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# Enhanced radar precipitation estimates using a combined clutter and beam blockage correction technique

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**Abstract.** Weather radar observations are currently the most reliable method for remote sensing of precipitation. However, a number of factors affect the quality of radar observations and may limit seriously automated quantitative applications of radar precipitation estimates such as those required in Numerical Weather Prediction (NWP) data assimilation or in hydrological models. In this paper, a technique to correct two different problems typically present in radar data is presented and evaluated. The aspects dealt with are non-precipitating echoes – caused either by permanent ground clutter or by anomalous propagation of the radar beam (anaprop echoes) – and also topographical beam blockage. The correction technique is based in the computation of realistic beam propagation trajectories based upon recent radiosonde observations instead of assuming standard radio propagation conditions. The correction consists of three different steps: 1) calculation of a Dynamic Elevation Map which provides the minimum clutter-free antenna elevation for each pixel within the radar coverage; 2) correction for residual anaprop, checking the vertical reflectivity gradients within the radar volume; and 3) topographical beam blockage estimation and correction using a geometric optics approach. The technique is evaluated with four case studies in the region of the Po Valley (N Italy) using a C-band Doppler radar and a network of raingauges providing hourly precipitation measurements. The case studies cover different seasons, different radio propagation conditions and also stratiform and convective precipitation type events. After applying the proposed correction, a comparison of the radar precipitation estimates with raingauges indicates a general reduction in both the root mean squared error and the fractional error variance indicating the efficiency and robustness of the procedure.

Moreover, the technique presented is not computationally expensive so it seems well suited to be implemented in an operational environment.

## 1 Introduction

During the last decades there has been an increasing need to improve the quality of weather radar observations. Large efforts have been devoted to bridge the gap between traditional qualitative uses, such as weather surveillance, to radar automated quantitative processing required nowadays by a number of applications. For example, the recent EU COST 717 concerted action has remarked the potential applications of weather radar systems (Rossa, 2000), addressing specifically the assimilation in Numerical Weather Prediction (NWP) models and also in hydrological forecast systems (Bruen, 2000). In particular, the automated use of radar data has impelled a deeper understanding of the plethora of phenomena affecting radar observations. Methods to detect and correct such effects need to be incorporated within the quality control of radar data (Alberoni et al., 2003).

One of the typical factors affecting the quality of radar images intended for quantitative precipitation estimates (QPE) is the presence of non-precipitating echoes. They are mostly caused by the back-scatter of radar energy intercepted by ground targets or sea waves, the so-called ground or sea clutter echoes (Collier, 1996, 1998). Sometimes the quantity and intensity of such echoes is increased because the vertical profile of air temperature and moisture contains sharp gradients that deviate or refract more than usual the radar beam trajectory to the ground, which under normal circumstances is bent away from the ground (Bean and Dutton, 1968). This phenomenon is known as super-refraction and the new echoes

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are commonly called AP, anaprop or anomalous propagation echoes. Ducting, an extreme case of super-refraction, is produced when the radar energy is trapped within an air layer and may travel for long distances before intercepting ground targets.

Another important factor that should be considered in weather radars operating in complex topography regions is the effect of radar beam blockage by the surrounding hills with similar or higher altitudes. The screening effect of mountains prevent the radar observing the complete volume of air which would be sampled if the system operated in a flat land region. Obviously the lower the radar antenna elevation angle considered, the larger the blockage produced. This is a dilemma difficult to solve because the lower antenna elevations are generally the most valuable for radar precipitation estimates (Joss and Waldvogel, 1990; Smith, 1998).

In the last years many techniques have been studied to diagnose and correct ground clutter and AP echoes. Apart of processor-level techniques for non-coherent and Doppler radars, many others have been developed as post-processing methodologies designed to remove AP or clutter echoes. They are typically based in the analysis of the reflectivity field structure (see, among others, Joss and Lee, 1995; da Silveira and Holt, 1997; Pamment and Conway, 1998; Fulton et al., 1998; Alberoni et al., 2001; Sánchez-Diezma et al., 2001; Steiner and Smith, 2002; Haddad et al., 2004). Other studies aimed at enhancing radar observations with satellite images to remove non-precipitating echoes (Pankiewicz et al., 2001; Michelson and Sunhede, 2004).

More recently, polarimetric techniques have been considered to overcome – among others – AP, clutter and beam blockage problems as described by Ryzhkov and Zrnic (1998), Illingworth (2003) or Ryzhkov et al. (2005). Besides, other contributions oriented to hydrological modelling have merged radar data with raingauge observations making use of geostatistical techniques to clean ground clutter or AP echoes from radar images (see, for example, Todini, 2001; or Wesson and Pegram, 2004).

In this paper a new post-processing technique to decontaminate AP and clutter echoes and to correct for topographical beam blockage single-polarisation radar observations is presented and evaluated. The organisation of the paper is as follows. In Sect. 2 a description of the new correction methodology is given. Section 3 gives an overview of the data sets employed and Sect. 4 introduces the statistical magnitudes used to examine the performance of the method. Section 5 contains four different case studies where the performance of the methodology is discussed. Finally, a summary and conclusions are given in Sect. 6.

## 2 The combined method of clutter and beam blocking correction

The methodology used to correct or reduce ground clutter and radar beam blockage is briefly detailed in this section. A preliminary description of the method can also be found in Fornasiero et al. (2005).

The procedure is applied on the polar volume of reflectivity acquired at short pulse mode – with the Doppler filter active – to a three dimensional dataset. The output is a bi-dimensional map of equivalent reflectivity values, also in polar coordinates azimuth-range, which is finally interpolated into a 2-D Cartesian grid of  $1\text{ km} \times 1\text{ km}$  resolution.

The main idea of the procedure is to choose, for each azimuth-range cell or range-bin, the elevation that is most representative of the reflectivity value for rainfall estimation, and projecting this value in the 2-D grid. These data are then corrected for beam blockage and clutter.

The working scheme consists of three different stages: choice of the scan elevation for each cell through a dynamic antenna elevation map; correction for residual anomalous propagation clutter through a vertical continuity test; and correction for the beam shielding.

### 2.1 Dynamic elevation map

The choice of the dynamic map is based on the search for the minimum antenna elevation simultaneously not affected from climatological clutter, with a beam shielding fraction lower than 0.5, and not contaminated by anomalous propagation clutter (the lower limit of the 3-dB beam should not intercept the ground). This map is built resuming different information: the static map of precipitation-free echoes, the beam blocking rates and the heights, relative to the ground, of the 3-dB beam lower limit. The last two parameters are obtained simulating the beam path through a multi-layer model (Fornasiero et al., 2005, 2006), where the refractivity gradient profile is retrieved from a linear weighted combination of the two nearest (in time) radiosoundings profile, collected by a station that is co-located with the radar. The resulting profile is then interpolated at vertical steps of 0.2 km.

### 2.2 Residual anaprop

The correction for residual anaprop clutter is performed with the method of Alberoni et al. (2001) modified by Fornasiero et al. (2005). This method compares, for each range-bin, the reflectivity value at the dynamic elevation map with those observed at the successive elevation scans at the same location. The basis of this method is founded on the expectation of higher vertical coherence of meteorological signal with respect to ground clutter, which presents stronger vertical gradients. This procedure is sometimes called vertical continuity test and is often included in the corrections performed

**Table 1.** Principal characteristics of San Pietro Capofiume radar.

Model	GPM 500 C		
Site	San Pietro Capofiume		
Latitude (°)	44.655		
Longitude (°)	11.624		
Antenna height (m a.s.l.)	30		
Frequency (GHz)	5.43–5.64		
Peak power (kW)	500		
Pulse width (μs)	0.5	1.5	3.0
Radial resolution (m)	250	500	1000
PRF (Hz)	1200	600	300
Maximum range (km)	125	250	500
Main lobe width (°)	0.9		
Antenna diameter (m)	5		
Antenna gain (dB)	46		
Azimuthal speed (°/s)	20		
Polarisation	Vertical or Horizontal		
Angles of scan strategy (°)	0.5 + k·0.9, with k=0, 1,..., 15		

over reflectivity radar data to obtain rainfall rate estimations (see for example Fulton et al., 1998).

