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Observations of total alkyl nitrates within the Sacramento Urban Plume

P. A. Cleary^{1,*}, J. G. Murphy¹, P. J. Wooldridge¹, D. A. Day^{1,}, D. B. Millet^{2,***},
M. McKay², A. H. Goldstein^{2,4}, and R. C. Cohen^{1,3,4}**

¹Department of Chemistry, University of California, Berkeley, Berkeley, CA 94720, USA

²Department of Environmental Science, Policy and Management, University of California, Berkeley, Berkeley, CA 94720, USA

³Department of Earth and Planetary Science, University of California, Berkeley, Berkeley, CA 94720, USA

⁴Environment Technologies Division, Lawrence Berkeley National Laboratory, Berkeley, CA 94720, USA

*now at: Department of Chemistry, University of Leeds, Leeds LS2 9JT, UK

**now at: Department of Land, Air, and Water Resources, University of California, Davis, Davis, CA, USA

***now at: Department of Earth and Planetary Sciences, Harvard University, Cambridge, MA 02138, USA

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Correspondence to: R. C. Cohen (cohen@cchem.berkeley.edu)

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Abstract

During the summer of 2001, NO_2 , total peroxy nitrates (ΣPNs), total alkyl nitrates (ΣANs), HNO_3 , volatile organic compounds (VOC), CO_2 , O_3 , and meteorological variables were measured at Granite Bay, CA. The diurnal variation in ΣPNs , ΣANs and HNO_3 were all strongly correlated with sunlight, indicating both that they are photochemically produced and that they have a lifetime of a few hours at this site. The mixing ratios of ΣANs ranged as high as 2 ppbv. Mixing ratios at night averaged 0.4 ppbv. Odd-oxygen ($\text{O}_x = \text{O}_3 + \text{NO}_2$) and ΣANs were strongly correlated reflecting both the common chemical source terms and the similar lifetimes of both species. Several approaches to interpreting the simultaneous variations of O_x and ΣANs are described, and used to derive a best estimate of the ΣAN yield from the VOC mixture at this site of 4.2% and an estimate of the range that is consistent with the observations of 3.9–5.8%. A yield of 4.2% implies termination of the HO_x catalytic cycle by ΣAN formation once every 24 cycles. Analysis of the HNO_3 observations in combination with the ΣAN and O_3 measurements suggests that NO_x terminations limit the HO_x chain length to between 4.7 and 6.3.

1. Introduction

The CATSUP (Chemistry and Transport of the Sacramento Urban Plume)-2001 experiment was designed to provide observations that constrain chemical reactions, transport and source terms affecting atmospheric composition within the Sacramento urban plume. The project combined two field sites: one at Granite Bay, CA, a town on the eastern edge of the Sacramento urban sprawl, and a second site 5 h downwind at the University of California Blodgett Forest Research Station (UC-BFRS) (Fig. 1). The two sites are routinely connected in a Lagrangian sense by the predominant air-flow trajectories, permitting detailed examination of effects of chemistry and transport on changes in composition of the Sacramento urban plume.

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In this manuscript we focus on the initial conditions of this evolving plume. We describe observations of nitrogen oxides at the Granite Bay location and we emphasize analyses of the total alkyl nitrate (Σ AN) measurements which are of interest because:

- 1) models show that production of Σ ANs from these precursors represents a major sink of odd-nitrogen in the eastern US (Liang et al., 1998; Luecken et al., 1999; Poschl et al., 2000; Shepson et al., 1996; Sillman et al., 1998; Trainer et al., 1991) and observations show they are a major component of odd-nitrogen on the western slopes of the Sierra Nevada Mountains (Day et al., 2003);
- 2) they are closely connected to RO_2 , molecules that are central to the HO_x catalytic cycle but that are not well documented by observations;
- 3) formation of Σ ANs can be a major termination of the HO_x and NO_x catalytic cycles; yet the effects of this termination has not been extensively studied, presumably because of the lack of observations to constrain it and
- 4) Σ ANs partially determine the extent to which urban emissions of NO_x are redistributed to larger spatial scales (Bertman et al., 1995).

Recent observations of Σ ANs using the Berkeley TD-LIF instrument have shown that Σ ANs are likely all, or at least a majority of, the ‘missing NO_y ’ (Day et al., 2003) and that they are important controls over the fate of NO_x in urban plumes (Rosen et al., 2004). Mechanistic analyses described in those manuscripts indicate that one of the key diagnostics for understanding the temporal and spatial behavior of Σ ANs both in models and in the atmosphere will be their correlation with O_x ($\text{O}_x = \text{O}_3 + \text{NO}_2$). Rosen et al. (2004) described observations obtained at La Porte, TX, where the observed correlation of O_x with Σ ANs from 09:00–12:00 local time had a slope of 29 implying a Σ AN yield of 6.5% for the VOC mix at the site. The calculated yield for the observed VOC mix was 4.5%. In this paper, we build on this prior work, using observations of meteorological variables, O_3 , nitrogen oxides and VOC to investigate Σ AN, O_3 and nitrogen oxide photochemistry at Granite Bay, CA. The O_3 and Σ AN data are used to

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assess the alkyl nitrate yield from the VOC mixture at Granite Bay, and this correlation is studied with respect to background or nighttime effects. The VOC observations are used to model the expected alkyl nitrate yield and net O₃ and ΣAN production rates. The production rates help to investigate how mixing effects the rise in O₃ and ΣANs over the morning. Finally, we also use a reduced representation of NO_x oxidation by the sum of only HNO₃ and ΣANs to estimate chain length of HO_x and NO_x cycles at this location.

2. Experimental

Observations of NO₂, total peroxy nitrates (ΣPNs), ΣANs, HNO₃, O₃, VOC, and meteorological variables were collected from 19 July–15 September 2001 on the property of Eureka Union School District in Granite Bay, CA (38° N 44.23' 121° W 12.01', 277 m above sea level). This site is located 30 km north-east of Sacramento, CA, at the eastern edge of the suburban sprawl (Fig. 1) and between two major highways: Interstate 80, 8 km to the north and Highway 50, 13 km to the south. The instruments were housed in a temperature-controlled trailer with inlets mounted on a rooftop tower 7 m above the ground (1–2 m above the trailer). The meteorological sensors for wind speed, wind direction, relative humidity and temperature (RM Young model 05103 and Vaisala RH/T sensor model HMP45A) were also located on the tower. Photosynthetically Active Radiation (PAR) was observed using a LICOR LI190SZ Quantum Sensor. O₃ was measured with a Dasibi 1008 ultraviolet photometric ozone analyzer.

Skies were clear on all but two days of the campaign. Wind patterns observed at the site were quite regular, at speeds of 2–2.4 m/s from the southwest (directly from Sacramento) during the afternoon (12:00–17:00) and from the southeast at 1.6 m/s (downslope from the Sierra Nevada) at night (20:00–06:00). On most days, the wind direction rotated smoothly and continuously from southeasterlies to southwesterlies between the hours of 0600–1200h and then rotated back between 17:00 and 20:00. On these days temperature usually was below 16°C in the morning. On days where

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nighttime cooling was not as strong, the change in wind direction was more abrupt with stagnating wind speeds at 05:00–07:00 followed by a rapid rotation from the nighttime to the daytime direction. Back trajectories were calculated using the HYSPLIT model (NOAA-ARL, 1997) at hourly intervals for several days of the campaign. Most of the 8 h back trajectories in the morning were similar with air masses originating southeast of the site. At 8 a.m., most trajectories travel linearly from the southeast of Granite Bay (130°) (Line 1, Fig. 1). Later in the morning, at roughly 10 a.m., the air parcels that were originally following streamlines from the southeast turn and arrive at Granite Bay from the southwest (200°) (Line 2, Fig. 1). By noon, the air parcels observed at Granite Bay were ones that arrived at the Sacramento urban core from the South and then turned toward Granite Bay traveling over the entire length of the Sacramento metropolitan region (Line 3, Fig. 1). On warmer days, the wind turns toward Granite Bay earlier in the day. Temperatures (mean $\pm 1\sigma$) at Granite Bay during the campaign were 16 (± 2.4)°C at 05:00 LT rising to 33 (± 3.6)°C at 16:00 LT. On a few occasions, the trajectories indicated that the air parcels arrived at Granite Bay after traveling across the central valley to the north or south of the Sacramento metropolitan area. There were clear differences in the ozone and nitrogen oxide observations on these days as described below.

2.1. Nitrogen oxides

NO₂ was measured continuously by LIF using a continuous-wave tunable diode laser instrument with sensitivity of 145 pptv/min as described by Cleary et al. (2002). The instrument uses a 15 mW cw diode laser to excite ambient NO₂ at an isolated rovibronic line near 640 nm. The subsequent fluorescence of NO₂ is imaged onto a photomultiplier tube at 90° from the laser path. The fluorescence signal is enhanced by the use of a supersonic expansion, cooling the NO₂ to 25–50 K with a consequent increase of about a factor of 30 in the population of the rotational state excited by the laser.

NO₂, ΣPNs, ΣANs and HNO₃ were measured using a two-channel Thermal-Dissociation Laser Induced Fluorescence (TD-LIF) instrument as described by Day et

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al. (2002). Briefly, an ambient sample flows rapidly through two ovens causing dissociation of NO_z ($\text{NO}_z \equiv \text{HNO}_3 + \Sigma\text{ANs} + \Sigma\text{PNs}$) species to NO_2 . The inlet tip is constructed from PFA Teflon tubing and fittings, with a port for calibration and zeroing followed by a split to the two ovens. The tip is kept as compact as possible, and is thermostatted at $>50^\circ\text{C}$ to ensure high HNO_3 transmission even under high relative humidity conditions. Three different temperatures are used: 180°C , 350°C and 550°C , to observe the three distinct classes of NO_z : ΣPNs , ΣANs and HNO_3 . The NO_2 signal is the sum of the NO_2 contained in all NO_z species that dissociate at or below the oven temperature. The difference between the NO_2 observed in two separate channels at adjacent temperature set points is associated with a specific class of compounds. For these TD-LIF measurements NO_2 was measured using a pulsed dye laser with time gated detection (Thornton et al., 2000). The sensitivity to NO_2 was 15 pptv s^{-1} and the sensitivity for ΣPNs , ΣANs and HNO_3 was typically 45 pptv min^{-1} , 54 pptv min^{-1} , and 70 pptv min^{-1} (Day et al., 2002; Rosen et al., 2004).

