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Tectonic and climatic control of the changes in the sedimentary record of the Karnali River section (Siwaliks of western Nepal)

PASCALE HUYGHE,^{1,*} JEAN-LOUIS MUGNIER,¹ ANANTA PRASAD GAJUREL^{1,2}
AND BERNARD DELCAILLAU³

¹Laboratoire de Géodynamique des Chaînes Alpines et Université Joseph Fourier, Maison des Géosciences, Grenoble Cedex, France (email: huyghe@ujf-grenoble.fr), ²Department of Geology, Tribhuvan University, Kathmandu, Nepal, and ³Laboratoire de Morphodynamique Continentale et Côtière, Université de Caen, 14032 Caen Cedex, France

Abstract

A multidisciplinary study was conducted on the section of the Siwalik Group sediments, approximately 5000 m thick, exposed along the Karnali River. Analysis of facies, clay mineralogy and neodymium isotope compositions revealed significant changes in the sedimentary record, allowing discussion of their tectonic or climatic origin. Two major changes within the sedimentary fill were detected: the change from a meandering to a braided river system at *ca* 9.5 Ma and the change from a deep sandy braided to a shallow sandy braided river system at *ca* 6.5 Ma. The 9.5-Ma change in fluvial style is contemporaneous with an abrupt increase of $\Sigma_{Nd}(0)$ values following a $\Sigma_{Nd}(0)$ minimum. This evolution indicates a change in source material and erosion of Lesser Himalayan rocks within the Karnali catchment basin between 13 and 10 Ma. The tectonic activity along the Ramgarh thrust caused this local exhumation. By changing the proximity and morphology of relief, the forward propagation of the basal detachment to the main boundary thrust was responsible for the high gradient and sediment load required for the development of the braided river system. The change from a deep sandy braided to a shallow sandy braided river system at approximately 6.5 Ma was contemporaneous with a change in clay mineralogy towards smectite-/kaolinite-dominant assemblages. As no source rock change and no burial effect are detected at that time, the change in clay mineralogy is interpreted as resulting from differences in environmental conditions. The facies analysis shows abruptly and frequently increasing discharges by 6.5 Ma, and could be linked to an increase in seasonality, induced by intensification of the monsoon climate. The major fluvial changes deciphered along the Karnali section have been recognized from central to western Nepal, although they are diachronous. The change in clay mineralogy towards smectite-/kaolinite-rich assemblages and the slight decrease of $\Sigma_{Nd}(0)$ have also been detected in the Bengal Fan sedimentary record, showing the extent and importance of the two major events recorded along the Karnali section.

Key words: clay mineralogy, climate, facies analysis, Himalaya, Karnali section, neodymium isotope, Neogene, Siwaliks of Nepal, tectonics.

INTRODUCTION

Peripheral foreland basins develop in response to the load of thickened crust resulting from continental collision (Beaumont 1981; Allen *et al.* 1986). The sedimentary infill of foreland basins records the interaction between the growth of the thrust wedge, the isostatic adjustments of the cratonic lithosphere to thrust loading and additional bending moments, eustasy and the surface processes that redistribute material from the mountain belt into the surrounding basins. As a consequence, studying the fill of an ancient setting could serve as a key for the reconstruction of both the growth of the orogen and climate evolution. However, it is increasingly obvious that these two phenomena interact (Molnar & England 1990) and their own geological signature is very difficult to decipher. Nonetheless, a multidisciplinary study should provide clues to distinguish between these phenomena.

The aim of the present paper was to analyze the major changes in the fill of the Tertiary Himalayan

Foreland Basin (Siwalik Group in western Nepal) in the light of a pluridisciplinary study and to try to determine their tectonic or climatic origin. Interpretation was mainly based on lithofacies analysis to characterize the fluvial style (hydraulic conditions), Nd isotope geochemistry as a source marker, and clay mineralogy, which provides information on silicate weathering (Derry & France- Lanord 1997).

GEOLOGICAL SETTING

The Siwalik Formations represent the Cenozoic Foreland Basin sequence of the Himalayan Belt. They are located beneath the Gangetic Plain or in the outer part of the fold and thrust belt (DeCelles *et al.* 1998). The Outer Himalaya, located between the Lesser Himalaya to the north and the Terai Plain to the south (Fig. 1), is usually called the 'Siwalik Belt' and is affected by several thrusts. These thrusts branch off the main decollement, which lies at a depth of 4–5 km beneath the outer belt (Mugnier *et al.* 1993, 2004; DeCelles *et al.* 1998; Powers *et al.* 1998) and roots in the middle crust beneath Tibet (DeCelles *et al.* 2001). The thrust system of the Himalayan belt is characterized by a general forward propagation of the tectonics, with the activity of the main central thrust (MCT), mainly before 18 Ma (DeCelles *et al.* 2001), and the main frontal thrust (MFT), which is has been active since 1.8–2.4 Ma (Mugnier *et al.* 2004). The detail of the thrust sequence is complex and involves numerous out-of-sequence reactivations (Mugnier *et al.* 1998).

The foreland basin sequence is classically subdivided into three units on the basis of lithostratigraphic criteria (Auden 1935; Rao *et al.* 1988). These are named the Lower, Middle and Upper Siwaliks and their ages range from Middle Miocene to Pleistocene (Gautam & Rösler 1999). The Siwalik Group of western Nepal represents a typical fluvial fining upward succession, with the lower unit consisting of fluvial channel sandstones alternating with oxidized calcareous paleosols, the middle unit consisting of very thick channel sandstones and drab-colored histosols, and the upper unit comprising mainly gravely braided river deposits (DeCelles *et al.* 1998; Nakayama & Ulak 1999). Although this succession looks like an autogenetic succession in a foreland basin (Miall 1985; Delcaillau 1992), the Himalayan tectonics and the climate of southeast Asia also control the sedimentation (Brozovic & Burbank 2000); the Siwalik deposition is contemporaneous with the propagation and/or the reactivation of thrust faults and with major climatic events, like the intensification of the Indian monsoon system at *ca* 8 Ma (Kroon *et al.* 1991; Prell *et al.* 1992) or the global Pliocene cooling Molnar & England 1990; Ravelo *et al.* 2004). Stable isotope geochemical studies on herbivore tooth enamel, on fossil organic matter and on paleosol carbonates (Quade *et al.* 1992, 1995) show a dramatic change from forest to tropical grassland and strongly suggest that a major climatic change occurred in the Siwalik Basin at 7–6 Ma throughout the region from Pakistan to Nepal. The Karnali River drains through the Siwaliks in western Nepal (Fig. 1). Its drainage basin has an area of approximately 52 000 km² and it is one of the most important rivers in Nepal. The basin comprises formations belonging to the High Himalaya, Lesser Himalaya and Neogene molasses of the Siwaliks. The Karnali River offers a stratigraphic section approximately 5000 m thick, representing the complete series from the Lower to the Upper Siwalik Formations (Mugnier *et al.* 1998). The lower 3560-m-thick part of the section has been accurately dated from 15.8 to 5.2 Ma by magnetostratigraphy (Gautam & Fujiwara 2000), which permitted assigning ages to the various trends and events recorded in the sedimentary succession. For the younger deposits, from the deposits dated at 5.2 Ma to the top, a constant mean sediment accumulation rate of 33 cm/ky was adapted for the present study (Gautam & Fujiwara 2000).

SEDIMENTOLOGICAL FIELD OBSERVATIONS: STRATAL ORGANIZATION AND DEPOSITIONAL ENVIRONMENTS

LITHOSTRATIGRAPHY AND FACIES ASSOCIATIONS

Formations were analyzed and classified into various facies associations from field observations. The classification used is mainly based on bedforms, the nature of the contacts of beds, assemblages of sedimentary structures, grain size, sandstone : mudstone ratio, and thickness of sandstone beds. Miall's (1978, 1985) lithofacies codes and architectural elements were followed in the present work. The above-mentioned field criteria were also chosen for a regional comparison of the evolution of the Neogene fluvial system, as carried out by Nakayama and Ulak (1999). Altogether, seven facies associations (KFA1–KFA7) have been distinguished (Table 1). The upper and/or lower limits of the facies KFA1–KFA7 are all positioned along the section with reference to the 476 samples collected by Gautam and Fujiwara (2000) (Fig. 2). In the present

paper, the term 'sandstone' is used; however, most of the observed sandstones have a variable silty-clayey matrix and some of them should be called greywacke.

KFA1 facies association

The KFA1 facies association (Table 1; Fig. 2) is characterized by a predominance of thick-bedded mudstone (1–3 m) interbedded with fine-grained gray calcareous sandstone (0.5–2.0 m). Typically, 1.5–5.0 m of fining upward sequences from sandstone to mudstone beds are found in this facies. The basal contact (with the underlying sequence) is marked by an erosion surface and sandstone gradually passes upward into mudstone. Calcareous nodules occur on the 10–20-cm-thick upper surfaces of the sandstone beds and occasionally in the bioturbated highly pigmented mudstones above. Rootlets may be found, indicating the development of vegetation at the top of the fining upward sequence. Laterally accreted architectures may be found in the thicker sandstone beds. Some sequences with reddish thicker beds of medium-coarse sandstone and minor mudstone beds are found in the KFA1 series. Their basal surface is more erosive than fine-grained calcareous sandstone and the base of each sandstone bed contains many intraformational mud clasts. For the overall KFA1 facies association, the sandstone to mudstone ratio, abundance of laterally accreted sandstone and intraformational clasts increase stratigraphically upward.

KFA2 facies association

The KFA2 facies association (Table 1; Fig. 2) is characterized by alternating medium-coarsegrained sandstone beds and a set of several finegrained muddy sandstone beds. The frequency of mudstone layers is reduced against the KFA1 association. The coarser sandstone beds are 2–10 m thick. They are frequently ribbon-lensoid in shape with lateral accretionary architectures. Their basal surface, which is slightly erosive, is overlain by many intraformational mud clasts, of centimeter to decimeter scale. The top parts of these sandstone beds consist of finer bioturbated sandstone. The finer-grained sandstone beds have a sheet-like geometry and tabular, sometimes slightly undulated, basal surfaces. They show centimeter-scale laminations and locally grade into mudstone beds, the top of which is truncated by the overlying fine-grained sandstones. Therefore, the fine-grained muddy sandstones could be amalgamated sandstones. The thickness of these fine-grained sandstone beds and intercalated mud layers varies, from 20 to 100 cm and from 10 to 20 cm, respectively.

KFA3 facies association

The KFA3 facies association (Table 1; Fig. 2) consists of an alternation of thick medium-coarse-grained gray sandstone beds and thinner sequences of fine-grained muddy sandstone and mudstone beds. The basal surface of coarser grained sandstone beds is erosional. Many centimeter-to decimeter-sized intraformational mud clasts and rounded indurated gravels derived from older Siwalik formations are commonly found. Horizontal stratification and mostly trough crossstratification are developed in the medium-coarsegrained sandstone beds, which form accretionary architectures. The coarser sandstone beds are 5–8 m thick. The muddy sandstone beds are locally ripple laminated and frequently overlie bioturbated green to gray mudstone beds. Both muddy sandstone and mudstone beds have a thickness of approximately 2–3 m. Some thin dark-gray mudstone beds occur stratigraphically upward.

