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Comparison of EISCAT and ionosonde electron densities: application to a ground-based ionospheric segment of a space weather programme

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Abstract. Space weather applications require real-time data and wide area observations from both ground- and space-based instrumentation. From space, the global navigation satellite system – GPS – is an important tool. From the ground the incoherent scatter (IS) radar technique permits a direct measurement up to the topside region, while ionosondes give good measurements of the lower part of the ionosphere. An important issue is the intercalibration of these various instruments.

In this paper, we address the intercomparison of the EISCAT IS radar and two ionosondes located at Tromsø (Norway), at times when GPS measurements were also available. We show that even EISCAT data calibrated using ionosonde data can lead to different values of total electron content (TEC) when compared to that obtained from GPS.

Key words. Ionosphere (active experiments; auroral ionosphere; instruments and techniques)

1 Introduction

Both ionospheric and thermospheric monitoring is important in the context of space weather, this having applications in radio communications, navigation and orbital prediction. Space weather requires real-time data and wide area observations but there are few regional or global instrumentation networks.

The TIMED (Thermosphere Ionosphere Mesosphere Energetics and Dynamics) spacecraft serves some of the thermospheric requirements. TIMED, launched in December

2001, is studying the influences of both the Sun and the anthropogenic effects on the mesosphere and lower thermosphere/ionosphere. It includes four instruments, of which the Doppler Interferometer (TIDI) globally measures wind and temperature profiles to investigate the dynamics and energy balance of the Earth's mesosphere and lower-thermosphere. However, the upper altitude of the TIDI experiment is only 180 km, which is relatively low and for most space weather applications it is necessary to use thermospheric models, calibrated via indirect (proxy) measurements (Liliensten, 2001).

The ionosphere is, on the other hand, monitored more comprehensively. The world-wide network of ionosondes and more recently, global navigation satellite system (GNSS) receivers provide a relatively dense ground-based measurement network. GNSS receivers measure the line of sight (slant) total electron content (TEC) (Leitinger, 1998). The world-wide number of two frequency semi-codeless ground global positioning satellite (GPS) receivers providing easily accessible data is 366 as of 28 March 2004 (<http://igsb.jpl.nasa.gov/network/list.html>) and is expected to increase with the advent of the Galileo system. For space weather applications, ionosondes also need to be networked (e.g. Galkin et al., 1999) – in Europe there are around ten such ionosondes providing real-time data directly or via data servers, e.g. Chilton (UK), Juliusruh (Germany), Rome (Italy), Athens (Greece), Pruhonice (Czech Republic) and Tromsø (Norway).

The strengths and weaknesses of these techniques need to be clearly understood. Liliensten and Cander (2003) used the peak electron density ($NmF2$) and its height ($hmF2$) measured by EISCAT as inputs to two profilers (i.e. models based on adjustments of a parameterised profile, expressed in terms of simple mathematical functions). The profilers were run to

Table 1. Characteristics of the five experiments (decimal hours).

| Start day | Start hour (LT) | End day | End hour (LT) | Ap | F _{10.7} |
|---------------|-----------------|---------------|---------------|----------|-------------------|
| 17 Aug. 1998 | 11.28 | 19 Aug. 1998 | 17.11 | 2 to 7 | 132 to 139 |
| 9 Feb. 1999 | 11.34 | 12 Feb. 1999 | 17.26 | 0 to 32 | 124 to 163 |
| 11 Aug. 1999 | 7.33 | 11 Aug. 1999 | 23.33 | 3 to 9 | 127 |
| 15 Sept. 1999 | 16.43 | 17 Sept. 1999 | 17.11 | 12 to 80 | 156 |
| 6 Dec. 1999 | 19.29 | 9 Dec. 1999 | 8.61 | 9 to 22 | 142 to 153 |

compute the corresponding modelled electron density profiles and the modelled TEC was compared to incoherent scatter (IS) radar measurements. This study indicated that when *hmF2* and *NmF2* data were available – but not the whole electron density – ionosonde data might be successfully used with profilers to retrieve the electron density profile and the TEC (see also Huang and Reinisch, 2001 and Belehazi et al., 2003). The intercomparison of ionosondes and positioning systems (at the same location) would also provide an indirect measurement of the protonosphere, located between the upper altitude of the profiler modelling (about 3000 km) and the altitude of the GPS satellites (22 200 km).