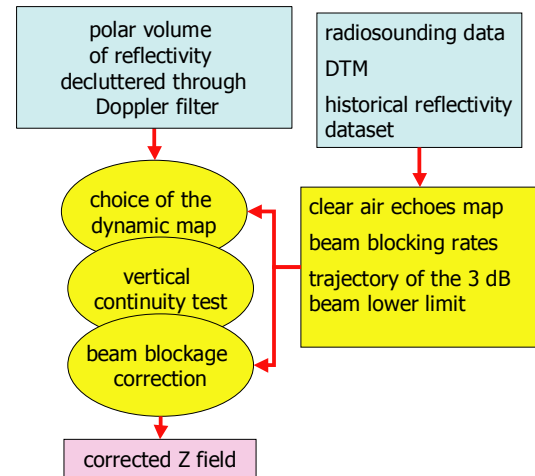
### 2.3 Beam shielding

Finally, the evaluation and correction for beam blocking is performed using a geometric optic approach based in the interception function proposed by Bech et al. (2003). In this study it has been applied the above-mentioned multi-layer model to retrieve the radar beam trajectory instead of assuming a homogeneous vertical refractivity gradient for the whole air-layer. This refinement allows a more detailed description of the radar beam behaviour introducing the possibility to simulate a wider variety of propagation effects such as beam-splitting.

In Fig. 1 is illustrated the complete correction procedure, emphasizing the input and output data and the additional required information. The scheme will be called hereinafter ‘BDA’ (acronym of the three procedures included: Beam blocking, Dynamic map and Anaprop correction).

## 3 Datasets description

To evaluate the efficiency of the described methodology it has been applied to the rainfall rate measured by the C-band Doppler radar of San Pietro Capofiume, located in northern Italy (see Table 1 for details). The radar precipitation estimates have later been compared with raingauge data.



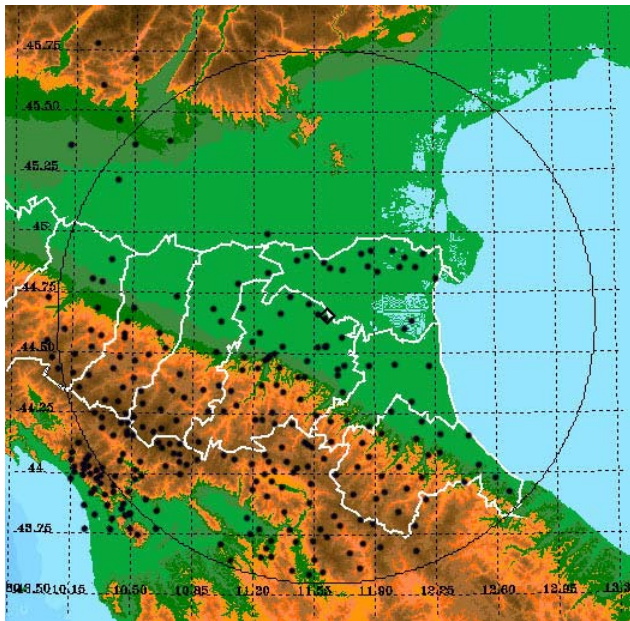
**Fig. 1.** BDA scheme of reflectivity data correction. The two blue rectangles are the input data, the yellow rectangle contains the intermediate products, the yellow circles are the three correction steps leading to the final output, the corrected reflectivity (Z) field.

The reflectivity data used in the analysis are provided every 15 min and collected using a short pulse of 0.5 microseconds and a pulse repetition frequency (PRF) of 1200 Hz. The radar pixels had a resolution of 250 m×0.9° and the measurements extended up to a maximum range of 125 km. As explained in the previous section, after the application of the correction the radar reflectivity is also contained in a bi-dimensional grid azimuth-range where each cell measures 250 m×0.9°. The data are then resampled into a temporal Cartesian grid of 250 m×250 m resolution, assigning the value of each polar cell to the nearest Cartesian cell. Thereafter this Cartesian grid is transformed to another one of lower resolution (1 km×1 km) more adequate to the original radar sampling, using the criterion of the maximum reflectivity value into a box composed by 16 elementary cells.

The reflectivity Z is converted into rainfall rate R using a Z-R power-law relation  $Z=aR^b$ , where the coefficients a and b are those indicated by Joss and Waldvogel (1970) for convective (a=500, b=1.5) and stratiform rain (a=250, b=1.5). Stratiform and convective events were distinguished subjectively by visual inspection on the radar data volumes. For sake of simplicity no attempt has been made to assign different Z-R relationships to different pixels of the same image distinguishing between stratiform and convective precipitation type. This could be done with a classification technique such as those described by Sánchez-Diezma et al. (2001) or Rigo and Llasat (2004).

Finally, the hourly cumulated rainfall amount is obtained through weighted average of the five precipitation intensity values at 00, 15, 30, 45, 60 min before the nominal hour, where the extreme minutes of each hour (00' and 60') weigh half of the others. This is done in order to take into account that half part of the 00' and 60' amounts do correspond to the





**Fig. 2.** Raingauges available for the analysis (black points) and radar coverage (black circle) in the North of Italy.

previous and next hour respectively unlike the other precipitation amounts, which are recorded completely within the nominal hour.

Raingauge data were provided every 60, 30, or 15 min (see Tables 2 and 3). The amounts reported every 30 and 15 min were summed to obtain the corresponding hourly value. Most of the stations were located in the mountainous area and behind the mountains – in the SW part of the radar coverage –, so that beam blockage corrections could be evaluated. The raingauges had tipping-bucket switches and were not warmed. Therefore some records might be mistaken during snow events. As explained later this has been taken into consideration in the analysis.

In Fig. 2 is represented the location of the raingauges and the radar maximum range circle.

#### 4 Evaluation parameters

The comparison between raingauges and radar data has been performed through the visual inspection of rain maps and through a quantitative-statistical analysis based on the calculation of some indices relative to the hourly-cumulated rain and to the event cumulated rain. The indices are: the bias, the root mean squared error, the fractional mean reduction and the fractional variance reduction.

Calling  $R_G$  the cumulated rain rate (over the hour or over the event) produced by the gauges and  $R_R$  that estimated from reflectivity radar data in the corresponding cells, the indices used can be written as (Marzano et al., 2004):

**Table 2.** Raingauges available for the analysis divided by integration time.

	1st type	2nd type	3rd type
Integration time (min)	15	30	60
Number of raingauges	184	176	41
Total raingauges	401		

**Table 3.** Raingauges taken into consideration for the analysis in each case study, detailing the number of observations in mountain blocked areas and in all areas (total).

Number of stations	Apr mount.	Dec mount.	July mount.	Set mount.	July total	Set total
Available	401	401	401	401	401	401
Used	142	135	176	179	221	239

The bias  $\langle \varepsilon_R \rangle$

$$\langle \varepsilon_R \rangle = \langle R_R - R_G \rangle \quad (1)$$

the root mean square error RMSE

$$\text{RMSE} = \sqrt{\langle \varepsilon^2 R \rangle} \quad (2)$$

the fractional mean reduction FMR

$$\text{FMR} = \frac{\langle R_G \rangle - \langle \varepsilon_R \rangle}{\langle R_G \rangle} \quad (3)$$

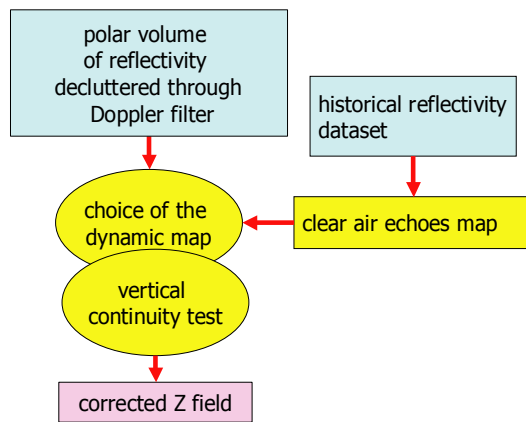
the fractional variance reduction FVR

$$\text{FVR} = \frac{\sigma^2 R_G - \sigma^2 \varepsilon_R}{\sigma^2 R_G} \quad (4)$$

where the angle brackets mean average over time and over the cells.