KFA4 facies association

The KFA4 facies association is characterized by light-gray, thick and coarse- to very coarsegrained horizontal and trough cross-stratified sandstone beds (Table 1; Fig. 2). Pebbly sandstones and gray-brown to gray-green mudstones are interbedded. In most cases, the coarse sandstone beds approximately 10 m thick have a nonerosional planar basal surface, whereas a few thinner coarse sandstone beds show several decimeter-scale deep erosional scours. At their bottom, the coarse sandstone beds contain layers of intraformational mud clasts, reaching up to 50 cm in thickness. Pluri-decimetric lenses of laminated clays and paleosols may also be found in these layers. The grain size of the thick sandstone beds decreases stratigraphically upward and the sandstone beds show ripples at their top. The thickness of interbedded pebbly sandstone ranges between 30 and 100 cm. The majority of the subrounded to rounded gravels and cobbles are of quartzite with a length and breadth of up to 5 and 15 cm, respectively, whereas the Siwalik cobbles show a length of up to 20 cm. Mudstone beds, when present, are 50 cm to 3 m thick. Rootlets have not been found in mudstones. Near the top of the KFA4 facies association, several sandy beds containing inclined, sometimes flat gravels and cobbles occur. These beds are approximately 60–100 cm thick and truncate the underlying light-gray sandstone. They occur every 2–10 m and no mudstone is observed in between. Gravels and cobbles consist mainly of quartzite, subangular sandstone and mudstone, the latter showing a 'tile-type' geometry with a common orientation (Fig. 2). No grading is observed in these beds and the matrix is composed predominantly of very coarse sand.

KFA5 facies association

The KFA5 facies association is characterized by light-gray and thick, coarse- to very coarse-grained sandstone beds, pebbly sandstones and gray mudstones (Table 1; Fig. 2). The coarse- to very coarse-grained sandstones are commonly planar and cross-stratified. Trough stratified gravel beds are also observed. Clasts consist essentially of 3–10-cm rounded quartzite and sandstone pebbles. The basal surface of coarse- to very coarse-grained sandstone beds is mostly nonerosive and almost flat, whereas it is erosive in the case of the pebbly sandstones. Mudstone beds are rare and their thickness varies from 30 to 100 cm. *KFA6 facies association* The KFA6 facies association consists of coarse buff sandstone beds interlayered with gray to green mudstones (Table 1; Fig. 2). The proportion of buff sandstones relative to mudstones increases stratigraphically upward. The basal surface of the very coarse-grained sandstones is erosive, and exotic gravels, mainly of quartzite, up to 10 cm in diameter occur at the base of these beds. Sandstones of the KFA6 facies association are homogenous, poorly indurated and contain fewer white mica in comparison with the sandstones of the former facies. Cross-stratifications and convolute features may be observed. The thickness of the coarse-grained sandstones ranges from a few decimetres to approximately 3 m. Mudstone beds have sizes varying between several decimeters to a few meters and show bioturbation. A few conglomerate beds, approximately 1 m thick, occur stratigraphically upward.

KFA7 facies association

The KFA7 facies association consists of moderately sorted pebble- and cobble-sized massive clast-supported conglomerates (up to 20 m thick). Erosional surfaces and cross-stratifications are observed locally. Subordinate lenses of sandstones, 1–5 m thick, may occur (Table 1; Fig. 2). The conglomerate clasts are rounded to subrounded in shape and of Himalayan crystalline and Siwalik origin. The yellowish matrix is coarse and micaceous.

DEPOSITIONAL ENVIRONMENTS

The KFA1–KFA5 facies associations are very similar to the FA1–FA5 facies associations defined by Nakayama and Ulak (1999) and Ulak and Nakayama (2001), who studied the Surai Khola, Tinau Khola and Bakiya Khola sections (Siwaliks in west and central Nepal). In addition, KFA7 is similar to the FA6 in the Tinau Khola and FA7 in the Surai Khola and Bakiya Khola. However, the KFA6 facies association that was observed along the Karnali section in the present study has only been recognized along the Surai section.

The facies associations FA1–FA6 identified by Ulak and Nakayama (2001) correspond, respectively, to a meandering system, a flood-flow-dominated meandering system, a flood-flow-dominated shallow sandy meandering system, a deep sandy braided system, a shallow sandy braided system and a gravely braided system. Using the analogy between the facies observed in the present study and those observed in Ulak and Nakayama's (2001) study, as well as further detailed analysis, the following evolution for the Karnali River system is proposed (Fig. 2): 1. From 15.8 to 13.1 Ma (KFA1), the sedimentation resulted from a fine-grained meandering river system. Frequent paleosols developed, which attest to long-term exposure of the flood plain. 2. From 13.1 to 12.6 Ma (KFA2), the meandering system remained quite stable, with the occurrence of flood flows as suggested by the amalgamated fine-grained sandstone. 3. From 12.6 to 9.5 Ma (KFA3), deposits led to a meandering system with increasing discharge. 4. From 9.5 to 6.4 Ma (KFA4), sedimentation resulted from a sandy braided system. Starting from approximately 7.5 Ma, several conglomeratic beds truncating the underlying deposits and without any sorting occur. They could correspond to erosional and immediately succeeding depositional episodes such as flooding events. This set of flooding events indicates an irregular stream discharge. 5. Between 6.4 and 5.5 Ma (KFA5), the sandy braided river system persisted with discontinuous sedimentation and frequent and episodic high energy discharge events. 6. From 5.5 to 4.0–3.7 Ma (KFA6), the difference in sandstone composition and increase of the relative proportion of the overbank deposits suggest an intermediate period when the drainage system was reorganized and changed to a river fed by piedmont streams, providing a more proximal supply. 7. From approximately 3.9–3.7 Ma (KFA7), a gravely braided river system developed.

EVOLUTION OF STREAM ENERGY

The form of the basal surface of sandstone beds results from a competition between discharge and sediment load. Although analyzed in 2-D, some of the strongly erosive sandstone bases observed along the Karnali River section could underline fluvial channels (Fig. 2). From the consistent increase in several

variables, such as the depth of the channels, the amount of intraclasts in the basal channels, bedload and the grain size, an increase of energy from KFA1 to KFA7 is noticeable in the fluvial system. For KFA1 and KFA2, the main sandy channels are slightly erosive, suggesting a gentle fluvial system with low discharge. The erosive basal surface of sandstone beds in KFA3 indicates the increasing energy of the stream. In contrast, the major deep channels occurring with the KFA4 facies association display nonerosional basal surfaces, suggesting an increase in sediment load relative to discharge capacity. Erosive pebbly sandstones, several decimetres thick, are recognized as large-magnitude flood interbeds. Their occurrence increases upsection within KFA4. Together with an increasing proportion of intraformational mudclasts, the pebbly sandstones suggest frequent increases in the magnitudes of hydrodynamic parameters, which could be linked to an increase in seasonal discharge. Both KFA5 and KFA7 show a regular increase in sediment load, with low erosive sandstone beds and massive gravels. KFA6 is associated with a higher proportion of mudstone, especially at the beginning of the deposition. The basal erosive channels indicate a decrease in sediment load relative to discharge. Channels and sand bodies have a much smaller size. These conditions could correspond to a lateral shift of the stream in an interfluvial zone and/or a decrease of sediment supply linked to a modification of the catchment basin, as suggested by Gupta (1997) and Kumar *et al.* (1999).

MINERALOGY AND GEOCHEMISTRY

CLAY MINERAL ASSEMBLAGES

Analytical methods

Clay mineral assemblages have been extracted from both coarse-grained and muddy layers. Analysis of the clay mineralogy was performed by X-ray diffraction at the University of Grenoble, using a Philips PW 1120/90 diffractometer. Samples were gently crushed to obtain grains of approximately 2 mm in diameter and were then decarbonated in 0.2 mol/L HCl. The excess acid was removed by repeated centrifuging followed by homogenization. The < 2- μ m fractions were collected by decantation after settling, and oriented aggregates were made on glass slides. The X-ray diffractograms were made using an untreated sample, a glycolated sample and a sample heated for 2 h at 490°C. The method of semiquantitative estimation is based on the peak heights of clay mineral peak, assuming that these weighted amounts added up to 100% (Capet *et al.* 1990). Final data are given in percent, the relative error being approximately $\pm 5\%$ (Holtzappel 1985). The values of illite crystallinity (KI) correspond to the full width at half maximum of the 10 Å peak, and this measurement (in $^{\circ}2\theta$) is made on the X-ray diffractogram obtained from the glycolated sample (Kübler & Goy-Eggenberger 2001). The KI measurement is a statistical means to determine the low metamorphic grade in sedimentary terrains. With increasing burial and very low-grade metamorphism, smectites, illite-smectite mixed-layers and illites (re-)crystallize and progressively transform into illites and finally into muscovite. These transformations, very complex in detail, lead to a progressive decrease in the width of the 10 Å peak as measured on an X-ray diffractogram. Therefore, low KI values correspond to well-crystallized minerals, whereas high KI values indicate poorly crystallized minerals. The KI measurements allow us to establish three subdivisions within a very low-grade metamorphism: diagenesis, anchizone and epizone. With the Philips PW 1120/90 diffractometer (Cu-K α radiation) used in Grenoble, the diagenesis-anchizone boundary is defined at 0.42 ($^{\circ}2\theta$, Cu- α) and the anchizone-epizone limit is defined at 0.25 ($^{\circ}2\theta$, Cu- α).

Results

The < 2- μ m fraction of the Siwalik Formations along the Karnali River shows a clayey series with illite, chlorite, kaolinite, smectites and mixed layers (illite-smectite/chlorite-smectite). From the mineralogical variations of the clay fraction, three main periods can be distinguished (Fig. 3a): (i) period 1 – sediments older than 9.5 Ma are dominated by illite- and chlorite-rich assemblages (more than 50% of (Ill + Chl)/ \odot clays); (ii) period 2 – sediments between 9.5 and 6.5 Ma old show rapid alternations of smectite-kaolinite- and illite-/chlorite-rich banks (KFA4 facies association); and (iii) period 3 – sediments younger than 6.5 Ma are dominated by smectite- and kaolinite-rich assemblages (KFA5–KFA7). The KI values vary from 0.14 to 0.52 ($^{\circ}2\theta$) – that is, from well-crystallized minerals, with crystallinity similar to that of the muscovites from the High Himalaya, to poorly crystallized minerals (Fig. 3b). From the base of the section (approximately 5300 m) to 3200 m (ca 9.5 Ma), KI values vary between 0.26 and 0.52 ($^{\circ}2\theta$), indicating conditions of anchizone or diagenesis. From 3200 m to approximately 2100 m (ca 6.5 Ma), KI values range between 0.14 and 0.46 ($^{\circ}2\theta$), suggesting conditions varying from epizone to diagenesis domains. Finally, from approximately 2100 m to the

upper part of the section, illite crystallinity improves, with KI values ranging from 0.30 to 0.14, suggesting a domain of epizone. The observed crystallinity deterioration from a depth of approximately 2100 m (samples older than 6.5 Ma) corresponds to an increasing proportion of illite–smectite-mixed layers (Fig. 4), which are linked to the illitization of smectites as a response to increasing temperature with burying (Lanson *et al.* 1995). Such a transformation begins at approximately 70–95°C (Dunoyer de Segonzac 1970), a domain of temperature that was reached at a depth of approximately 1800–2200 m in the Siwalik Foreland Basin, as determined by Robert (2005) from apatite fission tracks. However, if this burying effect is taken into account, the three periods of different dominant clay assemblages defined above are still discernible. In the case of samples that were buried to less than 2100 m (younger than 6.5 Ma), the proportion of illite–smectite-mixed layers is independent of KI values that fall in the domain of epizone (Fig. 4). Finally, the above-mentioned KI characteristics suggest that the clay fraction of the Siwalik Formations mainly contains ‘inherited illites’ that were affected by epizone-type metamorphism. From a depth of approximately 2100 m, illite particles consist of a mixture of inherited illites and illites derived from smectite–illite transformation.

