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**Biological control of
the terrestrial carbon
sink**

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Biological control of the terrestrial carbon sink¹

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

¹This paper is based on the Vernadsky Medal Lecture 2004

Abstract

This is a summary of the Vernadsky medal lecture given at the Nice EGU meeting in 2004. The lecture reviews the past (since the International Biological Program) and the future of our understanding of terrestrial carbon fluxes with focus on photosynthesis, respiration, primary, ecosystem, and biome productivity. Consideration is given to the interactions between biodiversity and biogeochemical processes.

1. Introduction

The interest in understanding global land surface processes has increased since the 1996 IPCC report (IPCC, 1996), stating that “the balance of evidence suggests a discernible human influence on global climate”. The human influence, however, is of rather diverse nature (Fig. 1). The human population affects natural ecosystems and climate via its size and its resource use through activities which range from agriculture to trade. All these activities lead to land transformations, such as forest clearing, to changes in biogeochemistry, such as the global carbon cycle, and to biotic changes, such as invasion and extinction of species. These changes affect global climate as well as biological diversity, and both feed back on the global biogeochemical cycles on the one hand, and on the human population on the other. This latter feedback, however, strongly depends on the socio-economic conditions, which, in contrast to the physical climate system, are difficult to predict for the future.

The overall effect of human activity on global biogeochemistry is most obvious from the rise in carbon dioxide (CO₂) as measured since the 1960s on Mauna Loa in Hawaii, and in a global network of stations since the 1970s. The difference between the annual fossil fuel consumption and the annual change of the rise in atmospheric CO₂ is caused by the activity of the surfaces of marine and terrestrial environments to assimilate CO₂, the so-called carbon sink (Fig. 2A). Despite large year-to-year variation, the sink activity has increased since 1960 (Prentice et al., 2001), and the variation of the global carbon

BGD

2, 1283–1329, 2005

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

sink is to a major part related to the occurrence of El Niño events. This has been interpreted as evidence for a strong biological regulation of the global carbon sink (Prentice et al., 2001). At the same time, Heimann (pers. communication) points out that the fraction of fossil fuel CO₂ that remains in the atmosphere has been constant since 1960 (Fig. 2B), which suggests that there might also be a strong physical control of this fraction.

Since atmospheric oxygen measurements have been used to partition the total global carbon sink into activities of the ocean and the terrestrial surface (Keeling et al., 1996), it has become clear, that the variation of the increase in atmospheric CO₂ is strongly related to processes at the land surface. Therefore, in the following the processes will be discussed that cause changes of the natural sources and sinks on land, and the effects of natural and anthropogenic disturbances on terrestrial ecosystems which are the basis for such sink activity. At the end of this paper I will add some thoughts about the future of fossil fuel emissions. This overview will not be complete, but a selection of mainly personal observations from the past decades. It is a summary of the Vernadsky Medal Lecture presented at the 2004 EGU meeting.

2. Natural sinks: CO₂ assimilation

Natural sinks in terrestrial systems result from photosynthesis by plants, and this process has attracted eco-physiologists since the International Biological Program (IBP), which brought together individual scientists from around the globe to study the performance of plants in different climatic regions (Schulze and Caldwell, 2000). In fact, in a historical perspective, it was the IBP which started field ecology and methodology in a coherent manner (Trebson, 1970).

We learnt from studies in IBP (Fig. 3) that plants operate at about 50% of their photosynthetic capacity (A_{\max}), and since CO₂ assimilation is directly related to stomatal conductance (Schulze et al., 1994), they also operate at about 50% of their maximum leaf conductance (g_{\max}). Different climatic factors cause this down-regulation: it is

mainly insufficient light in temperate regions (clouds, dawn, dusk), and it is the combination of light, temperature and air humidity in semi-arid regions, and it is mainly air humidity in a boreal *Larix* forest of Siberia.

5 It was a surprise to find that CO₂-assimilation was not related to the activity of the CO₂ fixing enzyme, RuBisCo. It is possible to down-regulate RuBisCo by 50% before it affects photosynthesis (Stitt and Schulze, 1993). Obviously RuBisCo has more physiological functions than only fixing CO₂. It is the enzyme that reduces CO₂ to carbohydrates, it protects the photosynthetic apparatus against damaging light through its demand for electrons in CO₂ reduction, and it is a storage product for nitrogen. Despite
10 these diverse functions, A_{max} (per gram dry weight) is linearly related to leaf nitrogen concentration (gN g_{dw}⁻¹) in a broad range of leaf types (Field and Mooney, 1986) (Fig. 4).

The strong effect of air humidity on CO₂ uptake which is shown in Fig. 3C results mainly from an effect of stomatal closure, which may even cause a secondary down-regulation of photosynthesis. Stomata operate in this case independent of CO₂ assimilation. Individual stomata can even be opened and closed when exposing them to moist or dry air (Fig. 5, Lange et al., 1971).
15

The effect of plant water stress on leaf conductance and CO₂ assimilation remained for a very long time an area of uncertainty. Leaf water potential was initially thought to control leaf performance, but the plant world is more complicated. As a general principle, plants have established many internal feedback mechanisms, but with respect
20 to responses towards changing environment, feedback responses may be too slow and thus too late. Damage may have already occurred before the feedback response reduces the stress. Thus, plants have developed also a multitude of feed-forward responses, such as the opening response of stomata at blue light, or their closure in dry air. The response to water stress is also such a feed-forward response. It is the
25 root tip than senses drought in the soil, and it sends a chemical signal (abscisic acid) to the shoot to close stomata. Leaf water potential can be shifted more than 1 MPa, by exposing the whole plant to dry or moist air, but stomata close independent of the water potential in the leaf by a response to dry soil (Fig. 6, Schulze, 1986; Schulze et

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

al., 1988).

The physiology of plant water relations has experienced a complete re-interpretation in recent years. Since Slatyer's book on plant water relations (Slatyer, 1967) it was a dogma that water moves along a water potential gradient in a soil-plant-atmosphere system. However, since then it became apparent that water does not move across membranes along a physico-chemical gradient, but that water movement is physiologically regulated by water channels (Tyermann et al., 1999), a unifying concept of plant water relations is being re-generated at molecular level (Schulze et al., 2005).

Moving from the leaf to the whole stand level, conductance and assimilation is additionally controlled by leaf area (Fig. 7). Maximum canopy conductance (G_{\max}) increases with leaf area index ($m_{\text{leaf area}}^2 m_{\text{ground area}}^{-2}$) in conifers at low G_{\max} and low leaf conductance, g_{\max} as well as in crop plants at high G_{\max} with high leaf conductance, g_{\max} . Surface conductance which is based on the sum of plant transpiration and on evaporation from the soil becomes important at low leaf area indices, where soil-surface conductance may be dominant and increase again the surface conductance. It is interesting to note, that a minor plant cover can reduce surface conductance significantly.

Based on this knowledge, we are able to compress the diversity of plant life forms and plant responses into very few equations, describing the sink performance (Fig. 8). Assimilation is linearly related to leaf nitrogen (nutrition in a general sense). Based on the interdependence of assimilation and conductance (Farquar et al., 1980), also stomatal conductance is related to leaf N. Thus, surface conductance (G , canopy plus soil conductance) is also related to stomatal conductance, i.e. plants can affect water loss beyond the demand of the physical climate. As a result A_{\max} is related to G_{\max} . Based on this knowledge we can draw a global map of A_{\max} based on leaf nitrogen. In a first approximation, actual CO_2 assimilation under given climatic conditions would be about 50% of A_{\max} .

We may conclude with respect to the control of CO_2 assimilation, that natural environments are always stressful. A_{\max} is basically never achieved. If light limitation is

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

released, then dry air or dry soil are limiting. If soils are wet, light is most likely limiting due to cloud cover. The operational point (50% of maximum performance) gives the greatest possible control in a variable environment. In fact, many enzyme systems operate at about 50% of their maximum performance in order to allow for a maximum range of regulation under variable conditions (Stitt, 1993).

In a historical perspective it is interesting to note, that photosynthesis research collapsed in the mid 1980s, because all leaf types of this globe had been enclosed into a measuring cuvette or porometer at least once, and plant scientists became increasingly aware, that CO_2 uptake is not the carbon balance of plants or canopies. Following this depression in research focus, the interest in A_{max} is rising again because mechanistic models of the ecosystem carbon flow require knowledge of A_{max} in a genuine way (Farquhar et al., 1980).

3. Respiration

Respiration entered into the photosynthesis centered research of ecophysicologists following the breakthrough of technology when the eddy covariance methodology was developed by micrometeorologists (Swinbank, 1951). Suddenly it became possible to quantify the net flux of a whole canopy. The exchange of CO_2 , water vapor and heat is measured by analyzing the content of air packages entering and leaving the ecosystem.

The surprising observation of the Euroflux network was (Valentini et al., 2000), that the annual ecosystem balance of assimilation and respiration, decreased with latitude (Fig. 9, uptake is denoted as negative number, because it is a depletion of the atmosphere by CO_2). Since growth studies did not support this result, the target of research became respiration. The full explanation and interpretation of this result is still under debate (Van Dijk et al., 2005), because of the difficulties to quantify respiration.

Ecosystem respiration contains several components, and only few examples exist, in which the partitioning of respiration has been untangled. Ecosystem respiration

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

consists of respiration by leaves or needles, respiration by woody stems, which may contain the largest amount of biomass in forests but the main part of which is dead, and of the soil. It still remains unclear, if leaves respire during photosynthesis or if respiratory CO₂ is immediately re-assimilated in leaves (Loreto et al., 1999), just as much as soil respiration is being re-assimilated in the canopy or the understory during the day. Figure 10A shows an example for a spruce forest in Germany, where needles contribute about 30%, woody stems contribute about 20%, and the soil almost 50% to ecosystem respiration. There is also a distinct seasonal and inter-annual variation.

