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# Analysis of the circulation and shelf-slope exchanges in the continental margin of the northwestern Mediterranean

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**Abstract.** In this paper, we present the results from a high horizontal resolution numerical simulation of the northwestern Mediterranean using a z-level, non-hydrostatic, primitive equation ocean model (DieCAST). The high resolution allows an accurate representation of the submarine canyons that presides in the region. The model is one-way coupled to a large scale model of the Mediterranean Sea through open boundaries and uses the atmospheric forcing fields provided in terms of HIRLAM outputs by the Spanish National Institute of Meteorology. Results show that the model can successfully reproduce the complex general circulation characteristics of the area, including the modifications induced by canyons in their vicinity and other phenomena observed such as instabilities and coastal trapped waves. The sea surface temperature is similar to satellite observations except that simulated temperatures are slightly warmer near the coast than observations and colder near the open boundaries. An important topic of this work is the computation of the shelf-slope exchanges, which are able to renew shelf waters in a few months.

## 1 Introduction

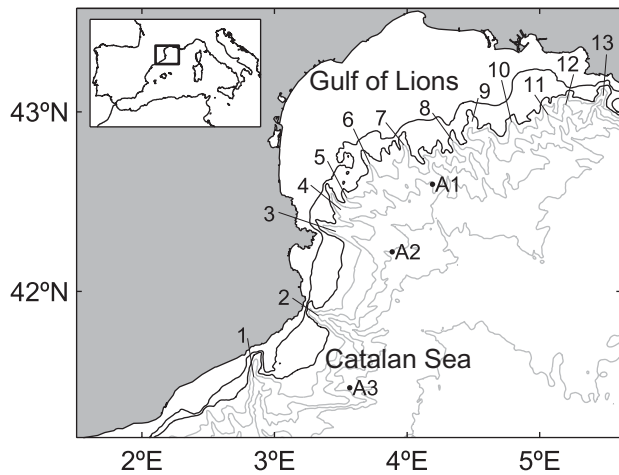
Shelf-ocean interactions are receiving increasing attention because of their enormous importance in aspects such as climate change, global biogeochemical cycles, marine ecosystems, resource exploitation and other human related activities. Much of the present understanding about shelf-slope exchanges derives from processes oriented studies carried out since the late 70s. It is not until recently that the oceanographic community has developed reliable tools to provide nowcasts and forecasts of marine environmental conditions, in support of a variety of activities at sea. This endeavor

has become possible both by the improvement of operational systems able to incorporate and analyze data in real time and by the progress in modeling and computing capabilities. An example of these efforts in providing operational information is the European project MFSTEP (Mediterranean Forecasting System Toward Environmental Predictions) which is focused on the Mediterranean Sea.

Observational and modeling based analysis of the Mediterranean reveal a complex and variable circulation that is strongly affected by frontal dynamics through mesoscale features (Millot, 1999; Send et al., 1999; Demirov and Pinardi, 2002). This variability is the result of interactions at basin, sub-basin and local scales that results in intense mesoscale variability. In the Mediterranean Sea, mesoscale and even sub-mesoscale and local processes have profound impact on ocean dynamics and, consequently, in ocean geochemistry and ecosystems (e.g. Levy et al., 2001). The necessity of predictive tools resolving processes at these scales encompasses an integral view of the ocean system and the use of nested schemes that allow for downscaling from general to local conditions.

The northwestern Mediterranean Sea is one of the areas of interest of the European cooperative oceanography project MFSTEP where nested operational models have been implemented. This region comprises the shelves and slopes of the Gulf of Lions and the Catalan Sea (Fig. 1). Bottom topography in the area is rather complex, varying from a broad shelf in the Gulf of Lions to a rather narrow continental margin in the Catalan Sea. The slope is indented by a series of canyons that cut across about 60% of the slope area (between 200 and 1000 m isobaths). Submarine canyons in the Gulf of Lions occur as deep and short indentations of the shelf edge separated by narrow open slopes. Conversely, the Palamós and Blanes canyons in the Catalan shelf are more prominent and expand to the proximity of the coast.

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**Fig. 1.** Model domain of the northwestern Mediterranean and bathymetry. The 100 and 200 m isobaths are plotted with black lines, and the 500, 1000, 1500, 2000 and 2500 m isobaths with gray lines. Black points represent the location of Argo profiles. Numbers indicate prominent submarine canyons: Blanes (1), Palamós (2), Cap de Creus (3), Lacaze-Duthiers (4), Pruyot (5), Boucart (6), l’Hérault (7), Aigüesmortes (8), la Petite Rhône (9), la Grand Rhône (10), Marseille (11), Planier (12), and Cassidaigne (13).