EVOLUTION OF THE NEODYMIUM ISOTOPE COMPOSITION

Analytical methods

Neodymium (Nd) isotope compositions, determined on whole rocks after classical acid digestion and liquid chromatographic separation, were measured on a Finnigan-MAT 262 in static mode (Galy *et al.* 1996; and references therein). Nd isotope ratios were normalized to $^{146}\text{Nd}/^{144}\text{Nd} = 0.7219$; the $^{143}\text{Nd}/^{144}\text{Nd}$ ratio of the La Jolla standard averaged to $0.511\,854 \pm 0.000\,016$ (1σ , $n = 47$). For samples or standards, the typical 2σ precision for a single run was 0.000008. Blank contributions are negligible for the amount of analyzed elements.

Results

Neodymium isotopes are useful for distinguishing rocks of the Lesser Himalaya from the Higher Himalaya (France-Lanord *et al.* 1993; Galy *et al.* 1996; Robinson *et al.* 2001) in central and western Nepal. They allow us to determine the main source area of the detritus deposited in the Siwalik Foreland Basin (Huyghe *et al.* 2001). 1. The $\Sigma_{\text{Nd}}(0)$ values of the sediments of the Karnali section range from 14.6 to 17.9‰ (Fig. 5; Table 2). They are close to those of the modern Higher Himalaya and suggest that the detrital input of the Siwalik sediments was predominantly derived from the erosion of formations analogous to the modern Higher Himalaya (France-Lanord *et al.* 1993; Ahmad *et al.* 2000), although a variable but minor proportion of Lesser Himalayan material was present. Two main points can be outlined: from the Middle Siwaliks (inclusive) to the end of the deposition of Siwalik sediments, $\Sigma_{\text{Nd}}(0)$ values decrease, implying an increase in the proportion of the Lesser Himalayan material with time. Similar results have been obtained in the Surai Khola section (Huyghe *et al.* 2001; Robinson *et al.* 2001). Assuming an average $\Sigma_{\text{Nd}}(0)$ composition of 15.2‰ for the Higher Himalaya and 22.7‰ for the Lesser Himalaya (Robinson *et al.* 2001) and using a mixing equation (Allegre 2005), the contribution of the Lesser Himalaya increased from < 10% for the Lower Siwalik samples to approximately 30–35% for the Middle and Upper Siwalik samples. 2. A high proportion (approximately 30%) of Lesser Himalayan input started abruptly at ca 10 Ma (at the end of facies KFA3) in the Karnali section, indicating local exhumation of Lesser Himalayan rocks in the catchment basin. The exhumation of the Lesser Himalayan rocks has been controlled by the thin-skinned tectonics of the Lesser Himalayan thrust system, and would have begun at 12–10 Ma, taking into account the delay for denudation (Huyghe *et al.* 2001; and references therein).

DISCUSSION

COMPARISON OF THE SEDIMENTARY RECORDS IN VARIOUS SECTIONS OF THE NEPALESE SIWALIKS AND THE BENGAL FAN

Regional evolution of the Neogene fluvial system –Bakiya, Tinau, Surai and Karnali sections

Comparison of the facies analysis performed along the Karnali River section with similar studies by Nakayama and Ulak (1999) and Ulak and Nakayama (2001) along the Surai, Tinau and Bakiya sections (Fig. 1a), shows similar changes in the fluvial system during the Upper Miocene in the Siwalik Basin from central to western Nepal (Fig. 6). Magnetostratigraphic ages are based on the studies of Appel *et al.* (1991), Gautam and Appel (1994) and Harrison *et al.* (1993), respectively, all of them updated to the geomagnetic polarity

timescale of Cande and Kent (1995) by Gautam and Rösler (1999). Fluvial changes are diachronous over a distance of approximately 500 km, with age gap varying from 1.1 to 4.5 Ma, which seems common for fluvial system dynamics in the Himalayan Foreland Basin (Brozovic & Burbank 2000). The fluvial changes are in general older along the Karnali River section (western Nepal) than along the other sections; this age difference could probably be a result of the larger size of the Karnali Catchment Basin. Owing to the large size, the Karnali depobasin was probably more sensitive to variations in discharge, sediment supply (larger surfaces affected by erosion and a larger variety of rocks) and tectonics from both the High Himalaya and Lesser Himalaya domains. The minimum age gap occurs for the change from sandy to shallow sandy braided system (KFA4–KFA5), which occurs between 6.6 Ma (along the Bakiya section) and 5.5 Ma (along the Tinau section). Keeping in mind the uncertainties of magnetostratigraphy, the 1.1-Ma age gap is maybe not meaningful and the change from sandy to shallow sandy braided system could be considered synchronous over the Siwalik Basin compared with the other changes in fluvial style.

Regional evolution of the Nd isotope composition and clay mineralogy – the Bengal Fan

The observed trends of the $\Sigma_{Nd}(0)$ in the Siwaliks are consistent with the slight decrease of $\Sigma_{Nd}(0)$ compositions of Bengal Fan sediments (since 8 Ma), where a comparable range of compositions was detected (France-Lanord *et al.* 1993; Galy *et al.* 1996). A Lesser Himalayan contribution is also inferred from radiogenic Os isotopic sediment in both Bengal Fan and Siwalik constituents (Pierson-Wickmann *et al.* 2000). The unique source of such radiogenic Os in Himalaya is the black shale of the Upper Lesser Himalaya (Pierson-Wickmann *et al.* 2000), implying that the Lesser Himalaya was locally exposed to erosion since at least *ca* 12 Ma. Results similar to those of the Karnali section have been obtained from the clay mineral assemblages of Ocean Drilling Project Leg 116 and Deep Sea Drilling Project Leg 22 (Bouquillon *et al.* 1990; Derry & France-Lanord 1997; Galy 2003), where the occurrences of smectite and kaolinite have been linked to variations in environmental conditions at approximately 7 Ma.

ORIGIN OF OBSERVED MAIN CHANGES

The sedimentological data described above show that two kinds of changes in the fluvial system are contemporaneous with major changes in the sediment characteristics within the Siwalik Group. First, a major change from a meandering river system to a braided river system occurred at *ca* 9.5 Ma between KFA3 and KFA4. This change followed an abrupt increase in the Lesser Himalayan detrital input inferred from Nd isotopes. Second, the change from a deep braided system to a shallow braided system occurred at *ca* 6.5 Ma between KFA4 and KFA5. It is contemporaneous with an increase of the smectite/kaolinite proportions of the clay mineralogy. A scenario for the tectonic evolution of far western Nepal is proposed from changes in the fluvial style along the Karnali River, clay mineralogy and Nd isotope geochemistry data. The controls of these fluvial changes are then discussed.

A tectonic scenario for the Himalaya of far western Nepal

The gradual decrease of the neodymium isotope values recorded between *ca* 13.0 and *ca* 9.5 Ma in the deposits of the Karnali River points to an increased input in the Lesser Himalayan rock material, an event that was attributed to initiation of the Lesser Himalayan tectonics (Huyghe *et al.* 2001). This hypothesis is supported by the occurrence of zircons of paleoproterozoic age from the Lesser Himalaya, and Cambrian–Ordovician age derived from the Dadeldhura granite, in the Middle Siwalik sandstones (DeCelles *et al.* 1998, 2001). In detail, it is proposed here that the propagation of the basal detachment beneath the Lesser Himalaya affected different thrust sheets in the following succession: the hanging wall of the Dadeldhura thrust, the Ramgarh thrust (RT) and the main boundary thrust (MBT) (Fig. 7). Such a scenario of propagation has already been proposed by DeCelles *et al.* (2001). It also agrees with the significant grain-size increase from Middle Siwalik time to the present day (Huyghe *et al.* 2001). Here, a more detailed description is provided by taking into account the present results along the Karnali section and by using the following rule based on the example of the recent uplift pattern in central Nepal (Lavé & Avouac 2001): the motion along the active ramps of the basal detachment induces zones of rapid and localized uplift (Fig. 7) and the proportion of High Himalayan and Lesser Himalayan rocks exposed in these zones of intense erosion controls the supply to the foreland basin and the mélange of source recorded in the Siwalik deposits of the Karnali River section. The main differences between this scenario and the interpretation of DeCelles *et al.* (2001) are given below: 1. The Dadeldhura sheet did not overly the whole of the Ramgarh sheet. Therefore, the activity of RT led to the erosion of Lesser Himalayan rocks as soon as the foreland basin cover was removed. This tectonic activity controlled the decreasing trend of the neodymium isotope values between *ca* 13.0 and *ca* 9.5 Ma in the deposits of the

Karnali River (Fig. 5); 2. Main boundary thrust initiation occurred at approximately 10–9 Ma and was synchronous with the RT motion. This early and weak activity of the MBT induced erosion of the syn-orogenic sediments located at its hanging wall, the sources of which were mainly derived from the High Himalaya. The continuous RT activity led to the consumption of most of the Lesser Himalayan rocks located at its hanging wall prior to 9.5 Ma, and caused erosion of the Dadeldhura sheet after 9.5 Ma. Therefore, the neodymium isotope values recorded at 10.0–9.5 Ma in the foreland basin increased as a result of the activities of both the MBT and RT and also the supply of eroded material mainly from the High Himalayan domain or with affinities close to those of the High Himalaya. The 10–9.5 Ma increase of neodymium isotope values was also synchronous with reactivation of the MCT, evidenced by MacFarlane 1993) in central Nepal. However, the 10–9-Ma period is marked by the maximum sediment accumulation rate (Gautam & Fujiwara 2000), a maximum that fits with an increase of the subsidence rate because of the rapid southward migration of the thrust load above the flexed lithosphere rather than a reactivation of the MCT, which would increase the distance between the thrust load and the foreland. Therefore, such a reactivation is not appropriate.

Tectonic control for the change from meandering to braided river system

From the large content of feldspars and biotites in the ‘salt and pepper’ sandstones of the Middle Siwaliks, Ulak and Nakayama (2001) argued that the onset of the braided river system was caused by rapid uplift of the Higher Himalayan terrain. However, thermochronologic detrital investigations suggest a steady-state erosion (Bernet *et al.* 2005) or a decrease (Szulc *et al.* 2004) of the exhumation by that time (10–8 Ma). Therefore our tectonic scenario proposes the following interpretation for the sedimentological evolution pointed out by Ulak and Nakayama (2001): the development of the duplex and the motion along the midcrustal ramp of the RT induced passive uplift and erosion of the High Himalaya and Dadeldhura thrust sheets, whereas the weak activity of the MBT induced erosion of syn-orogenic sediments and of a few Lesser Himalayan rocks; the initiation of the MBT decreased abruptly the distance from the Karnali section to the topographic front and induced, by increasing the river gradient, the transport of coarser sediment load, as revealed by grain-size analysis (Huyghe *et al.* 2001), and the development of braided river systems, as shown by the KFA3/KFA4 transition of the Karnali section. Following the tectonic sketch proposed above, the evolution from illite-/chlorite-rich assemblages prior to 9.5 Ma to a rapid alternation of illite/chlorite and smectite/kaolinite banks could be linked to a change of source material. Recycling of the early foreland basin sediments in the hanging wall of the MBT provided a source of rocks and minerals that consequently suffered two consecutive weathering episodes, then susceptible to produce more clay minerals linked to weathering (Fig. 7c). Some basic intrusives at the base of the Precambrian low-grade metasediments in the Ramgarh thrust sheet could also have been weathered and produced clays with smectites in the Middle Siwaliks. However, the proportion of basic intrusives is minor (Shrestha 1987) and other origins have to be proposed. For example, the formation of smectites from feldspar weathering has already been described in the Siwaliks when fluid circulation is involved (Huyghe *et al.* 1998). The porosity of the coarse and weakly cemented Middle Siwalik sandstones and the feldspar supply of up to 20% from 11 to 9 Ma (Tokuoka *et al.* 1986; Hisatomi 1990; DeCelles *et al.* 1998) are favorable to such a transformation. In such a case, smectites would form at shallow levels in the foreland basin and would result from an indirect effect of the change of source.

