



Multiwavelength and polarization lidar measurements of Asian dust layers over Tsukuba, Japan: a case study

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**Lidar measurements
of Asian dust**

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Multiwavelength and polarization lidar measurements of Asian dust layers over Tsukuba, Japan: a case study

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Abstract

Elastic and polarization lidar was used to measure the vertical profiles of aerosol backscattering coefficients at wavelengths of 355, 532, 735, and 1064 nm, and the depolarization ratio at 532 nm in order to study the aerosol properties in the free troposphere over Tsukuba, Japan, in 2006. An elevated dust layer was observed at altitudes between 3 and 8.5 km on 1 April during the Asian dust period. The wavelength exponents of the aerosol backscattering coefficient (k) were -0.1 to 0.5 , and the depolarization ratio (δ_p) was 25% for the dust layer, suggesting the predominance of supermicrometer-sized (coarse mode) nonspherical particles. An aerosol layer observed at altitudes between 1.5 and 5 km on 19 October during the less-dust period exhibited the values of $k=1.0$ to 1.6 and $\delta_p=1$ to 13%, suggesting the predominance of submicrometer-sized (fine mode) particles. In those layers, the values of k and δ_p varied with height; they were also negatively correlated, suggesting that the proportion of the coarse nonspherical particles to total particles varied. The particle size distributions estimated from the observed values and the theoretical computation revealed number mode radii of $0.3\text{ }\mu\text{m}$ for the coarse mode and $0.1\text{ }\mu\text{m}$ for the fine mode, assuming bimodal distribution. These results were consistent with those obtained from the sky-radiometer measurements, although they revealed another mode in the larger radius. The column volume concentration derived from the lidar was 48% lower than that derived from the sky-radiometer on 1 April and 16% lower on 19 October. The optical thickness derived from the lidar was 12% lower than that obtained from the sky-radiometer on 1 April and 29% higher on 19 October. Further case study is necessary to validate the method for estimating aerosol properties based on the lidar measurement.

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1 Introduction

Mineral dust particles that originate from the arid and semi-arid lands of the Asian Continent (Asian dust) are often transported and spread over East Asia and the North Pacific, chiefly during the spring season (Iwasaka et al., 1988; Duce, 1995). The number of studies and programs focusing on the Asian dust has grown over the past few decades as the importance of regional climate and global biogeochemical cycles has become increasingly apparent (Huebert et al., 2003; Arimoto et al., 2006; Mikami et al., 2006). The dust can affect the Earth's radiation balance by scattering and absorbing solar and terrestrial radiation (Aoki et al., 2005; Shi et al., 2005), and it can affect cloud properties by acting as cloud condensation nuclei and ice nuclei (e.g. Isono et al., 1959; Yamagata et al., 2004).

The impact of dust on the climate critically depends on particle properties such as number concentration and size distribution, shape, chemical composition, and vertical distribution (e.g. Liao and Seinfeld, 1998; Quijano et al., 2000). However, these properties have not been adequately researched, particularly in the free troposphere, because of difficulties in measurement. Lidar is a useful tool for measuring the vertical profiles of the aerosol optical properties. Multiwavelength and Raman lidar measurements provide the aerosol optical properties that can be used for estimating particle size distribution and the refractive index (e.g. Müller, 1999, 2000; Veselovskii et al., 2005). The polarization lidar measurement provides the particle shape and phase (Sassen, 2000). However, few studies have used multiwavelength and polarization data for estimating the aerosol properties.

This paper provides a case study of the multiwavelength and polarization lidar measurement of the aerosol optical properties in the free troposphere over Tsukuba, Japan. We studied the aerosol optical properties by focusing on the relation between the wavelength dependence of the backscattering coefficients and the depolarization ratio. Based on the results of the measurement and the theoretical computation of the aerosol optical properties, we estimated the fraction of fine and coarse particles in the

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backscattering coefficient and the size distribution. The estimated size distributions were compared with those obtained from the sky-radiometer measurement.

