



HAL
open science

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann, M. Butzin, A. Micheels, T. Bickert, V. Mosbrugger

► **To cite this version:**

G. Lohmann, M. Butzin, A. Micheels, T. Bickert, V. Mosbrugger. Effect of vegetation on the Late Miocene ocean circulation. *Climate of the Past Discussions*, 2006, 2 (4), pp.605-631. hal-00298144

HAL Id: hal-00298144

<https://hal.science/hal-00298144>

Submitted on 18 Jun 2008

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Climate of the Past Discussions is the access reviewed discussion forum of *Climate of the Past*

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann¹, M. Butzin², A. Micheels³, T. Bickert³, and V. Mosbrugger³

¹Alfred-Wegener-Institut für Polar- und Meeresforschung, Bussestr. 24, 27570 Bremerhaven, Germany

²University of Bremen, P.O. Box 330440, 28334 Bremen, Germany

³Senckenberg Research Institute and Natural History Museum, Senckenberganlage 25, 60325 Frankfurt am Main, Germany

Received: 4 August 2006 – Accepted: 17 August 2006 – Published: 23 August 2006

Correspondence to: G. Lohmann (Lohmann@awi-bremerhaven.de)

CPD

2, 605–631, 2006

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

Abstract

A weak and shallow thermohaline circulation in the North Atlantic Ocean is related to an open Central American gateway and exchange with fresh Pacific waters. We estimate the effect of vegetation on the ocean general circulation using the atmospheric circulation model simulations for the Late Miocene climate. Caused by an increase in net evaporation in the Miocene North Atlantic, the North Atlantic water becomes more saline which enhances the overturning circulation and thus the northward heat transport. This effect reveals a potentially important feedback between the ocean circulation, the hydrological cycle and the land surface cover for Cenozoic climate evolution.

1 Introduction

The Eocene-Oligocene and the Mid-Miocene climate transitions are two major cooling steps in the Cenozoic climate evolution (Zachos et al., 2001) from greenhouse to “icehouse” climate conditions. A drastic increase in the oxygen-isotopic composition measured in benthic foraminifer shells points to a combination of Antarctic ice growth and global cooling at 34 Ma and 14 Ma (Ma: million years before present), respectively, which is also indicated by the occurrence of Southern Ocean ice-rafted detritus and eustatic sea-level change (Miller et al., 1987; Kennett and Barker, 1990; Billups and Schrag, 2002). Ocean circulation changes and atmospheric pCO₂ variations are often cited as potential catalysts of these cooling events (DeConto and Pollard, 2003). Large-scale ocean circulation changes, caused by atmospheric circulation changes and/or by tectonic reorganizations of gateway regions, may have altered poleward transports of heat and moisture, which in turn may have resulted in Antarctic ice growth and global cooling (Kennett, 1977; Zachos et al., 2001). Ocean circulation hypotheses are supported by C-13 proxy evidence (e.g. Wright and Miller, 1996; Billups, 2002) and the timing of tectonic events at critical ocean pathways like the Drake Passage, the Tasmanian Seaway, the Indonesian Throughflow (Cane and Molnar, 2001; Lawver and Gaha-

CPD

2, 605–631, 2006

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

gan, 2003), the eastern Tethys (Flower and Kennett, 1994), and the Central American Seaway (e.g., Haug and Tiedemann, 1998).

Here, we will examine climate processes in connection with large-scale ocean circulation changes for a selected Cenozoic time slice, namely the Late Miocene or Tortonian (11–7 Ma). The Tortonian was characterized by intensive Antarctic glaciation and the buildup of ice sheets in the North Atlantic realm. Specifically, we focus on the spatial temperature distribution, which is a principal problem in understanding Cenozoic climate change. In the case of the Miocene, elevated global-mean surface temperatures and weak equator-to-pole temperature gradients are proposed (Greenwood and Wing, 1995; Crowley and Zachos, 2000). While numerical simulations exhibit rising global-mean temperatures for increasing greenhouse gas concentrations, they fall far short of attaining the reconstructed reduction in the meridional temperature gradient (Barron, 1987; Huber and Sloan, 2001). Since it seems that atmospheric carbon dioxide concentration hardly varied during the Miocene (Pagani et al., 1999, 2005, Pearson and Palmer, 2000), the mechanism which causes bipolar glaciation in the Tortonian remains even more enigmatic.

Some authors (Schmidt and Mysak, 1996; Hay et al., 1997) have suggested that atmospheric heat transport may have played an important role in resolving this “low gradient paradox”. It is plausible to expect a warmer atmosphere to transport more latent heat poleward, helping to reduce meridional temperature gradients. However, despite the exponential increase of saturation vapor pressure with temperature, this feedback becomes less powerful as temperature rises (Caballero and Langen, 2005). Other possible mechanisms located in the atmosphere involve the atmospheric stationary wave response due to changing paleogeography and sea level.

On the other side, marine proxy data indicate that ocean gateway changes and major reorganizations of the global ocean circulation (e.g. Kennett, 1977; Wright et al., 1992; Zachos et al., 2001) are consistent with a weakening of the ocean heat transport during the Miocene. Concerning the Tortonian, the (then) still opened Central American Seaway (CAS, i.e., the Panama Strait) allowed for the exchange of saline Atlantic water

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

with comparatively fresher Pacific water, and it has been shown that this leads to weakening of the thermohaline circulation in the North Atlantic Ocean (e.g., Mikolajewicz et al., 1993; Bice et al., 2000; Butzin et al., 2006¹). Therefore, the global ocean circulation seems not to be a proper candidate to be responsible for a weaker equator-to-pole temperature gradient.

The question of temperature gradients might be linked to other feedbacks in the climate system, such as changes in the hydrological cycle and vegetation cover. Paleontological and palynological data give evidence for drastic changes in vegetation and therefore climate during the Cenozoic (Retallack, 2001; Willis and Mc Elwain, 2002). For example, during the Eocene/Oligocene glaciation tropical rain forests virtually disappeared poleward of the northern and southern high pressure zones. Grasslands, which had begun to develop under dry conditions during the Eocene, became more and more widespread in the Oligocene. During the Mid-Miocene Climatic Optimum, moist warm forests expanded poleward of the subtropical high pressure zones for a short period. Following the global climatic deterioration after the Mid-Miocene Climatic Optimum, tropical rain forests withdrew again to the equatorial zone. Grasslands and deserts expanded through much of the lower mid-latitudes (Morley, 2000; Bredenkamp et al., 2002). C4 grasslands became widespread during the interval from about 8 to 5 Ma (Cerling et al., 1997; Freeman and Colarusso, 2001). During the Miocene, most of the climatically arranged vegetation belts developed ranging from rain forest along the equator to polar desert at high latitudes. However, to date, little is known about the role of continental vegetation for climate change during the Cenozoic. It is still an open question whether the vegetation just has adapted to hydrological changes or whether it has played an active role as a modifier of major climate transitions. In principle, the vegetation can contribute to a weaker-than-present meridional temperature gradient through modifying the local albedo (e.g., Dutton and Baron, 1997; Otto-Bliesner and

¹Butzin, M., Lohmann, G., and Bickert, T.: Effect of ocean gateways onto the evolution of the ocean circulation and marine carbon cycle during the Miocene: Sensitivity studies with an ocean-circulation carbon cycle model, *Paleoceanography*, submitted, 2006.