Climatic control for the change from deep sandy to shallow sandy braided river system

The change from deep sandy to shallow sandy braided river system observed at 6.5 Ma along the Karnali River section is characterized by discontinuous sedimentation and frequent and episodic high-energy discharge events. It was preceded by evidence of irregular stream discharge producing large-magnitude flood interbeds. Altogether, these characteristics could be linked to an increase in seasonal discharge. From the tectonic interpretation in the present study (Fig. 7), no major tectonic event occurred between approximately 8 and 5 Ma and the distance between the frontal relief of the Himalaya and the Karnali section remained approximately 50 km. This suggests that the change in fluvial style at 6.5 Ma could be of climatic origin. Moreover, the regional correlation of fluvial styles within the Siwalik Basin (Fig. 6) indicates that it is relatively synchronous over 500 km. Such characteristics also favor a climatic origin. From 7 to 5 Ma, an increase of seasonality has been identified in the Siwalik Basin from C and O isotopic analyses of pedogenic carbonates and mollusk shells (Harrison *et al.* 1993; Quade *et al.* 1995; Gajurel *et al.* 2003; Sanyal *et al.* 2004), and from paleovegetation analysis (Awasthi & Prasad 1990). This increase of seasonality has been attributed by these authors to a variation in monsoon intensity. From a mineralogical point of view, this change

in fluvial style occurs during the period of smectite-/kaolinite-dominant clay assemblages. Nd isotopes (Fig. 5) do not indicate any abrupt source rock change and no consequent source rock of smectites is noted from 6.5 Ma onward in the Karnali Catchment Basin. As indicated above, the minor surface exposure of basic intrusive is also not enough to explain the long period of smectite abundance in the deposits. As the observed feldspar content in these deposits is lower for the period from 6.5 Ma to the present than during 9.5–6.5 Ma, the formation of higher amounts of smectites as a result of feldspar weathering after 6.5 Ma is not plausible. All these arguments suggest that the abundance of smectites in the clay assemblages from 6.5 Ma onward is linked to strong seasonality that favored the formation of smectites from pedogenetic processes. The pedogenesis occurred in areas characterized by soils preserved from fast physical erosion (i.e. within the zones of weak uplift rate in the Lesser Himalaya and the foreland basin). This supply of smectite-/kaolinite-rich clay assemblages is added to the supply of illite-/chlorite-rich clay assemblages provided by the zones of high uplift, dominated by physical erosion in the High Himalaya.

CONCLUSIONS

The present multidisciplinary study comprising analysis of facies, clay minerals and Nd isotopes, conducted along the Karnali River section, permitted us to distinguish several changes in the fluvial system. Six classes of fluvial facies associations among the seven recognized along the Karnali River section have been described previously from three Siwalik sections, from western to central Nepal, stretched over a distance of approximately 500 km. It is notable, however, that the facies boundaries in these sections are diachronous, except in the case of the change from a deep sandy braided river system to a shallow sandy braided system where the time gap is only approximately 1 my. The present study reveals that the difference between fluvial changes induced by tectonics and fluvial changes induced by climatic variations is very small. Difficulty in separating tectonic and climate events is more pronounced in the case of the Himalayas, where tectonics and climate may interact. However, multidisciplinary studies jointly utilizing source markers, weathering indicators and sedimentological parameters give some clues for deciphering climatic and tectonic phenomena. Of the six changes in fluvial systems distinguished along the Karnali River, it is suggested here that at least one of them is linked to tectonic reorganization of the Himalayan belt and another to a climatic change:

1. A tectonic evolution compatible with the major change from a meandering to a braided river system and with the geochemical variations is proposed. The southward propagation of the basal detachment beneath the Lesser Himalaya from *ca* 13 Ma led to the successive uplift of different thrust sheets, the exhumation and erosion of which controlled the sediment supply to the foreland basin, as shown from the Nd-isotope composition of Siwalik deposits. The activity of the Ramgarh thrust led to erosion of the Lesser Himalayan rocks from *ca* 12 Ma in the catchment basin of the paleo-Karnali river. The southward propagation of the thrust and fold activities within the outer part of the Lesser Himalaya at approximately 10–9 Ma changed the morphology, the proximity of relief and the surface slope. It was also responsible for the high gradient and sediment load shown by the onset of braided river systems in the paleo-Ganga plain at *ca* 9.5 Ma.

2. Another significant change from a deep sandy braided to a shallow sandy braided system took place at *ca* 6.5 Ma. It seems to be of climatic origin as suggested by the discontinuous sedimentation and episodic high-energy discharge and by synchronism of such a change over approximately 500 km. It might be linked to an increase in seasonality also shown by the change from C3 to C4 vegetation and by the abundance of smectites among the clay minerals. Similar results obtained from clay minerals of the distal part of the Bengal Fan attest to the wide impact of this event controlled by the regional evolution climate. The observed indicators ascribed to a climatic change are not totally synchronous along the Karnali River section; the initiation of irregular stream discharge precedes the start of the 6.5 Ma change in the fluvial system by *ca* 1 Ma. This is an indication of a gradual environmental change from approximately 7.5 Ma.

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FIGURES

Fig. 1 Location map of the study area. (a) Structural sketch of the Himalayas with location of the area considered (western Nepal). Sections cited in the text are, from west to east: Karnali River (KR) section, Surai Khola (SK) section, Tinau Khola (TK) section, Bakiya Khola (BK) section. MBT, main boundary thrust; MCT, main central thrust; MFT, main frontal thrust; RT, Ramgarh thrust. (b) Structural and active-fault map of western Nepal (adapted from Upreti & Le Fort 1999). Location of Figure 7 is shown.

Fig. 2 Sedimentary facies recognized along the Karnali River section. The ages of the limits of facies KFA1–KFA6 are reported along the simplified lithologic column established by Gautam and Fujiwara (2000). ms., mudstone (clay and siltstone); ms. > ss., mudstone-dominated; pbl.ss., pebbly sandstone; ss., sandstone; ss. > ms., sandstone-dominated. Photographs with characteristic features of facies association KFA1–KFA7 are shown on the right side. Their locations are shown by red stars along the lithologic column. (a) Sedimentary facies KFA1–KFA4; (b) sedimentary facies KFA4–KFA7.

Fig. 3 Clay mineralogy of the Siwalik deposits along the Karnali River section. (a) Variation of the ratio (illite + chlorite)/(sum of clays: illite + chlorite + kaolinite + smectites + illite–smectite mixed layers) vs time. The limits of facies KFA1–KFA7 are reported from their stratigraphic ages. Values for modern samples are plotted along the upper horizontal axis: suspension material from Ganga River during monsoon in Bangladesh (gray triangle); suspension load from Bheri River during monsoon at approximately 700 m elevation (gray square); modern sands from Karnali River (open circle); rock flour collected at approximately 5050 m elevation (gray star). (b) Variation of illite crystallinity vs depth of the deposits (approximately 5300 m in stratigraphic thickness). Values for modern samples are shown along the upper horizontal axis by gray stars for rock flour collected at approximately 5050 m elevation and by gray rectangles for suspended materials from Bheri River down to the Ganga River during monsoon.

Fig. 4 Variation of the illite crystallinity vs the proportion of illite–smectite mixed layers. (□), Samples younger than 6.5 Ma; (·), samples older than 6.5 Ma.

Fig. 5 Nd-isotope composition $\Sigma_{Nd}(\text{‰})$ of samples from Siwalik sedimentary rocks of the Karnali River based on data from Huyghe *et al.* (2001) and new data (the present study). Along the base, the range of $\Sigma_{Nd}(0)$ values from the higher Himalayan (HH) zone and Lesser Himalayan (LH) zone from Robinson *et al.* (2001; and references therein) are shown with the average $\Sigma_{Nd}(0)$ value marked with a black bar and the standard deviation of the average marked with gray bars. The limits of facies KFA1–KFA7 have been reported from their stratigraphic ages.

Fig. 6 Regional evolution of the Neogene fluvial system in western and central Nepal from comparison of the Siwalik sedimentary facies identified along the Karnali (circles), Surai (squares), Tinau (triangles) and Bakiya (cross marks) sections. For each section, the following facies are found in order of increasing energy, from left to right: fine-grained meandering system, flood-flow meandering system, flood-flow-dominated meandering system, deep sandy braided system, shallow sandy braided system to gravelly braided system. Interpretation of depositional environments is based on Nakayama and Ulak (1999). The horizontal position of the sections is not to scale. (Ulak & Nakayama 2001.)

Fig. 7 Incremental balanced restoration of the Himalayan fold-thrust belt in far western Nepal. DT, Dadeldhura thrust; MBT, main boundary thrust; MCT, main central thrust; MFT, main frontal thrust; RT, Ramgarh thrust; 1, foreland basin; 2, Autochthonous Indian craton; 3, Lesser Himalayan thrust sheet; 4, Ramgarh thrust sheet; 5, duplex of the Lesser Himalaya; 6, Dadeldhura thrust sheet; 7, High Himalaya. Rocks of thrust sheets 2, 3, 4 and 5 have $\Sigma_{Nd}(0)$ values close to those of the Lesser Himalayan material as given in Figure 5; rocks of thrust sheets 6 and 7 have $\Sigma_{Nd}(0)$ values close to those of the High Himalayan material as given in Figure 5. The final geometry is adapted from the balanced cross-section of DeCelles *et al.* (2001) using the sequence of thrusting deduced from the present work. The heavy full lines indicate the active thrusts, and the vertical arrows refer to the zones of high uplift at the hanging wall of the ramps. The negative values above the horizontal arrow on the right side correspond to the amount of shortening; the italic values on the left side indicate the distance between the position of the present-day Karnali

section and the frontal relief of the Himalaya. (a) Geometry at ca 15 Ma – the Dadeldhura thrust sheet overthrust the Lesser Himalaya prior to 13 Ma, tectonics that led to erosion of rocks with High Himalayan affinities. (b) Geometry at ca 11 Ma – the emplacement of the Ramgarh Thrust between 13 and 10 Ma led to the erosion of the Lesser Himalayan rocks ahead of the Dadeldhura thrust sheet. (c) Geometry at ca 8 Ma – the duplexing of the lower part of the Lesser Himalaya after 10 Ma uplifted the Dadeldhura sheet and the High Himalaya, and eroded rocks with $\Sigma_{Nd}(0)$ values close to those of High Himalayan material. The simultaneous propagation of the detachment thrust to the MBT decreased the distance between the present-day Karnali section and the frontal relief and led to recycling of sediments of the pre-Siwalik Foreland Basin that had $\Sigma_{Nd}(0)$ values close to those of the High Himalayan rocks. (d) Geometry at ca 5 Ma – the main phase of growth of the Upper Lesser Himalayan duplex led to erosion of Lesser Himalayan and High Himalayan rocks. (e) Present-day (0 Ma) geometry – the uplift above the mid-crustal ramp leads to erosion of the Lesser Himalayan and High Himalayan rocks.

