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Retrieval of ozone column content from airborne Sun photometer measurements during SOLVE II: comparison with coincident satellite and aircraft measurements

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Abstract

During the 2003 SAGE (Stratospheric Aerosol and Gas Experiment) III Ozone Loss and Validation Experiment (SOLVE) II, the fourteen-channel NASA Ames Airborne Tracking Sunphotometer (AATS-14) was mounted on the NASA DC-8 aircraft and measured spectra of total and aerosol optical depth (TOD and AOD) during the sunlit portions of eight science flights. Values of ozone column content above the aircraft have been derived from the AATS-14 measurements by using a linear least squares method that exploits the differential ozone absorption in the seven AATS-14 channels located within the Chappuis band. We compare AATS-14 columnar ozone retrievals with temporally and spatially near-coincident measurements acquired by the SAGE III and the Polar Ozone and Aerosol Measurement (POAM) III satellite sensors during four solar occultation events observed by each satellite. RMS differences are 19 DU (6% of the AATS value) for AATS-SAGE and 10 DU (3% of the AATS value) for AATS-POAM. In these checks of consistency between AATS-14 and SAGE III or POAM III ozone results, the AATS-14 analyses use airmass factors derived from the relative vertical profiles of ozone and aerosol extinction obtained by SAGE III or POAM III.

We also compare AATS-14 ozone retrievals for measurements obtained during three DC-8 flights that included extended horizontal transects with total column ozone data acquired by the Total Ozone Mapping Spectrometer (TOMS) and the Global Ozone Monitoring Experiment (GOME) satellite sensors. To enable these comparisons, the amount of ozone in the column below the aircraft is estimated either by assuming a climatological model or by combining SAGE and/or POAM data with high resolution in-situ ozone measurements acquired by the NASA Langley Research Center chemiluminescent ozone sensor, FASTOZ, during the aircraft vertical profile at the start or end of each flight. Resultant total column ozone values agree with corresponding TOMS and GOME measurements to within 10–15 DU (~3%) for AATS data acquired during two flights – a longitudinal transect from Sweden to Greenland on 21 January, and a latitudinal transect from 47° N to 35° N on 6 February. For the round trip DC-8 latitudinal

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transect between 34° N and 22° N on 19–20 December 2002, resultant AATS-14 ozone retrievals plus below-aircraft ozone estimates yield a latitudinal gradient that is similar in shape to that observed by TOMS and GOME, but resultant AATS values exceed the corresponding satellite values by up to 30 DU at certain latitudes. These differences are unexplained, but they are attributed to spatial and temporal variability that was associated with the dynamics near the subtropical jet but was unresolved by the satellite sensors.

For selected cases, we also compare AATS-14 ozone retrievals with values derived from coincident measurements by the other two DC-8 based solar occultation instruments: the National Center for Atmospheric Research Direct beam Irradiance Airborne Spectrometer (DIAS) and the NASA Langley Research Center Gas and Aerosol Monitoring System (GAMS). AATS and DIAS retrievals agree to within RMS differences of 1% of the AATS values for the 21 January and 19–20 December flights, and 2.3% for the 6 February flight. Corresponding AATS-GAMS RMS differences are ~3% for the 21 January flight; GAMS data were not compared for the 6 February flight and were not available for the 19–20 December flight. Line of sight ozone retrievals from coincident measurements obtained by the three DC-8 solar occultation instruments during the SAGE III solar occultation event on 24 January yield RMS differences of 2.1% for AATS-DIAS and 4.2% for AATS-GAMS.

1. Introduction

The second SAGE III Ozone Loss and Validation Experiment (SOLVE II) was an international measurement campaign conducted during Arctic winter 2003 to acquire correlative data for validation of Stratospheric Aerosol and Gas Experiment (SAGE) III measurements, and to study polar winter ozone loss rates and related photochemical processes, polar stratospheric clouds, and polar winter transport and dynamics processes. During SOLVE II, the NASA DC-8 research aircraft deployed from NASA Dryden Flight Research Center (DFRC) at Edwards Air Force Base (AFB), California

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to Kiruna, Sweden on 9 January and flew 12 science missions out of Kiruna before its return to DFRC on 6 February. The DC-8 instrument payload consisted of a variety of in-situ and active and passive remote sensors, including three solar occultation instruments: the fourteen-channel NASA Ames Airborne Tracking Sun photometer (AATS-14), the NASA Langley Gas and Aerosol Monitoring System (GAMS)-Langley Airborne A-Band Spectrometer (LAABS) (Pitts et al., to be submitted, 2005¹) and the National Center for Atmospheric Research (NCAR) Direct beam Irradiance Airborne Spectrometer (DIAS) (Swartz et al., 2004; Shetter and Muller, 1999).

AATS-14 obtained measurements during the sunlit segments of 12 science flights, including two flights based out of DFRC prior to the deployment to Kiruna and the return transit flight from Kiruna to DFRC. In this paper, data from eight of these flights have been analyzed in detail to yield spectra of total and aerosol optical depth (TOD and AOD), and to retrieve values of ozone and water vapor column content above the DC-8 flight altitude. In a companion SOLVE II paper (Russell et al., 2004), AATS-14 slant transmission and AOD retrievals are compared with analogous SAGE III, POAM III, and DIAS measurements. In this paper, we compare AATS-14 ozone retrieval results with coincident or near-coincident measurements obtained with the SAGE III, POAM (Polar Ozone and Aerosol Measurement) III, TOMS (Total Ozone Mapping Spectrometer), and GOME (Global Ozone Monitoring Experiment) satellite sensors, and with the DC-8 based GAMS and DIAS instruments. To our knowledge, this is the first published report of a sub-orbital airborne Sun photometer having been used to retrieve columnar ozone.

¹Pitts, M. C., Thomason, L. W., et al.: Ozone observations by the Gas and Aerosol Measurement Sensor during SOLVE II, Atmos. Chem. Phys. Discuss., to be submitted, 2005.

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2. AATS-14 instrument description

The NASA Ames Airborne Tracking Sunphotometer (AATS-14) measures the direct beam solar transmission in 14 spectral channels from 354 to 2138 nm. AATS-14 azimuth and elevation motors, controlled by tracking-error signals derived from a quad-cell photodiode, rotate a tracking head to lock on to the solar beam and maintain detectors normal to it. During SOLVE II, the instrument was mounted on the NASA DC-8 in a custom manufactured fairing located on top of the aircraft just in front of the wing and slightly left of the aircraft centerline. This allowed the AATS-14 tracking head to be positioned outside and above the skin of the aircraft, which afforded an unobstructed view of the sun for all solar azimuth angles between 180° and 360° (i.e., to the left) relative to the aircraft heading for all solar zenith angles. For relative solar azimuth angles between 0° and 180° (i.e., to the right relative to aircraft heading), the view was unobstructed for all solar zenith angles <65°, but for solar zenith angles >65° it was partially obstructed by the aircraft tail, a VHF antenna, or a satellite communications antenna enclosure for certain azimuth angles. Because the other two solar occultation instruments on the DC-8 were located in the cabin and viewed out optical ports on the left side of the aircraft, all SOLVE II DC-8 flights were designed to include sun runs with the sun positioned to the left of the aircraft heading. AATS-14 was the only sun-tracking sensor that was able to acquire any data when the sun was located to the right of the aircraft heading, and useful science data have been derived from AATS-14 measurements obtained under such conditions. However, we have judiciously screened these data because analysis revealed that for certain angles these measurements were contaminated by diffuse light due to what is thought to have been reflection of the incoming solar beam off the aircraft fuselage or attached structures.

Each science channel in AATS-14 consists of a baffled entrance path, interference filter, photodiode detector, and integral preamplifier. The filter/detector/preamp sets are temperature controlled to avoid thermally induced calibration changes. The module that contains the science channels and the tracking photodiode is protected from the

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outside air by a quartz window that is exposed to the ambient air stream. In an effort to prevent window fogging or frost a foil heater is mounted on the inside surface of this window. The heater foil has a hole for each science channel and for the tracking detector, to permit the solar beam to pass through.

5 Despite the foil heater, during the first three science flights (12, 14, 16 January) out of Kiruna thin patches of frost formed on the internal surface of the quartz window and caused a decrease in the measured transmission of the incident solar radiation in 12 of the 14 channels during the flights on 12 and 14 January and in 13 channels on 16 January. This resulted in the retrieval of AOD values that were anomalously large in the affected channels and AOD spectra that were anomalously flat (for the frost-affected
10 channels). Because there were no in-flight pictures of the quartz window, the magnitude and spectral shape of the retrieved AOD spectra were initially the only evidence that the anomaly might, in fact, be an instrumental problem. The problem was correctly diagnosed as frost on the window during the second Kiruna science flight (14 January), when in-flight instrument field of view (FOV) measurements yielded anomalous
15 detector response curves, and pictures of the quartz window after the aircraft landed documented the presence of frost. The effect of frost can be modeled as a decrease in window transmission, but the actual magnitude and wavelength dependence of the attenuation attributable to the frost was unknown, so it was not possible to correct the affected data to retrieve accurate AODs. It is believed that the frost resulted from
20 residual moisture within the instrument head. The problem was solved by more extensively purging the instrument head with dry nitrogen between flights. As described in a companion paper by Russell et al. (2004) in this issue, in-flight FOV scans and post-acquisition diagnostic data analyses confirmed that window frost did not occur during
25 subsequent flights.

