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30 ka of Angolan  
vegetation development

# Thirty thousand years of vegetation development and climate change in Angola (Ocean Drilling Program Site 1078)

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## Abstract

ODP Site 1078 situated under the coast of Angola provides the first record of the vegetation history for Angola. The upper 11 m of the core covers the past 30 thousand years, which has been analysed palynologically in decadal to centennial resolution. Alkenone sea surface temperature estimates were analysed in centennial resolution. We studied sea surface temperatures and vegetation development during full glacial, deglacial, and interglacial conditions. During the glacial the vegetation in Angola was very open consisting of grass and heath lands, deserts and semi-deserts, which suggests a cool and dry climate. A change to warmer and more humid conditions is indicated by forest expansion starting in step with the earliest temperature rise in Antarctica, 22 thousand years ago. We infer that around the period of Heinrich Event 1 a northward excursion of the Angola Benguela Front and the Congolian Air Boundary resulted in cool sea surface temperatures and a northward extension of desert vegetation along the coast. Rain forest and dry forest returned 15 thousand years ago. During the Holocene, dry forests and Miombo woodlands expanded. Also in Angola globally recognised climate changes at 8 thousand and 4 thousand years ago had an impact on the vegetation. During the past 2 thousand years, savannah vegetation became dominant.

## 1 Introduction

The climate of tropical Africa is strongly coupled to the position of the Intertropical Convergence Zone (ITCZ) and its southern African branch, the Congolian Air Boundary (CAB) (Leroux, 1983; Tyson, 1986). Shifts in the latitudinal position of the ITCZ have been associated with the climate variability in the Sahara and Sahel, but much less is known about the connections with the climate in southern Africa (e.g. Nicholson, 2000). Our knowledge mainly comes from studies on east African lakes, which focus on the climate of East Africa and the Rift Mountains (Vincens et al., 1999, 2005, 2007; Bonnefille and Chalié, 2000; Johnson et al., 2004). They are, therefore, less suitable

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to monitor the effects of ITCZ and CAB variability on climate and vegetation of west southern Africa.

Recording and understanding climate change in southern Africa is important because it opens a window to study the variability of both tropical and subtropical climate systems of the Southern Hemisphere. How does the Atlantic monsoon in Angola change from glacial to interglacial? Does it have Northern or Southern Hemisphere timing? How much is the precipitation in Angola influenced by East Atlantic sea surface temperatures. What is the impact of changes in the Atlantic meridional overturning circulation?

The latitudinal position of the CAB is connected with the amount of rainfall and the duration of the rainy season. The land-ocean temperature contrast is associated with aridity, wind strength and wind direction. Both directly affect the vegetation. Earlier studies have shown that vegetation changes in Angola are sensitively recorded by pollen and spores in marine sediments along the coast (Shi and Dupont, 1997; Dupont and Behling, 2006). The marine material allows us to compare the pollen record with the sea surface temperature (SST) record from the same core and thus enabling us a direct land-sea correlation.

We studied a high resolution marine core (ODP Hole 1078C) spanning the past 30 ka encompassing the last glacial maximum, deglaciation including Heinrich event 1 and the Younger Dryas period, and the Holocene. We aim to interpret changes in the pollen assemblage in terms of vegetation change on the continent taking into account effects of transport of pollen by wind and rivers. The thus reconstructed vegetation changes are compared to other African environmental records. We aim at a conceptual model of changes in the latitudinal position of atmospheric and marine fronts explaining the environmental change in Angola that might be tested with numerical models.

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## 2 Regional climate and ocean circulation and its impact on the modern vegetation of west southern Africa

Present-day climate in Angola ranges from the permanently wet climate of the Congo basin in the north to the dry summer rain climates in the south. The Atlantic or west African monsoon controls the climate of the northernmost part during most of the year, while the middle and southern parts receive (Atlantic) moisture only during austral summer. Angola is the southernmost region where the West African monsoon brings rain (Fig. 1). The boundary is in the form of a low pressure area, the CAB (formerly Zaire Air Boundary), that annually migrates over Angola from its southern surface position around 20–15° S in January (austral summer) to its northern position around 9–6° S in July (austral winter)(Leroux, 1983). The CAB is part of the discontinuities that divide easterly trades and westerly monsoonal wind systems over Africa and can be regarded as a southern branch of the ITCZ. The latitudinal position and annual migration of these systems determines the timing and duration of the rainfall seasons. They are associated with the global thermal gradient and the strength and position of the Hadley Cell (Nicholson, 2000). However, sea surface temperatures and ocean circulation also have an impact on the position of the ITCZ.

The surface and shallow subsurface ocean circulation in the eastern Angola Basin is dominated by the South Atlantic cyclonic gyre consisting of the Benguela Ocean Current flowing northwards and westwards over the Walvis Ridge, the South Equatorial Counter Current flowing eastwards, and the Angola Current flowing southwards along the coast (Fig. 1). In the marine realm, as in the terrestrial one, an important frontal system exists off the coast of Angola in the form of the Angola Benguela Front (ABF) annually migrating between 17 and 14° S. South of the ABF, the surface Benguela Coastal Current being a tongue of the Benguela Ocean Current flows northwards along the coast. North of the ABF, the Angola Current flows southwards and the Benguela Coastal Current can be traced as a shallow subsurface current up to 5° S (Peterson and Stramma, 1991; Schneider et al., 1995, 1997).

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The ABF is recognizable by a sea surface temperature (SST) gradient of 4°C per 1° of latitude and identifiable to a depth of at least 200 m. The front is most intense within 250 km from the coast. The position and the strength of the front are maintained through a combination of factors including coastal orientation, wind stress, and opposing flows of the equatorward Benguela Coastal Current and the poleward Angola Current (Shannon and Nelson, 1996). Within the limits of the coastal orientation, there is a coupling between the position of the ABF and the CAB on land. South of the front the zonal component of the wind stress over the ocean is east (SE trades) and north of the front the zonal component is west (SW monsoon). Strong SE trade winds will trigger both a cooling south of the ABF by increasing the upwelling along the coast and a warming north of the ABF due to stronger intrusion of warm equatorial waters by the Angola Current (Kim et al., 2003). Intense coastal upwelling occurs only south of the ABF, whereas directly north of it upwelling is seasonal and relatively weak (Wefer et al., 1998).

The position of the climatic and oceanic frontal systems is reflected in the vegetation of southern Africa. The southern boundary of the Congolian rain forest corresponds to the latitudinal July (austral winter) position of the CAB. The Zambebian phytogeographical region is most extended consisting of different types of dry forest, woodland, savannah and edaphic grassland (White, 1983). In the north of Angola the transition between the Zambebian and the Congolian vegetation zones nowadays is mainly occupied by secondary grassland and savannahs. However, surviving cloud forests of Congolian affinity exist occurring at altitudes between 350 and 1000 m. Along the west coast, climate and vegetation are strongly influenced by the SST. Especially, the low SSTs of the upwelling area south of the ABF suppress coastal rain fall (Nicholson and Entekabi, 1987). As a result, the vegetation along the Angolan coast south of 15° S has affinities with the Namib Desert (desert of the Moçamedes) and south of 11° S with the Karoo-Namib shrubland (White, 1983).

The vegetation of Angola forms the transition between the tropical lowland rain forest of the Congo and the dry savannahs and deserts of the Kalahari and Namibia. Its

development is sensitive to climate change and shifts of the frontal systems described above. In this paper, we present a detailed pollen record of marine sediments of ODP Hole 1078C, which allows the study of the vegetation history of the past 30 ka during full glacial, deglacial, and interglacial conditions.

### 3 Material and methods

The material used is sediment from the upper 11 m ODP Site 1078 Hole C. The site is located outside the Bight of Angola at 11°55' S 13°24' E in 427 m deep water. The sediments are composed predominantly of a moderately bioturbated olive-gray silty clay with varying amounts of nannofossils and foraminifers (Wefer et al., 1998).

The age model is established after radiocarbon AMS measurements on foraminifer tests and molluscs carried out at the Leibniz-Laboratory, University of Kiel, Germany (Kim et al., 2003; Rühlemann et al., 2004). Extension of the age model uses eight dates (Table 1) calibrated after Fairbanks et al. (2005) using a marine reservoir effect of 400 years ( $\Delta R=0$ ). All ages are expressed in 1000 years calibrated BP abbreviated as ka. Dates at 1071 and 1021 m below sea floor (mbsf) have been discarded, because the shell fragments might have been displaced to lower levels by burrowing.