Both instruments were calibrated several times daily by dynamically diluting an NO_2 standard (Scott Specialty Gases, 19.2 ppb) in dry NO_x -free air ($\text{NO}_x = \text{NO} + \text{NO}_2$). The accuracy of the standard was established by comparison to a library of 6 other standards before and after the campaign. Standard addition of NO_2 was used as a check that the effects of humidity on NO_2 fluorescence and that small ($<15\%$) effects due to secondary chemistry that occurs in the inlet of the TD-LIF instrument were accurately accounted for in the data analysis (Rosen et al., 2004). The overall accuracy is $\pm 15\%$ for each compound class. About half of the systematic uncertainties in the individual classes of NO_{yi} observed by TD-LIF are correlated, in the sense that if ΣANs are too high, then NO_2 , ΣPNs , and HNO_3 are also too high by the same percentage. Laboratory experiments on aerosol NH_4NO_3 and NaNO_3 show that the HNO_3 measurements are the sum of gas and semi-volatile aerosol nitrate but that TD-LIF is blind to non-volatile salts such as NaNO_3 (Bertram and Cohen, 2003). Thus our measurements of ΣANs are the total ΣANs independent of whether they are gas and aerosol ΣANs and the HNO_3 we report is the sum of gas phase HNO_3 and semivolatile aerosol nitrate.

We did not measure NO or NO_y directly. For the purpose of this paper, we calculate NO using a photostationary state approximation (Eq. 1) and the concurrent NO₂ and O₃ measurements:

$$[\text{NO}]_{ss} = \frac{J[\text{NO}_2]}{k_{\text{O}_3+\text{NO}}[\text{O}_3]} \quad (1)$$

5 The photolysis rate of NO₂ is estimated using calculations with the TUV model (UCAR, 2002). The effect of peroxy radicals on the conversion of NO to NO₂ is neglected. Thus our estimates of NO are likely too large by 10–30%. ΣNO_{yi} is calculated as the sum of NO_{ss}+NO₂ + ΣPNs+ΣANs+HNO₃. As NO is usually a small fraction of NO_y at this site (5–10%), the uncertainty in NO does not affect any of the conclusions about ΣNO_{yi}
10 presented here. Also, O₃ and NO₂ are in a rapid steady-state making O_x (O_x≡NO₂+O₃) a more conserved quantity than O₃. For this reason we consider O_x rather than O₃ in the following analysis.

2.2. VOC Measurements

15 Volatile organic compounds were measured hourly with a fully automated, in situ, two-channel gas chromatograph/mass selective detector/flame ionization detector (GC/MSD/FID) system. This system is described in detail elsewhere (Millet et al., 2004). Briefly, the FID channel was configured for analysis of C₃-C₆ alkanes, alkenes, and alkynes, and the MSD channel for analysis of a range of other VOC, including aromatic, oxygenated and halogenated compounds. For 36 min out of every hour,
20 two subsample flows (15 ml/min) were drawn from the main sample line (4 liters/min) and passed through a preconditioning trap for the removal of water (–25°C cold trap). Carbon dioxide and ozone were scrubbed from the FID channel subsample (Ascarite II), and ozone was removed from the MSD channel subsample (KI impregnated glass wool). Preconcentration was accomplished using a combination of thermoelectric cooling (–15°C) and adsorbent trapping. Samples were injected into the GC by rapidly
25 heating the trap assemblies to 200°C. The instrument was calibrated several times

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daily by dynamic dilution (factor of 1000) of ppmv level standards (Scott Marin Inc., and Apel-Riemer Environmental Inc.) into zero air. Zero air was analyzed daily to check for blank problems and contamination for all measured compounds.

3. Observations

5 Measurements of Σ ANs, Σ PNs, HNO_3 , O_3 , and temperature from 27 July–9 September 2001 are shown in Fig. 2. Over 45 VOC were measured at Granite Bay by GC-MS. These observations will be described in more detail in forthcoming manuscripts. The median mixing ratios for the morning period (09:00–12:00, local time) along with our estimates (denoted by a *, see Appendix A) of the mixing ratios of other compounds likely to contribute to the production of O_x and Σ ANs in the region are listed in Table 1. The average diurnal cycle of the nitrogen oxides and O_x are shown in Fig. 3. On a daily timescale, these chemical species have patterns that closely follow solar illumination indicating rapid chemical production coupled to loss mechanisms (primarily mixing/dilution and deposition) that act on timescales of a few hours.

15 All of the observed species have variations on synoptic timescales that are correlated with large-scale transport patterns and temperature. Peak mixing ratios of O_x are typical of the region (CARB, 1999). 24 h average values of NO_x and NO_y at Granite Bay were 6 ppbv and 9 ppbv respectively. Mixing ratios of NO_x reported by the CARB at several nearby regional air quality monitoring sites have mean daily averages of 10–15 ppbv. The higher values at the CARB sites suggest that their NO_x observations by catalysis-chemiluminescence include a large fraction of NO_z ($\text{NO}_z = \Sigma\text{PNs} + \Sigma\text{ANs} + \text{HNO}_3 + \text{HONO} + \dots$), as has been reported in previous studies (Kolb et al., 2004; Navas et al., 1997). It is also possible that the CARB measurements have a substantial high bias due to being located near major traffic corridors. The Σ PN concentrations measured at Granite Bay exhibit a diurnal maximum of 1.2 ppbv at noon and with a mean of 0.68 ppbv during the entire study. Grosjean (2003) reviews PAN and PPN measurements made in southern California during the past 40 years. The

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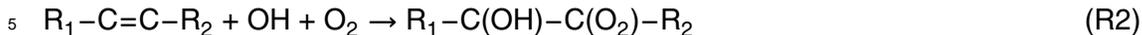
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most recent studies report PAN mixing ratios within 40% of our Σ PN measurements at Granite Bay. The mixing ratios observed at Granite Bay are also similar to those made in other US cities. For example, Roberts et al. (2002) report that the sum of PAN, PPN and MPAN, observed outside of Nashville in June 1999, had an average daytime mixing ratio of 1.14 ppbv and they report the sum of PAN, PPN, MPAN, PiBN and APAN was on average 1.18 ppbv at La Porte, Texas during August–September 2000 (Roberts et al., 2002, 2001). Our measurements of HNO_3 at Granite Bay with daytime peaks near 6 ppbv and daily averages of 2 ppbv are similar to the few prior urban measurements of HNO_3 . Bari et al observed gaseous HNO_3 in Manhattan in abundances ranging from 1–5 ppbv (Bari et al., 2003). Neuman et al. (2002) measured HNO_3 from an aircraft near La Porte, TX in the range of 4–7 ppbv. TD-LIF measurements of HNO_3 on the ground at La Porte, TX peaked at 5 ppbv and averaged 3.6 ppbv (Rosen et al., 2004), demonstrating that the semi-volatile aerosol nitrate component of this measurement in a highly polluted urban environment is likely to be a small addition to the gas-phase HNO_3 .

In contrast to the NO_2 and Σ PN measurements which are similar to prior measurements, the observations of Σ ANs are considerably larger than previous measurements of any individual alkyl nitrate or of the sum of individual alkyl nitrates (Fischer et al., 2000; Grossenbacher et al., 2001; Kastler et al., 2000; O'Brien et al., 1995, 1997; Ostling et al., 2001; Roberts et al., 1998). They are much larger than Σ AN measurements at the UC-Blodgett forest, but similar to TD-LIF observations from La Porte outside of Houston, Texas (Day et al., 2003, 2002; Rosen et al., 2004). The ratios of Σ ANs to NO_y (5–10%) or to NO_z (15–25%) are also larger than nearly all prior observations of the sums of individual nitrates but comparable to our measurements of Σ ANs in other locations (Day et al., 2003, 2002; Rosen et al., 2004).

4. Total alkyl nitrates

The chemical mechanism responsible for producing ozone and alkyl nitrates (Fig. 4) is as follows: ozone production is initiated by oxidation of VOC (Reactions R1–2):



The peroxy radical products of Reactions (R1) and (R2) (both will be denoted RO_2) then reacts with NO to form NO_2 and the alkoxy radical (RO):



10 Typically, the alkoxy radical reacts with O_2 to form a peroxy-radical, HO_2 , and a carbonyl compound:



The HO_2 reacts with NO to form a second NO_2 and regenerates OH :



Ozone is produced by the photolysis of NO_2 , and recombination of $\text{O}(^3\text{P})$ with O_2 :



Reaction (3) also proceeds along an alternate path, an association reaction producing alkyl nitrates (Reaction R3b), and terminating catalytic ozone production:



20 Thus for this typical oxidation path, either one alkyl nitrate is generated or 2 molecules of ozone are generated for each VOC oxidized. The branching ratio to alkyl nitrate

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formation is $\alpha = k_{3b} / (k_{3a} + k_{3b})$. Since the ozone production and total alkyl nitrate (Σ AN) production are both directly dependent on the same rate limiting steps (Reactions R1 or R2), relative production rates of O_3 and Σ ANs are strongly correlated.

