

# Shipborne solar absorption measurements of CO2, CH4, N2O and CO and comparison with SCIAMACHY WFM-DOAS retrievals

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# Shipborne solar absorption measurements of CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O and CO and comparison with SCIAMACHY WFM-DOAS retrievals

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### Abstract

CO, CH<sub>4</sub>, N<sub>2</sub>O and CO<sub>2</sub> were retrieved from high resolution solar absorption spectra obtained during a ship cruise from Capetown to Bremerhaven in January/February 2003 by Fourier Transform Infrared (FTIR) spectroscopy. Precisions of better than 0.5% for the column averaged volume mixing ratios (VMR) of  $CH_4$  and  $CO_2$  are achieved us-5 ing of O<sub>2</sub> as a reference gas. Shipborne FTIR-measurements of CO and data from SCIAMACHY/ENVISAT retrieved by the Weighting Function Modified Differential Optical Absorption Spectroscopy (WFM-DOAS) retrieval algorithm show qualitatively the same latitudinal variations. WFM-DOAS data of CH<sub>4</sub>, N<sub>2</sub>O and CO<sub>2</sub> measured over sea exhibit a great spread. The spread is significantly reduced for satellite measure-10 ments over land and a reasonable agreement can be obtained if the shipborne data is compared with the closest SCIAMACHY measurements over land. The number of comparisons is too small to draw conclusions. However, by including only WFM-DOAS data with small errors the shipborne and WFM-DOAS data compare within 5% for  $CH_{4}$ 

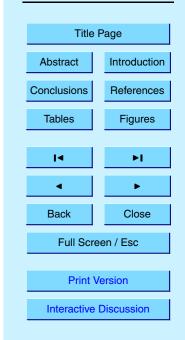
and  $CO_2$  and within 30% for  $N_2O$ . 15

#### Introduction 1.

Carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>) and nitrous oxide (N<sub>2</sub>O) are the most important human-influenced greenhouse gases. From pre-industrial times until 1998 atmospheric CO<sub>2</sub> has increased by approximately 90 ppmv to 367 ppmv, atmospheric  $CH_4$  has increased by a factor 2.5 to 1745 ppbv and N<sub>2</sub>O has increased by 45 ppbv 20 to 314 ppbv. Together,  $CO_2$ ,  $CH_4$  and  $N_2O$  contribute more than 80% to the radiative forcing from the changes in concentrations of all of the long-lived and globally mixed greenhouse gases. Due to their impact on global warming all three gases are addressed in the Kyoto protocol. While the increase of CO<sub>2</sub> has been mainly attributed to the combustion of fossil fuel and changes of land use, natural as well as anthro-25 pogenic sources are responsible for the increase of CH<sub>4</sub> and N<sub>2</sub>O. For atmospheric

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CO<sub>2</sub> concentrations exchange processes with the biosphere and the oceans are highly important because they result in an increase of atmospheric CO<sub>2</sub> at only about half the rate of CO<sub>2</sub> emissions from human activities (IPCC, 2001). However, the spatial distribution and temporal variability of the carbon sinks is still uncertain. In the case
 of methane the sources are poorly quantified and characterized (IPCC, 2001). Hence better constraints on the sources and sinks of CH<sub>4</sub> and CO<sub>2</sub> are needed to predict future atmospheric CO<sub>2</sub> and CH<sub>4</sub> concentrations that are essential for future climate predictions.

Carbon monoxide (CO) plays an important role in atmospheric chemistry and acts as an indirect greenhouse gas by perturbing the OH-CH4-O3 chemistry (IPCC, 2001). It is closely linked to CH<sub>4</sub>, because CH<sub>4</sub> oxidation is an important source for CO and the reaction with OH radicals is a common sink for CH<sub>4</sub> and CO. CO sources include fossil fuel combustion, oxidation of methane and nonmethane hydrocarbons and biomass burning.