Following this analysis the question still remains open, to what extent roots and microbes contribute to soil respiration. Since processes in soils are notoriously difficult to study, the partitioning of respiration in soils remained unclear until P. Högberg (Högberg et al., 2001) designed an experiment in which he was able to shut off root respiration, but maintain microbial respiration without disturbance, i.e. without changing soil structure and water status. Högberg interrupted the carbohydrate supply to the root, but maintained the water transport from the root by girdling the trees (taking off the bark). Buchmann (published in Schulze, 2003) repeated this experiment for spruce in Germany (Fig. 10B). It was estimated, that 40% of soil respiration originates from plant roots. Combining this information with Fig. 9, then plants contribute about 75% to total ecosystem respiration. Microbial respiration reaches a maximum in spring from mineralization of leaf litter, and respiration from the mineral soil. In contrast to the distinct seasonal course in root respiration, the seasonality in soil respiration is smaller. At this moment it is unclear if this is due to the buffered climate conditions in forest soils or to acclimation of soil organisms. The latter would mean, that using a constant Q₁₀ for predicting future behavior of soils under various scenarios of temperature rise may just be questionable.

Although the experimental approach to separate plant and microbial respiration by girdling the trees appears very elegant, it has also its problems. Schulze et al. (2005) identified the exiting organismic community using proteomics fingerprinting (Fig. 11). In the healthy control stand there was a wealth of organismic groups present, in total

9. However, in the girdled stand only 4 organismic groups remained. This shows first of all, how many compartments in an ecosystem directly depend on carbohydrates from the tree. It is more than just the root and its mycorrhiza. Second, the analysis shows, that the change in respiration was not representing just the activity of plant roots, but respiration of a whole group of associated organisms. Third, this analysis shows, that it is impossible to carry out a manipulative experiment without changing the soil biota, and that it appears impossible to maintain the soil biota constant. Finally, any manipulative experiment should follow closely the activity of the organismic community as a whole.

10 Investigating roots and soils has become a priority research (Science, 2004), but many aspects remain unclear. Where ends a root? Figure 12A shows the root system (orange part) and the associated mycorrhizal system (white hyphae) of a spruce seedling. Obviously, the plant's range of action is significantly enlarged by the mycorrhizal fungus, and it remains unclear, to what extent this fungus contributes to soil respiration, if it is strictly autotrophic or heterotrophic, and how much C is gained or lost by the plant through mycorrhizae. According to Read (Schulze, 2003) soil respiration decreases about 30% if roots and mycorrhizae are excluded. Julou et al. (2005) showed that there are plants that make use of the carbohydrate flow from the root to the mycorrhizae and tap this flow as parasites. Orchids seem to be especially capable of doing this as a basis for survival in the shade.

20 Roots do not only extend into the microbial community by fungal associations, but extend the active zone of biological activity also in terms of soil depth. Most ecosystem studies end at about 0.6 m soil depth, while roots explore soils in all ecosystems to a depth of more than 1 m, and may reach maximum measured depth at 50 to 70 m (Canadell et al., 1996). They may extend to about 100 m in arid climates. This part of the ecosystem remained almost unstudied so far (Fig. 12B), but the activity would be captured in the eddy covariance measurement.

25 It is only recently that one has become aware that the net canopy flux (NCF) as measured by eddy covariance is not the net flux of the ecosystem (NEE). There are

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Print Version](#)[Interactive Discussion](#)

lateral transports by thermal gradients and associated advection (Aubinet et al., 2003), and there are losses to the groundwater by dissolved organic carbon (DOC). The fate of both fluxes is important at a landscape scale. The laterally transported C will emerge elsewhere in the landscape. The DOC may be mineralized in the anaerobic zone of the profile and emerge as methane, or contribute to the formation of dissolved inorganic C (DIC). Thus, the C balance in soils remains unsettled.

4. The ecosystem approach

The change in research focus, away from CO₂ assimilation research towards the study of ecosystems processes was historically initiated by the sudden widespread occurrence of forest decline in Europe (Last and Watling, 1991). Initially the yellowing of needles was regarded as a direct effect of air pollutants, especially of SO₂ on CO₂ assimilation. However, it was shown that photosynthesis was independent of the S concentration in the leaf (Fig. 13, Lange et al., 1989). The mass balance of sulfur indicated (Köstner et al., 1998) that the main fraction of sulfate in leaves originates from uptake by roots from the soil, and not via uptake of SO₂ from the atmosphere. This finding underlined the necessity to link photosynthesis research with whole plant ecology and ecosystem science.

Although sulfur proved not to be damaging to needles, but it does affect the soils. It became obvious that other basic nutrients have an essential role in forest decline: cations and nitrogen. Leaves and shoots are not only a surface for exchange of water vapor and CO₂, but participate in the nutrient and cation exchange of plants just as much as roots. The processes that are involved in nitrogen exchange became especially complicated. Figure 14 shows the complexity of pathways that are involved in the canopy uptake of nitrogen. Nitrogen enters in gaseous form as NO_x or NH₃. Ammonia may also co-deposit with SO₂ on leaf surfaces, mainly catalyzed by Mn, and then enter in liquid form as ammonium cation or nitrate anion into the plant. Even sulfate may enter via this route. The ionic uptake is not regulated as in roots, but it is associated with

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

an exchange of protons or cations from the inside of the leaf, well known as canopy leaching (Ulrich, 1987). To my knowledge, the different pathways have not yet been measured independently and directly. There are only estimates that the wet deposition to ecosystems as measured in rain (on average 10 to 20 kg ha⁻¹ yr⁻¹) may increase by factor 2 to 3 if aboveground uptake of N into biomass is being considered (Harrison et al., 2000).

With respect to forest decline it emerged that the visible symptoms of yellow needles results from an interaction of N-uptake by unregulated uptake from the air under conditions of a limitation in cation availability in soils which results from a sulfate induced soil acidification. The experiment that demonstrated this interaction was surprisingly simple (Fig. 15A). Since N uptake affects primarily growth, the hypothesis was that if one would stop growth then cation uptake could balance the demand, and needles would not become yellow. Therefore, buds were clipped off or left for growth of new shoots on opposite twigs of the same branch (Lange et al., 1987). At high cation supply, clipping of buds had no visible effect, but under conditions of cation limitation needles became yellow. Clipping buds and avoiding growth restored the cation balance and lead to a re-greening of the shoot. In general terms, inhibiting growth affected the following chain of processes (Fig. 15B): If we consider two levels of cation supply (healthy green and yellow plants), and if both plants grow in relation to the amount of available nitrogen, then cation re-allocation takes place in the plant from old to young needles, in order to support the new growth. This leads to old needles reaching a threshold at which they start to turn yellow or die. This phenomenon is most clearly expressed with respect to Mg, but in principle the same processes are involved with Fe, K, Mn and other cation deficiencies.

In summary, the decade of acid rain had a very significant effect on ecosystem science. In fact it was the birth of an interdisciplinary approach to biogeochemistry in contrast to the “trophic level” approach of earlier ecosystem research. In addition, it was the acid rain research, which induced the awareness that research must go beyond ecosystems and even beyond national boundaries. The cross boundary transport of air

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

pollutants (FBW, 1989), and the concept of critical loads (Schulze et al., 1989) lead to European and global research networks, and the start of the International Geosphere-Biosphere Program (IGBP) with GCTE (Global Change and Terrestrial Ecosystems) as an important component (Walker and Steffen, 1996). In the last 2 decades IGBP-GCTE had been the driving force to create global networks (e.g. the CO₂- and flux-networks) and to concentrate international research along continental transects. These networks are the foundation for our research even today, e.g. the European Integrated project CarboEurope-IP emerged from IGBP-type research organization.

5. The role of disturbance

Ecosystems loose carbon not only in a regular response to carbohydrate supply and climate but also by events that are rather irregular in time, called disturbances. With respect to the carbon balance, disturbance is not only the event of destruction, but also the regular change as imposed by global temperature change that continually stresses an ecosystem. One may classify disturbances (Fig. 16) as continuous forcing of the C and N cycle, e.g. by temperature or CO₂ increase, which could result in acclimation as ecosystem response, or disturbance may act as disruption of the C and N cycle, e.g. by wind throw, where all plant material may remain on site if it is not extracted by forest management, or disturbances contain events, where C and N are exported from the ecosystem, such as by fire or harvest, where the C-loss bypasses the heterotrophic chain of respiration, and the oxidation of biomass into CO₂ may occur in a totally different place than the original ecosystem, e.g. in a city incinerator. In the latter case the ecosystem has to re-generate resources such as cations and nitrogen. Fire and harvest emerge as the dominating forms of disturbances, considering the fact that 30% of the terrestrial surface of the globe is used by agriculture, >30% are about used by grazing, and 20 to 30% are used by forestry.

The awareness that ecosystems are not closed systems where input by assimilation is balanced by respiration (Odum, 1956), but that major carbon accumulations and

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

losses take place which bypass respiration, requires a new terminology (Fig. 16, Schulze and Heimann, 1998). Gross primary production, GPP, denotes the CO₂ assimilation as discussed before. Net primary production, NPP, would be the balance between GPP and plant respiration. It is the growth rate of plants, and the target for agriculture and forestry. Net ecosystem productivity, NEP, includes heterotrophic soil respiration, and this quantity may be measured by eddy covariance methodology. In most ecosystems studies so far NEP were indicating a C sink. However, eddy covariance does not cover the export of carbon that bypasses respiration, mainly by harvest and by fire. Harvest for food and fiber is the main target of human land management. The combined process, including carbon losses off site and respiration has been termed Net biome productivity, NBP.

(1) $GPP = CO_2$ Assimilation (assuming that respiration in the light equals dark respiration)

(2) $NPP = GPP - R_a$ (R_a being respiration by the autotrophic plant)

(3) $NEP = NPP - R_h$ (R_h being the respiration of the heterotrophes)

(4) $NBP = NEP -$ (non-respiratory fluxes)

Non-respiratory fluxes are a whole suit of processes that contribute or distract C from an ecosystem, mainly fire and harvest.

All these quantities are fluxes (g or mol per area and time), which can be integrated over time and space as desired. Since non-respiratory losses generally occur patchy across the landscape, the “biome” served as entity, but NBP can also be measured at the ecosystem scale meanwhile Chapin et al. (2005) defined NECB, Net Ecosystem Carbon Balance, as NBP at the ecosystem level.