The reaction sequence outlined above is initiated by Reactions (R1) or (R2) and proceeds almost quantitatively, thus the instantaneous ozone production rate in an air parcel is given by:

$$P_{O_3} = \sum_i \gamma_i (1 - \alpha_i) k_{OH+VOC_i} [OH][VOC_i] \quad (2)$$

and the Σ AN production rate by:

$$P_{\Sigma ANs} = \sum_i \alpha_i k_{OH+VOC_i} [OH][VOC_i] \quad (3)$$

where γ_i is the number of peroxy radicals produced in the HO_x cycle from VOC_i , k_i is the rate coefficient for the reaction of VOC_i with OH, and α_i is the alkyl nitrate branching ratio for the $R_iO_2 + NO$ reaction. As shown in Fig. 4, γ is typically 2, although it is 1 for direct sources of HO_2 such as the reactions of OH with CO or H_2CO and larger than 2 for VOC whose initial RO_2 fragments produce a second RO_2 prior to HO_2 such as n-butane (Atkinson, 1997). HO_x - HO_x chain termination reactions, such as $HO_2 + HO_2$ or $HO_2 + RO_2$ also reduce γ and Σ AN formation. However, these effects are minor under high NO_x conditions such as were present in this data set, modifying the effective γ for the production rate of Σ ANs for the ambient VOC mixture by less than 10%. We neglect them in this analysis. Table 1 lists the constants needed to evaluate Eqs. (2) and (3). The branching ratios to nitrate formation are estimated by analogy to similar compounds where laboratory measurements are unavailable.

To a first approximation, the composition of Σ ANs in the atmosphere should closely reflect their source distribution in the morning (Rosen et al., 2004). The sources at Granite Bay are detailed in Table 1 and summarized by category in Table 2. Scaling the observed Σ ANs by the relative contribution of the individual sources, the morning (09:00–12:00 LT) Σ AN mixing ratio averaging 700 pptv corresponds to mixing ratios of

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190 pptv of isoprene nitrates, 140 pptv of aldehyde nitrates, 96 pptv of terpene nitrates, 71 pptv of hydroxyalkyl nitrates derived from alkenes, alkynes or aromatics and 72 pptv of multifunctional nitrates derived from isoprene oxidation products. Dinitrates formed by the OH reactions with alkyl nitrates are estimated to be 35 pptv. The sum of C₁-C₉ alkyl nitrates corresponding to the alkane precursors in Table 1 is estimated to be 56 pptv. Only 8% of the nitrates are calculated to be monofunctional species.

Shepson and coworkers have given the most attention to multifunctional nitrates. Direct measurements of isoprene nitrates by Grossenbacher et al during the PROPHET 1998 campaign in northern Michigan ranged from 0.5–35 pptv (Grossenbacher et al., 2001). During SOS 1999 experiments in central Tennessee, Grossenbacher measured 115 pptv of isoprene nitrates at midday. At the PROPHET site, low concentrations of isoprene nitrates were attributed by Grossenbacher et al to the combined effects of the close proximity to isoprene sources and to the large fraction of isoprene peroxy radicals removed by reactions with other peroxy radicals (Grossenbacher et al., 2004). Our estimate of 190 pptv averaged from 09:00–12:00 h, and 270 pptv at noon are larger than the observations by Grossenbacher. However, the difference is likely due in part to the short lifetime of isoprene nitrates before conversion to different nitrate species, which are indistinguishable from the parent nitrates in the ΣAN measurements but would not be identified as an isoprene nitrate in experiments measuring a specific compound. O'Brien et al. (1997) report the sum of 17 different C₃-C₆ alkyl nitrates and C₂-C₄ hydroxy-alkyl nitrates (not including isoprene nitrates), which ranged from 50–550 pptv in the lower Fraser Valley, British Columbia. O'Brien et al. (1995) measured the same 17 species of alkyl and hydroxy-alkyl nitrates and report a sum that ranged from 12–140 pptv at a rural site in Ontario. We scale the measured ΣANs by the relative rates of production from the VOC precursors to estimate that the sum of C₃-C₆ alkyl and C₂-C₄ hydroxy-alkyl nitrates are 6% of ΣANs (~40 pptv at 09:00–12:00 h) at Granite Bay. The other major efforts on multifunctional nitrates are from Ballschmiter and colleagues (Fischer et al., 2002, 2000; Kastler et al., 2000; Luxenhofer et al., 1996; Werner et al., 1999), who measured the sum of C₆-C₁₀ alkyl mononitrates in the range 2.9–11 pptv

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and the sum of C₂-C₄ hydroxy-alkyl nitrates in the range 7.3–28 pptv in urban smog at several sites. At Granite Bay we estimate these same compounds (C₆-C₁₀ alkyl±C₂-C₄ hydroxy-alkyl) nitrates total 30 pptv in the morning.

There are many more studies of the monofunctional nitrates. For example, Flocke et al. (1998) describe measurements of C₁-C₈ alkyl nitrates ranging from 30–620 pptv at a site in southern Germany (Flocke et al., 1998). Scaling our ΣAN observations to the production rate from the VOC precursors gives an estimate of 56 pptv for these species at Granite Bay. Ostling et al. (2001) report summertime C₃-C₅ alkyl nitrate measurements at the PROPHET site in northern Michigan that range from 3–66 pptv and average 18.6 pptv; our scaled observations for these species are 29 pptv at 09:00–12:00 and 25 pptv averaged over the entire campaign.

To summarize, the results presented in this section show that partitioning the observed ΣANs in proportion to their sources results in estimates of the mixing ratio of individual organic nitrates that are in the range of prior observations.

5. Correlations of ΣANs with ozone

The correlation of O_x with ΣANs is shown in Fig. 5. All of the observations are shown as open circles with those between 09:00 and 12:00 h local time shown filled with black and between 14:00 and 18:00 h local time filled with grey. The measured concentrations of O_x and of ΣANs at the surface and their correlation are not only affected by production, but also by deposition, by entrainment of air from the surroundings and by chemical losses. Changes in the ratio of O_x to ΣANs over time are thus represented by Eq. (4):

$$\frac{\Delta O_x}{\Delta \Sigma AN_s} = \frac{\int (P_{O_x} - D_{O_x} + E_{O_x} - L_{O_x}) dt}{\int (P_{\Sigma AN_s} - D_{\Sigma AN_s} + E_{\Sigma AN_s} - L_{\Sigma AN_s}) dt} \quad (4)$$

where *P* represents the chemical production, *D* the deposition, *E* the entrainment and *L* the chemical losses for each species. As illustrated by Rosen et al. (2004), if the

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VOC mix that produces O_x and ΣANs does not vary significantly with time and both deposition and entrainment rates are negligible, the correlation of O_x and ΣANs will be linear (the applicability of these assumptions will be examined below). When this is the case, the slope, $\Delta O_x / \Delta \Sigma ANs$, is then represented by the ratio of Eqs. (2) and (3) evaluated with the VOC mix from a single point in time which is equivalent to Eq. (5):

$$\frac{\Delta O_x}{\Delta \Sigma ANs} \approx \frac{P_{O_x}}{P_{\Sigma ANs}} \approx \frac{2(1-A)}{A} \approx \frac{2}{A} \quad (5)$$

where the constant A is the alkyl nitrate yield for the VOC mixture. Rosen et al. (2004) show that in the absence of mixing, observations early in the daily oxidation sequence should have an $O_x - \Sigma AN$ relationship that closely approximates that of the primary VOC mix in the region. In contrast, if deposition of ΣANs or changes in the VOC mixture that increase the rate of ozone production without increasing the rate of ΣAN production occurs, then the correlation will exhibit significant curvature.

Rosen et al. show that if observations are restricted to early in the day these effects are minimized and a linear correlation will still be observed. However, this analysis by Rosen et al. also made the implicit assumption that the initial values for O_x and ΣANs each day are nearly identical so that increases in the mixing ratios of O_x and ΣANs represent a vector originating at a single point on the correlation diagram, a situation that was true for the data set from La Porte, TX described by Rosen et al. (2004), but is not the case for the Granite Bay data set where we observe a large and highly variable mixing ratio of ΣANs in the early morning. As a result the correlation plot for data at Granite Bay is to some extent skewed by these initial conditions.

At Granite Bay, the slope of the observed correlation of O_x versus ΣANs in the morning (9:00–12:00 local time) is 34.6 (± 0.9 , 1σ) (reduced major axis fitting, $R^2=0.44$) (black symbols Fig. 5), corresponding to an alkyl nitrate yield, A, of 5.8 (± 0.2) %. The slope for data from the afternoon at Granite Bay (14:00–18:00 LT) is 49.2 (± 1.2) (grey symbols Fig. 5) corresponding to a yield of 4.0 (± 0.1)%. The slope derived from observations in the afternoon are higher than in the morning likely due the increasing

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importance of formaldehyde and other secondary organics with high ozone yields and near zero nitrate yields and possibly also to more rapid deposition of hydroxyalkyl nitrates than O_x . If instead of directly fitting the slope of the observed correlation plot we first subtract the median O_x and ΣANs observed in the hour surrounding 08:00 h we arrive at the representation of the data shown in Fig. 6. The figure includes data from 10:30 h to 12:00 h to allow enough time for chemical production to significantly increase the O_x and ΣANs above their 08:00 h value. Since each point in this figure represents $\Delta O_x / \Delta \Sigma ANs$ at different times, the variation in the data represents changes in that slope for different times of day, or for different production rates observed on different days. Figure 6 provides evidence for a range of different slopes, some of which do overlap with the slope inferred from Fig. 5. A fit to the data of Fig. 6 gives a slope of $48.1 (\pm 2.3) \Delta O_x / \Delta \Sigma ANs$, corresponding to a ΣAN yield of $4.2\% (\pm 0.2)$. In our opinion the 4.2% yield inferred from the fit to Fig. 6 is our most reliable estimate of the typical value of the yield at this site although one could make a reasonable argument for slopes as high as 65 using Fig. 5 as low as 29 using the median value of the data shown in Fig. 6. The large values of ΣANs observed at night at Granite Bay strongly influence the slope derived from the raw correlation of daytime measurements of O_x and ΣANs (Fig. 5). This large background obscures the effects of photochemistry on any given day. This is in contrast to observations reported by Rosen et al. (2004) where the nighttime ΣANs were typically less than 0.2 ppb; subtracting an initial value for O_x and ΣANs for each day of that data set has no effect on the correlation slope.