In this paper we undertake a further comparison of ionosonde and EISCAT data at times when GPS-TEC measurements were also available. The study particularly addresses the intercalibration of the IS radar electron density (see also Sedgemore et al., 1998).

2 Data sources

Ionosonde data have been obtained from two instruments located on the Tromsø EISCAT site, namely the Dynasonde and the Digisonde (DGS) (Reinisch et al., 1992). For the observations in this paper the Digisonde and the Dynasonde typically operated every 12 min with the parameters *foF2* – being an estimate of the highest frequency with ordinary polarisation reflected from the F2-region – and *hmF2*. Both have been automatically scaled on both instruments. The former is related to *NmF2* according to $foF2 = 8.98 \sqrt{NmF2}$, where *foF2* is in Hz and *NmF2* is in electrons m⁻³. The Digisonde data were autoscaled and were then manually scaled as a check. Weak echoes were obtained on many occasions making interpretation difficult.

In this paper, we focus on EISCAT data from the Tromsø UHF (931 MHz) radar (e.g. Rishbeth and Van Eyken, 1993). The details of the background theory have been reviewed in Bauer (1975). The analysis of the scattered signal provides height profiles of the ion and electron temperature, the ion velocity parallel to the magnetic field line, and the electron density. From the electron density profile, the true-height of the peak, corresponding to *hmF2*, its density value, corresponding to *NmF2*, and therefore its plasma frequency, corresponding to *foF2* can be determined. We call these quantities *foF2*(IS) and *NmF2*(IS).

The calibration to absolute density must be done by comparison with other techniques. Once done, this calibration is in principle valid for a long period, assuming no changes to the radar hardware.

3 EISCAT – GPS comparison: origin of the problem

In a recent study Lilensten and Cander (2003) found five periods with common EISCAT and GPS observations at the Tromsø site and in this paper, we again focus on these periods which correspond to well calibrated GPS-TEC and EISCAT CP1 data availability. In this EISCAT mode, the Tromsø beam is pointing parallel to the local magnetic field line. The integration time for processing the raw data is 1 or 2 min. The measurement of the ionospheric profiles is performed between 90 and 498 km with a computational accuracy of ~5% for the electron densities (Lathuillère, 1994; Lathuillère et al., 2002). These five experiments cover 257.75 h during 15 days. Table 1 shows that the experiments cover a wide range of solar and magnetic activities, from very quiet to disturbed.

The source of GPS data is the International GPS Service for Geodynamics receiver at Tromsø. TEC estimates along all GPS satellite links in view for elevation angles greater than 10° are derived using a technique developed by Ciralo (1993, 2000) and described recently in detail by Cander and Ciralo (2002). Assuming a single-layer approximation for the ionosphere, these slants TEC data are then converted to equivalent vertical values at the intersection point of the ray-path with an ionospheric shell fixed at 400 km height. The conversion of the slant TEC to the vertical TEC may of course be a source of error, especially at high latitudes where the shell approximation can break down. The accuracy of the TEC deduced from GPS data is of the order of 2 TECU (1 TECU=10¹⁶ electrons.m⁻²).

Figure 1 shows a comparison between GPS-TEC and the IS-TEC₄₉₈ (IS-TEC₄₉₈ denotes TEC up to a height of 498 km) derived from the EISCAT measurements. This last value has not been corrected by ionosonde calibration. In several cases, the TECs are close to or even smaller than the IS-TEC₄₉₈ – see, for example, 15 to 17 September 1999. This should never happen, since the radar observes between the ground and 498 km while the GPS-TEC is an integration between the ground and 20 200 km. In order to make