NBP would contribute to the fraction of CO₂ assimilation that accumulates in the

ecosystem over very long time, the ecosystem carbon pool. It would be the actual long-term C-sequestration by ecosystems which the atmosphere “sees”. It turns out, that the measurement of NBP is difficult, because we are dealing with slow uptake process accumulating C over very long time, and very fast, short-term C-losses. Körner (2003) characterized this behaviour as “slow in, fast out”-system, where a balance is difficult to assess. Probably it can only be quantified over large areas, such as a biome. The long-term effects of the non-respiratory, “fast out” C-losses may be dramatic for the C-pool. A model example is shown in Fig. 18, where boreal ecosystems are disturbed by fire in different frequency, and the total C accumulation after 6000 years differs by factor 10 between frequent and low fire events. The total C-accumulation is in this case neither related to respiration, nor to CO₂-assimilation, but related to the drain of C by fire.

Disturbances by land use change are particularly long-lasting. The carbon loss by change of forest into agriculture is detectable in forest soils, even after several hundreds of years of forest (Wirth et al., 2004).

The existing data of NBP were reviewed by Schulze et al. (2002). In a modeling exercise Ciais et al. (2005a) investigated the reasons which cause NEP and NBP to be positive (i.e. never reaching an equilibrium) (Fig. 19). It turns out that most continents are carbon sinks but for different reasons. It is increasing CO₂ which accelerates NEP in the Brazilian rain forest, it is increasing temperature which enhances growth in Siberia, and it is the deposition of nitrogen which accelerates NEP and NBP in Europe. However, recent accelerations of climate disturbances via drought and fire may even consume this sink (Ciais, 2005b).

6. The global carbon budget

Based on data in the IPCC (2002), Schimel et al. (2001) summarized the global C-budget and regional sinks (Table 1). 25% of the global fossil fuel emissions originated from the USA in the decade from 1990 to 2000, 17% from the EU, and 13% from

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Russia. On a global average, about 48% of the emissions are re-assimilated by vegetation. However, the terrestrial surface is also a source due to land use change. Thus, the global average sink shrinks to 22% of the emissions. Looking at the Euro-Siberian continent, it becomes apparent, that the total emissions of West Europe and Russia of 1.9 GtC yr^{-1} were balanced by a net carbon sink of 1.7 GtC yr^{-1} . Obviously, the Siberian boreal forest balances almost 90% of the Eurosiberian emissions.

This observation is remarkable, because the boreal forest of Siberia is to a large extent still a pristine forest, that has not been managed by wood extraction. Obviously, un-managed pristine forest is not in some kind of equilibrium of assimilation and respiration in the sense of Odum, but continues to accumulate carbon. This is despite the natural fires occurring in that region. The result has been confirmed by eddy correlation measurements (Schulze et al., 2000; Roeser et al., 2002), and a closer analysis of the tropospheric inversion models indicate (Ciais et al., 2005a) that the Siberian sink results from the temperature rise in that region. In contrast, if we were able to use the Siberian forest for timber and fiber, and if this forest would be logged within a short period of time, the data of Schulze et al. (2000) suggest that this region could turn into a carbon source, and the fossil fuel emissions of boreal Eurasia would remain in the atmosphere with unpredictable consequences for global climate.

The results of Siberian forest indicate, that land-use and land-use-change can affect global climate significantly, and the preservation of carbon stocks, and the maintenance of previously un-managed regions may be just as important as the reduction of fossil fuel emissions in view of global climate. The Kyoto protocol does not allow accounting of this natural sink, and even encourages its destruction by stating in the preamble that “the mere presence of carbon stock be not accountable”.

However, besides of this scenario of immediate destruction, it must also be considered that forests can be managed in a sustainable way. The Euroflux project (Fig. 9) is based on managed forests only, and all tower stations indicate a sink, but this does not account the export of C by harvest which takes place every 10 years by thinning and every 100 years by final harvest. Thus the net carbon sink of Europe appears to

be very small, mainly due to emissions from agriculture (Janssens, 2003).

It remains a task to separate management effects from environmental change, because the Kyoto protocol in its Article 3.4 will only allow accounting of human induced changes. At present there is no consensus, how to separate the different driving factors. The management effect can only be estimated by modeling. Using the Biome BGC-model Wirth et al. (2004) compared inventory and model simulations which indicate, that 22 to 37% of the observed change in carbon stocks of the state forest of Thuringia were the result of management. However, since German forest management operates at a high level of carbon stocks, the options for a further increase in stocks are very limited, and the high carbon stocks are not accountable since they existed before 1990. The forest sink could be increased by large scale harvest and the associated unaccounted C-loss. Model calculations show, that fast growing plantations would reduce the inventory stocks. Extending the rotation time would also be a mechanism to “buy for time”, but it would not increase the stocks in a sustainable way. The expected increase from changing softwood into hardwood is limited, and it would start to take place after about a half rotation period (50 years). Selective cutting systems would have limited effects, but protection of forest from wood extraction would have a long-lasting and surprisingly large effect, but it is not accountable in the Kyoto context.

7. Biodiversity effects

Up to now, the carbon cycle was treated independent of the species involved, and we are at the very beginning to observe and understand the effects of biodiversity on the carbon cycle. The most powerful observation was made by the “Biodepth” experiment of the EU (Hector et al., 1999), where experimental plots with variable biodiversity were established as grassland across Europe (Fig. 20A). It emerged, that plant biodiversity had an effect in all climatic regions of Europe. Loosing half of the plant species did decrease productivity by 20% (Fig. 20B). Also this result is under debate. It may hold only for early successional stages while in later successions the relation may reverse

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

(Thompson et al., 2005)

Biodepth focused on plant species, which is not enough for understanding the value of biodiversity, but we are at the very beginning to even detect biodiversity across all heterotrophes. A new approach was introduced by Schulze et al. (2004) using proteomics to show the presence of organismic groups in ecosystems, and their response to environmental change (Fig. 21). In fact, the plant cover contributes only very little to the total ecosystem biodiversity, which is dominated by microbial organisms. Although we must admit, that we do know very little about the functional effect of this diversity, proteomics give us an initial tool to detect the origin of a number of key extracellular enzymes such as laccases, cellulose or peroxidase. Based on this type of information, we hope that we can link the biological control of C sinks to the effects of specific species or functional groups in future.

8. The future C cycle

Energy consumption is very uneven across the globe. When relating energy consumption to GDP (Gross Domestic Product) the huge difference between developed and underdeveloped nations becomes obvious (Fig. 22). This is further underlined, if we investigate the number of persons without access to electricity across the globe (WBGU, 2004). It is to be expected, that the underdeveloped world will request its fair share in energy supply in the future. This will be amplified by the increase of population in these countries. Global human population is expected to reach 9 Billion in 2050. Thus, it is expected, that the total energy use will increase from about 400 EJ/yr (about 6 Gt C/yr) in 2004 to reach possibly more than 1700 EJ/yr (about 25 GtC/yr) in 2050. We are certain that biology will not be able to compensate such an increase in emissions, if this energy was produced from fossil fuels.

If the energy would be produced from fossil fuel, and if renewable energy came into place only after real shortage of fossil fuel becomes visible, then temperatures are expected to rise 3 to 4 K, and reach a maximum change of 0.4 K/decade. Clearly

BGD

2, 1283–1329, 2005

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

this would be damaging to many ecosystems independent if they were managed or not (IPCC-WG2, 2002). It could endanger the food supply to mankind. If one would try to maintain temperatures in a tolerable window of about 2 K change at a rate of maximal 0.2 K/decade, then dramatic changes in the energy production are required now, i.e. immediately. Figure 23 shows the traces of global temperature change as predicted by the A1 scenario of IPCC, and the observed rate of global climate change. Presently global change exceeds the most aggressive scenario of IPCC.

Considering, how this required energy could be produced, then relatively few options remain. Coal and oil will get increasingly depleted by 2050. Gas may still remain as major fossil fuel. Biomass for energy will not increase, because fertile land will be needed for agriculture. Wind power will increase, but the total amount of energy delivered remains limited, nuclear power could increase but this option contains also the thread of weapon proliferation. Solar energy appears the only viable energy source that could be developed to an extent to meet the demand, but it remains uncertain, if the energy required to build these systems will drop to an extent, that the systems become a net source.

The total energy requirement of fossil fuel is estimated to reach 25 Gt/yr, this large number is small when compared to the carbon pools that are still stored in soils and biomass (about 2100 Gt). If only 50% were activated through land use and land use change over the next 50 years, than an additional 21 Gt/yr could potentially be emitted. This would almost be as much as the peak energy requirement in 2050, and this would bring all attempts to heal the world by renewable energy to a fall. Thus, trying to solve the problem of global climate change without protecting terrestrial global carbon pools, and neglecting the biology of the carbon cycle will most likely fail.

9. Conclusions

- Climate change has stimulated C cycle research in an un-preceded manner

- The focus of such research has moved from CO₂ assimilation and net primary production research in the times of IBP towards ecosystem research during the decade of acid rain towards global understanding of sources and sinks
- Biodiversity has entered the biogeochemical research as inherent and important component
- Predictions of the future C cycle indicate that biology will not be able to compensate fossil fuel emissions. However, neglecting the potential emissions from land use and land use change may offset any attempts to mitigate climate change by technical measures of reduced fossil fuel use.

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Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

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**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

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BGD

2, 1283–1329, 2005

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

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BGD

2, 1283–1329, 2005

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

**Biological control of
the terrestrial carbon
sink**E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

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**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

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Biological control of the terrestrial carbon sink

E.-D. Schulze

Table 1. The global carbon budget of fossil fuel emissions and of the terrestrial regional C sink (Schimel et al., 2001).