Although the slope of 48 is derived from a fit to the data in Fig. 6, the data shown do include events corresponding to slopes as low as 30 and with ozone increases at the same time as ΣAN decreases. The latter events are most likely dominated by mixing and not by local photochemical production. The variance in the observed correlation is partly due to variation in meteorology as illustrated by Fig. 7, which shows observations from typical and atypical meteorological episodes 29 July 2001 (Fig. 7a), is representative of most of the days of the campaign where the winds shifted to 200° (from Sacramento) by late morning. O_x is 40 ppbv at 08:00 LT, 75 ppbv at noon and

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rises to above 100 ppbv in the late afternoon and Σ ANs are 0.45 ppbv at 08:00 h rising to 1.3 ppbv at noon. The difference between the morning and noontime values gives $\Delta O_x / \Delta \Sigma$ ANs equal to 35. In contrast, on 31 July (Fig. 7b), the winds turn very slowly and were not arriving from the direction of Sacramento (200°) until the late afternoon.

5 O_x peaked at 40 ppbv and Σ ANs did not increase on this day. On 7 September (Fig. 7c), the wind direction swings from the east at 45° to due west at 275° at 06:00 h. Air arriving at Granite Bay directly from the west is not from urban Sacramento, but rather the rural region to the north of the city. The Σ ANs are unusually high (1.2 ppb) even before sunrise that morning, suggesting that production by NO_3 chemistry during the night
10 might have been important. Σ ANs then decrease for most of the day rising again only when the winds change direction late in the afternoon altering the source region for those observations.

In addition to, or perhaps coupled with, meteorological variability, there is photochemical variability in the VOC sources that drive the O_x / Σ ANs correlation. On some
15 days, we observe a high ratio of isoprene oxidation products to the parent compound at 09:00 LT suggesting there has already been significant oxidation of reactive VOC. As a result, we expect high concentrations of secondary photoproducts, such as H_2CO , a compound that contributes to O_x production but not Σ AN production. Variation in the contribution of early morning oxidation is an effect that biases the spread of the
20 O_x / Σ ANs correlation to higher slopes (lower yields) than the slope representing the primary VOC mixture. If this was the main source of bias, then it is likely that the slope derived from Figs. 5 and 6 are not representative of the primary VOC mixture and that the Σ AN yield of that mixture might be higher than the 4.2–5.8% inferred from those figures.

25 6. Calculating the Σ AN yield

Direct calculation of the Σ AN yield for the VOC mixture at the site shows that the estimate of 4.2–5.8% for the Σ AN yield derived from correlations with O_x above is

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reasonable. P_{O_x} and $P_{\Sigma ANs}$ at Granite Bay are calculated for each individual VOC (Table 1) using an assumed mean OH concentration of 5×10^6 molecules cm^{-3} from 09:00–12:00 LT (Cleary, 2003; Dillon et al., 2002a; Dreyfus et al., 2002; Schade and Goldstein, 2001). Note the absolute OH concentration is important to calculating the rate of increase in O_x or ΣANs as we do below, but is unimportant to a calculation of the slope of O_x vs. ΣANs because daytime chemical production of both of these species are proportional to OH. Table 2 summarizes the contributions for each of the major categories of compounds as well as a few individual compounds. The major sources of ozone are isoprene (22.4%), C_2 and larger aldehydes (20.8%), isoprene oxidation products (12.5%) and formaldehyde (12.3%). The major sources of ΣANs are biogenic compounds which are together responsible for 50% of the ΣANs including isoprene (27%) and its oxidation products (10%) and terpenes and their oxidation products (14%). There is also a large contribution from the higher aldehydes (20%), the largest portion of which comes from nitrates derived from nonanal (75% of aldehyde nitrates, 15% of total alkyl nitrates). Using this data and Eqs. (10)–(12), we calculate the slope of $\Delta O_3/\Delta \Sigma ANs$ to be 52 and an alkyl nitrate yield, A, of 3.7%, for the VOC mixture at Granite Bay. This yield is similar to the 4.2% yield that is our best estimate derived from Fig. 6.

Notable uncertainties in the calculated yield are 1) completeness of the observed and estimated VOC mixture and 2) uncertainty about the nitrate yield of individual molecules. With respect to the first item, recent evidence suggests that there may be considerable emissions of sesquiterpenes and other highly reactive species that are not represented in current inventories or in our estimates (Di Carlo et al., 2004; Goldstein et al., 2004). Little is known about the net source of O_x or ΣANs these species represent. With respect to the second item, the isoprene nitrate yield has been reported at values ranging from 4.4 to 15% (Chen et al., 1998; Chuong and Stevens, 2002; Sprengnether et al., 2002; Tuazon and Atkinson, 1990). As isoprene represents 26.7% of the calculated ΣAN source this is a major uncertainty. Our calculation of a 3.7% effective yield for the VOC mixture at Granite Bay used the lowest end of this

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range, because the branching ratio was determined from direct measurements of all isoprene nitrate isomers. If instead we use 12%, a number near the high end of the range, we find that isoprene nitrates represent 50% of the nitrate source and the total yield increases to 5.7%. This is at the high end of what might be compatible with the observations. The uncertainty in the isoprene nitrate yield is by far the largest quantifiable uncertainty in our understanding of Σ AN yields. However, there are other important terms to the overall yield such as the nitrate yields of isoprene oxidation products and the mixing ratio of formaldehyde that are also quite uncertain. The sensitivity of the calculated yield to these terms is much lower than the sensitivity to the isoprene nitrate yield but still significant. For example, doubling (8 ppbv) or halving (2 ppbv) the mixing ratio of formaldehyde results in calculations of the Σ AN yields for the VOC mixture of 3.4% and 4.1%, respectively. Halving or doubling the nitrate yields of isoprene oxidation products puts the calculated yields in the range 3.6% to 4.3%. In combination, an isoprene nitrate yield of 12% and an H_2CO mixing ratio of 2 ppb results in a Σ AN yield of 6.1%. Direct calculation of $\text{P}(\text{O}_x)$ and $\text{P}(\Sigma\text{ANs})$ for the VOC mixture at this site thus gives a Σ AN yield of 3.4–6.1% with a best estimate of 3.7%. This is consistent with the range calculated from the observed $\text{O}_x/\Sigma\text{AN}$ correlation (4.2–5.8%, best estimate of 4.2%).

7. The morning rise of Σ ANs and ozone

Although the analyses above lend support to the interpretation of the rise in Σ ANs as implying a 4.2% yield of Σ ANs during the oxidation of the VOCs at this site, it has the weakness that it neglects entrainment, deposition and chemical losses. In this section we present a more quantitative investigation of the effects of these factors on Σ ANs and O_x by adapting a simple Lagrangian model described by Dillon et al. (2002a) to calculate the morning rise of O_x and Σ ANs.

We use a model that represents production and entrainment to calculate the simultaneous increases in O_x and Σ ANs. The model assumes chemical evolution of a plume

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with O_x and ΣAN production rates (P_{O_x} and $P_{\Sigma ANs}$) initialized at the time, t , as it mixes with a background atmosphere with values of $[O_x]_b$ and $[\Sigma ANs]_b$ as governed by a dilution time constant (K). This model has 5 parameters: K , $P_{\Sigma ANs}$, P_{O_x} , $[O_x]_b$ and $[\Sigma ANs]_b$:

$$\frac{d[O_x]}{dt} = -K ([O_x] - [O_x]_b) + P_{O_x} \quad (6)$$

$$\frac{d[\Sigma ANs]}{dt} = -K ([\Sigma ANs] - [\Sigma ANs]_b) + P_{\Sigma ANs} \quad (7)$$

Both the chemical losses and the deposition described in equation 5 can be neglected. In the morning we estimate a boundary layer height of 500 m near Granite Bay. Assuming a deposition velocity of 0.5 cm-s^{-1} the removal of O_x and ΣANs by deposition is about $3.5\% \cdot \text{h}^{-1}$. An estimate of the chemical losses of O_x is $2\% \cdot \text{h}^{-1}$ and we assume reactions of specific ΣANs largely convert those species to other ΣANs for net zero chemical loss of the sum. These terms are roughly a factor of 10 smaller than the chemical production or entrainment rates.

Dillon et al. (2002a) used a model of this form to derive mixing rates and the OH concentration as they affect VOC during afternoon transport of the Sacramento regional plume from the city into the foothills of the Sierra Nevada. That analysis derived a value for K of 0.23 h^{-1} and an OH concentration of 1.4×10^7 molecules cm^{-3} as an average representing the 5 afternoon hours downwind of the Sacramento suburbs during July 1997. In that region, NO_x concentrations are decreasing from their peak near Granite Bay. We adopt a lower value of OH of 5×10^6 molecules- cm^{-3} to reflect the fact that OH is lower in the morning than at midday and to account for suppression of OH by high NO_x conditions along the trajectories arriving at Granite Bay. The OH we use is also consistent with the observed $\Sigma \text{PNs}/\text{NO}_2$ ratio at Granite Bay (Cleary et al., 2005¹) and

¹Cleary, P. A., Wooldridge, P. J., Millet, D., Goldstein, A. H., and Cohen, R. C.: On the photostationary state of PANs, Aldehydes, NO_x and OH, J. Geophys. Res., in preparation, 2005.

with the HNO_3/NO_2 ratios at Granite bay (J. Murphy, personal communication).

With an average OH concentration of 5×10^6 molecules- cm^{-3} and the VOC listed in Table 1 the calculated production rate is 12.5 ppbv-h^{-1} for O_x , and $0.242 \text{ ppbv-h}^{-1}$ for ΣANs , both numbers that are roughly double the observed rates of increase shown in Fig. 7. Thus we require mixing to slow the rise in the concentrations of O_x and ΣANs . This implies the background we are mixing with has concentrations below those of the plume. Ozone concentrations above the boundary layer are known to be 50–60 ppb (Carroll and Dixon, 1998) thus the primary mixing affecting ozone at Granite Bay must be within the surface layer or a residual nocturnal layer that has experienced significant ozone loss overnight. We set the O_x background to 30 ppbv based on the nighttime observations in the surface layer. Whereas at locations farther downwind from the Sacramento urban plume, the ΣANs tend to have larger nighttime and ‘background’ abundances (as high as 600 ppt) (Day et al., 2004), the events driving ΣAN background and nighttime concentrations at this site are not well known, requiring that we use the observations and model to derive it.