Preparation of palynological samples follows standard procedures. In brief, 3–4 cm<sup>3</sup> sediment was decalcified with diluted HCl (ca. 5%) and treated with HF (ca. 20%) for several days to remove silicates. Samples were sieved over a ca. 10 μm mesh using ultrasonic sieving. Pollen samples were mounted in glycerine jelly, microscopically examined by H. Behling and L.M. Dupont. Pollen grains were identified using several publications (Ybert, 1979; Bonnefille and Riollet, 1980; Sowunmi, 1973, 1995), the African Pollen Database (<http://medias.obs-mip.fr/pollen/>), and a reference collection. Charcoal particles have been counted from the pollen slides. Samples are counted routinely up to 300 pollen grains. However, some samples turned out to be too poor to reach that goal. Percentages are expressed on the total number of pollen and fern spores, which is denoted in Fig. 4.

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Detailed information on analytical techniques used for the determination of long-chain alkenones can be found in the study by Kim et al. (2002). We calculated the alkenone unsaturation index from  $U_{37}^{K'} = [C_{37:2}] / ([C_{37:2}] + [C_{37:3}])$ , where  $C_{37:2}$  and  $C_{37:3}$  represent the di- and triunsaturated  $C_{37}$  alkenones, respectively (Prahl and Wakeham, 1987). The  $U_{37}^{K'}$  values were converted into temperature values applying the culture calibration ( $U_{37}^{K'} = 0.034 \cdot T + 0.039$ ) by Prahl et al. (1988), which has also been validated by core-top compilations (Müller et al., 1998). The precision of the measurements ( $\pm 1\sigma$ ) was better than  $0.003 U_{37}^{K'}$  units (or  $0.1^\circ\text{C}$ ), based on replicate extractions on different days of two laboratory internal reference sediments (CC 2107-3 and CC 1706-2) from the South Atlantic.

#### 4 Age model and sedimentation rates

The age depth model is shown in Fig. 2. The younger part (0–22 ka) is after Kim et al. (2003) and Rühlemann et al. (2004) and has been calibrated using Calib and the marine calibration curve (Stuiver and Reimer, 1993; Hughen et al., 2004). The lower dates (Table 2) are outside the range of the marine calibration curve and have been calibrated using Fairbanks et al. (2005). Calibration ranges increase with age and the uncertainty of the calibration. Some radiocarbon ages seem to occur at different levels, which can be explained by burrowing down of shells and shell fragments. We therefore consequently chose the upper ones to include in the age model.

For the calibration of the radiocarbon dates, we have applied a reservoir effect of 400 years as suggested by the present day marine reservoir correction database (Butzin et al., 2005). Although there are no direct measurements in the vicinity of ODP Site 1078, we think this is justifiable because the site is just outside the coastal upwelling range and the influence of the open ocean is more pronounced than at sites north or south of it (Wefer et al., 1998). Simulations for the glacial ocean indicate an increase in the surface reservoir age to maximal 600 years (Butzin et al., 2005). An extra shift of 200

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years in the glacial dates would not significantly change our age model.

The sedimentation rates for ODP Site 1078 that result from linear interpolation between the calibrated radiocarbon ages are also given in Fig. 2. Until 27 ka sedimentation rates range between 22 and 35 cm/ka, but are rather imprecise. Between 10–9.8 mbsf (27–22 ka) sedimentation rates drop to low values of about 4 cm/ka. This estimate is supported by three ages. Moreover, a concentration of shells and molluscs have been found at those levels both in Hole C, used in this study, and in the neighbouring Hole B (Wefer et al., 1998). The very low sedimentation might have occurred during the last glacial lowest sea-level period. This seems to be much too early, but a recent publication provided evidence that the period of lowest sea-level must have occurred between 26 and 21 ka ago (Peltier and Fairbanks, 2006).

After 22 ka sedimentation rates increase again to levels between 20 and 40 cm/ka and jump to levels around 1 m/ka shortly before 14 ka. We associate the strong increase in sedimentation rates with the rapid sea-level rise during Meltwater Pulse 1a dated between 14.6–14.3 ka (Fairbanks, 1989; Hanebuth et al., 2000). The levels of high sedimentation rates between 8 and 3 mbsf (14–9 ka) shows raised *Rhizophora* pollen percentages. This increased relative abundance of mangrove pollen comes together with increased amounts of other mangrove materials and is interpreted as sedimentation of eroded mangrove peat (Kim et al., 2005). Sedimentation levels remain high until 9 ka, when they drop again to levels between 25 and 40 cm/ka.

Thus the history of sea-level change at ODP Site 1078 as reconstructed with our age model fits well into the global picture, which makes us more confident about the age model itself.

## 5 Results

Alkenone-derived SSTs showed a typical glacial-interglacial warming trend, ranging from 21° to 24°C over the last 30 ka (Fig. 3). During the glacial period, alkenone-derived SSTs varied between 21° and 23°C and decreased to the lowest value of

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20.9°C occurred at 16.3 ka during Heinrich Event 1. Afterwards, alkenone-derived SSTs increased towards the Holocene, showing a warm period during the Younger Dryas. During the Holocene, alkenone-derived SSTs continued to increase to a core-top value of 24.9°C with a slight cooling around 5 ka.

5 Based on the relative abundance of the most conclusive pollen taxa, seven pollen zones (PZ) have been distinguished. In this section, each zone is briefly characterised. Average percentage values per pollen zone are given for selected pollen taxa in Table 2. Percentages of groups and selected pollen taxa are given in Figs. 4–6. Appendix A lists the pollen taxa within each group.

10 PZ 1: 25 samples between 11.16 (lowermost sample analysed) and 9.70 m, dated between 31.7–21.9 ka. The record is poor between 9.99 and 9.81 m (26.3–22.4 ka), but not systematically different from the rest of the zone (see Table 2). Radiocarbon dates indicate very low sedimentation rates (~4 cm/ka) for this part of the zone. Maximum percentages of the mountain elements including pollen of Ericaceae, *Myrica* and *Ilex* cf. *mitis*, as well as spores of the fern *Anemia* and liverworts (*Phaeoceros*, *Anthoceros*) are found. High pollen percentages of Asteroideae, *Stoebe*-type, Poaceae, and other desert elements, e.g. *Tribulus* occur as well as raised percentages of *Brachystegia*.

15 PZ 2: 19 samples between 9.70–8.80 m, dated between 21.9–18.8 ka. Gradual increase in the pollen concentration (Fig. 3) and a decline in percentages of desert and mountain elements are found. *Myrica* percentages are still relatively high. First rise in *Rhizophora* pollen percentages. Increase in percentages of *Podocarpus* and rain forest elements including *Mallotus*-type.

20 PZ 3: 23 samples between 8.80–8.02 m, dated between 18.8–15.4 ka. Maximum *Podocarpus* percentages occur parallel to a minimum in sea surface temperatures as inferred from alkenones (Kim et al., 2003; Fig. 3). Of the rain forest elements only percentages of *Mallotus*-type and *Tetrorchidium* are raised. *Rhizophora* percentages also increase.

25 PZ 4: 43 samples between 8.02–4.20 m, dated between 15.4–10.0 ka. Pollen accumulation rates are dominated by the sedimentation rates (Fig. 3). Increase of pollen

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percentages of *Uapaca* and *Brachystegia* and the sum of Miombo elements. Rain forest percentages are relatively high and so are the values for a number of constituents: *Tetrorchidium*, *Lophira*, *Celtis*, and *Alchornea*. *Rhizophora* percentages also rise, parallel to the global eustatic sea-level curve (Fleming et al., 1998)

5 PZ 5: 15 samples between 4.20–2.80 m, dated between 10.0–7.8 ka. Percentages of *Olea* decrease, those of Miombo elements including *Brachystegia* and *Uapaca* are high, percentages of dry forest with *Burkea*-type increase, and percentages of rain forest elements decline with the exception of *Alchornea*. High percentages also of *Artemisia afra* also occur. *Rhizophora* and Cyperaceae percentages are maximal between 9–8 ka. Pollen concentrations are high and exceed on average 10 000 grains per ml.

15 PZ 6: 17 samples between 2.80–1.50 m, dated between 7.8–3.7 ka. Maximum percentages are found for *Podocarpus*, Miombo elements including *Brachystegia*, *Azalia*, *Berlinia*-type, and *Uapaca*. Furthermore percentages of Rhamnaceae (*Ziziphus*) pollen and *Pteris* spores are raised. From the dry forest elements *Adansonia digitata* and *Lannea*-type have higher values. Minimal Poaceae and *Rhizophora* percentages occur and Cyperaceae percentages are low. Pollen influx (accumulation rates) of *Podocarpus* and Miombo elements stay at the same level, but decline for Poaceae, Cyperaceae, *Rhizophora*, dry forest and rain forest elements.