Satellite total column measurements of CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O and CO will provide quantitive information about these gases with a global coverage. These measurements are needed for the prediction of future global warming. In the case of the "Kyoto gases" CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O the measurements can be used to verify national reduction targets of theses gases. One important aspect is that CH<sub>4</sub> and CO<sub>2</sub> total column measurements are likely to become important constraints on the geographic and temporal distribution of CH<sub>4</sub> and CO<sub>2</sub> sources and sinks in the future. Currently no satellite instrument is able to achieve the precisions needed to compete with the existing surface

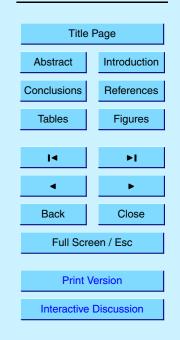
network. Among the instruments that are currently in space SCIAMACHY on ENVISAT has the highest potential to achieve this goal, but the retrieval is just in an experimental phase (Howelling et al., 2003).

Precise measurements from the ground are essential for the validation of satellite data. It is desired to validate satellite measurements over sea and over land because the low surface albedo over sea is a potential source of error in satellite data. FTIR spectroscopy offers the unique ability to measure column averaged volume mixing ra-

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tios of about 20 different trace gases simultaneously and for some gases it is possible to retrieve vertical profiles. Recently it has been shown that  $CH_4$  and  $CO_2$  can be measured from the ground with a precision of 0.5% in the near-infrared using  $O_2$  as a reference gas (Yang et al., 2002; Washenfelder et al., 2003).

<sup>5</sup> In this paper we show that CH<sub>4</sub> and CO<sub>2</sub> can be retrieved from shipborne FTIR measurements with precisions of better than 0.5%, which is essential for the validation of high precision satellite measurements. In addition we compare total columns data for CO<sub>2</sub>, CH<sub>4</sub>, CO, N<sub>2</sub>O retrieved from SCIAMACHY by the WFM-DOAS algorithm (Buchwitz et al., 2000; Buchwitz and Burrows, 2004).

### 10 2. Experimental and data analysis

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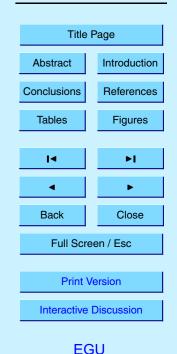
Solar absorption Fourier-transform (FT) measurements aboard the research vessel Polarstern have been performed during a cruise on the Atlantic, starting in Capetown (33.9° S, 18.4° E) on 25 January 2003 and ending in Bremerhaven (53.5° N, 8.6° E) on 17 February 2003. Details about the cruise (ANTXX/3) as well as meteorological data can be obtained via the internet from the Alfred-Wegener-Institute (http://www.awi-bremerhaven.de/MET/Polarstern/GraphInter.html). The experimental setup is described elsewhere (Notholt et al., 2000).

The spectra were analysed using the line-by-line codes GFIT, developed at NASA/JPL (e.g. Toon et al., 1992). Profiles of pressure, temperature, ozone and relative humidity up to 30 km are taken from ozonesondes that were launched daily from the ship. For  $CO_2$  a constant initial vmr-profile of 370 ppmv was assumed. The initial vmr-profiles for all other trace gases are based on balloon observations (Toon et al., 1999). The spectral line parameters were taken from the ATMOS database (Brown et al., 1996).

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#### 3. Shipborne solar absorption measurements

The retrieval of carbon monoxide (CO) and nitrous oxide (N<sub>2</sub>O) from shipborne measurements has been discussed previously (Notholt et al., 2000). Both gases were analysed in the mid-infrared and the resolution of the spectra is  $0.005 \text{ cm}^{-1}$ . The microwindows used for the retrieval and the fitted interfering gases are given in Table1.