2 Measurements

Multiple wavelength and polarization lidar measurements were carried out over Tsukuba (36.0° N, 140.1° E), which is located 60 km northeast of Tokyo. The characteristics of the lidar are presented in Table 1. The multiwavelength lidar transmits laser pulses at wavelengths of 355 nm, 532 nm, 735 nm, and 1064 nm by use of three laser sources: 1) Nd:YAG laser that transmits laser pulses with a repetition rate of 20 Hz at 355 nm with a pulse energy of 80 mJ, and 532 nm with a pulse energy of 40 mJ; 2) Nd:YAG laser that transmits pulses at 1064 nm with 550 mJ/pulse at 20 Hz; and 3) Nd:YAG pumped-dye laser that transmits pulses at 735 nm with a 16 mJ/pulse at 10 Hz. The light backscattered from atmospheric molecules and particles at wavelengths of 355 nm, 532 nm, and 1064 nm is collected by using a telescope with a diameter of 50 cm and a 3.0 mrad field of view (FOV). The light backscattered at a wavelength of 735 nm is collected by using a telescope with a diameter of 35 cm and a FOV of 2.2 mrad. The polarization lidar transmits laser pulses at 532 nm with 300 mJ/pulse and 10 Hz. The two orthogonal components of the linearly polarized elastic backscattering at 532 nm are collected with telescopes with a diameter of 20 cm (FOV of 3.0 mrad) and 40 cm (FOV of 2.0 mrad). The geometrical configurations of the lidar are in biaxial except for the 20-cm telescope of the polarization lidar, which is coaxial. The lowest altitude of the measurement where the laser beam fully overlaps with receiver's FOV is 1.45 km for the multiwavelength lidar and 0.1 km for the polarization lidar. The collected light is detected with photomultiplier tubes (PMTs) that are operated in photon counting (PC) mode and analog-to-digital conversion (A/D) using transient recorders (Licel GbR) with resolutions of 7.5 m in the vertical direction, and 1.5 or 3 min in time. We use PC data above an altitude of 2 km and A/D data below that altitude. The A/D data are connected to PC data, multiplying a constant that is obtained by dividing the

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photon count by analog voltage at the height point of the connection.

The quantities obtainable with the multiwavelength lidar are the backscattering ratio (R) and the particle backscattering coefficients (β_p) at 355 nm, 532 nm, 735 nm, and 1064 nm. The value of R is proportional to the aerosol mixing ratio, and that of β_p is proportional to the concentration, assuming that the aerosol size distribution and the backscattering efficiency remain constant. To calculate these values, we used the iterative method (Fernald et al., 1972; Ismail et al., 1998). In the calculation, we used the height-dependent particle extinction-to-backscatter ratios (lidar ratios, S_p) that were estimated from the particle depolarization ratios. The estimation method is described in Sect. 4. The backscatter signals were calibrated at altitudes of 10 to 15 km, assuming that the R was 1.01 at a wavelength of 355 nm, 1.03 at a wavelength of 532 nm, 1.09 at a wavelength of 735 nm, and 1.27 at a wavelength of 1064 nm, based on the observations at midlatitude (Russell et al., 1979, 1982) and the wavelength dependence of the backscattering coefficient for the continental aerosol model proposed by Ackermann (1998). The molecular backscattering and extinction coefficients were calculated from the atmospheric temperature and pressure data that were obtained with the radiosondes launched 250 m northeast of the lidar at 20:30 Japan standard time (JST). The vertical resolution of the analyzed data was reduced to 150 m to improve the signal-to-noise ratio. The uncertainty in the lidar-derived quantities was estimated using Poisson statistics for the observed photon counts.

We calculated the wavelength exponents (k) of the aerosol backscattering coefficient between the four wavelengths (355 to 532 nm, 532 to 735 nm, and 532 to 1064 nm), assuming the relation $\beta_p(\lambda) \propto \lambda^{-k}$. This value was negatively correlated with the mean particle size; it was higher for the smaller particles and lower for the larger particles (e.g. $k \sim 1.5$ for the submicrometer-sized particles and $k \sim 0$ for the supermicrometer-sized particles). We noted that the measured values of β_p at 1064 nm had large uncertainty in this study because of the low signal intensity at 1064 nm at the calibration height. The measurement uncertainties in β_p were on average 25% at 355 nm, 50% at 532 nm, 6% at 735 nm, and over 50% at 1064 nm for the studied altitude ranges.

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The quantities obtainable with the polarization lidar are β_p and the particle linear depolarization ratio (δ_p) at 532 nm, defined by the ratio of the perpendicular to the total components in β_p with respect to the polarization plane of the emitted laser light. This value is a measure of the particles' nonsphericity if they are comparable to or larger than the measurement wavelength (e.g. Mishchenko and Sassen, 1998). For example, mineral dust usually has high δ_p values (>10%) because it is mostly supermicrometer-sized and nonspherical (Okada et al., 2001). In contrast, spherical particles (e.g. droplets) and submicrometer-sized particles have low δ_p values (<10%).

3 Results

3.1 Case on 1 April 2006: Asian dust layer

Figure 1a presents the vertical distributions of R , k , and δ_p that were obtained from 00:01 to 05:03 JST on 1 April 2006. High values of R exceeding 1.5 at 532 nm were obtained at altitudes between 3.0 and 8.5 km, with a peak value of 2.45 at an altitude of 3.8 km. The optical thickness was estimated to be 0.27 at 532 nm for this altitude range. The values of k were low, ranging from -0.1 to 0.5 (thin lines in the right panel of Fig. 1a), suggesting that supermicrometer-sized (coarse) particles were predominant. The values of δ_p were as high as 25% (thick solid line in the right panel of Fig. 1a), indicating that coarse nonspherical particles were predominant. It should be noted that the values of δ_p and k varied with height, and they were negatively correlated (see also Fig. 3). These results suggested that the fraction of coarse nonspherical particles that have higher δ_p and lower k values than those of the fine particles varied with height. This is discussed further in Sect. 4. The coarse nonspherical particles were probably mineral dust particles originating from the arid or semi-arid lands of the Asian continent: back trajectory analysis revealed that the air parcels had been transported over the arid and semi-arid lands of the Asian continent and over the northern part of the Taklimakan Desert three days earlier (Fig. 2a). Details of the method for