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Upchurch, 1997).

In the light of these findings, we investigate whether such a feedback was effective during the Late Miocene. We examine if the Tortonian vegetation significantly enhanced the hydrological cycle with increased precipitation rates over continental areas providing for a greener land surface. In particular, we are interested in the climate sensitivity of the thermohaline circulation (THC) with the vegetation cover and associated hydrological cycle. For that purpose, we apply an atmospheric circulation model (AGCM) in combination with a coarse resolution model of the ocean. A dynamical vegetation model is used to evaluate the consistency between reconstructed and simulated vegetation cover. The models and experiments are briefly described in the following Section.

2 Methods

2.1 Atmospheric circulation model

For the Late Miocene climate simulations, we apply the atmosphere general circulation model ECHAM4 (Roeckner et al., 1996). The prognostic variables are calculated in the spectral domain with a triangular truncation at wave number 30 (T30), which corresponds to a Gaussian longitude-latitude grid of approximately 3.75° . The vertical domain is represented by 19 hybrid sigma-pressure (terrain following) levels with the highest level at 10 hPa. The model is coupled to a 50 m slab ocean. This allows a prescription of the Miocene ocean heat transport consistent with proxy data (Steppuhn et al., 2006). Furthermore, the orography is adapted to the Tortonian when the height of mountain ranges was generally reduced. For example, Greenland reaches only about a tenth of its recent elevation. In addition to the above-described boundary conditions, the atmospheric CO_2 is specified with the present-day level of 353 ppmv for all experiments. This lies within the spectrum of values which are given for the Miocene (Cerling et al., 1997, Pagani et al., 1999; Pearson and Palmer, 2000). For the land surface, sen-

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

sitivity experiments were performed which are described below. Each model simulation with the AGCM was run over 20 years. The model reaches an equilibrium state after 5 years, and the last 10 years are taken into account for further analysis. A list of the experiments is given here:

5 **CTRL:** Present day control simulation (Roeckner et al., 1996).

TGEO: Tortonian simulation with adapted geography (Steppuhn et al., 2006). The global vegetation represents modern conditions, except that the recent Greenland ice cap is replaced by tundra vegetation.

10 **TVEG:** Tortonian simulation with adapted geography as in TGEO and reconstructed vegetation cover. The Tortonian vegetation was reconstructed on the basis of palaeobotanical data such as fossil pollen and leaf data, and fossil car-poflora (Micheels, 2003). Figure 1 shows the resulting reconstruction of the global Tortonian vegetation. The Tortonian palaeovegetation was generally more lush as compared to today, tropical forests expanded and their margins shifted further poleward. According to the reconstruction of the Tortonian vegetation, land sur-
15 face parameters are adapted. To consider the changed vegetation in the model, data for the albedo, the leaf area index, the vegetation and forest cover, and the maximum soil water capacity are changed.

2.2 Dynamical vegetation model

20 The LPJ dynamical vegetation model (Sitch et al., 2003) combines process-based descriptions of terrestrial ecosystem structure (vegetation composition, biomass and height) and function (energy absorption, carbon cycling). Vegetation composition is described by nine different plant functional types (PFTs), which are distinguished ac-
25 cording to their physiological (C3, C4 photosynthesis), morphological (tree, grass) and phenological (deciduous, evergreen) attributes. The model is run on a grid cell basis with input of soil texture, monthly fields of temperature, precipitation, as well as short

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

and long wave radiation. Each grid cell is divided into fractions covered by the PFTs and bare ground. Both the presence and the covered fraction of PFTs within a grid cell depend on their specific environmental limits and on resource competition among the PFTs. Carbon isotope fractionation is included in the model (Kaplan et al., 2002; Scholze et al., 2003). The model is run on a horizontal $2^\circ \times 2^\circ$ grid, directly forced with the output of the AGCM experiments.

2.3 Ocean circulation model

Our ocean model is an updated version of the LSG circulation model developed by Maier-Reimer et al. (1993). We implemented some significant improvements such as a new advection scheme for tracers (Schäfer-Neth and Paul, 2001; Prange et al., 2003) as well as an overflow parametrization for the bottom boundary layer (Lohmann, 1998; Lohmann and Schulz, 2000). The spatial resolution is $3.5^\circ \times 3.5^\circ$ in the horizontal and 22 levels in the vertical. We calibrated the model by simulating anthropogenic C-14 (Butzin et al., 2005). The ocean is forced by ten-year averaged monthly fields of wind stress, surface air temperature, and freshwater flux, which serve as background climatology and originate from the simulations with the atmosphere general circulation model ECHAM4 described in Sect. 2.1. A surface heat flux formulation based on atmospheric energy balance model considerations permits that sea surface temperatures (SST) can freely adjust to ocean circulation changes (e.g., see Prange et al., 2003; Knorr and Lohmann, 2003; Butzin et al., 2005). The hydrological cycle is closed by a runoff scheme which considers continental catchment areas and allows for variable land-sea distributions, which permits that sea surface salinities (SSS) can freely evolve. The total integration time of each experiment is 5000 years. For the late Miocene simulations, we assumed a 500 m deep and three gridpoints wide (between 9° N and 18° N) gateway between the Atlantic and Pacific Oceans.

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

3 Results

3.1 Hydrological cycle and vegetation cover

The control climate simulates the mean hydrological cycle reasonably well as shown by Arpe et al. (2000) and is in agreement with observations (e.g., Peixoto and Oort, 1992; Zaucker and Broecker, 1992). The subtropical highs over the North and South Atlantic and Pacific oceans provide a moisture transport from the subtropics to higher latitudes. In the tropics between 20° S and 20° N, strong easterlies are observed, especially over the Atlantic and Pacific Oceans.

Figure 2 indicates strong changes in the hydrological cycle when comparing TVEG and CTRL. Boreal summer precipitation over the Sahel region is strongly increased for the green Sahara (compare Fig. 1). In accordance with the removal of the inland ice of Greenland, sea ice is drastically reduced caused by a considerably increasing surface temperatures and the ice-albedo feedback, and local precipitation is increased over Northern Greenland (Fig. 2c). The Icelandic Low is slightly shifted to the south-east leading to more precipitation off western Europe and less precipitation between Greenland and Iceland (Fig. 2).

The reduced ocean heat transport causes a southward migration of the thermal equator in both Tortonian simulations TGEO and TVEG. When comparing TVEG with CTRL in Fig. 2, the Intertropical Convergence Zone moves southward resulting in enhanced water vapor export out of the Atlantic catchment area. We evaluate an additional moisture transport from the Atlantic to the Pacific Ocean accounting for an increase of net Atlantic evaporation (0.12 and 0.31 Sv for TGEO and TVEG, respectively). The unit 1 Sv corresponds to a mass transport of 10^9 kg s^{-1} , equivalent to a volume transport of $10^6 \text{ m}^3 \text{ s}^{-1}$ liquid water.