AATS-14 acquired “frost-free” measurements in all channels during the sunlit portions of eight science flights during the mission. Table 1 lists the dates and times of measurements for these eight flights, which included one flight that was staged out of Edwards AFB prior to the aircraft deployment to Kiruna, Sweden, six flights out of

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Kiruna, and the return transit from Kiruna to Edwards AFB. AATS-14 samples at three Hz. Every four seconds it records data consisting of an average and standard deviation of nine samples taken during the first three of the four seconds. These data are stored together with data on instrument tracking and temperature control, aircraft location, and ambient temperature, dewpoint, and static pressure. The standard deviations of all the channels are used subsequently in a cloud-screening algorithm, as described by Schmid et al. (2003). Data are transmitted serially from a computer within the instrument to a remote operator station (laptop computer). The science data are then combined with previously determined radiometric calibration values to calculate and display aerosol (particulate) optical depth, $\tau_p(\lambda)$, and columnar water vapor, CWV, in real time at the operator station.

Our methods for data reduction, calibration, and error analysis have been described in detail previously (Russell et al., 1993a, b; Schmid and Wehrli, 1995; Schmid et al., 1996, 2001, 2003). The AATS-14 channels are chosen to permit separation of aerosol, water vapor, and ozone attenuation along the measured slant path. From these slant-path transmissions we can retrieve $\tau_p(\lambda)$ in 13 narrow wavelength bands (centered at 354, 380, 453, 499, 519, 604, 675, 778, 865, 1019, 1241, 1558, and 2139 nm, with full-width-half-maximum bandwidths of 2.0, 4.6, 5.6, 5.4, 5.4, 4.9, 5.2, 4.5, 5.0, 5.1, 5.1, 4.9, and 17.3 nm, respectively) and the columnar amounts of water vapor (from the channel centered at 941 nm) and ozone. SOLVE II is the first mission in which the 2139-nm channel was flown, but these data will not be presented because we have not rigorously characterized gaseous absorption in that passband yet. In Fig. 1, the wavelengths of the AATS-14 channels are shown together with atmospheric attenuation spectra calculated using the Moderate Resolution Transmission (MODTRAN4v3.0) Code (Kneizys et al., 1996) for a view path from 10.4 km to the Sun at a solar zenith angle of 90°. This figure displays effective slant path optical thickness, $\ln(1/T)$, where T is transmission, as a function of wavelength for the primary attenuating species, including Rayleigh and aerosol scattering, and gaseous absorption. It can be seen in Fig. 1 that, in addition to the corrections for Rayleigh scattering (for which the Bucholtz,

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1995, cross sections are used) and O₃ absorption, measurements in some AATS-14 channels require corrections for NO₂, H₂O and O₂-O₂ absorption in order to retrieve AOD. Gas absorption cross-sections for the AATS-14 channels were computed using LBLRTM 6.01 (Clough and Iacono, 1995) with the CKD 2.4.1 continuum model using the HITRAN 2000 (v11.0) line-list (Rothman et al., 2003) that includes an update for water vapor from 2001. NO₂ cross-sections not included in LBLRTM 6.01 were taken from Harder et al. (1997). For results in this paper NO₂ column content was assumed constant at 2×10^{15} molecules cm⁻², which corresponds to optical depths of 0.0011 and 0.0008 in the 380-nm and 453-nm channels, respectively. Optical depth due to absorption by O₂-O₂ was less than 0.0002 in every channel except 1241 nm, where the contribution was ~0.0013 for a flight altitude of 10 km. Optical depth due to H₂O absorption was negligible ($< 5 \times 10^{-6}$) in all aerosol channels (that is, excluding 941 nm) for a flight altitude of 10 km.

Calibration of AATS-14 was achieved by analysis of sunrise measurements acquired at Mauna Loa Observatory (MLO), Hawaii, for six sunrises in November 2002 prior to SOLVE II and for seven sunrises in March 2003 after SOLVE II. Exoatmospheric detector voltages, V₀, were derived using the Langley plot technique (e.g., Russell et al., 1993a, b; Schmid and Wehrli, 1995) for all channels except 941 nm, for which a modified Langley technique (Reagan et al., 1995; Michalsky et al., 1995; Schmid et al., 1996, 2001) was employed to account for water vapor absorption. In a manner analogous to that used by Schmid et al. (2003) in analyzing AATS-14 data acquired during the 2001 Asian Pacific Regional Aerosol Characterization Experiment (ACE-Asia), we used high altitude AOD spectra acquired during DC-8 flights in December 2002 to adjust the November MLO calibration voltages for seven channels to yield “smooth” $\tau_p(\lambda)$ spectra. These adjustments were <0.3% of the unadjusted November-March mean values for four channels (675, 778, 1019, 1240 nm), and 0.6% (499 nm), 0.8% (1558 nm), and 1.3% (864 nm) of the unadjusted means for the remaining three channels. The adjusted November MLO V₀ values were used in the analysis of the December 2002 flights, but the final V₀ values applied to the January-February SOLVE

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5 II data set were simple arithmetic means of the adjusted November MLO values and the March MLO values. Corresponding V_0 uncertainties include statistical (sampling) errors that span the range of V_0 values included in the averages. Table 2 presents the MLO calibration results. With the exception of the 941-nm channel, sample standard deviations in V_0 were $\sim 0.3\%$ or less in 12 of the 13 remaining channels for the November calibration data set, and $\sim 0.2\%$ or less for all channels except 2139 nm for the March calibration data set. Most channels exhibited changes of less than 0.7% from November to March, but the 499-nm and 864-nm channels exhibited increases of 1.9% and 1.3%, respectively.

10 3. AATS-14 ozone retrieval methodology

King and Byrne (1976) developed a weighted least squares method for retrieving columnar ozone η from spectral solar irradiance measurements. Their method exploits the known spectral behavior of ozone absorption within the Chappuis absorption band (see Fig. 1) and assumes that the spectral variation in τ_p can be described by a quadratic fitting function of the form

$$\log \tau_p(\lambda) = a_0 + a_1 \log \lambda + a_2 (\log \lambda)^2, \quad (1)$$

15 where λ is wavelength and the a_i are the fitting coefficients to be determined. Figure 2, which demonstrates the method, shows an example of Eq. (1) fitted to values determined from AATS-14 measurements on the DC-8 flight of 21 January 2003. The King and Byrne method minimizes the statistic χ^2 defined as

$$\chi^2 = \sum_i \frac{1}{\sigma_i'^2} \left[\log \tau_p(\lambda_i, P, \eta) - a_0 - a_1 \log \lambda_i - a_2 (\log \lambda_i)^2 \right]^2, \quad (2)$$

20 where the summation extends over all λ_i for which no additional molecular absorption occurs, P is atmospheric pressure, and σ_i' are the standard deviations of the $\log \tau_p(\lambda_i$,

P, η) values, or

$$\sigma_i'^2 = \sigma_i^2 \left[\frac{\log e}{\tau_t(\lambda_i) - \tau_R(\lambda_i, P) - \eta a(\lambda_i)} \right]^2, \quad (3)$$

where τ_t is measured total optical depth, τ_R is calculated Rayleigh optical depth, σ_i represents the uncertainty $\delta\tau_p(\lambda_i)$ in $\tau_p(\lambda_i)$, and $a(\lambda_i)$ is the ozone absorption coefficient at wavelength λ_i . Hence, the weighting factors σ_i' are proportional to the relative uncertainties in $\tau_p(\lambda_i)$. Unfortunately, as King and Byrne note, because σ_i' is a function of η , the set of equations that results after setting the partial derivative of χ^2 with respect to each of the unknown coefficients (a_0, a_1, a_2, η) equal to zero is nonlinear in the coefficients. However, for a specific value of η , the set of simultaneous equations that results is linear in the unknowns a_0, a_1, a_2 . Hence, the problem can be solved by varying the ozone column content η until the minimum χ^2 is found where $\partial\chi^2/\partial\eta$ equals zero, for which $\eta = \eta_0$. Finally, King and Byrne point out that, in practice, the uncertainty in $\tau_p(\lambda_i)$ can be approximated well by the measurement uncertainty in $\tau_t(\lambda_i)$ in the region of the minimum χ^2 .

King and Byrne applied their method to ground-based sunphotometer measurements taken in Tucson, Arizona under light to moderate aerosol loadings (King et al., 1980). Other methods for retrieving columnar ozone have also been reported in the literature. Flittner et al. (1993) applied a second-derivative smoothing approach to retrieve total ozone, and reported results comparable with those achieved with the King and Byrne approach when both were compared with TOMS overpass measurements. Taha and Box (1999) proposed a separate method, originally outlined by Box et al. (1996), based on an eigenvalue analysis for retrieving total column ozone. Unlike the King and Byrne approach, the Taha and Box method requires a priori assumptions of the type and optical properties of the aerosol in order to construct the kernel covariance matrix necessary for the eigenanalysis. They applied both methods to one year of daily measurements obtained at Sydney, Australia with a Multifilter Rotating Shadowband Radiometer (MFRSR) (Harrison et al., 1994) and found that the eigenanalysis

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gave better agreement with TOMS overpass ozone retrievals than the King and Byrne least squares method.

We have applied the King and Byrne method to the SOLVE II AATS-14 optical depth measurements. As shown in Fig. 1 and in Fig. 2, AATS-14 includes seven channels centered between 453 and 864 nm located within the Chappuis band, including one with its center wavelength (604.4 nm) near the peak, and three with center wavelengths (499.4, 519.4 and 675.1 nm) where ozone absorption is ~25–35% of that at the peak. For the typical DC-8 SOLVE II cruising altitudes of ~8–12 km and the background stratospheric aerosol conditions that prevailed during SOLVE II, ozone absorption comprised a significant fraction of the aerosol-plus-ozone optical depth measured in the four AATS-14 channels centered between 499.4 and 675.1 nm. Typical AODs above the DC-8 ranged from 0.003–0.009 in these channels. For comparison, an ozone overburden of 0.3 atm-cm (300 DU) translates to ozone optical depths of 0.009, 0.014, 0.041, and 0.012, respectively, at these same wavelengths.