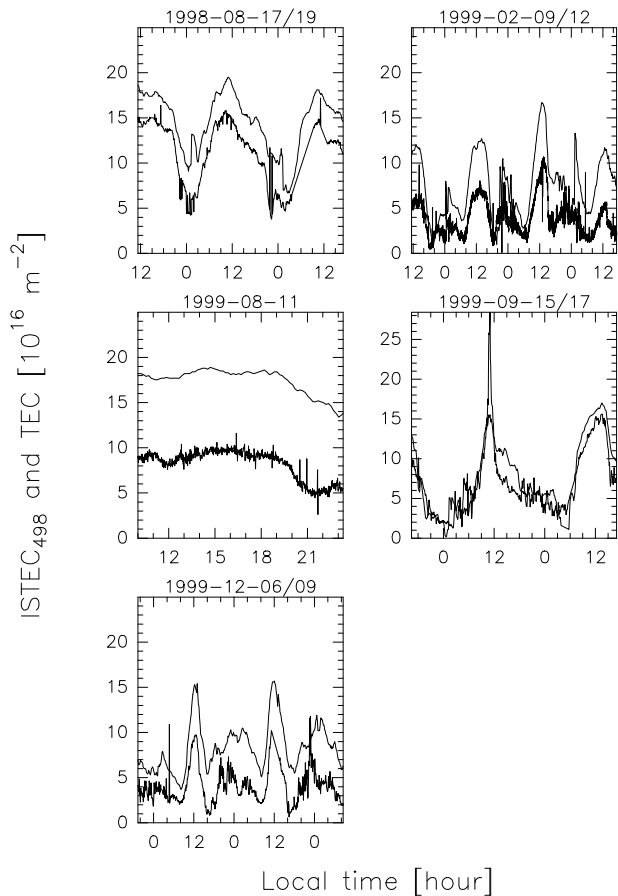


Fig. 1. Comparison between GPS TEC (thin lines) and the ISTECS₄₉₈ derived from the uncalibrated EISCAT measurements (bold lines).

sure that this is not an artefact of the uncalibrated IS data, it is necessary to compare the uncalibrated IS radar data with ionosonde data.

4 EISCAT–ionosonde comparisons

In this study, the Digisonde only has two common periods, 15–17 September 1999, and part of 6–9 December 1999. The Dynasonde data covers all five experimental periods.

In the following figures, we show comparisons between the IS measurements and the ionosondes. We compare *fo*F₂, the squared ratio being the correction factor that should be used for the IS-TEC. For ease of plotting, the hours are local time (LT) on the first day of the experiment, LT + 24 on the second day, LT + 48 on the third day, etc.

For the period 15–17 September 1999 (Fig. 2) *fo*F₂ is generally closer to the IS measurements than the Dynasonde but both ionosondes give values smaller than those measured by the IS radar. When the least reliable ionosonde measurements have been eliminated, the mean *fo*F₂(NmF₂) ratio between the ionosonde and the IS-*fo*F₂ is 1.1 (1.23) for the Dynasonde and 1.07 (1.14) for the Digisonde. When we

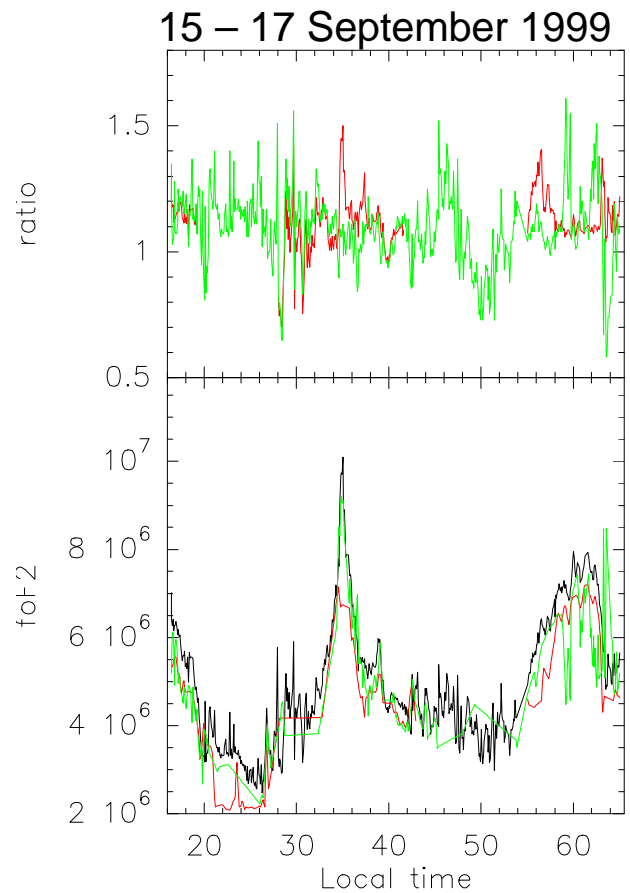


Fig. 2. *fo*F₂ comparison between the IS radar data (black), the Dynasonde (red) and the DGS (green). The top panel shows the ratio between the ionosonde-*fo*F₂ data and the IS-*fo*F₂ data in the same colour scheme.