20 PZ 7: 34 samples between 1.50–0 m, dated between 3.7 ka and modern. Maximum percentages of Poaceae and Cyperaceae occur. Poaceae percentages reach their maximum after 2 ka. Percentages are raised again for *Rhizophora* and rain forest elements, in particular *Alchornea* and *Elaeis guineensis*. Also dry forest elements are better represented including *Hymenocardia*, *Artemisia afra*, *Protea*, *Burkea*-type, and Combretaceae.

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## 6 Vegetation development in Angola in relation to climate change

ODP Site 1078 located on the lower shelf only 50 km from the shore is in an ideal position to collect pollen from the mainland. Fluvial transport of pollen is important, because the site is situated in front off the mouth of the Balombo River, which drains the highest range of the Huambe mountains. The river collects pollen from a wide range of vegetation types passing at present through Afromontane vegetation, wetter Zambezan Miombo woodland and bushy Karoo-Namib shrubland along the coast (White, 1983). Almost 2° north of the site, the southernmost mangrove swamps of the Atlantic coast of Africa are situated along the Longa River (Spalding et al., 1997). Southwest of ODP Site 1078, Moçamedes bushland and desert stretches along the coast of southern Angola and Namibia. Further inland in Angola, Mopane scrub woodlands occur. Aeolian transport of pollen to ODP Site 1078 is mainly expected during austral fall and winter when the lower tropospheric flow is from the SE (Dupont and Wyputta, 2003). Source areas for fluvial and aeolian pollen largely overlap.

### 6.1 Glacial

The lowermost part of the record, PZ1, covers the glacial between 31.7 and 21.9 ka. As discussed above, it includes the last glacial maximum (defined as the period of lowest global sea-level). From the high pollen percentages of mountainous elements, including Ericaceae, *Myrica* and *Ilex* accompanied with high relative abundances of pollen of Poaceae and Asteraceae, we infer that open mountainous vegetation with heather species (Ericaceae), composites (Asteraceae), and grasses (Poaceae) was wide-spread (Fig. 5). Ericaceae and *Ilex mitis* are typical elements from the Afromontane vegetation (White, 1983). Palgrave (2002) mentions several *Myrica* species from mountain areas. Apart from the Afromontane vegetation, the existence of dry woodlands, probably at lower altitudes is suggested by the occurrence of pollen of *Brachystegia* and spores of *Anemia*. The Flora Zambesiaca (Schelpe, 1970) describes the present habitat of *Anemia angolensis* in the undergrowth of dry woodland at altitudes

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between 600–1525 m.

Pollen from composites and grasses probably had their sources not only from Afromontane vegetation but from desert and semi-desert vegetation along the coast, as well. Extension of desert and semi-desert is also indicated by the regular occurrence of *Tribulus* and *Stoebe*-type pollen. Is this extension a result of a northward shift of the Angola Benguela Front (ABF) and increased upwelling along the coast? Alkenone measurements indicate SSTs (Fig. 3) were around 22°C (Kim et al., 2003) 2 to 3° lower than today. Increased upwelling, however, should have lowered the SST below the general cooling of the glacial ocean, which is estimated at 4–5°C for the east equatorial Atlantic (Niebler et al., 2003). Comparison of several SST profiles with a coarser timely resolution between 5° and 20° S did not indicate a shift of the ABF north of 12° S during the glacial (Schneider et al., 1995). Thus, there is little indication of upwelling of cooler intermediate waters.

The open mountainous vegetation and the expansion of desert and semi-desert suggest a cooler and drier climate in Angola during the glacial. Our data corroborate terrestrial mountain records, especially in east Africa, of an open glacial landscape covered with a grass- and ericaceous-rich vegetation (e.g. Coetzee, 1967; Bonnefille and Riollet, 1988; Bonnefille and Chalié, 2000). However, in Congo, at lower altitudes and nearer to the equator, no mountain elements were found (Elenga et al., 1994). Forest growth in the mountains might have been impeded by low temperatures at higher altitudes, dry conditions, and/or due to low glacial atmospheric carbon dioxide levels (Jolly and Haxeltine, 1997).

Between 21.9 and 18.8 ka, PZ2, the vegetation gradually changes. The relative abundance of pollen from open vegetation types (Afromontane and desert) declines, and tree pollen percentages of *Podocarpus* (Fig. 4) and *Mallotus*-type (Fig. 6) rise. Probably, *Podocarpus* trees were spreading in the mountains, while rainforest started to expand in the northern lowlands with pioneer trees such as *Mallotus* (Maley and Brenac, 1998). Other Afromontane trees such as *Ilex mitis* and *Myrica* are still present (Fig. 5). The gradual increase in pollen concentration suggests higher terrestrial –

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probably fluvial – input. Around 22 ka, the lake level of Lake Malawi rises after the glacial lowstand indicating increased precipitation in the East African mountains (Johnson et al., 2004). Furthermore, semi-deciduous forest elements became established around Lake Masoko just north of Lake Malawi suggesting the dry season became as short as 3 to 4 months (Vincens et al., 2007). Glacial woodland and bushland with Afromontane elements around Lake Rukwa in Tanzania (almost 2° of latitudes further north), however, indicates that the climate was still rather cool and dry (Vincens et al., 2005).

Expansion of rain forest in the northern lowlands of Angola also suggests increased precipitation, which is possible if the winter migration of the CAB migrated less northwards than in the period before. Then the CAB position would remain sufficiently to the south to allow Atlantic rain penetrating during most of the year in the northern lowlands of Angola. The spread of forest between 22 and 19 ka is too early to be an effect of rising carbon dioxide levels, because the CO<sub>2</sub> rise after the glacial is dated by Monnin et al. (2001) between 17 and 15.4 ka. Moreover, the pollen diagram from Lake Barombi in Cameroon indicates an opening of the forest and increase of grasslands from ca. 24 ka on (20<sup>14</sup>C ka; Maley and Brenac, 1998). The development north and south of the equator proceeding in opposite directions would fit with a southern average position of the ITCZ coupled to a southern position of the CAB (Fig. 7). Our interpretation fits the study of Johnson et al. (2004) inferring a southern average position of the ITCZ between 23 and 19 ka based on the increased productivity in Lake Malawi coupled to northerly winds over the lake.

## 6.2 Deglaciation

The pollen diagram for the period between 18.8 to 15.4 ka, PZ3, covering Heinrich Event 1, is dominated by *Podocarpus* (Fig. 4). *Podocarpus* trees produce much aerodynamically well-equipped pollen and are, therefore, overrepresented in the pollen record, especially when wind transport is important. Still, there are few other elements represented save for *Mallotus* and *Tetrorchidium* from the rain forest (Fig. 6).

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Thus, although rain forest remained present in the northern lowlands, *Podocarpus* forest might have been locally abundant in the mountains. Unfortunately, the marine record of *Podocarpus* for this period is not consistent, because a comparable *Podocarpus* pollen maximum is missing in the nearby core of GeoB1016, but present in core GIK16867 at 3° S (Dupont et al., 1996). The terrestrial pollen records in Cameroon and Congo indicate a mixture of semi-deciduous forests and grasslands but neither an open savannah nor a *Podocarpus* forest (Elenga et al., 1994; Maley and Brenac, 1998; Elenga et al., 2004). Precipitation estimates based on deuterium of terrestrial plant waxes from a core off the Congo River mouth indicate minimal humidity in the Congo Basin (Schefuß et al., 2005).

During this period SSTs reach a minimum of 21°C and the temperature difference with data from a site located 5° of latitude further south (GeoB1023) is minimal (Kim et al., 2003). Kim et al. (2003) concluded that because of the small SST difference between 12 and 17° S, the SE trade winds must have been weak. During the same period, biomarker based annual temperature estimates for the Congo Basin indicate lowest temperatures of the past 25 ka (Weijers et al., 2007). The low SSTs at 12° S suggest a northern position of the ABF, over or even north of the location of the site. The CAB reached a northern position again, at least during austral summer. The associated SE trade winds probably reached further northward and might have brought more *Podocarpus* pollen grains with them to the site increasing the representation of *Podocarpus* both in absolute values (pollen concentration, Fig. 3) and relative ones (pollen percentages, Fig. 4). Another implication is that upwelling over the site intensified. This upwelling would have increased the reservoir effect in the radiocarbon dates and the material consequently might be several centuries younger than dated by our age model. The pollen zone, however, would still include the period of Heinrich Event 1.