	(Gt C a ⁻¹)	(%)
Fossil fuel emissions	6.3±0.4	
USA	1.6	25
EU-15	1.1	17
Russia	0.8	13
Germany	0.3	4.8
Atmospheric CO2 increase	3.2±0.1	
Net uptake of the oceans	1.7±0.5	
Net uptake of the continents	1.4±0.7	
Emissions due to land-use change	1.6±0.8	
Assimilation by vegetation	3.0	
USA	0.8	27
Europe	0.4	13
Siberia	1.3	43
Tropics	0.5	17

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

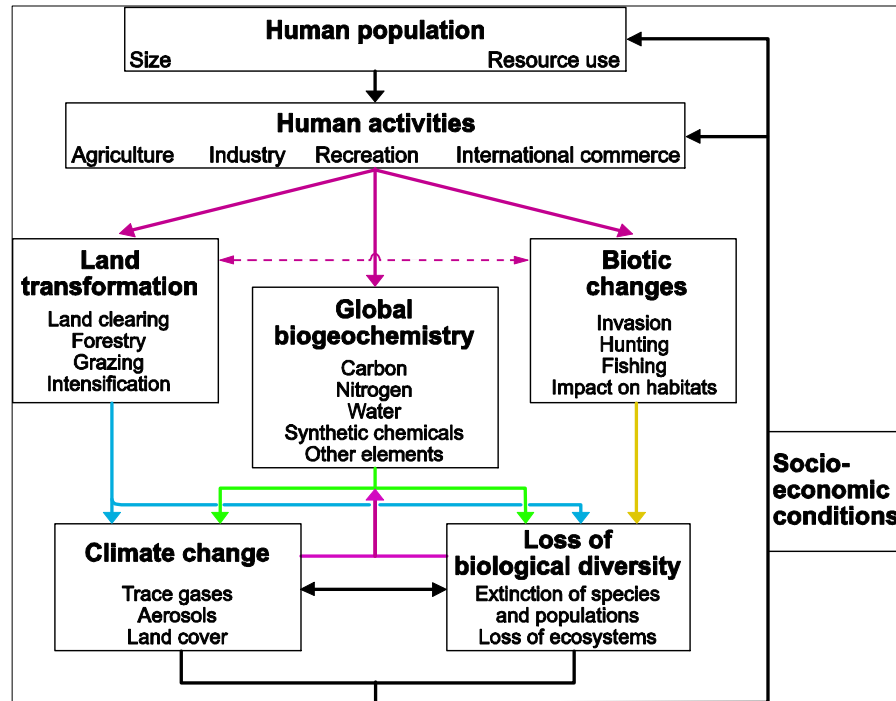


Fig. 1. Schematic presentation of the interactions between the human population and Global Change (Schulze, 2000).

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	

Print Version
Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

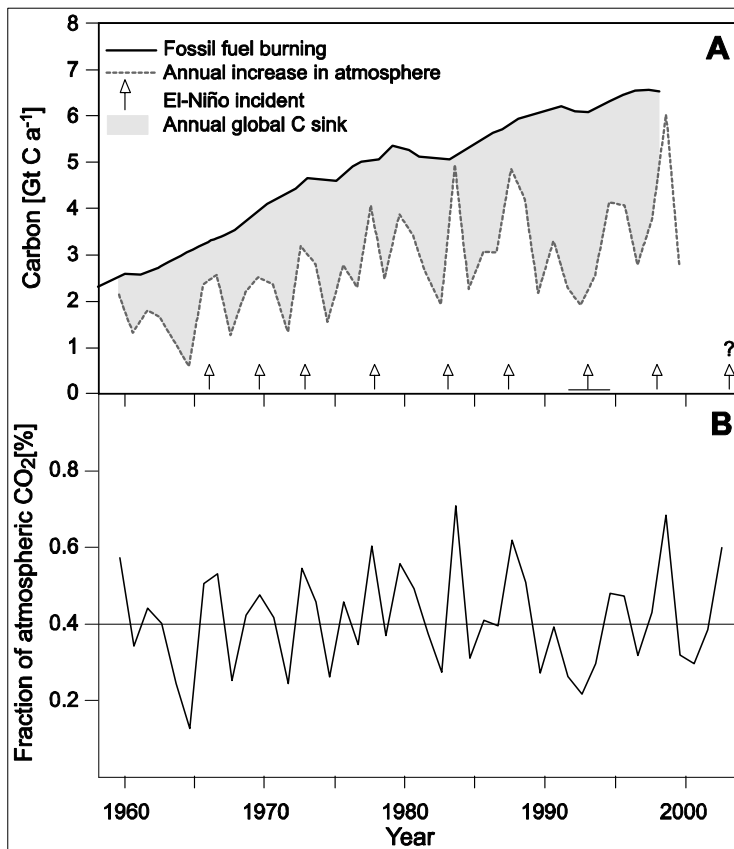


Fig. 2. (A) The annual change of fossil fuel burning and the increase in atmospheric CO₂ concentration. The difference between the two lines represents the global carbon sink. The variation is to a large extent related to El Niño events (Prentice et al., 2002), (B) The fraction of fossil fuel C that remains in the atmosphere over time (Heimann, personal communication).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

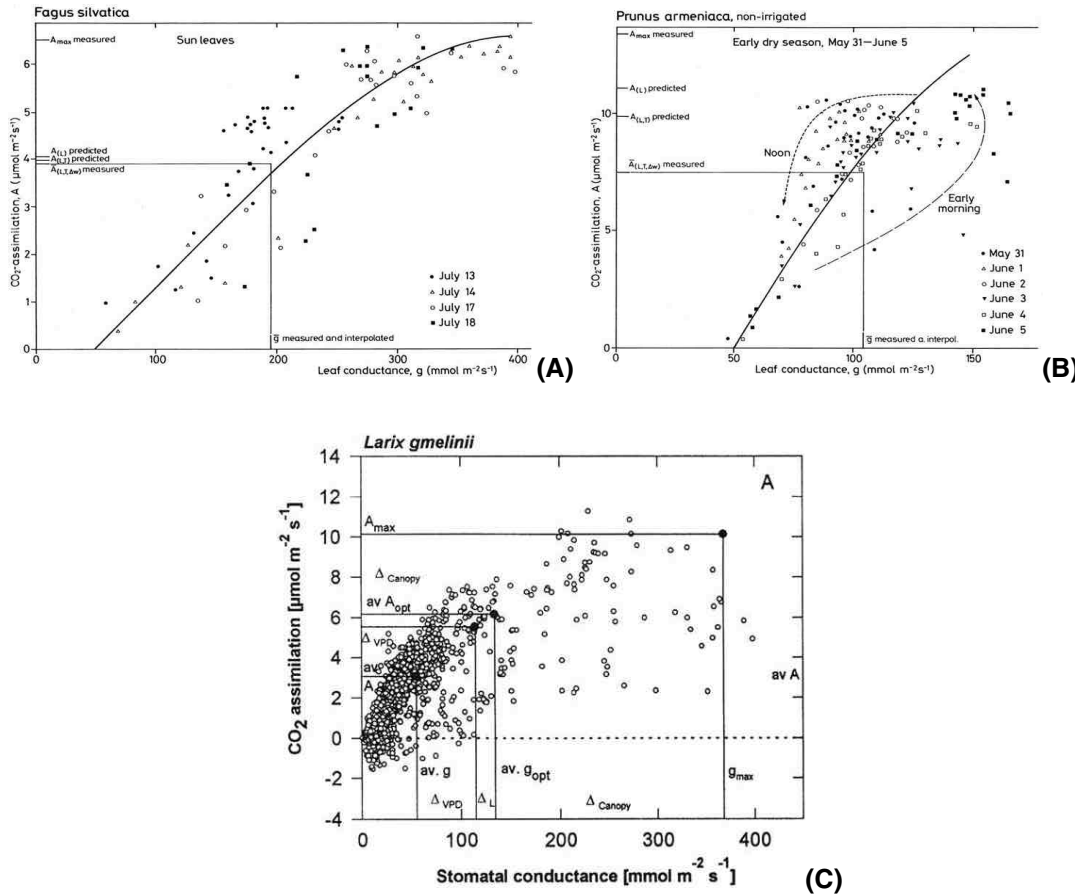


Fig. 3. CO₂ assimilation as related to leaf conductance for (A) *Fagus sylvatica* in the temperate deciduous forest of Germany (Schulze, 1970), (B) *Prunus armeniaca* in a run-off Fram of the Negev (Schulze and Hall, 1982), and (C) *Larix gmelinii* in East Siberia (Vygodskaya et al., 1996).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

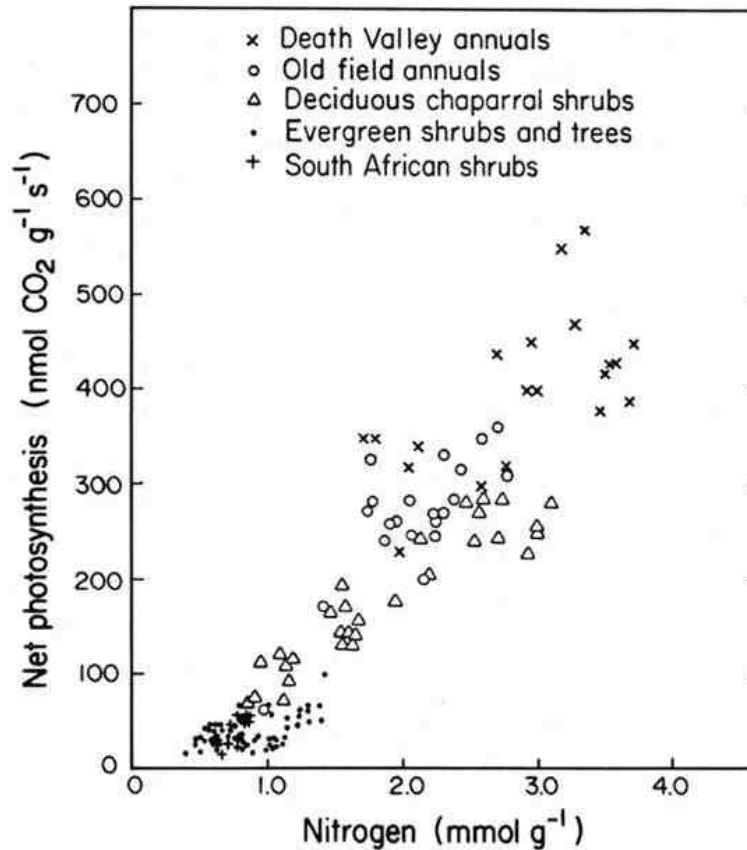


Fig. 4. CO₂-assimilation of leaves in different vegetation types as related to leaf nitrogen concentration (Field and Mooney, 1986).