The model is initialized at 08:00 h with O_x of 32 ppbv and ΣANs of 0.58 ppbv and constrained to reproduce the median noontime values of O_x and ΣANs of 54 ppbv and 0.99 ppbv, respectively. We examine three scenarios that illustrate the constraints on our understanding and the remaining uncertainties. In scenario 1, we fix P_{O_x} at 12.5 ppbv h^{-1} , $P_{\Sigma\text{ANs}}$ at $0.242 \text{ ppbv h}^{-1}$ and the background O_x at 30 ppbv. We then solve for K and the background ΣANs and the resulting parameters are: $K=0.44$; and $[\Sigma\text{ANs}]_b=0.53 \text{ ppbv}$. In this scenario the ratio of the production rates is the same as calculated from Table 1 and is within the range indicated by the data. However, this scenario requires a mixing rate almost double that inferred for the afternoon by Dillon et al. (2002b) of 0.23 and the background ΣAN mixing ratio derived is larger than observed during a typical night.

For scenario 2, we fix K at the 0.23 h^{-1} value derived by Dillon et al. (2002), $P_{\Sigma\text{ANs}}$ from Table 1 at $0.242 \text{ ppbv h}^{-1}$ and the background of O_x at 30 ppbv, and derive the production rate of O_x as 9.0 ppbv h^{-1} and the background ΣANs as 0.220 ppbv. The

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ratio of the production rates in this scenario is 37.2 which corresponds to a 5.3% Σ AN yield. The background Σ ANs derived in this scenario is about half the median nighttime value of 0.4 ppbv. The relative production rates of O_x and Σ ANs derived for this scenario could be achieved by a combination of increases in the nitrate yields for isoprene and its oxidation products and decreases in H_2CO . The absolute rates could be achieved by reducing OH by 28%, a change that is within the uncertainty of our estimates of OH although at the high end of that uncertainty. For a third scenario, we fix the background O_x and Σ ANs at their 08:00 h values of 32 ppbv and 0.58 ppbv and fix the dilution rate at 0.23 h^{-1} . Here we derive P_{O_x} as $8.5\text{ ppbv}\cdot\text{h}^{-1}$ and $P_{\Sigma ANs}$ as $0.16\text{ ppbv}\cdot\text{h}^{-1}$. These production rates correspond to a slope of 53 and a Σ AN yield of 3.8% values consistent with the observations and require reducing OH to 66% of our estimate.

All three of these scenarios reproduce the observed rise in O_x and Σ ANs. Taken together, they suggest a range of parameters that are consistent with the observed time evolution of O_x and Σ ANs at Granite Bay are: the mixing term K in the range $0.23\text{--}0.43\text{ h}^{-1}$, OH in the range $3.4\text{--}5\times 10^6$, P_{O_x} in the range $8.5\text{--}12.5\text{ ppbv}\cdot\text{h}^{-1}$, $P_{\Sigma ANs}$ in the range $0.16\text{--}0.242\text{ ppbv}\cdot\text{h}^{-1}$, an O_x background near 30 ppbv and a Σ ANs background in the range $0.27\text{--}0.55\text{ ppbv}$, and a corresponding AN yield in the range of 3.9–5.3.

8. HO_x and NO_x chain termination

One of the major reasons that Σ AN yields are of interest is that Σ AN formation is a termination step for both the HO_x and NO_x catalytic cycles. At low NO_x where OH concentrations are low, the formation rate of HNO_3 slows and under conditions with high VOC reactivity, Σ AN formation is the main termination step for the NO_x catalytic cycle. Similarly, Σ AN formation remains an important contributor to HO_x chain termination at low NO_x because mixing ratios of RO_2 vary much more slowly than OH at low NO_x mixing ratios, even when HO_x self reactions become the dominant chain termination step (e.g. Thornton et al., 2002).

A variety of methods have been used to combine calculations and observations to

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infer free radical chain lengths for the HO_x and NO_x cycles. One common method (Sillman, 1995, 1999) and references therein) is to examine the slope of O_x vs. NO_z, where NO_z is usually defined as NO_y-NO_x and is approximately equal to the sum of HNO₃, ΣANs and ΣPNs. The slope is used to indicate the relative rates of O₃ production versus chain termination, in the form of NO_x oxidation to NO_z. Typical slopes of O_x vs. NO_z for urban plumes are in the range 5–15. During summer, when ΣPN lifetimes are short, this definition overestimates the rate of chain termination. Here we use O_x vs. HNO₃ + ΣANs as our indicator of the relative rates of ozone production and chain termination. The slope of O_x vs. HNO₃ + ΣANs at Granite Bay is 12.5 in the morning (09:00–12:00 h) when we expect deposition to have the smallest effects on our interpretation of the correlation, and 16 for the full dataset. Neglecting deposition or dilution, these slopes are equal to the NO_x chain length and are twice the HO_x chain length. Typically ΣANs are between 1/4 and 1/5 of the sum of ΣANs and HNO₃ suggesting that ΣANs represent 20–25% of the chain termination events.

A second approach to determining the NO_x and HO_x chain lengths is to calculate them directly from the observed hydrocarbons and speciated NO_y measurements. Using the VOC data from Table 1 and the median NO₂ observed during the corresponding 09:00–12:00 h time period of 5.62 ppb, we find the loss rate of OH to VOC is 8.21 s⁻¹, the loss rate to NO₂ is 1.42 s⁻¹. Adding in the additional termination step to ΣAN formation corresponding to 4.2% of the loss to VOC (i.e. a termination rate to ΣAN formation of 0.345 s⁻¹) results in a calculated HO_x chain length of 4.65 and thus a NO_x chain length of 9.30. ΣAN formation is calculated to be 20% of the chain termination events. These numbers are shorter than derived from the correlation of O_x vs. HNO₃ + ΣANs, as we expect based on the assumption that HNO₃ and hydroxyalkyl-nitrates deposit faster than does O_x.

9. Conclusions

We have described observations of sources and products of the major catalytic cycles governing urban and regional photochemistry: VOC, NO₂, ΣANs, HNO₃ and O₃. Our observations show that ΣANs are a photoproduct of oxidation of urban hydrocarbon mixtures with approximately one alkyl nitrate produced for every 25 molecules of O₃. The modeled alkyl nitrate yield of 3.7% is within the uncertainty of the observed yield (best estimate 4.2% and an estimated range of 4–6%) analysis of O_x and ΣANs. These measurements emphasize the importance of ΣAN formation as a chain termination step and suggest that a more complete and quantitative description of ΣAN formation than is present in current reduced mechanisms (Luecken et al., 1999) will be required for accurate models of urban and regional O₃.

Appendix A

In order to account for well-known sources of ozone, we estimate abundances of several VOC. We do not have measurements of the lightest hydrocarbons: CO, CH₄, ethane, ethene, acetylene or formaldehyde. We use the ratio of CO/NO_y of 20 ppbv/ppbv observed in Sacramento in the morning (CARB, 1999) to estimate CO to be 190 ppbv using the ΣNO_{y*i*} measurements. We estimate CH₄ as 1.7 ppmv assuming it is equal to the global average mixing ratio (Dlugokencky et al., 1998). We use the average hydrocarbon ratios observed in experiments at comparable locations (Lamanna and Goldstein, 1999; Parrish et al., 1998) to estimate mixing ratios of 0.4 and 0.1 ppbv for ethene and acetylene. We estimate 4 ppbv formaldehyde based on observations from Sauer et al. (2003) in Parker, CA. The mixing ratio for ethane is estimated from the ratio of 2:1 ethane to propane observed by Goldstein et al. (1995). Concentrations of propanal, octanal, heptanal and nonanal are estimated using ratios of the reactivity of these compounds to the reactivity of the largest aldehyde we did measure, hexanal. We scale to ratios of reactivities reported in Bowman et al. (2003).

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We also do not have observations of many heavy alkanes, aldehydes or β -pinene. We estimate the mixing ratios of octane and nonane using their ratios to i-pentane and set the ratio to the median of the range described by Parrish, et al. (1998). Observations of the ratio of β -pinene to α -pinene, which we did measure, range from 1–3 (Ciccioli et al., 1999; McKeen et al., 1997; Schade et al., 1999). The value of 3 from McKeen is more likely comparable to our site because of many similarities in VOC ratios, giving an estimate of β -pinene of 0.04 ppbv and its contribution to the production of O_3 of 0.72%. There is limited data available for other members of the terpene family. Based on results from Schade et al. (1999) and other observations made at Blodgett Forest (A. Lee, personal communication), the sum of all terpenes is likely twice the sum of α - and β -pinene and we estimate them as such. We also estimate the contribution to ozone production from the photolysis of formaldehyde and acetone (Table 2). Other secondary oxidation products, including and analogous to the C5-hydroxyaldehydes produced during isoprene oxidation, are important contributors to P_{O_3} . We derive the mixing ratio of the C5-hydroxyaldehydes using the ratio to MVK from Sprengnether et al. (2002). Little is known about their reactivity. We estimate that it is similar to that of glycoaldehyde (Bacher et al., 2001). The sum of total terpenoid oxidation products is estimated by analogy to be in the same ratio as the sum of isoprene oxidation products to isoprene. We also include estimates of oxygenated species like MBO, glyoxal and hydroxyacetone as similar to those from measurements at Blodgett Forest (Spaulding et al., 2003). Since methyl glyoxal is a product of MVK oxidation, and we have an estimation of 20% loss of MVK to OH, and the branching ratio of 0.3 to methyl glyoxal formation in the oxidation of MVK, we calculate from an average morning concentration of MVK of 0.53 ppbv, a methyl glyoxal concentration of 0.030 ppbv. This abundance is roughly $\frac{1}{4}$ the amount measured at Blodgett Forest. We then scale all other oxygenated species measured by Spaulding et al. (2003) at Blodgett Forest by this same 25% to represent the fewer biogenic emissions around Granite Bay.