According to our age model, sedimentation rates increase strongly to values around 1 m/ka at 14.3 ka which correlates with Meltwater Pulse 1a and global sea-level rise (Fairbanks, 1989; Hanebuth et al., 2000). Higher sedimentation rates increased the

accumulation of all pollen types, but *Rhizophora* pollen from the mangroves is over-represented (Fig. 4). A raised representation of mangroves both in relative pollen abundance and by biomarkers during this time and is explained by erosion of coastal mangrove peat during transgression and sea-level rise (Kim et al., 2005).

After Heinrich Event 1, between 15.4 and 10.0 ka (PZ4), the representation of the lowland rain forest becomes stronger. Not only pioneer trees such as *Mallotus*, *Celtis*, and *Alchornea* occurred, but also *Tetrorchidium* and *Lophira*, indicating increase and maturing of the forest. During this period – just before the Holocene – the lowland rain forest probably reached its southernmost extension (Fig. 6). Apart from the expansion of the lowland rain forest, elements from dry forest and woodlands, such as *Burkea*-type and Combretaceae, and from the Miombo woodlands, such as *Uapaca* and *Brachystegia*, are increasingly represented. The *Podocarpus* record goes slightly down and that of *Olea* rises indicating a change in the mountain forest composition. The Younger Dryas period is not distinguished as such in our pollen record, but part of a longer period of favourable conditions for forest growth.

Most African records show a first return to moister conditions after the glacial around 15 ka, or slightly later (e.g. Maley and Brenac, 1998; Gasse, 2000; Gasse and Van Campo, 2001; Barker et al., 2004). Deuterium based precipitation estimates are high between 15 and 12 and low between 12 and 11 ka (Schefuß et al., 2005). In the Congo Basin increased precipitation resulted subsequently (at 13.5 ka) in a large palaeodischarge event (Marret et al., 2001). Interruption and return to arid conditions during the Younger Dryas period is clear in the records from north of the equator (e.g. Hoelzmann et al., 2004). South of the equator, Lake Malawi shows an increase of productivity during the Younger Dryas period (Johnson et al., 2004). The pollen record of Lake Masoko shows an abrupt and irreversible change in the flora of the Younger Dryas period suggesting the climate changed from a mild short (3–4 months) dry season to a distinctive dry period of no less than 5 to 6 months (Vincens et al., 2007). Other records from southern Africa either indicate wetter conditions (Shi et al., 2000) or no specific changes that might be associated with the Younger Dryas.

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The strong development of the rain forest and the initial development of the woodlands in Angola suggest an early southward shifting of the CAB during winter and summer (Fig. 7). The summer latitude might have been comparable to those of today and the winter latitude might have been somewhat further to the south causing a southward shift of the area with more or less humid conditions all year round. SSTs rose strongly indicating that the ABF, and with it the area of coastal upwelling, probably shifted poleward, south of the site. The large difference in SSTs between 12 and 17° S suggest strong trade winds, in particular during the Younger Dryas period (Kim et al., 2003).

### 6.3 Holocene

During the early Holocene, between 10.0 and 7.80 ka (PZ5), dry forest and Miombo woodlands became increasingly important while rain forest elements retreated, but for the pioneer tree *Alchornea* (Fig. 6). Miombo woodland is nowadays the prevalent vegetation on well drained soils of the main plateau in Angola and other parts of the Zambebian region (White, 1983). The woodlands are dominated by species of *Brachystegia*, often accompanied by *Julbernardia* or *Isoberlinia* (included in the *Berlinia*-type). Associates are *Afzelia*, *Uapaca*, and *Pericopsis*. Decline in the representation of most rain forest elements suggest a northward shift of the winter CAB more or less to its modern position. The occurrence of *Artemisia afra* throughout the period and that of *Protea* around 8 ka might indicate disturbance of mountainous vegetation. Mangrove peat erosion increased during Meltwater Pulse 1b around 10 ka (Flemming et al., 1998) and the high total pollen input afterwards is associated with more river discharge.

Forest composition changed at 7.8 ka. Until 3.7 ka (PZ6) *Podocarpus* dominates in the mountains and other mountainous elements, such as *Olea*, are not represented (Figs. 4, 6). The representation of grasses is minimal indicating that open vegetation had become rare (Fig. 5). Dry forests included *Adansonia digitata* and *Lannea* (Fig. 6). The high representation of *Brachystegia*, *Afzelia*, *Berlinia*-type, and *Uapaca* suggests a large expansion of the Miombo woodland. However, pollen percentages of

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Rhamnaceae, probably from *Ziziphus*, and spore percentages of *Pteris* are also raised, indicating increased swamp forest vegetation. Among the genus *Brachystegia* there is one species *B. laurentii* occurring in wetter types of the rain forest. Also among *Azelia*, the rain forest species *A. africana* might be present, but this species occurs nowadays in the Congolian rain forest only north of the equator (White, 1983). The genus *Uapaca* includes rain forest species, as well. Thus, an alternative explanation to the spreading of Miombo woodland would be the expansion of wetter types of rain forest. In both interpretations, savannahs were replaced by closed forest.

At 3.7 ka, the vegetation drastically changed again. The occurrence of *Rhizophora* pollen grains (Fig. 4) during this period indicates pollen transport from the north as the modern southernmost distribution of mangroves is along the River Longa (10° S, Spalding et al., 1997). Supply of *Rhizophora* pollen through erosion of older mangrove peat strata is less likely, because sea-level did not rise much during this period.

Typical for the past 3.7 ka (PZ7) is the dramatic increase in the representation of Cyperaceae and grasses (Fig. 4). After 2 ka, the relative abundance of Poaceae pollen is even higher than during glacial times. The increase in Cyperaceae pollen percentages might be a localised phenomenon, the increase in grass pollen percentages is more regional as it is found as well in GeoB1016, located further offshore on the lower continental slope at 3411 m water depth (Shi and Dupont, 1997; Dupont et al., 2007). We find the high representation of grasses associated with more forest elements but not with increased relative abundance of desert or mountain elements. We, therefore, infer a strong increase in savannah. The modern savannah forest mosaic of northern Angola might have come into existence around 2000 years ago. The patchy character of the forests is underlined by the increased representation of light-loving trees such as *Alchornea* and the oil palm, *Elaeis guineensis* (Fig. 6).

Other changes are found in the composition of the dry forest, in which *Burkea*, *Hymenocardia*, *Protea*, and *Artemisia afra* must have been prominent (Fig. 6). The latter three taxa are associated with fire. *Artemisia afra* often colonizes burnt areas in mountains (Beentje, 1994), *Protea* grows in mountainous grasslands that are burnt, and

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*Hymenocardia acida* is fire hardy (White, 1983). However, there is no increase in the amount of charcoal particles relative to the amount of pollen and spores. The charcoal record is especially high during the glacial (Fig. 6), but indicates that fire always played a role in the vegetation development. We can only explain the occurrence of more fire hardy elements by assuming a change but not necessarily an increase in the fire regime that favoured these specific trees.

The development in Angola fits the general trend of a humid early and middle Holocene followed by a drier late Holocene. The savannah minimum and probably most humid phase in Angola from 7.8 to 3.6 falls during the maximum northward extension of the Sahelian vegetation from 7.8 to 5.7 ka (Hoelzmann et al., 2004) but lasts longer. The humid phase in Angola ended around the same time as the second of two aridification steps (around 6.7–5.5 ka and 4.0–3.6 ka) indicated by lake levels of the Sahara-Sahel region (Gasse, 2000). Only the first aridification step is recorded in the dust flux record of ODP Hole 658C as the end of the African Humid period (DeMenocal et al., 2000; Adkins et al., 2006). Also west equatorial pollen records and the deuterium record for the Congo Basin indicate a sharp drop in humidity around 3 ka, which favoured the opening of the forest and the spread of savannahs (Elenga et al., 2004; Schefuß et al., 2005). A reforestation after 1 ka as suggested by the pollen records from Congolian sites (Vincens et al., 1999) is not found in our record. In East Africa Holocene precipitation levels became increasingly variable after 4 ka and declined after 2 until 0.5 ka (Bonnefille and Chalié, 2000). For southern Africa numerous proxies indicate a temperature optimum shortly before 7.5 ka and increase in summer rainfall. Aridification set in soon after 5 ka (Scott and Lee-Thorp, 2004).

Thus, the Holocene climatic changes proceed largely parallel in northern, southern and equatorial Africa. Important changes in the average position or latitudinal migration of the tropical frontal systems (ITZC and CAB), however, should have caused changes in the opposite direction. It seems that these frontal systems were in place during the Holocene and that the aridification of the past 4 to 2 ka cannot be explained by a southward shift of the average position of the CAB and ITZC over Africa.