An efficient correction method should reduce the RMSE towards the “optimum limit” of 0, and it should produce values of FMR and FVR tending to 1. In the analysis of the case studies, the RMSE calculated on the event-cumulated rain, is called RMSECUM. Because of the linear relation between bias and FMR, in the next section, it has been considered only the FMR that is normalized with respect to the mean gauges rainfall field. Moreover, when FMR is higher or lower than 1, the bias is respectively negative or positive.

The evaluation is limited to the raingauges located beyond 20 km from the radar site, to avoid the radar antenna secondary lobes effect. Moreover, two different coverage areas are taken into consideration: the whole 360° angle and that limited between angles 135° and 270°, with respect to the north direction (in this case the minimum range is increased to 40 km). The second area represents the mountainous area, considered to highlight the behaviour of the algorithm in complex orography.



**Fig. 3.** Scheme of reflectivity data correction operative at ARPA-SIM (SA scheme).

## 5 Results and discussion

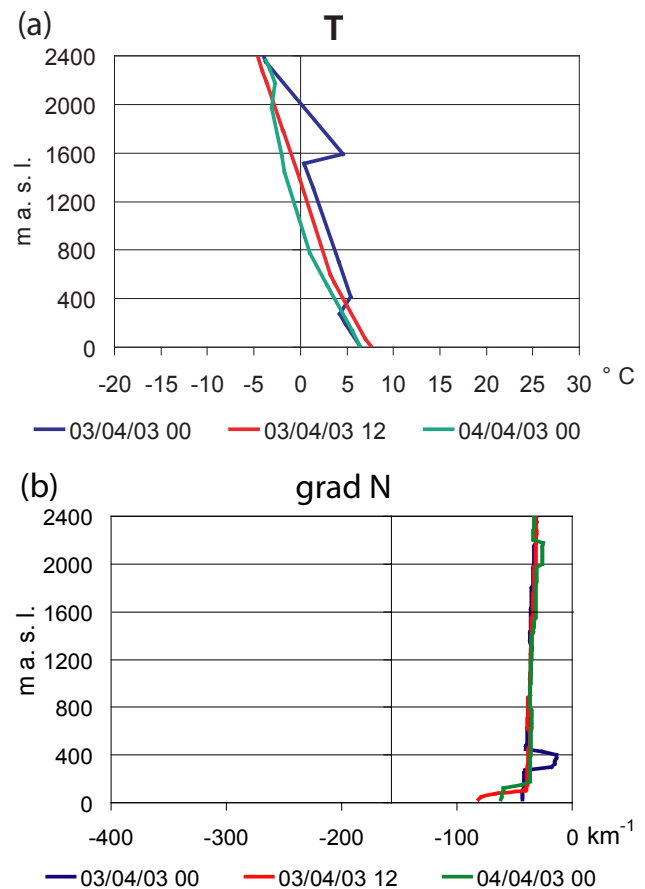
To evaluate the performance of the new BDA correction method it has been compared with another reference method that does not correct the radar data for beam blockage and does not consider either the approximate path of the beam to reduce the clutter.

This reference algorithm of clutter removal is actually operative at ARPA-SIM to correct data of San Pietro Capofiume and Gattatico radars. The algorithm is based on the use of a static clutter filter derived from a climatological map of non-precipitating echoes employed to select the antenna elevation. Besides, the removal of residual anaprop clutter is done applying a vertical continuity test to the radar echo volumetric observations (see Fig. 3). This reference method will be called hereinafter “SA”.

The evaluation of the performance of the algorithms by using raingauges is uncertain at low temperatures when snow is present and tipping-bucket raingauges are not warmed. To address this problem, in the stratiform precipitation cases, comparisons were restricted to raingauges located at least 100 m below the minimum 0° level during the event considered and having cumulated amounts higher than 1 mm. This restriction tries to avoid including snowfall observations mistaken by the gauges though some sleet is likely to be included in the higher radar observations considered. However this approach does not remove all the stations of interest, because raingauges sited in the valleys and beyond the mountains are conserved.

The first two events considered were stratiform cases: one occurred in spring, the other one in winter. In these cases, the analysis was limited to the mountainous and blocked areas, because the propagation conditions were nearly standard and hence no significant changes between the BDA and SA outputs are produced in flat land areas.

The last two cases were convective summer events. In those cases superrefraction and ducting conditions were



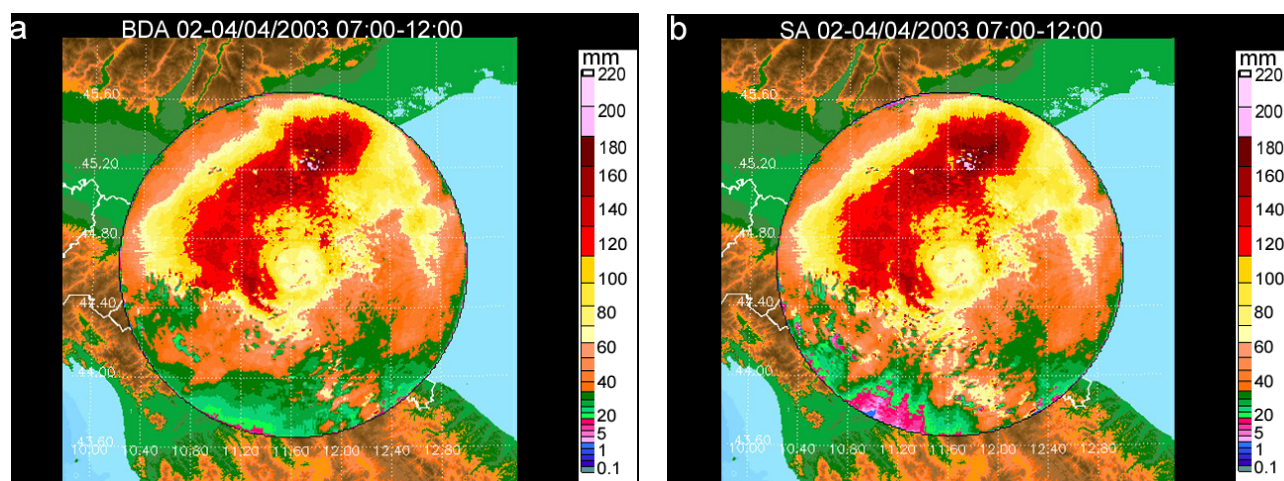
**Fig. 4.** Temperature (a) and refractivity vertical gradient profiles (b) obtained from radiosonde ascents during the 24-h period from 3rd 00:00 Z to 4th 00:00 Z April 2003.

clearly present due to the effect of a surface thermal inversion causing a strong surface duct. Therefore, these two cases are the most interesting to illustrate the effects of BDA methodology.