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computing the trajectory are given by Sakai et al. (2000). The reasons why we calculated many trajectories located around the lidar site (25 points in steps of 0.5 degrees for longitude and latitude) are to consider the possibility of the mixing of different air masses that had passed over the source areas as well as to reduce the uncertainty of the trajectory that can be arisen from the coarse resolution of the meteorological data (1.25 degrees for longitude and latitude) used in the computation. In addition to the trajectory analysis, the aerosol transport model of Navy Aerosol Analysis and Prediction System (NAAPS) also predicted the presence of dust with a mass concentration of 15 to 50 μgm^{-3} between 2 and 7 km in height and sulphates with 1–2 μgm^{-3} between 2 and 3 km over the lidar site (http://www.nrlmry.navy.mil/aerosol_web/globaer/ops.01/tsukub/200604/2006032700_2006040100_tsukub.gif). For comparison of the optical properties of the Asian dust, Murayama et al. (2004) observed k of 0.3 to 1.2 at 355 to 532 nm and δ_p of 18% at altitudes between 3 and 5 km over Tokyo during the Asian dust period. Sugimoto and Lee (2006) observed $k \sim 0.5$ at 532 to 1064 nm and $\delta_p \sim 10\%$ at an altitude of 2 km for the Asian dust layer over Korea. Our measured values were a little lower for k (–0.1 to 0.5) and higher for δ_p (25%) than their reported values. This difference was possibly due to the higher fraction of coarse particles in the dust layer that we measured. Ansmann et al. (2002) observed k ranging from –0.5 to 0 at 355 to 532 nm and δ_p ranging from 10% to 20% at altitudes between 1 and 5 km for the Saharan dust layers over Europe. These values were a little lower than our results, possibly due to the differences in the fraction of coarse particles and the optical properties between the Saharan dust and Asian dust over the remote sites.

3.2 Case on 19 October: fine particle layer

Figure 1b plots the vertical profiles obtained during the less-dust period for 02:01 to 06:45 JST on 19 October 2006. The profile of R indicated peak values of 1.3 at an altitude of 4.3 km, and 1.9 at 2 km at 532 nm. The optical thickness was estimated to be 0.09 at 532 nm between 1.45 and 4.7 km. The values of k ranged from 1.0 to 1.6 in those regions, suggesting that submicrometer-sized (fine) particles were predominant.

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The values of δ_p were 1 to 13%, indicating that spherical and/or small (compared with the laser wavelength of 532 nm) particles were predominant. It should again be noted that δ_p and k were negatively correlated (Figs. 1 and 3). Clustered backward trajectory analysis revealed that the majority of the air parcels for that altitude region had passed over the industrial regions of the coastal area in China (30 to 35° N, 105 to 120° E), and two parcels had passed over the Taklimakan Desert regions (Fig. 2b). The NAAPS model predicted the presence of a relatively small amount of dust (15 to 25 μgm^{-3}) and sulphates (1 to 2 μgm^{-3}) below 4 km over the lidar site. These results suggested that the Asian dust and sulphates were both present in the aerosol layers measured with the lidar.

4 Estimation of particle size distribution

We estimated the particle size distributions of the aerosol layers measured with the lidar by comparison with an aerosol model based on the results of the airborne measurements in the free troposphere near the lidar site (Takamura et al., 1990; Matsuki et al., 2003; Sakai et al., 2003; Murayama et al., 2004). The model assumes that the optically effective aerosols were composed primarily of two components: submicrometer-sized ammonium bisulphates in the fine mode and supermicrometer-sized mineral dust in the coarse mode. In the calculation of the optical properties of these particles, the sulphate particles were assumed to be spherical droplets with a refractive index of $1.46+0i$ (Tang and Munkelwitz, 1994) and a lognormal size distribution with a geometric standard deviation (σ_g) of 1.66 (Sakai et al., 2003). The dust particles were assumed to be triaxial spheroids with a length-to-width ratio of 1.4 and a height-to-width ratio ranging from 0.2 to 1.0, based on the morphological study of the Asian dust by Okada et al. (2001). The refractive index of $1.5+0.001i$ (Aoki et al., 2005; Kalashnikova and Sokolik, 2004; Arimoto et al., 2006) and a lognormal distribution with $\sigma_g=2.0$ (Sokolik and Toon, 1999) were assumed for this particle mode. To compute the aerosol optical properties, we used the Mie code (Bohren and Huffman, 1983) for the fine mode and

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the combined field integration (CFIE) method (Mano, 2000) for the coarse mode. The geometric number mean radii (r_g^N) of the two modes were determined in an effort to minimize the difference between the computed values of k and δ_p and those obtained with the lidar by varying r_g^N as described below.