3.1. Dependence on relative optical airmass

Accurate retrieval of columnar ozone and/or AOD depends critically on the use of appropriate relative optical airmass factors for all attenuating species along the instrument line-of-sight (LOS). Airmass is defined as the ratio of LOS optical thickness (OT) to vertical OT (i.e., OD) for any given constituent and/or wavelength. When $\text{SZA} > \sim 85^\circ$, as it was for most AATS-14 measurements in SOLVE II, airmass can be highly dependent upon the shape of the vertical distribution of the attenuating species (e.g., Thomason et al., 1983; Russell et al., 1993b). Hence, airmass uncertainties result from uncertainties in requisite assumed vertical profiles of number density or extinction or simply from lack of available data. Most AATS-14 measurements during SOLVE II were acquired at large solar zenith angles (SZA) near 90° , where refraction is significant and small changes in SZA can lead to large changes in airmass.

In our analysis of the SOLVE II AATS-14 data set, we have implemented a relatively new airmass algorithm (Yee et al., 2002; DeMajistre and Yee, 2002; Ver-

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vack et al., 2002) originally developed for inversion of combined extinctive and re-
fractive stellar occultation measurements. We have modified the computer routine
that implements the algorithm to accept user-input profiles of aerosol extinction and
molecular and ozone number density. For flights out of Kiruna, we used profiles re-
trieved by the SAGE III or POAM III sensors during SOLVE II to calculate aerosol
and ozone airmass, respectively. At altitudes below the lower limit of the satel-
lite retrievals, the satellite-retrieved ozone profiles were combined with DC-8 based
in-situ measurements of ozone concentration obtained by the NASA Langley fast
response chemiluminescent ozone sensor, FASTOZ (Avery et al., 2001; see also
http://cloud1.arc.nasa.gov/solveII/instrument_files/O3.pdf), during aircraft ascent im-
mediately after takeoff from Kiruna at the start of each flight. For the December flights
out of DFRC and the return transit flight from Kiruna to DFRC where measurements
were taken at latitudes where no SAGE or POAM profiles were available, it was neces-
sary to construct best estimate models of aerosol and ozone number density profiles
from a mixture of climatological data, SAGE or POAM data taken at higher latitudes,
and coincident lidar data when available. The MODTRAN subarctic winter model atmo-
spheric density and water vapor profiles were used for calculating Rayleigh and water
vapor airmass factors, respectively. Comparison of Rayleigh airmass factors calculated
using the subarctic winter model with those derived from assimilated meteorological
measurements (e.g., UKMO) generally yielded differences of 1% or less. The Yee et
al. (2002) algorithm accounts for refraction of the incoming solar radiation by assuming
horizontal homogeneity of attenuating species within atmospheric shells of specified
thickness (typically, 0.5 km). Recently, Yee et al. (in preparation, 2005)² have modified
the algorithm to account for three-dimensional variations in species density along the
viewing path, but our results did not use this formulation. The companion SOLVE II
paper by Russell et al. (2004) presents input profiles and results from our application

²Yee, J.-H., Swartz, W. H., DeMajistre, R., Shetter, R. E., Randall, C. E., and Lloyd, S. A.:
Effects of field inhomogeneity on high-latitude ozone measurements from spaceborne remote
sensing, in preparation, 2005.

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of the algorithm. In practice, we derived airmass factors for analyzing AATS-14 data by using interpolation in matrices of airmass factors pre-calculated as a function of inertial SZA (at 0.1° resolution) and aircraft altitude (at 0.5-km resolution) for a particular SAGE or POAM occultation event. Russell et al. (2004) give details.

5 3.2. Uncertainty in ozone retrieval

King and Byrne (1976) show that the variance in the derived ozone column content η_0 may be approximated by

$$\sigma_{\eta_0}^2 \approx \frac{1}{\sum_i \frac{a^2(\lambda_i)}{\sigma_i^2}}, \quad (4)$$

where, as noted above, σ_i is the uncertainty $\delta\tau_p(\lambda_i)$ in $\tau_p(\lambda_i)$ and $a(\lambda_i)$ is the ozone absorption coefficient at wavelength λ_i . They note that if the absolute uncertainties are about the same at all wavelengths then the wavelengths with the largest ozone absorption have the largest influence on the variance, and that the magnitude of the absolute AOD error at that wavelength has nearly a linear effect on the magnitude of σ_{η_0} . Their variance expression does not account for any uncertainty in the assumed ozone absorption spectroscopy. We show below that the uncertainty due to temperature dependence of the ozone absorption cross sections within the Chappuis band is probably negligible. Their variance expression assumes further that the uncertainties $\delta\tau_p$ are well characterized. There are a number of sources of $\delta\tau_p$: uncertainties in instantaneous measured voltages $V(\lambda_i)$, calibration – i.e., exoatmospheric detector voltages $V_0(\lambda_i)$, airmass, Rayleigh optical depth, and gas absorber optical depths. For the bulk of our SOLVE II data, the effect of uncertainties in $V(\lambda_i)$ and $V_0(\lambda_i)$ on $\delta\tau_p$ is minimized because these uncertainties are inversely proportional to airmass (Russell et al., 1993a) and most of our measurements were acquired at large airmass values. However, for large airmass values, the uncertainties in airmass (Sect. 3.1) become relatively more significant, although we have made every effort to minimize these un-

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certainties through rigorous calculation of airmass factors, as described above. We believe the Rayleigh optical depth uncertainty is well known because we have an accurate measure of the ambient atmospheric pressure and the Rayleigh cross sections are well known. Because the ozone optical depth was a significant fraction of the total for the channels near the peak of the Chappuis band, the primary contributor to $\delta\tau_p$ in these channels is actually the uncertainty in the retrieved ozone column content, which depends critically on the use of correct airmass factors both for ozone and for aerosol and cannot be known a priori. In the absence of any available measurements of the vertical distributions of aerosol extinction and ozone number density, the choice (construct) of inappropriate models can lead to large errors in airmass and, hence, in the retrieved ozone and τ_p .

We calculated ozone cross sections for the AATS-14 channels by convolving the AATS-14 filter functions with output from LBLRTM for a series of runs for a range of SZAs. Figure 3 shows the LBLRTM and the SAGE III database cross sections for the 590–615 nm wavelength region, which encompasses the peak of the Chappuis band, for temperatures of 220 K and 298 K. We have overplotted our calculated cross section for the 604.4-nm AATS filter. Like the SAGE III inversion algorithm (NASA Langley Research Center, 2002), LBLRTM accounts for the temperature dependence of the ozone absorption cross sections by taking into account the distribution of ozone and atmospheric density along the LOS. As can be seen in Fig. 3, the temperature dependence of ozone absorption in the Chappuis band is small (<1% near the peak, even for the large range of temperatures shown). In fact, for our retrievals, the temperature effect is negligible since the vertical distribution of ozone peaks at altitudes where the range of temperatures is significantly less than both the range shown in Fig. 3 and the variation of temperatures within those shells that account for most of the ozone-weighted path length along the LOS. For an ozone overburden of 0.3 atm cm (300 DU), even a 1% uncertainty in the assumed ozone cross section would result in a maximum uncertainty in retrieved columnar ozone of only 0.003 (3 DU) if the only AATS measurement used in the inversion were the 604-nm data, which is not the case.

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Although most of our SOLVE II measurements were taken at large airmass values (typically >28 for aerosol and >20 for ozone), as noted above, we do present results for two flight segments where airmasses were <10 and as low as 3. As airmass decreases, uncertainties in $V_0(\lambda_i)$ become relatively more important and can lead to a bias in the retrieved ozone column. This bias does not seem to be captured adequately by the estimated variance expression (Eq. 3). We have quantified this effect empirically by comparing results from ozone retrievals performed for AATS-14 data acquired at MLO with coincident columnar ozone measurements by Brewer and Dobson spectrometers. These results are presented in Fig. 4, which shows the temporal behavior of retrieved columnar ozone for four MLO sunrise calibration events each before (November 2002, Fig. 4a) and after (March 2003, Fig. 4b) SOLVE II. Results are shown for the case where the V_0 values derived from the Langley plot analysis (which used the measured Brewer or Dobson columnar ozone) for the particular day were used in the retrieval, and for the case where the mean V_0 values (mean of March and adjusted November values) were used. For the latter, $\pm 1\sigma$ uncertainties are also plotted. In addition, both the Brewer and the Dobson measurements and curves for ozone and aerosol airmass are also shown. AATS retrievals agree with the Brewer and Dobson measurements to better than 5 DU when the calibration voltages for the specific day were used. When the mean SOLVE II calibration voltages are used, AATS retrievals for the November period also agree with the corresponding Brewer/Dobson data to within about 5 DU. However, AATS retrievals for the March data set decrease markedly with decreasing airmass and significantly underestimate the Brewer/Dobson measurements (by up to 25 DU for an airmass of ~ 2). In summary, the apparent bias in retrieved ozone column content due to calibration uncertainty is negligible for large airmass values (e.g., near 20), a few DU for airmass factors ~ 10 , but can be significant (up to 25 DU) for small airmass values (<3).