combine the data from both ionosondes and keep the points temporally closest to the IS points, the ratio is 1.01.

For the period 6–9 December 1999 (Fig. 3) we obtain the opposite behaviour: both ionosondes estimate *fo*F₂ to be higher than the IS radar. The Digisonde data were only available for the last two days of the experiment. On average, the *fo*F₂ ratio between the Dynasonde and the IS-measurement is 0.91 (0.83) and 0.88 (0.77) with the Digisonde. When we use the data from both ionosondes and keep the points temporally closest to the IS points, the ratios are 0.94 (0.88).

For the period 17–19 August 1998 (Fig. 4) we only have data from the Dynasonde (red). It gives a ratio of 1.13 (*fo*F₂) or 1.28 (NmF₂). This ratio is computed with the part of the experiment shown in the upper panel of Fig. 4 but even when we include the complete period (and, therefore, some less reliable points), the ratio remains approximately the same.

The 9–12 September 1999 experiment (Fig. 5) is one of those where the TEC measured by the IS radar is very close to the TEC obtained from the GPS signals. It is interesting to notice that the *fo*F₂ measured by the Dynasonde fluctuates around the *fo*F₂ measured by the radar. On average, the ratio is very close to 1. However, if we were using a ratio

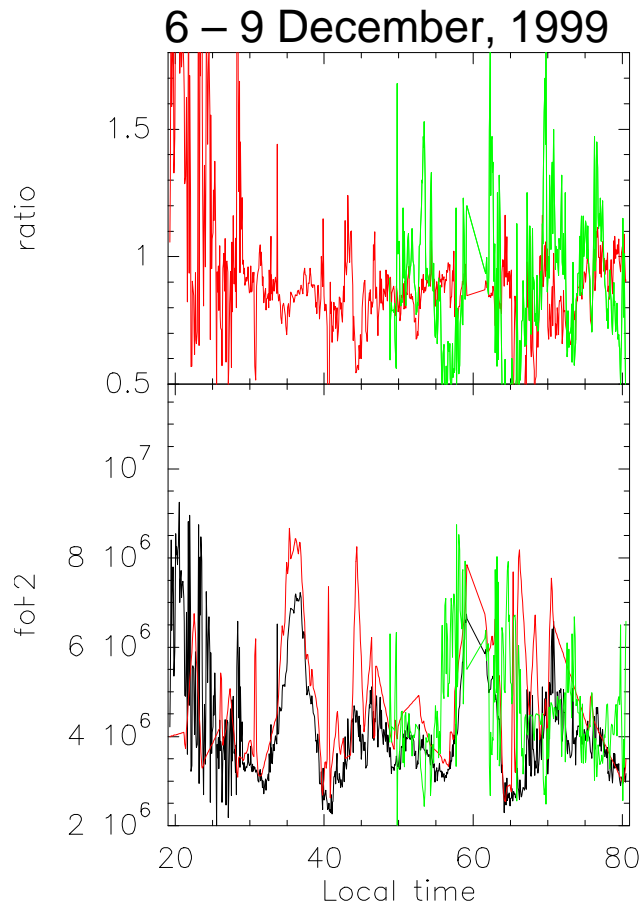


Fig. 3. $foF2$ comparison between the IS radar data (black), the Dynasonde (red) and the DGS (green).

dependent on time, we would obtain an increase of the IS- TEC_{498} when it is already close to or larger than the GPS data. This is especially true during the period 65 LT to 75 LT (i.e. 11 February, around 20 LT).