In order to check the consistency of the reconstructed vegetation distribution with the modelled climate in TVEG, we apply the dynamical vegetation model LPJ. We use the monthly output of the last 10 years of the CTRL and TVEG simulations, iterating these simulations 200 times in order to get an equilibrium of the dynamical vegetation model

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Effect of vegetation
on the Late Miocene
ocean circulation**G. Lohmann et al.

after 2000 model years. We build an average over the last 500 years and identify the spatial patterns of the PFTs (in %) for the Tortonian and present-day vegetation cover (Fig. 3). For the late Miocene, tropical trees are spread in the subtropical Africa (North and South) and parts of Australia (Fig. 3a), whereas temperate trees are extended over Asia (Fig. 3b) relative to present conditions. The extension of boreal forests far into the northern high latitudes during the Tortonian (Fig. 3c) is in accordance with proxy data (Boulter and Manum, 1997). Grassland is extended into subtropical areas, over Greenland and over Alaska. The Sahara desert is smaller than today and consists of steppe and open grassland rather than sand desert which is consistent to fossil data (Le Houerou, 1997; Schuster et al., 2006).

3.2 Ocean circulation

In the ocean circulation experiments, we employ a hybrid coupled modeling approach, which allows an adjustment of surface temperatures and salinity to changes in the ocean circulation, based on an atmospheric energy balance model (Lohmann and Gerdes, 1998; Prange et al., 2002). No flux correction is applied for present day and other climate conditions. The control experiment for present-day conditions (Fig. 4a) reasonably reflects the modern Atlantic Ocean circulation with a southward water export of 16 Sv at 30° S and a heat transport of 0.96 PW (1 PW=10¹⁵ W) at 30° N, which is in the range of oceanographic observations (Schmitz, 1995; Macdonald and Wunsch, 1996).

A comparison of the control run with the Tortonian experiments (TGEO, TVEG) reveals significant changes in the meridional overturning circulation (Fig. 4bc): The formation of deep water in the North Atlantic is strongly reduced (TGEO) when the Central American Seaway (CAS) is open (Fig. 4). The meridional circulation is only 3 Sv and represents a “mini-conveyor belt” circulation with an ocean heat transport at 30° N of 0.19 PW (Fig. 4b). In experiment TVEG, the circulation strength is similar to the present-day circulation (14 Sv export at 30° S, 0.83 PW at 30° N), but slightly shallower than under present-day conditions. The reason might be the increased flow of bottom

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

water from the Antarctic (Fig. 4c).

A detailed analysis of the flow patterns in various depths of the Panamanian gateway shows an export of surface water from the Atlantic to the Pacific Ocean (Fig. 5a). An import of thermocline and intermediate layer water from the Pacific to the Atlantic Ocean is responsible for a reversal of the Northeast Brazil Current (Fig. 5a). The net flux of Pacific water through the CAS into the Atlantic leads to relatively low-salinity thermocline water which hinders deep water formation in the North Atlantic. In T GEO, the surface winds and net freshwater flux in the North Atlantic are not able to overcome this freshening (Fig. 4b), whereas the background conditions in TVEG with stronger northward flow (Fig. 5b) and increase in net evaporation are sufficient to push the ocean circulation into a present-day-like circulation mode (Fig. 4c). Both the increased ocean circulation with a northward shift of the Arctic sea ice, and a local warming associated to the land surface quantities, induce an anomalous warming between TVEG and T GEO of up to 8°C (Fig. 5b).

Caused by the drop in ocean circulation in T GEO relative to CTRL, the sea surface salinity in the North Atlantic is considerably reduced (Fig. 6a). Due to the exchange of surface water close to the CAS, the surface water in the tropical Pacific becomes more saline. In contrast, the stronger ocean circulation for TVEG as compared to T GEO and the increased net evaporation yield considerably higher sea surface salinities in the North Atlantic Ocean (Fig. 6b). The strong increase in North Atlantic upper 500 m salinity is clearly emphasized in the Atlantic zonal-mean salinity distribution (Fig. 7a). The surface and subsurface warming of TVEG relative to T GEO is strongest in the subtropics and polar latitudes (Fig. 7b). At northern polar latitudes, the warming is associated to strong poleward surface currents (Fig. 5b), sea ice retreat, and meridional heat transport.

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

4 Discussion

The vegetation effect on the ocean circulation may be an important mechanism for the relatively warm late Miocene climate over Europe as reconstructed by terrestrial proxy data (e.g., Mai, 1995; Wolfe, 1994; Fortelius et al., 2003). Sensitivity experiments with atmospheric general circulation models demonstrate that the late Miocene vegetation contributes to a warming of particularly the high latitudes (Dutton and Barron, 1997; Micheels, 2003; Francois et al., 2006; Micheels et al., 2006).

Here, we analyse the hydrological cycle and associated vegetation cover onto the ocean circulation. The atmospheric hydrological cycle has a high mobility and links the THC with the Earth's water budget. We find that the Intertropical Convergence Zone moves southward resulting in enhanced water vapor export out of the Atlantic catchment area. A similar effect has been proposed for tropical water vapor transport during glacials (Lohmann and Lorenz, 2000) and Heinrich events (Lohmann, 2003) which may be responsible for an additional sea surface salinity contrast between the Atlantic and Pacific/Indian Oceans (Broecker, 1992), as well as for El Niño conditions (Schmittner et al., 2000; Soden, 2000; Latif et al., 2000). We find that the water vapor transport out of the Atlantic area is enhanced for the Tortonian climate relative to the control experiment. The net Atlantic freshwater forcing has been recognized as an important external parameter of ocean sensitivity studies (e.g., Birchfield, 1989; Zaucker et al., 1994; Rahmstorf, 1996). The increased export at Central America is caused by an increase in the zonal moisture transport associated to the Atlantic trade winds.

As pointed out by Steppuhn et al. (2006), there is a significant warming of more than 2°C at the eastern margin of the Pacific Ocean associated with a decreased upwelling in this area. This is again linked to the southward shift of the thermal equator, the ITCZ and weaker equatorial Walker circulation. The latter gives rise to a Tortonian permanent El-Niño state. This aspect will be analyzed in a subsequent study using a coupled atmosphere-ocean circulation model for the late Miocene. Fedorov et al. (2006) proposed that a permanent El-Niño state may be important for Pliocene glaciation and

CPD

2, 605–631, 2006

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

Cenozoic climate evolution.

We find that the vegetation effect on the ocean circulation can be an important mechanism for the relatively warm late Miocene climate over Europe. Caused by high salinities at northern high latitudes, the sea ice edge is moved poleward which is in general agreement with proxy data (e.g., Wolf and Thiede, 1991). In addition, it is possible that other mechanisms not included in the present generation of GCMs also had an important impact on Tortonian climate, such as high-latitude radiative warming by polar stratospheric clouds (Sloan and Pollard, 1998), increased ocean heat transport driven by tropical cyclone-induced mixing (Emanuel, 2002; Huber, personal communication), or increased levels of methane. Methane can be estimated through stable carbon isotopes (biological processes preferentially incorporate C-12) and areas of wetlands as calculated from the land surface scheme including the vegetation distribution.

5 Conclusions

The Cenozoic climate evolution includes significant changes in the oceanic transports which are ultimately linked to the paleotopography and opening/closing of passages. The open Central American Seaway leads to an exchange of fresh Pacific water with saline Atlantic water thereby reducing the density in the North Atlantic Ocean and weakening of the large-scale ocean circulation. For the Late Miocene, we find that the modified vegetation cover can compensate this gateway effect by changes in the subtropical wind system and by more net-evaporation in the Atlantic Ocean. This increases North Atlantic salinity, ocean circulation and poleward heat transport to the north. Due to a “greener” Tortonian land surface and associated atmospheric and oceanic circulation changes, the Tortonian Atlantic meridional overturning and heat transport have almost their present strengths.