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4. Measurement results

4.1. Comparison of AATS-14 with SAGE III and POAM III

In this section, we compare AATS-14 retrievals of columnar ozone above aircraft altitude with near-coincident measurements by SAGE III and by POAM III for four occultation events each. The SAGE III occultation events occurred on four separate days: 19, 24, 29 and 31 January. The four POAM III events occurred on two days: 19 and 21 January. The satellite measurements obtained on 19 January are unsatisfactory for comparing SAGE to POAM, as they were separated by ~600 km and 3.7 h. For each occultation event, AATS-14 retrievals have been averaged during time periods when the DC-8 flew at constant altitude, and the near-coincident satellite profile of ozone number density has been integrated from the aircraft altitude to the top of the profile. Table 3 lists the times and locations of the satellite observations, and the range of AATS-14 measurement times, DC-8 altitudes, and horizontal distances from the DC-8 to the sub-orbital satellite LOS tangent point corresponding to the AATS-14 data included in each average. As can be seen in Table 3, all AATS-14 measurements used in the SAGE comparisons were acquired within 45 min of the corresponding SAGE observation, and they were separated by no more than ~200 km from the SAGE 10-km LOS tangent point. AATS-14 data used in the POAM comparisons were separated from the POAM 20-km LOS tangent point by a few minutes to a maximum of ~1.5 h and by 20–300 km.

Figure 5 shows the measured satellite ozone number density profiles and corresponding scatterplots of satellite versus AATS retrieved column ozone. The AATS mean ozone value along the flight track during each time period is plotted with a color-coded (by altitude) dot, and the range (variability) of AATS values along the flight track is shown with a horizontal bar. For the SAGE/AATS comparison, separate values are given for two aircraft altitudes (lower limit of integration for the SAGE profile) for the 19 January and 24 January events. RMS differences are 19 DU (6%) for (AATS-SAGE)/AATS and 10 DU (3%) for (AATS-POAM)/AATS comparisons. As shown by

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the horizontal bars in Fig. 5, spatial variability along the AATS flight track is likely a significant contributor to these RMS differences.

This spatial variability is evident in Fig. 6, which examines the 24 January case in more detail. The temporal variation in AATS AOD is presented in Fig. 6a for the time period 9.8–12.2 UT, and the corresponding variation in vertical columnar ozone, aircraft altitude, and SZA is shown in Fig. 6b. The SAGE III solar occultation occurred at 10.98 UT. AATS LOS ozone values (airmass times vertical column ozone) are overplotted with coincident DIAS ozone retrievals and with GAMS relative ozone retrievals (derived from measurements acquired at two different solar zenith angles and normalized to AATS values at 11.52 UT) in Fig. 6c. Mean (AATS-DIAS)/AATS differences are –1.75%, with an RMS difference of 2.1%. Mean (AATS-GAMS)/AATS differences are <1%, but with an RMS difference of 4.2%. The largest differences occurred during the period before 10 UT, when SZAs exceeded 90°.

Figure 6d shows the variation in AATS-retrieved columnar ozone along the DC-8 flight track above the DC-8 flight altitude of 9.9 km. It includes data from two flight segments, 10.6–10.9 UT and 11.8–12.2 UT, as shown in Fig. 6b. Only data acquired during the first flight segment (which is closer to the SAGE III occultation time of 10.98 UT) were averaged for the 9.9-km data point shown in Fig. 5c. The location of the SAGE III 10-km tangent altitude suborbital location is shown in Fig. 6d with a large dot that is color coded to reflect the SAGE III columnar ozone value of 422.7 DU, which was calculated by integrating the SAGE III ozone number density profile above 9.9 km. Corresponding AATS-14 retrievals range from 401 to 423 DU for the two flight segments, with the maximum measured at ~10.65 UT just to the southwest of the SAGE tangent location. The AATS-14 columnar ozone retrieval at the 9.9-km altitude DC-8 location closest in time (~10.6 UT) and space (~8 km) to the SAGE III 10-km tangent point is 417 DU, or ~6 DU (1.3%) less than the SAGE III result.

Figure 6e presents analogous AATS-14 results for measurements acquired at an altitude of ~8.74 km during the 10.98–11.71 UT period. These data are included in Fig. 5c as a separate data point. During this flight segment AATS ozone retrievals for

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measurements closest in time and space to the SAGE III 10-km tangent location were 436 DU. The lower limit of the SAGE III retrieval was 9.5 km, and the integrated value above this altitude was 425 DU. Assuming a constant ozone density between 8.74 km and 9.5 km equal to the SAGE retrieval at 9.5 km yields a columnar ozone value of 428 DU above 8.74 km, or 8 DU (1.8%) less than the AATS-14 8.74-km result (436 DU) closest in time and space to SAGE III.

4.2. Comparison of AATS-14 with GOME and TOMS

In this section, we compare values of overlying columnar ozone retrieved from AATS-14 measurements acquired during extended horizontal transects during three flights (19–20 December, 21 January, 6 February) with corresponding TOMS and GOME total ozone column retrievals. For TOMS, the Version 8 (V8) Level-3 global gridded (1° latitude by 1.5° longitude) daily mean product (McPeters et al., 1998) has been used. For GOME, one or both of the harmonic analysis daily mean product or the Kalman-filtered assimilated product (Bittner et al., 1997; Meisner et al., 1999) has/have been used.

In order to compare the AATS retrievals with the total column satellite retrievals, it is necessary to estimate the amount of ozone below the aircraft altitude. We have made this estimate in two ways: by using an assumed climatological model (either the 1976 standard atmosphere ozone model or a seasonal/latitudinal ozone model (McPeters et al., 2003) which is a composite of balloon, SAGE I and SAGE II measurements), and/or by incorporating in-situ data obtained by the FASTOZ instrument during DC-8 ascent at the beginning or aircraft descent at the end of each science flight. Typically, the FASTOZ profiles were separated from the satellite observations by several hours. The relative contribution of the amount of ozone below the aircraft to the total column will be discussed in the case analyses below.

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4.2.1. DC-8 longitudinal transect on 21 January 2003

Figure 7 presents results from the DC-8 flight from Sweden to Greenland on 21 January. Corresponding TOMS and GOME total column ozone retrievals are shown in color contour maps in Fig. 7a and b, respectively, and the DC-8 flight track is overplotted in magenta. The GOME data represent the daily harmonic retrieval product.

Figure 7c compares ozone retrievals by AATS-14 with those from the satellite sensors. At DC-8 flight locations between 16° W and 56° W (14.6–17.1 UT) columnar ozone retrievals by the three DC-8 solar occultation sensors (AATS-14, DIAS, and GAMS) are in good agreement, with RMS differences of <1% of the AATS value for AATS-DIAS and <3% for AATS-GAMS. DIAS retrievals are shown in Fig. 7c, but GAMS results are not presented here. LOS retrievals from measurements acquired by all three DC-8 solar occultation sensors during this flight can be found in Pitts et al. (to be submitted, 2005)¹. For comparison of AATS-14 results with GOME and TOMS total column retrievals, the ozone distribution below aircraft altitude has been assumed equal to that measured by FASTOZ during the DC-8 ascent out of Kiruna, and an amount equal to the integral of the FASTOZ measurements between the surface and the aircraft altitude has been added to each AATS-14 ozone retrieval. The resultant values are displayed in Fig. 7c with a thick black line. The amount of ozone below the aircraft increased the AATS-14 retrievals by ~6.5–12% (20–37 DU), which was ~7–13.7% of the total columnar ozone. Corresponding GOME and TOMS retrievals are shown in Fig. 7c for comparison. Only TOMS data for grid points located within 120 km of the DC-8 flight track have been included in the comparison. Satellite and DC-8 based results appear mutually consistent, in light of the different spatial and temporal resolutions of the measurements. The RMS difference between AATS-14 and GOME retrievals interpolated to the DC-8 flight track is 3.3% (12 DU).

DC-8 based solar occultation measurements during the 21 January flight were obtained at refracted SZAs that ranged from ~84° to ~89° (Fig. 7d), with SZAs less than 87° over the longitude range (0–50° W) for which GOME retrievals were available along

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the DC-8 flight track. These values represented the smallest SZAs during any of the Kiruna-based science flights. For analysis of the AATS-14 measurements, ozone air-mass factors were calculated using a single profile constructed from the FASTOZ measurements taken below 10 km during the ascent (12.2 UT, 67.8° N, 20.3° E to 12.8 UT, 67.2° N, 12.8° E) after takeoff from Kiruna and from the SAGE III ozone retrieval at and above 10 km for the 12.77 UT (69.77° N, 19.19° W) occultation. Aerosol airmass factors were calculated using POAM III aerosol extinction retrievals from the 18.02 UT (65.16° N, 48.90° W) satellite solar occultation event. Ozone airmasses ranged from ~8.0–17.3, and aerosol airmasses ranged from ~8.6–35.4. Russell et al. (2004) show time series of these airmasses.

4.2.2. DC-8 latitudinal transect on 19–20 December 2002

On 19 December 2002, the DC-8 flew a roundtrip latitudinal transect from Edwards AFB (34.96° N, 117.78° W) to 22.6° N, 116.1° W. As shown in Fig. 8, the flight track was designed to traverse the subtropical jet core and include a short segment in stratospheric air near 31° N during the southward transect, which was flown at altitudes between 10.1 and 10.5 km, and a longer segment in the stratosphere north of ~30° N during the northward transect, which was flown at altitudes between 12.4 and 12.1 km. During the southbound flight segment AATS-14 data were acquired at SZAs that increased from 63.8° at 33.5° N (21.9 UT) to 70.3° at the southernmost point of 22.6° N (23.3 UT). During sunset on the northbound return segment AATS-14 successfully tracked the sun down to a SZA of 91.2° (~31.3° N, 25.0 UT). However, due to uncertainties in ozone and aerosol airmass factors, AATS-14 ozone retrieval results are presented in Fig. 9 only for SZAs <89°, and they are compared with corresponding TOMS and GOME retrievals and with smoothed (using a five-point running mean filter on data at 30-sec resolution) DIAS measurements, which were obtained only during the northbound segment. DIAS columnar ozone results were calculated from the DIAS LOS ozone retrievals by applying the same ozone airmass factors that were used in the analysis of AATS-14 data. DIAS and AATS-14 columnar ozone results agree to within 1%. No GAMS data

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are available for this flight. For direct comparison with TOMS and GOME total column amounts, the AATS retrievals have been increased by an amount of ozone (23–26 DU) below aircraft altitude that has been estimated using the McPeters et al. (2003) climatological results for December, 20°–30° N. These contributions represented ~7.5–10% of the total columnar ozone.