The agreement between the ionosonde and the IS- $foF2$ is almost perfect during most of the 11 August 1999 experiment (Fig. 6) but at the end of the experiment, the ionosonde critical frequency drops faster than that from the IS radar.

The drop corresponds to weak echoes and non-continuous traces such that the cusp near the penetration frequency is not properly measured. Under such conditions the ionograms may not give $foF2$ reliably, this being most likely due to absorption, possibly compounded by broadcast or other interference.

5 Effect of ionosonde calibration on the IS- TEC estimates

We conclude that if we calibrate EISCAT using the ionosonde measurements, we should use the following factors: 17–18 August: 1.28; 9–12 February: 1; 11 August: 1; 15–17 September: 1.14; 6–9 December: 0.83.

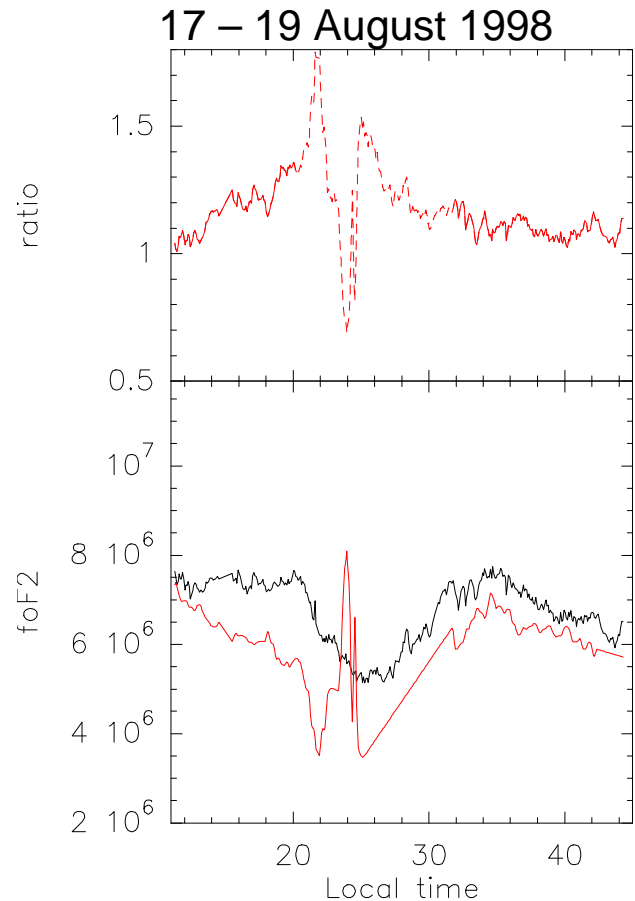


Fig. 4. $foF2$ comparison between the IS radar data in black and the IS-ionosonde in red. In the upper panel, the full line shows the points that have been kept to compute the mean ratio.

Figure 7 is similar to Fig. 1 but now we plot the IS- TEC_{498} measured by the EISCAT after correction by the ionosondes together with the TEC measured by GPS.

Even after calibration, the periods when the IS- TEC_{498} is larger than the GPS- TEC remain. Moreover, the last experiment in December 1999 (bottom panel) shows that the IS data are sometimes identical to the GPS data, which is impossible, since the IS measures only up to 498 km while the satellites measure between the ground and 22 200 km. This last experiment is the only one to have a correction factor smaller than 1. The ratio between GPS- TEC and IS- TEC is about 2 before correction and 1.7 afterwards. A value of ~ 2 would be expected because the electrons below 500 km, on the one hand, and in the range 500–22000 km, on the other hand, account for half of the total each (Lilensten and Bllely, 2002).

6 Interpretation

There may be practical reasons for observing discrepancies between the ionosondes and the IS radar measurements of $foF2$. For example, the region illuminated by the ionosonde