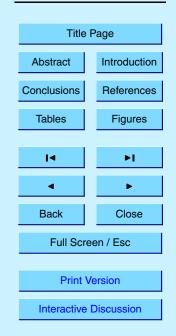
- The CO<sub>2</sub> and CH<sub>4</sub> measurements are converted to column averaged vmrs to achieve a higher precision. The column averaged vmr is also available from WFM-DOAS data. It has been shown that CO<sub>2</sub> and CH<sub>4</sub> can be retrieved from ground based solar absorption spectra with precisions better than 0.5% (Yang et al., 2002; Washenfelder et al., 2003). For the shipborne measurements we follow essentially the ideas of Yang
- <sup>10</sup> al., 2003). For the sinpborne measurements we follow essentially the ideas of Yang et al. (2002) and Washenfelder et al. (2003). The spectra used for the  $CO_2$  and  $CH_4$ retrieval cover the spectral range from 5800 to 8000 cm<sup>-1</sup>. In this spectral range also  $O_2$  can be retrieved. This allows to calculate the column averaged VMR's by scaling the  $CO_2/O_2$  ( $CH_4/O_2$ ) column ratio with the mean  $O_2$  VMR (0.2095). The resolution
- <sup>15</sup> of the near-IR spectra is 0.075 cm<sup>-1</sup> (resolution = 0.9/retardation). It was aimed to average fifty scans for each spectrum to produce high signal to noise ratios. However, during most days weather conditions permitted only averages of 10–20 scans. 77 spectra were obtained on 11 days during the cruise from Capetown to Bremerhaven. Depending on weather conditions, the time difference between the recordings ranges
- from minutes to up to 5 hours. The microwindows used for the retrieval are different from those by Yang et al. (2002) and Washenfelder et al. (2003). For CO<sub>2</sub> two microwindows, each containing one CO<sub>2</sub>-band were used (Table 1). CH<sub>4</sub> was analysed in between 6020 cm<sup>-1</sup> and 6120 cm<sup>-1</sup>. For our spectra this spectral window proved to be advantageous compared with the one used by Washenfelder et al. (2003), presum-25 ably because of the low efficiency of our detector (GaAs-diode) below 6000 cm<sup>-1</sup>.

The diurnal variation of the column ratios  $CO_2/O_2$  and  $CH_4/O_2$  is reduced by up to a factor 5 compared to the diurnal variation of the  $CO_2$  and  $CH_4$  columns. This is expected because many systematic errors are common to  $CO_2$  (CH<sub>4</sub>) and  $O_2$  (e.g.

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surface pressure, zenith angle, instrumental line shape) and will partially cancel in the  $CO_2/O_2$  (CH<sub>4</sub>/O<sub>2</sub>) ratio. Assuming a constant O<sub>2</sub> volume mixing ratio (VMR) in the atmosphere the column averaged volume mixing ratios of  $CO_2$  (CH<sub>4</sub>) can be obtained by scaling the column ratio  $CO_2/O_2$  (CH<sub>4</sub>/O<sub>2</sub>) by the VMR of O<sub>2</sub>. The diurnal variation is smaller than 0.4% for  $CO_2/O_2$  and smaller than 0.75% for the CH<sub>4</sub>/O<sub>2</sub> (Fig. 1).

- <sup>5</sup> tion is smaller than 0.4% for  $CO_2/O_2$  and smaller than 0.75% for the  $CH_4/O_2$  (Fig. 1). However, for the vast majority of the  $CH_4/O_2$  measurements the diurnal variation is smaller than 0.5%. Using the daily variation as a measure of precision it can be concluded that the column averaged VMR of  $CH_4$  can be measured with a precision of better than 0.5% and the column averaged VMR of  $CO_2$  with a precision of better than
- 10 0.4%. Currently no satellite instrument can achieve this precision and it is also sufficient for the validation of proposed future satellite instruments that are dedicated to CO<sub>2</sub> measurements.

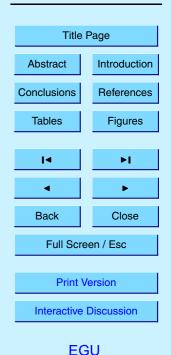
Surface in-situ measurements of CO<sub>2</sub> and CH<sub>4</sub> are routinely performed by the National Oceanic and Atmospheric Administration Climate Monitoring and Diagnostics Laboratory (NOAA/CMDL) at land-based sampling stations. These measurements are highly accurate (0.1 ppmv). Data have been collected approximately during the time of the passage of the ship from several NOAA/CMDL-stations along the route (CCGG, 2003). These stations include Mace Head (Ireland), Terceira Island, Tenerife Island, Ascension Island and Crozet Island. The sampling altitude varies between sea level

- and 2360 m a.s.l. and the stations are in some cases far away from the ships position. However, these stations provide the best available information about the expected latitudinal variation of  $CH_4$  and  $CO_2$  for the time of the cruise. The retrieved shipborne column averaged volume mixing ratios are unreasonable high compared to the surface data. This probably arises from errors in the spectroscopic linelist used for the retrieval
- of the shipborne data. The retrieved shipborne column averaged volume mixing ratio of  $CH_4$  is scaled with a factor 0.9 and the  $CO_2$  with 0.957. This factor was empirically determined by minimizing the bias between the surface and the column data.