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Table 1. Mean of the measured VOC mixing ratios (09:00–12:00 LT) and estimates of other VOC along with rate constants with OH, the number of ozone produced per VOC oxidized, γ , and branching ratios for alkyl nitrate formation, α , are also listed.

compound	mixing ratio (ppbv)	$k_{\text{OH}}^a \times 10^{11}$ molec. $\text{cm}^{-3}\text{s}^{-1}$	γ^a	α	P_{ozone} (ppbv hr^{-1})	$P_{\Sigma\text{ANs}}$ (ppbv hr^{-1})
Alkanes						
CH ₄	1700*	0.0006	2	0.005	0.384	0.00096
CH ₂ Cl ₂	0.077	0.0082 ^l	2	0*	0.0002	0.00000
CHCl ₃	0.017	0.0088 ^l	2	0*	0.0001	0.00000
CH ₃ Cl	0.600	0.0053	2	0*	0.001	0.00000
ethane	2 ^e	0.0257	2	0.006 ^y	0.018	0.00009
propane	1.005	0.115	2	0.036 ^w	0.040	0.00075
n-butane	0.315	0.254	2.85 ^{rf}	0.073 ^x	0.038	0.00105
iso-butane	0.228	0.233	2.85 ^{rf}	0.255 ^y	0.020	0.00244
pentane	0.199	0.394	2.85 ^{rf}	0.105 ^x	0.036	0.00148
iso-pentane	0.541	0.39	2.85 ^{rf}	0.07 ^y	0.095	0.00456
cyclopentane	0.030	0.508	2.85 ^{rf}	0.10*	0.007	0.00028
2,2-dimethylpentane	0.005	0.340	2.85 ^{rf}	0.12*	0.001	0.00004
2-methylpentane	0.252	0.560	2.85 ^{rf}	0.14*	0.062	0.00356
methylcyclopentane	0.053	0.56 ^{fm}	2.85 ^{rf}	0.14*	0.012	0.00067
hexane	0.089	0.561	2.85 ^{rf}	0.141 ^x	0.022	0.00126
heptane	0.088	0.715	2.85 ^{rf}	0.178 ^x	0.026	0.00201
octane	0.022 ^c	0.87	2.85 ^{rf}	0.226 ^x	0.007	0.00076
nonane	0.011 ^c	1.02	2.85 ^{rf}	0.25*	0.004	0.00049
Alkenes/alkynes/aromatics						
ethene	0.4 ^b	0.854	2	0.0086 ^s	0.122	0.00053
C ₂ Cl ₄	0.052	0.0173	2	0.01*	0.0003	0
propene	0.125	2.63	2	0.015 ^s	0.126	0.00089
1-butene	0.014	3.14	2	0.025 ^s	0.016	0.00020
methylpropene	0.035	5.14	2	0.02*	0.063	0.00064
t2-butene	0.003	6.40	2	0.03*	0.008	0.00012
c2-butene	0.010	5.60	2	0.034 ^s	0.019	0.00033
3-methyl-1-butene	0.015	3.18	2	0.03*	0.016	0.00025
1-pentene	0.017	3.14	2	0.05*	0.019	0.00049
t2-pentene	0.006	6.70	2	0.05*	0.013	0.00034
2-methyl-2-butene	0.010	8.69	2	0.05*	0.022	0.00057
c2-pentene	0.003	6.50	2	0.05*	0.006	0.00017

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Table 1. Continued.

compound	mixing ratio (ppbv)	$k_{\text{OH}}^a \times 10^{11}$ molec. $\text{cm}^{-3}\text{s}^{-1}$	γ^a	α	P_{Ozone} (ppbv hr^{-1})	$P_{\Sigma\text{ANs}}$ (ppbv hr^{-1})
Alkenes/alkynes/aromatics						
cyclopentene	0.002	6.70	2	0.05*	0.005	0.00011
benzene	0.116	0.123	2	0.10*	0.005	0.00026
1,2,4-trimethylbenzene	0.070 ^c	0.380	2	0.10*	0.068	0.00380
1,3,5-trimethylbenzene	0.024 ^c	0.253	2	0.10*	0.046	0.00253
ethylbenzene	0.052	0.710	2	0.10*	0.012	0.00067
1,3-butadiene	0.007	6.60	2	0.11 ¹	0.015	0.00092
toluene	0.347	0.596	2	0.10*	0.067	0.00372
o-xylene	0.066	1.37	2	0.10*	0.029	0.00163
m-xylene	0.104	2.36	2	0.10*	0.080	0.00443
p-xylene	0.070	1.43	2	0.10*	0.032	0.00180
acetylene	0.1 ^b	0.09	1.2	0.01*	0.002	0.00002
propyne	0.016	0.592	0.24	0.05*	0.0004	0.00009
Isoprene and oxidation products						
isoprene	0.807	10.1	2	0.044 ^u	2.80	0.06452
methyl vinyl ketone	0.570	1.88	2	0.05*	0.367	0.00965
methacrolein	0.336	3.35	4 ^q	0.05*	0.962	0.01010
3-methylfuran	0.012	9.35 ^k	2	0.10*	0.034	0.00195
hydroxycarbonyls	0.53 ^d	1.1 ^d	2	0.03*	0.204	0.00315
Terpenes						
α -pinene	0.013	5.37	2	0.18 ^z	0.021	0.00233
limonene	0.010	17.1	2	0.20*	0.050	0.00628
3-carene	0.003	8.80	2	0.20*	0.007	0.00082
β -pinene	0.04 th	7.8	2	0.20*	0.090	0.01123
other terpenes	0.04 ^{ti}	7 ^{ti}	2	0.20*	0.081	0.01008
terpene oxidation products	0.07 ^{ti}	1 ^{ti}	2	0.20*	0.020	0.00252
n-aldehydes						
formaldehyde	4 ^f	0.937	1	0	0.675	0.00000
acetaldehyde	1.318	1.58	3 ^q	0*	1.13	0.00000
propanal	0.22 ^g	1.96	3	0.001*	0.232	0.00012
pentanal	0.090	2.99 ⁿ	3	0.01*	0.143	0.00048
hexanal	0.134	3.17 ⁿ	3	0.03*	0.222	0.00229

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Table 1. Continued.

compound	mixing ratio (ppbv)	$k_{\text{OH}}^a \times 10^{11}$ molec. $\text{cm}^{-3}\text{s}^{-1}$	γ^a	α	P_{Ozone} (ppbv hr^{-1})	$P_{\Sigma\text{ANs}}$ (ppbv hr^{-1})
n-aldehydes						
heptanal	0.14 ^g	3.03 ^m	3	0.05 [*]	0.222	0.00390
octanal	0.13 ^g	3.17 ^m	3	0.07 [*]	0.222	0.00558
nonanal	0.23 ^g	3.6 ^g	3	0.20 ^g	0.445	0.03705
Oxygenated VOC						
methyl ethyl ketone	0.245	0.115	2	0.05 [*]	0.010	0.00025
acetone	3.196	0.0219	2	0.03 [*]	0.025	0.00000
methanol	7.142	0.094	2	0 [*]	0.241	0.00000
ethanol	1.783	0.327	2	0.0086 [*]	0.208	0.00090
MTBE	1.435	0.296	2	0.10 [*]	0.137	0.00759
MBO	0.175 ^j	3.9	2	0.03 [*]	0.238	0.00369
glycoaldehyde	0.16 ^j	1.1 ^o	2	0.01 [*]	0.062	0.00031
hydroxyacetone	0.095 ^j	0.3 ^p	2	0.01 [*]	0.010	0.00005
glyoxal	0.006 ^j	1.2 ^p	2	0.01 [*]	0.003	0.00001
methyl glyoxal	0.03 ^j	1.7 ^p	2	0.01 [*]	0.018	0.00009
2-hydroxy- 2-methylpropanal	0.0.053 ^j	2.0 ^p	2	0.01 [*]	0.038	0.00019
Other						
ΣANs	0.7	1.6	2	0.06 [*]	0.379	0.01210
CO	190 [*]	0.015	1	0	0.821	0.00000
Photolysis						
		j^{\ddagger}				
formaldehyde	4	3×10^{-5}	2	0	0.864	0.00000
acetone	3.423	5.3×10^{-7}	5	0	0.011	0.00000
Total					12.5	0.242

^{*} estimated; ^a Atkinson (1994); ^b estimated from VOC ratios from Lammana and Goldstein (1999): [ethene]=0.5×[propane], [acetylene]=[benzene]; ^c estimated from VOC ratios from Parrish et al. (1998): [n-octane]=0.04×[i-pentane], [n-nonane]=0.02×[i-pentane], [1,2,4 trimethylbenzene]=0.2×[benzene], [1,3,5 trimethylbenzene]=0.07×[benzene]; ^d hydroxycarbonyls are estimated from the theoretical production ratio to MVK: [hydroxycarbonyls]=[$(0.41-\alpha)/0.32$]×[MVK], where $\alpha=0.12$, as the lower limit to hydroxycarbonyl yields. The k_{OH} for these compounds is assumed to equal to that of glycoaldehyde: 1.1×10^{-11} molecules $\text{cm}^{-3} \text{s}^{-1}$; ^e estimated from ratio of 2:1 [propane]:[ethane] from Goldstein et al. (1995); ^f estimated from measurements by Sauer et al. (2003); ^g estimated abundances from Hurst Bowman et al. (2003), where the reactivities of heptanal, propanal and octanal are equal to the reactivity to that of hexanal and the reactivity of nonanal is twice that of hexanal. The rate coefficient, k_{OH} , for nonanal is as measured by Hurst Bowman et al. (2003); ^h estimated from McKeen et al. (1997), where [β -pinene]=3×[α -pinene]; ⁱ other terpenes estimated as = β -pinene with OH rate constant $\sim \beta$ -pinene reaction rate, terpene oxidation products are estimated = 60% of total terpenes; **4834**

^j estimated from measurements taken by Spaulding et al. (2003) at Blodgett Forest at $\sim 0.25\times$ the Blodgett Forest abundances for all species; ^k Atkinson et al. (1989); ^l Hsu and Demore (1994); ^m estimated from Kwok and Atkinson (1995); ⁿ Papagni et al. (2000); ^o Bacher et al. (2001); ^p Spaulding et al. (2003) and references therein; ^q PAN and MPAN are assumed to be in steady state with their production from acetaldehyde and methacrolein, thus the peroxyacetyl nitrate formation is considered a temporary reservoir and the total O₃ molecules generated per oxidation of aldehyde includes those formed from the net loss of PAN and MPAN (Cleary, 2003); ^r estimated from the photooxidation mechanism for n-butane from Seinfeld et al. (1998); ^s O'Brien et al. (1998); ^t Sprengnether et al. (2002); ^u Chen et al. (1998); ^v Ranschaert et al. (2000); ^w Atkinson et al. (1982); ^x Arey et al. (2001); ^y Lightfoot et al. (1992) and references therein; ^z Noziere et al. (1999). [†] Photolysis rates are calculated as an average from between 08:00–12:00 h using the TUV model (UCAR, 2002). [‡] in s⁻¹.