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## 7 Conclusions

The marine pollen record of ODP Hole 1078C provides a detailed history of the Angolan vegetation development since the last glacial, which we interpret as follows.

- From 31.7 ka (the beginning of the record) until 21.9 ka open mountainous vegetation with heather species, composites, and grasses was important, but dry woodlands also existed. Desert and semi-desert extended rather far to the north.
- Between 21.9 and 18.8 ka *Podocarpus* trees spread in the mountains, while rain forest started to expand in the northern lowlands.
- From 18.8 to 15.4 ka rain forest remained present in the northern lowlands and *Podocarpus* forest might have been locally abundant in the mountains.
- After 15.4 ka the *Olea* record indicates a change in the mountain forest composition. Until 10.0 ka rain forest was important, but woodland elements were also present.
- The Holocene started with dry forest and Miombo woodlands expanding while rain forest elements retreated.
- After a disturbance around 8 ka *Podocarpus* forest dominated between 7.8 and 3.7 ka. During this period, either Miombo woodland or wetter types of rain forest expanded and replaced open savannah vegetation.
- Savannahs became increasingly important after 3.7 and again after 2 ka. Forests probably became patchy growing light-loving and fire hardy trees.

Glacial SSTs fluctuate around 22°C until 19 ka and show a prominent minimum of ca. 21°C at 16 ka. From this minimum SSTs rise continuously to mid and late Holocene levels between 24 and 25°C. During the past 7 ka, a slight cooling by half a degree occurred between 5 and 4 ka.

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After the cool and dry climate of the last glacial the change to more humid conditions started very early in Angola (22 ka). The early amelioration is probably related to the early warming on Antarctica. The CAB might have shifted to the south (Fig. 7).

Between 18.8 and 15.4 (around Heinrich Event 1), low SSTs indicate a northern position of the ABF implying permanent upwelling along the coast and strong SE trade winds. Conditions were cool but probably not arid. During this early stage there is a different development in Angola than in the Congo Basin or in West Africa where the climate was cold and dry.

After Heinrich Event 1, the climate in Angola became wetter and warmer and possibly even wetter than during most of the Holocene. No return to glacial conditions was found for the Younger Dryas period. The average position and the latitudinal migration of the CAB and ITCZ might have been further south than during the Holocene (Fig. 7).

During the early Holocene the average position of the CAB was probably already close to its present position and the modern West African monsoon system in place.

## Appendix A

### Grouping of pollen types

#### Steppe/desert/semi-desert

Acanthaceae (*Barleria*, *Justicia/Monechma*, *Petalidium*), Aizoaceae, Amaranthaceae (*Gomphrena*-type), Amaranthaceae/Chenopodiaceae, Arecaceae (*Phoenix*), Asteraceae (*Atractylis*-type, Asteroideae, Cichoroideae, *Pentzia*-type, *Stoebe*-type, *Tar-chonanthus*-type), Boraginaceae (*Cordia*, *Heliotropium*), Burseraceae (*Boswellia*), Campanulaceae, Capparaceae (*Cadaba*, *Cleome*, *Maerua*-type), Caryophyllaceae, Convolvulaceae (*Convolvulus*-type, *Ipomoea*-type, *Merremia*), Crassulaceae (*Kalan-choe*), Cucurbitaceae (*Citrullus*), Euphorbiaceae (*Chrozophora*, *Euphorbia*), Fabaceae (*Crotalaria*, *Indigofera*-type, *Parkinsonia*, *Rhynchosia*-type, *Tephrosia*), Gerani-

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aceae (*Pelargonium*), Liliaceae (*Aloe*-type), Malvaceae, Neuradaceae, Nyctaginaceae (*Boerhavia*), Plantaginaceae (*Plantago*), Polygonaceae (*Calligonum*), Salvadoraceae (*Salvadora persica*), Sapindaceae (*Dodonaea viscosa*), Scrophulariaceae (*Walafrida*), Solanaceae (*Solanum*), Sterculiaceae (*Hermannia*, *Hermannia stricta*-type), Thymelaeaceae, Tiliaceae (*Corchorus*), Verbenaceae (*Phyla nodiflora*), Welwitschiaceae (*Welwitschia*), Zygophyllaceae (*Nitraria*, *Tribulus*, *Zygophyllum*)

Dry forest and savannah

Acanthaceae (*Blepharis*-type, *Hypoestes*-type) Anacardiaceae (*Lannea*-type, *Sclerocarya*-type *birrea*), Annonaceae, Araliaceae (*Cussonia*), Arecaceae (*Cocos nucifera*), Asteraceae (*Artemisia afra*-type), Balanitaceae (*Balanites*), Bignoniaceae (*Spathodea*), Bombacaceae (*Adansonia*, *Bombax*, *Ceiba pentandra*), Caesalpiniaceae (*Bauhinia*-type, *Burkea*, *Cassia*-type, *Colophospermum mopane*, *Detarium*, *Dialium*-type, *Piliostigma*), Capparaceae (*Boscia*-type, *Capparis*), Celastraceae (*Maytenus*-type), Clusiaceae (*Garcinia*), Conneraceae (*Cnestis*-type), Cucurbitaceae (*Coccoloba*, *Luffa*), Euphorbiaceae (*Acalypha*, *Antidesma*, *Bridelia*, *Croton*-type, *Hymenocardia*, *Securinega*), Fabaceae (*Baphia*-type, *Dalbergia*, *Erythrina*, *Ormocarpum*, Papilionoideae p.p., *Pterocarpus*-type, *Sesbania*-type), Flacourtiaceae (*Flacourtia*), Linaceae (*Hugonia*), Loganiaceae (*Strychnos*), Malphigiaceae (*Triapsis*), Meliaceae (*Pseudocedrela*), Mimosaceae (*Mimosa pigra*), Myrsinaceae (*Maesa lanceolata*-type), Nyctaginaceae p.p., Olacaceae (*Olax*), Oleaceae (*Jasminum/Schrebera*), Proteaceae (*Protea*), Ranunculaceae (*Clematis*-type), Rubiaceae (*Canthium*, *Kohautia*, *Psydrax*-type *subcordata*, Rubiaceae monade, *Rytigynia*-type, *Spermacoce*-type, *Stipularia africana*, *Tarenna*), Rutaceae (*Teclea*, *Zanthoxylum*), Sapindaceae (*Allophylus*, *Cardiospermum*), Sapotaceae (*Butyrospermum*), Sterculiaceae (*Dombeya*, *Sterculia*-type), Tiliaceae (*Grewia*)

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## Miombo woodland

Caesalpiniaceae (*Afzelia*, *Berlinia*-type, *Brachystegia*, *Paramacrolobium*), Euphorbiaceae (*Uapaca*), Fabaceae (*Pericopsis*)

## Mountain elements

- 5 *Aquifoliaceae* (*Ilex* cf. *I. mitis*), *Ericaceae*, *Geraniaceae* (*Geranium*), *Lamiaceae* (*Ocimum*), *Myricaceae* (*Myrica*), *Myrsinaceae* (*Myrsine africana*), *Oleaceae* (*Chionanthus*-type, *Olea*, *Olea capensis*), *Podocarpaceae* (*Podocarpus*), *Restionaceae*, *Sapotaceae* (*Chrysophyllum*),

## Wet lowland forest (rain forest)

- 10 *Anacardiaceae* (*Sorindeia*-type), *Apocynaceae* (*Adenium*, *Funtumia*-type), *Arecaceae* (*Elaeis guineensis*), *Bignoniaceae* (*Stereospermum*), *Caesalpiniaceae* (*Daniellia*-type, *Duparquetia orchidacea*, *Mezoneuron*-type, *Tamarindus/Cryptosepalum*, *Tessmannia*), *Celastraceae/Hippocrateaceae*, *Euphorbiaceae* (*Alchornea*, *Klaineanthus*, *Malolotus*-type, *Martretia*, *Phyllanthus*, *Tetrorchidium*),  *Icacinaceae* (*Iodes*), *Leeaceae*
- 15 (*Leea*), *Loranthaceae* (*Tapinanthus*-type), *Meliaceae* (*Khaya*-type, *Manilkara*), *Mimosaceae* (*Entada*-type, *Pentaclethra*), *Moraceae* p.p., (*Bosqueia*), *Myristicaceae* (*Pycnanthus*), *Myrtaceae* (*Syzygium*), *Ochnaceae* (*Lophira*, *Sauvagesia*), *Araliaceae* (*Schefflera*), *Rubiaceae* (*Coffea*-type, *Crossopteryx*, *Cuviera*, *Gaertnera*, *Ixora*-type, *Morinda*, *Mussaenda*, *Psydrax*-type *parviflora*, *Sabicea*, *Sherbournea*), *Sapindaceae*
- 20 (*Blighia*-type), *Sterculiaceae* (*Cola cordifolia*, *Cola nitida*-type, *Melochia*), *Ulmaceae* (*Celtis*, *Holoptelea*), *Urticaceae*, *Vitaceae* (*Cissus*)