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# 4. Comparison of shipborne measurements with WFM-DOAS retrievals from SCIAMACHY near infrared spectra

For the retrievals of CO,  $CH_4$ ,  $CO_2$ ,  $H_2O$ , and  $N_2O$  from SCIAMACHY/ENVISAT nearinfrared spectra the Weighting Function Modified Differential Optical Absorption Spec-

- <sup>5</sup> troscopy (WFM-DOAS) retrieval algorithm has been developed at the University of Bremen (Buchwitz et al., 2000; Buchwitz and Burrows, 2004). WFM-DOAS version 0.4 scientific products recently released (available from http://www.iup.physik.uni-bremen. de/sciamachy/NIR\_NADIR\_WFM\_DOAS/) have been used for the comparison of CO, CH<sub>4</sub>, CO<sub>2</sub>, and N<sub>2</sub>O vertical columns.
- <sup>10</sup> WFM-DOAS data are available for 5 days of the cruise. Only the longitudinally closest WFM-DOAS retrievals within 3°-latitude bands around the ships position have been taken into consideration. The considered WFM-DOAS retrievals are indicated by the blue areas in Fig. 2. Each area consists of a high number of WFM-DOAS retrievals. The blue areas indicate measurements over land and over sea. Land and sea WFM-
- DOAS data have been compared separately with the shipborne data. The WFM-DOAS data also states an error for each datapoint. Only data with an error below a certain threshold was included. A higher threshold results in a larger spread of the data and a smaller threshold results in a lower number of included datapoints. This demonstrates that the error indicated in the WFM-DOAS data is a qualitatively sensible measure of
- <sup>20</sup> the precision. In our case the error threshold of the included data is a trade-off between the spread in the data and the number of days for which data can be included. The error threshold was empirically determined by minimising the spread of the data and ensuring that at least three of the five days are included. The error threshold was determined for each dataset separately and is for CO<sub>2</sub> 7% over sea and 7.4% over land 60% for CO ever and 7% for CH ever and 20% for N O ever land.
- $_{25}$  land, 60% for CO over sea, 7% for CH<sub>4</sub> over sea and 30% for N<sub>2</sub>O over land. For CO and CH<sub>4</sub> over land the spread did not depend on the threshold and all data were included.

For all compared gases the spread in the WFM-DOAS data is larger over sea than

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over land (Fig. 3). This is due to the lower surface albedo over the oceans compared to land resulting in a lower signal to noise ratio in the measurements over the oceans.

Among the investigated gases the best agreement between the shipborne and the WFM-DOAS data is observed for CO. Deviations between the datasets are smaller than

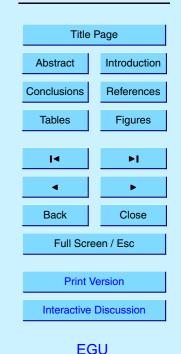
- the observed latitudinal variation (Fig. 3). The most striking feature in both datasets is the CO enhancement between the equator and 10° N. Trajectory calculations reveal biomass burning in West Africa as the apparent source of this enhancement (Fig. 2). WMFDOAS data shows that during the time of the cruise enhanced CO can be observed over large parts of the African continent in the latitude band between the equa-
- tor and 10° N (Fig. 4). This coincides with the biomass burning (Fig. 2) which is typical for that region during the considered time of the year. The deviations between the WFM-DOAS data and the shipborne data between the equator and 10° N result most likely from the different sampling locations. Between 12° S and 15° S the WFM-DOAS CO data shows a larger spread than at other latitudes, which is also valid for the other investigated gases.