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze



Fig. 5. Individual stomata of *Polypodium vulgare* respond to dry and humid air. In this case, the epidermis had been peeled from the leaf and exposed on a water surface, maintaining a small air bubble below few stomata, to simulate the sub-stomatal cavity. The response is reversible, i.e. the closed stomata open with humid air and the open stomata close with dry air (Lange et al., 1971).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Print Version](#)[Interactive Discussion](#)

EGU

Biological control of the terrestrial carbon sink

E.-D. Schulze

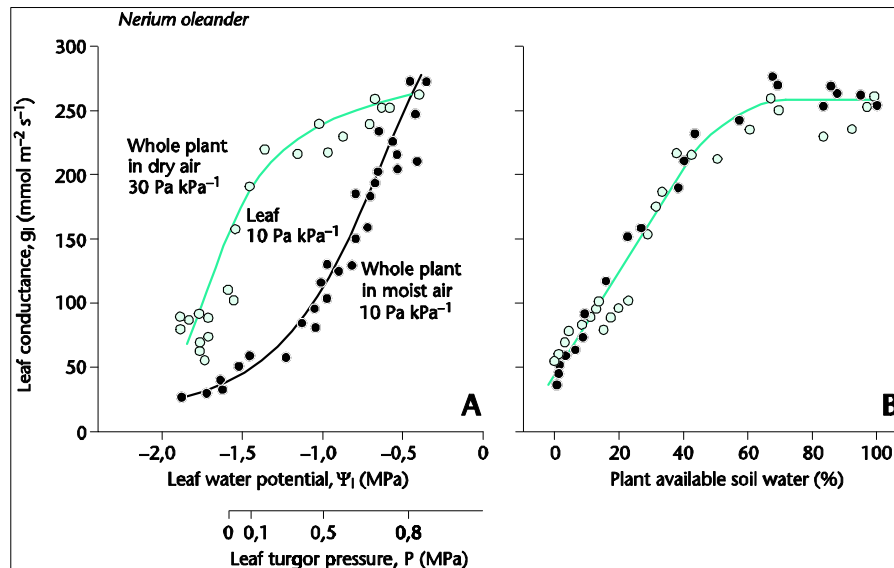


Fig. 6. Leaf conductance as related (A) to leaf water potential and (B) to plant available soil water. The whole plant was exposed either to dry or to moist air, which shifted the leaf water potential by 1 MPa because of differences in whole plant transpiration in dry versus moist air. In both situations, a single leaf was maintained at constant air humidity, in order to avoid stomatal closure. Then, then the plant was dried out, and the figures (A) and (B) show the traces of the dry out cycle. In both situations (exposure to dry and humid air) the plant responds to the common signal of plant available soil water (Turner et al., 1985; Schulze, 1986).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

Biological control of the terrestrial carbon sink

E.-D. Schulze

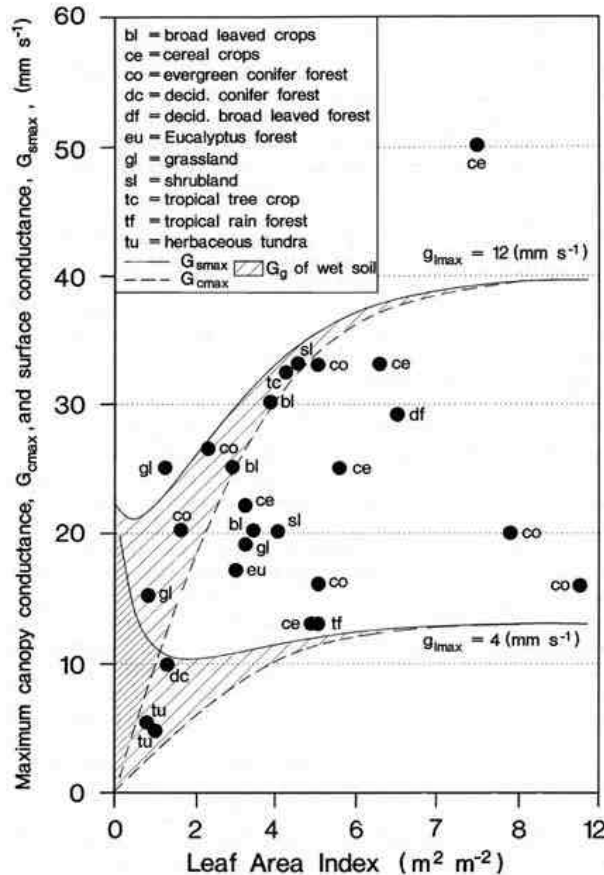


Fig. 7. Maximum canopy conductance (dashed line) and surface conductance (solid line) as related to leaf area index. The difference between canopy and surface conductance is a measure of the contribution of the soil surface (Kelliher et al., 1993).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

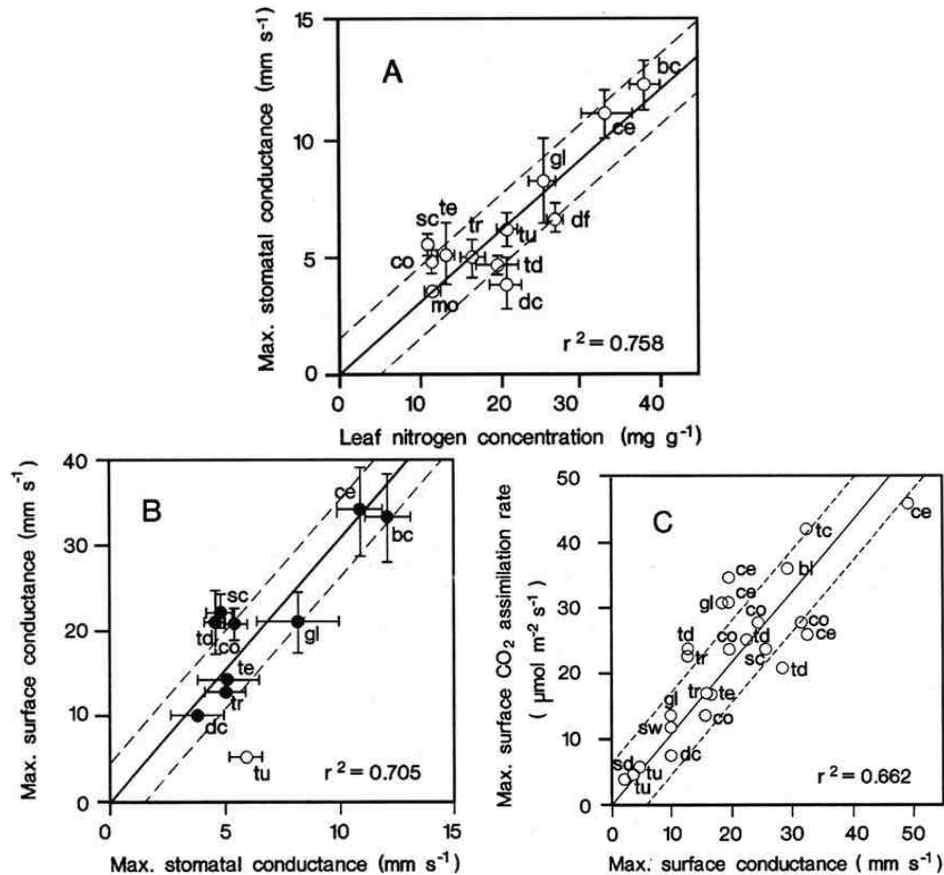


Fig. 8. (A) Maximum stomatal conductance as related to leaf nitrogen concentration, (B) maximum surface conductance as related to stomatal conductance, and (C) maximum CO_2 assimilation as related to maximum surface conductance for leaves from different vegetation types (Schulze et al., 1994).

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Print Version](#)
[Interactive Discussion](#)

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

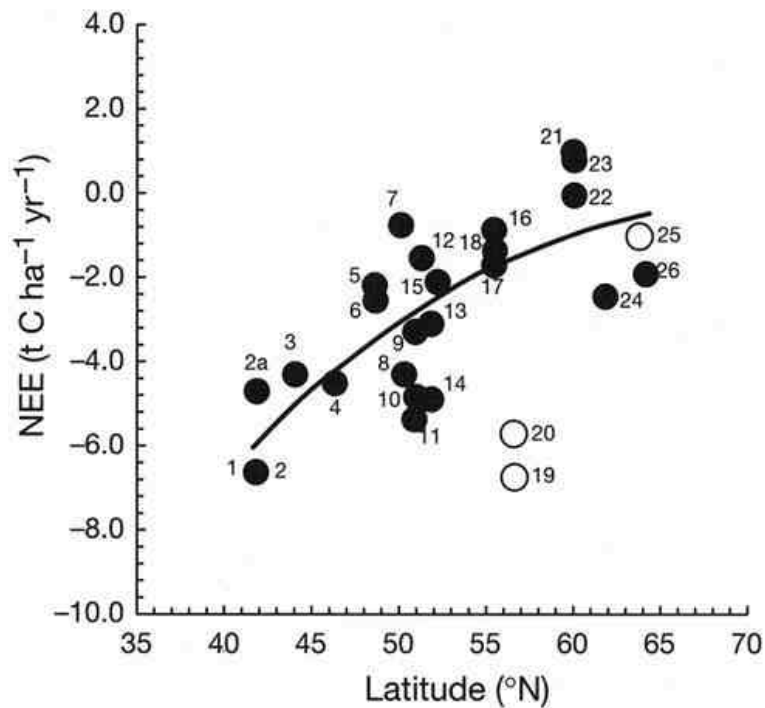


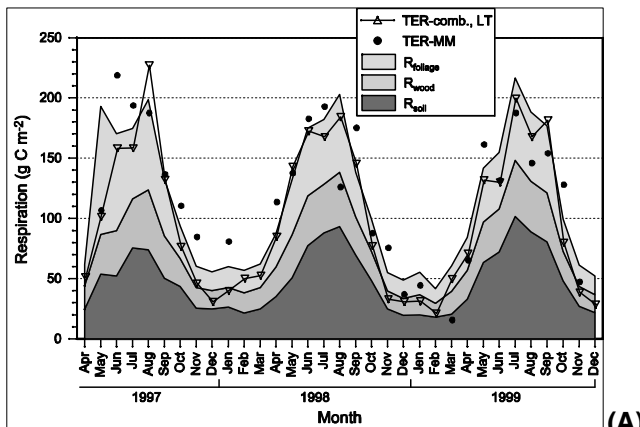
Fig. 9. The annual balance of net ecosystem exchange as related to latitude at 26 stations of the Euroflux network (Valentini et al., 2000).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Print Version](#)[Interactive Discussion](#)