To interpret the observed aerosol optical properties, we expressed the aerosol backscattering coefficient observed with the lidar (β_{obs}) by

$$\beta_{\text{obs}} = \beta_f + \beta_c, \quad (1)$$

where the subscript f refers to the fine mode particles, and c refers to the coarse mode particles. If we define the fraction of β of the coarse mode particles in total (fine + coarse) particles by

$$f_c = \frac{\beta_c}{\beta_f + \beta_c} = \frac{\beta_c}{\beta_{\text{obs}}}, \quad (2)$$

then the wavelength exponent of the backscattering coefficient observed with the lidar (k_{obs}) is expressed by

$$\left(\frac{\lambda_2}{\lambda_1}\right)^{-k_{\text{obs}}} = (1 - f_c) \left(\frac{\lambda_2}{\lambda_1}\right)^{-k_f} + f_c \left(\frac{\lambda_2}{\lambda_1}\right)^{-k_c}, \quad (3)$$

where $\lambda_1 = 532$ nm and $\lambda_2 = 355, 735, \text{ or } 1064$ nm. The particle depolarization ratio observed with the lidar (δ_{obs}) is expressed by

$$\delta_{\text{obs}} = (1 - f_c)\delta_f + f_c\delta_c. \quad (4)$$

To determine k_f , k_c , and δ_c (δ_f is zero because the fine mode particles were assumed to be spherical droplets), we compared the lidar-derived k and δ_p with those computed for the aerosol optical model. Figure 3 presents the scatter diagram of k as a function of δ_p obtained with the lidar and those computed by the model. We determined k_c , δ_c , and k_f in order to minimize the difference between the observed and modelled values. The lidar-derived values (solid diamonds in Fig. 3) fit to a line of the model for which f_c

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varied between 0 and 1 (thick solid line with open squares and numbers in Fig. 3). The r_g^N varied from 0.02 to 0.4 μm in steps of 0.02 μm (thin line with open circles) for the fine mode and from 0.2 to 2.0 μm in steps of 0.1 μm for the coarse mode (thin line with open triangles).

We found that the lidar-derived values most closely fit the computed values when $r_g^N=0.1 \mu\text{m}$ for the fine mode and $r_g^N=0.3 \mu\text{m}$ for the coarse mode. These radii were consistent for the fine mode and a little smaller than the previously reported values obtained from the airborne optical particle counter measurements (0.05 to 0.1 μm for the fine mode and 0.4 to 0.5 μm or 1 μm for the coarse mode) (Takamura et al., 1990; Sakai et al, 2003; Matsuki et al., 2003), and consistent with those derived from the sky-radiometer measurements (Fig. 5). The optical properties computed for these two modes are given in Table 2. For comparison of the computed aerosol optical properties with those obtained from field experiments, Takamura et al. (1994) obtained S_p of 61 sr at 532 nm by use of the elastic lidar and sunphotometer data over Tsukuba in November 1990, when the Asian dust was rare. This value was close to our computed value for the fine mode. Murayama et al. (2004) obtained $S_p \sim 49$ sr at 355 nm, $S_p \sim 43$ sr at 532 nm, and $\delta_p \sim 20\%$ for the Asian dust layers over Tokyo by use of the Raman and polarization lidar. Sakai et al. (2003) reported $S_p=47 \pm 18$ sr at 532 nm and $\delta_p=20 \pm 7\%$ for the Asian dust layers over Tsukuba by the same technique. These values were close to or a little lower than our computed values for the coarse mode. It is necessary to determine the S_p and δ_p of the Asian dust in a consistent manner such like laboratory experiment using a chamber.

For the estimation of f_c averaged for the aerosol layers, Fig. 4 compares k at the three wavelength pairs (355 to 532 nm, 532 to 735 nm, and 532 to 1064 nm) and δ_p obtained with the lidar (solid symbols) and those calculated from the model (open symbols) for the aerosol layers observed for the two periods. The averaged f_c values were 0.84 for the aerosol layer on 1 April (height range between 1.5 and 9.4 km) and 0.17 on 19 October (height range between 1.5 and 4.7 km).