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Table 2. Summary of ozone and Σ AN production by major categories of VOC.

Compound	ppb	P_{ozone} (ppbv hr ⁻¹)	% P_{ozone}	$P_{\Sigma ANs}$ (ppbv hr ⁻¹)	% $P_{\Sigma ANs}$
Alkanes	5.47	0.392	3.1	0.020	8.0
Alkenes/alkynes/aromatics	1.63	0.779	6.2	0.024	10.1
Isoprene	0.886	2.80	22.4	0.065	26.7
Isoprene oxidation products	1.39	1.57	12.5	0.025	10.3
Terpenes + oxidation products	0.071	0.269	2.1	0.033	13.5
Aldehydes	1.56	2.61	20.8	0.049	20.4
Other OVOC	13.7	0.990	8.0	0.013	5.4
Formaldehyde	4.0	1.54	12.3	0.00	0
CO + CH ₄	1900	1.20	9.6	0.00	0
Organic Nitrates	0.7	0.379	3.0	0.012	5.0
Total		12.5	100	0.242	100

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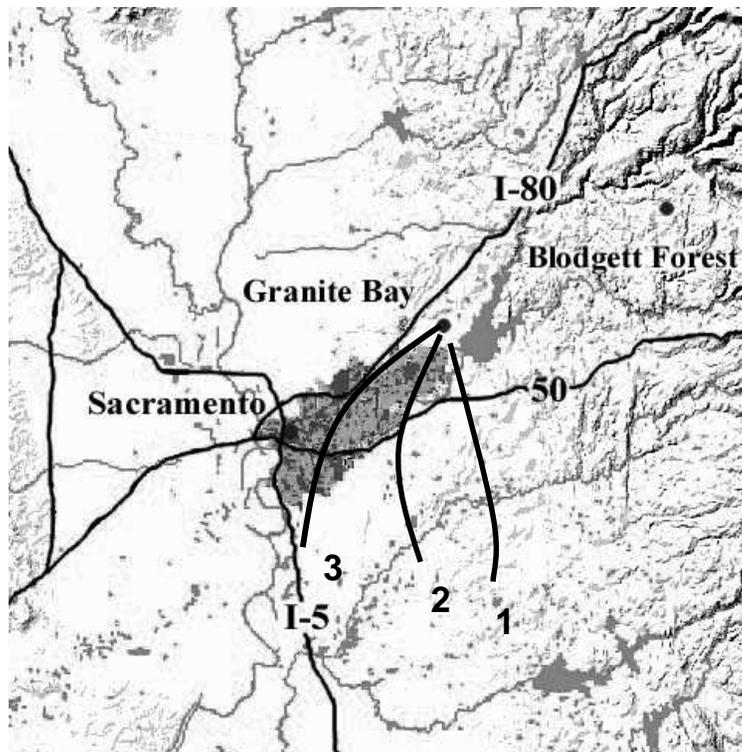


Fig. 1. Topographical map of Sacramento and surrounding areas. Three locations, downtown Sacramento, Granite Bay and UC-BFRS are indicated with black dots. The Granite Bay research site is located 20 miles from the Sacramento urban core, on the outskirts of the Sacramento suburbs. The Blodgett forest research station is located 20 miles to the NE of the Granite Bay site. The major highways, routes 5, 50 and 80, rivers and reservoirs are shown. The gray region around the Sacramento metropolitan area (total population = 1.6 million) marks the region of population density greater than 1900 people km^{-2} . The trajectory lines represent the 4 h back trajectories for air parcels arriving at the site at 08:00 LT (1), 10:00 LT (2) and 12:00 LT (3).

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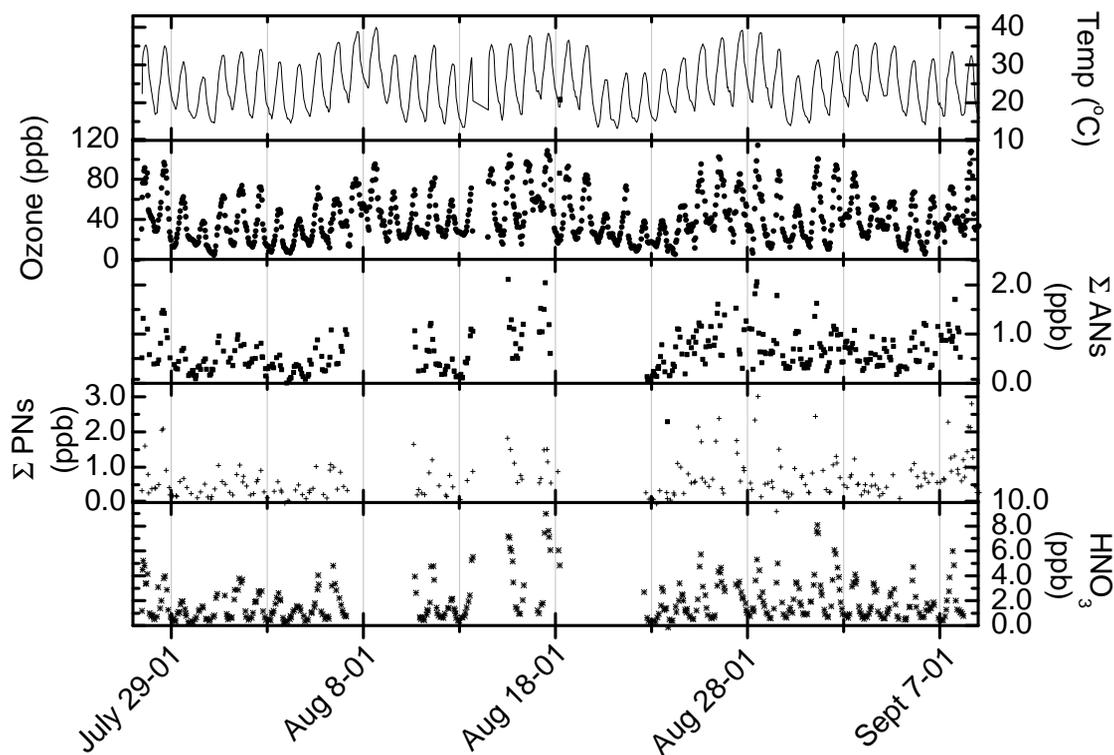


Fig. 2. Measurements (60 min average) of temperature (-), ozone (•), total alkyl nitrates (■), total peroxy nitrates (+), and nitric acid (*) from 27 July to 9 September 2001.

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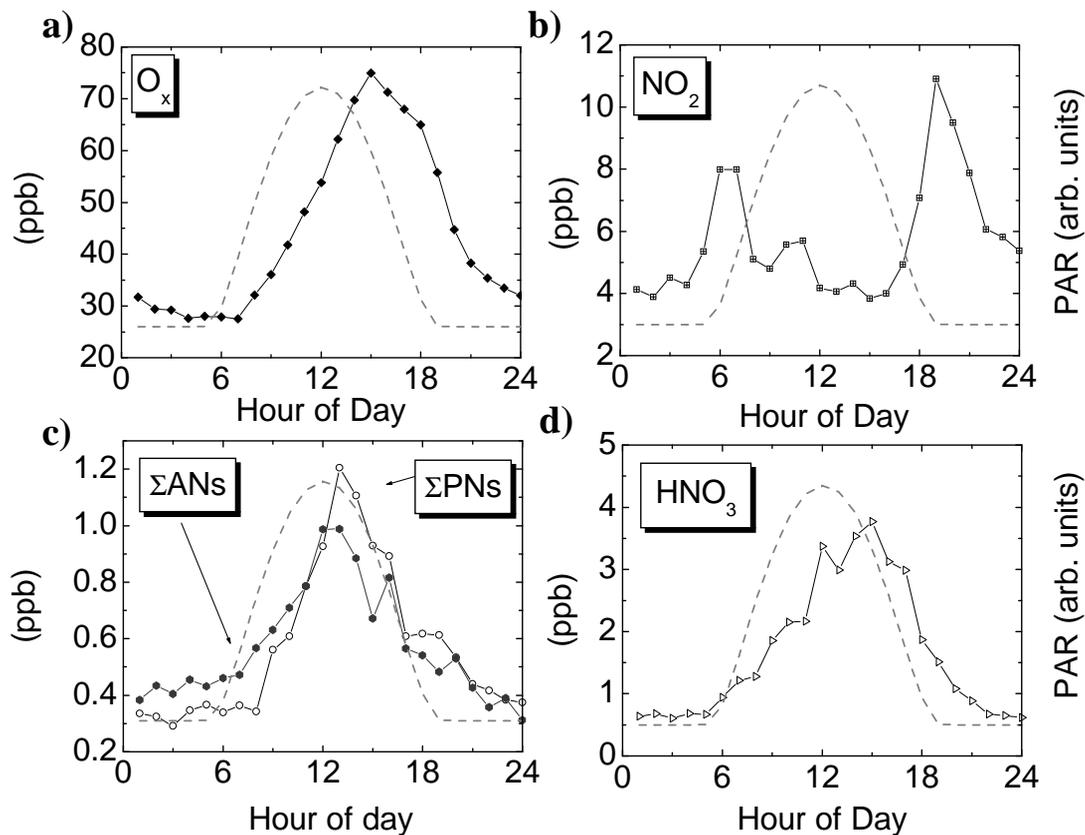


Fig. 3. The median diurnal measurements of (a) O_x (O_3+NO_2), (b) NO_2 , (c) ΣPN s (••) and ΣAN s (—), (d) HNO_3 with PAR representing solar flux (---).