The failure of AGCMs to simulate the reduced pole-to-equator gradients of warm climate intervals is a long-standing problem in paleoclimate modeling (e.g., Sloan et al., 1995). The interaction between land surface cover, atmospheric as well as oceanic

CPD

2, 605–631, 2006

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

circulation could be the so far unknown mechanism for increasing ocean heat transport at a time when meridional surface temperature and vertical temperature gradients were greatly reduced relative to the modern (Bice et al., 2000).

Based on our finding, it is conceivable that reorganizations of the global ocean circulation, large-scale shifts of vegetation zones, topographical changes and changes in the global carbon chemistry play a dominant role for the major Cenozoic climate transitions. Consequently, it is of utmost importance not only to understand the behaviour of these individual systems in better detail but also to investigate the full dynamics, feedbacks, and synergisms of the coupled system. The results presented show a possible strong connection between the hydrological cycle, vegetation cover, and the ocean circulation. Future work will address the numerous interactions between the climate system components by use of a global atmosphere-ocean-vegetation-carbon cycle model. Dutton and Barron (1997) applied a palaeo-vegetation in a modelling study of the Miocene which led to a significant warming suggesting that vegetation and vegetation-climate feedbacks could be a significant component of the Cenozoic climate evolution. Feedback analysis including synergisms shall be performed to consider the dynamics of the climate system in a similar way as for the Quaternary climate variations (e.g., Ganopolski et al., 1998; Kubatzki et al., 2000). A focus can be on the relative roles of the thermohaline circulation, the atmospheric dynamics including high latitude and monsoon circulation, as well as land surface effects caused by changed vegetation distribution.

In order to further investigate major developments during the Miocene a combined approach between modeling and establishing proxy records from selected key locations is needed. Model results on changing patterns of heat transport can be validated by temperature reconstructions (Mg/Ca, alkenones, TEX86), both from the deep (benthic fauna) and the shallow (planktonic) ocean (Lear et al., 2003; Billups et al., 2002; Sluijs et al., 2006). Major changes in ocean circulation can be traced by using water mass characteristic proxies like Cd/Ca, Nd isotopes, and C-13 (Frank et al., 1999; Frank et al., 2002; Delaney and Boyle, 1987). Combination of temperature reconstruc-

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

tions with O-18 gives evidence on changes in salinity and may provide indications on the high salinities in the northern North Atlantic, the position of the ITCZ and associated teleconnections.

Acknowledgements. The work was supported by the Deutsche Forschungsgemeinschaft through the project “Miocene Carbonate Deposition”.

The model ECHAM4 was provided from the Max Planck Institute of Meteorology, Hamburg. Computer simulations were performed at the Deutsches Klimarechenzentrum DKRZ, Hamburg.

References

Arpe, K., Bengtsson, L., Golitsyn, G. S., Mokhov, I. I., Semenov, V. A., and Sporyshev, P. V.: Connection between Caspian Sea level variability and ENSO, *Geophys. Res. Lett.* 27, 2693–2697, 2000.

Barron, E. J.: Eocene equator-to-pole surface ocean temperature: A significant climate problem?, *Paleoceanography*, 2, 729–739, 1987.

Bice, K. L., Scotese, C. R., Seidov, D., and Barron, E. J.: Quantifying the role of geographic change in Cenozoic ocean heat transport using uncoupled atmosphere and ocean models, *Paleogeography, Paleoclimatology, Paleoecology*, 161, 295–310, 2000.

Billups, K.: Late Miocene through early Pliocene deep water circulation and climate change viewed from the sub-Antarctic South Atlantic, *Paleogeography, Paleoclimatology, Paleoecology*, 185, 287–307, 2002.

Billups, K. and Schrag, D. P.: Paleotemperatures and ice volume of the past 27 Myr revisited with paired Mg/Ca and 18O/16O measurements on benthic foraminifera, *Paleoceanography*, 17, 1003–1013, 2002.

Billups, K., Channell, J. E. T., and Zachos: Late Oligocene to early Miocene geochronology and paleoceanography. *Paleoceanography*, 17, 1004–1014, 2002.

Birchfield, G. E.: A coupled ocean-atmosphere climate model: temperature versus salinity effects on the thermohaline circulation, *Climate Dynamics*, 4, 57–71, 1989.

Boulter, M. C. and Manum, S. B.: A lost continent in temperate Arctic, *Endavour*, 21, 105–108, 1997.

CPD

2, 605–631, 2006

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

**Effect of vegetation
on the Late Miocene
ocean circulation**G. Lohmann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Bredenkamp, G. J., Spada, F., and Kazmierczak, E.: On the origin of northern and southern hemisphere grasslands, *Plant Ecol.*, 163, 209–229, 2002.
- Broecker, W. S.: The salinity contrast between the Atlantic and Pacific during glacial time, *Paleoceanogr.*, 4, 207–212, 1992.
- 5 Butzin, M., Prange, M., and Lohmann, G.: Radiocarbon simulations for the glacial ocean: the effects of wind stress, Southern Ocean sea ice and Heinrich events, *Earth Planet. Sci. Lett.*, 235, 45–61. doi:10.1016/j.epsl.2005.03.003, 2005.
- Caballero, R. and Langen, P.: The dynamic range of poleward energy transport in an atmospheric general circulation model, *Geophys. Res. Lett.*, 32, L02705, doi:10.1029/2004GL021581, 2005.
- 10 Cane, M. A. and Molnar, P.: Closing of the Indonesian seaway as a precursor to east African aridification around 3–4 million years ago, *Nature*, 411, 157–161, 2001.
- Cerling, T. E., Harris, J. M., MacFadden, B. J., Leakey, M. G., Quade, J., Eisenmann, V., and Ehleringer, J. R.: Global vegetation change through the Miocene/Pliocene boundary, *Nature*, 389, 153–159, 1997.
- 15 Crowley, T. J. and Zachos, J. C.: Comparison of zonal temperature profiles for past warm periods, in *Warm Climates in Earth History*, edited by: Huber, B., MacLeod, K. G., and Wing, S. C., 50–76, Cambridge Univ. Press, New York, 2000.
- DeConto, R. M. and Pollard, D.: Rapid Cenozoic glaciation of Antarctica induced by declining atmospheric CO₂, *Nature*, 421, 245–249, 2003.
- 20 Delaney, M. and Boyle, E. A.: Cd/Ca in Late Miocene benthic foraminifera and changes in the global organic carbon budget, *Nature*, 330, 156–159, 1987.
- Dutton, J. F. and Barron, E. J.: Miocene to present vegetation changes: A possible piece of the Cenozoic puzzle, *Geology*, 25, 39–41, 1997.
- 25 Emanuel, K.: A simple model for multiple climate regimes, *J. Geophys. Res.*, 107(D9), 4077, doi:10.1029/2001/JD001002, 2002.
- Fedorov, V., Dekens, P. S., McCarthy, M., Ravelo, A. C., deMenocal, P. B., Barreiro, M., Pacanowski, R. C., and Philander, S. G.: The Pliocene paradox (Mechanisms for a permanent El Niño), *Science*, 9, 312, 5779, 1485–1489, 2006, doi:10.1126/science.1122666.
- 30 Flower, B. P. and Kennett, J. P.: The middle Miocene climatic transition: East Antarctic ice sheet development, deep ocean circulation and global carbon cycling, *Palaeogeography, Palaeoclimatology, Palaeoecology*, 108, 537–555, 1994.
- Fortelius, M., Bruch, A., Butzin, M., van Dam, J., Damuth, J., Francois, L., Kvacek, Z., Meu-

**Effect of vegetation
on the Late Miocene
ocean circulation**G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

lenkamp, J., Micheels, A., Popescu, S., Suc, J.-P., and Utescher, T.: Report of the Second EEDEN workshop of proxy and modelling groups, Bratislava, 6 pp., 2003.