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In order to determine S_p , we expressed S_p in a similar manner to δ_p in Eq. (4) by

$$S_p = (1 - f_c)S_f + f_c S_c. \quad (5)$$

By substituting S_f and S_c given in Table 2 and f_c obtained from the lidar-derived δ_p profile (by Eq. 4) into Eq. (5), we obtained S_p as a function of height. The obtained values were used to compute the vertical profiles of β_p at the four wavelengths. We used δ_p Eq. (4) instead of k Eq. (5) to estimate f_c because the measurement uncertainty in δ_p is lower than that in k . It was necessary to iterate the computation to derive these values (β_p , δ_p , and S_p) because they are mutually dependent. We found that f_c ranged from 0.81 to 1 on 1 April and from 0.07 to 0.36 on 19 October (Fig. 3). The corresponding S_p values were 55–56, 62, 67–69, and 71–78 sr on 1 April and 50–52, 61, 59–62, and 45–55 sr on 19 October at 355, 532, 735, and 1064 nm. It should be noted that f_c decreased with height above an altitude of 3 km on 1 April (which could be expected from the decrease in δ_p with height in Fig. 1a), suggesting that the fraction of the coarse-mode particles decreased with height. This result was possibly due to the faster gravitational settling of the coarse particles than that of the fine particles.

Based on these results, we estimated the column particle concentration and size distribution. The column total particle number concentrations (N) were estimated from β_p and f_c by

$$N = N_c + N_f, \quad (6)$$

where

$$N_c = \int_0^{z_{\text{top}}} f_c(z) \beta_p(z) dz \bigg/ \frac{d\sigma_c}{d\Omega}, \quad (7)$$

for the coarse mode and

$$N_f = \int_0^{z_{\text{top}}} [1 - f_c(z)] \beta_p(z) dz \bigg/ \frac{d\sigma_f}{d\Omega} \quad (8)$$

for the fine mode, where $\frac{d\sigma_x}{d\Omega}$ is the mean volume backscattering cross section of x -mode particles calculated from the model (Table 2). We used β_ρ at 532 nm because the minimum level of the measurement was the lowest (0.1 km in height) and the measurement uncertainty in the signal was the smallest among the four wavelengths.

We set z_{top} 15 km and used the values at 0.1 km for the height integration between the ground and 0.1 km. The column number concentrations were calculated to be $1.1 \times 10^8 \text{ cm}^{-2}$ ($N_c = 0.2 \times 10^8 \text{ cm}^{-2}$ and $N_f = 0.9 \times 10^8 \text{ cm}^{-2}$) on 1 April and $7.3 \times 10^8 \text{ cm}^{-2}$ ($N_c = 2.4 \times 10^6 \text{ cm}^{-2}$ and $N_f = 7.3 \times 10^8 \text{ cm}^{-2}$) on 19 October.

5 Comparison with sky-radiometer measurement

Figure 5 plots the columnar volume size distributions for the two periods. It indicates the predominance of the coarse mode on 1 April and the fine mode on 19 October. For comparison of the size distribution, Fig. 5 also presents the size distributions derived from the sky-radiometer measurement 180 m east of the lidar at 06:40 JST on 1 April and 07:00 JST on 19 October 2006. The method for retrieving the size distribution from the sky-radiometer measurement is described by Uchiyama et al. (2005a, b) and Kobayashi et al. (2006). The size distributions derived from the sky-radiometer indicated a tri-modal distribution with volume mode radii (r_g^V) of $0.25 \mu\text{m}$, $1.16\text{--}1.69 \mu\text{m}$, and $3.62 \mu\text{m}$. The first and second mode radii were close to those estimated from the lidar data for the fine mode with $r_g^V = 0.2 \mu\text{m}$ (computed using the relation $\ln r_g^V = \ln r_g^N + \ln^2 \sigma_g$) and coarse mode with $r_g^V = 1.3 \mu\text{m}$. The third mode radius could not be captured by the lidar data because we assumed a bimodal distribution. The total volume concentrations (V) were $0.19 \mu\text{m}^3 \mu\text{m}^{-2}$ ($V_f = 0.01$ and $V_c = 0.18 \mu\text{m}^3 \mu\text{m}^{-2}$) on 1 April and $0.12 \mu\text{m}^3 \mu\text{m}^{-2}$ ($V_f = 0.1$ and $V_c = 0.02 \mu\text{m}^3 \mu\text{m}^{-2}$), whereas those obtained from the sky-radiometer were $0.37 \mu\text{m}^3 \mu\text{m}^{-2}$ ($V_f = 0.03$ and $V_c = 0.34 \mu\text{m}^3 \mu\text{m}^{-2}$) on 1 April and $0.14 \mu\text{m}^3 \mu\text{m}^{-2}$ ($V_f = 0.06$ and $V_c = 0.08 \mu\text{m}^3 \mu\text{m}^{-2}$) on 19 October. The lidar-derived V were lower than those derived by the sky-radiometer by 48% on 1 April and

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16% on 19 October. The reason for the difference between the lidar and sky-radiometer data is currently unresolved but is probably related to the uncertainty in the assumptions made for calculating the aerosol properties from the measurement data and the spatial and temporal variability of the aerosol properties between the two measurements. The optical thicknesses (τ) at 532 nm derived from the lidar data were 0.44 on 1 April and 0.67 on 19 October, whereas those obtained from the sky-radiometer data were 0.50 on 1 April and 0.52 on 19 October, calculated by interpolating τ at 500 and 650 nm. The lidar-derived τ were 12% lower than that obtained with the sky-radiometer on 1 April and 29% higher on 19 October.