AATS-14 retrievals indicate a decrease of ~40 DU above the aircraft between 33° N and 23° N during the southward transect, including a decrease of ~20 DU between 30° N and 23° N. A slightly larger increase was found over this same latitude range during the northbound return at the higher altitude. AATS-14 and DIAS retrievals agree to within a few DU except near the boundaries of the DIAS data, where larger differences lead to a composite RMS difference of 5.8 DU. Corresponding TOMS and GOME retrievals interpolated to the DC-8 flight track are shown for different time periods that encompass the DC-8 flight times. The daily mean TOMS retrievals show an increase of ~10–30 DU at similar latitudes north of 27° N between 19 and 20 December. The GOME data, which include the harmonic analysis retrievals for 19 and 20 December (labeled 12 UT), and the Kalman-filtered product for 23 UT 19 December, reflect a much smaller increase (~10 DU) in columnar ozone north of 28° N than reported by TOMS.

Retrievals from both satellites exhibit the same general latitudinal behavior – namely, higher ozone to the north, as observed by AATS-14. However, after correction for an estimated amount of ozone below aircraft altitude, the AATS-14 data exceed GOME and TOMS 19 December data by ~20–30 DU at all latitudes, but agree to within 10 DU or less with 20 December TOMS data north of 28° N.

The DC-8 based NASA Langley Differential Absorption Lidar (DIAL) instrument (Browell et al., 1998; Grant et al., 1998) measured ozone concentrations at altitudes between 14 and 26 km during the northbound transect for latitudes north of 25° N. The lidar data indicate an increasing ozone concentration to the north above 20 km, with a broad peak in concentration centered at ~22 km altitude north of ~29.4° N. Because the lidar data cover a limited vertical range, vertical integration of the lidar data set

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yields layer ozone values much less than the columnar content (above the aircraft) measured by AATS and DIAS. However, the lidar data are extremely useful because they provide the only measurement of the vertical distribution of ozone through the peak of the stratospheric ozone layer along the DC-8 flight track. The data have been combined with the McPeters et al. (2003) climatological ozone profile at altitudes above and below to construct a composite profile that was used to calculate ozone airmass values that best reflect the actual vertical distribution of ozone in that region.

4.2.3. DC-8 transit from Kiruna to DFRC on 6 February 2003

Figure 10 compares AATS-14 and DIAS ozone retrievals from measurements acquired during the final segment of the DC-8 transit back to California on 6 February with TOMS and GOME values. AATS-14 results are shown for data acquired during the period 15.0–16.8 UT, when SZAs decreased from 90° just south of the U.S./Canada border to 68.5° upon landing at Edwards AFB. This corresponded to ozone airmasses of 16 to 2.6. DIAS data are only available as far south as 37° N. Again, DIAS columnar ozone results were calculated from the DIAS LOS retrievals by applying the same ozone airmass values used in the AATS-14 data analysis. The mean (AATS-DIAS)/AATS difference was 0.6%, with an rms difference of 2.3%. AATS-14 values are shown with dark blue and red symbols, respectively, for data acquired at about 12.4 km constant altitude and during the descent into Edwards AFB. The thick yellow line shows AATS-14 results obtained by adding an estimated amount of ozone below DC-8 altitude by using FASTOZ measurements during descent into DFRC or the 1976 standard atmosphere ozone model profile after normalization to AATS-14 inferred columnar ozone above aircraft altitude. AATS retrievals show a distinct peak in columnar ozone of ~470 DU above the aircraft (12.3–12.5 km) from 45° to 40° N, and a sharp latitudinal gradient with a minimum value of ~290 DU at 35.5° N. During the descent from 12.5 km to 1.1 km (altitude of Edwards AFB) AATS observed a 60 DU increase in ozone to ~350 DU.

Gridded data from TOMS and GOME have been interpolated to the locations of the DC-8 flight track. Two sets of GOME retrievals are shown: the 12 UT harmonic

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analysis composite product, and the 18 UT Kalman filter data assimilation. The composite retrieval resolves more of the latitudinal variation than does the Kalman product. The agreement between the latitudinal variation in columnar ozone measured by AATS retrievals and that measured by both TOMS and GOME retrievals is striking. After
5 correcting AATS values for the amount of ozone below the aircraft (42–58 DU, representing ~11–13% of the total column, for the DC-8 maximum altitude of 12.4 km), AATS and TOMS retrievals agree to better than 10 DU, except at the surface, where the AATS value of 350 DU is about 30 DU greater than the interpolated TOMS value. Upon landing, the ozone airmass value was 2.6, which, in light of the uncertainty discussion in Sect. 3.2 above, could result in an uncertainty in the AATS value measured
10 at the surface of up to 25 DU.

The descent into Edwards AFB permitted a direct comparison of coincident AATS-14 and FASTOZ measurements. Two approaches have been taken. Columnar ozone values are compared in Fig. 11a and ozone concentrations in Fig. 11b. In Fig. 11a, the
15 FASTOZ ozone column at 12.2 km has been set equal to the AATS-14 ozone overburden retrieved at that altitude. The AATS-14 retrieval overestimates the amount of ozone between 12.1 and 1.6 km, as measured by FASTOZ, by ~12 DU (or $3.22 \times 10^{17} \text{ cm}^{-2}$). In Fig. 11b, differentiation of a spline fit to the AATS-14 columnar ozone profile, shown in Fig. 11a, yields a vertical profile of ozone concentration that is unable to reproduce
20 the fine-scale structure measured by FASTOZ. This result is not surprising, as ozone concentration is a second order product of the AATS-14 ozone retrieval.

5. Summary and conclusions

Measurements with AATS-14 during DC-8 flights in SOLVE II have been analyzed using least squares methodology to retrieve values of columnar ozone above aircraft altitude.
25 These values have been compared with measurements by SAGE III and by POAM III for four coincident occultation events each, yielding RMS differences of 19 DU (6% of the AATS value) for AATS-SAGE and 10 DU (3%) for AATS-POAM. Line of sight ozone

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retrievals from coincident measurements obtained by the three DC-8 solar occultation instruments (AATS-14, DIAS, and GAMS) during the SAGE III solar occultation event on 24 January yield RMS differences of 2.1% for AATS-DIAS and 4.2% for AATS-GAMS.

5 AATS-14 columnar ozone retrievals from data acquired during extended horizontal transects during three flights (19–20 December, 21 January, 6 February) have been compared with corresponding TOMS and GOME total column ozone retrievals. In these comparisons, the AATS values are increased by the estimated amount of ozone below the DC-8 altitude, obtained either from a climatological model or from mea-
10 surements acquired with the on-board fast response chemiluminescent ozone sensor (FASTOZ) during the aircraft ascent at the start of each science flight or descent at the end of each science flight. Resultant AATS and satellite retrievals are found to be mutually consistent for the 21 January and 6 February flights, with RMS differences generally <10–15 DU. AATS and DIAS retrievals agree to within an RMS difference of
15 1% of the AATS value for the 21 January flight and 2.3% for the 6 February flight. Corresponding AATS-GAMS RMS differences are ~3% for the 21 January flight, but are not compared for the 6 February flight. For the round trip DC-8 latitudinal transect from 34° N to 22° N on 19–20 December, coincident DC-8 based measurements by AATS-14 and by DIAS yield ozone retrievals that agree with each other to ~1%. After adding an
20 estimated amount of ozone below the aircraft to the AATS-14 retrievals, the resultant total column ozone values yield a latitudinal gradient that is similar in shape to that observed by TOMS and GOME, but they exceed the satellite values by up to 30 DU at certain latitudes. These differences in absolute magnitude are attributed to spatial and temporal variability that was associated with the dynamics near the subtropical jet but
25 was unresolved by the satellite sensors.

AATS-14 ozone retrievals have been compared directly with high time resolution in-situ ozone measurements (the NASA Langley FASTOZ instrument) during the DC-8 descent into Edwards AFB on 6 February. Corresponding layer ozone values agree to within ~12 DU for the layer between 12.4 and 1.1 km altitude. Differentiation of

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a spline fit to the AATS-14 columnar ozone values yields a coarse vertical profile of ozone number density that is unable to reproduce the fine structure seen by FASTOZ.

In this study, we have investigated the retrieval of ozone column content from airborne sunphotometer aerosol optical depth spectra through application of the King and Byrne (1976) least squares methodology only. Regardless of the retrieval scheme employed, we conclude that retrieval of ozone column content from multiwavelength airborne sunphotometer measurements in the Chappuis absorption band can provide a valuable means of validating temporally and spatially coincident or near-coincident satellite sensor ozone measurements, but only if retrieval uncertainties can be minimized by acquiring data under all of the following conditions:

1. The midvisible aerosol optical depth above the aircraft is less than or equal to the corresponding overlying ozone optical depth, which is ~ 0.04 for an ozone overburden of 0.3 atm-cm (300 DU).
2. Refracted SZAs are $> \sim 80^\circ$, corresponding to aerosol and ozone airmass factors $> \sim 5.8$.
3. The relative shapes of the overlying vertical distributions of aerosol extinction, ozone number density, and neutral molecular number density are known or can be modeled to within a few percent.
4. For SZAs $> \sim 83^\circ$, the vertical temperature distribution is sufficiently well known or modeled to permit accurate calculation of atmospheric refraction along the instrument to Sun line of sight.
5. The line of sight is cloud-free.

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Table 1. DC-8 flight dates and times of frost-free AATS-14 measurements during SOLVE II.