The major aspects of the water mass structure and flow dynamics of this region are fairly well known and have been thoroughly described by several authors (e.g. Millot, 1999). The general circulation is dominated by a cyclonic along-slope density front, called the Northern Current, flowing from northeast to southwest (Font et al., 1988). Previous studies have found that this current is characterized by relevant spatial and temporal variability, with a large variety of mesoscale features (such as meanders and instabilities) strongly interacting with the basin-scale circulation (La Violette et al., 1990). These instabilities have been observed to evolve into very energetic filaments and shedding eddy-like structures (Wang et al., 1988; Tintoré et al., 1990). In addition, very intense ocean-atmosphere interaction mechanisms, characterized by strong episodes of northerly winds, have been described in this area (Estournel et al., 2003).

The complex topography and the strong dynamical processes that characterize the Northwestern Mediterranean provide a good opportunity to analyze the effect of physical transport processes on the shelf-slope exchanges of water, solutes and particles. In this context, the objectives of the present paper are twofold: (1) to validate the ocean circulation model implemented in the continental margin of the northwestern Mediterranean in the basis of previous knowledge and available data, and (2) to provide estimations of the shelf-slope exchanges resulting from the interaction between the Northern Current and the topographic irregularities under real forcing conditions.

## 2 Model description

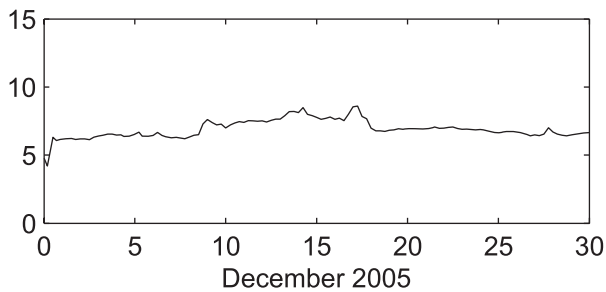
For the purpose of this study, the Dietrich/Center for Air-Sea Technology (DieCAST) ocean model was run for the northwestern Mediterranean at high horizontal resolution following a downscaling approach from a basin-scale model of the Mediterranean Sea. DieCAST is a z-level, non-hydrostatic, primitive equation ocean model (Dietrich, 1997; Dietrich and Lin, 2002), which provides high computational accuracy and low numerical dissipation and dispersion (Dietrich et al., 2004). The model uses a rigid lid approximation, with a sea surface pressure formulation, and the Boussinesq approximation. The governing equations are solved using a blend of collocated and staggered grid structures (Arakawa A and C grids) and fourth-order-accurate control volume approximation for the horizontal pressure gradient and advection terms, except in zones adjacent to boundaries where conventional second order accuracy is used (Sanderson, 1998; Sanderson and Brassington, 1998).

The hydrostatic approximation may become inaccurate for features whose horizontal scale is not large compared to their vertical scale. These features may have significant vertical acceleration compared to the buoyancy force. Rapid changes due to steep bathymetric features like submarine canyons in the presence of strong flows may cause vertical acceleration, which violates the hydrostatic approximation, as shown by Tseng et al. (2005) in the Monterey Submarine Canyon. Thus we use a non-hydrostatic model where the solution of the non-hydrostatic equation is obtained iteratively from the hydrostatic solution, as described in more detail in (Dietrich and Lin, 2002).

Simulations are run with horizontal viscosity and diffusivity coefficients set to a constant value of  $10 \text{ m}^2/\text{s}$ . For the vertical viscosity and diffusivity, a variable formulation which includes the Richardson number is used, as described and used in Staneva et al. (2001), with background values set at near-molecular values ( $10^{-6}$  and  $2 \cdot 10^{-7} \text{ m}^2/\text{s}$  respectively). The bottom dissipation is represented by a conventional nonlinear bottom drag with a coefficient of  $2 \cdot 10^{-3}$ . At the lateral walls we assume free-slip boundary condition. The DieCAST model has been recently used to study the regional circulation in the vicinity of the Monterey Submarine Canyon (Haney et al., 2001; Tseng et al., 2005) and to study the mesoscale, seasonal and interannual variability in the Mediterranean Sea (Fernández et al., 2005).

### 2.1 Model domain

The model domain includes the continental margins of the Gulf of Lions and the Catalan Sea (northwestern Mediterranean) extending from  $41.2^\circ \text{ N}$  to  $43.6^\circ \text{ N}$  in latitude and from  $1.5^\circ \text{ E}$  to  $5.6^\circ \text{ E}$  in longitude (Fig. 1). The model horizontal resolution is  $0.75'$ , which is equivalent to 1389 m in the meridional direction and up to 1036 m in the zonal direction. Note that the Coriolis parameter  $f$  is variable



**Fig. 2.** Time evolution of the volume averaged kinetic energy ( $10^8$  J)

and, in consequence, the resolution in the zonal direction is not constant (although the variation is not significant). The bathymetry on the model grid is derived from the bathymetric chart of the northwestern Mediterranean (<http://www.icm.csic.es/geo/gma/MCB>) by a bilinear interpolation without filtering or smoothing. In the vertical, the model grid has 30 non uniform vertical levels concentrated towards the surface with the top layer thickness being 10 m, and increasing smoothly to 300 m at the bottom.