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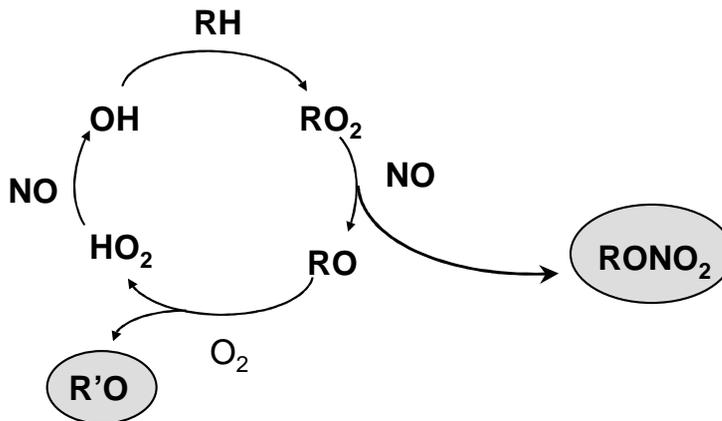
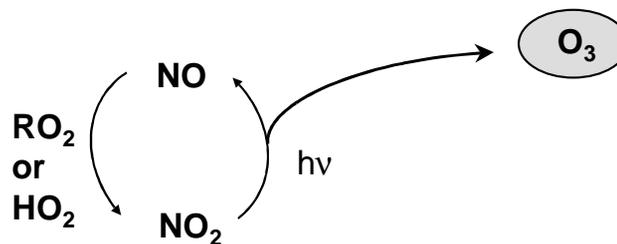
HO_x Cycle**NO_x Cycle**

Fig. 4. A schematic diagram of tropospheric HO_x and NO_x cycles. Ozone, alkyl nitrates, and carbonyl compounds are products of the coupling of both cycles. Typically, as shown in this figure, two molecules of ozone are produced per VOC oxidized.

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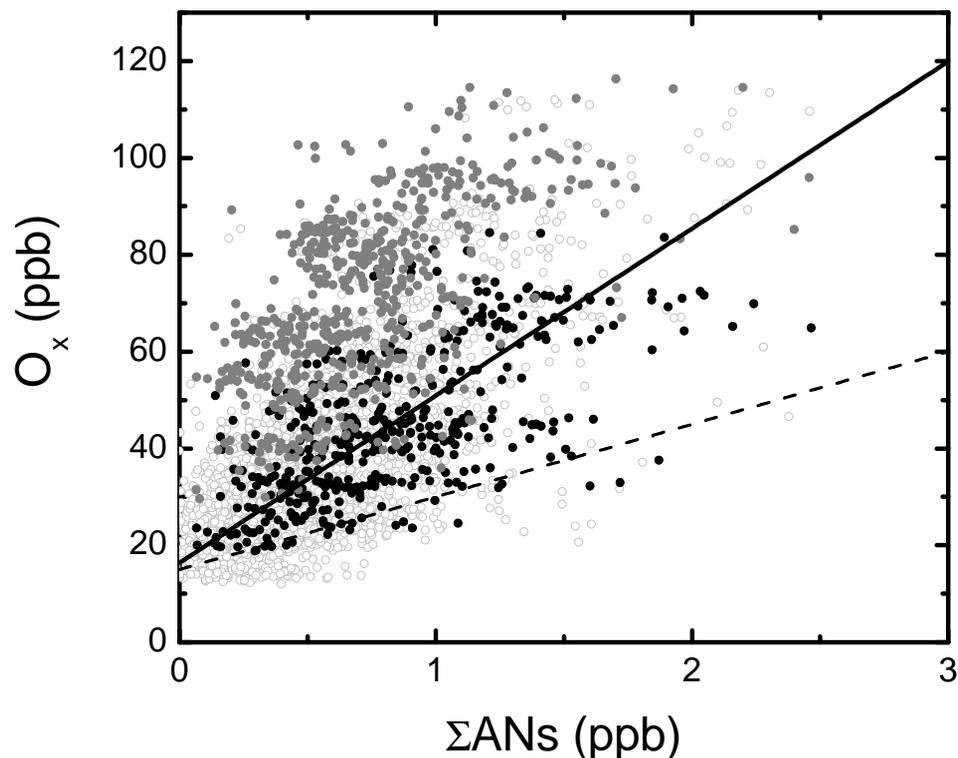


Fig. 5. Observations of O_x versus ΣANs . All points (\circ), 09:00–12:00 (solid black, \bullet), 14:00–18:00 (grey, \bullet). The solid line represents the linear regression fits to the 09:00–12:00 data ($y=34.6x+16.4$). The dashed line represents a minimum slope for the 09:00–12:00 data ($y=15x+15$).

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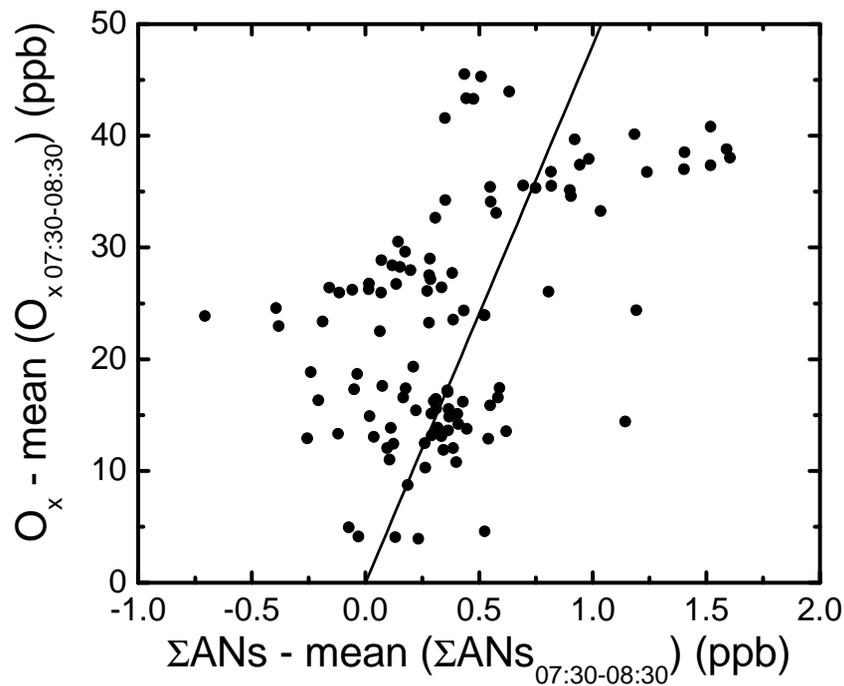


Fig. 6. Observations of O_x versus the increase in ΣANs over the 08:00 h mixing ratio. The observations from 10:30–12:00 h are plotted.

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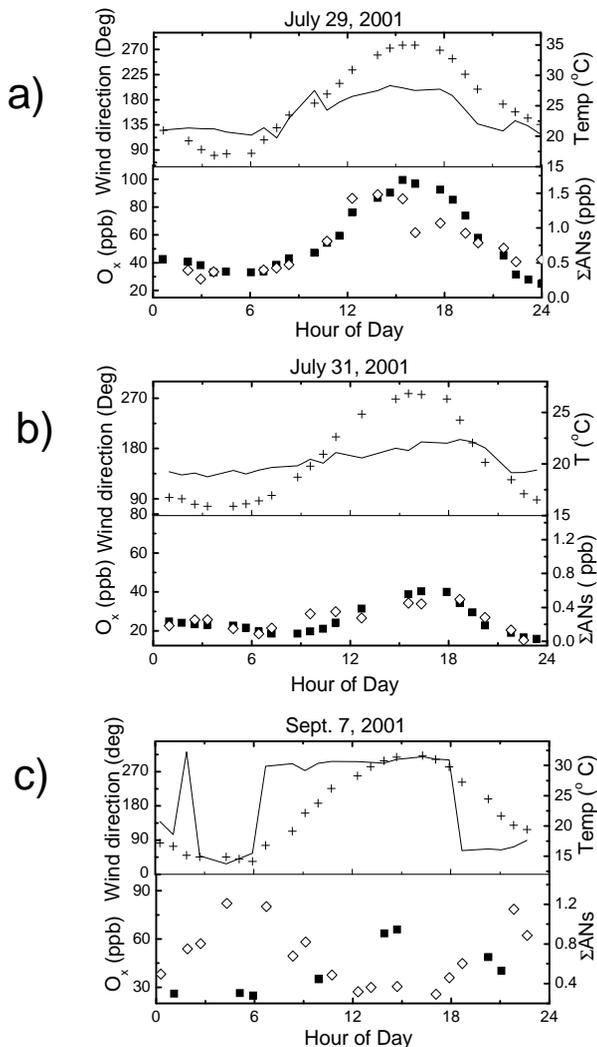


Fig. 7. Three example days of the rise in ΣANs (\diamond) and O_x (\blacksquare) along with wind direction (-) and temperature (+): **(a)** 29 July 2001 where the winds change smoothly from southeasterlies (100°) to southwesterlies (200°), a high temperature of 35°C and maximum O_x and ΣANs of 100 and 1.5 ppbv, respectively; **(b)** 31 July 2001 where the winds shift slowly over to southwesterlies later in the evening with a high temperature of 27°C , and maximum O_x and ΣANs of 40 and 0.5 ppbv, respectively; and **(c)** 7 September 2001 where the winds change abruptly in the early morning from north easterlies (45°) to westerlies (270°) with a high temperature of 31°C and maximum O_x and ΣANs of 65 and 1.2 ppbv, respectively. The ΣANs peak in the early morning on this day and are at a minimum at noon at 0.30 ppbv.

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