Date ^a	Start time ^b	Stop time ^b	No. of measurements
19/12/2002	21:54:00	24:52:46	2142
19/01/2003	09:27:04	13:34:19	2720
21/01/2003	12:30:15	17:46:51	4286
24/01/2003	10:34:09	12:11:32	1163
29/01/2003	10:18:00	11:27:33	717
31/01/2003	08:30:14	10:03:18	1076
02/02/2003	09:54:41	12:55:36	2364
02/06/2003	15:00:03	16:49:47	1551

^a All dates in dd/mm/yyyy format.

^b All times in hh:mm:ss UT format.

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Table 2. Results of Langley plot analysis of AATS-14 measurements acquired during sunrises at Mauna Loa Observatory before and after SOLVE II. Mean and standard deviation of exoatmospheric detector voltages, V_0 , are listed. November results include six sunrises; March results include seven sunrises.

Wavelength (nm):	353.5	380.0	452.6	499.4	519.4	604.4	675.1	778.4	864.5	940.6	1019.1	1241.3	1557.8	2139.3
November 2002														
Mean:	10.893	9.819	8.966	8.469	8.918	6.345	7.348	6.976	7.765	5.776	8.114	8.106	9.394	8.424
Relative SD (%):	0.23	0.46	0.21	0.31	0.11	0.18	0.14	0.21	0.13	2.84	0.14	0.08	0.10	0.13
Adjusted November 2002														
Mean:	10.893	9.819	8.966	8.511	8.918	6.345	7.333	6.986	7.866	5.776	8.098	8.122	9.470	8.424
March 2003														
Mean:	10.878	9.796	8.994	8.669	8.935	6.362	7.283	6.945	7.727	5.642	8.065	8.024	9.395	8.489
Relative SD (%):	0.20	0.23	0.15	0.13	0.12	0.08	0.14	0.06	0.13	1.71	0.08	0.12	0.10	0.27
Combined March and adjusted November														
Mean:	10.886	9.808	8.980	8.590	8.927	6.353	7.308	6.965	7.814	5.709	8.081	8.073	9.447	8.456

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Table 3. Times and locations of SAGE III and POAM III satellite occultation events coincident with DC-8 based AATS-14 measurements; AATS-14 data times with corresponding DC-8 altitudes, and DC-8 distances from sub-track satellite occultation tangent altitude point (SAGE – 10 km, POAM – 20 km).

DATE	SAGE UT (hr)	SAGE Lat (deg)	SAGE Lon (deg)	AATS UT (hr)	DC-8 Altitude (km)	Horizontal Separation (km)
19 Jan.	9.87	69.31° N	24.21° E	9.69–9.80	9.60–9.80	39–99
				10.32–10.62	10.41–10.45	12–99
24 Jan.	10.98	70.61° N	11.09° E	10.60–10.88	9.91–9.94	7–205
				11.00–11.50	8.71–8.76	3–178
29 Jan.	10.33	72.00° N	25.74° W	10.13–10.43	10.28–10.32	20–100
31 Jan.	9.72	72.55° N	36.90° W	9.38–9.82	9.05–9.08	24–150

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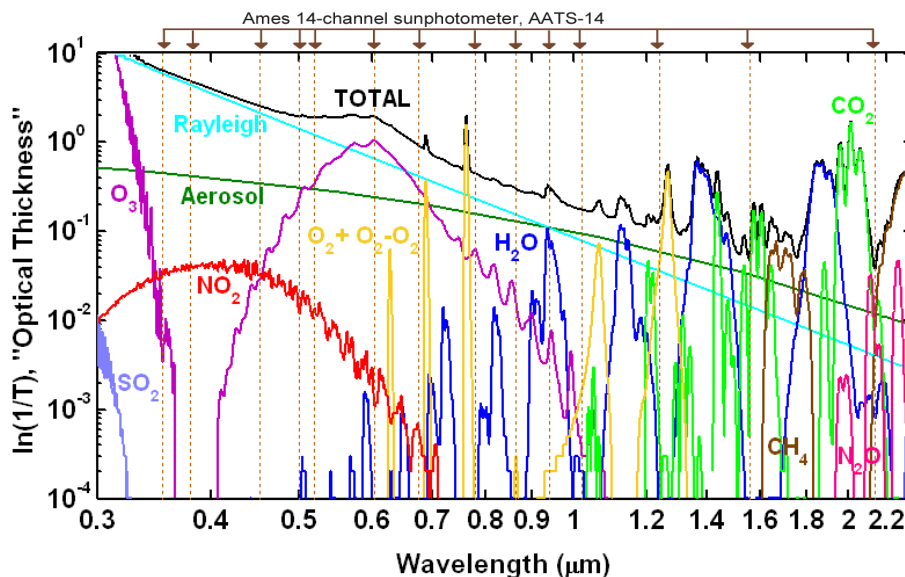


Fig. 1. Atmospheric attenuation (T is transmission) spectra calculated using MODTRAN4v3.0 for a view path from 10.4 km to the Sun at a solar zenith angle of 90° . The passband centers of the AATS-14 channel filters are also shown.

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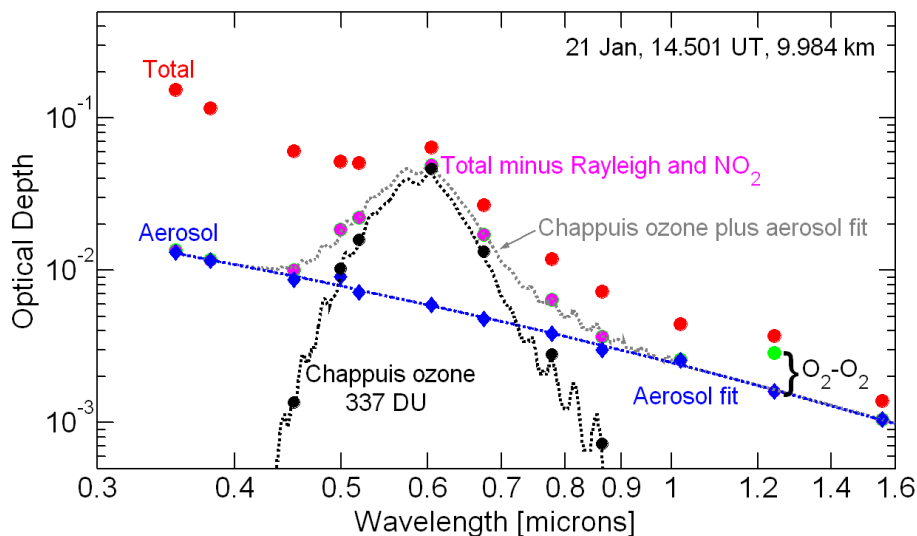


Fig. 2. Ozone retrieval methodology applied to optical depth spectrum measured by AATS-14 on 21 January 2003. The optimal least squares fit yields an ozone column content of 337 DU.

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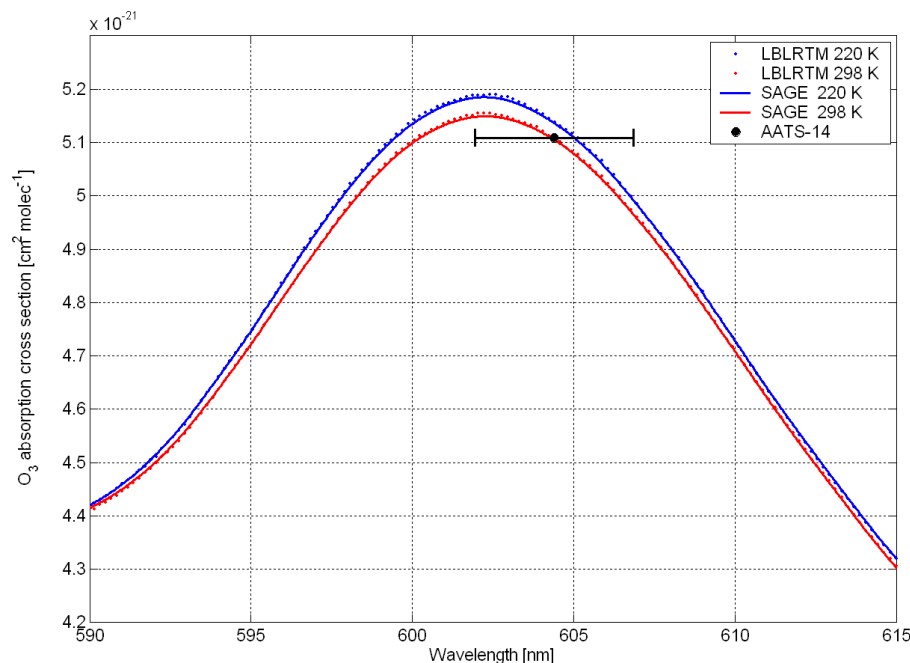


Fig. 3. Ozone cross sections from LBLRTM and the SAGE III database near the peak of the Chappuis absorption band. The black dot shows the value calculated by convolving the spectral response curve for the AATS-14 channel centered at 604.1 nm (with a FWHM of 4.9 nm, as shown by the horizontal bar) with output from a series of LBLRTM runs for a range of SZAs.