## Swamps

*Acanthaceae* (*Hygrophila*-type), *Agavaceae* (*Dracaena*), *Alismataceae* (*Alisma plan-tago-aquatica* ), *Arecaceae* (*Borassus*-type, *Hyphaene*-type, *Raphia*-type), *Bignoni-*

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aceae (*Kigelia*-type), Burseraceae (*Canarium*), Caesalpiniaceae (*Monopetalanthus*-  
type), Clusiaceae/Guttiferae (*Symphonia globulifera*), Dispacaceae (*Cephalaria*),  
Droseraceae (*Drosera*), Euphorbiaceae (*Macaranga*), Fabaceae (*Aeschynomene*),  
Halorrhagaceae (*Myriophyllum*), Irvingiaceae (*Irvingia*), Meliaceae (*Trichilia*), Mi-  
mosaceae (*Pseudoprosopis*), Nymphaeaceae (*Nymphaea*), Onagraceae (*Ludwigia*),  
Passifloraceae (*Adenia*), Potamogetaceae (*Potamogeton*), Rhamnaceae, Rubiaceae  
(*Gardenia*, *Psychotria*), Salicaceae (*Salix*), Sapindaceae (*Aporrhiza*), Typhaceae (*Ty-  
pha angustifolia*-type, *Typha latifolia*), Utriculariaceae (*Utricularia*), Xyridaceae (*Xyris*)

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## L. M. Dupont et al.

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vegetation development**Table 1.** AMS radiocarbon dates of ODP Site 1078C below 980 mbsf. Calibration using  $R=400$  after Fairbanks et al. (2005).

Lab. No	Core/ Section	Interval (cm)	Depth (mbsf)	$^{14}\text{C}$ Age (a)	Error	Cal. Age range (ka)
KIA13030	2H-3	20–22	991	21 340	+220/–210	24 799–25 415
KIA13008	2H-3	30–32	1001	22 780	+260/–250	26 586–27 146
KIA13028*	2H-3	50–52	1021	16 960	+/- 130	19 522–19 838
KIA13027*	2H-3	100–102	1071	22 850	+260/–250	26 661–27 220
KIA13023	2H-4	70–72	1191	30 250	+630/–590	34 282–35 567
KIA13038	2H-5	60–62	1331	33 880	+1010/–890	37 357–39 780
KIA13007	2H-5	130–132	1401	36 780	+1560/–1300	40 444–42 791
KIA13006	2H-CC	17–19	1648	44 850	+5190/–3130	>45 000

\* dates not used in the age model

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**Table 2.** Percentage mean and standard deviation per zone for selected pollen taxa and groups (listed in Appendix A). “LGM” period of lowest sea level at the site (see Sect. 4).

Pollen zone	Holocene				Deglaciation				Glacial				“LGM”			
	7	6	5	4	3	2	1	“LGM”	7	6	5	4	3	2	1	“LGM”
upper boundary (m)	0.00	1.50	2.80	4.20	8.02	8.80	9.70	9.81–9.99								
upper boundary (ka)	top	3.7	7.8	10.0	15.4	18.8	21.9	22.4–26.3								
Pollen taxon	mean	stdev	mean	stdev	mean	stdev	mean	stdev	mean	stdev	mean	stdev	mean	stdev	mean	stdev
Mountain el.	1.1	0.7	0.8	0.6	1.5	0.8	2.6	1.0	2.8	1.7	<b>4.2</b>	1.4	<b>7.4</b>	1.5	5.9	1.6
Ericaceae	0.2	0.3	0.6	0.6	0.6	0.4	0.9	0.7	1.4	1.1	<b>2.2</b>	1.1	<b>5.4</b>	1.4	4.3	1.1
<i>Phaecoceros</i>	0.2	0.3	0.4	0.3	0.5	0.4	0.6	0.6	2.1	1.2	<b>2.8</b>	1.3	<b>3.0</b>	0.9	3.0	1.0
<i>Brachystegia</i>	0.9	0.6	<b>5.1</b>	2.0	<b>1.9</b>	0.9	<b>1.5</b>	0.9	0.8	0.6	1.0	0.6	<b>1.5</b>	0.9	1.4	0.9
Desert el.	3.0	1.0	2.4	1.0	2.0	0.9	3.7	1.6	6.4	1.9	<b>7.6</b>	1.6	<b>8.3</b>	1.6	9.5	1.1
Asteroidae	1.1	0.5	0.7	0.4	0.6	0.4	1.6	0.8	2.6	1.2	<b>3.3</b>	1.1	<b>4.0</b>	1.2	4.5	1.4
Poaceae	<b>38.2</b>	12.5	9.6	1.5	17.1	3.3	25.6	3.8	26.5	7.0	<b>31.7</b>	6.7	<b>35.2</b>	3.6	33.2	2.6
<i>Rhizophora</i>	<b>2.7</b>	1.4	0.6	0.6	<b>6.0</b>	2.3	<b>3.2</b>	1.3	<b>1.5</b>	1.0	<b>0.9</b>	0.7	0.1	0.2	0.1	0.2
Rainforest el.	<b>4.5</b>	1.2	2.3	1.0	<b>3.5</b>	1.3	<b>3.9</b>	1.7	<b>2.8</b>	1.1	<b>3.1</b>	1.2	1.6	0.7	1.7	0.8
<i>Mallotus</i> -type	0.7	0.6	0.4	0.3	0.6	0.5	<b>2.5</b>	1.7	<b>1.5</b>	0.9	<b>1.6</b>	1.0	0.3	0.4	0.6	0.7
<i>Myrica</i>	0.2	0.3	0.0	0.1	0.3	0.2	0.5	0.3	0.8	0.6	<b>1.2</b>	0.6	0.8	0.5	0.9	0.4
<i>Podocarpus</i>	13.2	10.0	<b>40.9</b>	4.1	28.3	3.7	33.0	4.8	<b>33.1</b>	11.5	22.0	5.7	14.8	3.0	16.7	2.6
Dry forest el.	<b>7.6</b>	1.8	<b>5.7</b>	2.3	<b>5.7</b>	1.8	3.4	1.3	2.8	1.1	2.9	1.1	3.1	1.3	3.9	1.2
<i>Burkea</i> -type	<b>1.4</b>	1.0	0.4	0.4	<b>1.2</b>	0.7	<b>0.7</b>	0.6	0.2	0.2	0.2	0.2	0.1	0.2	0.1	0.2
Combretaceae	<b>0.6</b>	0.5	0.3	0.3	0.5	0.4	<b>0.6</b>	0.4	0.1	0.2	0.1	0.2	0.2	0.3	0.2	0.2
Miombo el.	2.8	1.3	<b>8.8</b>	2.3	<b>5.1</b>	1.8	<b>3.2</b>	1.5	1.7	1.0	1.7	1.1	2.2	0.9	2.0	0.8
<i>Uapaca</i>	1.2	0.7	<b>1.7</b>	1.3	<b>2.4</b>	0.8	<b>1.3</b>	0.7	0.3	0.4	0.1	0.2	0.2	0.2	0.1	0.1
Cyperaceae	<b>12.9</b>	2.6	5.6	1.1	<b>8.4</b>	2.1	5.6	1.9	8.2	2.8	10.0	2.3	9.2	2.5	10.2	3.4
<i>Alchornea</i>	<b>1.9</b>	1.0	0.6	0.5	<b>1.1</b>	0.7	0.3	0.3	0.1	0.2	0.1	0.2	0.1	0.2	0.2	0.4
<i>Artemisia afra</i>	<b>1.4</b>	1.0	0.1	0.2	<b>1.0</b>	0.8	0.0	0.1	0.1	0.2	0.1	0.2	0.2	0.3	0.2	0.3
<i>Azella</i>	0.5	0.5	<b>1.6</b>	1.0	0.4	0.5	0.2	0.3	0.4	0.4	0.5	0.5	0.3	0.3	0.3	0.2
<i>Berlinia</i> -type	0.2	0.4	<b>0.5</b>	0.6	0.1	0.2	0.0	0.1	0.0	0.1	0.0	0.1	0.1	0.1	0.0	0.0
Rhamnaceae	0.2	0.2	<b>0.6</b>	0.5	0.2	0.3	0.0	0.0	0.1	0.2	0.1	0.2	0.1	0.2	0.1	0.2
<i>Pteris</i>	0.6	0.5	<b>3.4</b>	1.5	1.7	0.8	1.4	0.9	1.1	0.6	1.2	0.7	1.2	0.5	0.9	0.6
<i>Hymenocardia</i>	<b>1.2</b>	0.8	0.1	0.2	0.2	0.3	0.0	0.1	0.0	0.1	0.0	0.1	0.0	0.1	0.0	0.0