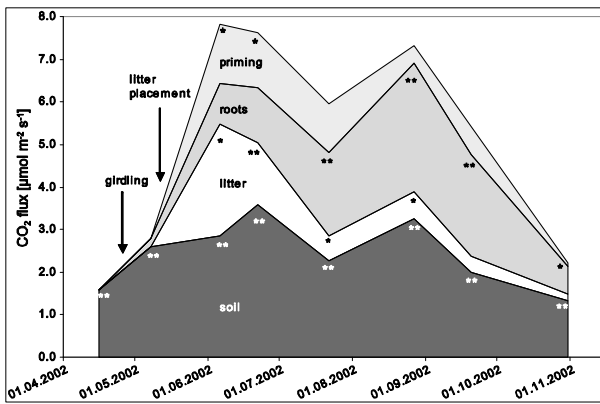
EGU

Biological control of the terrestrial carbon sink

E.-D. Schulze



(A)



(B)

Fig. 10. (A) Ecosystem respiration as measured by eddy covariance (TER-MM) and modeled by temperature (TER-comb) and the partitioning into respiration of needles, wood and soil as related to season at the Waldstein site (Rebmann, 2004). **(B)** Components of soil respiration in a girdled and un-girdled forest at the Wetzstein site. This figure also shows the effect of adding litter on soil respiration (Buchmann, in Schulze, 2003).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

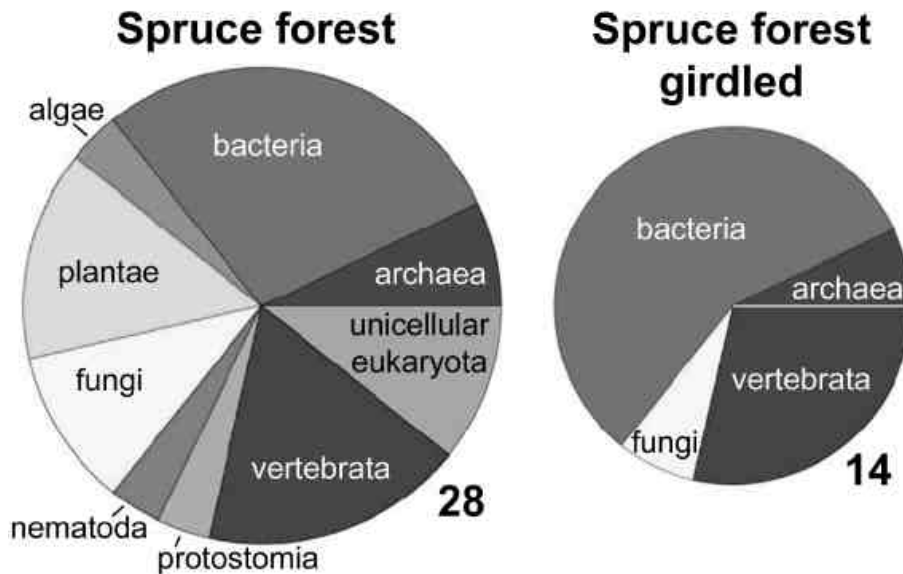


Fig. 11. Comparison of the phylogenetic contribution in protein pools derived from a natural and a girdled spruce forest. Area and number next to the pie charts represent the number of proteins (Schulze et al., 2004).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Print Version](#)[Interactive Discussion](#)

EGU

Biological control of the terrestrial carbon sink

E.-D. Schulze

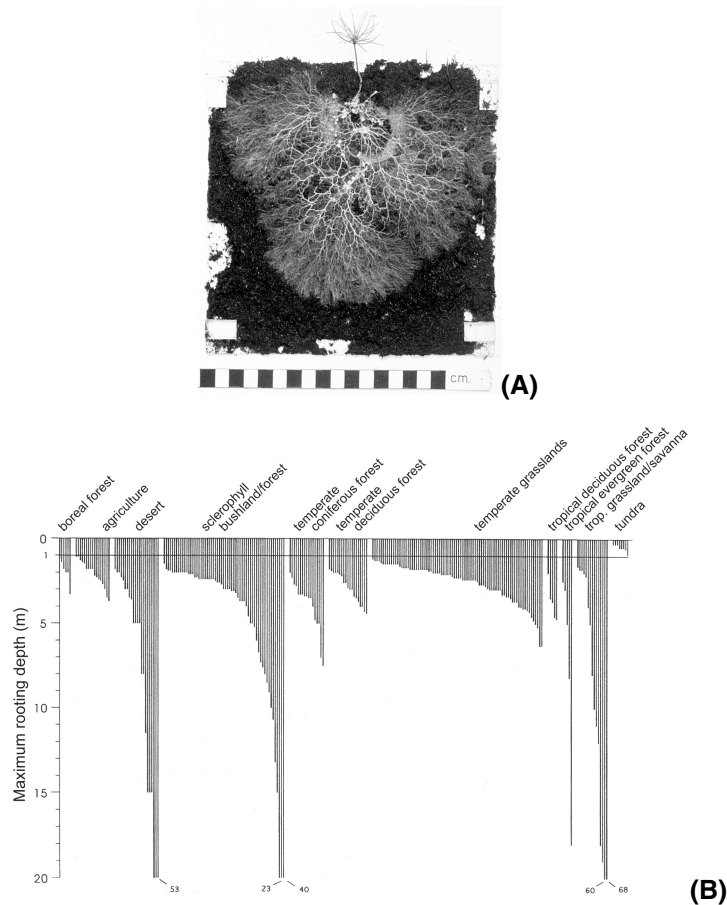


Fig. 12. (A) The distribution of biomass in a seedling of *Picea abies* growing in a Petri dish into aboveground biomass, roots and mycorrhizae (by courtesy of D. Read), **(B)** Maximum rooting depth in different biomes (Canadell et al., 1996).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

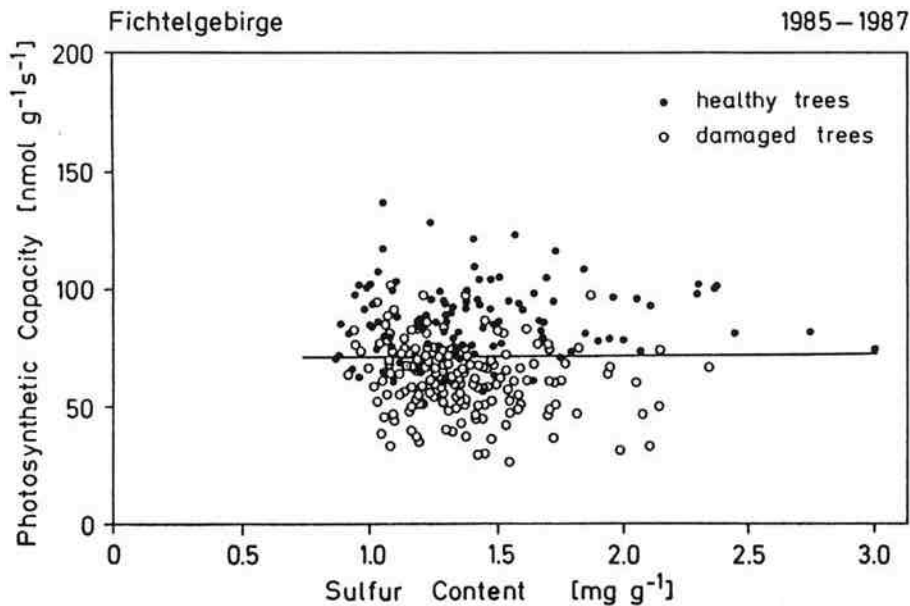


Fig. 13. The relation between photosynthetic capacity, as related to dry weight, and sulfur concentration of healthy and damaged needles (Lange et al., 1989).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Print Version](#)[Interactive Discussion](#)

EGU

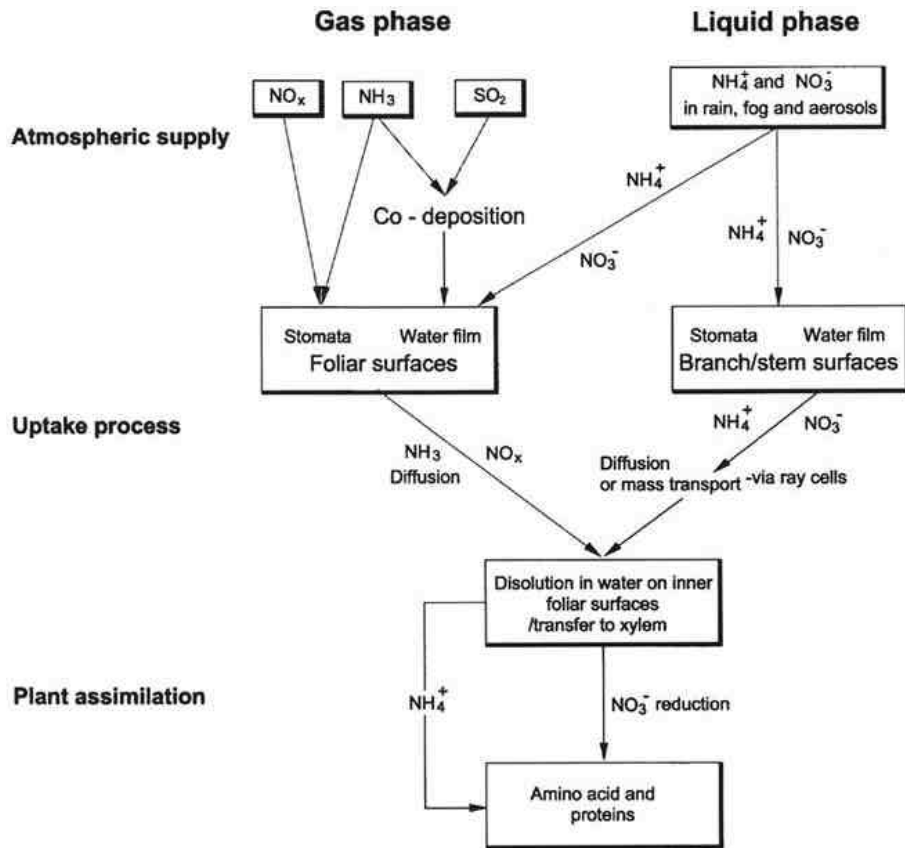


Fig. 14. Pathways of uptake of ammonia and nitrate through aboveground biomass of leaves and stem surfaces (Harrison et al., 2000).