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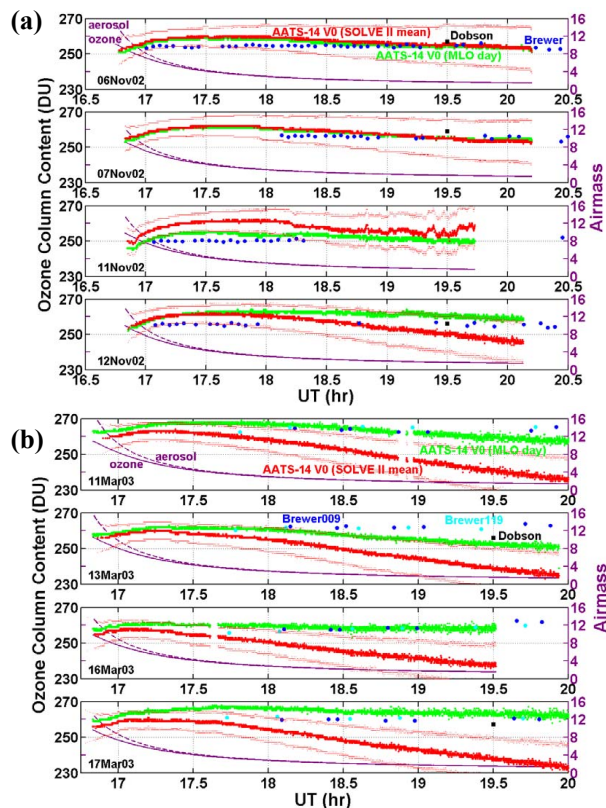


Fig. 4. Ozone column content retrieved from AATS-14 measurements after sunrise during instrument calibration at Mauna Loa Observatory, Hawaii in **(a)** November 2002 and **(b)** March 2003. Green points show retrievals that use exoatmospheric detector voltages (V0) calculated from the Langley plot analysis for the particular day; thick red points use the mean V0 values from the November and March data sets, and thin red points give the \pm one standard deviation uncertainty returned by the fitting procedure. Coincident Brewer (blue and cyan) and Dobson (black) spectrometer ozone retrievals are also shown. Aerosol and ozone airmass values are plotted in purple.

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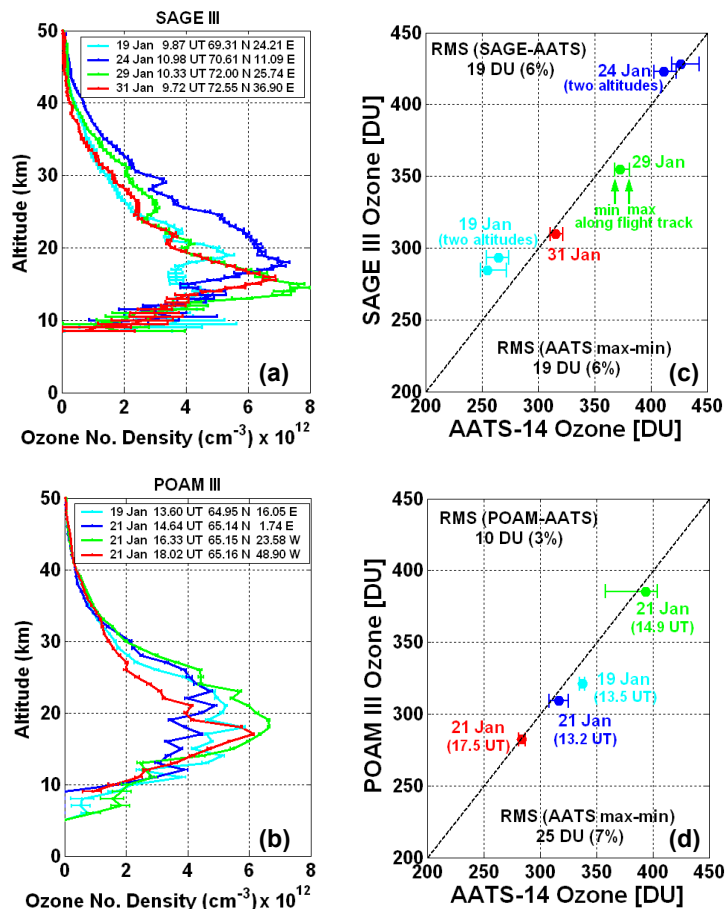


Fig. 5. SAGE III (a) and POAM III (b) ozone profiles; comparison of columnar ozone measured by (c) SAGE III and AATS-14 and (d) POAM III and AATS-14. Bars represent minimum and maximum AATS-14 retrievals along the DC-8 flight track.

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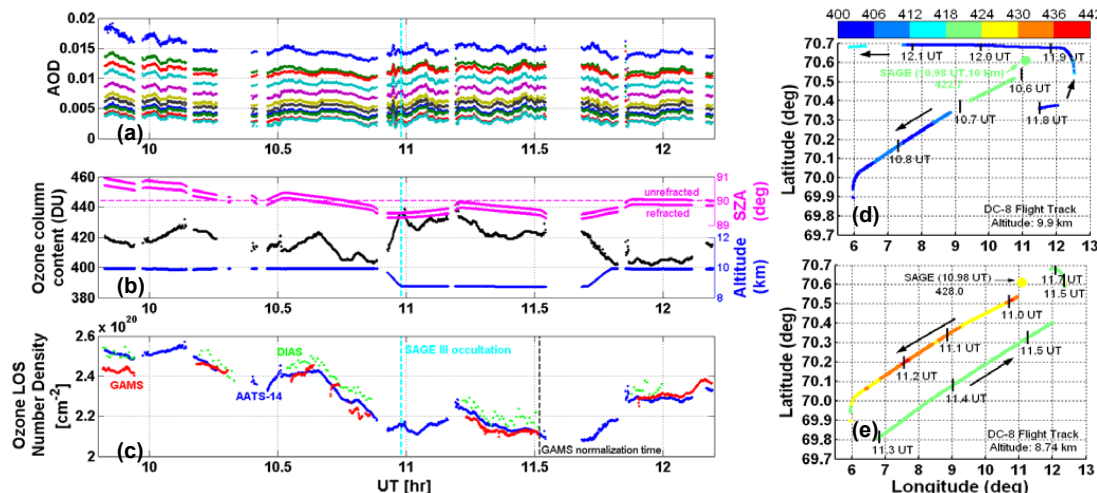


Fig. 6. Frames (a) and (b) show the variation in AOD and columnar ozone, respectively, measured by AATS-14 during the 24 January 2003 flight. AATS-14, DIAS, and GAMS retrievals of ozone number density along the line of sight (LOS) are plotted in (c). Rms differences are $\sim 2.1\%$ for AATS and DIAS, and $\sim 4.2\%$ for AATS and GAMS. In (d) and (e), the spatio-temporal variation in AATS-14 columnar ozone along the DC-8 flight track is shown for data acquired at flight altitudes of 9.9 km (10.6–10.9 UT and 11.8–12.2 UT) and 8.74 km (11.0–11.7 UT), respectively. Values are consistent with the integrated 10.98 UT SAGE III retrieval (color coded dot).

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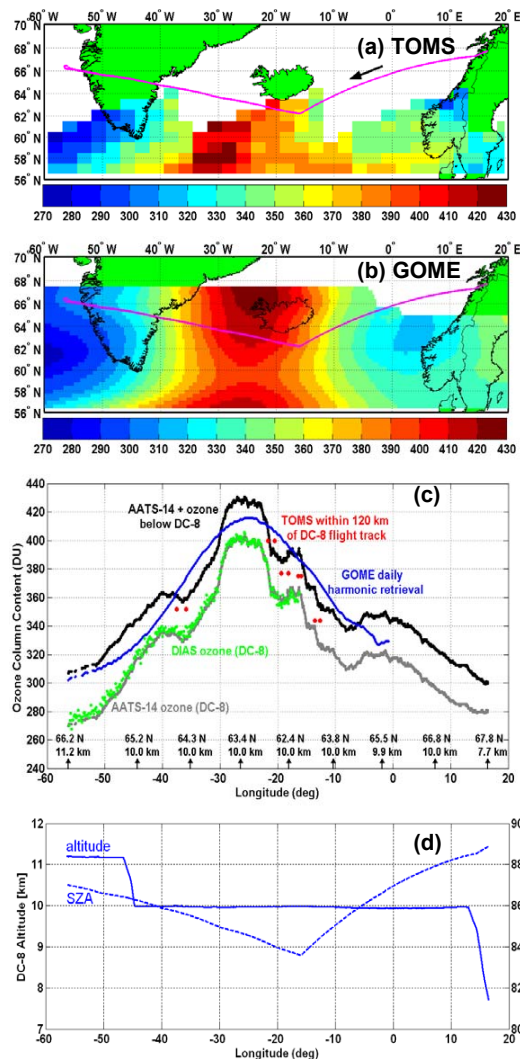


Fig. 7. Comparison of columnar ozone retrieved from AATS-14 measurements during an extended horizontal transect by the DC-8 on 21 January 2003. Contours of ozone column content retrieved from (a) TOMS and (b) GOME measurements overplotted on the DC-8 flight track (magenta). Comparison of AATS-14 (gray) and DIAS (green) ozone above DC-8 altitude, and AATS-14 (black) with along-track GOME (blue) and near-track TOMS (red) total column ozone in (c). Solar zenith angle (dashed) and DC-8 altitude (solid) are shown in (d).

