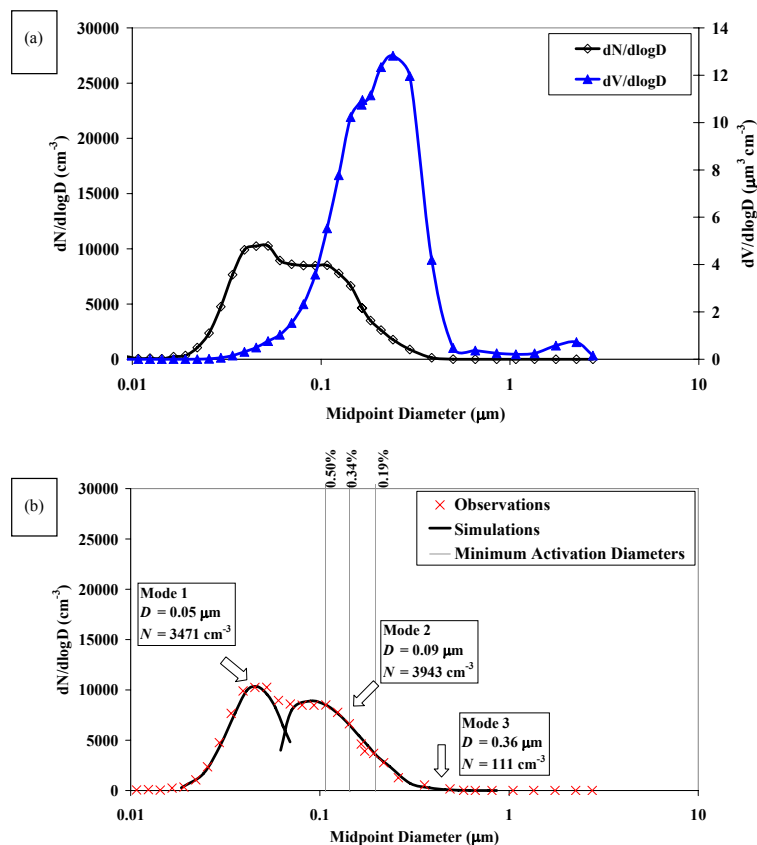


Fig. 11. (a) Golden Ears 11 August 2001 number and volume distributions from SMPS and PCASP. (b) The observed size distribution and fitted size distribution used in the CCNc simulations. The total number concentration for this period is 7525 cm^{-3} . The vertical lines show activation for the insoluble case. See Fig. 2a for further details.

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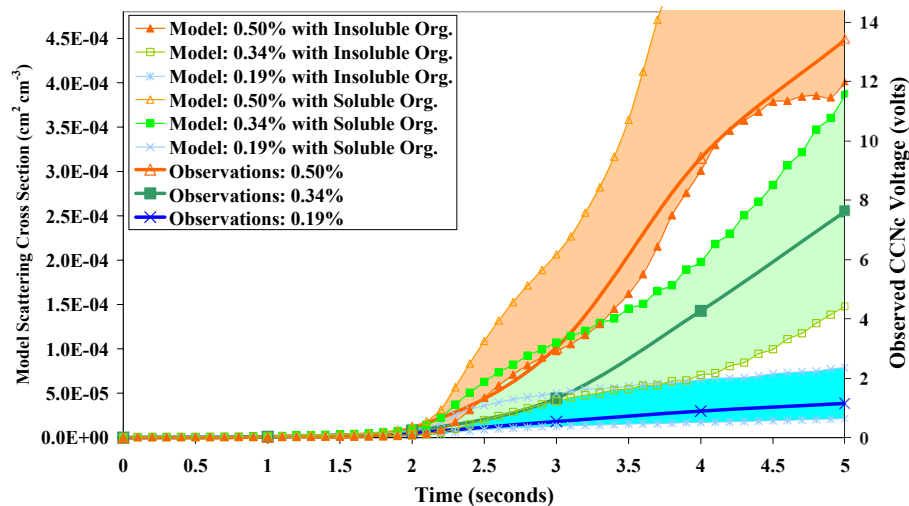


Fig. 12. Golden Ears 11 August 2001 observations and simulations of CCNc droplet growth rates, showing soluble and insoluble scenarios. See Figs. 3 and 7b figure captions for details.

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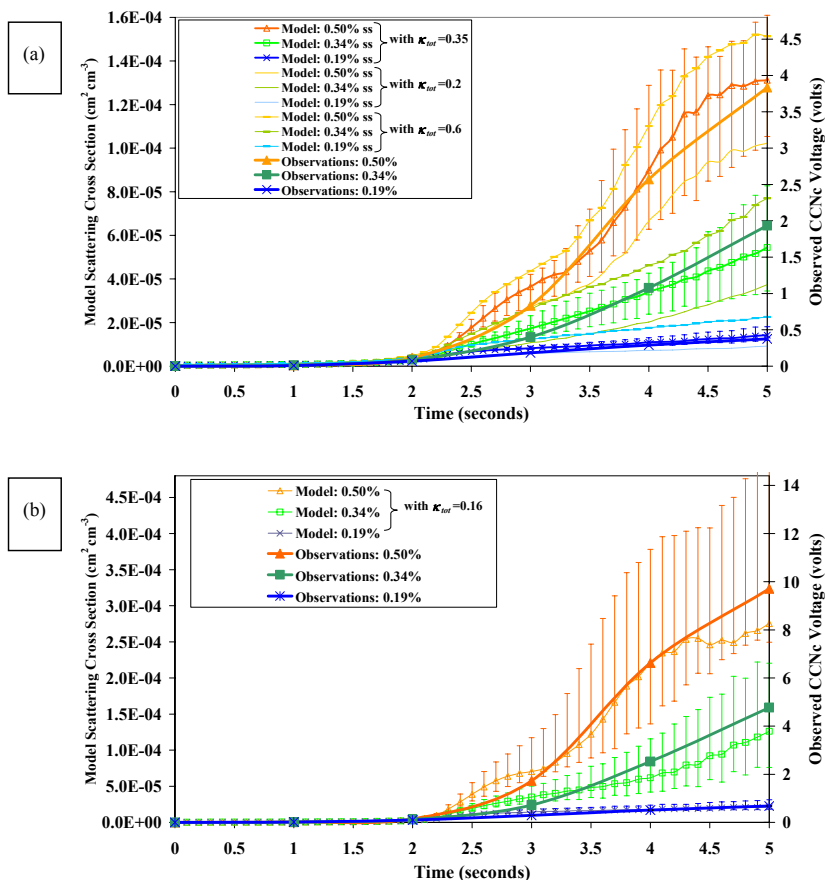


Fig. 13. Observations and simulations of CCNc droplet growth rates assuming a κ_{tot} for the entire distribution for **(a)** C-SOLAS 18 July and **(b)** Golden Ears 7 August.

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