The WFM-DOAS data for CO<sub>2</sub> over sea shows a high spread of up to 90 ppmv. This high spread cannot be real considering a maximal seasonal amplitude of the column of less than 15 ppm. Over land CO<sub>2</sub> ranges from 362 ppmv to 368 ppmv at 25° S–28° S and from 375 ppmv to 391 ppmv at 25° N–28° N. The mean of the included data at 25° S–28° S -28° S is 365±2.2 ppmv and 384.6±3.8 ppmv at 25° N–28° N. The deviation between the WFM-DOAS and the shipborne data is less than 4% at these latitudes. The CH<sub>4</sub> WFM-DOAS data over sea exhibits a large spread and the majority of datapoints

- is higher than the shipborne data. Similar to CO the  $CH_4$  data over land shows a larger spread between 12° S and 15° S than at other latitudes. A deviation between satellite
- <sup>25</sup> and shipborne measurement of up to 20% is observed in this latitude band. By lowering the error threshold to 4% the data between 12°S and 15°S is excluded. At the remaining two latitude bands the WFM-DOAS data over land and the shipborne  $CH_4$ agree within 2.5%. In the case of N<sub>2</sub>O the WFM-DOAS data over land is significantly lower than the shipborne data and its spread is higher than the latitudinal variation.

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Due to its high spread the WFM-DOAS data over sea is not shown.

### 5. Conclusions

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High resolution solar absorption spectra were obtained during a ship cruise from Capetown to Bremerhaven in January/February 2003 by FTIR spectroscopy. It has been demonstrated that column averaged volume mixing ratios (VMR) of  $CH_4$  and  $CO_2$  can be retrieved with precisions of better than 0.5% from these spectra by using  $O_2$  as a reference gas.

Recently released SCIAMACHY WFM-DOAS scientific products of CO,  $CH_4$ ,  $N_2O$ and  $CO_2$  are compared with the shipborne data. In the case of CO the agreement between shipborne and satellite data is very good in the sense that the data show qualitatively the same latitudinal variation and that the difference between shipborne and satellite data is smaller than the latitudinal variation. WFM-DOAS data of  $CH_4$ ,  $N_2O$  and  $CO_2$  measured over sea exhibit a large spread. A better agreement can be obtained if the shipborne data is compared with the closest SCIAMACHY measure-

ments over land. A crucial point for the comparison is the definition of an error threshold filter applied on the WFM-DOAS results. At the expense of the number of included datapoints the shipborne and the WFM-DOAS data for CH<sub>4</sub> and CO<sub>2</sub> can be brought into an agreement of 5%. N<sub>2</sub>O shows the strongest deviation between shipborne and satellite data of up to 30%. However, for all gases the number of comparisons is too
 small for statistical analyses in order to achieve further quantitative conclusions about the agreement.

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 developed by NASA's Goddard Space Flight Center and the University of Maryland for the fire data. We are grateful to the NOAA Climate Monitoring and Diagnostics Laboratory for providing free access of surface data. Last but not least we would like to thank the German Ministry for

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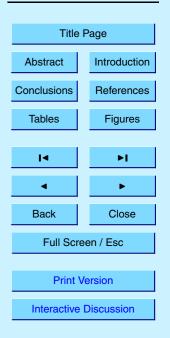
 Table 1. Microwindows and fitted interfering gases used in the retrieval.

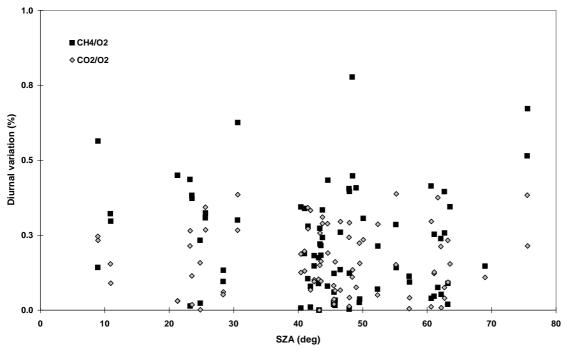
Retrieved gas	Spectral window for fit in cm <sup>-1</sup>	Fitted interfering gases
CO <sub>2</sub>	6175–6265 6297–6382	H <sub>2</sub> O H <sub>2</sub> O
CH <sub>4</sub>	6020–6120	H <sub>2</sub> O, CO <sub>2</sub>
O <sub>2</sub>	7765–8005	0O <sub>2</sub> , H <sub>2</sub> O
CO	2111–2112.1 2157.3–2159.15 2157–2163	H <sub>2</sub> O, O <sub>3</sub> , N <sub>2</sub> O H <sub>2</sub> O, O <sub>3</sub> , N <sub>2</sub> O H <sub>2</sub> O, O <sub>3</sub> , N <sub>2</sub> O
N <sub>2</sub> O	2441.8–2444.4 2481.2–2482.5 2806.05–2806.55	CO <sub>2</sub>