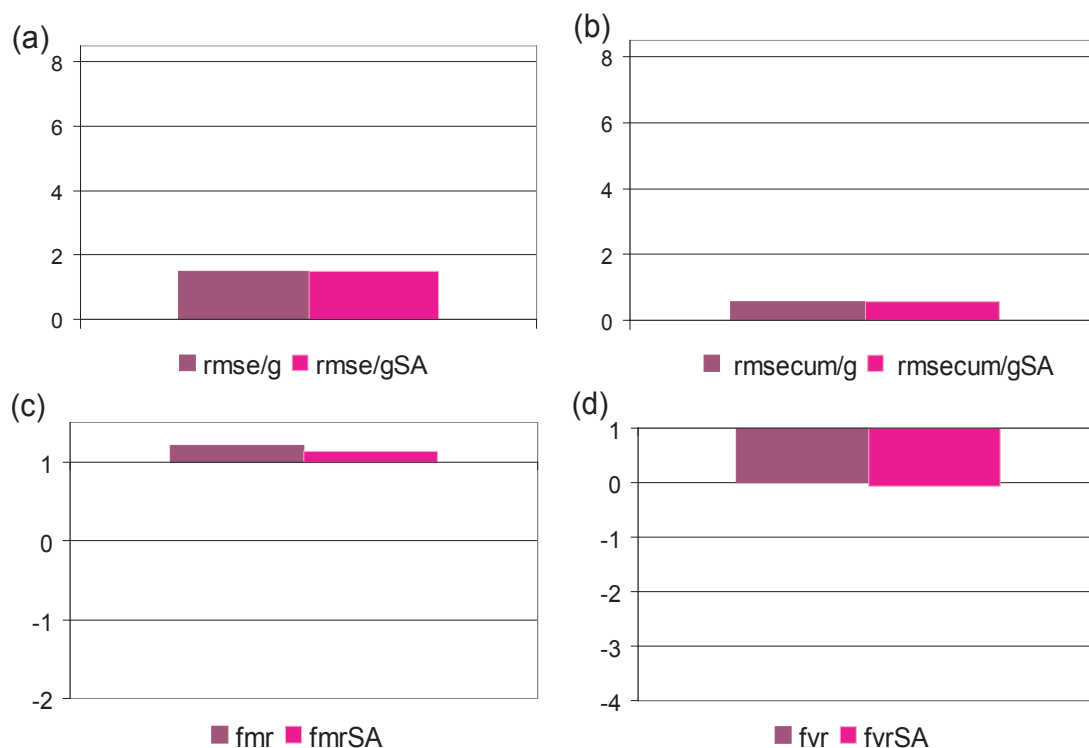
### 5.1 Case study 1: April 2003

The first case presented took place from 2 to 4 April 2003. It was a clear stratiform event with minimum 0°C level at around 1000 m (measured by radiosoundings) and normal refraction conditions, as it is visible in Fig. 4. The cumulated rainfall field, obtained using BDA method, seems to be more coherent and realistic as compared to SA output, because the shielding effect of mountains is reduced (Fig. 5). However, no significant difference is shown by the raingauge-comparison RMSE and FVR indices, and the FMR is even slightly worse (Fig. 6).

One reason for this behaviour is probably the lack of correction for vertical reflectivity variation. In fact, to avoid the mountain shielding effect in BDA method, the elevation is increased and the considered reflectivity values become less



**Fig. 5.** Cumulated rain of the April 2003 event, obtained using BDA (a) and SA (b) methods. In the SA precipitation field it is visible the shielding effect of the mountains; in BDA, after beam blockage correction, the field seems more coherent in the mountainous area and beyond.

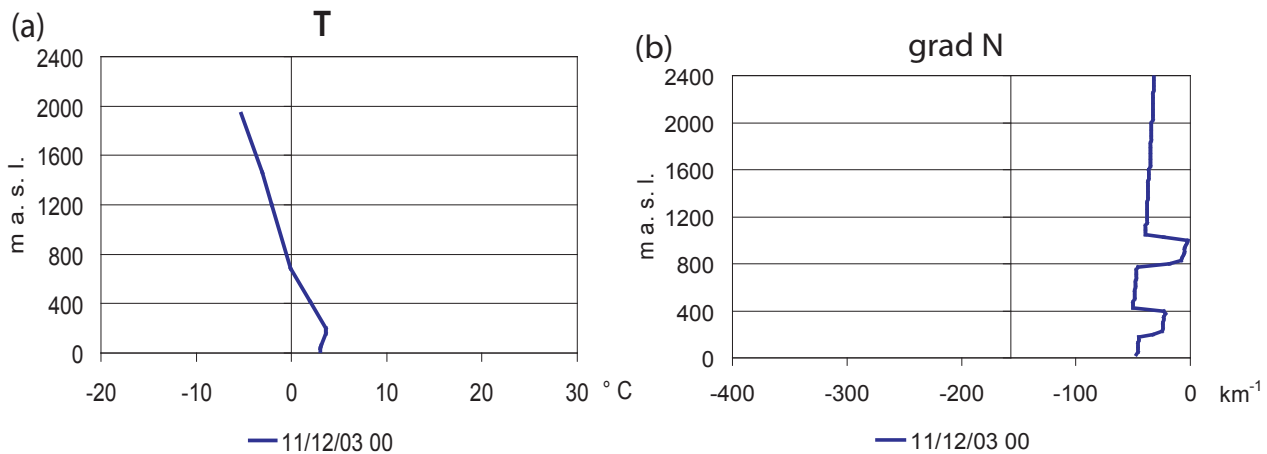


**Fig. 6.** Radar-raingauges comparison indices for the April event showing the performance of the BDA (dark pink) and SA (fuchsia) methods. Only raingauges located below 900 m a.m.s.l. (100 m below the minimum  $0^{\circ}$  level) and in the mountainous sector (azimuth  $[135^{\circ}, 270^{\circ}]$ , range  $>40$  km) are considered. (a) RMSE normalised with respect to the mean gauge rainfall field for the hourly; (b) as a for the event cumulated rain rate; (c) FMR; (d) FVR indices.

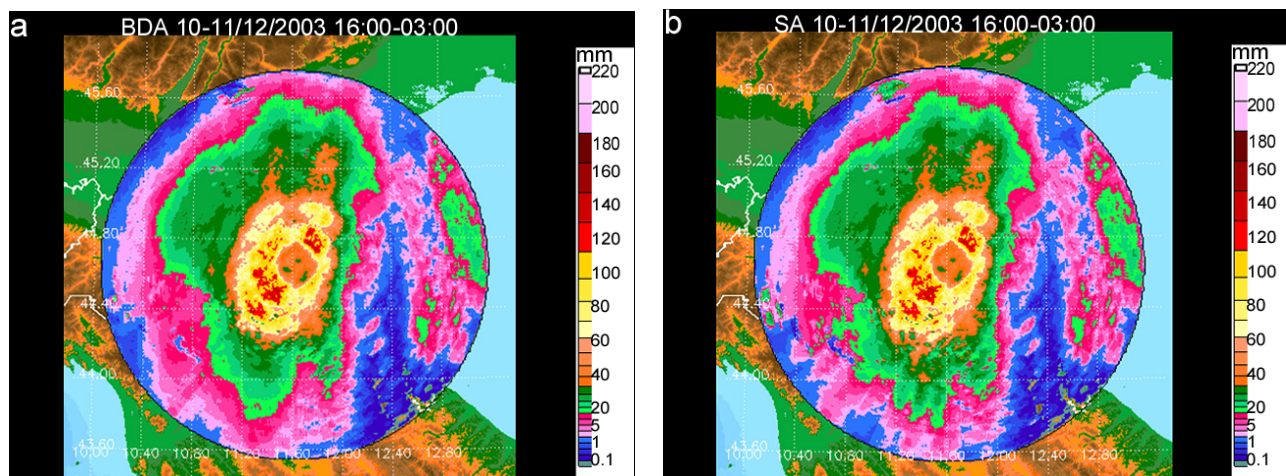
representative of the ground value. So it seems plausible that during this event, at least in some cases, a blocked radar observation nearer the ground might provide a more accurate

rainfall estimate than an unblocked observation aloft because of the presence of snowfall and sleet at higher altitudes.





**Fig. 7.** Temperature (a) and refractivity vertical gradient profile (b) obtained from radiosonde data during the 11th 00:00 Z December 2003.



**Fig. 8.** Cumulated rain of 10–11 December 2003 event, obtained using BDA (a) and SA (b) methods.

## 5.2 Case study 2: December 2003

The second event took place during 10 and 11 December 2003 and the minimum 0°C level was approximately 600 m, while the propagation conditions were normal (Fig. 7). Hence, the raingauges taken into consideration were those up to 500 m.

As in the previous case, the BDA output seems to be more coherent at the visual inspection (Fig. 8). In this case all the scores, calculated in the mountainous sector, were improved when the BDA methodology was applied: the RMSE and RMSECUM were reduced, patent improvement is visible in the FVR, the bias (positive, because FMR is lower than 1) decreased. The considerable reduction in FVR confirms the higher spatial coherence of BDA field beyond the orography obstacles.