Table 1 Facies associations recognized in the Siwalik series along the Karnali River section

Magnetostratigraphic sites †	Age (Ma)	Lithofacies and sedimentary structures ‡	Facies association	Architectural elements and form of the basal surface of main sandstone beds §	Interpretation ¶
3980 +/- 30 m	~3.9	Gm, Gt – St – (Fm) PBL>SS	KFA7 Upper Siw.	CH, GB, DA	Gravelly braided fluvial system
3980 +/- 30 m	~3.9	Gp, Gt – Sp, St, Cv – Fm, Mb SS>PBL	KFA6 Middle Siw.	CH, GB, DA, OF Base erosive	Anastomosed system
G10	5.5-5.6				
G10	5.5-5.6	Sp, St, Gt, Se, (Sl) - Sr – Fm, (P) “Salt & pepper” SS	KFA5 Middle Siw.	CH, SB, DA, (LA), OF	Shallow sandy braided fluvial system
E20	6.4				
E20	6.4	Sp, Se, St – (Sh), Sr – Fm, P SS>SL+CL	KFA4 Middle Siw.	CH, LA, SB, DA, OF Base slightly erosive	Deep sandy braided fluvial system
C43	9.5	“Salt & pepper” SS	Siw.		
C43	9.5	St, Sp, (Sl), Se, Ss – Fm, Fl, Fsc – P, Fr SS>SL+CL	KFA3 Lower Siw.	CH, SB, DA, (LA), amalg, LS, OF Base erosive	Sandy flood flow- dominated meandering fluvial system
A158	12.6-12.7	“Salt & pepper” SS	Siw.		
A158	12.6-12.7	Sp, Sl, Se – Fm, Fl - (Fr, P) SL>SS	KFA2 Lower Siw.	CH, SB, LA, (amalg), OF Base tabular and slightly erosive	flood flow-dominated meandering fluvial system
A122	13.1				
A122	13.1	St, Sp, (Sl) – Fm, Fl, Fsm, (Sh) – P, Fr	KFA1 Lower Siw.	CH, SB, LA, OF Base slightly erosive	Fine-grained meandering fluvial system with low- discharge and relief
A1	15.8	SL+ CL >SS	Siw.		

† from Gautam and Fujiwara (2000)

‡Fm: mud, silt (massive), Fsc: silt, mud (massive), Gm: massive or crudely bedded gravels, Gt: gravels with trough crossbeds, Gp: gravels with planar crossbeds, Cv : Convolute, Sr: sand, very fine to coarse with ripple marks, Sh: sand very fine to very coarse with horizontal laminations, Ss: sand, fine to coarse, with broad, shallow scours, St: sand, medium to very coarse with trough crossbeds, Sp: sand, medium to very coarse with planar crossbeds, Se: erosional scours with intraclasts, Fl: fine laminations and very small ripples within sand, silt and mud, Mb: bioturbated mud, P: carbonate pedogenic features, Fr: rootlets. CL: mudstone, SL: siltstone, SS: sandstone, PBL: conglomerate

§ CH: channel element, DA: downstream accretion, GB: gravel bar, LA: lateral accretion, LS: laminar sand sheet, OF: overbank fines, SB: sandy bedforms, amalg: amalgamated sandstone

Brackets indicate that lithofacies, sedimentary structures or architectural elements have been observed in the field but are of minor importance.

¶ from Nakayama and Ulak (1999)

Table 2 Sr-Nd isotopic compositions of Nepalese Siwaliks samples

Sample #	Age (Ma)	Sr (ppm)	$^{87}\text{Sr}/^{86}\text{Sr}$	Nd (ppm)	Sm (ppm)	$^{143}\text{Nd}/^{144}\text{Nd}$	$^{147}\text{Sm}/^{144}\text{Nd}$	$\epsilon_{\text{Nd}}(0)$ (‰)
† ka 25	1.4	N.D.	0.75207	N.D.	N.D.	0.511753	N.D.	-17.3
† ka 27	3.4	N.D.	0.73252	N.D.	N.D.	0.511803	N.D.	-16.3
† ka 29	4.1	N.D.	0.73034	N.D.	N.D.	0.511755	N.D.	-17.2
‡ Kar 13	5.2	113.1	0.748910	33.95	2.76	0.511829	0.118	-17.6
† ka 34	6.4	N.D.	N.D.	17.3	3.46	0.511831	0.127	-15.8
† ka 35	7.15	N.D.	N.D.	37.7	7.30	0.511846	0.123	-15.5
† ka 36	7.5	N.D.	N.D.	33.7	6.35	0.511840	0.120	-15.6
† ka 37	8.1	N.D.	N.D.	29.8	5.74	0.511808	0.122	-16.2
† ka 38	9.1	N.D.	N.D.	31.7	6.01	0.511830	0.120	-15.8
† ka 39	10	N.D.	N.D.	33.4	6.43	0.511720	0.122	-17.9
† ka 40	10.4	N.D.	N.D.	38.1	7.36	0.511736	0.123	-17.6
† ka 41	11.5	N.D.	N.D.	15.0	3.38	0.511831	0.143	-15.8
† ka 42	12.35	N.D.	N.D.	N.D.	N.D.	0.511891	N.D.	-14.6
‡ Kar 3	14.1	80.9	0.741119	33.95	6.38	0.511829	0.114	-15.8
‡ Kar 1	15.9	49.1	0.753261	14.06	2.76	0.511865	0.119	-15.1

† : Karnali River section (Huyghe *et al.* 2001) ; ‡ : Karnali River section (this study)

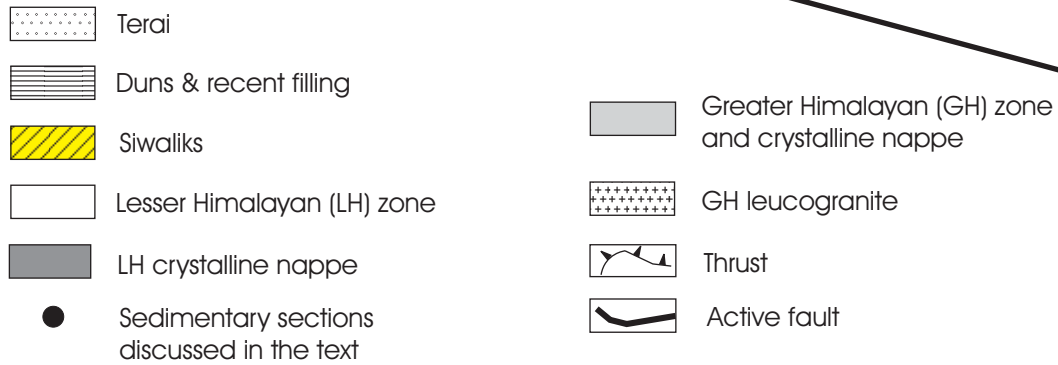
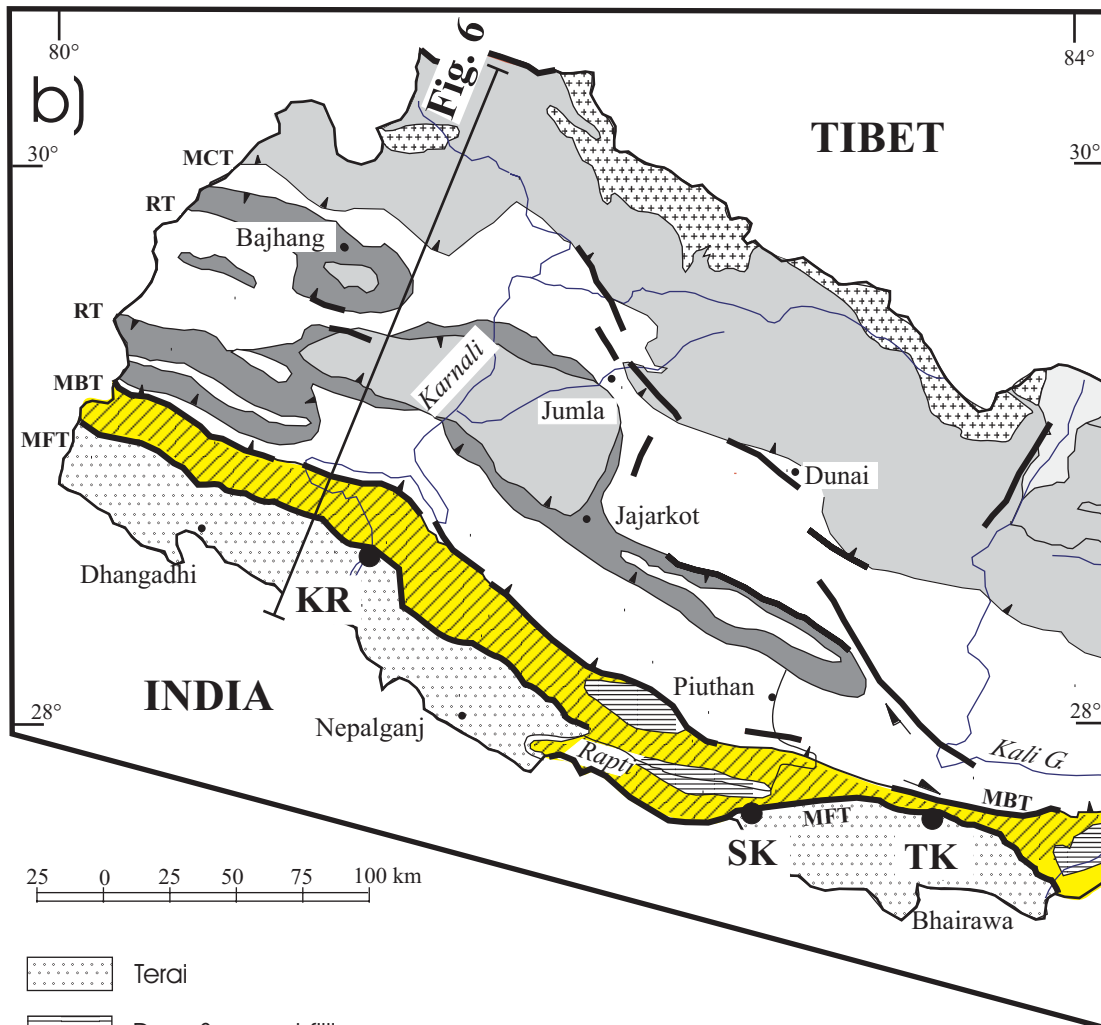
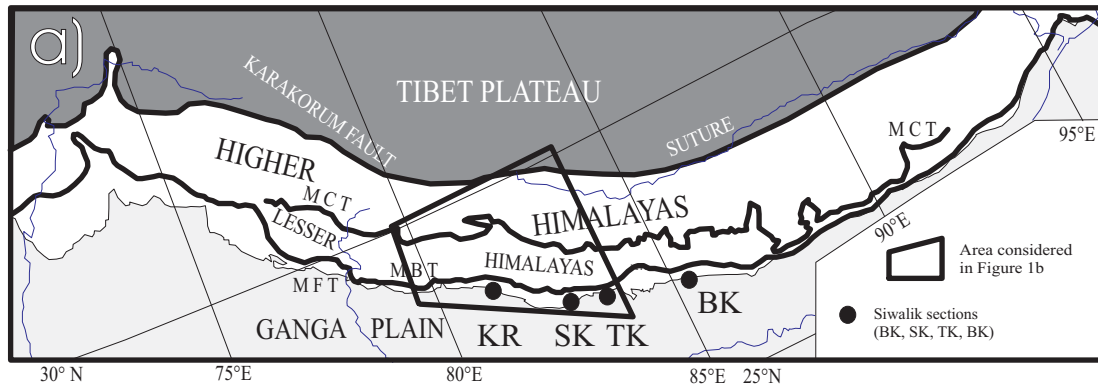
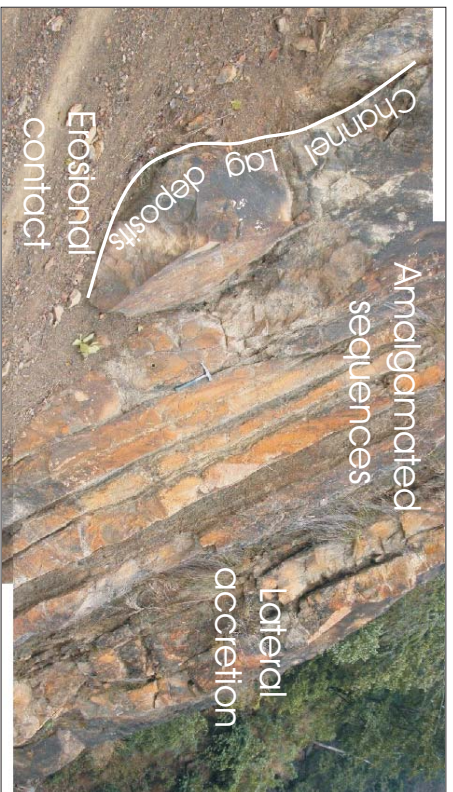
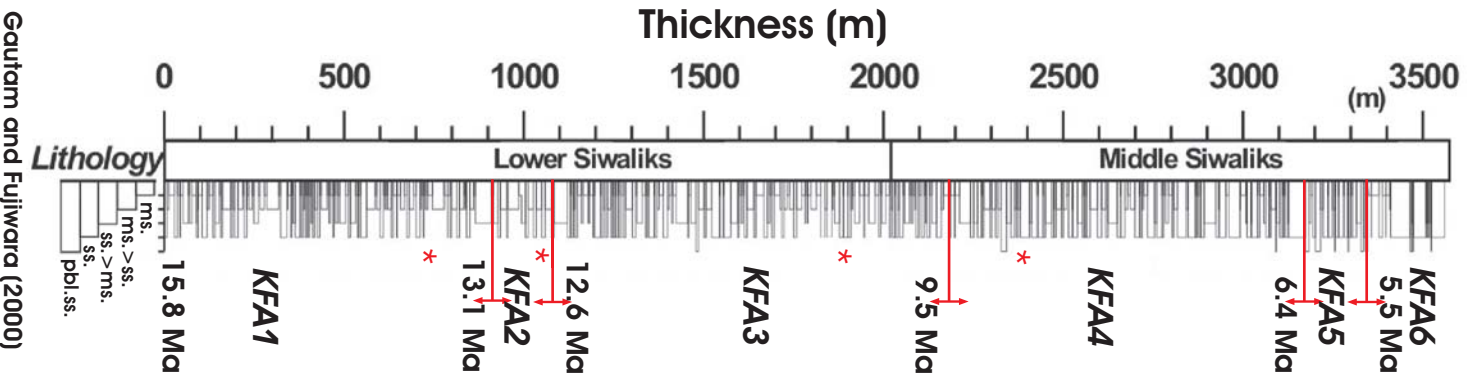
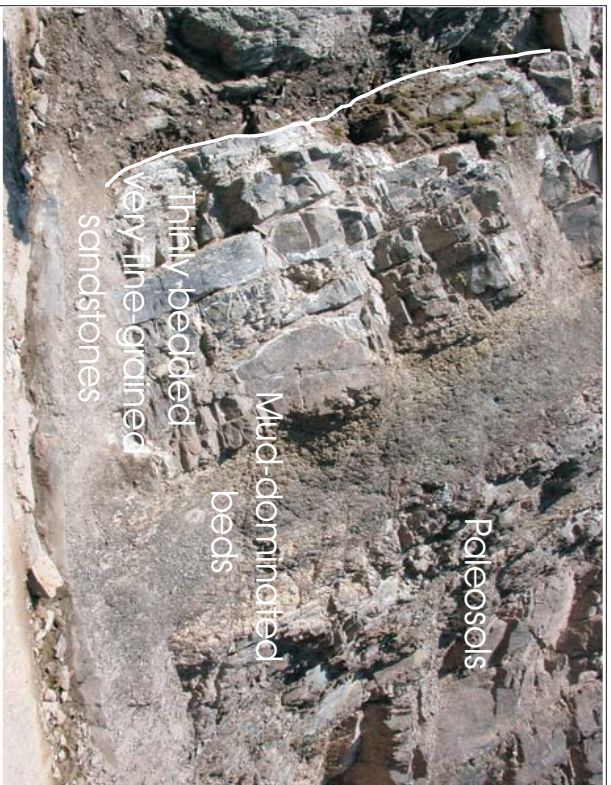