François, L., Ghislain, M., Otto, D., and Micheels, A.: Late Miocene vegetation reconstruction with the CARAIB model, *Palaeogeography, Palaeoclimatology, Palaeoecology*, in press, doi:10.1016/j.palaeo.2006.03.034, <http://dx.doi.org/10.1016/j.palaeo.2006.03.034>, 2006.

Frank, M., O’Nions, R. K., Hein, J. R., and Banakar, V. K.: 60 Ma records of major elements and Pb-Nd isotopes from hydrogenous ferromanganese crusts: Reconstruction of seawater paleochemistry, *Geochim. Cosmochim. Acta*, 63, 1689–1708, 1999.

Frank, M., Whiteley, N., Kasten, S., Hein, J. R., and O’Nions, R. K.: North Atlantic Deep Water export to the Southern Ocean over the past 14 Myr: Evidence from Nd and Pb isotopes in ferromanganese crusts, *Paleoceanography*, 17, 1022–1031, 2002.

Freeman, K. H. and Colarusso, L. A.: Molecular and isotopic records of C4 grassland expansion in the late Miocene, *Geochimica et Cosmochimica Acta*, 65, 1439–1454, 2001.

Ganopolski, A., Kubatzki, C., Claussen, M., Brovkin, V., and Petoukhov, V.: The influence of vegetation–atmosphere–ocean interaction on climate during the Mid-Holocene, *Science*, 280, 1916–1919, 1998.

Greenwood, D. R., and Wing, S. L.: Eocene continental climates and latitudinal temperature gradients, *Geology*, 23(11), 1044–1048, 1995.

Haug, G. and Tiedemann, R.: Effect of the formation of the Isthmus of Panama on Atlantic Ocean thermohaline circulation, *Nature*, 393, 673–676, 1998.

Hay, W. W., DeConto, R. M., and Wold, C. N.: Climate: Is the past key to the future?, *Geol. Rundsch.*, 86, 471–491, 1997.

Huber, M. and Sloan, L.: Heat transport, deep waters and thermal gradients: Coupled climate simulation of an Eocene greenhouse climate, *Geophys. Res. Lett.*, 28, 3481–3484, 2001.

Kaplan, J. O., Prentice, I. C., Knorr, W., and Valdes, P. J.: Modelling the dynamics of terrestrial carbon storage since the Last Glacial Maximum, *Geophys. Res. Lett.*, 29, 2074–2078, 2002.

Kennett, J. P.: Cenozoic evolution of Antarctic glaciation, the Circum Antarctic Ocean, and their impact on global paleoceanography, *J. Geophys. Res.*, 82, 3843–3860, 1977.

Kennett, J. P. and Barker, P. F.: Latest Cretaceous to Cenozoic climate and oceanographic developments in the Weddell Sea, Antarctica: An ocean-drilling perspective, in: *Proceedings of the Ocean Drilling Program*, edited by: Barker, P. F. and Kennett, J. P., 113, *Scientific Reports, Weddell Sea, Antarctica*, 937–960, 1990.

Knorr, G. and Lohmann, G.: Southern Ocean origin for resumption of Atlantic thermohaline

**Effect of vegetation
on the Late Miocene
ocean circulation**G. Lohmann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

circulation during deglaciation, *Nature*, 424, 532–536, 2003.

Kubatzki, C., Montoya, M., Rahmstorf, S., Ganopolski, A., and Claussen, M.: Comparison of the last interglacial climate simulated by a coupled global model of intermediate complexity and AOGCM, *Climate Dynamics*, 16, 799–814, 2000.

5 Latif, M., Roeckner, E., Mikolajewicz, U., and Voss, R.: Tropical stabilization of the thermohaline circulation in a greenhouse warming simulation, *J. Climate*, 13, 1809–1813, 2000.

Lawver, L. A. and Gahagan, L. M.: Evolution of Cenozoic seaways in the circum-Antarctic region. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 198, 11–37, 2003.

Le Houerou, H. N.: Climate, flora and fauna changes in the Sahara over the past 500 million years, *J. Arid Environ.*, 37, 619–647, 1997.

Lear, C. H., Rosenthal, Y., and Wright, J. D.: The closing of a seaway: Ocean water masses and global climate change, *Earth Planet. Sci. Lett.*, 210, 425–436, 2003.

Lohmann, G.: The influence of a near-bottom transport parameterization on the sensitivity of the thermohaline circulation, *J. Phys. Oceanogr.*, 28, 2095–2103, 1998.

15 Lohmann, G.: Atmospheric and oceanic freshwater transport during weak Atlantic overturning circulation, *Tellus 55 A*, 438–449, 2003.

Lohmann, G. and Gerdes, R.: Sea ice effects on the sensitivity of the thermohaline circulation in simplified atmosphere-ocean-sea ice models, *J. Climate*, 11, 2789–2803, 1998.

Lohmann, G. and Lorenz, S.: On the hydrological cycle under paleoclimatic conditions as derived from AGCM simulations, *J. Geophys. Res.*, 105(D13), 17, 417–436, 2000.

20 Lohmann, G. and Schulz, M.: Reconciling Bølling warmth with peak deglacial meltwater discharge, *Paleoceanography*, 15(5), 537–540, 2000.

Macdonald, A. M. and Wunsch, C.: An estimate of global ocean circulation and heat fluxes, *Nature*, 382, 436–439, 1996.

25 Mai, D. H.: *Tertiäre Vegetationsgeschichte Europas*. Gustav Fischer, Jena, 1995.

Maier-Reimer, E., Mikolajewicz, U., and Hasselmann, K.: Mean circulation of the Hamburg LSG OGCM and its sensitivity to the thermohaline surface forcing, *J. Phys. Oceanogr.*, 23, 731–757, 1993.

Micheels, A.: Late Miocene climate modelling with ECHAM4/ML – the effects of the palaeovegetation on the Tortonian climate, PhD Thesis University of Tübingen, 110 pp., 2003.

30 Micheels, A., Bruch, A. A., Uhl, D., Utescher, T., and Mosbrugger, V.: A Late Miocene climate model simulation with ECHAM4/ML and its quantitative validation with terrestrial proxy data, *Palaeogeography, Palaeoclimatology, Palaeoecology*, in press, 2006.