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

EGU

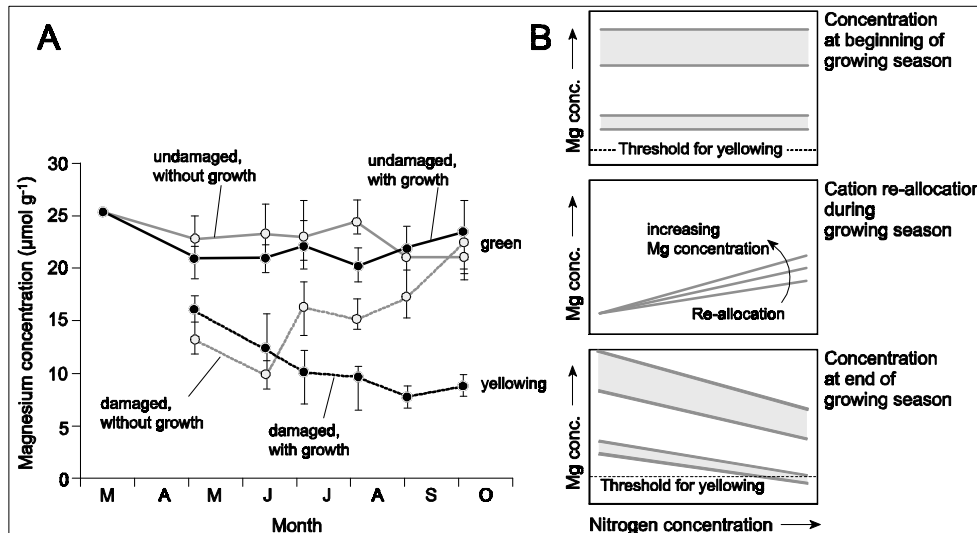


Fig. 15. (A) The response of needles with respect to magnesium concentration after clipping of buds and thus inhibiting growth and on intact twigs on trees showing symptoms of forest decline with yellow old needles and on healthy looking trees. The experiment shows that yellow needles become green if growth is inhibited (Lange et al., 1987). (B) A conceptual model of magnesium-nitrogen interaction. Re-allocation of magnesium from old to young needles results in needle yellowing if the supply of Mg is limiting (Oren et al., 1989).

Biological control of the terrestrial carbon sink

E.-D. Schulze

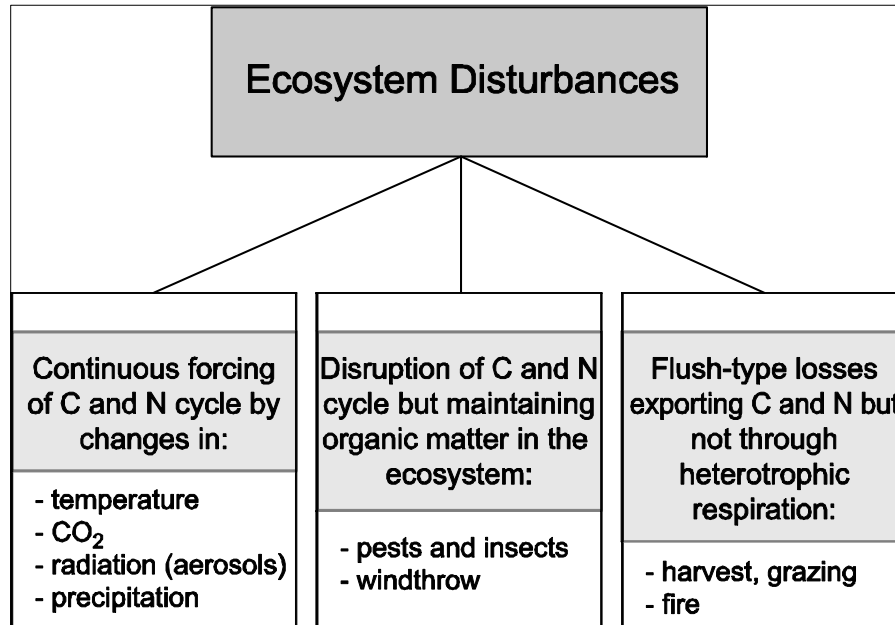


Fig. 16. A carbon-related classification of ecosystem disturbances (Schulze et al., 1999).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

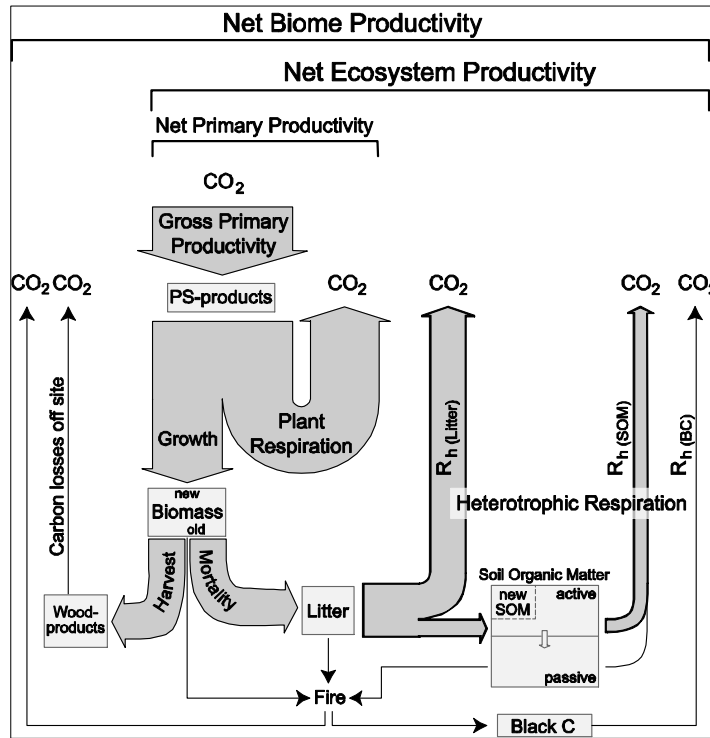


Fig. 17. A conceptual scheme of the carbon flow through ecosystems, with input by photosynthesis (Gross primary productivity) and C-losses plant respiration. The balance describes plant growth (Net primary productivity, NPP). At the same time, plants shed biomass that reaches the soil and which is decomposed by microorganisms and a microbial respiratory C-loss. The balance of CO₂-assimilation and all respiratory processes is defined as Net ecosystem productivity (NEP). However, besides respiratory losses, there are carbon losses that bypass respiration. These are by fire, harvest and grazing. The total carbon balance, including respiratory and non-respiratory losses, is termed Net Biome Productivity (NBP) (Schulze and Heimann, 1998; Schulze et al., 2000).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

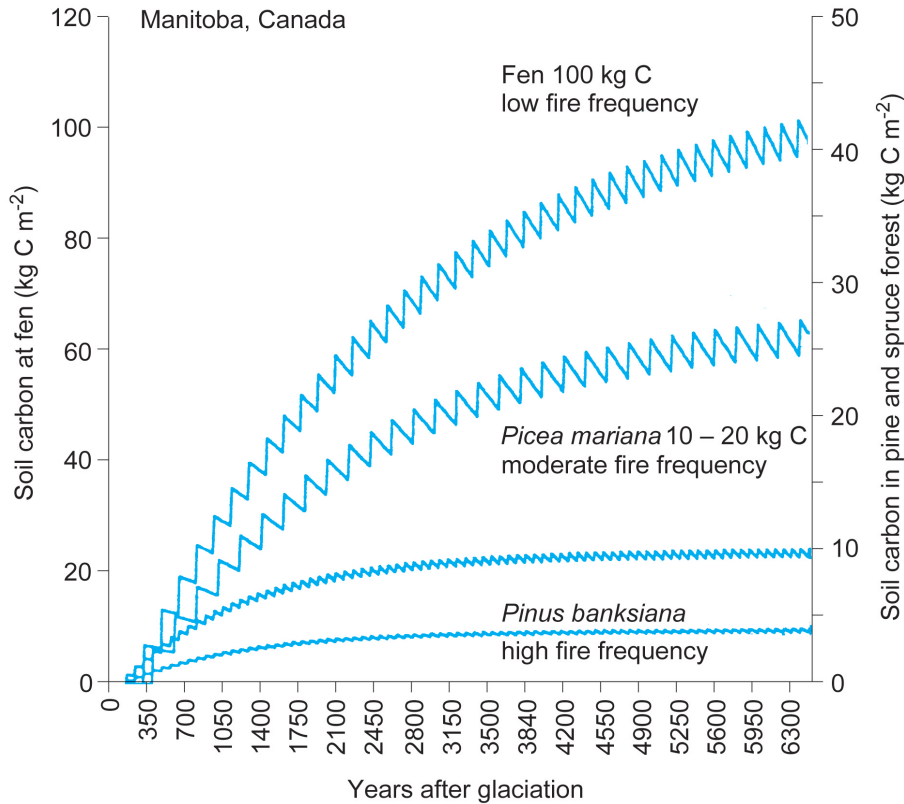


Fig. 18. Soil carbon accumulation in boreal forest ecosystems with low and high fire frequency (Harden et al., 2000).