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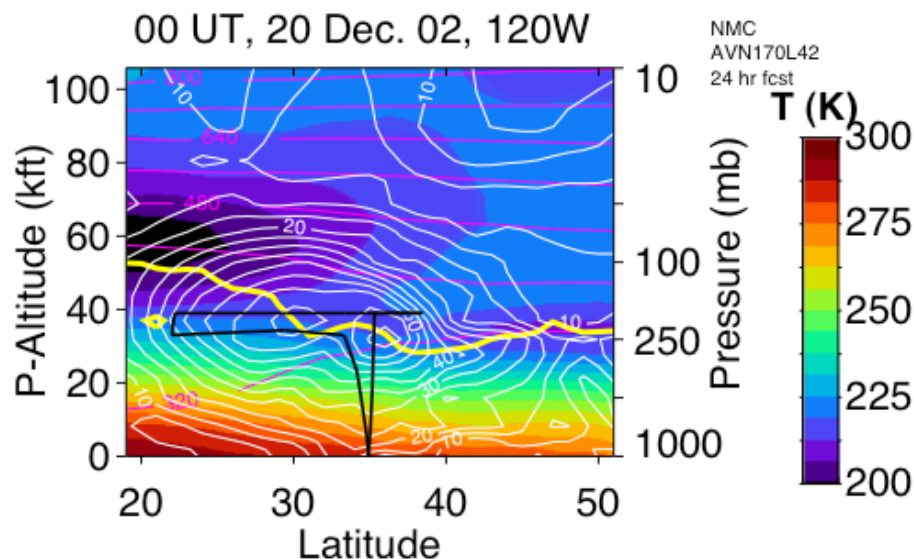


Fig. 8. Planned 19–20 December 2002 DC-8 flight track (solid black) overplotted on 24-h NMC AVN forecast of winds (white contours), isentropic surfaces (magenta lines), and tropopause location (solid yellow line) along 120° W longitude. The DC-8 southbound transit was at the lower altitude. Maps were produced by L. Lait, M. Schoeberl, and P. Newman.

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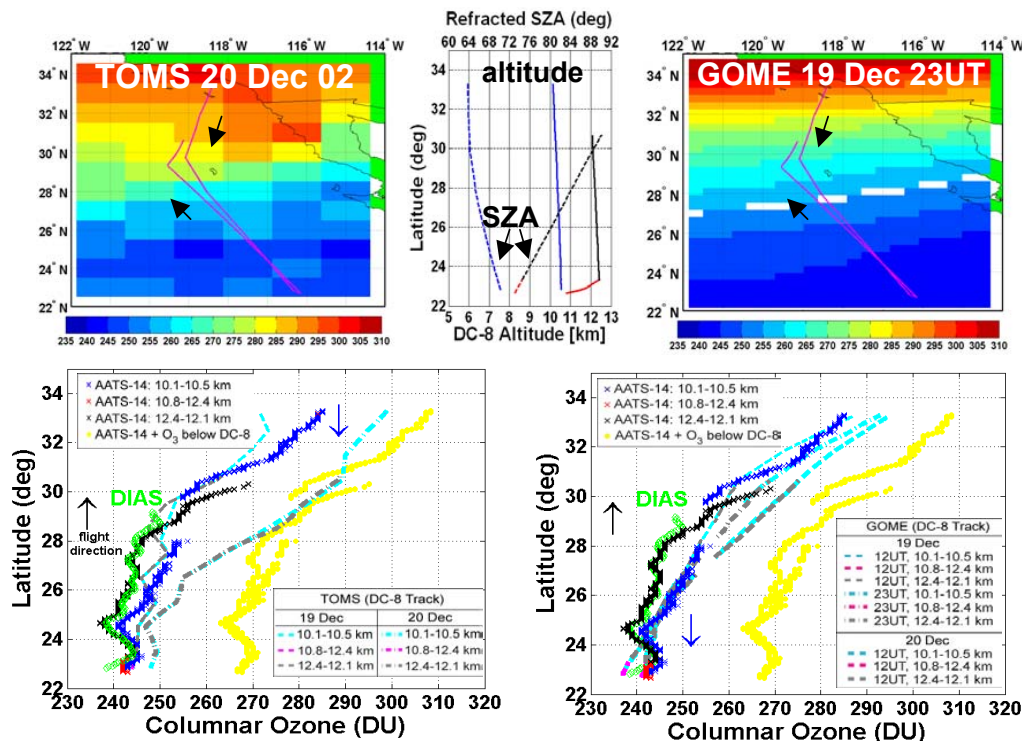


Fig. 9. Comparison of AATS and DIAS columnar ozone retrievals for DC-8 flight of 19–20 December 2002 with corresponding TOMS and GOME retrievals. The actual DC-8 flight track (magenta) is overplotted on contours of TOMS and GOME columnar ozone in the left and right upper frames; the upper middle frame shows the refracted SZA and DC-8 altitude during the southbound transit (blue), the climb from 10.8 km to 12.4 km altitude at the southernmost point (red), and the northbound transit (black). The same color scheme is used for the AATS-14 ozone retrievals shown in the lower frames. The yellow points show AATS ozone retrievals plus an estimated ozone amount below DC-8 altitude.

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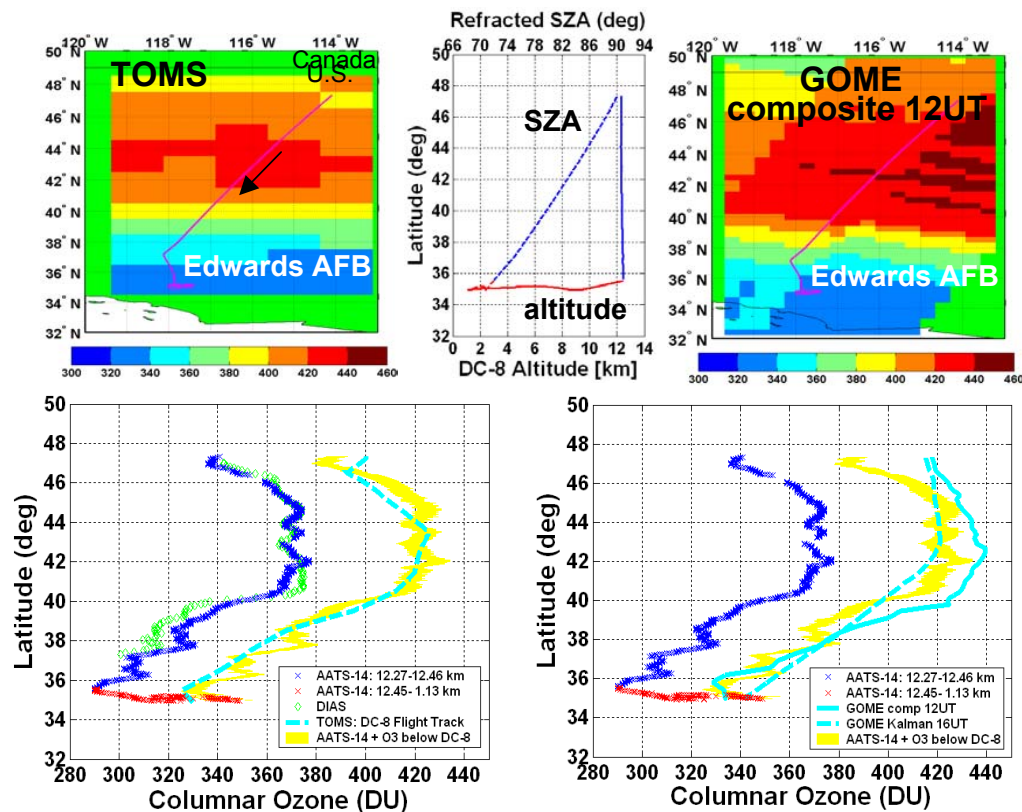


Fig. 10. Comparison of AATS and DIAS columnar ozone retrievals for DC-8 flight of 6 February 2003 with corresponding TOMS and GOME retrievals. Yellow band shows AATS ozone retrieval plus an estimated amount below DC-8 altitude.

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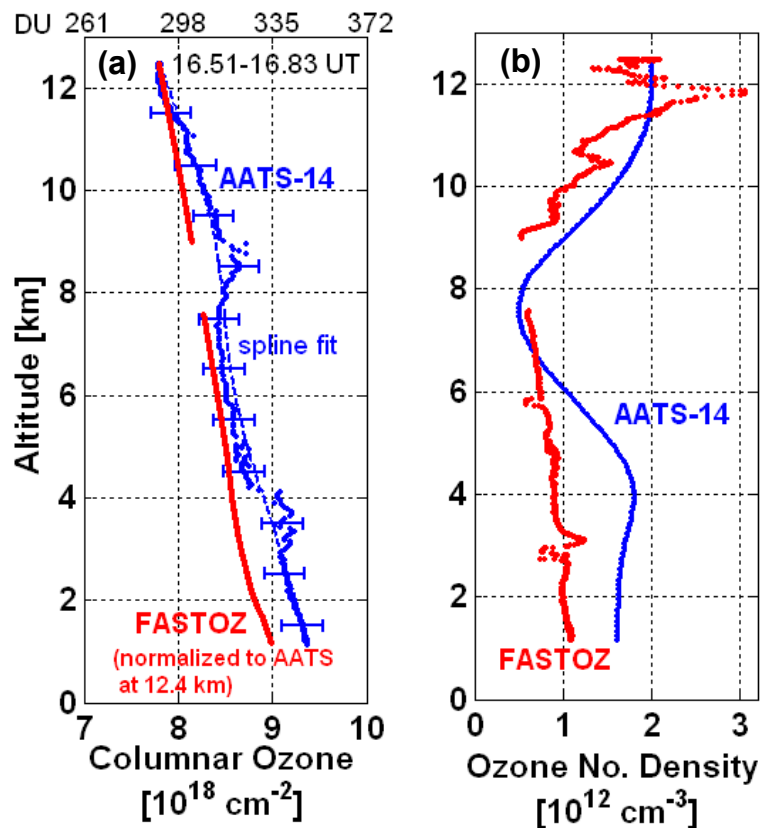


Fig. 11. Comparison of AATS-14 and FASTOZ ozone measurements during DC-8 descent into Edwards AFB on 6 February 2003. FASTOZ measurements of ozone concentration **(b)** have been normalized to the AATS-14 columnar ozone retrieval at 12.4 km altitude and vertically integrated for comparison to AATS-14 retrievals in **(a)**. Differentiation of a cubic spline fit to AATS-14 columnar ozone retrievals (a) yields ozone concentrations that are compared to FASTOZ ozone number density in (b).

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