- Mikolajewicz, U., Maier-Reimer, E., Crowley, T. J., and Kim, K. J.: Effect of Drake and Panamanian gateways on the circulation of an ocean model, *Palaeoceanography*, 8, 409–426, 1993.
- 5 Miller, K. G., Fairbanks, R. G., and Mountain, G.: Cenozoic oxygen isotope synthesis, sea-level history, and continental margin erosion, *Paleoceanography*, 2, 1–20, 1987.
- Morley, R. J.: *Origin and evolution of tropical rain forests*, Wiley, Chichester, 362 p, 2000.
- Otto-Bliesner, B. and Upchurch, G. R.: Vegetation-induced warming of high-latitude regions during the Late Cretaceous period, *Nature*, 385, 804–807, 1997.
- 10 Pagani, M., Arthur, M. A., and Freeman, K. H.: Miocene evolution of atmospheric carbon dioxide, *Paleoceanography*, 14, 273–292, 1999.
- Pagani, M., Zachos, J. C., Freeman, K. H., Tipple, B., and Bohaty, S.: Marked decline in atmospheric carbon dioxide concentrations during the Paleogene, *Science*, 309, 600–603, 2005.
- Pearson, P. N. and Palmer, M. R.: Atmospheric carbon dioxide concentrations over the past 15 600 million years, *Nature*, 406, 695–699, 2000.
- Peixoto, J. P. and Oort, A. H.: *Physics of climate*, American Institute of Physics, 520 pp., 1992.
- Prange, M., Romanova, V., and Lohmann, G.: The glacial thermohaline circulation: stable or unstable?, *Geophys. Res. Lett.*, 29(21), 2028, doi:10.1029/2002GL015337, 2002.
- Prange, M., Lohmann, G., and Paul, A.: Influence of vertical mixing on the thermohaline hysteresis: Analyses of an OGCM, *J. Phys. Oceanogr.*, 33(8), 1707–1721, 2003.
- 20 Rahmstorf, S.: On the freshwater forcing and transport of the Atlantic thermohaline circulation, *Climate Dynamics*, 12, 799–811, 1996.
- Retallack, G.: Cenozoic expansion of grasslands and climatic cooling. *J. Geol.*, 109, 407–426, 2001.
- 25 Roeckner E., Arpe, K., Bengtsson, L., Christoph, M., Claussen, M., Dümenil, L., Esch, M., Giorgetta, M., Schlese, U., and Schulzweida, U.: The atmospheric general circulation model ECHAM-4: Model description and simulation of present-day climate, Max-Planck-Institut für Meteorologie, MPI Report No. 218, 1996.
- Schäfer-Neth, C. and Paul, A.: Circulation of the glacial Atlantic: a synthesis of global and regional modeling, edited by: Schäfer, P., Ritzrau, W., Schlüter, M., and Thiede, J., in: *The northern North Atlantic: A changing environment*, Springer, Berlin, 446–462, 2001.
- 30 Schmidt, G. A. and Mysak, L. A.: Can increased poleward oceanic heat flux explain the warm Cretaceous climate?, *Paleoceanography*, 11, 579–593, 1996.

Effect of vegetation on the Late Miocene ocean circulationG. Lohmann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Schmitz, W. J.: On the interbasin scale thermohaline circulation, *Rev. Geophys.*, 33, 151–173, 1995.
- Schmittner A., Appenzeller, C., and Stocker, T. F.: Enhanced Atlantic freshwater export during El Niño, *Geophys. Res. Lett.*, 27, 1163–1166, 2000.
- 5 Scholze, M., Knorr, W., and Heimann, M.: Modelling terrestrial vegetation dynamics and carbon cycling for an abrupt climate change event, *The Holocene*, 13(3), 327–333, 2003.
- Schuster, M., Düringer, P., Ghienne, J.-F., Vignaud, P., Mackaye, H. T., Likius, A., and Brunet, M.: The Age of the Sahara, *Science*, 311, 821, 2006.
- 10 Sitch, S., Smith, B., Prentice, I. C., Arneth, A., Bondeau, A., Cramer, W., Kaplan, J. O., Levis, S., Lucht, W., Sykes, M. T., Thonicke, K., and Venevsky, S.: Evaluation of ecosystem dynamics, plant geography and terrestrial carbon cycling in the LPJ dynamic global vegetation model, *Glob. Change Biol.*, 9, 161–185, 2003.
- Sloan, L. C. and Pollard, D.: Polar stratospheric clouds: A high latitude warming mechanism in an ancient greenhouse world, *Geophys. Res. Lett.*, 25, 3517–3520, 1998.
- 15 Sloan, L. C., Walker, J. C. G., and Moore, T. C.: Possible role of ocean heat transport in early Eocene climate, *Paleoceanography*, 10, 347–356, 1995.
- Sluijs, A., Schouten, S., Pagani, M., Wolterring, M., Brinkhuis, H., Sinninghe Damste, J. S., Dickens, G. R., Huber, M., Reichert, G.-J., Stein, R., Matthiessen, J., Lourens, L. J., Penderguch, N., Backman, J., Moran, K., and the Expedition 302 Scientists, 2006: Subtropical Arctic Ocean temperatures during the Palaeocene/Eocene thermal maximum, *Science*, 441, 610–613, 2006.
- 20 Soden, B. J.: The sensitivity of the tropical hydrological cycle to ENSO, *J. Climate*, 13, 538–549, 2000.
- Steppuhn, A., Micheels, A., Geiger, G., and Mosbrugger, V.: Reconstructing the Late Miocene Climate and oceanic heat flux using the AGCM ECHAM4 coupled to a mixed-layer ocean model with adjusted flux correction, *Palaeogeography, Palaeoclimatology, Palaeoecology*, doi:10.1016/j.palaeo.2006.03.037, <http://dx.doi.org/10.1016/j.palaeo.2006.03.037>, in press, 2006.
- 25 Willis, K. J. and McElwain, J. C.: *The evolution of plants*. Oxford University Press, Oxford, 378 pp, 2002.
- Wolf, T. C. W. and Thiede, J.: History of terrigenous sedimentation during the past 10 m.y. in the North Atlantic (ODP Legs 104 and 105 and DSDP Leg 81), *Mar. Geol.*, 101, 83–102, 1991.

Effect of vegetation on the Late Miocene ocean circulationG. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

- Wolfe, J. A.: An analysis of Neogene climates in Beringia, *Palaeo. Palaeo. Palaeo.*, 108, 207–216, 1994.
- Wright, J. D. and Miller, K. G.: Control of North Atlantic Deep Water circulation by the Greenland-Scotland Ridge, *Paleoceanography*, 11, 157–169, 1996.
- 5 Wright, J. D., Miller, K. G., and Fairbanks, R. G.: Miocene stable isotopes: Implications for deepwater circulation and climate, *Paleoceanography*, 7, 357–389, 1992.
- Zachos, J., Pagani, M., Sloan, L., Thomas, E., and Billups, K.: Trends, rhythms, and aberrations in global climate 65 Ma to Present, *Science*, 29, 686–693, 2001.
- 10 Zaucker, F. and Broecker, W. S.: The influence of atmospheric moisture transport on fresh water balance of the Atlantic drainage basin: General circulation model simulations and observations, *J. Geophys. Res.*, 97, 2765–2773, 1992.
- Zaucker, F., Stocker, T. F., and Broecker, W. S.: Atmospheric freshwater fluxes and their effect on the global thermohaline circulation, *J. Geophys. Res.*, 99, 12 443–12 457, 1994.