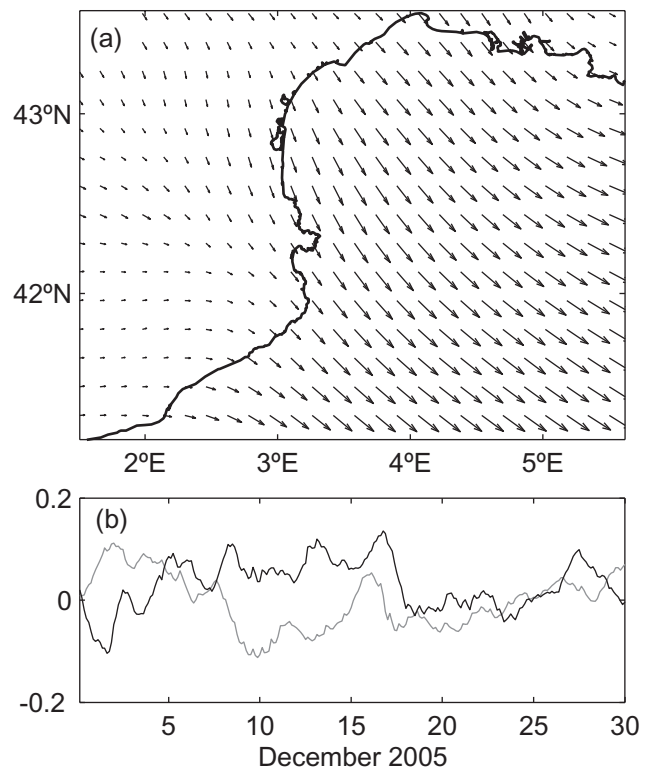
## 2.2 Initialization and boundaries

The model is one-way coupled to a large scale global Mediterranean model through the open boundaries of the domain. The open boundaries allow perturbations generated inside the computational domain to leave it without deterioration of the inner model solution and allows physically important external information to advect inward using a pure up-wind advective scheme (Tseng et al., 2005). Results from the global Mediterranean model (Tonani et al., 2006<sup>1</sup>) are used to force the two lateral open boundaries (south and east) and all quantities are interpolated from daily global Mediterranean model output and updated every time step (150 s). The scale relation between both models is 1/10, taking into account that the resolution of global Mediterranean model is  $1/8^\circ$ .

The model runs for 30 days between 1 December 2005 00:00:00 UTC and 31 December 2005 00:00:00 UTC. This period has been selected to demonstrate the feasibility and product quality of the model system. In addition, a strong episode of northerly wind took place during that period as it is shown in next section. Therefore, December 2005 was an adequate period to study the interaction between the different dynamical processes that characterize the area and to estimate the associated shelf-slope exchanges.

The global Mediterranean model outputs of horizontal velocity, temperature and salinity on 1 December 2005 are used to initialize the model by bilinear interpolation in the hor-

<sup>1</sup>Tonani, M., Pinardi, N., Dobricic, S., and Adani, M.: A High Resolution Free Surface Model on the Mediterranean Sea, *Ocean Sci.*, to be submitted, 2006.



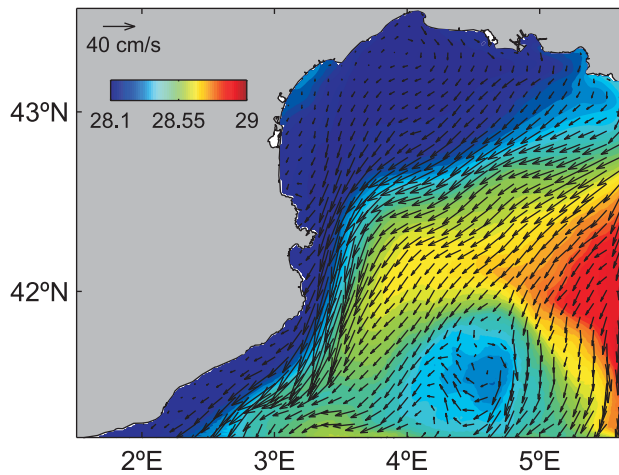
**Fig. 3.** (a) Spatial pattern of the most relevant EOF from the time variability of wind forcing at 10 m. (b) Corresponding amplitude functions. The gray line represent the amplitude function for the u-component of the wind forcing, and the black line is the amplitude function for the v-component.

izontal grid and linear interpolation in the vertical levels. These fields represent daily averaged outputs from the global Mediterranean model and, in consequence, they should be almost in balance. The full dynamical balance is achieved after an adjustment phase where the flow field is subject to inertial oscillations. The duration of this phase is very short (a few hours) and does not change the initial mass field (Fig. 2).