More case studies are necessary to validate the method for estimating the particle size distribution from the lidar data because we used a simple model for calculating the aerosol optical properties. We assumed the two aerosol components of which particle shape, refractive index, and the size distribution shape (σ_g of the lognormal distribution) were constant over the height; only their proportion (f_c) varied with height. Moreover, the measurement data used in this study were taken for only the two periods when the transportation pathways of the air parcels were similar (westerly winds prevailed over the lidar site), so that potential source regions and the predominant aerosol constituents might also be similar. If the predominating aerosol constituents were different from this case study, the parameters (k_c , δ_c , and k_f) of the aerosol model should be modified. In addition, the lidar-derived values must be compared with those obtained with the in-situ aerosol instruments (e.g. particle counters) to validate the results.

6 Conclusions

The vertical profiles of aerosol backscattering coefficients and depolarization ratio were measured by use of multiwavelength and polarization lidar over Tsukuba, Japan, in 2006. The Asian dust layer between altitudes of 3 and 8.5 km on 1 April revealed $k = -0.1$ to 0.5 and $\delta_p \sim 25\%$. The aerosol layer between 1.5 and 5 km on 19 October revealed $k = 1.0$ to 1.6 and $\delta_p = 1$ to 13%. The values of k and δ_p varied with height, and

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they were negatively correlated in those layers, suggesting that the proportion of the coarse nonspherical dust particles to the total particles varied with height. The fraction of the coarse particles in total backscattering was estimated to vary from 0.81 to 1 with an average of 0.84 on 1 April, and from 0.07 to 0.36 with an average of 0.17 on 19 October. The mode radii of these particles were estimated to be $0.3\text{ }\mu\text{m}$ for the coarse mode and $0.1\text{ }\mu\text{m}$ for the fine mode based on k and δ_p values obtained with the lidar and the aerosol model. The column volume concentrations derived from the lidar were $0.19\text{ }\mu\text{m}^3\text{ }\mu\text{m}^{-2}$ on 1 April and $0.12\text{ }\mu\text{m}^3\text{ }\mu\text{m}^{-2}$ on 19 October. These concentrations were 48% lower than those derived from the sky-radiometer on 1 April and 16% lower on 19 October. The optical thicknesses derived from the lidar were 0.44 on 1 April and 0.67 on 19 October, which were lower than that derived from the sky-radiometer by 12% on 1 April and higher by 29% on 19 October. The reason for the differences is not resolved but probably the incorrectness of assumptions of the aerosol model used for calculating the aerosol properties from the measured data and temporal and spatial variability of the aerosol properties between the measurements. Further case studies and comparisons with in-situ measurements are needed to validate the method for estimating the aerosol properties from the lidar measurement.

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Table 1. Specification of lidar used in this study.

Lidar type	Multiwavelength lidar				Polarization lidar
Transmitter					
Wavelength (nm)	355	532	1064	735	532
Laser type (Continuum, Inc.)	Nd:YAG (PL8020)		Nd:YAG (YG581C)	ND:YAG-pumped Dye (TDL60)	Nd:YAG (SL2)
Repetition rate (Hz)	20		20	10	10
Energy/pulse (mJ)	80	40	550	16	300
Beam divergence (mrad)	0.125	0.125	0.125	2.0	0.2
Receiver					
Telescope type	Nasmyth			Schmidt-Cassegrain	Schmidt-Cassegrain
Diameter (m)	0.5			0.35	0.2, 0.4
Field of view (mrad)	3.0			2.2	3.0, 2.0
Detection wavelength (nm)	355	532	1064	735	532
Detection species	Elastic	Elastic	Elastic	Elastic	Elastic and Polarization
Filter bandwidth (nm)	0.50	0.57	0.46	3.00	0.50
Photomultiplier tubes (Hamamatsu Photonics)	R1332	R1332	R3236-01	R1333	R3234-01
Detection mode*	PC and A/D			PC and A/D	PC and A/D
Range resolution (m)	7.5			7.5	7.5
Temporal resolution (min)	1.5			3	3

* PC: Photon counting. A/D: analog-to-digital conversion.

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Table 2. Optical properties computed for the aerosol model.

Mode	Fine				Coarse			
Species	ammonium bisulphate				mineral dust			
Refractive index	1.46+0i				1.5+0.001i			
Size distribution								
geometric mean radius(μm)*	0.1				0.3			
geometric standard deviation	1.66				2			
Shape	sphere				triaxial ellipsoid**			
Wavelength (nm)	355	532	735	1064	355	532	735	1064
k (532 nm- λ)	1.55	–	1.47	1.38	–0.15	–	0.22	0.38
δ_ρ (%)	–	0	–	–	–	27	–	–
S_ρ (sr)	50	61	58	42	56	62	69	78
$d\sigma/d\Omega$ ($10^{-14} \text{ m}^2 \text{ sr}^{-1}$)	1.2	0.5	0.6	0.4	3.6	3.8	3.6	3.0

* Values were estimated from the lidar measurements (see Sect. 3).