CPD

2, 605–631, 2006

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

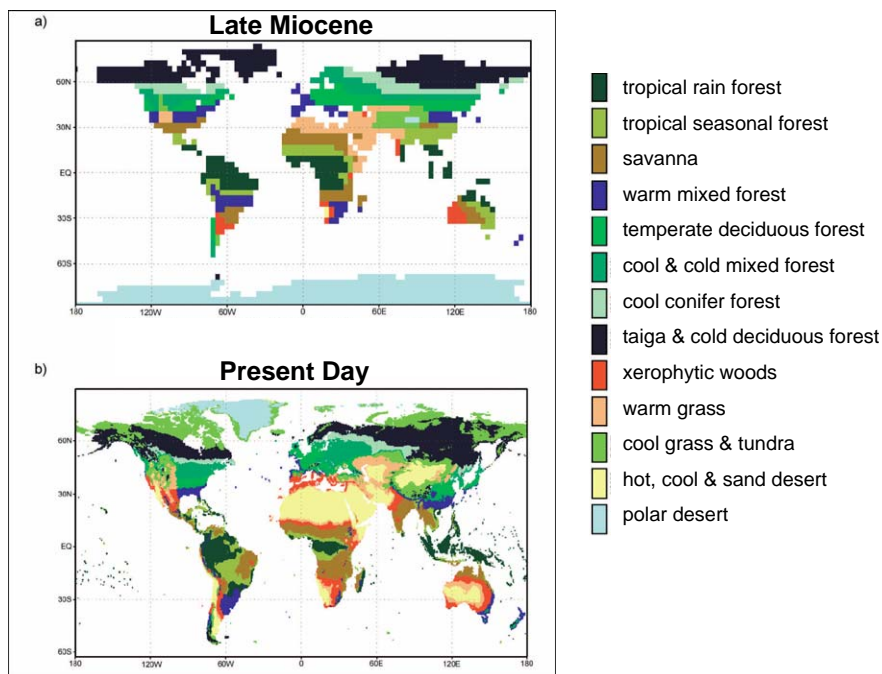


Fig. 1. (a) The proxy-based reconstructed Tortonian vegetation, and (b) the present-day's vegetation (New et al., 1999). These maps serve as an input into the AGCM experiments TVEG, TGEO and CTRL, respectively.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Effect of vegetation
on the Late Miocene
ocean circulation**

G. Lohmann et al.

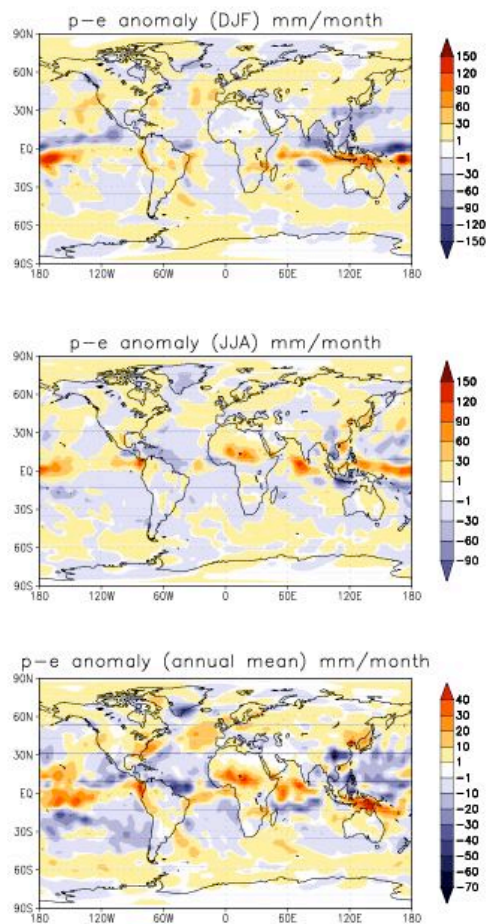


Fig. 2. Changes in the net precipitation minus evaporation for TVEG minus CTRL. Units are mm/month.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

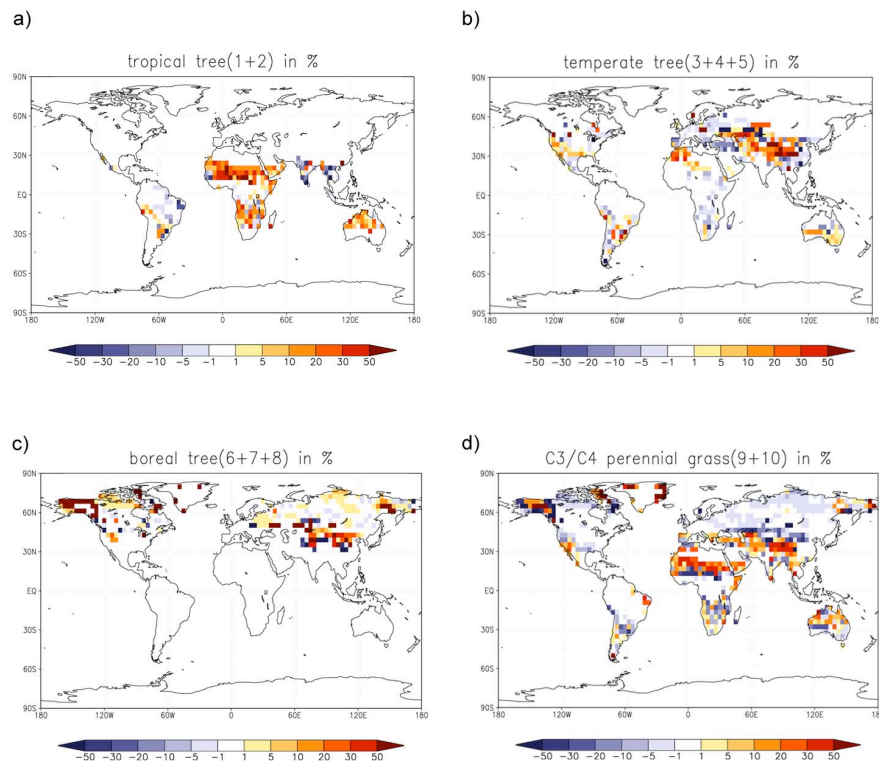


Fig. 3. Change of the late Miocene relative to the present vegetation (TVEG minus CTRL). Fractional coverage (in percent): **(a)** Tropical tree, **(b)** temperate tree, **(c)** boreal tree, **(d)** grass.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