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Shipborne solar absorption measurements of CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O and CO

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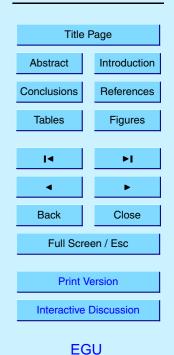


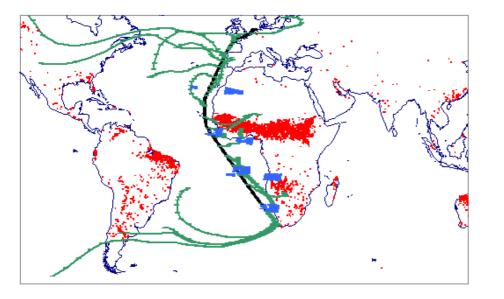
**Fig. 1.** Diurnal variations of the CO<sub>2</sub> and CH<sub>4</sub> ratios. The diurnal variation is defined as  $100 \times abs(\frac{x}{\langle x \rangle} - 1)$ , where x is one measurement,  $\langle x \rangle$  the mean of the day and *abs* stands for absolute value.

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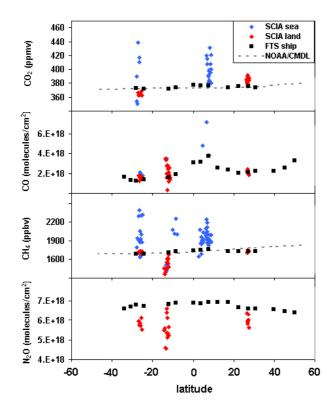
Shipborne solar absorption measurements of  $CO_2$ ,  $CH_4$ ,  $N_2O$  and CO

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**Fig. 2.** Ship track (black), backward trajectories (green) at 850 hPa and closest available WFM-DOAS-retrievals (blue). The red dots indicate fires (from MODIS Rapid Response).

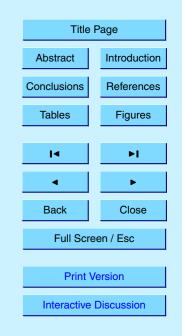


**Fig. 3.** Comparison of WFM-DOAS-retrievals (blue over sea and red over land) with shipborne data (black rectangles). Only cloudfree WFM-DOAS data with an error smaller than a certain value were taken into account. The maximal errors for the data that were included are for  $CO_2$  7% over sea and 7.4% over land, 60% for CO over sea, 7% for CH<sub>4</sub> over sea and 30% for N<sub>2</sub>O over land. For CO and CH<sub>4</sub> over land all data were included. Due to its large spread no N<sub>2</sub>O over sea were included. The shipborne data for  $CO_2$  is scaled with a factor 0.957 and the CH<sub>4</sub> data with a factor 0.9. The dotted grey line represents NOAA/CMDL surface sampling data connects data measured during the cruise at stations closed to the cruise track. Stations included are Mace Head (Ireland), Terceira Island, Tenerife Island, Ascension Is. and Crozet Island.

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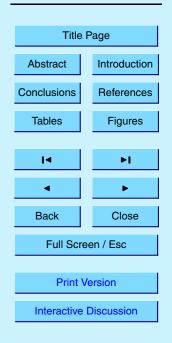
Shipborne solar absorption measurements of CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O and CO

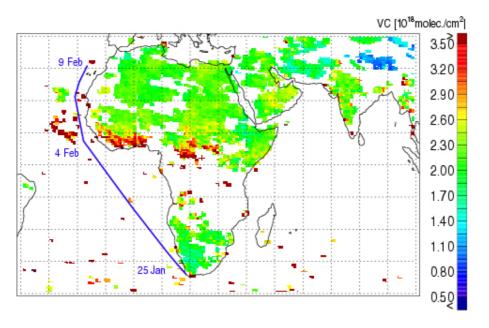


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**Fig. 4.** CO vertical columns from SCIAMACHY WFM-DOAS-retrievals (Version 0.4). The plot includes all available WFM-DOAS data for the duration of the cruise (24, 27, 30, 31 January 2003 and 3, 4, 8 February 2003). Only cloudfree measurements with an error <60% were used.