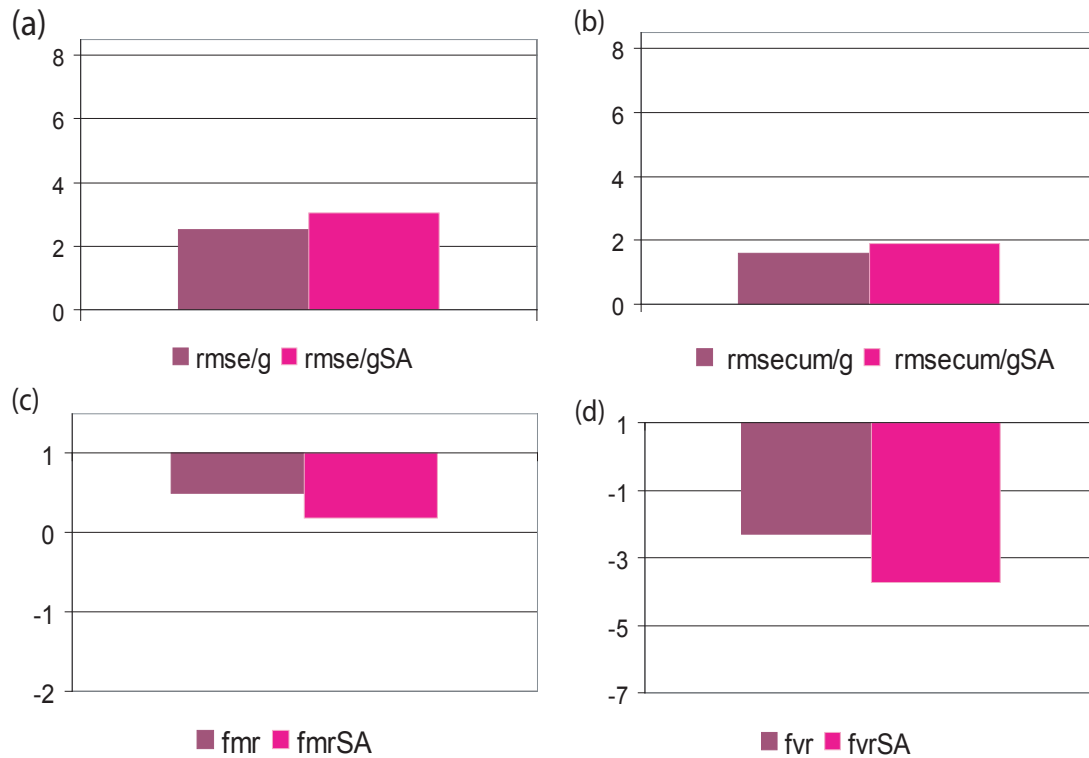
Otherwise, the results, especially in this case, should be carefully considered. In fact, in the centre of the Fig. 8, that represents the event cumulated rainfall rate, a remark-

able concentric increase seems present, that evidences a clear bright band. Therefore, at longer ranges, the precipitation is probably snow and sleet. The bright band contribution is confirmed by the positive bias and by the height of the midnight 0° level (Fig. 9).

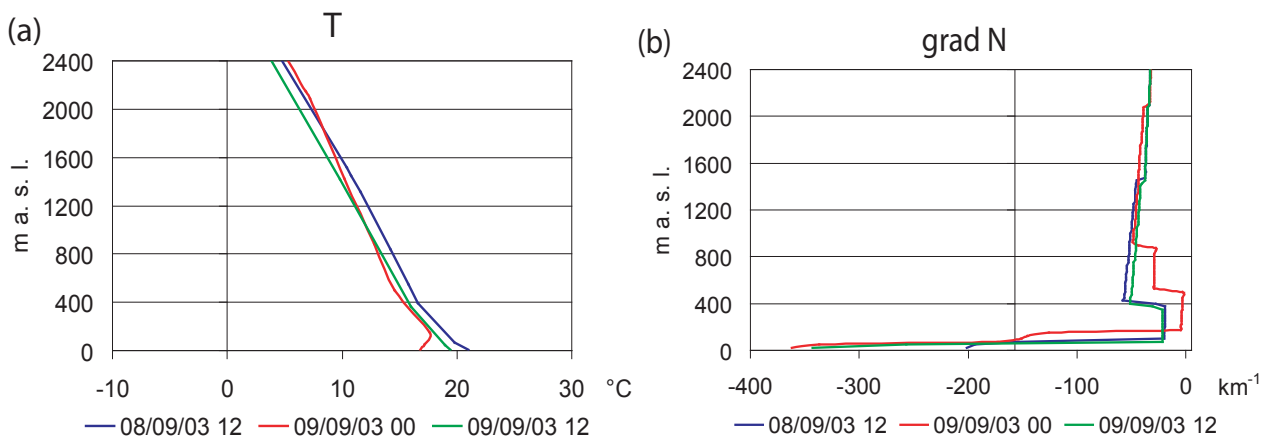
## 5.3 Case study 3: September 2003

The third case took place during 8 and 9 September 2003: the N gradient profiles, retrieved from the radiosoundings representative of the event, show deep superrefraction especially in the second day (moreover at midnight a temperature inversion is present), as pointed out in Fig. 10.

All scores yielded by the BDA method improved SA results (except the FMR that does not provide relevant changes) either in the mountain region (Fig. 12) or in the 360° domain (Fig. 13) presented separately for this case study. The improvement, though still exists as pointed out by the scores, is less evident in the 360° coverage: in fact, dealing with beam



**Fig. 9.** Radar-raingauges comparison indices for the December event. Only the raingauges located below 500 m a.s.l. (100 m below the 0° level at midnight) and in the mountainous sector are considered. (a) RMSE normalised with respect to the mean gauge rainfall field for the hourly; (b) as a for the event cumulated rain rate; (c) FMR; (d) FVR indices.

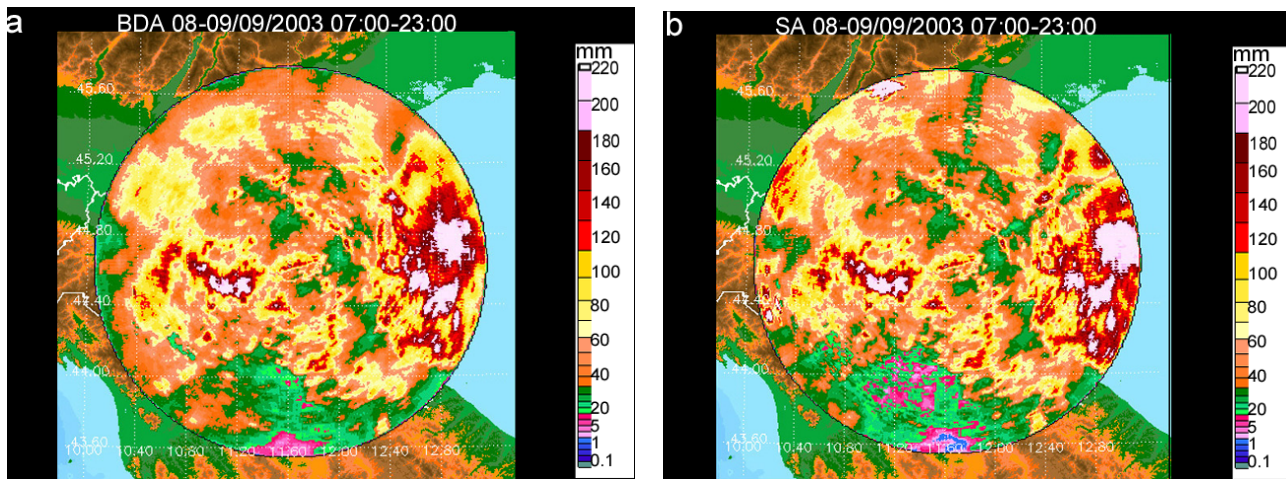


**Fig. 10.** Temperature (a) and refractivity vertical gradient profiles (b) obtained from radiosonde ascents during the 24-h period from 8th 12:00 Z to 9th 12:00 Z September 2003.