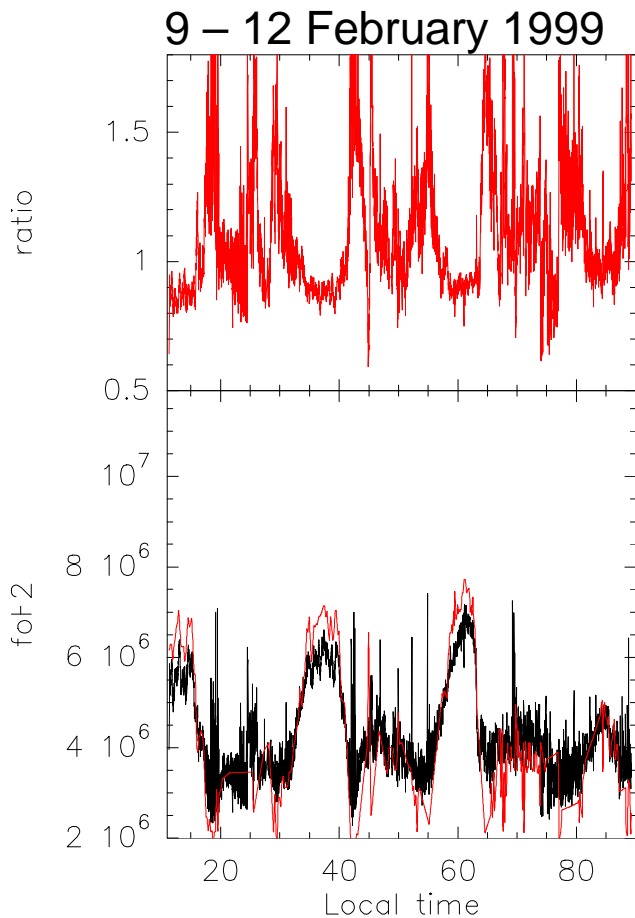


Fig. 5. $foF2$ comparison between the IS radar data in black and the Dynasonde.

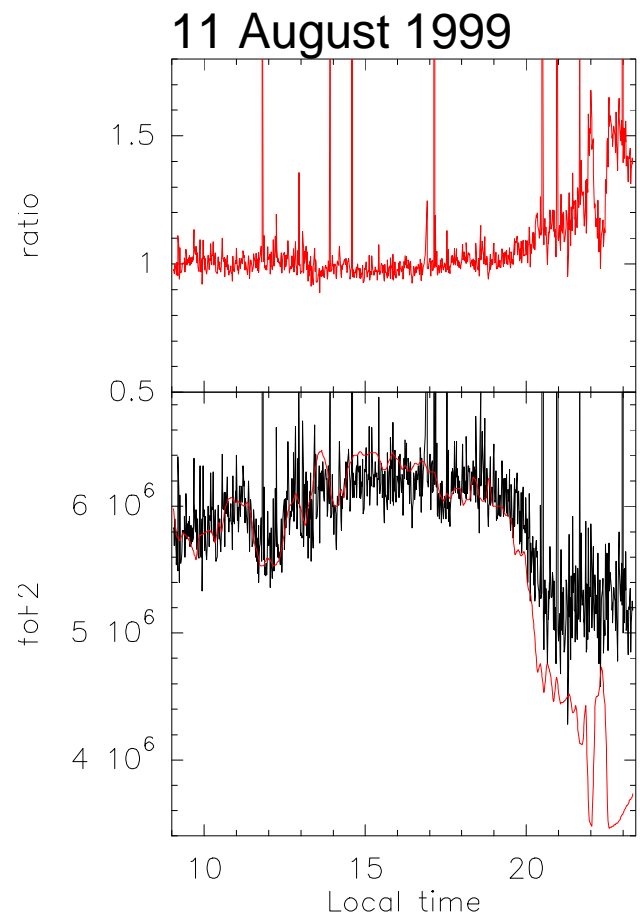


Fig. 6. $foF2$ comparison between the IS radar data in black and the Dynasonde (red).

is wider than that associated with EISCAT measurements. Ionosonde echoes are only returned from regions where the refractive index surface lies perpendicular to the ray path. In a spherically stratified ionosphere this is a relatively narrow region overhead, appropriate to the Fresnel radius at the measurement frequency. In a nonuniform ionosphere signals can originate from off the vertical - indeed from several regions - none of which necessarily correspond to the EISCAT volume. The ionosonde echo might originate from localised high-density structures which may result in an $foF2$ measured by ionosonde that is higher than a measurement based on the mean density measured by EISCAT.