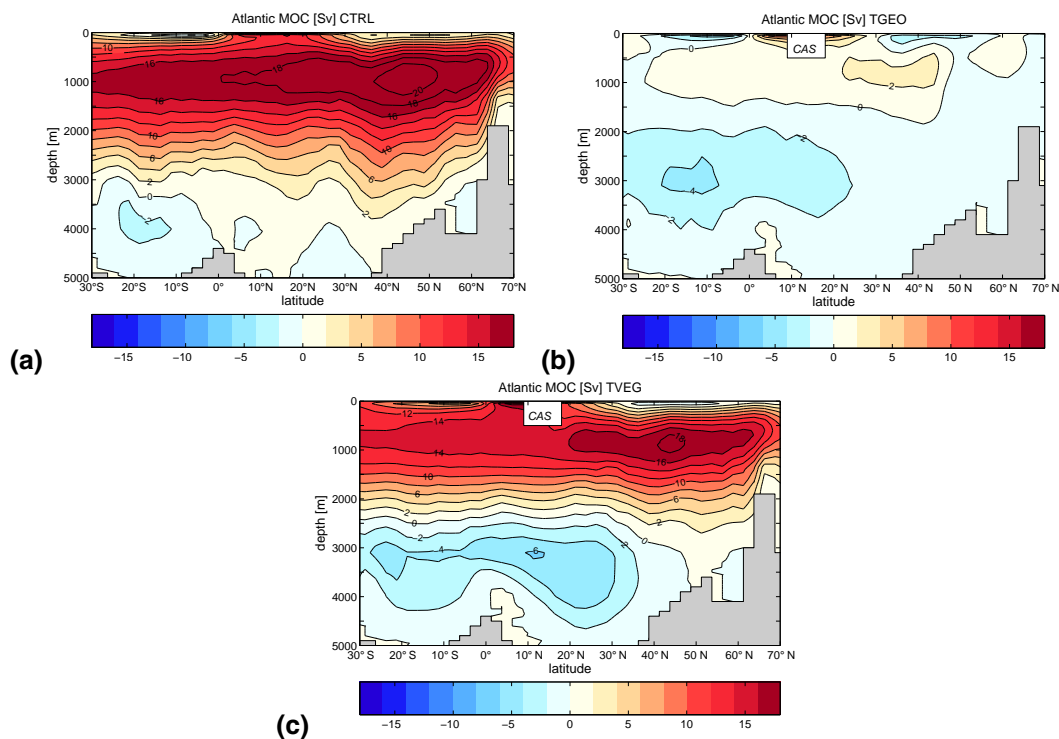


Fig. 4. Atlantic meridional overturning circulation ($\text{Sv} = 10^6 \text{ m}^3/\text{s}$) for present-day (a), and the late Miocene configuration with open Central American Seaway (CAS). (b) with present vegetation cover (T GEO), and (c) with reconstructed vegetation cover (T VEG).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

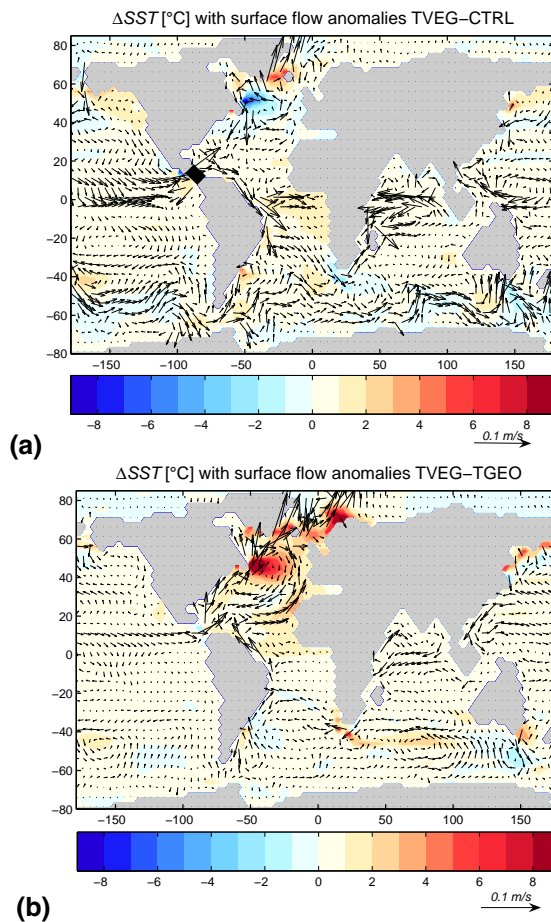


Fig. 5. Modelled sea surface temperature anomalies [°C] and surface flow [m/s]. **(a)** difference between TVEG and CTRL, **(b)** difference between TVEG and TGEO.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

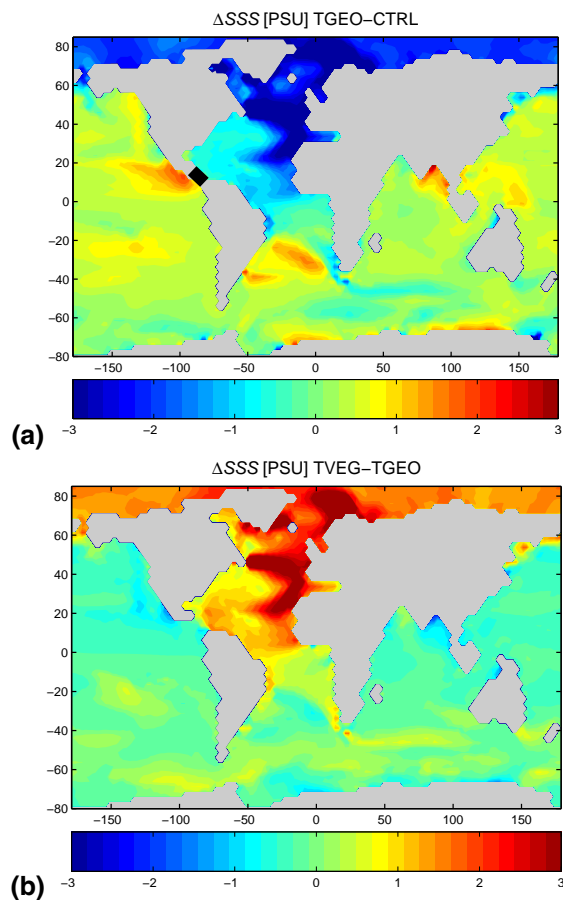


Fig. 6. Modelled sea surface salinity anomalies [PSU]. **(a)** difference between TGE0 and CTRL, **(b)** difference between TVEG and TGE0.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effect of vegetation on the Late Miocene ocean circulation

G. Lohmann et al.

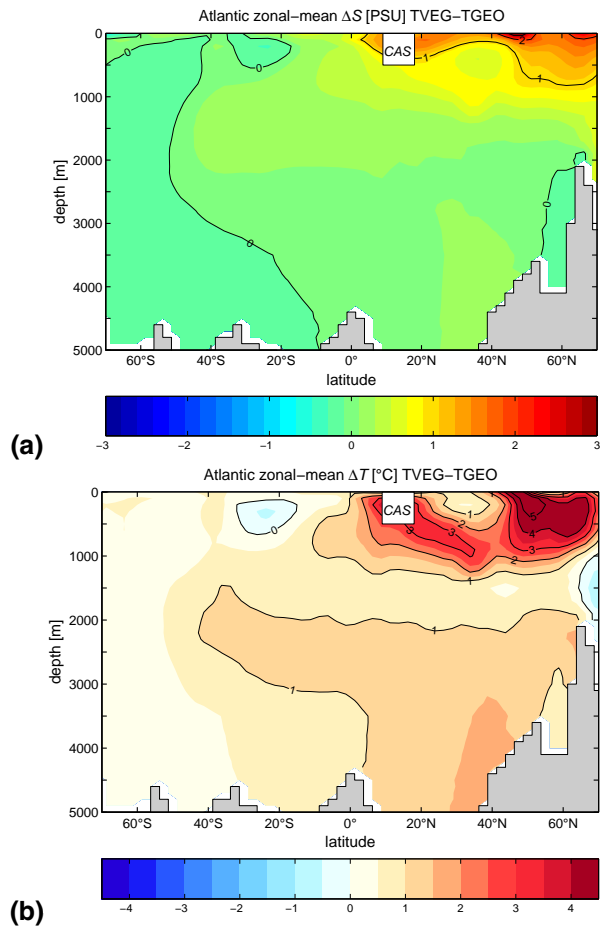


Fig. 7. Zonal mean difference between TVEG and TGeo in the Atlantic Ocean: **(a)** salinity, **(b)** temperature.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)