Figure 1



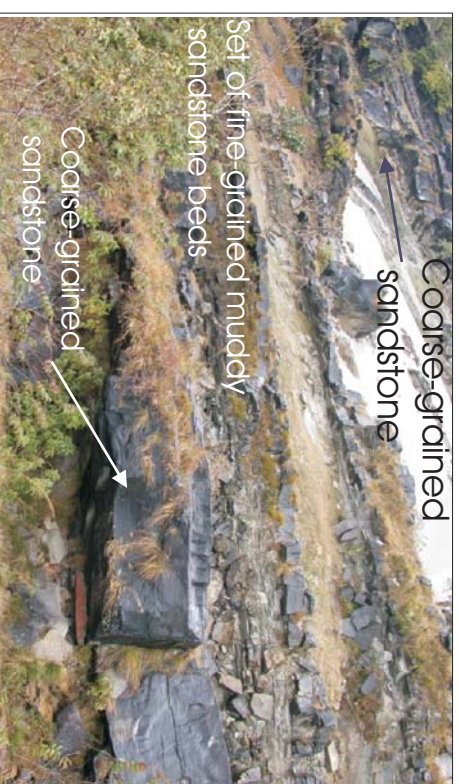
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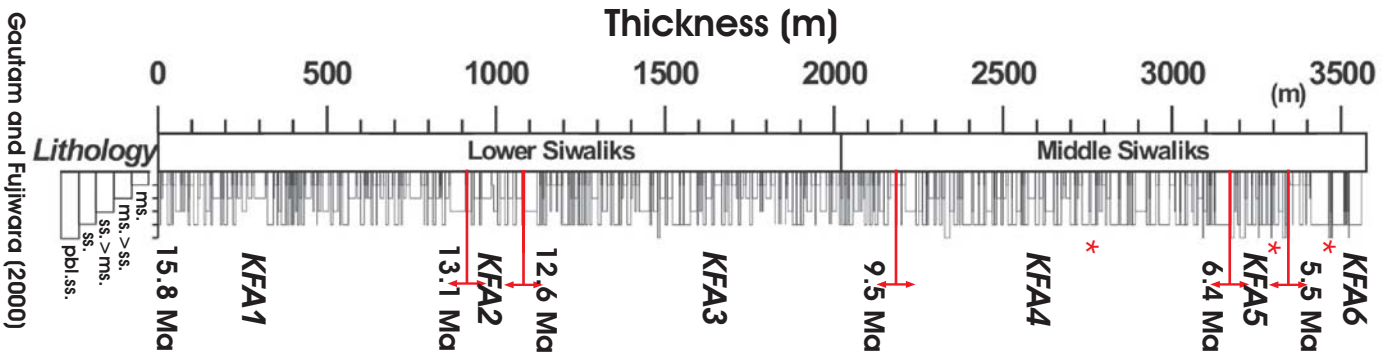


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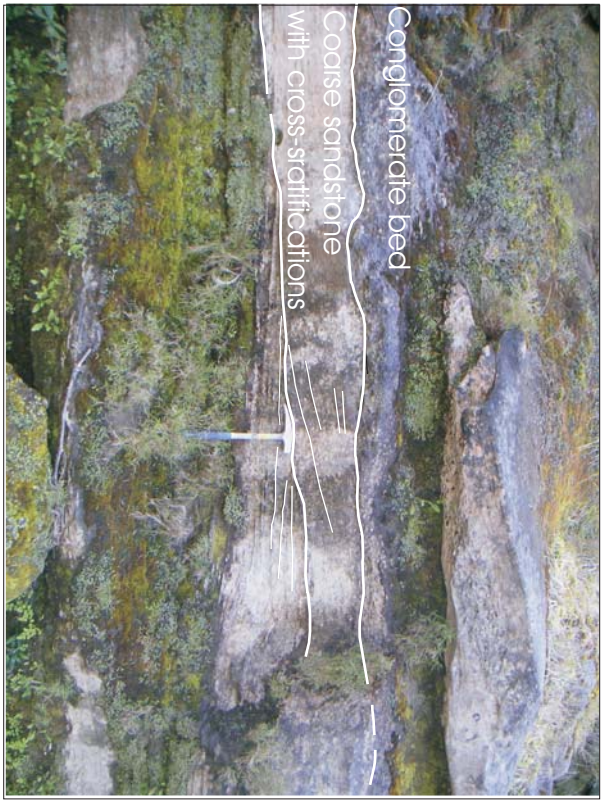
Figure 2a