## 2.3 Atmospheric forcing

At the sea surface, the model is driven by the wind stress and surface fluxes (heat and water) provided directly by the Spanish National Institute of Meteorology in terms of 3 hourly fields of the High Resolution Limited Area Model (HIRLAM, <http://www.hirlam.org>). The horizontal resolution of wind forcing at 10 m above the surface and net heat and water fluxes is  $0.16^\circ$ . There is no relaxation of the surface temperature or salinity.

To provide a general picture of the evolution of wind forcing at 10 m, temporal variance complex EOF modes and corresponding amplitude functions are computed. The spatial pattern and its corresponding amplitude functions of the most relevant EOF are shown in Fig. 3a. This pattern accounts



**Fig. 4.** Time-averaged horizontal velocity and density ( $\sigma_t$  units) at 5 m depth. Vectors are plotted at every eighth grid points.

for 64.4% of the total variance of the data, while the second mode (not shown) accounts for 11.3%. The spatial distribution is characterized by a northwesterly wind blowing in the entire basin, with maximum speeds offshore.

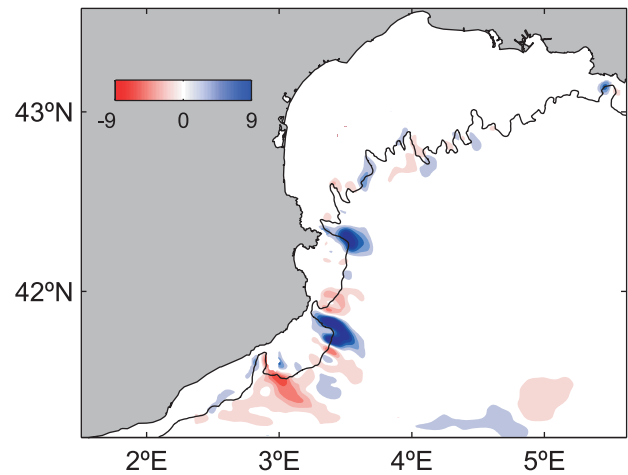
The amplitude functions describe the dynamics of the spatial modes (Fig. 3b). During the first days the wind blows from the southwest and it progressively rotates clockwise to the northeast increasing the intensity. Vector velocities exceeding 12 m/s were registered between 8 and 15 December 2005, rotating to the northwest at the end of this period. A secondary peak from the northwest was observed in last days of December.

### 3 Results and discussion

#### 3.1 General circulation

A temporal average of surface circulation over the integration period based on 6-hourly model outputs reproduces the characteristic features of the circulation in northwestern Mediterranean Sea. Along the shelf break, the flow is constrained to follow the topography and strong currents (>40 cm/s) develop cyclonically flowing from the Gulf of Lions to the Catalan Sea (Fig. 4). This feature is consistent with descriptions of the Northern Current although current velocities are above reported values of 20–30 cm/s (García-Ladona et al., 1996; Send et al., 1999; Lapouyade and Durrieu de Madron, 2001). However it should be noted that, as mentioned above, strong northeasterly winds were experienced during a significant part of the analyzed period.

Another conspicuous structure is the presence of an anticyclonic eddy to the south of the model domain (between 4° E and 5° E). Anticyclonic eddy formation appears to be favored between adjacent coastal upwelling cells on the Gulf of Lions



**Fig. 5.** Time-averaged vertical velocity at a depth of 95 m. The units are m/day. The 200 m isobath is plotted with a black line.

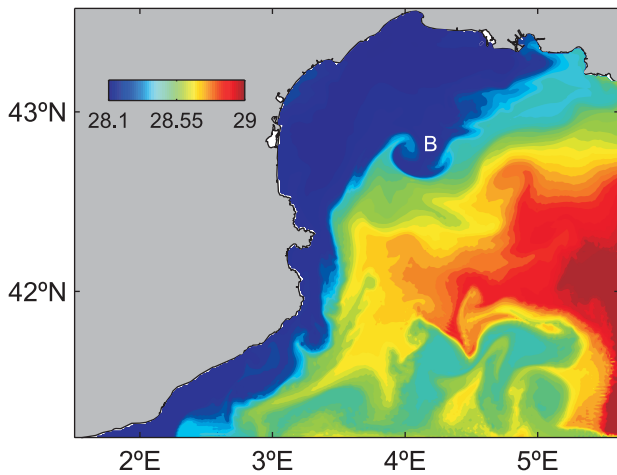
under Mistral forcing (Millot, 1990), which was predominant during the model integration period, and the structure tends to be rather persistent. Indeed, Pascual et al. (2002) tracked a similar structure in that position from middle September 1998 to middle March 1999 using a combination of in situ and satellite data. Onken et al. (2005)<sup>2</sup> also describe a strong anticyclonic vortex at about 41° N exhibiting maximum surface flow speeds of up to 65 cm/s, on fall 2002.