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

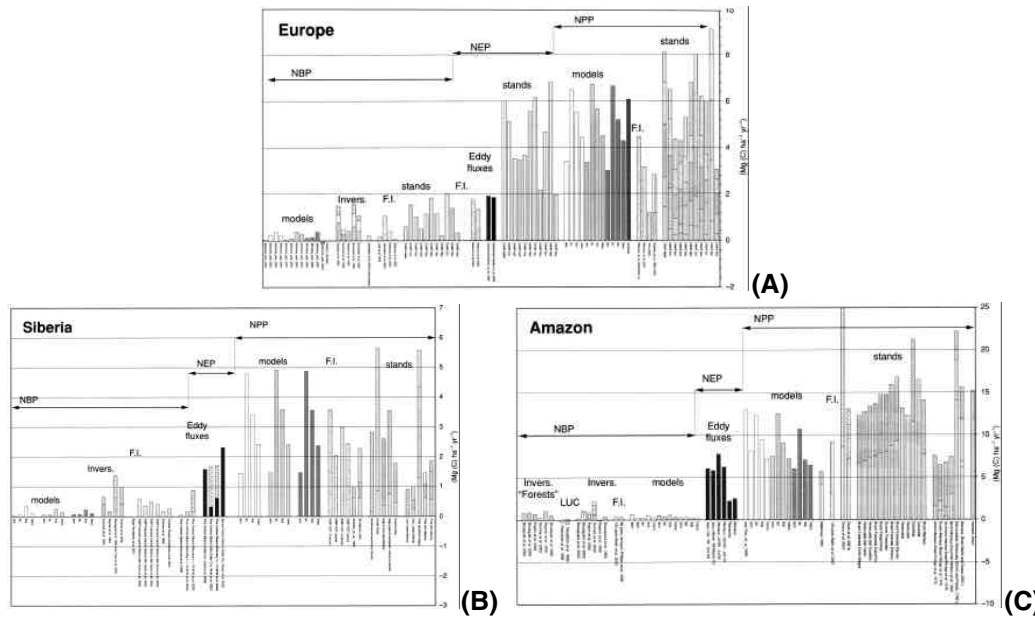


Fig. 19. A summary of existing data of NPP, NEP and NBP over **(A)** Western Europe, **(B)** Siberia and **(C)** Amazonia. NPP components are coarse roots, fine roots, stems, twigs and branches, leaves. F.I.: forest inventories. Models: four global models HRBM, IBIS, LPJ, TEM from left to right prescribed with rising CO_2 , climate variability and rising CO_2 , climate variability and rising CO_2 and crop establishment. Eddy fluxes: up-scaled Euroflux towers using satellite and gridded climate fields. Inverse: atmospheric inversions are “forest NBP” obtained from subtraction of cropland and pasture fluxes (Ciais, 2005).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

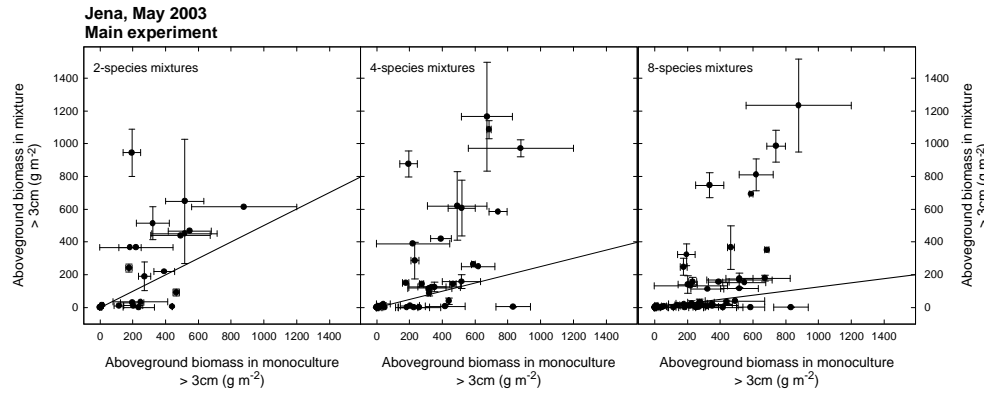
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Full Screen / Esc

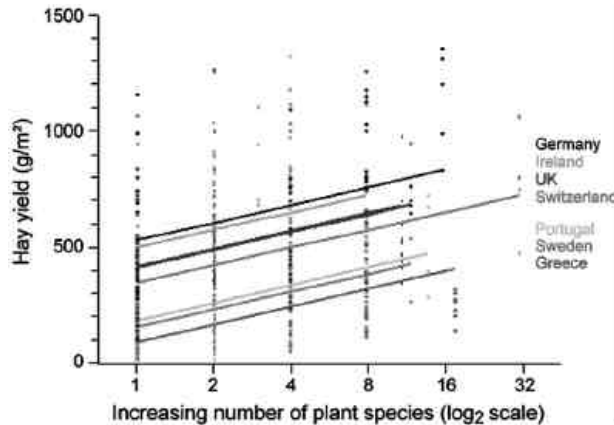
Print Version

Interactive Discussion

EGU



(A)



(B)

Fig. 20. (A) The Jena Biodiversity experiment. 440 plots differing in their diversity between 1 and 60 species were established. A part of the plots combined the most dominant species (Roscher et al., 2004). (B) The relation between hay yield and plant biodiversity in experimental fields across Europe (Hector et al., 1999).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

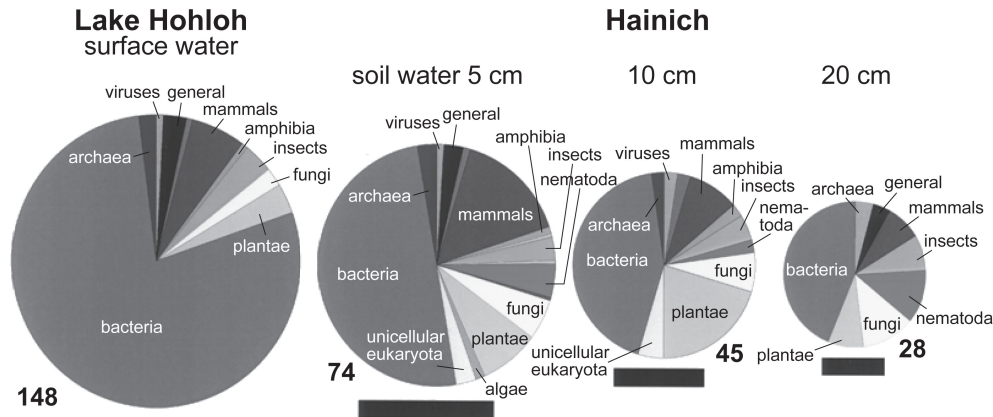


Fig. 21. Identification of phylogenetic groups in a lake and in soil water of a diverse deciduous forest (Schulze et al., 2004).

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

Print Version

Interactive Discussion

Biological control of the terrestrial carbon sink

E.-D. Schulze

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Print Version

Interactive Discussion

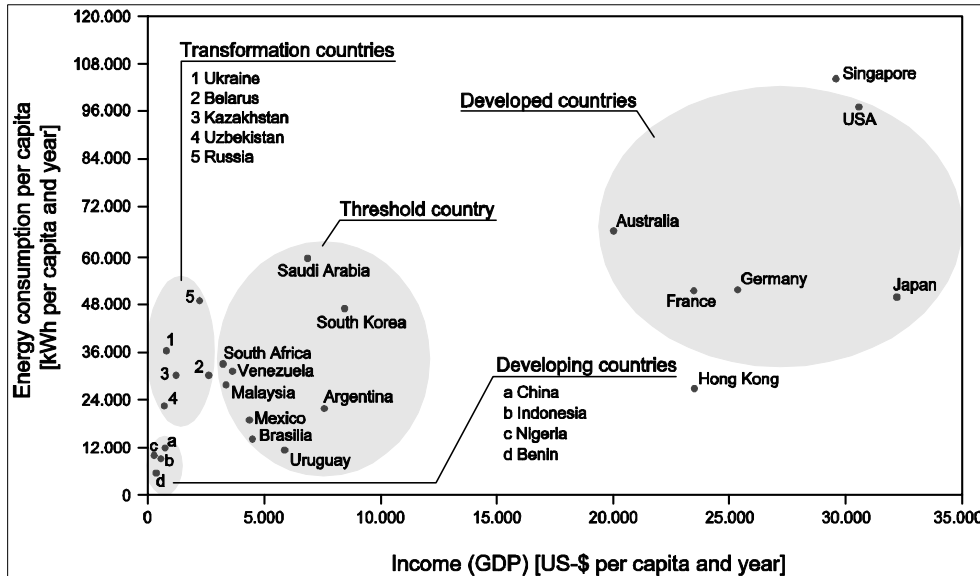


Fig. 22. Energy consumption per capita versus income per capita in various nations (WBGU, 2004).

**Biological control of
the terrestrial carbon
sink**

E.-D. Schulze

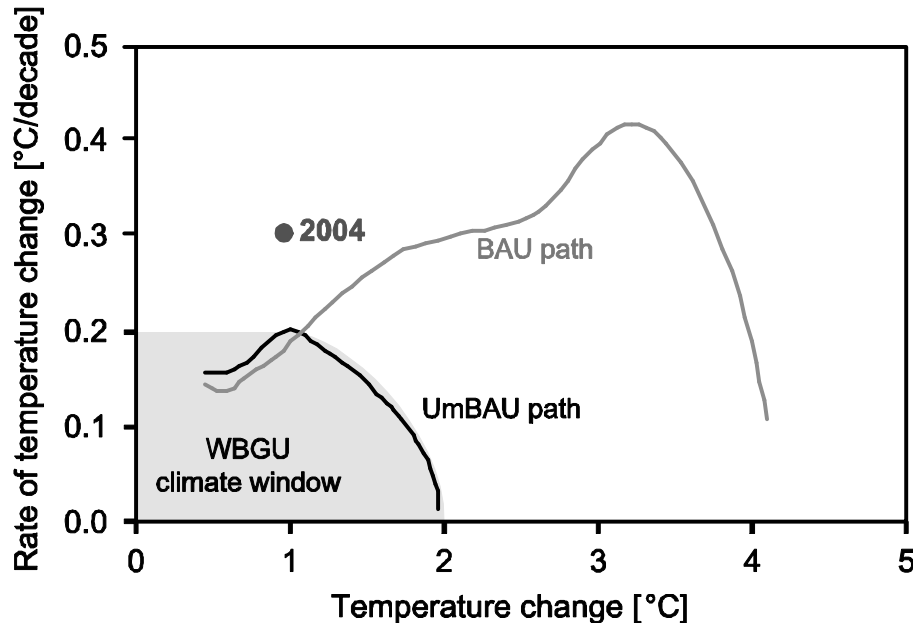


Fig. 23. Rate of temperature change and absolute change in temperature for an energy production without change in energy source (Business as usual BAU-path) or with change in energy source towards renewable energies (UmBAU-path) according to WBGU (2004).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Print Version](#)[Interactive Discussion](#)

EGU