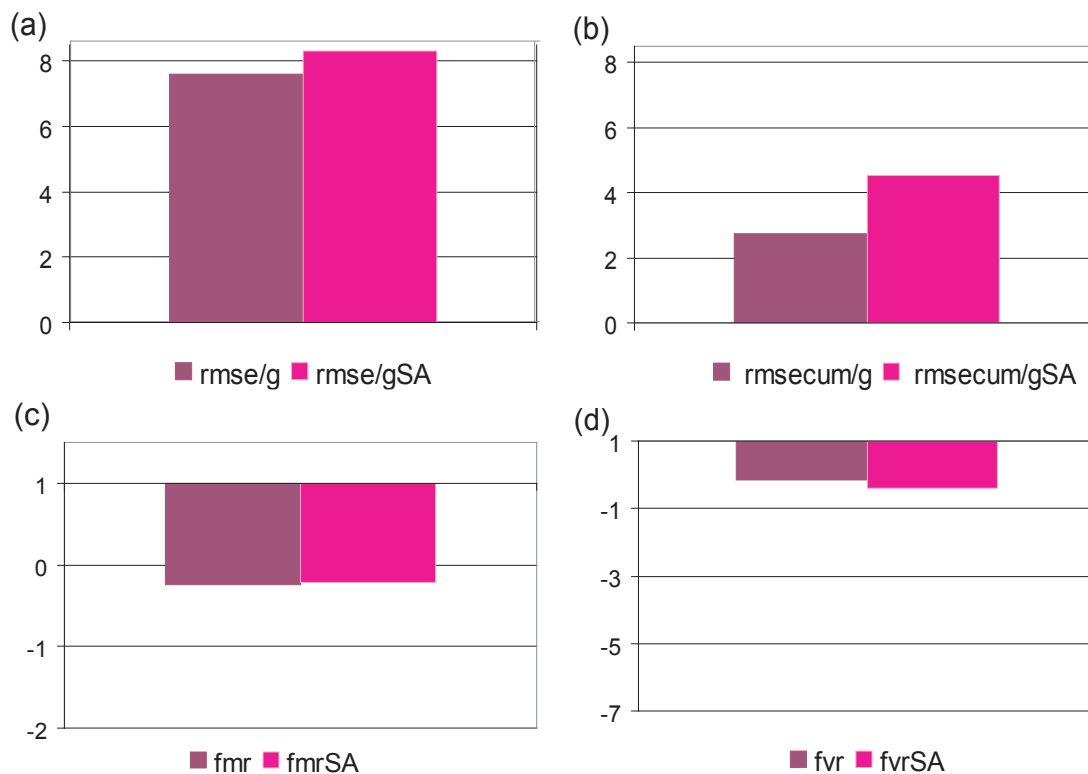
propagation and beam blocking, the BDA method produces the major effect in complex orography areas, and therefore, the scores improvement is smoothed by averaging over the whole radar domain (that includes a large flat land share). In either cases the higher impact is on RMSECUM, that is moreover, two-three times smaller than RMSE. This fact proves the higher usefulness of radar data in the estimate of

cumulated rainfall field, as compared to the instantaneous one.

Finally, the visual inspection of Fig. 11 reveals that the BDA precipitation field appears less affected from orography shielding (note the 170–270° and the 5–10° sectors of the image) than that obtained with the reference method SA.



**Fig. 11.** Cumulated rain of September 2003 event, obtained using BDA (left) and SA (right) methods. In the SA output, sectors  $5^{\circ}$ – $10^{\circ}$  and  $170^{\circ}$ – $270^{\circ}$  show the shielding effect of mountains, much less apparent in BDA field after the beam blockage correction was applied.



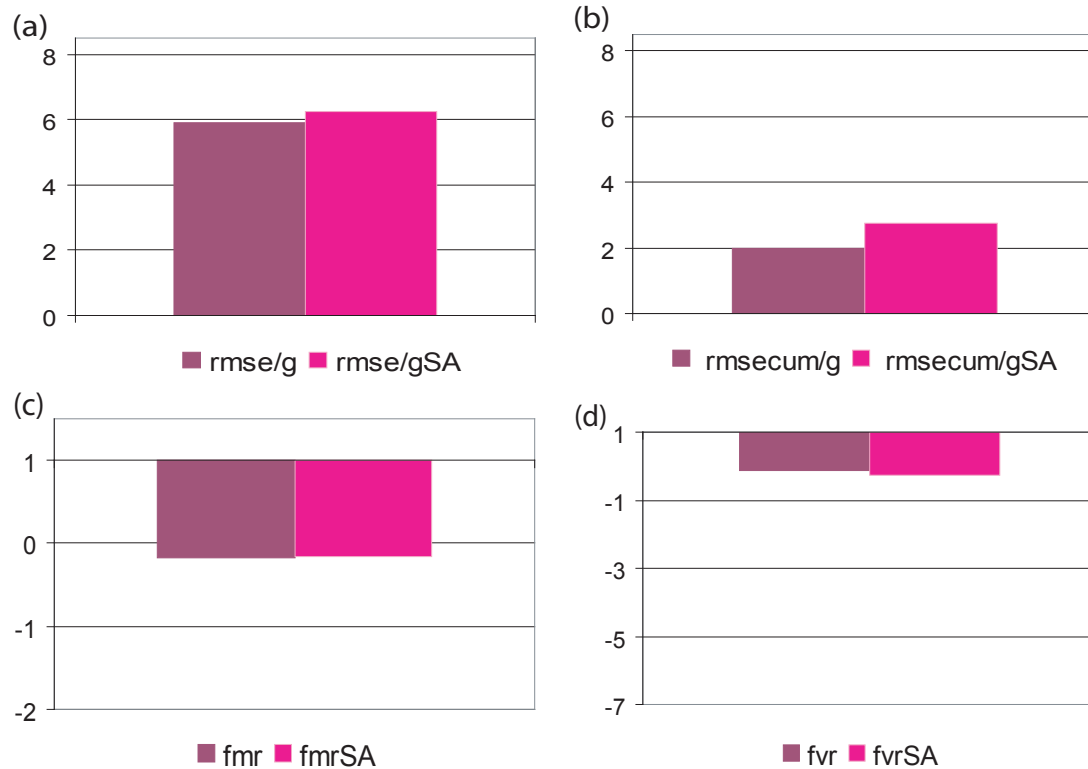
**Fig. 12.** Radar-raingauges comparison indices for the September event, calculated in the mountainous sector. (a) RMSE normalised with respect to the mean gauge rainfall field for the hourly; (b) as a for the event cumulated rain rate; (c) FMR; (d) FVR indices.

#### 5.4 Case study 4: July 2003

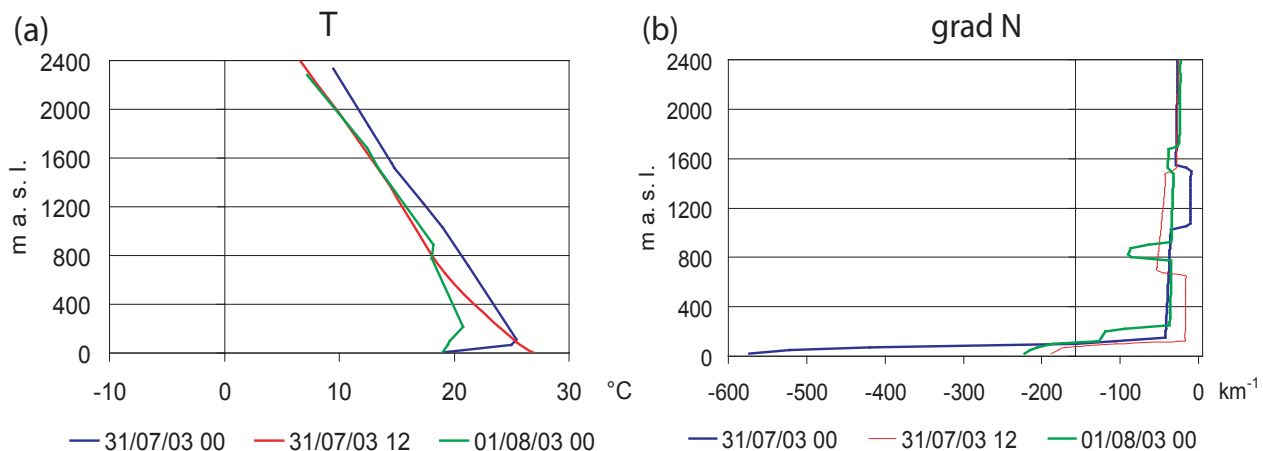
In 31 July 2003 the results were not so evident. The propagation conditions – showing an intense surface duct – are shown in Fig. 14 and the rainfall fields in Fig. 15. In particular, in the mountainous area the scores seemed to be worse (Fig. 16),

more specifically both the RMSE and FMR calculated from hourly values. A closer analysis revealed substantial temporal variations throughout the event of the performance of the BDA algorithm.

Comparing the 17 p.m. hourly radar rainfall field obtained through BDA and SA methods with the rain gauges



**Fig. 13.** Radar-raingauges comparison indices for the September event, calculated in the 360° sector and beyond 20 km from radar site. (a) RMSE normalised with respect to the mean gauge rainfall field for the hourly; (b) as a for the event cumulated rain rate; (c) FMR; (d) FVR indices.