Further agreement concerns the impact of topographic irregularities on the Northern Current. It is well known that submarine canyons affect the spatial pattern of regional vertical motions (Hickey, 1997), particularly in subsurface layers. Figure 5 shows the temporal average of vertical velocity at 95 m depth. A close relation between vertical movement and canyons is evidenced with downwelling on the upstream walls of the canyons and upward motions on the downstream edges. This antisymmetrical structure of vertical motions was previously observed by Pascual et al. (2004) in the Palamós Canyon and simulated by Jordi et al. (2005a) in the same canyon and by Arduin et al. (1999) in the Blanes Canyon.

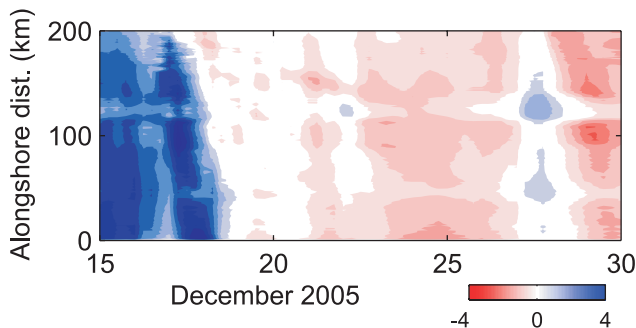
It is important to comment that the Northern Current and the anticyclonic eddy to the south of the domain were already present in the initial condition fields from the global Mediterranean model. However, our model allows the generation and propagation of small-scale structures associated with the Northern Current as we report in the next section. The anticyclonic eddy is reinforced and moves to the southwest during the model run (not shown), probably by the action of the northerly winds. Conversely, the vertical motions

<sup>2</sup>Onken, R., Tintoré, J., Fernández, V., Vizoso, G., Basterretxea, G., and Haley, P.: A forecast system of the Balearic Sea, *J. Marine Syst.*, submitted, 2005.





**Fig. 6.** Surface density ( $\sigma_t$  units) for 23 December 2005. A filament with a mushroom-like structure is indicated with B.

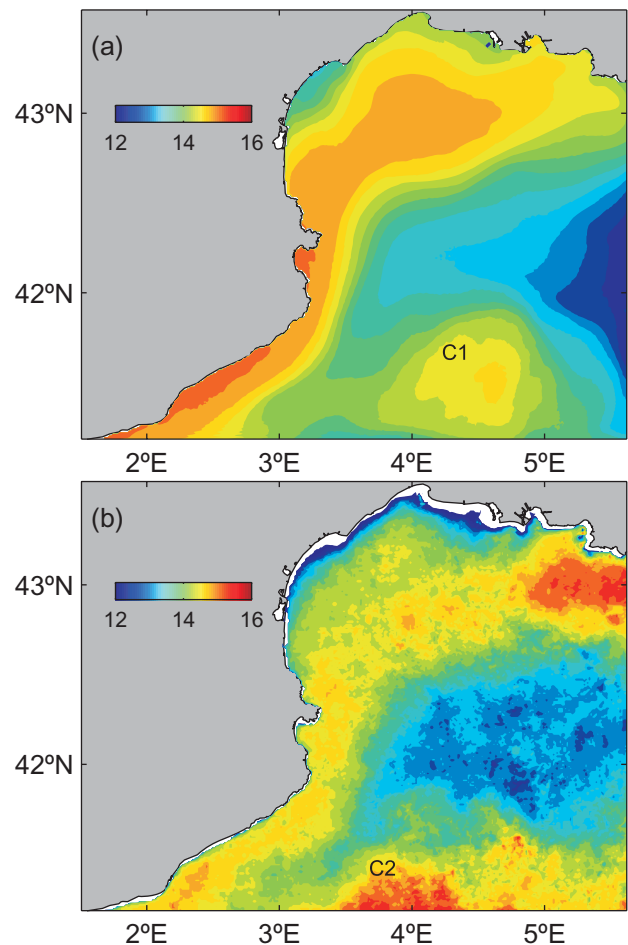


**Fig. 7.** Evolution of sea level anomaly (SLA, in cm) along the 100 m isobath in the Catalan Sea from 15 to 30 December 2005. The zero along-shore distance is taken at the southernmost point of the 100 m isobath in the model domain. The 200 km along the 100 m isobath corresponds approximately to Cap de Creus, which separates the Catalan Sea and the Gulf of Lions (see Fig. 1).

related to the submarine canyons develops fully during the model run.