** Length-to-height ratio was 1.4, and height-to-width ratio ranged from 0.2 to 1.0, based on microscopic measurement of Asian dust by Okada et al. (2001).

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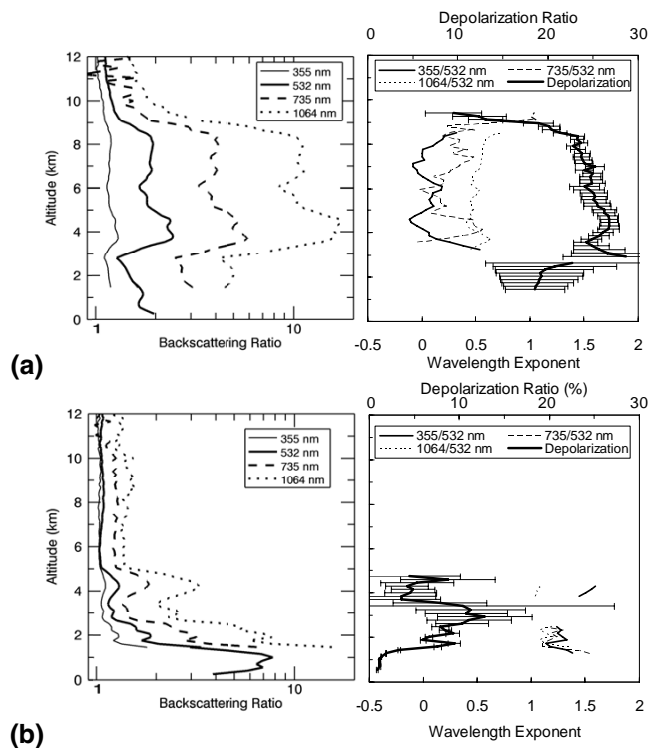


Fig. 1. Vertical distribution of the particle optical properties that were obtained with the lidar for the period 00:01–05:03 JST on 1 April 2006 **(a)** and 02:01–06:45 JST on 19 October 2006 **(b)** over Tsukuba. The left panels show the backscattering ratios at wavelengths of 355, 532, 735, and 1064 nm. The right panels show the wavelength exponent of the backscattering coefficients and the particle depolarization ratio at 532 nm.

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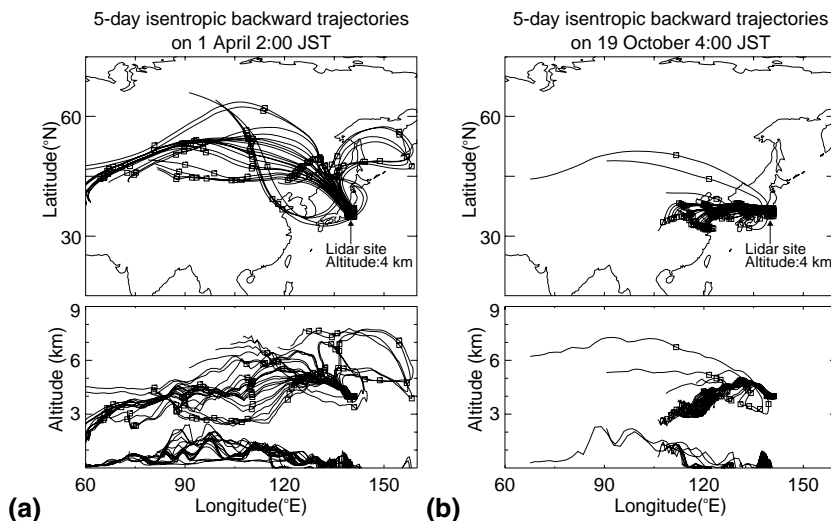


Fig. 2. Clustered isentropic backward trajectories of the air parcels arriving over the lidar site at an altitude of 4 km on 1 April (a) and 19 October (b) in 2006. Symbols are plotted at 24-h intervals. The lowermost lines in the vertical cross section show the land elevations.