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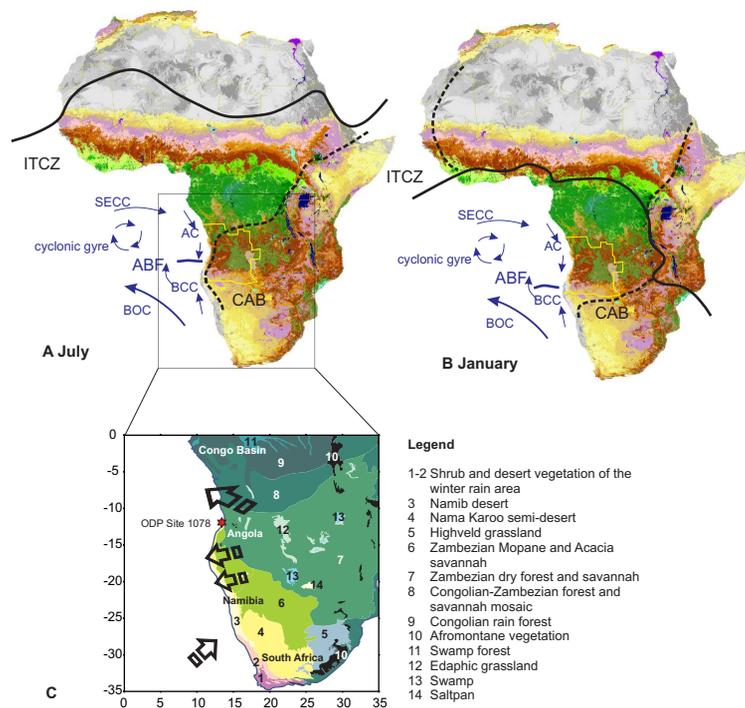
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**Fig. 1.** Landcover of Africa (Bartholome et al., 2002), simplified phytogeography of southern Africa (White, 1983), ocean surface currents of the eastern South Atlantic (Shannon and Nelsson, 1996), and summer and winter position of important tropical frontal systems (Leroux, 1983). ITCZ, Intertropical Convergence Zone; CAB, Congo Air Boundary; ABF, Angola Benguela Front; SECC, South equatorial Counter Current; AC, Angola Current; BCC, Benguela Coastal Current; BOC, Benguela Ocean Current. **(A):** Austral winter (July) position of the ITCZ, CAB and ABF. **(B):** Austral summer (January) position of the ITCZ, CAB, and ABF. **(C):** Location of ODP Site 1078, general surface (air pressure level of 850 hPa) wind directions in austral winter (June, July, August), and biomes of southern Africa (see legend).

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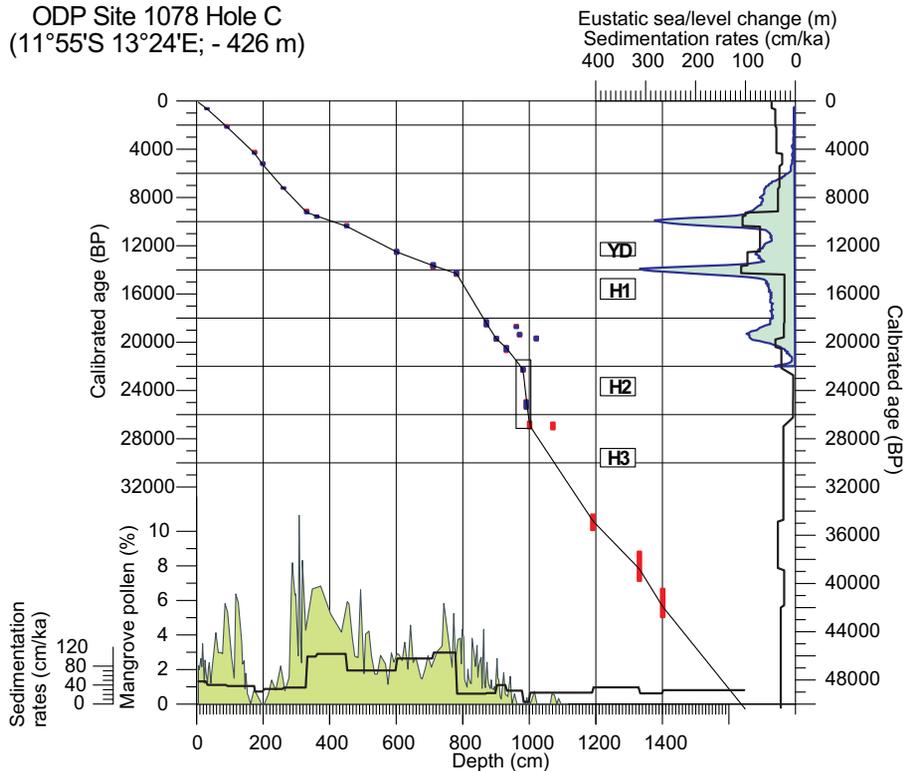
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ODP Site 1078 Hole C  
(11°55'S 13°24'E; - 426 m)



**Fig. 2.** Age model for ODP Hole 1078C. Denoted in blue are the age ranges calibrated with Calib using the marine04 calibration set (Hughen et al., 2004). Denoted in red are the age ranges calibrated after Fairbanks et al. (2005). Bottom shaded curve and left axis, pollen percent of *Rhizophora* (mangrove tree); bottom line and far left axis, sedimentation rates against depth; right shaded curve and upper axis, eustatic sea-level change after Fleming et al. (1998); right line and upper axis, sedimentation rates against age. YD, Younger Dryas; H1, H2, H3, Heinrich Events 1, 2, 3, respectively.

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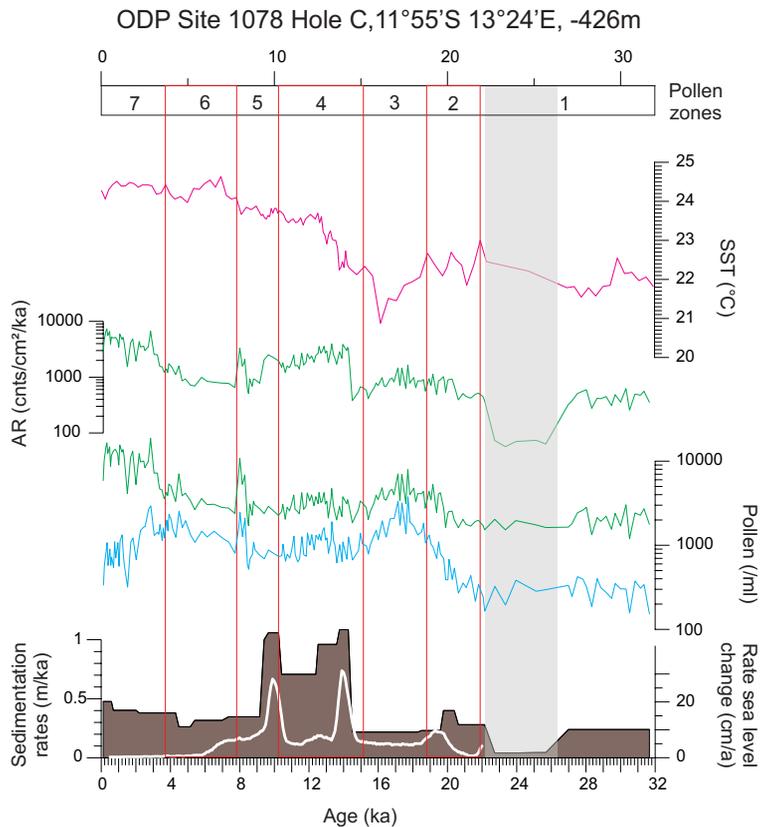
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**Fig. 3.** From top to bottom. Pollen zones. Alkenone based sea surface temperatures (red line, upper right axis). Accumulation rates of total pollen per cm<sup>2</sup> and ka (green line, upper left axis). Total pollen concentration (green line) and Podocarpus pollen concentration (blue line) per ml sediment (middle right axis). Sedimentation rates in m per ka (shaded, lower left axis) and rates of sea-level change (line, lower right axis) after Fleming et al. (1998).