### 3.2 Frontal structures and coastal trapped waves

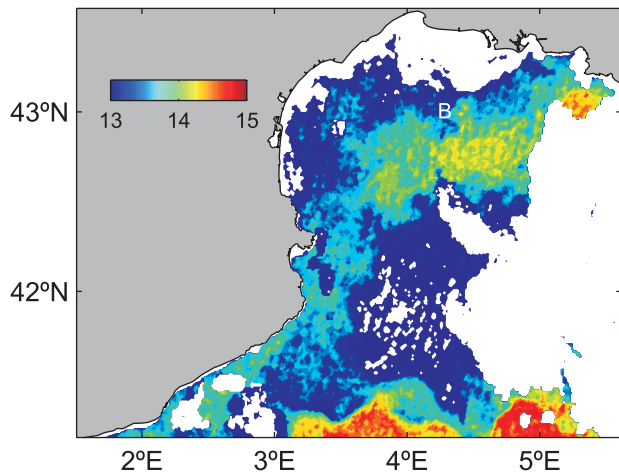
The southwestern propagation of meanders and eddies along the slope, associated with the Northern Current has been frequently observed moving along the slope in the Gulf of Lions and Catalonia (La Violette et al., 1990; Tintoré et al., 1990). Meanders of about 30–40 km propagating at phase speeds of 5–15 km/day that grow as filaments and develop into isolated eddies have been reported (Flexas et al., 2002; Rubio et al., 2005). Similar southwestward-propagating frontal structures are observed in the present simulation. Figure 6 reveals several meanders propagating along the slope. Southward propagation at phase speeds of about 9 km/day was estimated



**Fig. 8.** (a) Time-averaged model surface temperature ( $^{\circ}\text{C}$ ). The anticyclonic eddy is indicated with C1. (b) Monthly-averaged satellite (AVHRR) SST ( $^{\circ}\text{C}$ ) for December 2005. The anticyclonic eddy is indicated with C2.

from visual analysis of the complete image set (not shown). Off the slope of the Gulf of Lions ( $3.9^{\circ}\text{E}$ ,  $42.8^{\circ}\text{N}$ ) a filament with a mushroom-like structure (B) develops during the analyzed period.

The model is also able to reproduce coastal trapped waves which propagate cyclonically along the continental margins of the Gulf of Lions and the Catalan Sea. The presence of coastal trapped waves in the area have been recently identified by Jordi et al. (2005b) using alongshore currents and sea level time series. The time evolution of the sea level anomaly (SLA) predicted by the model along the 100 m isobath in the Catalan Sea is shown in Fig. 7. In spite that SLA contours are disturbed by small scale features, they present a constant slope that evidences the southwestward propagation of disturbances. The phase speed of these disturbances is about 250 cm/s, which corresponds to the phase speed of the first coastal trapped wave mode in the area.



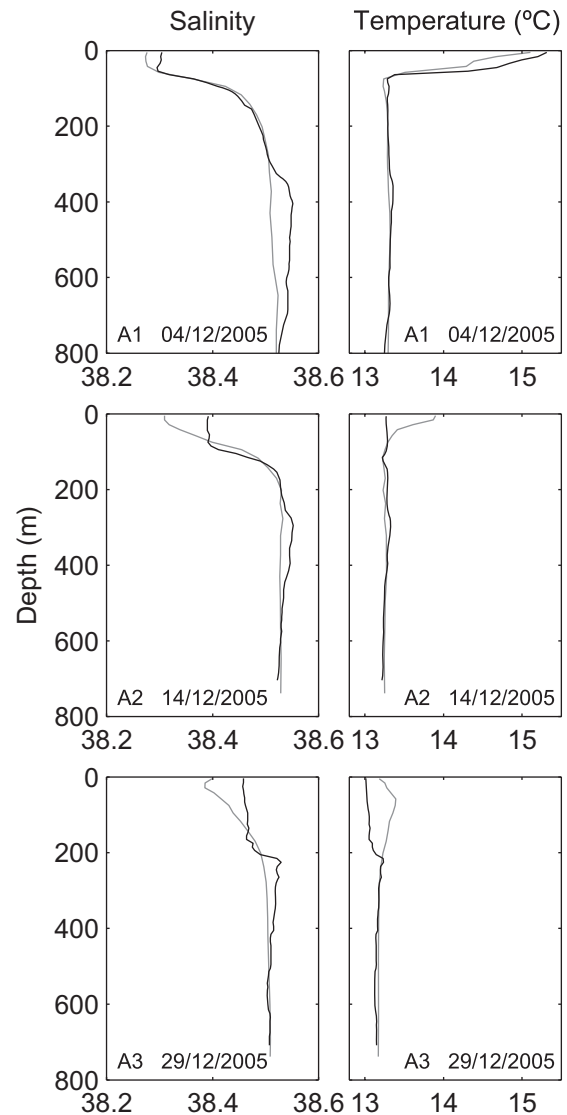
**Fig. 9.** Daily-averaged satellite (AVHRR) SST ( $^{\circ}\text{C}$ ) for 23 December 2005. A filament with a mushroom-like structure is indicated with B and corresponds to that indicated in Fig. 6.

### 3.3 Comparison with observations

Satellite images of sea surface temperature provide a useful tool to test the performance of numerical results in the upper ocean. A monthly averaged SST image of the study area, for December 2005, was obtained from the German Aerospace Research Center (DLR). Figure 8 shows the comparison between predicted sea surface temperature and observed data. Although this comparison is not direct because the first model layer is 10 m depth and the satellite gives a skin temperature, the large-scale pattern matches well with observations with some exceptions (i.e. local warming in the northeastern edge of the domain). Obviously, satellite temperature displays small-scale variability that is not captured by the model where all the features appear smeared. The model's temperature along the coast is generally warmer than the observations probably because of the lack of riverine inputs in the model. The location of the anticyclonic eddy in the model (C1) is displaced to the northeast and presents a slightly colder core than satellite values (C2).