By comparison, EISCAT measures the average density in a narrower cone: the EISCAT aperture is 0.9° and the height integration in the F-region is 3 km with the power profile data and 22.5 km with the long pulse. In other words, EISCAT measures in a volume, whereas ionosondes measure reflection heights which are relatively thin, but within large areas.

In selecting the echoes from the Dynasonde to compare with the IS radar, there was no strict requirement that the echoes should come from exactly within the EISCAT volume, but most echoes had a direction of arrival of less than 20° from zenith.

The discrepancies, although small, between the two ionosondes located only 100 m apart raises another issue. Ionospheric absorption, perhaps by underlying auroral-E, broadcast or other radio interference, can limit the strength and quality of echoes. The determination of $foF2$ depends on the strength of the echoes, on the software and hardware which registers these echoes and on the quality of the automatic or human data reduction approach, especially under conditions of spread-F.

As shown in this study, the calibration of the EISCAT Incoherent Scatter radar through ionosonde data is not a trivial task. However, it is important to calibrate EISCAT as often as possible, even continuously. The best way to proceed is to use the plasma line (Perkins and Salpeter, 1965; Perkins et al., 1965). In principle, it is a measurement that can be made by any IS radar, provided that the plasma line is enhanced by photo- or auroral electrons or artificially by high-power HF waves. It is directly linked to the absolute electron density, but varies in strength with time and altitude. Most of the plasma lines experiments have been difficult to perform because the plasma line can vary quickly (Nilsson et al., 1996; Guio et al., 1998). The new data-taking program tau2_pl at EISCAT now allows for routine measurements of the plasma

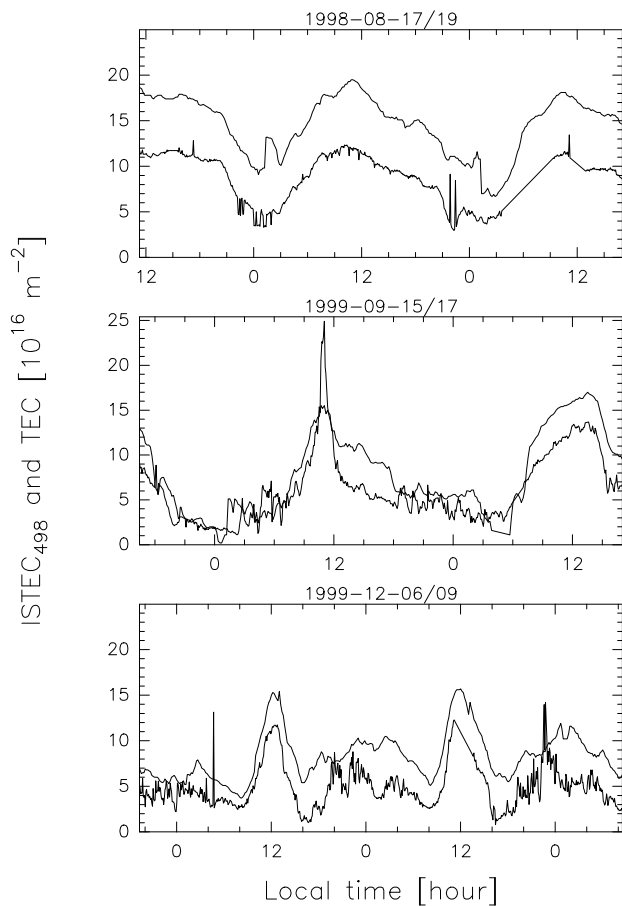


Fig. 7. Comparison between the TECs derived from the GPS stations (thin lines) and the calibrated ISTEC₄₉₈ derived from the EISCAT measurements (bold lines). Only the three corrected periods are shown.

line and, therefore, the direct and continuous calibration of the ion line.

The importance of ionosonde and GPS measurements to the Space Weather programme is well accepted. However, a better understanding the generality of both TEC derived from *hmF2* and *NmF2* profilers and GPS instrumentation used in regions where the thin shell approximation is likely to break down is needed. The IS technique provides a third reference technique to study this. Therefore, we suggest that it is important to (i) calibrate several IS radars with the plasma line technique, and (ii) compare GPS, ionosonde and IS radar data over a large set of cases in order to better understand the differences.

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