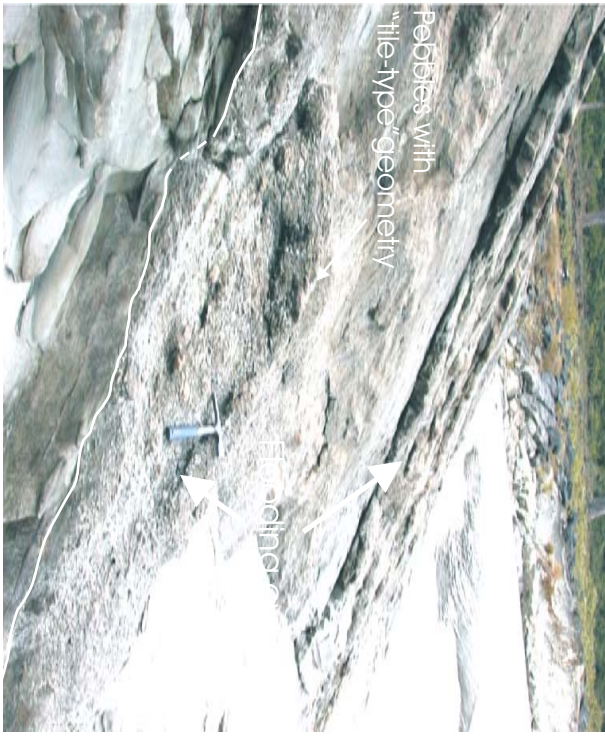
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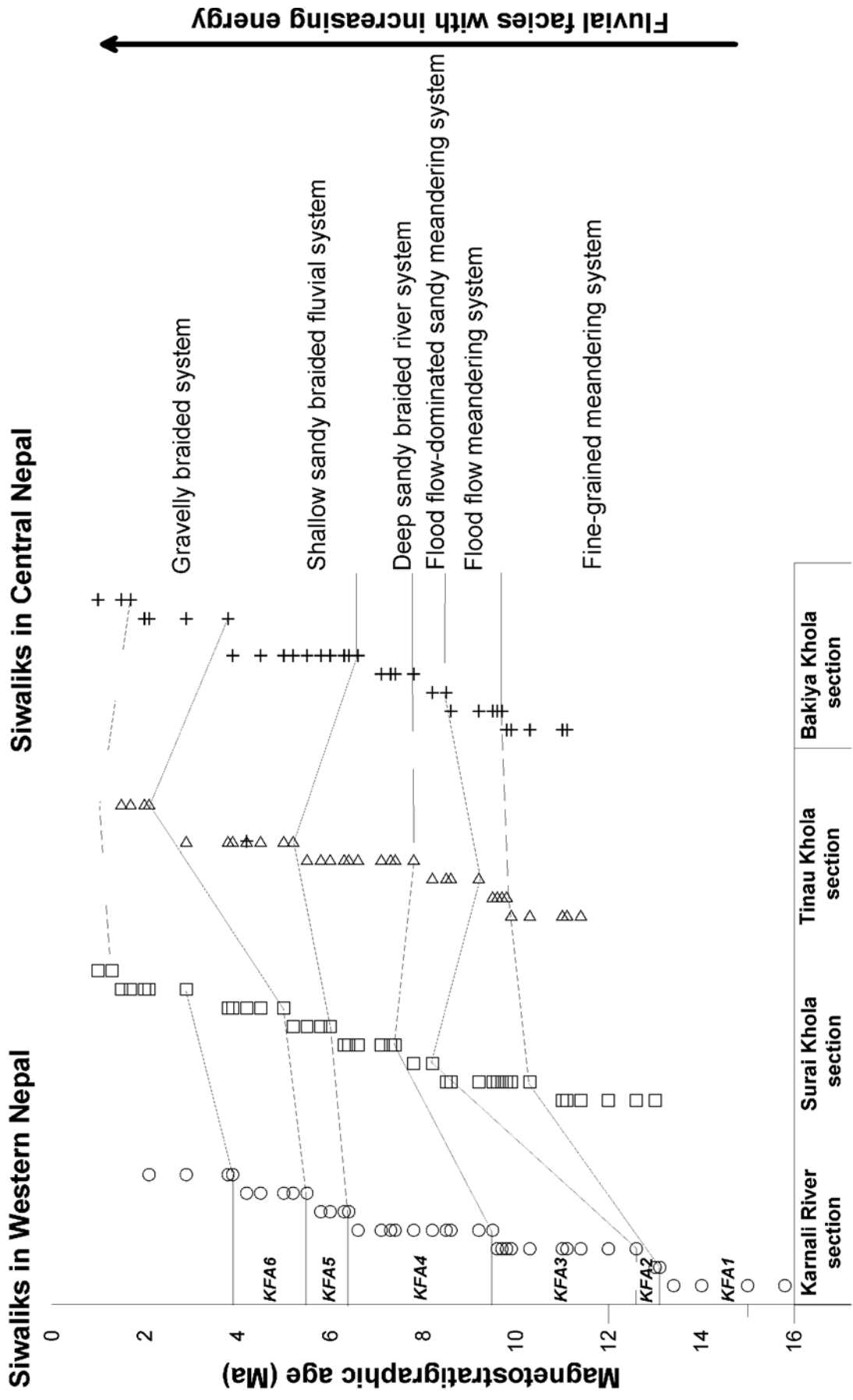


Facies KFA6



Facies KFA4

Figure 2b



This study
 Nakayama & Ulak (1999)
 Ulak & Nakayama (2001)
 Nakayama & Ulak (1999)

Figure 6

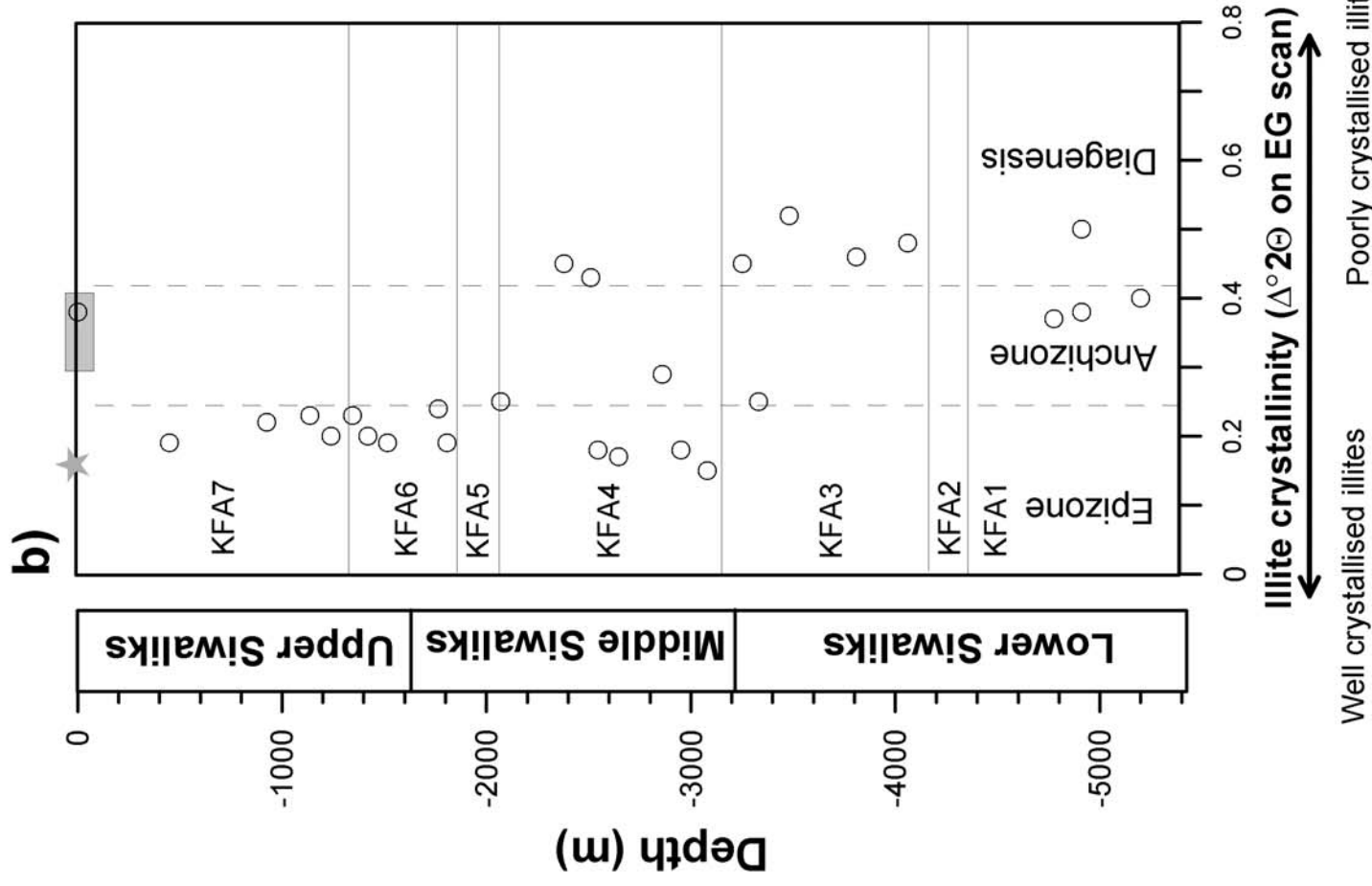
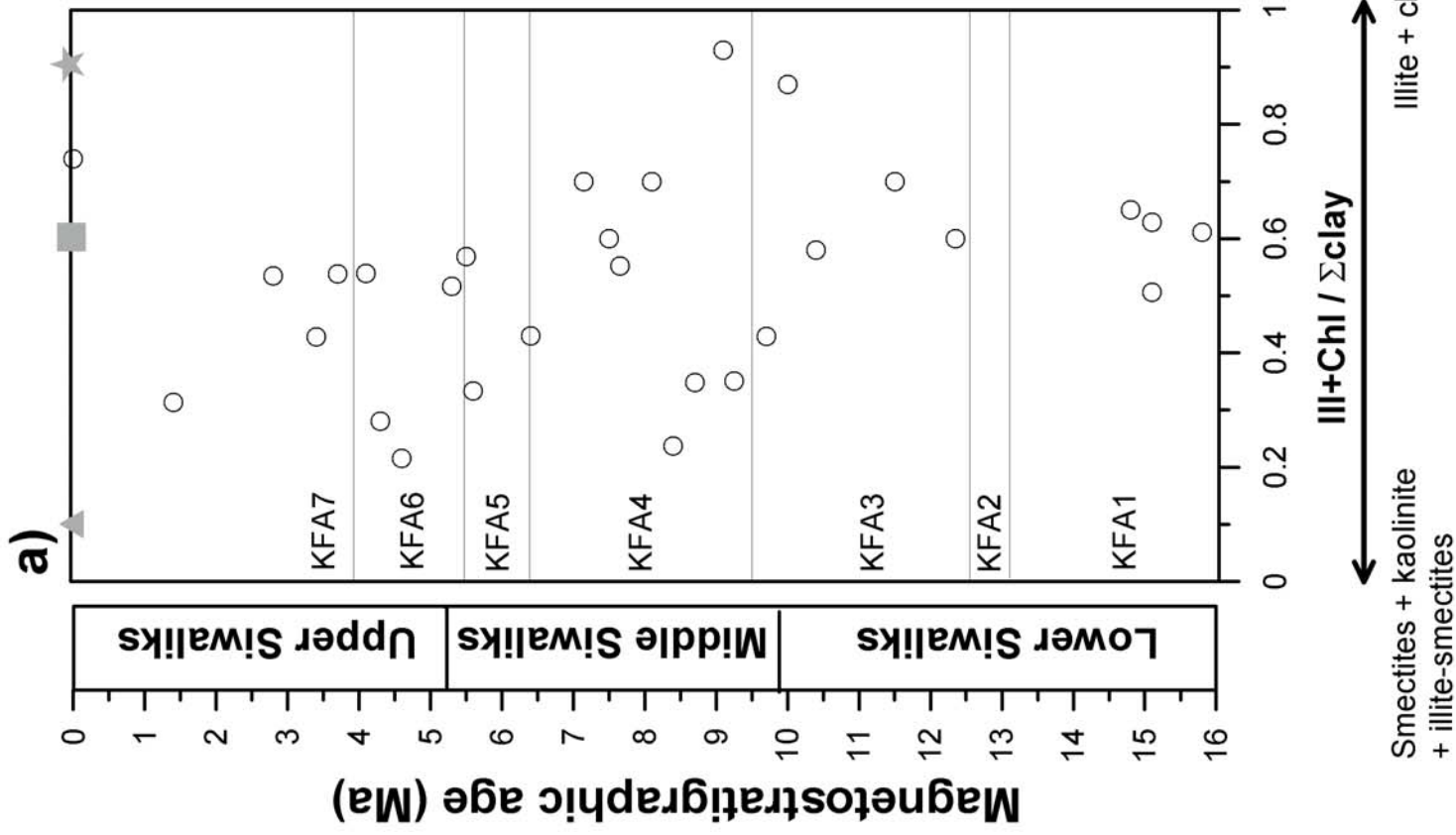