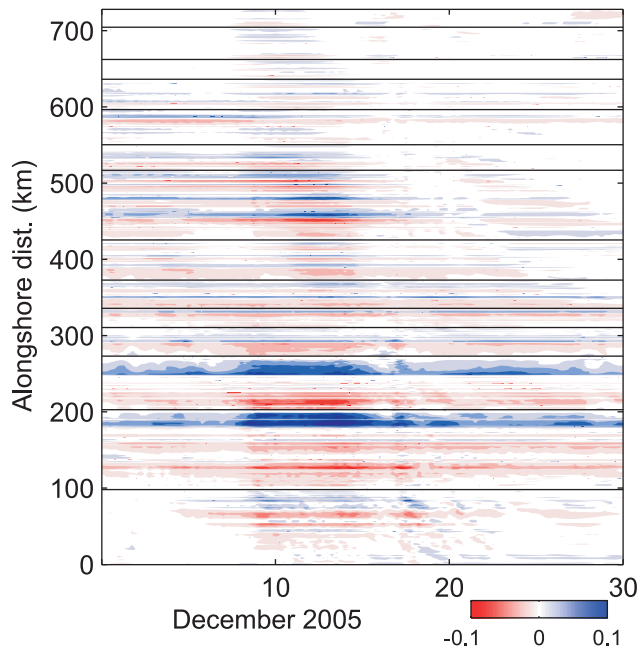
Frontal structures developing and propagating with the Northern Current can be also observed on satellite images of sea surface temperature since the Northern Current is mainly driven by salinity differences and temperature represents an excellent tracer (La Violette et al., 1990). Figure 9 shows the daily-averaged satellite image for 23 December 2005. The filament with a mushroom-like structure (B) located off the slope of the Gulf of Lions is contemporaneous to the predicted by the model (see Fig. 6).

An alternative tool to check the model performance of temperature and salinity predictions in the water column is the set of MedArgo floats (<http://www.coriolis.eu.org/cdc/argo.htm>). Figure 10 compares the predicted temperature



**Fig. 10.** Comparison between model temperature and salinity profiles (grey line) and those measured by Argo floats (black line). The date of each profile is indicated and its location is shown in Fig. 1.

and salinity profiles with those measured by Argo floats (see Fig. 1 for the location). At deeper levels, both modeled and observed fields are very similar, as we expected because ARGO profiles had been already assimilated into the global Mediterranean model, from which the initial condition has been taken. However, discrepancies appear in the mixed layer with differences of the order of  $0.2^{\circ}\text{C}$  in temperature and 0.03 in salinity on 4 December 2005. On 14 and 29 December 2005, observed surface temperature and salinity are cooler and saltier, respectively. Although further analysis is required, these discrepancies could be attributed to errors on the atmospheric forcing or to inaccuracies on the model's vertical turbulence formulation.



**Fig. 11.** Transport of volume (Sv) across the shelf break (defined by the 200 m isobath) at each individual grid point along the shelf break. Positive values indicate transport from the shelf to open ocean. The alongshore distance is the distance (in km) following the shelf-break from the southernmost point in the Catalan Sea (0 km) to the eastern edge of the Gulf of Lions at a distance of 720 km. Horizontal black lines represent the location of the canyons heads, which correspond to the canyons indicated in Fig. 1 (the upper black line is the Cassidaigne Canyon (13) in Fig. 1).

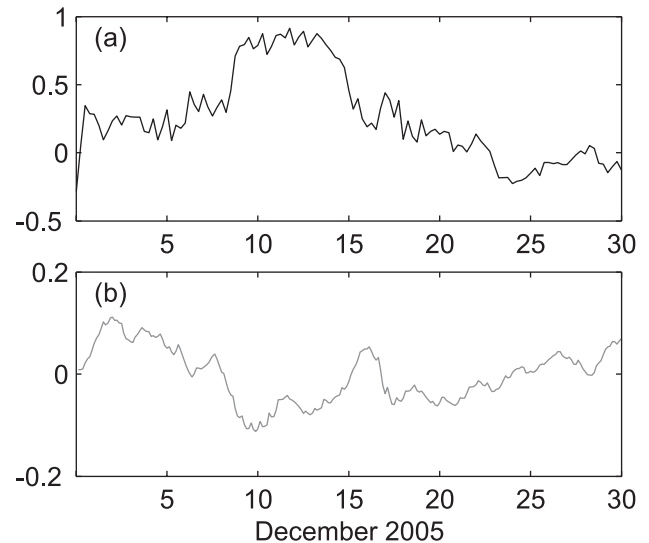
### 3.4 Shelf-slope exchange

To study the shelf-slope exchanges, we compute horizontal transports across a vertical plane formed by the shelf break (defined by the 200 m isobath) and its projection on the surface. Horizontal transport is estimated by integrating the cross-slope velocity across this plane with positive transport being directed offshore. According to Dinniman et al. (2003), the cross-slope velocity at each individual grid point along the shelf break is calculated as

$$U_{\text{cross}} = \mathbf{U} \cdot \left( \frac{\nabla H}{|\nabla H|} \right) \quad (1)$$

where  $\mathbf{U}$  is the horizontal velocity vector and  $H$  is the bathymetry.