**Fig. 14.** Temperature (a) and refractivity vertical gradient profiles (b) obtained from radiosonde ascents during the 24-h period from 31st 00:00 Z July to 1st 00:00 Z August 2003.

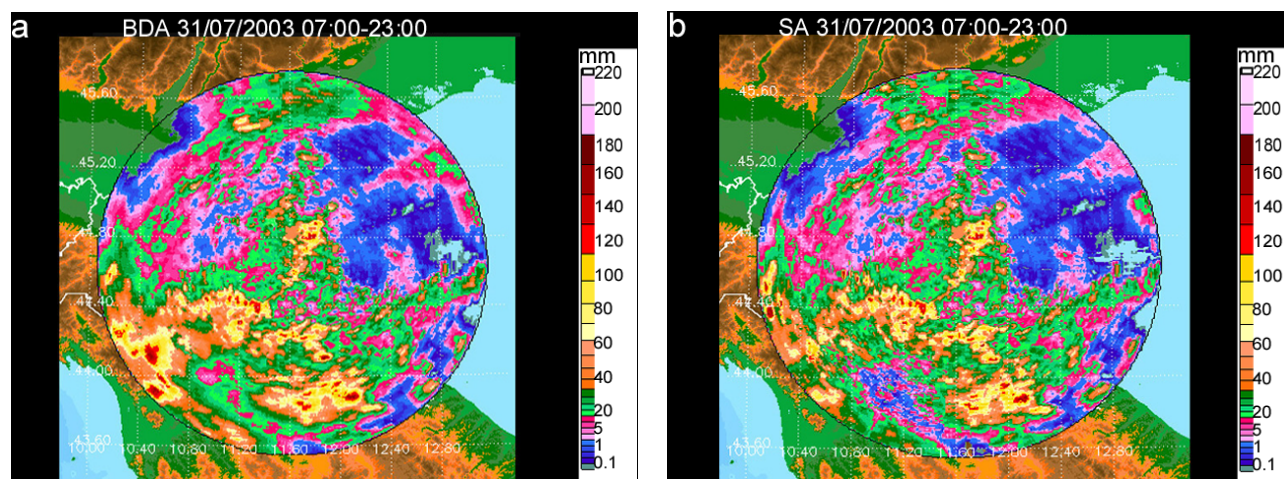
measurements, in the sector around the green line (−179° azimuth) shown by Fig. 17, it can be observed an overestimation, more evident in the new method. However, in that moment raingauges did not measure precipitation.

This was an unexpected result because at this time, due to anaprop risk, the BDA scheme chose nearly always the

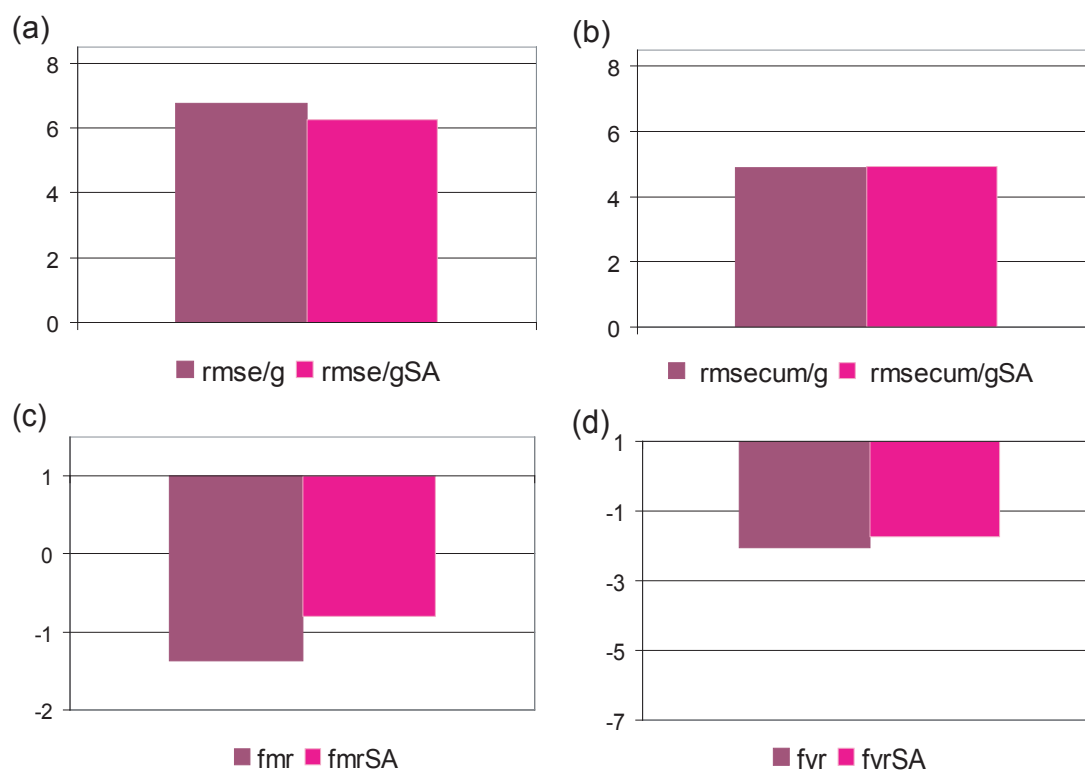
second elevation scan, while in SA the first was mostly considered (see the right panel of Fig. 18), and they were also expected lower reflectivity values.

The reason is probably another one: as it is visible at −179° RHI (left panel of Fig. 18), overhanging precipitation is present and it is intercepted more widely from second





**Fig. 15.** Cumulated rain of July 2003 event, obtained using BDA (left) and SA (right) methods.



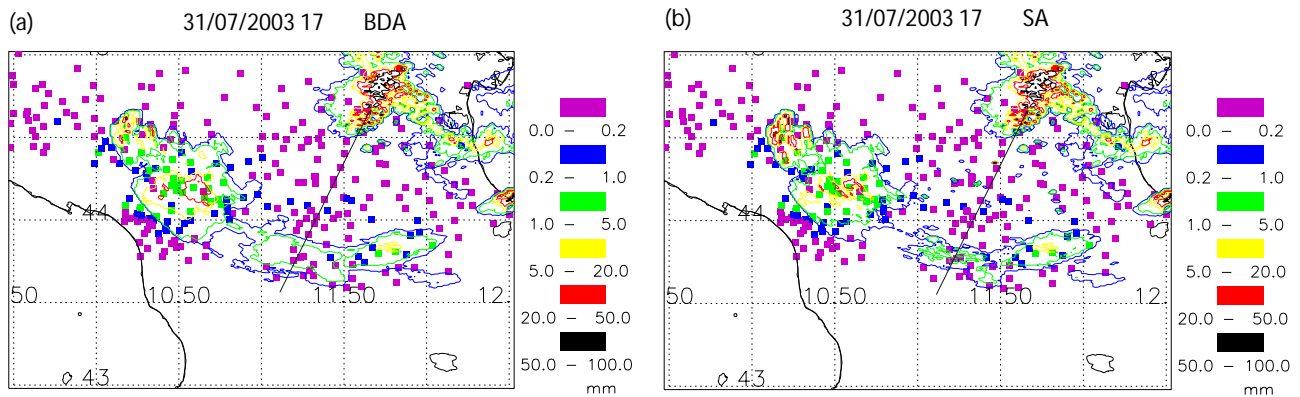
**Fig. 16.** Radar-raingauges comparison indices for the July event, calculated in the mountainous sector. (a) RMSE normalised with respect to the mean gauge rainfall field for the hourly; (b) as a for the event cumulated rain rate; (c) FMR; (d) FVR indices.

elevation scan than from the first one. Therefore, a method to recognise such features could be very helpful in this case.