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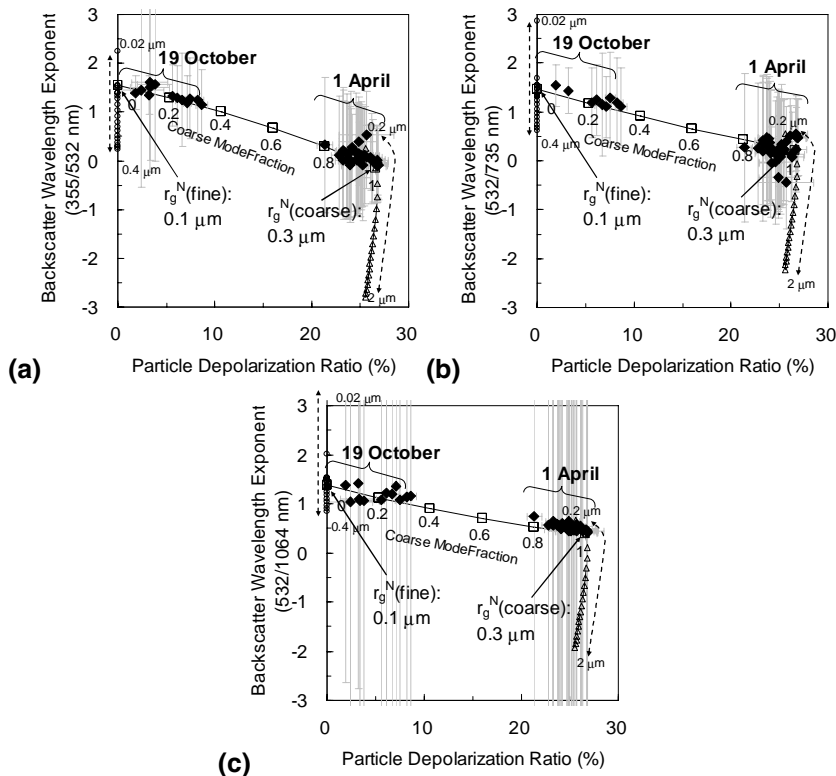


Fig. 3. Scatter diagrams of wavelength exponent of backscattering coefficients as a function of depolarization ratio. The wavelength exponents were between 355 and 532 nm **(a)**, 532 and 735 nm **(b)**, and 532 and 1064 nm **(c)**. Solid diamonds indicate the measured values. Open symbols (circles for the fine mode and triangles for the coarse mode) denote the computed values for which the mode radii varied from 0.02 to 0.4 μ m and 0.2 to 2.0 μ m. Numbers (0 to 1) on a solid line with open squares represent the ratio of coarse to total particles; the mode radii was 0.3 μ m for the coarse mode and 0.1 μ m for the fine mode.

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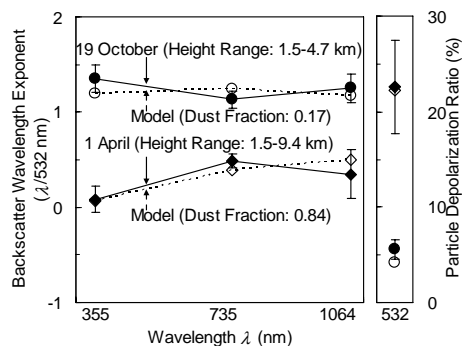


Fig. 4. Comparison of wavelength exponent of backscattering coefficients between 355 and 1064 nm (left panel) and particle depolarization ratio (right panel) between the lidar-derived values and the models. Solid symbols indicate the lidar-derived values averaged for heights of 1.5 to 9.4 km on 1 April (diamonds) and of 1.5 to 4.7 km on 19 October (circles). Open symbols denote the values obtained using the model, assuming the dust fraction of 0.84 (diamonds) and 0.17 (circles) that fit the measurements. The error bar indicates the standard deviation in the height ranges.

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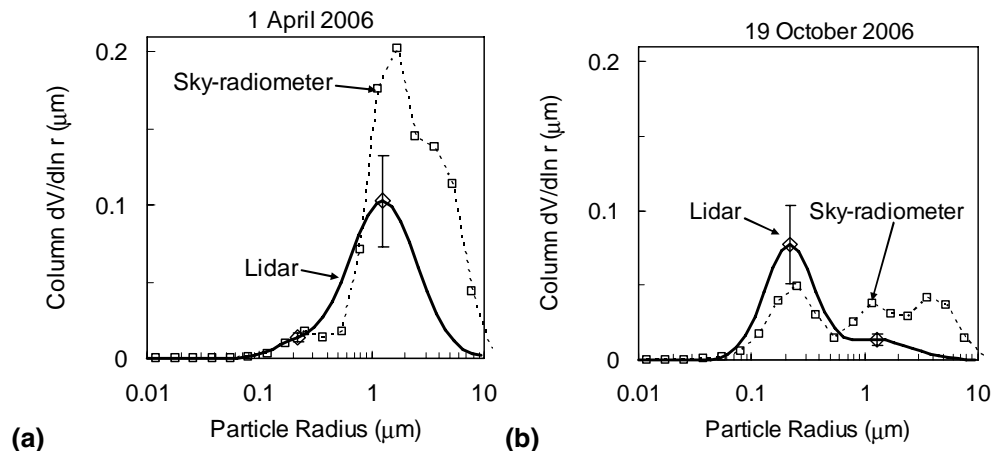


Fig. 5. Column particle volume size distributions estimated from the lidar measurements (solid line with diamonds) and those derived from sky-radiometer measurement (dotted line with squares) over Tsukuba on 1 April (a) and 19 October (b) in 2006. The error bars denote the uncertainties estimated from the lidar signal.

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