Figure 11 shows the cross-slope transport, during December 2005, through the vertical plane formed by the projection of the shelf-break. Despite that the cross-slope transport is very close to zero at several locations, clear enhancements occur related to the location of submarine canyons, in particular along the Catalan shelf. Furthermore, onshore and offshore flows are observed on the upstream and downstream side of the canyons respectively. This effect of submarine



**Fig. 12.** (a) Total transport of volume (Sv) across the shelf break (200 m isobath) during December 2005. Positive values indicate transport from the shelf to open ocean. (b) Amplitude function of the most relevant EOF from the time variability of the u-component of wind forcing at 10 m.

canyons on cross-shelf transports has been previously observed and modeled by several authors (Skliris et al., 2002; Jordi et al., 2005a). The maximum cross-shelf transport occurs in the vicinity of the Palamós Canyon, which is the most prominent canyon of the area.

As suggested by Fig. 11, the cross-shelf transports are greater in the period from 8 to 15 December 2005. A budget of the volume transport across the shelf-break confirms this point and shows good correspondence with the amplitude function of the most relevant EOF from the u-component of wind forcing (Fig. 12). During this wind enhanced period, larger transports correspond to winds from the north-east. This is in agreement with the results of Arduin et al. (1999) for the Blanes Canyon. In fact, the combined effects of canyon topography and of the wind forcing during energetic storms are responsible for a large increase of both cross-shore and vertical transports (Skliris et al., 2004). Recently, a large increase in downward particle fluxes was measured in the Palamós Canyon during a severe storm with northeasterly winds (Palanques et al., 2005).

The total volume of water transported from the shelf to the open ocean during the integration period reached  $10.8 \cdot 10^{12} \text{ m}^3$ . Similarly, the volume transport from the open ocean to the shelf is estimated to be  $9.2 \cdot 10^{12} \text{ m}^3$  and hence a net offshore transport is suggested by these values. Nevertheless, not all cross-shelf transported water is effectively exchanged (Jordi et al., 2006<sup>3</sup>), since to a large extent these fluxes

<sup>3</sup>Jordi, A., Klinck, J. M., Basterretxea, G., Orfila, A., and Tintoré, J.: Estimation of shelf-slope exchanges induced by frontal



could correspond to onshore-offshore fluctuations. Jordi et al. (2006)<sup>3</sup> suggest that only a half of water transported across the shelf break produces an effective exchange of water. Following this estimation, the effective exchange in the Gulf of Lions and the Catalan Sea for December 2005 would be of about  $5 \cdot 10^{12} \text{ m}^3$ . Assuming a similar magnitude for each winter month, it could be concluded that waters of the Gulf of Lions and the Catalan Sea are completely renewed during this season. This result is consistent with the estimations obtained by Jordi et al. (2005a) for the northwestern Mediterranean and with the values reported by She and Klinck (2000) off Oregon.

#### 4 Conclusions

Within the framework of the MFSTEP, a high-resolution nonhydrostatic model has been developed in the northwestern Mediterranean coupled to a basin-scale Mediterranean Sea model. The numerical study is focused on the continental margins of the Gulf of Lions and the Catalan Sea, where the circulation is dominated by the Northern Current and shelf bathymetry is incised by numerous submarine canyons. The model reproduces well several known features of the general circulation in the area, although slight discrepancies were observed with satellite derived SST and Argo floats measurements in the mixed layer. Deficiencies in the atmospheric forcing and in the model's vertical viscosity and diffusivity formulation are suggested as plausible causes for these differences. Major efforts are presently centered in the use of air-sea coupling through bulk formulae, the implementation of a second-order turbulence closure scheme, and the development of data assimilation schemes to sequentially update the model forecast with new observations.

A relevant topic of this study is the estimation of shelf-slope exchanges. Volume transports across the shelf-break have been evaluated for December 2005. Our results indicate that submarine canyons contribute significantly to the exchange of water between the continental margin and the open ocean, specially during enhanced wind period. These estimations which suggest that shelf water may be renewed in a few months, are consistent with previously reported values in the vicinity of submarine canyons.

*Acknowledgements.* This work has been partially supported by ESEOO (VEM200320577C1408) project funded by CYCIT and MFSTEP (EVK3CT200200175) project funded by European Commission (D. G. Research).

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instability near submarine canyons, *J. Geophys. Res.*, submitted, 2006.

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