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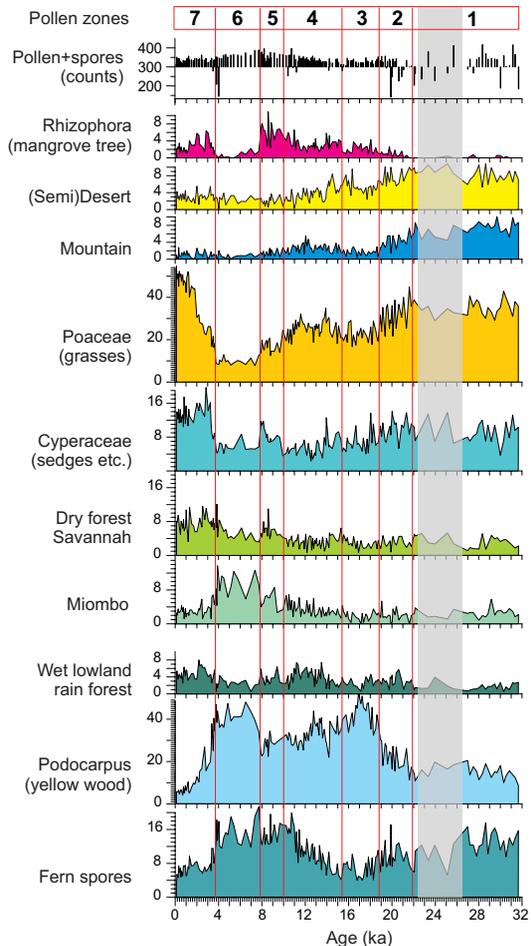
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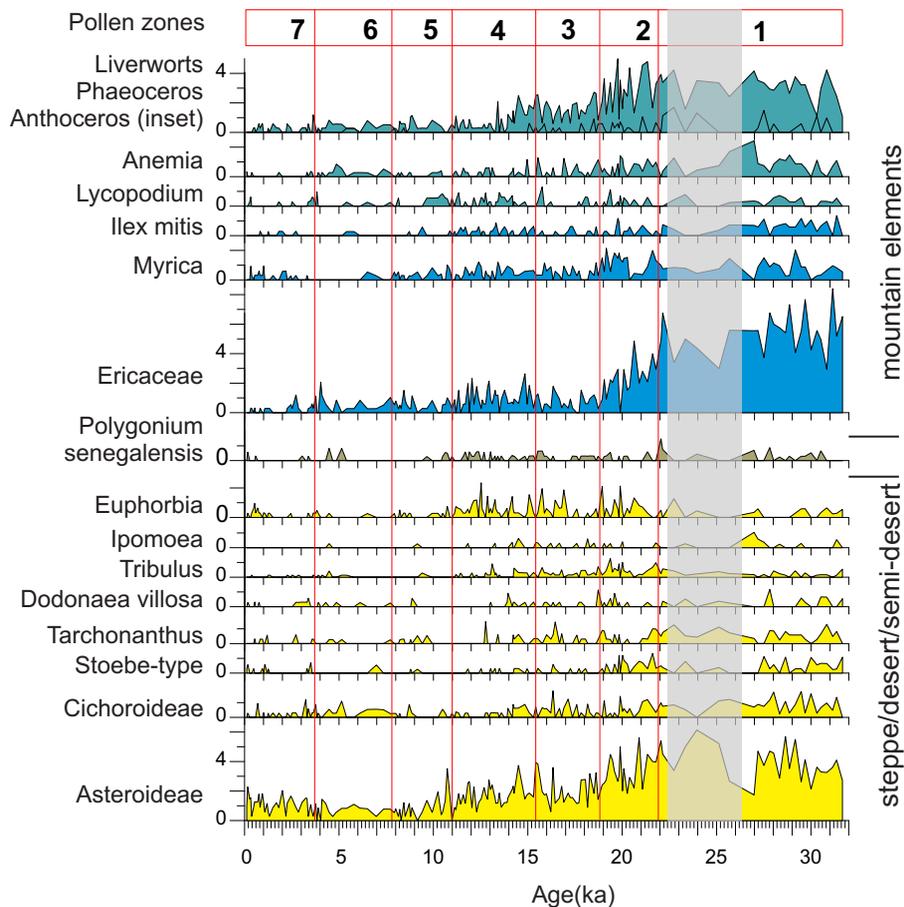
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**Fig. 4.** Percentages of pollen taxa and groups (listed in Appendix A). Counts of total pollen and fern spores being the basis of the percent calculation are given on top.

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**Fig. 5.** Percentages of selected pollen taxa from desert, semi-desert, or open mountainous vegetation (pollen sums are give in Fig. 4).

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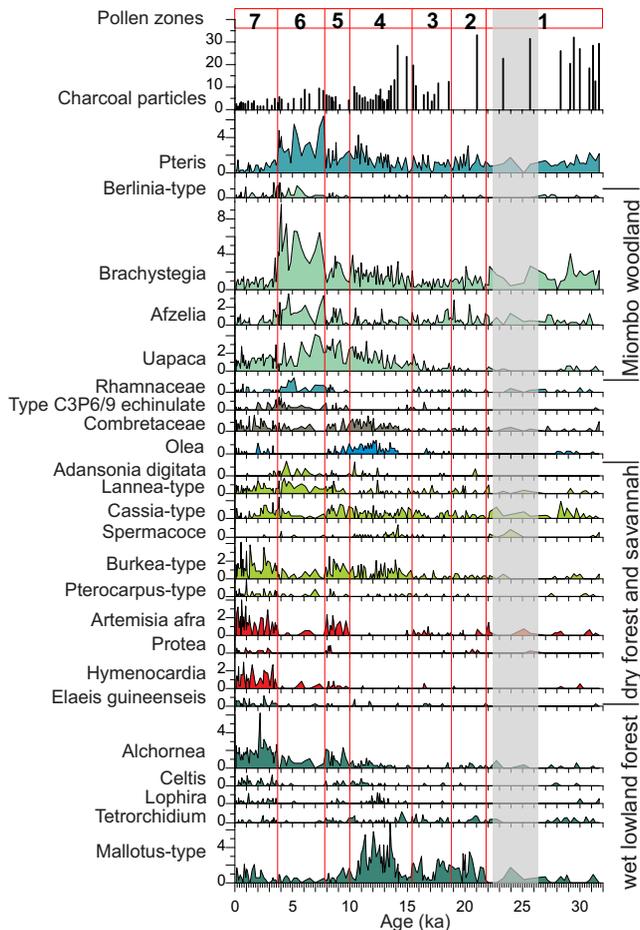
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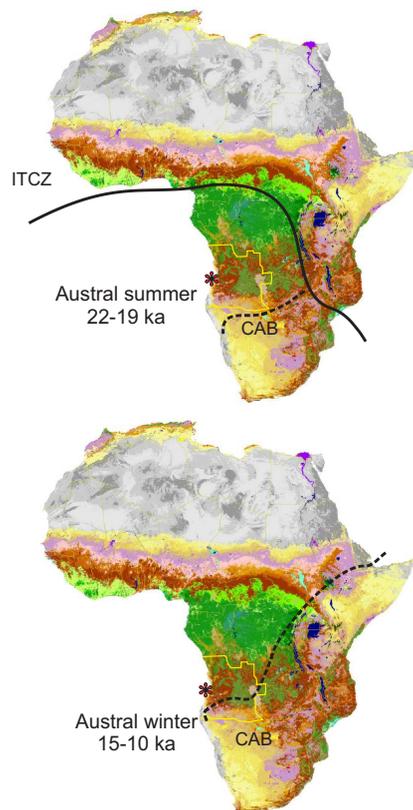
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**Fig. 6.** Percentages of selected pollen taxa from rain forest, dry forest and woodland, and Miombo woodland (pollen sums are given in Fig. 4). Charcoal percentages are also expressed on the total of pollen and spores.

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**Fig. 7.** Conceptual model of the latitudinal positions of the Congo Air Boundary (CAB) and the Intertropical Convergence Zone (ITCZ) during the deglaciation projected on the modern land cover map. Top, a southern position of the ITCZ during austral summer would favour forest development in Angola and savannahs in Cameroon. Bottom, a southern position of the CAB during austral winter would bring moisture during most of the year to large areas in Angola favouring an early establishment of rain forest.

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