Figure 3

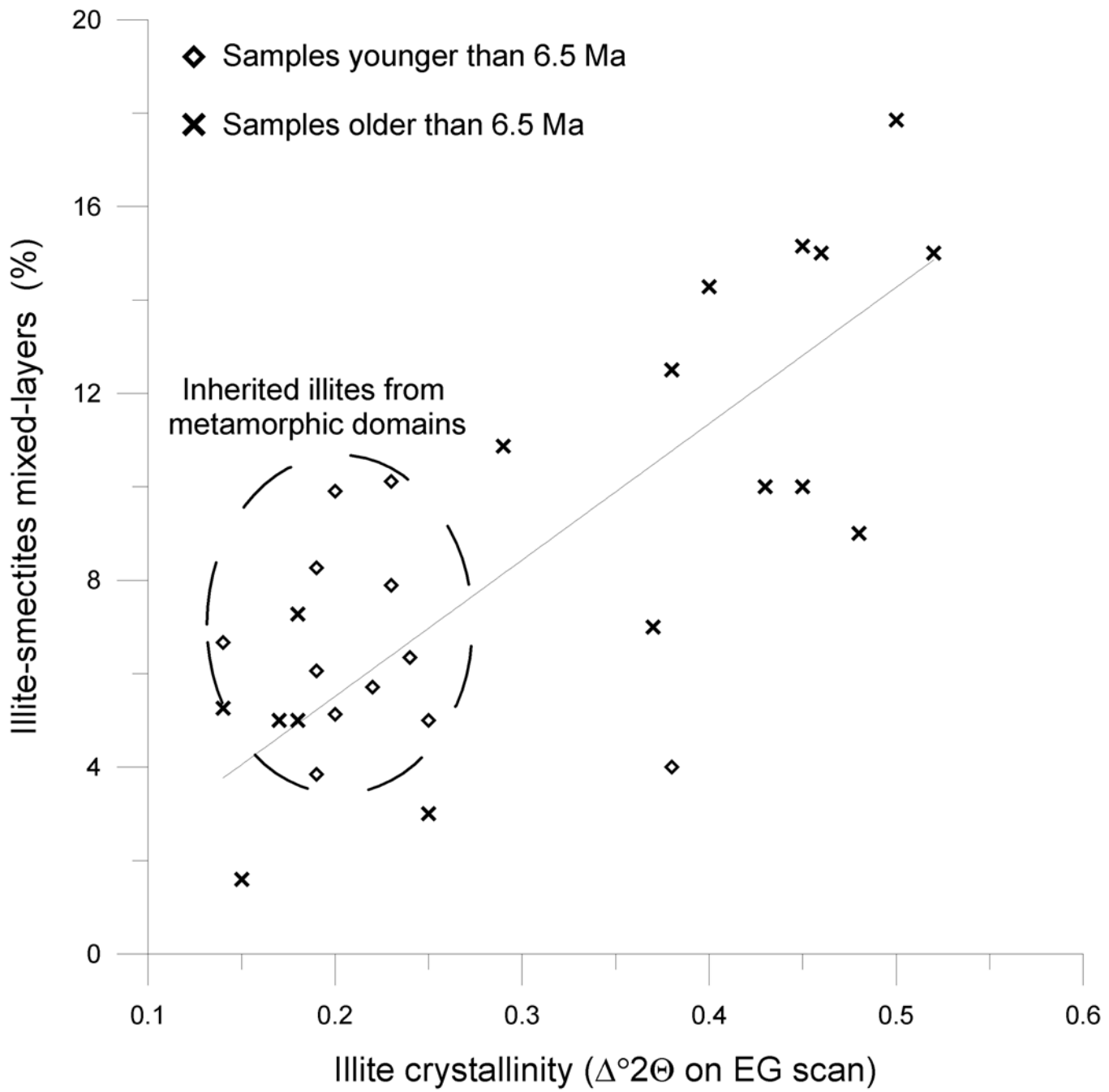


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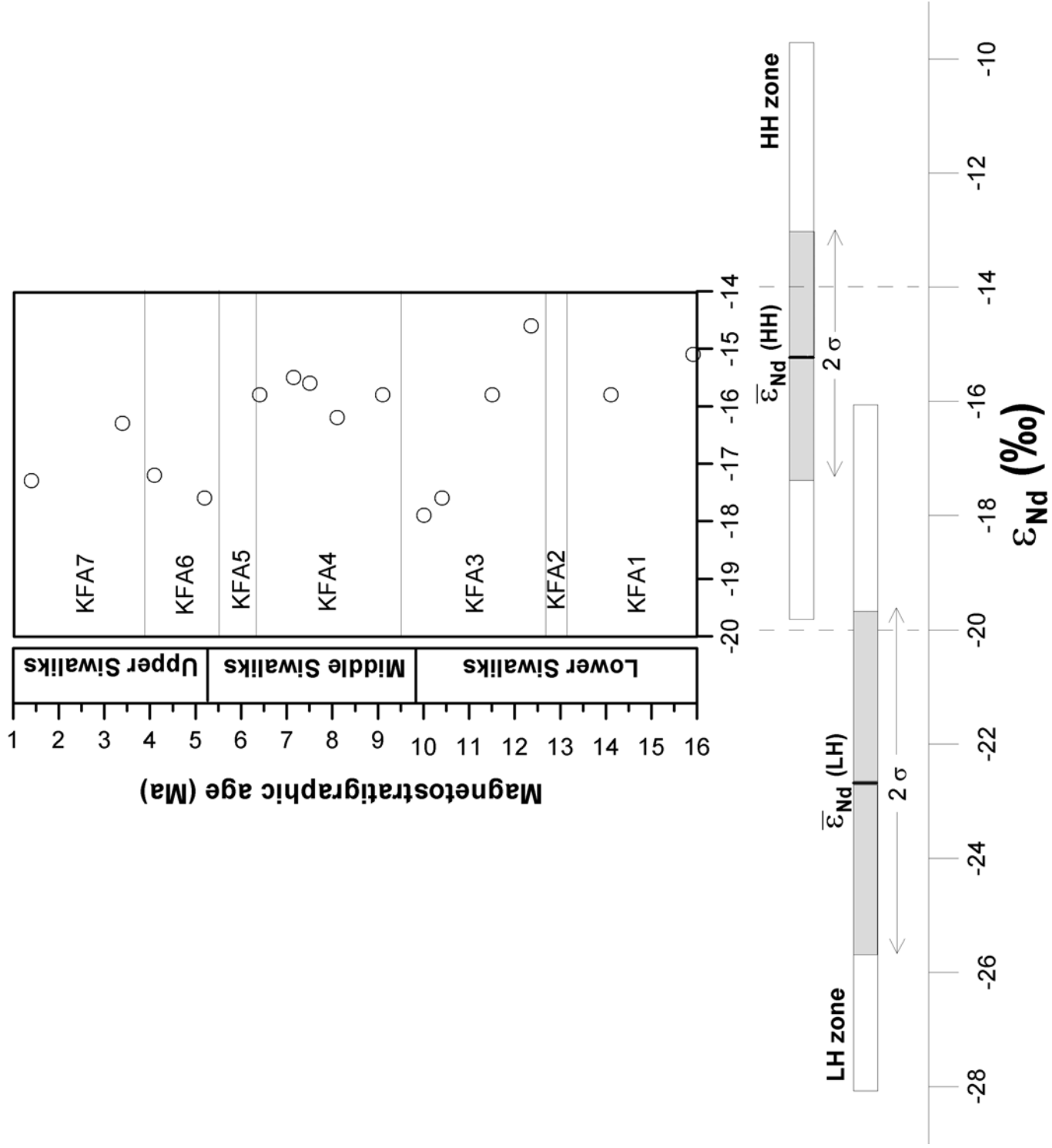


Figure 5

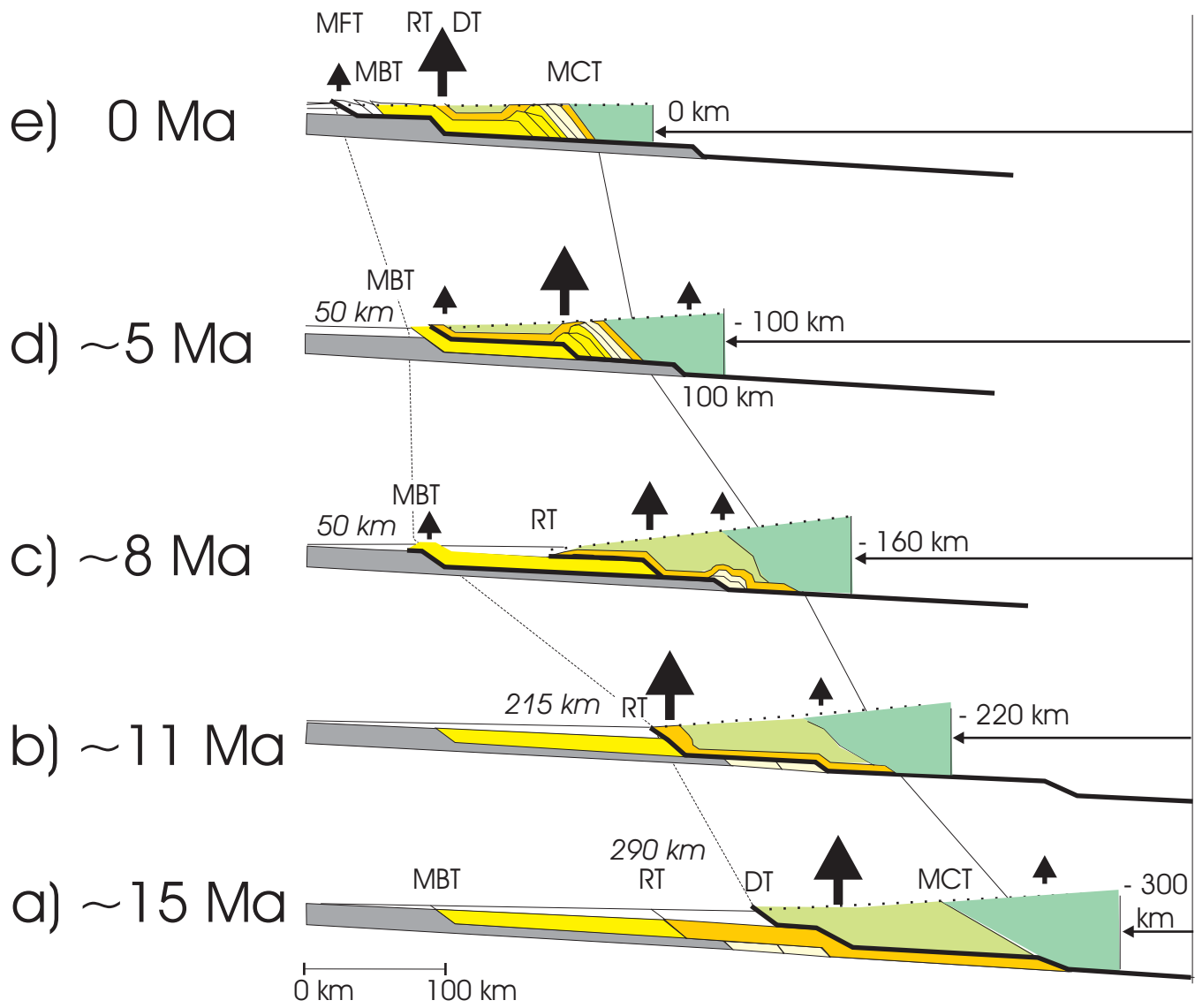


Figure 7