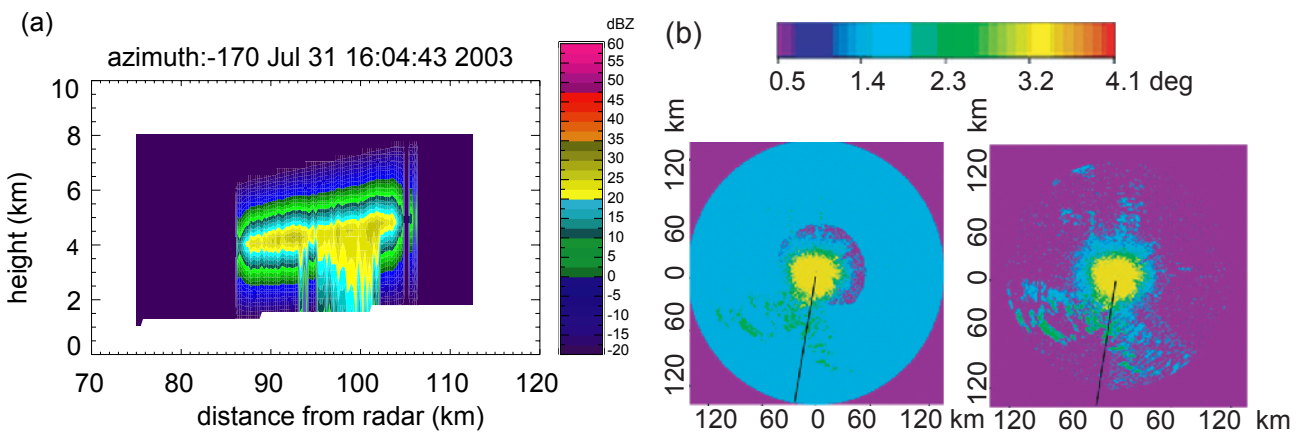
Extending the analysis to the whole radar domain, and focalising it on the evening hours, Fig. 19, the indices become better in BDA with respect to SA output (except the FMR that is directly related to the bias). It seems clear that in the evening, when the anaprop contribution is stronger, the new methodology works well and the effects are visible too in flat

land areas. The improvement is partially due to the use of radiosounding information and to the modelling of the 3-dB beam lower limit path that conditions the elevation choice. This thesis is confirmed, in the same figure, by the yellow bars which represent the indices calculated on the output of a method that works analogous to BDA, except for assuming standard propagation to reproduce the beam path.





**Fig. 17.** Comparison between rainfall rates measured by the radar and by the raingauges at 31 July 2003 17:00 Z and using BDA (a) and SA (b) algorithms. The green line represents the RHI section illustrated in Fig. 18. In the area surrounding this line BDA produces precipitation where raingauges measure 0 mm. SA output is less affected from this problem.



**Fig. 18.** Radar observations recorded on the 31 July 2003 16:04 Z (a): RHI at azimuth  $-170^\circ$ . (b): maps of antenna elevations chosen to retrieve rainfall rate from reflectivity data, for BDA and SA correction methods. The black line represents the RHI section. In the BDA map it is chosen the second elevation, nearly everywhere, which seems to intercept overhanging precipitation as is evidenced in the Fig. 17 from the comparison with the raingauges.

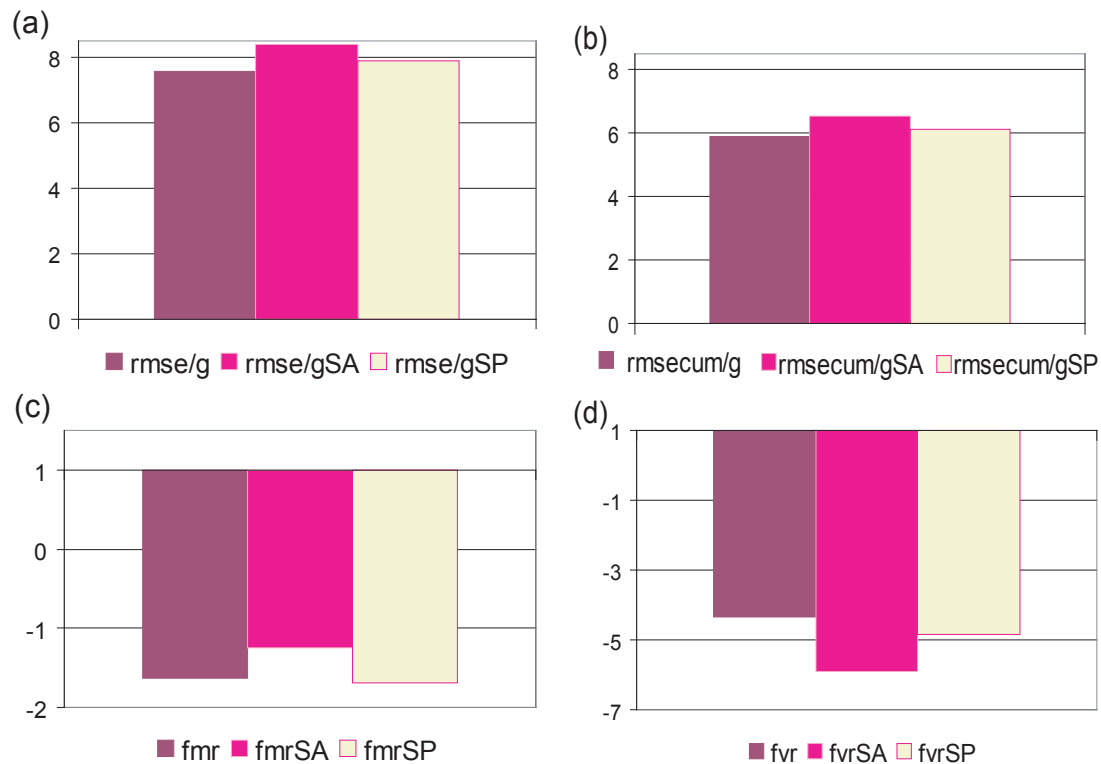
## 6 Conclusions

A post-processing operational method to correct single-polarisation radar reflectivity data for ground clutter and beam blocking has been evaluated in this paper. The method, called “BDA”, has among its main characteristics the detailed analysis of the beam shape and trajectory to build the correction and the use of radiosounding information as input data in the calculation of the radar beam path.

The advantages and limitations of the BDA method have been studied through the comparison with another previous method called “SA” based on static map clutter elimination and anaprop removal, which neglects beam blocking errors and assumes standard refraction conditions for the radar beam. The two methods have been applied to four case studies that took place at different seasons in the Po Valley (Italy). The analysis of the case studies was based on two tools: vi-

sual inspection and calculation of some statistic indices of comparison between the radar rainfall field and raingauge measures (FMR, FVR, RMSE).

In summary, through a comparison with raingauges the proposed BDA shows a tendency of the BDA method to reduce the RMSE and the variance, and a growth of the bias (except for the December event). The choice of higher antenna elevation measurements of BDA with respect to SA and the lack of Vertical Profile of Reflectivity (VPR) correction, is probably the reason of this bias increment. The improvement in the other indices, mainly in FVR, proves the higher coherency of the field obtained using BDA, confirmed too from visual inspection of cumulated rainfall fields. It is also expected that, once removed the bias through adequate VPR correction the new method works consistently better than the older one. Moreover, taking into account realistic propagation conditions, this combined methodology is



**Fig. 19.** Radar-raingauges comparison indices for the July event, calculated in the 360° sector beyond 20 km from radar site and during the evening (from 17:00 Z to 23:00 Z) performance of the BDA (dark pink) and SA (fuchsia) methods. In yellow are the indices calculated on the output of a correction method analogous to BDA but considering standard propagation conditions during the event (SP). **(a)** RMSE normalised with respect to the mean gauge rainfall field for the hourly; **(b)** as a for the event cumulated rain rate; **(c)** FMR; **(d)** FVR indices.

efficient to avoid and remove anomalous propagation echoes. Finally, it should be taken into account the probability of overhanging rain, mainly during summer, as shown by one of the case studies. This could be done considering for example the Pohjola and Koistinen (2004) method to recognise such features through analysis of the VPR shape.

The version of the algorithm studied in the present work assumes as input the atmospheric data acquired in the 12 hours before and after the time of interest. This is a constraint for real time applications. A possible solution is the use of NWP model forecasts of the refractivity field tendency (Bech et al., 2004), which could be complemented with the introduction of anaprop statistics in the simulation of the refractivity profile. Otherwise, it has been observed that the use of a detailed beam description produces the major effects in anaprop cases: a possible choice could be to use standard propagation or simply the average refractivity gradient, when anaprop has low probability, or when NWP model data are not reliable. After including these modifications, taking into account that the proposed methodology is not computationally expensive, this procedure is well suited for real-time application.

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